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THE UNIVERSITY OF ARIZONA

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Communications of the Lunar and Planetary Laboratory

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No. 166 ARIZONA-NASA ATLAS OF THE INFRARED SOLAR SPECTRUM, REPORT X

by D. C. Benner, G. P. Kuiper, L. Randić, and A. B. Thomson

May 1, 1970

ABSTRACT

This report concludes the series of the Arizona-NASA Infrared Solar Spectrum Atlas; it follows LPL Comm. No. 162 (which contains Charts B1-B9) and thereby completes the B series. The wavelength interval covered is 13350-34100 Å.

The present paper contains the second and last set of B-spectrometer records, obtained on the NASA CV-990, July 1968, with a 600-line per mm grating. There is a small overlap with the 1200 line/mm records; B9 terminates at 14098 Å, whereas the present series begins at 13350 Å. The value of the present series lies primarily in serving as reference to high-altitude planetary spectra taken with similar resolutions, 6-10,000. The 4-meter records, with their 4X higher resolution, give a more detailed and valuable description of the solar spectrum as such and, incidentally, assist in the interpretation of the intensity tracings here reproduced. Because of inherent uncertainties in spectrometer traces recorded in flight, the present records contain valuable checks on the solar and telluric spectra published before.

The wavelength scale and the classification of absorption lines are taken from the 4-meter spectra. For the interval 13350-28500 Å, two good independent B records were available and are here reproduced. For the 28500-34000 Å, mostly only one good record was available, with the second set used only to verify the set here published. For the 3.3 μ region (ν_3 of CH₄) a B-spectrometer trace had already been reproduced in *LPL Comm.* No. 125, page 223 (having slightly better resolution).

Table I contains the relevant data on each of the charts. In some spectra the continuum has a somewhat wavy appearance (e.g. on B15), due either to minor interference effects in the optical train (possibly a thin filter so oriented as to cause the double reflection to strike the detector); or else, to minor guiding errors (guiding was on the center of the solar disk but occasional uncompensated motions of the aircraft caused the solar image on the slit to shift. An extreme case on B13 at 17500 Å has been marked "g"). On B13-16 and B19 some cases occur of lines slightly displaced and distorted (apparently due to a sharply localized defect in the grating drive that may show its appearance after one or more revolutions, some 480 Å apart); they are marked "b" and are not marked "real" (dot below) if displaced (in such a case a real solar line will be missing among the numbered lines). On B12 and B13, a and b, the 1 μ filter (Corning 2540) was used, causing a minor second-order leak near 1.8 μ , similar to that discussed for the 4-meter records, LPL Comm. No. 163, p. 66. Attention is drawn to P α on B13, still cut by telluric H₂O. Good records were obtained of the three telluric CO₂ bands at 2 μ and the two stronger bands at 2.69 and 2.77 μ (including isotopic components).

Identifications have been added where deemed useful. As in the other Reports, all absorptions regarded real, either solar or telluric, are indicated by dots *below* the traces, numbered consecutively for future reference. Dots *above* the traces designate telluric H₂O absorptions; open triangles, CH₄; short vertical lines, CO₂; crossed open circles, N₂O; and asterisks, CO. Solar identifications are added where known and considered appropriate. Partial contributions, in excess of $\frac{1}{4}$ of the total intensity, are placed in parentheses. Acknowledgments — We are indebted to NASA Hq. and NASA-Ames for their support of this program; to Messrs. J. Percy and B. McClendon for assistance with the electronics during the flights; to Mrs. A. Agnieray and Mr. S. Larson for their assistance in the preparation of the charts for publication, which in this case was especially laborious. The program was supported by NASA through Grants NsG 161-61 and NGR-03-002-091.

ERRATA IN REPORTS I - IX

In the preparation of Report X ($\lambda > 13400A$), a few rather minor errata were detected in the earlier reports, which are listed below. The verification was not exhaustive but indicates very few errors among the thousands of identifications.

CHART	LINE	CORRECTION
25a	30	not H ₂ O
25d	30	partly H ₂ O
26a	5	H-O
26a	20	H ₂ O
26a	77	H2O
26a	84	H ₂ O
26a	97	not H₂O
27b	29	partly H ₂ O
27Ь	35	H₂O
34d	19, 21	not H₂O
34d	54	omit (:)
38a	75	add · below
39a	28	omit ()
47c	10	add (🕅)
B8c	5	add 🖸
B9c, d	2	partly H ₂ O
B9c, d	22	omit ()
B9c	39	H₂O
B9c	48	omit ()

FIG.	Снаят	λ(Å)	1968 Date	UT	ALT. (FT.)	Outside Temp, (°C)	CABIN Alt, (FT.)	Gain
1	B10 a* b* c* d*	13345-14200 13410-14200 14200-15033 14200-15033	July 17 July 17 July 17 July 17 July 17	19:52 19:39 19:55 19:43	39,000 39,000 39,000 39,000 39,000	51 51 51 51	8500 8500 8500 8500	4-1 4-1 4-1 4-1
2	B11 a* b* c* d*	15033-15850 15033-15850 15850-16664 15850-16664	July 17 July 17 July 17 July 17 July 17	19:58 19:46 20:01 19:48,20:06	39,000 39,000 39,000 39,000	51 51 52 51,52	8500 8500 8500 8500	4-1 4-1 4-1 4-1, -1, -2
3	B12 a* / b* c* d*	16664-17483 16664-17483 17483-18282 17483-18282	July 17 July 17 July 17 July 17 July 17	20:09 20:26 20:30 20:12	39,000 36,500 35,000 39,000	52 52 50 52	8500 8500 8500 8500	4-2 4-2 4-2 4-2
4	B13 a* b* c d	18282-19070 18282-19070 19070-19800 19070-19800	July 17 July 17 July 15 July 15	20:16 20:33 20:15 19:59	39,000 35,000 40,000 40,000	-52 -50 -54 -54	8500 8500 8000 8000	4-2 4-2 4-4 4-4
5	B14 a	19800-20529 20529-21250 20520-21250	July 15 July 15	20:18 20:32	40,000 40,000	-54 -54	8000 8000	4-4 4-4
6	B15 a b c d	21250-21250 21250-21986 21250-21986 21986-22660 21986-22660	July 15 July 17 July 17 July 17 July 17 July 15	18:25 18:20 18:29 20:39	39,000 39,000 39,000 39,000 40,000	54 49 49 54	8500 8500 8500 8500 8000	4-4 4-1 4-1 4-1 4-5
7	B16 a b c d	22660-23345 22660-23345 23345-24006 23345-24006	July 17 July 15 July 17 July 17 July 17	18:31,18:37 20:42 18:52 18:41	39,000 40,000 39,000 39,000	-49 -45 -49 -49	8500 8000 8500 8500	4-1 4-5 4-2 4-1, 4-2
8	B17 a b c d	24006-24645 24006-24645 24645-25270 24645-25270	July 17 July 17 July 17 July 17 July 17, 18	18:44 18:55 18:58 18:47,19:09	39,000 39,000 39,000 39,000 39,000	-49 -49 -50 -49, -56	8500 8500 8500 8500	4-2 4-2, 4-3 4-3 4-2, 4-3
9	B18 a b c d	25270-25873 25270-25873 25873-26440 25873-26440	July 18 July 18 July 18 July 18 July 15	19:02 19:11 19:05 19:15	39,000 39,000 39,000 40,000	56 56 56 50	8500 8500 8500 8000	4-3, 4-4 4-3, 4-4 4-4 5-4
10	B19 a b c d	26440-26965 26440-26965 26965-27461 26965-27461	July 18 July 15 July 15 July 18	19:26 19:17,19:21 19:23 19:29	39,000 40,000 40,000 39,000	56 50 51 56	8500 8000 8000 8500	5-5 5-4, -5, -6 5-6 5-5
11	B20 a b c d	27461-27990 27461-27990 27990-28498 27990-28498	July 15 July 15 July 15 July 15 July 15	19:25, 19:39 19:29 19:42, 19:54 19:31, 19:09	40,000 40,000 40,000 40,000	-51, -52 -52 -53 -52, -50	8000 8000 8000 8000	5-6, 5-5 6-2 5-5, -5, -6 6-2, 5-5
12	821 a b† c† d†	28498-28980 28980-30705 30705-32430 32430-34100	July 15 July 19 July 19 July 19 July 19	19:56 19:42 19:45 19:45	40,000 39,000 39,000 39,000	$ \begin{array}{r} -53 \\ -52 \\ -52 \\ -52 \end{array} $	8000 8500 8500 8500	5-6 6-2 6-2, 6-3 6-3

TABLE I Solar Spectrum Records, B-Spectrometer, NASA CV-990 Jet 2.5 μ Grating (600 lines/mm), Slit and Cell 0.10 mm, τ = 0.12, 2 μ Filters

* 1 μ filter, † 4 μ grating (300 l/mm), 2.4 μ filter, slit 0.30 mm, $\tau = 0.6$ sec.



Fig. 1 B-spectrometer record of solar spectrum $\lambda\lambda$ 13345-15033.



Fig. 2 B-spectrometer record of solar spectrum $\lambda\lambda$ 15033-16664.



Fig. 3 B-spectrometer record of solar spectrum λλ 16664-18282.



Fig. 4 B-spectrometer record of solar spectrum λλ 18282-19800.



Fig. 5 B-spectrometer record of solar spectrum λλ 19800-21250.

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Fig. 7 B-spectrometer record of solar spectrum $\lambda\lambda$ 22660-24006.



Fig. 8 B-spectrometer record of solar spectrum λλ 24006-25270.





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Fig. 10 B-spectrometer record of solar spectrum λλ 26440-27461.



Fig. 11 B-spectrometer record of solar spectrum λλ 27461-28498.

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INFRARED SOLAR SPECTRUM

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No. 167 NARROW-BAND PHOTOMETRY OF THE GALILEAN SATELLITES

by Willem WAMSTEKER

May 30, 1972

ABSTRACT

Observations were made of the Galilean satellites in a narrow-band photometric system with a resolution of ~ 30 that covers the wavelength range from 0.3-1.1 μ . The absolute scale of the system is determined from stellar model computations. It is concluded that 35 Leo, the main comparison star used, has an energy distribution similar to the sun. An attempt is made to interpret the albedo curves in terms of surface constituents.

1. Instrumentation

This paper is an attempt to use low-resolution spectrometry in the study of the surfaces of the Jovian satellites. This method was, among others, used successfully by McCord *et al.* (1970) on the asteroid Vesta, which led to the identification of pyroxene.

The Laboratory had acquired a set of 42 narrowband interference filters from Infrared Industries, Thin Film Products Division, 62 Fourth Avenue, Waltham, Massachusetts. The bandwidth of the filters had been selected to represent a constant resolution of $\lambda/\Delta\lambda \sim 30$ over the full wavelength range $0.3-1.1\mu$. The filters are mounted in three separate wheels each containing 14 filters, with the last filter repeated in the next set. The filter wheels were adapted to a single-channel photometer of the standard Johnson type (Johnson and Mitchell 1962).

An S-20 cathode was used for $0.30-0.72\mu$ and an S-1 for $0.72-1.1\mu$. The observations were made with the 61- and 40-inch telescopes of the Catalina Observatory, around the 1971 opposition of Jupiter. Owing to the low declination of Jupiter and the unfavorable summer observing conditions, the coverage is not as extensive as was intended. However all satellite data presented are an average of at least two observations on different nights. Atmospheric extinction was allowed for by average extinction coefficients derived independently, checked on some nights by the observation of stars at small and large air masses. Allowance was also made for nights with high and low humidity within the well-known atmospheric water-vapor absorptions.

Table I gives the observations as "normalized" albedos, with the maximum put at unity. The values are ratios of the satellite intensities in terms of HR 4030 (35 Leo). Wavelength regions observed at different orbital phases of the satellites were adjusted to each other by means of the filters common to both data sets. The wavelengths in the first column of Table I are those midway between the 50 percent transmission points of the filters, assuming a flat radiation source.

2. Comparison Stars

A number of G2V-stars were observed to obtain sources with spectra similar to the sun. However, the average solar-type star is not identical to the sun, as shown by Code (1960). To acquire an independent estimate of the similarity of the spectral distributions of sun and comparison stars, we also observed a number of stars included in the absolute stellar calibration by Hayes (1970). The disadvantage of this calibration is that it applies only to regions which are free of spectral lines. To overcome this problem, we used Mihalas' (1966) hydrogen line-blanketed models. The star which was most suited for this purpose was HR 4468 (θ Crt, B9V), which is included in Hayes' calibration, has a spectral type covered by Mihalas' (log g, θ)-range, and is early enough to be considered still undisturbed by metal absorptions at the resolution of our photometric system.

A first direct comparison between our photometric system and Hayes' calibration, in which the energy distributions relative to HR 4468 are taken, showed both systems to be consistent within 2 percent in regions free of lines. The spectral range covered by this comparison was from B3V to A1V. Shortward of 0.36μ the deviations were somewhat larger for B3V but still did not exceed 5 percent.

The stellar model chosen for HR 4468 is specified by Mihalas with the parameters log g=4, θ =0.45. To convert the model into our photometric system, the filter bandpasses were assumed to be square.

With the information from the model it is possible to construct an energy distribution for each

Table I

Normalized Relative Reflectivities of Galilean Satellites

λ(μ)	То	Europa	Ganymede	Callisto
0.300	0.216	0.332	0.476	0.392
0.308	0.108	0.408	0.389	0.436
0.318	0.075	0.370	0.375	0.382
0.328	0.096	0.378	0.386	0.403
0.340	0.096	0.410	0.446	0.440
0.350	0.097	0,422	0.490	0.483
0.358	0.104	0.448	0.547	0.515
0.368	0.132	0.491	0.582	0.547
0.378	0.123	0.510	0.591	0.552
0.390	0.164	0.570	0.639	0.612
0.399	0.213	0.592	0.683	0.639
0.423	0.331	0.673	0.723	0.683
Q.444	0.462	0.751	0.785	0.756
0.457	0.566	0.810	0.833	0.783
0.478	0.671	0.857	0.870	0.823
0.498	0.749	0.891	0.905	0.858
0.518	0,776	0.895	0.919	0.885
0.538	0,794	0.898	0,920	0.906
0.560	0.801	0.920	0.942	0.934
0.578	0.813	0.937	0.945	0.951
0.599	0.847	0.964	0.969	0.975
0.615	0.877	0.968	0.995	0.991
0.637	0.954	0.983	0.993	0.996
0.652	0.954	0.994	1.000	0.986
0.672	0.983	1.000	0.997	0.998
0.698	1,000	0.998	0.985	0.989
0.717	0.994	0,994	0.978	0.986
0.746	0.986	0.986	0.987	1,000
0.767	0.979	0.996	0.967	0.984
0.787	0.963	0.975	0.960	0.981
0,808	0.960	0.988	0.966	0.988
0.838	0.925	0,960	0.931	0.962
0.862	0.948	0,992	0.950	0.985
0.894	0.933	0.976	0.912	0,958
0.923	0.911	0.998	0,910	0,951
0.950	0.921	0,982	0.894	0.960
0.996	0.887	0,956	0.864	0.931
1.032	0,868	0,910	0.830	0.934
1.080	0.831	0.874	0.823	0.916
1.105	0.805	0.860	0.800	0.883
				•

observed star. The energy distributions thus derived for a B3V star and an A1V (resp. HR 3454 and HR 4963) are from 0.45 μ to 0.70 μ , similar to Hayes' calibration for the regions where the influence of the hydrogen lines is negligible. Over this region the deviations do not exceed 2 percent. Shortward of 0.36μ the deviations were larger, if the energy distributions were matched over the range from 0.45μ to 0.70μ . This, however, can easily be caused by a slight mismatch of the model parameters. This seems to be the case, because the deviations are of the same order of magnitude and sign both for the A1V star and the B3V star. Since the deviations in this case were somewhat larger than the deviations between our system and Hayes' as mentioned before, at $\lambda < 0.36\mu$ we applied a correction for this.

The corrections, which are listed in Table II, are such that the energy distribution derived by the model has to be multiplied by the correction factor to be consistent with Hayes' calibrations.

Energy Distribution					
λ(μ)	Correction Factor				
0.300	0.88				
0.308	0.88				
0.318	0.88				
0.328	1.00				
0.340	1.05				
0.350	1.50				

TABLE II Correction Factors Applied to Model-Derived Energy Distribution

The two shortest wavelengths listed in Table II were assumed to need the same correction factor as determined for $\lambda = 0.318\mu$; this may not be strictly correct in view of the trend longward of 0.318μ . At the long wavelength end of the spectrum there still remains the Paschen jump and the Paschen lines. However, the filter transmission bands are wider

in that region $(\Delta \lambda \sim 0.03 \mu)$. The Paschen jump, which is not very large, will tend to be smoothed out by the bandwidth and the effect of overlapping lines. For these reasons, we smoothed the energy distribution of the stellar model (log g = 4, θ = 0.45) through towards the longer wavelengths.

Figure 1 shows the energy distribution of HR 4030 (35 Leo) thus derived, from Mihalas' model (log g = 4, $\theta = 0.45$) and the observations of HR 4030 and HR 4468. The dotted line is the energy distribution of the sun from Labs and Neckel (1968). Although in detail there are some discrepancies, the overall agreement is very good. At the two shortest wavelengths HR 4030 seems too high, but that could be caused by the fact that the correction factor derived for 0.318 was used. The match for the other solar-type stars observed was considerably worse. These will be discussed elsewhere.



Fig. 1. Comparison between flux distributions of the sun (dotted line: Labs and Neckel 1968) and 35 Leo (HR 4030). Unit of F_{λ} is Wcm^{-2µ-1} x const.

3. The Albedo Curves

Figure 2 shows the albedo wavelength curves for the Galilean satellites. The left-and-right hand ordinates are respectively normalized and geometrical albedos. The geometrical albedos used here are described in Appendix 1. The error bars drawn in Figure 2 are the average deviations from the mean. The normalized reflection curves for all four satellites look similar. They all show a decrease in the albedo towards shorter wavelengths. For Io the decrease starts around 0.7μ and for the three outer ones at ~ 0.6μ . From 0.6μ longward the albedo curves are more or less flat with a slight decrease at the longest wavelengths. Apart from a shallow absorption feature between 0.5μ and 0.6μ , the



Fig. 2 Wavelength dependence of geometrical albedos of Galilean satellites; normalized albedos, P_n at left; geometrical albedos, P, derived in Appendix I, at right. Representative error bars show average deviations from the mean.

albedo curves are without any distinct features. This shallow feature is not present in the curve for J IV.

The curves are similar to the albedo-wavelength dependence found by Johnson (1971), but not identical. In the present data the absorption feature at $\lambda \sim 0.55\mu$ shown by Johnson for Io seems to be present also in the reflectivity curves for Europa and Ganymede, although less pronounced. Further,

Johnson's data show a stronger decrease in albedo longward of 0.95μ .

The similarity between the four satellites in the albedo-wavelength dependence exists only in the "normalized" curves. If one considers the geometrical albedos, a quite different picture emerges. In fact, the right-hand ordinates of Figure 2 show that the satellite albedos differ by a factor of 4.



Fig. 2, continued.

Table III gives the highest and lowest albedos measured for each satellite, together with the wavelengths at which they occur.

TABLE III Maximum and Minimum Albedos of Galilean Satellites $(0.3\mu < \lambda < 1.1\mu)$

	λ ΜΑΧ	P MAX	λ ΜΙΝ	P MIN			
JI	0.70	1.18	0.32	0.09			
1 II	0.67	0.98	0.30	0.32			
J III	0.65	0.58	0.32	0.22			
J IV	0.75	0.30	0.32	0.12			

Table III further indicates that, although the "normalized" albedo curves are similar in appearance, this does not imply similar surface materials, unless one assumes for each satellite an appropriate amount of a featureless greying agent. The high geometrical albedos make this assumption dubious. Also, the broad-band infrared data of Lee (1972) show dissimilarities between the satellites. The absence of distinct absorption features between 0.5μ and 1.1μ suggests the following possibilities (Adams 1968):

(a) The absence of iron-bearing minerals.

(b) Very small particles $(<10\mu)$ which reduce absorption because of reduced penetration.

(c) The presence of substantial amounts of a highly absorbing material.

In view of the high albedos for JI, JII and JIII possibility (c) is improbable for them. The weakening of the absorptions by very small particles might be present. The absence of iron-bearing minerals seems confirmed. The most conspicuous common spectral feature of the satellites is the nearly uniform drop in reflectivity from $0.5-0.3\mu$. This has not been explained. Sulfur and some of its compounds exhibit this property, but its common presence on other solar-system bodies (the Moon, Mercury, Saturn's Rings) suggest other possibilities, such as radiation damage.

In the following we will see what additional conclusions can be drawn from the albedo wavelength dependence from 0.3μ to 1.1μ . We will also make use of the broad-band infrared data in the Johnson JHKL-system (Lee, *op. cit.*)

4. Io

The strong UV absorption appears to rule out a cover by ice. Although various substances studied by Sill (Sill 1972 and Kuiper 1969) show UV absorptions, most of these are not very strong. Of the more abundant substances only sulfur and some of its compounds match the general run of the Io albedo curve. While some other compounds also show fairly strong UV absorptions, they are less probable on grounds of abundance. Now the slope of the UV sulfur absorption is dependent on the specific compounds; and, as is shown by Sill (1972), the position and slope of the ultraviolet absorption are temperature-dependent. The reflectivity of sulfur compounds increases with decreasing temperature, which might explain the high albedo of Io. Sill's reflection curves of ammonia compounds show a number of absorptions which are not seen in the albedo-wavelength dependence of Io. NH4 · HS gives rise to absorptions centered at $\sim 0.6\mu$, $\sim 1.4\mu$, ~1.7 μ and ~2.3 μ . The narrow-band data do show an absorption at $\sim 0.55\mu$ close to the 0.6 μ absorption. However, the absorptions at $\lambda > 1.1 \mu$ are of such strength and position that they would give rise to noticeable decrease in the observed albedo even in the broad-band data. The IR data of Lee do not show any depression. Thus there is no evidence for the presence of $NH_4 \cdot HS$, but S may be present.

5. Europa

The maximum albedo of JII is somewhat lower than that of Io, but still high. However, the UV absorption is considerably less strong. An absorption around 0.5μ seems to be present but is not very strong. The strength of the UV absorption of Europa is such that there exists a large range of compositional possibilities. Although the ferrous ion in most samples studied by Sill shows a UV absorption of the same order of magnitude as JII, most ferrous compounds show a low albedo. However, there exists observational evidence that Europa has polar caps (Veverka 1970). The region $\lambda > 1\mu$ must be consulted for additional clues. Lee's IR data suggest a close similarity between JII and JIII.

6. Ganymede and Callisto

The normalized albedos of these two satellites are very similar from 0.32μ to 0.92μ . With the observations having an accuracy of 2 percent, the ratio p(JIII)/p(JIV) is found constant within the expected error limits.

However, both the absolute visual and the IR albedos are considerably different. Lee finds that the IR albedo of JIII is compatible with H₂O ice. If one makes a detailed comparison between the reflection spectrum of H₂O ice at -190° C (Kuiper *et al.*), the agreement is very good. The probable presence of H₂O on Europa and Ganymede was discovered by Kuiper (1957).

Acknowledgments. The work here reported was supported by NASA Grant No. NGL-03-002-002.

APPENDIX I

The Geometrical Albedo of The Galilean Satellites

Table IV lists the values of the geometrical albedos of the Jovian satellites at $\lambda_{eff} = 0.55\mu$, as derived from both narrow- and broad-band observations. For the M_v of the sun we used Johnson's (1965) value of -27.64. The radii used are from Kuiper (1952). A change in radius of ΔR will change the albedo with a factor $(1+\Delta R/R)^2$. It is obvious from Table IV that the observational accuracy is fairly high. The main uncertainty in these numbers is still caused by the difficulty in measuring the diameter of the satellites. In determining the average value of the albedos, the measurement of Io by Lee is not taken into account. For the three other satellites all four data sets are used. The relative accuracy of the observed geometrical albedo thus derived is 3 percent (m.e.).

TABLE IV

Geometric Albedos of the Galilean Satellites at $\lambda = 0.55 \ \mu$

	1) Radius (Km.)	2) Harris	3) Leb	4) Johnson	THIS Paper	Average
Io	1625	0.97	(0.73)	0.91	0.94	0.94
Europa	1440	0.89	0.93	0.92	0.83	0.89
Ganymede	2510	0.52	0.51	0.57	0.54	0.54
Callisto	2230	0.27	0.27	0.30	0.27	0.28

¹⁾ Kuiper (1952); 2) Harris (1961); 3) Lee (1972); 4) Johnson and McCord (1970).

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No. 168 SPECTRAL ALBEDOS OF THE GALILEAN SATELLITES

by Thomas Lee

July 21, 1971

ABSTRACT

The wavelength dependence of the albedo of the Galilean satellites has been determined from wide-band photometric observations extending from 0.36μ to 3.4μ .

he author has observed Jupiter's Galilean satel-L lites with two telescopes: the 60-inch on Site II, Catalina Observatory (el. 8450 ft = 2580 m); and the 40-inch on Site IV (el. 8580 ft = 2620 m). Two standard photometers were used, one for UBVRI (with two photomultipliers, a 1P21 and a RCA 7102); and one for JHKL, using a liquid N₂-cooled PbS cell. The data are summarized in Table I. This table gives the photometric band designation, the effective wavelength, and the mean observed magnitude for each satellite. In all cases, the quoted magnitudes are means of several independent measures obtained within a few days of the 1971 opposition. Two sets of UBVRI observations were obtained on three nights and an average of 4.2 JHKL measures on six additional nights.

The scatter in the individual measures, particularly in the infrared (JHKL), is larger than found for routine stellar observations; e.g., the spread is $0^{m}2$ to $0^{m}3$ in the K magnitude. The satellites are known to vary in the visual region (Harris 1961); our observations suggest that in the infrared the variations may be somewhat larger and may be correlated with orbital phase. Because of the number of measures obtained, the average JHKL magnitudes given in Table I should be good to better than $\pm 0^{m}10$ (mean error). At shorter wavelengths, the errors will be smaller.

The observations were analyzed to determine the wavelength dependence of the albedo of these objects. The computed albedos, shown in Figure 1, are normalized to 0.5 in the visual band at 0.55μ .

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MAGNITUDES OF GALILEAN SATELLITES									
Designation Effective λ (μ)	U 0.36	В 0.44	V 0.55	R 0.70	I 0.90	J 1.25	н 1.62	K 2.2	L 3.4
JI(Io)	8.19	6.38	5.33	4.61	4.13	3.75	3.44	3.44	3.42
J II (Europa)	6.60	6.18	5.32	4.72	4.37	4.26	4.68	5.26	7.30
J III (Ganymede)	6.11	5.60	4.78	4.16	3.79	3.65	3.77	3.85	5.67
J IV (Callisto)	7.27	6.71	5.86	5.24	4.86	4,54	4.26	4.24	5.16

TABLE I

The albedos followed from: (a) the observed fluxes derived from the measured magnitudes; (b) the incident solar fluxes for each satellite; and (c) the assumption that the reflected radiation is spread uniformly over a hemisphere. The satellite radii adopted are taken from Kaula (1968); the mean heliocentric and geocentric distances of Jupiter were 5.365 and 4.372 A.U., respectively; the magnitude and colors of the sun were taken from Johnson (1964); and the photometric calibrations from Johnson (1966). The albedo thus obtained is a simplified Lamberttype, not equal either to the Bond or the geometrical albedo commonly used (Kaula 1968). The wavelength dependence shown in Figure 1 is not affected, of course. The visual albedos ($\lambda = 0.55\mu$) relative to Io are:



Fig. 1. Normalized Albedos of Galilean Satellites. The curves show the wavelength dependence of the albedo from $\lambda = 0.36\mu$ to 3.4 μ . The various photometric band positions and their half widths are shown. The curves are normalized to A = 0.50 at 0.55μ .

JII/JI	=	1.32
J III/J I	=	0.71
JIV/JI	=	0.31

These ratios should also apply to the Bond and geometric albedos.

The curves of Figure 1 show a variety of features. For Io the albedo falls off rapidly to the ultraviolet, but remains high and nearly constant from the visual to $\lambda = 3.4 \,\mu$; the high value at 3.4 μ does not favor a surface composition of NH₄Cl, but is compatible with a sulfur compound. The drop in the curves for Europa and Ganymede confirms earlier conclusions that H₂O ice is present on these satellites (Kuiper 1957), while the decrease seen at 3.4 μ for Callisto could be due to $NH_4 \cdot Cl$.

Acknowledgments — The author wishes to thank Dr. G. P. Kuiper for several interesting discussions; he also gratefully acknowledges financial support from NASA (Grant NGR 03-002-180, Suppl. #2) for this work.

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No. 169 MAP OF THE GALACTIC NUCLEUS AT 10μ

by G. H. RIEKE AND F. J. LOW

September 2, 1971

ABSTRACT

The galactic nucleus has been mapped at 10 microns with a resolution of a few arc seconds. Four discrete sources were resolved superimposed on an extended background.

We have scanned the center of our galaxy at a wavelength of 10 microns with a beam 5.5 arcsec in diameter. The scans have been combined to produce the map shown in Figure 1. At 10kpc, commonly assumed to be the distance to this region, 5.5 arcsec corresponds to 0.3 pc.

These observations were taken during the day on February 15, 1971, with the sixty-one-inch telescope at the Catalina Observatory. Because of the imprecision in the drive of the telescope, the scans could not be located accurately in declination; the telescope also drifted slowly in right ascension. These problems were overcome by comparing the scans with some made earlier by one of us (FJL) and D. E. Kleinmann with the Steward Observatory 90-inch telescope. Minor distortions may persist in the map. All of the features described below stand out clearly above the noise, which has a peak-to-peak value of about 6 x 10^{-17} W/m² Hz ster.

Four sources can be distinguished and are numbered in Figure 1. An additional broad zone of emission can be seen extending to the north. Sources 2 and 4 are partially resolved and have diameters of about 0.5 pc. There are indications on the individual scans that Source 2 would break up into two or three sources at higher resolution. Source 1 is unresolved and must be less than 0.2 pc in diameter. Source 3 may also be unresolved with our beam, although the ratio of signal to noise is inadequate to permit a strong statement.



Fig. 1 Map of the galactic center at 10 microns. The peakto-peak noise of the observations is $6 \ge 10^{-17}$ W/m² Hz ster. The contours are at levels of 1, 2, 4, 6, 8, 10, 12, 16 and 20 times this value. The brightest 10-micron sources are numbered in order of decreasing maximum flux. The beam used for these observations has a full width at half maximum of 5.5 arcsec. It is shown with the same contour levels as are used for the map. The location of the 2.2-micron "pointlike" source is indicated by a set of error bars which correspond to the uncertainty in our measurement of its position. The two most prominent field stars have been plotted and labelled A and B. The direction of the galactic plane is indicated by the dashed line.

We have also made observations at 2.2, 5 and 22 microns. A detailed description of all this work will be published after we have analyzed it more completely. We have located the "point-like" source discovered at 2.2 microns by Becklin and Neugebauer (1968) relative to our map; to within the accuracy of the measurements, its position agrees with that of Source 3. At 5 microns, scans across the complex of sources have been made at five declinations. Although these scans have a relatively low ratio of signal to noise and are insufficient in number to construct a map, they do show the same general features found on the 10-micron map.

At 22 microns, we have constructed a map with the same resolution and nearly the same ratio of signal to noise as for the 10-micron map. The noise for this map has a peak-to-peak value of about 1.8 x 10^{-16} W/m² Hz ster; at this level, Sources 2 and 3 were not detected. Source 1 stands out prominently. It lies near the center of an extended source that has a radius of about 20 arcsec.

Two field stars are plotted in Figure 1 and labeled A and B. Additional scans at 10 microns enabled us to locate our map relative to field star A. The coordinates of this star were measured from a Palomar Sky Atlas plate, allowing us to determine the absolute position of the 10-micron sources to within about 2 arcsec. (A finding chart for the galactic center has been published by Spinrad, *et al.*, (1971) (their Fig. 1). Star A is 2.5 mm south and 2.0 mm east of the indicated position of the galactic center on this chart and star B is 1.5 mm north and 1.0 mm west. It should be noted that the position of the 10-micron sources differs somewhat from the one indicated on the published finding chart.)

From our map, we estimate the total flux from the galactic center to be $480 \pm 50 \ge 10^{-26} \text{ W/m}^2 \text{ Hz}$ at 10 microns. The 22-micron map indicates that the flux at this wavelength is $2300 \pm 250 \ge 10^{-26} \text{ W/m}^2$ Hz, of which about 60% or $1400 \pm 150 \ge 10^{-26}$ W/m² Hz would fall within a field of view 25 arcsec in diameter. These results are in reasonably good agreement with earlier measurements made with a 25 arcsec field of view of $550 \pm 60 \ge 10^{-26} \text{ W/m}^2 \text{ Hz}$ at 10 microns, and $1700 \pm 200 \ge 10^{-26} \text{ W/m}^2 \text{ Hz}$ at 22 microns (Low, Kleinmann, Forbes and Aumann, 1969).

Maps of the galactic center at 2.2 microns have already been published (Becklin and Neugebauer, 1968). At this wavelength the detected flux arises almost entirely from unresolved stars. Except for the "point-like" source, the brightest features at 2.2 microns would produce a flux of 0.2×10^{-26} W/m² Hz in a field of view the size of ours. At 10 microns, the flux from these stars should be an order of magnitude lower. Our peak-to-peak noise is about 3.5 x 10^{-26} W/m² Hz; thus, the general stellar background detected at 2.2 microns makes virtually no contribution to the 10-micron map. Becklin and Neugebauer (1969) have published the results of a scan across the galactic center at 10 microns. Their scan appears to have passed near the center of Source 2.

The 10-micron sources fall within the region of maximum intensity of the extended 2.2-micron source, which is centered (Becklin and Neugebauer, 1968) at α (1950) = 17^{h} 42^m 30^s \pm 1^s and δ (1950) = -28° 59.4 \pm 0.1. Their position also agrees with that of the radio source Sgr A (Maxwell

and Taylor 1968), α (1950) = 17^h 42^m 30^s6 ± 1^s and δ (1950) = -28° 59' 14" ± 15".

Despite the consistency of these positions, significant ambiguities remain if one looks closely at the data. Our 22-micron map suggests that source 1 lies at the galactic center. On the other hand, a scan across the maximum of the 2.2-micron extended source, published by Becklin and Neugebauer (1969), places the greatest density of stars near Source 2. The 2.2-micron "point-like" object can be identified with Source 3; if it is, the 2.2-micron scan must be shifted in a direction which improves the correspondence with Source 2. The existence of Source 2 indicates that objects considerably cooler than normal stars lie very near the maximum 2.2-micron intensity. The contribution of these objects to the 2.2-micron flux is not known, leaving the position of the point of maximum stellar density uncertain.

Resolution of 6 arcsec at radio frequencies (Downes, private communication) reveals that Sgr A consists of two objects, one to the east and one to the west of our position for the infrared sources. The published position reflects the contribution of both sources.

Acknowledgments. The first measurements of the discrete sources in the galactic nucleus were made by

Low and D. E. Kleinmann (1970). These data were essential to our present efforts. We wish to thank A. A. Hoag for giving us a plate from the Palomar Atlas and E. Roemer for assistance in carrying out the measurements.

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No. 170 PRELIMINARY REPORT ON OPTICAL SEEING TESTS AT MT. LEMMON, MARCH-JUNE 1971

by George V. Coyne, S.J.

August 1, 1971

ABSTRACT

One of the Lick Observatory Polaris telescopes was used to test seeing conditions at Mt. Lemmon. The preliminary results indicate that the seeing is not unlike that at Kitt Peak. Soundings of the air flow patterns across Mt. Lemmon under winter conditions were also made by flying smoke pots on tethered balloons to elevations of 500 feet. There does not appear to be any serious local turbulence.

As part of the preliminary development of the Mt. Lemmon Observatory, an optical seeing test program was begun with one of the six-inch Polaris telescopes used by Lick Observatory under the direction of Dr. Merle Walker for the testing of various prospective observing sites. The description of this instrument is given by Harlan and Walker (1965) and the results of tests made at various California sites (Walker, 1970) and at sites in Arizona, Baja California, Chile and Australia (Walker, 1971) have been published.

The test telescope, which had been most recently used at Kitt Peak, was obtained by us in January, 1971. It was mounted by Frecker and Arthur at Mt. Lemmon on the radar drum proposed as the site of the photopolarimetric 70-inch telescope (see Figure 22 of Kuiper, 1970). The telescope was first mounted on the catwalk at the west of the drum but vibrations were too severe there and it was moved to the north edge of the drum where it is mounted on a steel I-beam structure. The objective protrudes slightly over the north edge of the drum and is at an elevation of approximately 30 feet above the ground level. There are no wind breaks about the telescope or mounting.

During the period March 3 to March 16, six exposures of 15 minutes each were made per night, two each at 21^{h} , 24^{h} and 3^{h} (MST); after March 16 three 15-minute exposures were made, one at each of the stated times.

TABL	ΕI
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Journal of Observations

Date 1971	Time (Start) MST	Fil Roll	m Exp.	Temp. °F	Rel. Hum. %	Win Vel.	d Dir.	Seeing (Sec. of Arc)	Obs.
March 3/4	21:05 21:20 00:20 00:35 03:05 03:20	1 1 1 1 1 1	1 2 3 4 5 6	34 34 33 22 32 32 32	40 40 45 45 45 45 45	0 0-5 5-7 5-7 5-7 5-7 5-7	SSW SSW SSW SW SW	<0.7 0.7 0.8 1.2 1.2 1.8	TK JG TK TK TK JG
4/5	21:26 21:43 00:05 00:20	1 1 1 1	7 8 9 10	31 31 30 30	53 53 62 62	15 15 20 20	SW SW SW SW	2.0 2.0 2.5 3.0	MH MH MH MH
5/6	21:00 21:05 21:15 23:50 00:15 02:55 03:10	1 1 1 1 1 1	11 12 13 14 16 17 18	20 20 25 25 25 25 25	55 55 25 25 18 18	5 5 10-20 10-20 2-9 2-9	NW NW NW NNW NNW NNW	6.0 6.0 5.0 5.0 1.8 2.0	JF JF JF JF JF JF JF
9/10	21:28 00:14 00:30 03:09 03:21	2 2 2 2 2	2 3 4 6 7	29 28 28 28 28 28	41 46 46 45 45	0-5 0-5 0-5 0-5 0-5	NW NW NW NW	1.3 1.5 1.6 2.0 2.0	JG JG JG JG
11/12	21:08	2	8	32	44	0-5	SSW	1.5	ΤK
12/13	21:00 21:16 00:00 00:30 03:00 03:15	2 2 2 2 2 2 2 2	9 10 11 13 14 15	35 35 34 34 34 34 34	56 56 74 74 82 82	8-10 8-10 10-15 10-15 10-15 10-15	S S S S S S	2.0 2.0 1.9 1.6 4.0 41.0-5.0	MH MH MH MH MH MH MH
16/17	20:50 21:10 21:15 23:50 00:00 02:46 03:00	3 3 3 3 3 3 3 3 3	1 2 3 4 5 6 7	40 40 38 38 38 38 38	23 23 23 24 24 24 24 24 24	8 8 8 8 8 8 8 8	S S S S S S S	1.0 0.8 1.0 1.9 1.2 1.2 1.7	JF JF JF JF JF JF JF
17/18	21:04	3	8	34	29	15,gusts	SW	5.0	МН
	00:00	3	9	32	34	20,gusts to 30	W	5.0	МН
	03:04	3	10	32	38	15,gusts to 20	พ	5.0	мн

TABLE]
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Continued

		·					1			
Date 1971		Time (Start) MST	Fi Roll	lm Exp.	Temp. °F	Rel. Hum. %	Wi Vel.	nd Dir.	Seeing (Sec. of Arc)	Obs.
March	20/21	21:35 01:05 03:05	3 3 3	111 12 13	32 31 29	25 25 25	0-5 5 0-5	W SW SW	1.5 1.3 1.6	JF JF JF
	27/28	20:50 23:45 02:50	3 3 3	14 15 16	42 39 36	56 90 72	8 8 6	W W W	1.3 1.5 1.5	JF JF JF
April	3/4	21:00 00:00 03:00	4 4 4	2 3 4	38 37 37	28 27 30	20 5 6	N N N	4.0 4.0 4.0	JF JF JF
	5/6	20:55 23:56 03:05 03:20	4 4 4 4	5 6 7 8	42 36 39 39	32 30 32 32	0 0 3 3-6	- - W W	$0.9 \\ 0.8 \\ 1.0 \\ 1.0$	ТК ТК ТК ТК
	6/7	21:10 23:45 02:45	4 4 4	9 10 11	43 42 39	38 43 38	7 10 4	SE SE SSW	1.5 1.7 1.2	JF JF JF
	7/8	23:50 02:55	4 4	13 14	37 36	27 38	5-25 10-30	SW SW	6.0 4.0	JG JG
	8/9	21:25 23:58 02:45	4 4 4	15 16 16	48 45 42	26 28 34	10-15 10-15 10-15	S₩ S S	1.5 1.6 1.5	GC GC TK
May	19/20	21:02 00:00 03:00	5 5 5	1 2 3	47 - -	14 - -	4 5-10 7-10	S S S	1.0 1.0 1.7	MH MH MH
June	3/4	21:00 00:00 03:00	5 5 5	5 6 7	36 34 36	58 76 58	8-10 10-12 10-12	₩ S S	5.0 1.8 1.5	MH MH MH
	9/10	21:00 00:00 02:50	5 5 5	8 9 10	46 42 40	52 79 70	8-10 15-20 15-20	S SE SE	2.5 1.7 3.0	MH MH MH
	29/30	21:06 00:00 03:00	5 5 5	11 13 14	55 54 54	30 32 30	0 4-6 7-10	- S S	1.5 1.0 1.0	MH MH MH

The following observers (identified by initials in Table I) have participated in this program: George Coyne, Jack Frecker, John Gradie, Morris Howes and Thomas Kunkle. The exposed film rolls were sent to Dr. Walker who developed the film, compared the star trails to his standard calibrated trails, and sent to me the results listed in column 9 of Table I. Temperature and relative humidity were read by each observer from the hygrothermographs, which are monitored weekly by Ewen Whitaker. Wind velocity and direction were measured from a handheld meter or read from a wind vane mounted at the west edge of the drum. The 25 percentile of the relative humidity by Table I is 27% (instrument reading); however, checks made with a psychrometer show that the instrument gave readings averaging about 4% too high.

Observations were made during a total of 19 nights during March-June, 1971. A journal of observations is given in Table I and the distribution of the average seeing for each night is given at the bottom of Figure 1. For each night the average seeing over all exposures was determined. However, for the nights of March 5/6, March 12/13 and June 3/4 the seeing changed by a large factor in the course of the night so that the last two exposures are treated as



Fig. 1 Distribution of the average seeing (in seconds of arc) per night at Mt. Lemmon and at Kitt Peak during the same season but for different years.

a separate night for each of these three nights. A total of 22 nights are, therefore, plotted in Figure 1. The average seeing over all nights is 2".1.

There are obviously too few observations from which to make any conclusive statements on the seeing at Mt. Lemmon. As a preliminary step, we make a comparison in Figure 1 between our Mt. Lemmon observations and those made at Kitt Peak with the same instrument for the same months, but for the two previous years. The information for Kitt Peak is from Walker (1971). At Kitt Peak the Polaris telescope was mounted on a concrete pier just west (up-wind) of the 36-inch No. 1 telescope, with its objective about 9 feet above ground level. The two years on Kitt Peak are sufficiently different to caution against general conclusions. All one can say is that the 1971 Mt. Lemmon run is not unlike the Kitt Peak runs made in 1969 and 1970.

It was feared that vibrations of the telescope by the wind might be introducing false seeing effects into the star trails, although there was no direct evidence for this, visually or photographically. In Figure 2 we have plotted the seeing for each individual exposure against the average wind at the time of that exposure. At wind velocities of 15 mph and less, a wind vibration effect cannot be very large, since both large and small seeing disks are observed. We cannot say what happens at higher velocities. In general, higher wind velocities are a sign of more turbulent conditions and poorer seeing. Therefore, we should expect some correlation between seeing and wind velocity even for a stationary telescope. For the present, the telescope appears to be adequately mounted for the average wind velocities encountered.

In addition to the seeing tests with the Polaris telescope, a sampling was made of air motion across the Mt. Lemmon summit. This was done by flying smoke pots on tethered balloons to elevations of 500 feet and observing the flow pattern of the smoke at various elevations. All flights were made from or near the same radar drum on which the Polaris telescope is mounted. The general results are:

- (1) The flow from the eastern quadrants is laminar for wind velocities up to 30 mph.
- (2) The air flowing up the steep western slope remains parallel to the ground and does not appear to be turbulent. It passes the drum on the summit, some 30 meters (70 ft.) from the west rim, at the same slope angle as the western slope itself, i.e. it does not curl over the drum.



Fig. 2 Estimated seeing (in seconds of arc) plotted versus wind velocity for each individual observation.

- (3) The air from the west above this boundary layer flows horizontally at elevations examined up to 500 feet above the drum.
- (4) By observing yellow smoke moving upslope at ground level and black smoke moving in from the west horizontally at the level of the drum, we attempted to study what happened at the interface of these two currents. The results are not altogether conclusive but there does not appear to be any serious turbulence. There were certainly no observable eddies and curls where the two currents met at the drum.
- (5) All of these tests were made under winter conditions (free air temperatures less than 60°F); warmer air may behave differently.

There is no prominent local air turbulence at the location tested.

Acknowledgment. I would like to thank Dr. Merle Walker for the use of the Polaris telescope, for the development and reduction of the film rolls, and for the use of the seeing data from Kitt Peak antecedent to publication.

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No. 171 SULFURIC ACID IN THE VENUS CLOUDS

by Godfrey T. Sill, O. Carm.

December 15, 1972

ABSTRACT

The extremely dry nature of the Venus upper atmosphere appears to demand the presence of an efficient desiccating agent as the chief constituent of the clouds of Venus. On the basis of polarization measures one would expect this substance to be present as spherical droplets, $1-2\mu$ in diameter, with a refractive index *n* of 1.46 ± 0.02 at 3500Å in the observed region of the atmosphere, with $T \sim 235^{\circ}$ K. This substance must have ultraviolet, visible and infrared reflection properties not inconsistent with the observed spectrum of Venus. Sulfuric acid, of about 86% by weight composition, roughly fulfills the first of these properties. The visible and ultraviolet transmission features of a thin layer of elemental bromine and hydrobromic acid dissolved in sulfuric acid somewhat resemble the Venus spectrum, up to 14μ . The chemical process postulated for forming sulfuric acid involves the oxidation of sulfur and its compounds to sulfuric acid through the agency of elemental bromine, produced by the photolytic decomposition of hydrogen bromide.

1. Current data on the Venus Clouds

Substantial data exist on the spectroscopic and polarimetric properties of the Venus clouds. Yet the identity of the clouds is still uncertain. Venus shows high reflection of visible light, with an albedo approaching 100%. There are strong absorption features in the ultraviolet and infrared, independent of the atmospheric absorptions caused by CO_2 . Ultraviolet photography of the planet shows regions of variable darkness, attributed to different cloud layers on the planet, revealing a dynamic system of production and destruction of a strong ultraviolet absorber, produced photolytically in the upper atmosphere. It is important to remember that the high resolution UV spectroscopy of the planet utilizes the average reflection of the whole planet in the ultraviolet. Figure 1 shows the albedo of Venus from 0.2 to 4.0μ , as taken from Kuiper's 1969 article on the clouds of Venus, with the UV region modified by Wallace *et al.* (1972), and the $3-4\mu$ region to allow for the CO₂ absorption $3.7-4.0\mu$ observed by Beer *et al.* (1972). A double absorption near 3.8μ ,



Fig. 1 Venus Bond albedos vs. wavelength, $0.2-4 \mu$.

assumed by these authors to be due to solid particles, is actually due to CO_2 gas (Kuiper 1972).

Polarization studies by Coffeen (1968) led him to conclude that the particles in the clouds were probably spherical, with a refractive index *n* of 1.43-1.60 and a mean particle size of 2.5μ . Further work by Hansen and Arking (1971) and Coffeen and Hansen (1972) give a mean particle size of 1.1μ and *n* of 1.46 \pm 0.02 at 0.35 μ decreasing to 1.43 at 1.0 μ . The particles are spherical, and show other properties of spheres, the rainbow and possibly the glory.

Another important property of the Venus atmosphere at the level of the clouds, namely at 50 mb pressure $(10^{-1.30} \text{ atm})$ and 235° K, is the extremely low water vapor mixing ratio of $10^{-6.0}$ (P_{II₂0} = $10^{-4.30}$ mb). The vapor pressure of ice is higher than this value by orders of magnitude. If H₂O is a constituent of the clouds, it must be present with a strong desiccating agent. Kuiper (1969) in suggesting FeCl₂·2H₂O as a constituent of the Venus clouds made the cogent observation that the vapor pressure of water in FeCl₂·2H₂O at 250°K was compatible with the observed vapor pressure of water. This property is not unique to FeCl₂, however. Other possible desiccating agents can fulfill this function.

Is there any one substance that can account for all the above properties of the clouds as well as being compatible with the albedo of Venus over the spectrum? FeCl₂ does a creditable job for some parts of the spectrum, but not all. FeCl₂ does not satisfy the polarimetric properties because it tends to form hexagonal platy crystals, not spheres.

The author was intrigued by the fact that Fe SO₄ hydrate, as well as other hydrated sulfates, showed low reflectivity between 3.0 to 4.0μ in the infrared. No other hydrates did as well in satisfying the low reflectivity of Venus in this region. An obvious candidate for consideration as a cloud constituent was

 H_2SO_4 , which also showed strong absorption in the 3 to 4μ region.

The above considerations led to an interim report dated 24 March 1972, sent to NASA and scientific colleagues, prepared by Dr. Kuiper, from which we quote:

"Mr. Sill found that elemental sulfur may form a cloud layer in the deeper atmosphere. In any case, sulfur-dioxide will penetrate into the upper atmosphere, where it could be oxidized to sulfuric acid by free halogens, particularly iodine; and there possibly form small cloud particles in the upper Venus atmosphere com-posed of sulfuric acid, of 88% concentration. This concentration would be consistent with the observed H₂O mixing ratio. The refractive index of this sulfuric acid solution would be about 1.45 at 240°K, in agreement with the polarization data. This material would supercool into spheres, as also indicated by the observed polarization properties of the planet. This explanation for the Venus clouds could be either an alternative to the FeCl₃·2H₂O explanation advanced in LPL Comm. No. 101 or possibly apply to the UV cloud layer only and thus be in addition to the ferrous-chloride explanation for the yellow lower clouds. In either case, the reflection spectra should be obtained next year of sulfuric acid fogs produced in the laboratory, both with and without dissolved halogens."

Another section of this report is also quoted:

"Mr. Sill found that the observed CO/CO₂ ratio of about $10^{-4.3}$ may be explained as due to an equilibrium reaction involving the weathering of iron-bearing olivines and pyroxenes by carbon dioxide:

$$\mathrm{CO}_{3} + \mathrm{Fe}_{2}\mathrm{SiO}_{4} = \mathrm{Fe}_{2}\mathrm{O}_{3} + \mathrm{SiO}_{2} + \mathrm{CO}_{4}$$
(1)

At $T_* = 750^{\circ}$ K and $P_* = 100$ atm. the CO/CO₂ ratio according to thermodynamics should indeed be about $10^{-4.3}$.

10^{-4.3} "Mr Sill found that the abundances of the acid halides (HF and HCl) may be explicable if one considers an equilibrium with a surface deposit of calcite, the latter being called upon to explain the abundance of CO₂. Carbonates are attacked by acid halides,

$$CaCO_3 + 2HX = CaX_2 + H_2O + CO_2,$$
 (2)

where X = F, Cl, Br, or I. With the ratio $H_2O/CO_2 = 10^{-6}$ found in the upper atmosphere, HCl and HF would indeed have the mixing ratios of 10^{-6} and 10^{-8} , as observed. If on the other hand, elemental abundances were solely responsible one would expect HCl to be only 5 times more abundant than HF. Therefore, it would follow that the process of eq. (2), rather than the abundance ratios, determines the relative amounts of HCl and HF. If the H_2O/CO_2 ratio in the lower atmosphere is well in excess of 10^{-6} , this could imply that the acid halides also would be more abundant there.

"If HBr and HI are used in eq. (2), then the equilibrium amounts of HBr and HI are found to be an order of magnitude more than HCl; though here elemental abundances may affect the total bromine and iodine availability. Furthermore, HBr and HI are less stable than HCl, and tend to break down upon exposure to UV (sun) light around 2500Å. "An observational program was therefore initiated

"An observational program was therefore initiated for the detection of Br_2 or HBr on Venus. The strong HBr absorption at 2.0 lies in the middle of very intense Venus CO₂ absorptions. However, Br_1 has a strong electronic band in the accessible region, 3600-5500Å, with a maximum absorption at 4100Å. This absorption causes the gas to be strongly colored visually. The band structure of Br_2 is most readily seen between 53005400Å, since the absorption is nearly continuous near its maximum. If the yellow tint of Venus were due to gaseous Br, only $\frac{1}{2}$ mm atm. would be required. This amount would just be detectable on Venus with an existing LPL spectrograph having 4A/mm dispersion, as observed between the sun's Fraunhofer lines. Spectra of Venus were obtained with this instrument on the 61-inch telescope, Catalina Observatory. No detectable Br₂ absorptions were found between 5100-5400Å. However, the Br₂ absorptions are so numerous and sharp that high resolution on Venus is called for. This problem is being pursued."

The present publication updates these preliminary results.

The first property to verify was the vapor pressure of water in equilibrium with various H_2SO_4 solutions at 235°K. The INTERNATIONAL CRIT-ICAL TABLES give values of P_{II_2O} over H_2SO_4 solutions from 10 to 95% H_2SO_4 . The appropriate values seemed to be in the high H_2SO_4 concentrations. At 235°K the following were calculated:

% (wt) H ₂ SO ₄	$\text{Log P}_{\text{H}_{2}0}$ of soln. (mb)
90	-5.03
85	4.14
80	-3.56

Since log P_{H_2O} (mb) at the cloud level is -4.3, the 85% H_2SO_4 solution is close to the observed value, with 86% solution a good fit. More exhaustive analysis of water vapor regulated by H_2SO_4 was performed by Fink, Larson, *et al.* (1972). They found H_2SO_4 of about 80% composition can dry the upper atmosphere to give good agreement with the measured abundance of H_2O .

The second property to consider is refractive index n. The INTERNATIONAL CRITICAL TABLES list refractive index for H₂SO₄ solutions (particularly 95-96% solutions) at various temperatures and wavelengths. From 301 to 387°K the refractive index decreases linearly with temperature increase. Extrapolating back to 235°K, and correcting for wavelength, n for 95-96% H₂SO₄ at 3500Å is 1.440. With the *n* value of 1.405 for $H_2SO_4 \cdot 2H_2O$ (73%) at 20°C (HANDBOOK OF CHEMISTRY AND PHYSICS), the 235° K *n* value should be 1,419. The 86% H_2SO_4 solution would therefore be calculated as n = 1.431 at 235°K. This value is close to the 1.46 \pm 0.02 of Coffeen and Hansen, but somewhat outside the limit. Solutions of H₂SO₄ > 95% are compatible with the polarization measurements. It well might be that the spherical droplets are actually frozen. This would raise the refractive index of 86% H₂SO₄ to about 1.453, well within the acceptable limits. H₂SO₄ freezes in a complex manner, with various eutectics ($H_2SO_4 \cdot H_2O$, $H_2SO_4 \cdot 2H_2O$, $H_2SO_4 \cdot 4H_2O$) freezing at various temperatures, and mixtures of them freezing at lower temperatures. Pure H_2SO_4 freezes at 10.5°C, the monohydrate at 8.6, the dihydrate at -38.9°, the tetrahydrate at -24.5°.

The author has observed in the laboratory the freezing and melting properties of an 88% H₂SO₄ solution. The frozen H₂SO₄ seemed to have an ill-defined melting point, it appeared to just become less viscous as it melted, behaving almost like a glass. A thin layer of the acid was frozen in dry ice and was perfectly transparent. The surface looked glassy. A cold metal spatula seemed to dent the surface; no fractures were observed as is common with water ice. The cloud particles of Venus could therefore be of this "glassy" type H₂SO₄, the droplets frozen into spheres, with a refractive index indicative of this denser state, namely about 1.45.

The third property of H_2SO_4 to compare with the clouds of Venus is spectral reflectivity. As of this time it has not been possible to obtain reflection spectra from fogs of H_2SO_4 . As a second best choice, transmission spectra of thin layers of H_2SO_4 might simulate to a degree the complex absorption and scattering occurring in a fog of droplets. The thin layer of H₂SO₄ was obtained by putting a drop of the acid between two plates and letting them seal together. Excess acid was usually squeezed out around the edges of the plates. It is estimated that the layer of acid is approximately 0.05 mm thick. Two plates of optical quartz, each 2 mm thick were used in the transmission spectra from 0.2 to 3.5μ . AgCl plates, ground and polished, were utilized in the spectral region 2.5 to 15.0μ . Also used in this region, with less success, were drops of H₂SO₄ sandwiched between layers of polyethylene and Saran. The absorptions of the plastics were added to H_2SO_4 , of course. This was partially cancelled in the dual beam spectrophotometer by placing two layers of plastic in the reference beam, but complete cancellation of the plastic film does not appear to be possible. Instrumental effects near the strong and sharp absorption features of the plastic film produce spurious features in the H₂SO₄ spectrum. The 0.2 to 3.5μ spectra were obtained from a Beckman DK-2A dual-beam quartz spectrophotometer, and the 2.5 to 15.0µ spectra from a Perkin-Elmer 137 NaCl spectrophotometer, both used by courtesy of Dr. B. Nagy of the University of Arizona's Department of Organic Geochemistry.



Figure 2 shows the spectral transmittance of a thin layer of 88% H₂SO₄. The two quartz plates were used with an acid layer about 0.05 mm thickness between the two plates. The reference beam was the air path. The spectrum needed only very small correction for the absorption of the two quartz plates. By comparison with the Venus spectrum, Figure 1, it can be seen that there is a good match from about 1-3 μ . Obviously the ultraviolet and visible spectral region must be explained by some other substance(s). Among possible candidates for ultraviolet absorptivity is SO₂ dissolved in cold H₂SO₄. Since S is oxidized to the 6+ oxidation state in H_2SO_4 , the intermediate oxidation state of $4 + in SO_2$ should be expected. The visible yellow color of the Venus clouds presents a problem as well. There are not too many gaseous substances which have a yellow tint. One is brown-yellow NO₂. But at the temperature of the Venus clouds the dimmer N_2O_4 is highly favored, with a consequent drastic lessening of color intensity. Furthermore, it may be extremely difficult to oxidize N in the atmosphere of Venus. A better candidate is Br₂; a red-brown liquid or gas, which if diluted and dissolved in H₂SO₄ to a moderate degree produces a yellow solution; higher concentrations give a red-brown solution. The intensity of the color of Br₂ varies with temperature. Br₂ saturated in 88%

 H_2SO_4 solution at room temperature produces a red-brown solution; when cooled to dry ice temperatures, the frozen solid is yellow. The chemical aspects of Br_2 and H_2SO_4 will be discussed below.

Figure 3 shows the inferred Bond albedo of Venus in expanded scale from 0.2 to 0.6μ . Wallace (personal communication), from his geometric albedo in the rocket UV, estimates that the Bond albedo of Venus is about 0.4 at 2500Å. Figure 4 shows the transmittance of Br2 and HBr dissolved in H_2SO_4 ; HBr was added to the Br_2 to dissolve more Br_2 than water alone would dissolve. The $HBr + Br_2$ solution was added to cold concentrated H₂SO₄ to obtain a solution of 88% H₂SO₄. Outgassing of HBr from the solution was rather vigorous. Solution Bwas more concentrated than A in Br_2 and showed more yellow color. Evidently even more coloring agent (Br_2) could be present than in solution B to match Venus $0.3-0.5\mu$; while the higher UV reflectance of Venus could be due to scattering. However, Venus has a UV peak near 2500Å, whereas the Br₂ solution peaks at 2200Å and the UV scattering would tend to displace it shortward. The Venus spectrum shows a dip at about 2900Å, whereas Br₂ solution dips at 2700Å, which would also be slightly displaced shortward by UV scattering.

The SO₂ spectrum is shown in Figure 5. In curve A, SO₂ gas was bubbled into cold 88% H₂SO₄ and a layer about 0.1 mm thick was observed through the quartz plates. The SO₂ does show a peak at 2400Å and a dip centered at 2800Å, somewhat closer to the Venus absorption and reflection. To observe the effect of a greater amount of dissolved SO₂ than could be held by the H₂SO₄ at room temperature, a cell of 88% H₂SO₄, 1 cm thick, was utilized with the SO₂ at about 1/200 of saturation. A similar cell with pure 88% H₂SO₄ was used in the reference beam of the Beckman DK-2A.



Fig. 3 Venus Bond albedo vs. wavelength, $0.2-0.6 \mu$.



Fig. 4 HBr + Br₂ dissolved in 88% H₂SO₄ (two concentrations).



Fig. 5 SO₂ dissolved in 88% H_2SO_4 (two concentrations).

When Br_2 vapor is poured onto a $SO_2-H_2SO_4$ solution, immediate bleaching of the Br_2 is observed, as well as a diminution of the absorption features of SO_2 . Depending on which constituent survives, the spectrum either resembles the SO_2 or the Br_2 spectrum more closely. Evidently the Br_2 oxidizes the SO_2 to H_2SO_4 and the Br_2 is reduced to HBr.

Figure 6 shows the infrared spectrum from 4.0 to 15.0μ of 88% H₂SO₄ (solid line) with the left hand ordinate of transmittance. The sample was one drop of acid squeezed between AgCl plates, with similar plates in the reference beam. The dashed spectrum of Venus is that of Gillett, Low, and Stein (1968) with the right-hand ordinate of log power at the detector. The dotted lines are, first, the reflection of sunlight from a body of unit albedo, and second, the black-body radiation curve at 225°K. If there is H_2SO_4 in the clouds cooler than the emitting surface, then indeed H₂SO₄ should show in absorption. There are some similarities between the absorptions of the Venus spectrum 7-12 μ and the absorptions of 88% H₂SO₄, especially in the absorption feature centered at 11.2μ , prominent in Hanel's Venus spectrum (1968).

2. Chemical Production of the H₂SO₄ Clouds

The mode of production of H_2SO_4 in the Venus atmosphere poses a complex problem. First of all no S compounds have been detected on Venus as gases. Upper limits to the mixing ratio of S compounds are available, and the values are (Kuiper 1969): $SO_2 < 10^{-7.5}$, $COS < 10^{-8}$, $H_2S < 10^{-3.7}$. Two of these compounds can probably be ruled out because of their instability to ultraviolet radiation in the upper atmosphere of Venus. H₂S was exposed to the UV radiation of a quartz mercury penlight (most effective radiation at the Hg 2536.5Å). Within minutes, noticeable deposition of S particles on the walls of the flask was observed as well as a fine smoke of S in the gas. Within an hour decomposition was virtually complete. COS is almost as unstable to UV as H_2S , decomposing into CO and S. This decomposition also approaches 100% in hours. Since the upper Venus atmosphere appears transparent to $\lambda > 2000$ Å, H₂S and COS can be ruled out. This leaves one with SO₂ and S as the most likely carriers of S to the upper atmosphere.

 SO_2 has S in the 4+ oxidation state; in order to produce H_2SO_4 it is necessary to oxidize S to the 6+ state, as found in H_2SO_4 , or in the anhydride of H_2SO_4 , SO_3 . (The SO₃ would quickly absorb water

and form H_2SO_4). Is it possible to produce SO_3 in the Venus atmosphere? The process would be:

Fig. 6 Solid line: Transmission spectrum with 88% H₂SO₄ solution (left ordinates). Dashed line: Infrared

spectrum of Venus with black body and unit albedo refer-

ences (right ordinates).

$$SO_2 + CO_2 \equiv CO + SO_3. \tag{1}$$

The resulting log K values are -32.64, -24.43, -16.40, and -13.14 for the temperatures of 298°, 400°, 600°, 750° K, respectively. The thermodynamic data for calculating the equilibrium constant, K, are from Robie and Waldbaum (1968). With the CO₂/CO ratio of 10^{4.34} of Venus, presumed to be constant throughout the atmosphere, it is possible to calculate the SO₂/SO₃ ratio for the lower atmosphere (T>400°K) where thermodynamics can furnish valid information The ratios for 298°K are included simply to show the trend:

Temp. °K 298 400 600 750
$$SO_2/SO_3$$
 ratio $10^{28.32}$ $10^{25.09}$ $10^{12.06}$ $10^{8.80}$

The ratio is *unfavorable* for SO₃ production *at any level*, let alone the upper atmosphere. In the lower hotter atmosphere it appears that the most abundant S species would be COS, SO₂, and S₂, in that order. It is impossible to give accurate ratios of the S compounds without knowing what the total S abundance is, and whether or not the CO₂/CO ratio remains constant. One possible surface reaction which might determine the abundance of atmospheric SO₂ involves the equilibrium decomposition of anhydrite, CaSO₄, in a CO₂-CO atmosphere:

$$CaSO_4 + CO = CaCO_3 + SO_2.$$
 (2)

At a surface temperature of 750° K, thermodynamics predicts a CO/SO₂ ratio of 10. The CaSO₄-CaCO₃ equilibrium was chosen because CaCO₃ is often considered as the major weathered mineral in a CO₂



atmosphere. Furthermore CaSO₄ is often found as a principal S compound in geothermal regions, indicating its ability to survive under moderately high temperature conditions. If the CO/SO₂ ratio of 10 should survive to the upper atmosphere there would be an abundant supply of SO₂ for forming H₂SO₄, provided there is a mechanism for further oxidizing the S to SO₃-H₂SO₄.

The oxidizing agent in the earth's atmosphere for converting SO₂ to SO₃ is O₂. On Venus this is ruled out, and the most likely agents are molecular halogens, namely F₂, Cl₂, Br₂ and I₂. F₂ can be rejected because of the extreme chemical stability of HF. Likewise Cl₂ gas must be rejected because HCl does not photolytically dissociate for $\lambda > 2000$ Å. The spectra of HCl solutions show no notable absorptions in the ultraviolet between 2000-4000Å. HBr and HI are strong absorbers. Both gases are readily dissociated into hydrogen gas and the molecular halogen under Hg vapor irradiation, and a few hour's radiation brings about equilibrium. In the case of HBr:

$$2HBr \stackrel{\text{lif}}{=} H_2 + Br_2, \tag{3}$$

equilibrium quantities of the gases indicate that 7% of HBr is decomposed in the gas-gas equilibrium. The aqueous solution of HBr also shows a noticeable darkening due to formation of Br₂ under UV exposure.

Mention was made of the bleaching effect of SO_2 gas dissolved in H_2SO_4 when Br_2 fumes were poured onto the surface of the acid- SO_2 solution. This apparently shows the ease at which Br_2 will oxidize SO_2 to H_2SO_4 in an acid medium. The thermodynamics also indicate a favorable reaction (Latimer 1952):

$$Br_{2}(liq) + H_{2}SO_{3}(aq) + H_{2}O = 2Br^{-}(aq) + 4H^{+}(aq) + SO_{4}^{2-}(aq) E_{o} = +0.89 \text{ volts.}$$
(4)

With a positive potential of 0.89 v, the reaction should and does proceed rapidly at room temperature. The two reactions can be summarized as:

$$2HBr = H_2 + Br_2$$
(5)
Br_2 + SO_2 + 2H_2O = H_2SO_4 + 2HBr.

The HBr is reconstituted in the process, and in effect serves as a carrier catalyst for the reaction. The overall reaction therefore involves destruction of H_2O into $H_2(g)$, which presumably escapes from Venus, and oxygen incorporated into H_2SO_4 :

$$SO_2 + 2H_2O = H_2SO_4 + H_2.$$
 (6)

Reaction (6) is not favored thermodynamically and could never occur at the low temperature of the Venus clouds, but the intermediate reaction with Br₂ (and presumably the other halogen, I_2) make the process possible. None of the intermediate substances has been identified in the atmosphere of Venus, namely SO₂, HBr and Br₂. HBr is a difficult problem spectroscopically because its absorption bands at 2 and 4 microns fall into the heavy CO₂ bands. Br₂ (gas) was examined at our laboratory with the helpful assistance of Allen Thomson and Thomas N. Gautier. Moderately strong absorption bands were found in the region of 5364Å. Sill, Gautier, and Kuiper attempted to detect gaseous Br₂ in the atmosphere of Venus first with the 61-inch telescope of the Catalina Observatory, then with the 107-inch telescope of the McDonald Observatory. Edwin S. Barker of the University of Texas graciously aided us in obtaining high resolution spectra at the Coudé focus of the 107-inch telescope with its attendant echelle grating spectrograph. Spectra of Venus, the sun, and the sun + Br₂ (gas) were obtained in the region of 5364Å. The results were negative. From the spectra it is estimated that the smallest quantity of Br₂ that could be seen in Venus is about 0.13 mmatm. An upper limit to the Venus mixing ratio $Br_2/$ CO_2 can be set at $< 10^{-7.3}$. This implies a very small quantity of Br₂ vapor, much less than the 4 mb vapor pressure of elemental Br₂ at 235°K. On the other hand Br₂ is very soluble in cold H₂SO₄, especially in the presence of HBr.

Both Br_2 and SO_2 have mixing ratios with upper limits $< 10^{-7}$. This may simply be a sign of their high reactivity with each other. The reaction is fast and involves reconstituting HBr. Since no upper limit on HBr has been established, the whole postulated mechanism cannot be proven at this time. The empirical data discussed above indicate that the visible and UV albedos of Venus may be consistent with Br_2 dissolved in H_2SO_4 . The highly sensitive test for gaseous Br_2 is not applicable to dissolved Br_2 as the sharp lines of the gas are smeared into the broad band of a liquid.

Chemical considerations would lead one to suspect that HBr is a gaseous constituent of the atmosphere of Venus, chiefly because HF and HCl are present. In general one halogen implies the presence of the others, taking due account of their chemical activity. A perfect case in point is the relative abundance of HCl and HF in the Venus atmosphere, with mixing ratios of 10^{-6.0} and 10^{-8.0} respectively. The Cl/F ratio in the atmosphere is 100, whereas the cosmic elemental abundance ratio of Cl/F is 5.5 (Suess-Urey) or 1.8 (Cameron) according to the HANDBOOK OF PHYSICAL CONSTANTS. Something is making F less abundant in the atmosphere. Sea water on earth shows a similar disparity with a Cl/F ratio of 7000. In this latter case the explanation is due to the higher solubility of chlorides vs fluorides, plus the relative stability of fluoride micaceous minerals versus the chloride. As far as Venus is concerned, the relative solubilities of chlorides and fluorides in water have no bearing on the case, but rather the chemical stability of chloride and fluoride minerals in the presence of a hot CO_2 atmosphere. Pertinent reactions might be:

$$CaCl_2 + CO_2 + H_2O = CaCO_3 + 2HCl$$

$$CaF_2 + CO_3 + H_3O = CaCO_3 + 2HF.$$
(7)

At 750°K, with P_{CO_2} at 100 atm, the relative amounts of gaseous hydrogen halides, HCl/HF, is about 100, which is the observed ratio in the upper atmosphere. The simple cosmic abundance ratio is not valid. The same argument can be applied to the equivalent bromide:

$$CaBr_2 + CO_2 + H_2O = CaCO_3 + 2HBr. \quad (8)$$

The ratio HCl/HBr is less than 1. It is possible that HBr might be more abundant than HCl in the Venus atmosphere if the surface reactions are the dominant influence in atmospheric abundance. This may be the case, otherwise HCl should be more abundant than the observed mixing ratio of $10^{-6.0}$. Cl in the carth's oceans amounts to 8.6×10^{22} moles, whereas in Venus' atmosphere there is about 1.3 \times 10^{16} moles. Unless severe depletion of Cl occurred on Venus vs earth, the excess Cl on Venus should be locked in the surface rocks in a way similar to equation (7). When one considers Br on Venus, there might appear to be a problem of the absolute abundance of Br. The Cl/Br ratio in the carth's oceans is 660, which ratio is the one assumed by Suess-Urey as the relative cosmic abundance. If the absolute abundance of Br on Venus matches that in the earth, there is more than enough Br to make the equilibrium of equation (8) applicable, namely HBr would be more abundant in the atmosphere than HCl. Similar arguments can be applied to I, and HI would be thermally unstable and decompose.

3. Forming H₂SO₃ Synthetically in the Laboratory

Tests were conducted in the laboratory to see if H₂SO₄ could be formed by the irradition of HBr and S compounds. Acid solutions of HBr and various S gases, particularly SO₂ and COS were irradiated with the quartz Hg lamp. COS vigorously decomposed, producing large quantities of S. Some H₂SO₄ was produced, which was analyzed by precipitating BaSO₄ on the addition of BaCl₂ solution. The conversion of COS to H₂SO₄ was about 3%. SO₂ showed slightly more conversion to H_2SO_4 . In one day's irradiation about 5% converted. Without HBr, conversion to H₂SO₄ by SO₂ was negligible. The amount of BaSO4 precipitated was minute. Even elemental Br₂ and S, mixed together in water solution showed a high conversion to H₂SO₄, but since this reaction occurred in the open air undoubtedy dissolved O₂ facilitated the reaction. Therefore, H₂SO₄ is readily produced under postulated Venus atmospheric conditions.

4. Conclusion

The lack of detectable S compounds in the atmosphere of Venus has always been a problem. Whereas the atmosphere of Earth is dominated by O_2 and H_2O , the Venus atmosphere could be dominated by the halogens. The lack of S is explained by postulating that the clouds of Venus are frozen droplets of 86% H₂SO₄, formed by the oxidation of SO₂ by elemental Br₂ in the upper atmosphere of Venus.

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