

1 **Cloudy sounding and cloud-top height retrieval from AIRS alone single**  
2 **field-of-view radiance measurements**

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24  
25 **Abstract**

26 High-spectral resolution measurements from the Atmospheric Infrared Sounder (AIRS)  
27 onboard the EOS (Earth Observing System) Aqua satellite provide unique information  
28 about atmospheric state, surface and cloud properties. This paper presents an AIRS alone  
29 single field-of-view (SFOV) retrieval algorithm to simultaneously retrieve temperature,  
30 humidity and ozone profiles under all weather conditions, as well as cloud top pressure  
31 (CTP) and cloud optical thickness (COT) under cloudy skies. For optically thick cloud  
32 conditions the above-cloud soundings are derived, whereas for clear skies and optically  
33 thin cloud conditions the profiles are retrieved from 0.005 hPa down to the earth's  
34 surface. Initial validation has been conducted by using the operational MODIS (Moderate

35 Resolution Imaging Spectroradiometer) product, ECMWF (European Center of Medium-  
36 range Weather Forecasts) analysis fields and radiosonde observations (RAOBs). These  
37 inter-comparisons clearly demonstrate the potential of this algorithm to process data from  
38 high-spectral infrared (IR) sounder instruments.

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58 **1. Introduction**

59 The AIRS instrument measures radiances in 2378 spectral channels within the spectral  
60 range from  $650\text{ cm}^{-1}$  to  $2675\text{ cm}^{-1}$  (corresponding to  $3.74\text{ }\mu\text{m}$  to  $15.4\text{ }\mu\text{m}$ ). The spectral  
61 coverage includes strong  $\text{CO}_2$  absorption necessary for temperature profile retrievals,  
62 window regions that are used for retrieving the surface and cloud properties, and a strong  
63 water vapor absorption band for humidity soundings. The maximum scanning angle of  
64 AIRS is 49.5 degrees, the swath width is 1650 km, and the footprint size is 13.5 km at  
65 nadir. More specifications about the AIRS instrument can be found elsewhere [e.g.,  
66 *Aumann et al.*, 2003; *Chahine et al.*, 2006]. Since one footprint (due to its size) contains  
67 clear and/or cloudy scenes with varying properties (e.g., fraction, height, phase),  
68 sounding retrievals using AIRS-only measurements under all sky conditions are quite  
69 challenging. According to *Smith et al.* [2005] there are essentially three ways to deal with  
70 cloudy radiances: (1) assuming opaque cloud conditions, (2) cloud clearing, and (3)  
71 making use of a physically based radiative transfer model. In approach (1) the sounding  
72 retrievals can be derived down to the cloud top level. Cloud clearing combines the cloudy  
73 radiances with clear measurements from another instrument; for example, AMSU  
74 (Advanced Microwave Sounding Unit) radiances are used in the AIRS operational  
75 retrieval system [*Susskind et al.*, 2003], and MODIS measurements are used for AIRS  
76 single FOV cloud clearing [*Smith et al.*, 2004, *Li et al.*, 2005b]. Approach (3) was first  
77 applied on AIRS measurements in *Smith et al.* [2005]. Using measurements from the  
78 aircraft sounder NPOESS (National Polar-orbiting Operational Environmental Satellite  
79 System) Airborne Sounder Testbed – Interferometer (NAST-I) with high spatial  
80 resolution (2 km at nadir) offers the advantage of cloudy FOVs with less varying cloud

81 height and optical properties. Statistical and physical inversion methods using a cloudy  
82 radiative transfer model have been developed to process NAST-I radiances for accurate  
83 retrieval of temperature and moisture profiles below optically thin clouds [Smith et al.,  
84 2005a; Zhou et al., 2005; 2007a; 2007b]. The algorithm for NAST-I cloudy sounding has  
85 been adjusted to be suitable for AIRS footprint in this paper. The main differences  
86 between the NAST-I method and AIRS cloudy sounding algorithm, as presented in this  
87 paper, are that (1) the latter uses a different training set (global instead of regional and  
88 seasonal), (2) assigns ice cloud-top and/or water cloud-top to each profile, the cloud-top  
89 assignment is a little different from that in NASTI algorithm which assumes two cloud  
90 levels and alters the profile to be isothermal below the lower cloud level, (3) performs  
91 different classification procedures in retrieval, (4) uses an IR technique to obtain the  
92 cloud phase in the retrieval step, and (5) uses MODIS product for independent  
93 comparisons. Initial results of inter-comparisons with the operational MODIS cloud-top  
94 pressure (CTP) product, ECMWF analysis and radiosonde observations are promising for  
95 the processing of data from future advanced IR sounder instruments like IASI (Infrared  
96 Atmospheric Sounding Interferometer) and CrIS (Cross-track Infrared Sounder).

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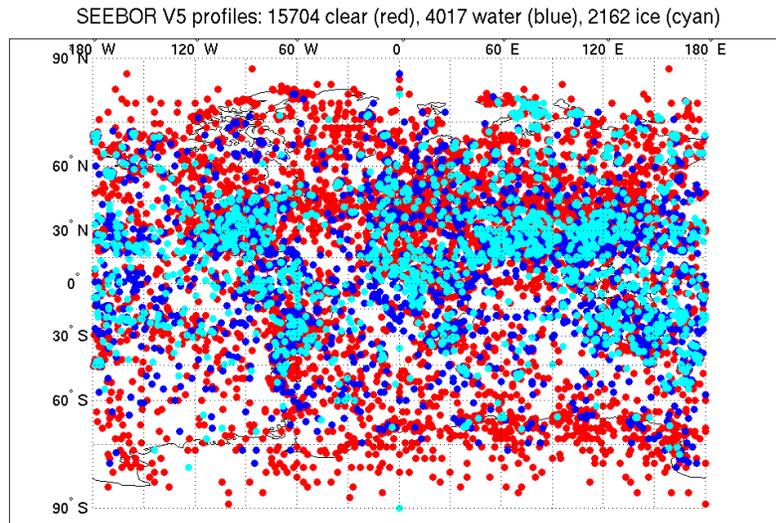
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## 108 **2. Methodology used for AIRS alone SFOV cloudy sounding**

109 The IMAPP AIRS retrieval software (latest version v1.3 was released in November  
110 2006), which delivers atmospheric and surface parameters with a validity restricted to  
111 clear skies, was used as the starting point for the cloudy retrieval methodology. The clear  
112 sky algorithm is based on eigenvector regression. The regression training set [*Borbas et*  
113 *al.*, 2005] consists of approximately 15000 globally distributed profiles (including surface  
114 parameters) and their associated **computed** radiances, which were generated by the Stand-  
115 alone Radiative Transfer Algorithm (SARTA v106, *Strow et al.*, 2003). The training set  
116 is classified based on the brightness temperature (BT) in the longwave window region  
117 and the AIRS viewing angles. In addition to the simulated IR radiances, the surface  
118 pressure (**extracted from analysis data provided by the National Centers of Environmental**  
119 **Prediction, NCEP**) and solar zenith angle are also used as predictors. The sounding  
120 retrieval product, obtained at AIRS single FOV resolution, includes temperature,  
121 humidity and ozone profiles, as well as surface skin temperature and surface emissivities.  
122 The surface IR emissivities are retrieved at 10 IR wavenumber points. Detailed  
123 information about the IMAPP AIRS retrieval algorithm under clear skies can be found in  
124 *Weisz et al.* [2003; 2006].

125 From the **regression training set** ~6200 profiles can be assigned with a CTP between  
126 900 and 200 hPa according to their relative humidity (RH). Out of these ~2160 profiles

127 are assumed to be ice clouds (those with CTP < 450 hPa) whereas ~4010 profiles are  
128 assumed to be water clouds (CTP > 400 hPa). Figure 1 displays the locations of the clear  
129 profiles, as well as profiles with water clouds and ice clouds.



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131 Figure 1. Global distribution of training set profiles. Clear sky pixels, water cloud and ice  
132 cloud pixels are indicated as red, blue and cyan dots, respectively.

133 Cloud optical thickness (COT) values between 0.001 and 2 were assigned randomly  
134 to each profile in both classes (water and ice). For ice clouds the effective particle radius  
135 ( $R_e$ ) was computed inserting COT into an equation that is given in *Heymsfield et al.*  
136 [2003]. With a random error of 10% added to  $R_e$ , a range between 10 and 30  $\mu\text{m}$  was  
137 obtained. For water clouds an effective particle radius distributed between 5 to 25  $\mu\text{m}$   
138 was randomly assigned to the profiles. The assignment of COT and  $R_e$  to the a profile is  
139 similar to *Zhou et al.* [2005], although the CTP assignment is different.

140 Through joint efforts of the University of Wisconsin-Madison and Texas A&M  
141 University, a fast radiative transfer cloud model for hyperspectral IR sounder  
142 measurements has been developed [*Wei et al.*, 2004]. For ice clouds, the bulk single-

143 scattering properties of ice crystals are derived by assuming aggregates for large particles  
144 ( $>300\ \mu\text{m}$ ), hexagonal geometries for moderate particles ( $50 - 300\ \mu\text{m}$ ) and droxtals for  
145 small particles ( $0 - 50\ \mu\text{m}$ ). For water clouds, spherical water droplets are assumed, and  
146 the classical Lorenz-Mie theory is used to compute their single-scattering properties. In  
147 the model input, the cloud optical thickness is specified in terms of its visible optical  
148 thickness at  $0.55\ \mu\text{m}$ . The IR COT for each AIRS channel can be derived from the  
149 visible COT. The cloudy radiance for a given AIRS channel can be computed by  
150 coupling the clear sky optical thickness and the cloud optical effects. The clear sky  
151 optical thickness is derived from the fast radiative transfer model SARTA. Once the  
152 cloudy radiances have been calculated, a regression is performed to output two sets of  
153 regression coefficients (water and ice). In addition to this classification based on the  
154 cloud phase, the viewing angle classification is also applied in the cloudy retrieval  
155 process.

156 A cloud phase detection method based on an IR technique [*Strabala et al., 1994*] is  
157 applied to the AIRS BT spectrum for identifying clear, ice clouds, water clouds and  
158 mixed phase clouds for a given AIRS footprint. If the pixel is clear, then the clear  
159 regression coefficients are applied to the AIRS BT spectrum, and the retrieval is  
160 performed as in version 1.3 of the IMAPP AIRS retrieval algorithm. One improvement to  
161 the clear sky algorithm involves using emissivity eigenvectors in the regression [*Zhou et*  
162 *al., 2001; Smith et al., 2005b*], and a hyperspectral emissivity spectrum is simultaneously  
163 obtained along with the sounding products; a manuscript on handling hyperspectral  
164 emissivities in sounding retrieval is under preparation.

165        If the pixel is cloudy, then the appropriate set of coefficients is used depending on  
166 the cloud phase. If the cloud phase is mixed, then the clouds are treated as ice clouds. If  
167 the retrieved cloud optical thickness is less than 1.5 (i.e., optically thin clouds), the  
168 sounding parameters are output from the top of the atmosphere down to the surface. In all  
169 other cloud cases (i.e., optically thick clouds), the sounding parameters are retrieved  
170 down to the cloud top pressure (CTP) level. In addition to temperature, humidity and  
171 ozone, COT and CTP are retrieved for every cloudy pixel in a granule.

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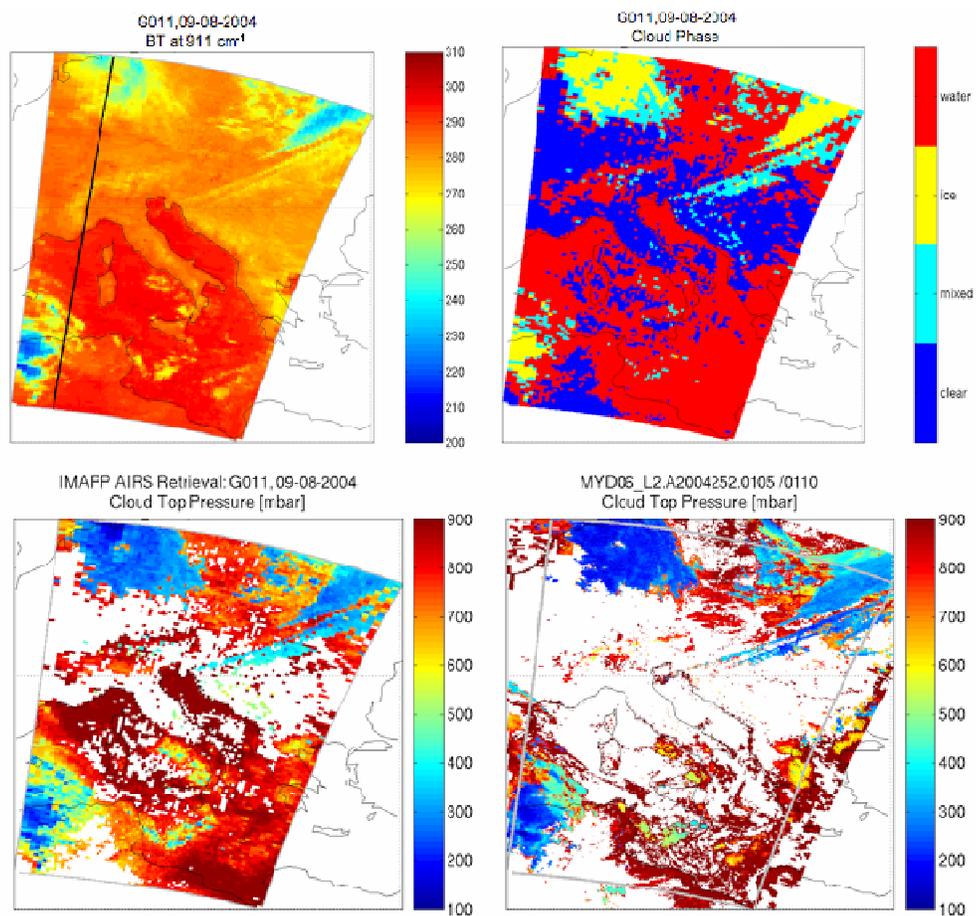
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### 192 3. Results and preliminary validation

193 Nighttime granule 11 on 08 September 2004 was chosen to illustrate the retrieval results.

194 Figure 2 shows the AIRS BT at wavenumber 911  $\text{cm}^{-1}$  (top left) and the **retrieved** cloud

195 phase (top right) for this granule.



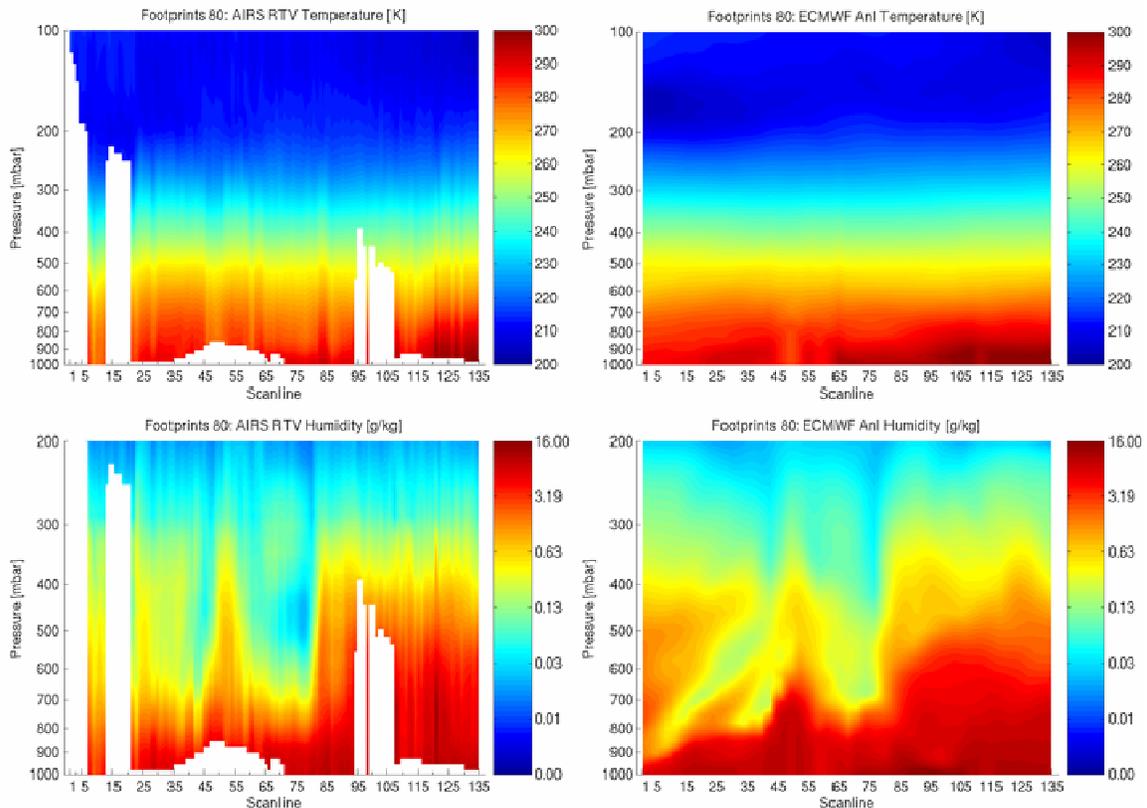
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197 Figure 2. AIRS BT at wavenumber  $911\text{ cm}^{-1}$  (top left), retrieved cloud phase (top right),  
198 AIRS retrieved CTP (bottom left) and operational MODIS CTP (MYD06) product  
199 (bottom right) for granule 11 on 08 September 2004.

200 The bottom panels of Figure 2 display the CTP retrieved by AIRS (left) and by  
201 MODIS (right). It should be mentioned that the operational MODIS (MYD06) CTP  
202 product uses sounding profiles from global forecasts, whereas the CTP from AIRS is  
203 simultaneously retrieved with the sounding (temperature, moisture and ozone) profiles.  
204 The CTP retrievals from AIRS agree very well with the operational MODIS CTP  
205 product. Specifically, over the ocean (Mediterranean Sea) the AIRS single FOV  
206 algorithm is capable of retrieving very reasonable values for lower clouds, whereas  
207 MODIS provides no retrievals due to limited spectral information. The circular feature of  
208 the thick cloud in the upper left corner of the granule depicts different cloud heights of  
209 mixed and ice clouds as seen in the cloud phase panel of Figure 2.

210 To assess performance of the sounding retrieval algorithm under cloudy conditions a  
211 cross section from north to south (as indicated in the BT panel of Figure 2 as a solid black  
212 line) is examined and evaluated by comparing with the ECMWF model analysis (see  
213 Figure 3). The ECMWF analysis data has been interpolated horizontally to the AIRS  
214 pixels and vertically to 101 pressure levels that are used in the AIRS radiative transfer  
215 calculation. The difference in time between the ECMWF analysis and the AIRS  
216 measurements is about 70 minutes. The spatial resolution of the ECMWF analysis is 0.5  
217 degrees. As mentioned above, the parameters are only retrieved to the CTP level when  
218 optically thick clouds are present. For temperature (top panels of Figure 3) some minor  
219 differences can be seen; for example, ECMWF analysis finds colder temperatures

220 between ~170 and 200 hPa for scanline 1 to 45. Nevertheless, the general pattern of the  
221 AIRS retrieved temperature field compares favorably with the ECMWF model.  
222 Furthermore, relatively accurate temperature soundings are obtained under thin clouds as  
223 can be seen in the area beyond scanline 105.



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225 Figure 3. Temperature (top) and humidity (bottom) from the AIRS cloudy sounding  
226 retrieval (left) and the ECMWF model analysis (right) for granule 11 on 08 September  
227 2004.

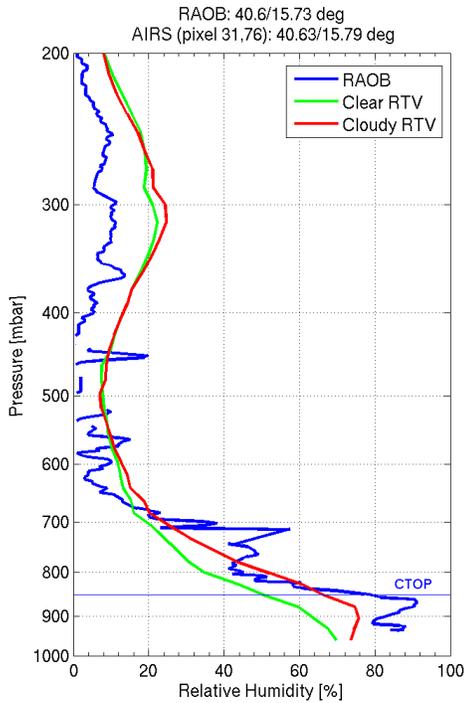
228 The same inter-comparison between AIRS and ECMWF was conducted for humidity  
229 (bottom panels of Figure 3). The AIRS sounding retrieval successfully reproduces the  
230 humidity variation shown in the ECMWF model analysis. Again in areas of thin or  
231 broken clouds (e.g., between scanlines 7 and 13, and between 105 and 125) the cloudy

232 retrieval achieves very reasonable values, which a clear sky sounding retrieval algorithm  
233 would not be able to accomplish.

234       Root-mean-square errors of the retrieval deviations from the ECMWF fields (not  
235 shown) offer further reassurance that the cloudy sounding algorithm is reasonably  
236 accurate beneath broken and optically thin clouds. Differences between ECWMF and  
237 AIRS are partly caused by the different spatial (horizontally and vertically) resolution and  
238 by the difference in time.

239       It is also worth noting that atmospheric profile retrievals (including ozone) of the  
240 layers above the cloud are not affected by the clouds below. That is particularly evident  
241 for ozone profiles (not shown), where a clear sky only method yields unrealistic  
242 stratospheric ozone values caused by clouds from lower levels. These disturbances are  
243 not seen when using the cloudy sounding algorithm.

244       To further investigate the performance of the sounding retrieval, a co-located  
245 radiosonde measurement within granule 11 (southern Italy) was used. The retrieved CTP  
246 is 849.8 hPa. For this particular thin cloudy case the temperature profile (not shown) is  
247 not significantly affected by clouds when compared with that from the clear sky method.  
248 However, for water vapor the improvement is significant when applying the cloudy  
249 sounding algorithm. This is illustrated in Figure 4 for relative humidity. The cloudy  
250 retrieval captures the atmospheric moisture variation as shown by the radiosonde very  
251 well, in particular below the cloud top level where improvements in relative humidity  
252 values larger than 10 % can be achieved.



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254 Figure 4. AIRS retrieved sounding profiles (green and red lines refer to the results from  
 255 the clear sky retrieval and the cloudy retrieval method, respectively) for relative humidity  
 256 (in percentage from 0 – 100%) compared with one co-located radiosonde measurement  
 257 (blue). The retrieved cloud top pressure is indicated as a thin blue line.

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#### 267 4. Summary

268 The probability of having clouds in an AIRS footprint, which is 13.5 km at nadir, is  
269 relatively high. An approach to retrieve sounding parameters along with cloud top  
270 pressure (CTP) and cloud optical thickness (COT) under cloudy skies is described in this  
271 paper. A fast cloudy radiative transfer model accounting for clouds of various phases,  
272 cloud particle sizes, and optical thicknesses was used to simulate cloudy radiances  
273 representing the regression training set. The simulations are performed for a subset of  
274 profiles from the global database with COT and effective particle size randomly assigned  
275 to each profile. An eigenvector regression retrieval method is applied to obtain two sets  
276 of regression coefficients (one for water clouds and one for ice clouds). The retrieval  
277 product includes temperature, humidity and ozone from 0.005 to either the surface for  
278 clear skies, and cloudy skies with broken and/or optically thin clouds or to the cloud top  
279 level when optically thick clouds are present. AIRS retrieved CTP agrees very well with  
280 the operational MODIS (MYD06) product. As for preliminary validation of the sounding  
281 products, ECMWF analysis fields were used. The spatial features are well reproduced by  
282 the AIRS cloudy sounding profiles. The case study involving a co-located radiosonde  
283 measurement further endorses the capability and accuracy of this algorithm.

284 Future work includes more validation of the current product, sounding improvement  
285 for mixed phase clouds, and sounding enhancement by using an iterative physical  
286 retrieval scheme **at AIRS SFOV resolution**. Studies have been conducted on the  
287 combination of MODIS and AIRS for cloud property retrieval [Li et al. 2004, 2005a],  
288 and cloud clearing [Li et al., 2005b]; a direct sounding approach using MODIS and AIRS  
289 will also be investigated. The long-term goal is to apply this methodology to new

290 instruments like IASI and CrIS, and to support the development of other high-spectral IR  
291 sounders.

292

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