

1 **Title**

2 Using Schumann Resonance measurements for constraining the water abundance on the giant  
3 planets – implications for the Solar System formation

4  
5 **Running Title**

6 Schumann resonances in the giant planets

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21  
22  
23 **Abstract**

24 The formation and evolution of the Solar System is closely related to the abundance of volatiles,  
25 namely water, ammonia, and methane in the protoplanetary disk. Accurate measurement of  
26 volatiles in the Solar System is therefore important to understand not only the nebular hypothesis  
27 and origin of life but also planetary cosmogony as a whole. In this work, we propose a new,  
28 remote sensing technique to infer the outer planets water content by measuring Tremendously  
29 and Extremely Low Frequency (TLF-ELF) electromagnetic wave characteristics (Schumann  
30 resonances) excited by lightning in their gaseous envelopes. Schumann resonance detection can  
31 be potentially used for constraining the uncertainty of volatiles of the giant planets, mainly  
32 Uranus and Neptune, because such TLF-ELF wave signatures are closely related to the electric  
33 conductivity profile and water content.

34

35 **Keywords**

36 Planets and satellites: composition; planets and satellites: formation; planets and satellites:  
37 physical evolution; protoplanetary disks; space vehicles: instruments; waves.

38 The authors suggest including ‘Schumann resonance’ in the keywords for indexing purposes.

39

40 **1. INTRODUCTION**

41 The nebular hypothesis is the prevailing model to explain the formation and evolution of the  
42 Solar System; specifically, the Solar Nebular Model receives most attention and, to some extent,  
43 is able to explain several characteristics of the Solar System planets, namely distribution and  
44 migration, and the composition of the initial protoplanetary disk and subsequent accretion  
45 processes (e.g., Tsiganis et al. 2005). According to theory, the accretion processes induce  
46 formation of silicates and grains of ice and dust that eventually coagulate in small planetesimals  
47 and planetary embryos. Detailed analyses of these processes are not the aim of this work;  
48 thorough, comprehensive descriptions can be found elsewhere (e.g., Benz et al. 2000;  
49 Kallenbach et al. 2003). Nonetheless, it is important to mention that the water vapor and ice  
50 contents in the gaseous giants, and consequently in the protoplanetary disk volatile inventory,  
51 remain largely unknown. Measurements of the water content in the atmosphere of Jupiter and  
52 Saturn have been made by various spacecraft (Mahaffy et al. 2000; Baines et al. 2009), but  
53 generalization to the entire fluid envelope of the two planets is not possible. Because of limited  
54 in situ measurements, even the accuracy of the Jovian planets (Jupiter and Saturn) aeronomy  
55 models cannot be validated, and the water content in Uranus and Neptune is still more uncertain.  
56 Since most volatiles in the core of the primordial solar nebula are dissociated or diffused toward  
57 the outer regions during the accretion process due to temperature increase, accurate estimates of  
58 water in the giant planets and beyond would be valuable to assess volatile inventories in the  
59 protoplanetary disk. For example, bombardment of the inner planets by comets and asteroids  
60 originating from the outer Solar System delivers water and other volatiles; terrestrial water  
61 originates, substantially, from cometary bombardment, possibly including the building blocks of  
62 life (e.g., Encrenaz 2008; Cooper et al. 2001).

63 Constraining initial parameterization of the protoplanetary disk is important for a better  
64 understanding of the Solar System origin and evolution. For example, the distribution of rocky,

65 icy, and gaseous bodies resulting from the protosolar nebula is linked to volatiles abundance and  
66 to the location of the “snow line”. The snow line, also known as ice or frost line, establishes the  
67 boundary in the protoplanetary disk beyond which hydrogenated molecules, namely water,  
68 methane, and ammonia, were cool enough (~150 K) to condense and form ice grains. During the  
69 early stages of the Solar System formation, the snow line was presumably located several AU  
70 from the protosun, separating the inner rocky (metals and silicates) and outer icy/gaseous  
71 (hydrogen, helium, and ices) regions (e.g., Lodders 2004).

72 Electromagnetic waves are able to penetrate into the shallow interior of gaseous planets and  
73 respond to the depth-dependent electrical conductivity of the atmosphere. This is set by the local  
74 thermodynamic state at any depth within the interior of the planet. In the case of the Earth, the  
75 surface and ionosphere form a closed cavity where electromagnetic waves can propagate. The  
76 closed cavity supports a set of normal electromagnetic modes with characteristics that depend on  
77 the physical dimensions of the cavity. When an impulsive electrical current such as lightning  
78 occurs within the cavity, the normal modes are excited to form the Schumann resonance  
79 spectrum. The Schumann resonances have been extensively used to investigate the lightning-  
80 thunderstorm connection. They have been conjectured to be excited in planetary environments  
81 that possess an ionosphere, from Venus to Neptune, as well as Titan, the largest moon of Saturn  
82 with a thick atmosphere (Simões et al. 2008a). Unlike Venus, Mars, and Titan, where  
83 atmospheric electric discharging phenomena remain uncertain, excitation of Tremendously and  
84 Extremely Low Frequency (TLF-ELF)<sup>1</sup> electromagnetic normal modes in the atmospheres of the  
85 outer planets is thought to be of highly probable because both Very Low Frequency (VLF) and  
86 optical signatures attributed to lightning have been detected on these planets (see Yair et al.  
87 2008, for a review). The normal mode frequencies of the Schumann resonances are related to  
88 cavity radius and medium conductivity, which in turn is dependent on the water vapor and ice  
89 abundances in the fluid envelope (Sentman 1990b; Liu 2006; Simões et al. 2008b). Detection of  
90 the Schumann resonances and a study of their properties could therefore indirectly yield the  
91 water content in the shallow interiors of these planets by way of its effects on the conductivity  
92 profile.

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<sup>1</sup> The nomenclature Ultra Low Frequency (ULF) is often used in ionospheric and magnetospheric sciences to designate frequencies below 3 Hz. For the sake of clarity and since this work is mostly related to electromagnetic wave propagation, we use the acronym TLF-ELF to define the frequency range 0.3-30 Hz, following the high frequency radio band classification analogy of the International Telecommunication Union.

93 Recent detection of the terrestrial Schumann resonances from orbit by the  
94 Communications/Navigation Outage Forecasting System (C/NOFS) satellite (Simões et al.  
95 2011a) unveils new capabilities for the investigation of planetary atmospheric electricity in other  
96 planets from orbit. Previously, Schumann resonance assessments required descent probes,  
97 balloons, or landers, but the C/NOFS results provide an original, remote sensing technique for  
98 TLF-ELF wave detection onboard orbiters; measurements in orbit are generally more versatile  
99 than *in situ* measurements.

100 In the following, after a brief theoretical description of the phenomenon, we examine the  
101 suitability of using Schumann resonances for determining the water/ice content in the fluid  
102 envelopes of the outer planets and, consequently, proving constraint on the volatile inventory of  
103 the protosolar nebula, hence to provide new constraints for Solar System formation models.

104

## 105 2. SCHUMANN RESONANCE THEORY

106 The propagation of low frequency electromagnetic waves within the cavity formed by two,  
107 highly conductive, concentric, spherical shells, such as those formed by the surface and the  
108 ionosphere of Earth, was first studied by Schumann (1952), and the resonance signatures of the  
109 cavity subsequently were observed in ELF spectra by Balser and Wagner (1960). Such a closed  
110 cavity supports both electric and magnetic normal modes. The lowest frequency of these modes  
111 is the transverse-magnetic mode of order zero (TM<sub>0</sub>) also sometimes called the transverse-  
112 electromagnetic (TEM) mode. These normal modes have an electric polarization that is radial,  
113 and a magnetic polarization that is perpendicular to the electric field and tangent to the surface of  
114 the planet. The modes may be excited by impulsive current sources within the cavity that, when  
115 observed as banded spectra, are known as the Schumann resonances. This phenomenon has been  
116 extensively used in atmospheric electricity investigations on Earth.

117 The normal mode frequencies (eigenfrequencies) of order  $n$ ,  $f_n$ , of a lossless, thin spherical  
118 cavity can be computed from (Schumann 1952)

119

$$120 f_n = \frac{c}{2\pi R} \sqrt{n(n+1)}, \quad (1)$$

121

122 where  $c$  is the velocity of light in vacuum,  $R$  the radius of the cavity (planet), and  $n=1,2,3,\dots$  the  
 123 corresponding order of the eigenmode. Taking into account the cavity thickness and medium  
 124 losses, a more accurate approximation of the eigenfrequencies yields

$$126 \quad f_n \approx \frac{c}{2\pi R} \sqrt{n(n+1) \frac{1-\frac{h}{R}}{\epsilon_r \left(1+i\frac{\sigma}{\epsilon_r \epsilon_0 2\pi f_n}\right)}}, \quad (2)$$

127  
 128 where  $h$  is the effective height of the ionosphere,  $\epsilon_r$  the relative permittivity and  $\sigma$  the  
 129 conductivity of a uniform medium, and  $\epsilon_0$  the permittivity of vacuum. The outer boundary is  
 130 chosen such that the skin depth,  $\delta_h$ , is much smaller at its location than the effective height of the  
 131 ionosphere, and

$$132 \quad \delta_h \cong \sqrt{\frac{2}{\mu_0 \omega \sigma}} \ll h, \quad (3)$$

133  
 134 with  $\mu_0$  the permeability of vacuum and  $\omega$  the angular frequency of the normal mode. Although  
 135 merely valid under the assumptions  $h \ll R$ ,  $\sigma \ll \omega \epsilon_0$ , and medium uniformity, Equation (2)  
 136 provides a simple method for assessing the eigenfrequency variation with cavity thickness and  
 137 medium losses; increasing these two parameters decreases the eigenfrequencies.  
 138

139 In addition to the eigenfrequencies, the cavity is characterized by a second parameter, known  
 140 as Q-factor, which measures the ratio of the accumulated field power to the power lost during  
 141 one oscillation period. The Q-factor measures the wave attenuation in the cavity and is defined  
 142 by

$$143 \quad Q_n \equiv \frac{Re(f_n)}{2 Im(f_n)} \approx \frac{f_n^p}{\Delta f_n^p}, \quad (4)$$

144  
 145 where  $Re$  and  $Im$  are the real and imaginary parts of  $f_n$ , and  $\Delta f_n^p$  is the full width at half  
 146 maximum of peak  $n$ ,  $f_n^p$ , in the Schumann resonance frequency. Good propagation conditions in  
 147 the cavity ( $Im(f_n) \rightarrow 0$ ), i.e., low wave attenuation, imply narrow spectral lines and, consequently,  
 148 high Q-factors.  
 149

150 Modeling ELF wave propagation on Earth is relatively straightforward compared to other  
151 planetary environments because several approximations are acceptable, namely (i) the cavity is  
152 thin, (ii) the surface is a perfect electric conductor, (iii) the conductivity profile is approximately  
153 exponential, and (iv) atmospheric permittivity corrections can be neglected. These conditions  
154 allow for longitudinal and transverse modes decoupling and simplify the analytical calculations  
155 (Greifinger & Greifinger 1978; Sentman 1990a). Unlike that of Earth, thick cavities containing  
156 dense atmospheres and sometimes undefined surfaces are more difficult to investigate, and  
157 numerical modeling is required. In this work, we use a finite element model previously employed  
158 to the study of TLF-ELF wave propagation in planetary cavities. This model solves Maxwell  
159 equations under full wave harmonic propagation (Simões 2007; Simões et al. 2007), eigenmode  
160 (Simões et al. 2008b, 2008c), and transient formalisms (Simões et al. 2009).

161

162

### 163 **3. SCHUMANN RESONANCES ON EARTH AND PLANETARY CONTEXT**

#### 164 **3.1 Earth Results Summary**

165 In the last decades, significant accomplishments have been reported in Schumann resonance  
166 measurements and modeling. Indeed, continuous monitoring of ELF waves from multiple  
167 stations around the world has been used to investigate lightning-thunderstorm and tropospheric-  
168 ionospheric connections, because Schumann resonance signatures are mostly driven by lightning  
169 activity and ionosphere variability. The interaction between the solar wind and the ionosphere  
170 distorts and modulates the upper boundary and the resultant cavity eigenfrequencies. The  
171 Schumann resonance signatures therefore vary over the 11-year solar cycle, as well as shorter  
172 temporal events such as solar flares; observations also show that the resonance amplitude,  
173 frequency, and cavity  $Q$ -factor vary during solar proton events. Another major interest of  
174 Schumann resonance studies on Earth is concerned with the processes linking lightning and  
175 thunderstorm activity to the global electric circuit. Currently, Schumann resonance studies of the  
176 Earth-ionosphere cavity are driven by three major research fields related to atmospheric  
177 electricity, specifically (i) the global electric circuit and transient luminous events such as sprites,  
178 (ii) tropospheric weather and climate change, and (iii) space weather effects. The most important  
179 Schumann resonance characteristics measured on the ground include:  $f \cong 7.8, 14.3, 20.8, 27.3,$   
180  $33.8 \text{ Hz}, \dots, Q \sim 5, E \sim 0.3 \text{ mVm}^{-1}\text{Hz}^{-1/2},$  and  $B \sim 1 \text{ pT}$ , where  $E$  and  $B$  are the electric and

181 magnetic fields, respectively. This work is focused on the outer planets, so we shall not elaborate  
182 further on Earth Schumann resonance matters, but the interested reader can find additional  
183 details in several reviews (Galejs 1972; Bliokh et al. 1980; Sentman 1995; Nickolaenko &  
184 Hayakawa 2002; Simões et al. 2011b).

185

### 186 **3.2 Planetary Environments**

187 The existence of Schumann resonances has been conjectured for most planets and a few  
188 moons. Mercury and our Moon, where lack of any significant atmosphere prevents the formation  
189 of a surface-ionosphere cavity, are obvious exceptions. Since a detailed description of each  
190 environment is not fundamental at this stage, we shall summarize the results relevant to this work  
191 only; additional information can be found elsewhere (see Simões et al. 2008a, for a review). In  
192 theory, normal modes of any cavity can be excited, provided a sufficiently strong impulsive  
193 excitation source is present to generate them. If the modes are not critically damped by high  
194 conductivity within the cavity, they would form a spectrum of distinct lines at the  
195 eigenfrequencies. One of the first questions to be answered is therefore the nature of the  
196 conductivity within the ionospheric-atmospheric cavities of the various planets of the solar  
197 system.

198

#### 199 **3.2.1 Venus**

200 Three major characteristics distinguish the cavity of Venus from that of the Earth: (i) the  
201 surface is not a perfect reflector of ELF waves, (ii) the cavity is more asymmetric, and (iii) the  
202 atmospheric density is larger. Moreover, although new reports by Russell et al. (2010) based on  
203 Venus Express data suggest that lightning activity is prevalent on Venus, the issue still remains  
204 controversial. This is mainly due to the lack of unequivocal optical observations of flashes in the  
205 clouds and a plausible required charging mechanism that will generate strong enough electrical  
206 fields to ensure breakdown in relatively short times to match the postulated rate (e.g., Yair et al.  
207 2008). There are nevertheless at least two works claiming observation of optical lightning on  
208 Venus: one performed onboard Venera 9 (Krasnopol'sky 1980) and another with a terrestrial  
209 telescope (Hansell et al. 1995). Although the expected Schumann eigenfrequencies are similar to  
210 those of Earth, surface losses can possibly lower the frequencies by as much as  $\sim 1$  Hz compared  
211 to those expected in a cavity with perfectly reflecting surface. Cavity asymmetry partially

212 removes eigenmode degeneracy and line splitting should be more marked than on Earth ( $\sim 1$  Hz).  
213 ELF wave attenuation is smaller than on Earth and, consequently, higher Q-factors are expected  
214 ( $Q \sim 10$ ). The most interesting feature, however, might concern the electric field altitude profile.  
215 Because of a significant atmospheric density, it is predicted that the Schumann resonance electric  
216 field profile should show a maximum at an altitude of  $\sim 32$  km, induced by refraction phenomena  
217 (Simões et al. 2008c), instead of a monotonic profile like on Earth. At this altitude, cavity  
218 curvature is balanced by atmospheric refraction and the wave vector is horizontal; this  
219 phenomenon is also predicted when the Fermat principle or ray tracing techniques are employed  
220 for much shorter wavelengths: at  $\sim 32$  km, a horizontal light beam propagates horizontally around  
221 the planet if scattering is negligible.

222

### 223 3.2.2 Mars

224 The electric environment of Mars remains uncertain despite the significant amount of data  
225 provided by several orbiters and landers over recent decades. Additionally, a highly  
226 heterogeneous surface and irregular magnetic field make models more complex and unreliable.  
227 Although the Martian cavity radius suggests higher eigenfrequencies than on Earth, the  
228 significant atmospheric conductivity decreases Schumann resonance frequencies and Q-factors  
229 as well. The fundamental eigenfrequency probably lies in the range 8-13 Hz; the most significant  
230 result, though, is a low Q-factor ( $Q \sim 2$ ) that implies significant wave attenuation. Thus, it is not  
231 clear whether triboelectric phenomena, even in massive dust storms, can sustain ELF resonances  
232 in the cavity. Interestingly, Schumann resonance monitoring could contribute to the study of a  
233 sporadic ionospheric layer probably induced by meteoroids (Molina-Cuberos et al. 2006).  
234 Attempts to remotely-sense the electromagnetic signature of the postulated electrical activity on  
235 Mars have been undertaken from Mars orbit and from Earth-based instruments. Ruf et al. (2009)  
236 conducted daily 5 h measurements using a new instrument on the Deep Space Network radio-  
237 telescope, and reported the detection of non-thermal radiation for a few hours that coincided with  
238 the occurrence of a deep dust storm on Mars. The spectrum of the non-thermal radiation showed  
239 significant peaks around predicted values of the lowest three modes of the Martian Schumann  
240 resonance (e.g., Pechony & Price 2004). Since Schumann resonance radiation is formed by  
241 discharges exciting the surface-ionosphere cavity, Ruf et al. (2009) interpreted their observations  
242 as indicative for the occurrence of lightning within the dust storm. However, the ELF peaks



243 reported imply large Q-factors ( $Q > 100$ ) and are almost equally spaced over the frequency range,  
244 contradicting a straightforward Schumann resonance interpretation. Anderson et al. (2011) used  
245 the Allen Telescope Array in an attempt to corroborate the previous results but did not detect any  
246 non-thermal emission associated with electrostatic discharges; it is nevertheless important to  
247 emphasize that they did not detect large-scale dust storms either. Gurnett et al. (2010) used the  
248 Mars Express MARSIS instrument to look for impulsive radio signals from lightning discharges  
249 of Martian dust storms and reported negative results. The search covered  $\sim 5$  years of data and  
250 spanned altitudes from 275 km to 1400 km and frequencies from 4.0 to 5.5 MHz, with a time  
251 resolution of  $91.4 \mu\text{s}$  and a detection threshold of  $2.8 \times 10^{-18} \text{ W m}^{-2} \text{ Hz}^{-1}$ . At comparable altitudes  
252 the intensity of terrestrial lightning is several orders of magnitude above this threshold. Although  
253 two major dust storms and many small storms occurred during the search period, no credible  
254 detections of radio signals from lightning were observed. The claim of Schumann resonance  
255 detection on Mars must be interpreted with extreme caution and requires confirmation.

256

### 257 3.2.3 Titan

258 Titan, the largest moon of Saturn, is the only body, other than Earth, where in situ  
259 measurements related to Schumann resonance have been attempted. Although convective clouds  
260 and storm systems have been detected in Cassini images, their composition, dynamics, and  
261 microphysics seem to be un-conducive to the emergence of electrical activity (Barth and Rafkin,  
262 2010). And indeed, despite repeated passages near Titan, Cassini did not detect any radio  
263 signature that can be attributed to lightning (Fischer & Gurnett 2011). The Huygens Probe did  
264 record ELF spectra during the descent upon Titan that exhibit a peak close to 36 Hz (Fulchignoni  
265 et al. 2005; Grard et al. 2006). Several laboratory tests on the flight spare and mockup models,  
266 including antenna boom vibration at cryogenic temperatures, revealed no artifact at the same  
267 frequency. In spite of progresses in Titan cavity modeling, the nature of this signal remained  
268 unclear for a while because the electric field signature was not fully consistent with that of a  
269 Schumann resonance (Simões et al. 2007; Béghin et al. 2007). The few VLF events recorded by  
270 Huygens, if related at all to lightning activity, imply a much lower flash rate than on Earth (Hofe  
271 2007; Simões 2007), inconsistent with the magnitude of the 36 Hz spectral line (Béghin et al.  
272 2007). Presently, the most promising mechanism that could explain the Huygens measurements  
273 involves an ion-acoustic turbulence resulting from the interaction of Titan with the

274 magnetosphere of Saturn (Béghin et al. 2007, 2009). Since Titan surface is a weak reflector ( $\delta_h >$   
275  $10^3$  km), ELF waves would propagate in the subsurface down to a depth where they would be  
276 reflected ( $\delta_h < 10$  km) by a water-ammonia liquid interface (Simões et al. 2007). Theoretical  
277 models predict the existence of a subsurface ocean (e.g., Lunine & Stevenson 1987; Tobie et al.  
278 2005), and the Huygens Probe measurements have been used for constraining the solid-liquid  
279 interface depth (Béghin et al. 2010). From a comparative planetology perspective, the surface  
280 properties of Titan fall between those of a perfect reflector, like on Earth, and those of a fuzzy,  
281 ill-defined surface, like on the giant planets.

282

### 283 3.2.4 Giant Planets

284 To our knowledge, only two works on TLF-ELF wave propagation and Schumann resonance  
285 in the giant planets have previously been published. Sentman (1990b) calculated the Schumann  
286 resonance parameters for Jupiter by computing from first principles the conductivity profile of  
287 shallow interior, then by assuming a perfectly conducting ionosphere estimating the  
288 eigenfrequencies and Q-factors. Since Jupiter's radius is one order of magnitude larger than that  
289 of Earth, the expected Schumann resonances are about tenfold smaller (Equation (1)). Simões et  
290 al. (2008b) considered improved conductivity profiles and also included the permittivity  
291 contribution because the cavity's inner boundary is located deep within the gaseous envelope,  
292 where refraction phenomena play a role. In the latter work, the wave propagation model was  
293 generalized to the other giant planets because similar conditions apply. Unlike the Jovian planets  
294 where measurements provided some atmospheric composition constrains, the water content  
295 uncertainty in the fluid envelopes of Uranus and Neptune is significant, implying electric  
296 conductivity profiles possibly differing by several orders of magnitude (Liu 2006; Liu et al  
297 2008). Simões et al. (2008b) showed that Schumann resonance measurements could be used to  
298 constrain the conductivity profile and the water content. The detection by C/NOFS of ELF waves  
299 leaking into space from the Earth surface-ionosphere cavity prompts a new approach for the  
300 investigation of Schumann resonances in other planets and, consequently, of the water content in  
301 their gaseous envelopes.

302

### 303 3.3 C/NOFS Measurements

304 The Vector Electric Field Instrument (VEFI) on the C/NOFS satellite offers new capabilities  
305 for the investigation of planetary atmospheric electricity, demonstrating that ELF wave detection  
306 no longer requires in situ techniques. VEFI consists primarily of three orthogonal 20 m tip-to-tip  
307 double probe antennas (Pfaff 1996) and is dedicated to the investigation of ionospheric  
308 irregularities, namely spread-F and related phenomena, and to the improvement of space weather  
309 forecast. The instrument measures AC and DC electric and magnetic fields; it also includes  
310 lightning optical detectors and a Langmuir probe (Pfaff et al. 2010). In the nominal mode, the  
311 VEFI electric field sampling is  $512 \text{ s}^{-1}$ , with sensitivity better than  $10 \text{ nVm}^{-1}\text{Hz}^{-1/2}$ . Remarkably,  
312 C/NOFS detected Schumann resonances from orbit, in the altitude range 400-850 km, above the  
313 ionospheric F-peak, i.e., outside the surface-ionosphere cavity. These signatures are  
314 unambiguous, and more perceptible and clear under specific conditions: in a quiet ionosphere,  
315 during nighttime, over equatorial regions developing mesoscale convective systems, while  
316 intense lightning bursts are seen. Figure 1 shows typical electric field data recorded on 2008 May  
317 31 during minimum solar activity. Spectrograms of the meridional/vertical and zonal/horizontal  
318 components are presented, as well as mean spectra integrated through the whole orbit for better  
319 peak visualization. Data are calibrated but intentionally not filtered to illustrate VEFI  
320 measurements robustness, namely instrument sensitivity and Schumann resonance features  
321 resolution. The Schumann resonance amplitude varies between about 0.01-0.1 and 0.1-3  $\mu\text{Vm}^{-1}$   
322  $\text{Hz}^{-1/2}$  during day and nighttime, respectively. During nighttime, the average electric field is  
323  $\sim 0.25 \mu\text{Vm}^{-1}\text{Hz}^{-1/2}$  in the altitude range covered by C/NOFS. Based on modeling, plasma  
324 anisotropy seems to allow ELF wave propagation through the ionosphere in the plane  
325 perpendicular to the magnetic field (e.g., Madden & Thompson 1965; Grimalsky et al. 2005),  
326 bearing resemblance to resonance tunneling phenomena of waves in stratified cold plasma  
327 (Budden 1979). Although more elaborate modeling is necessary to understand the leakage  
328 mechanism thoroughly, propagation in the whistler and extraordinary modes seem compatible  
329 with the observed results. These C/NOFS findings suggest that new remote sensing capabilities  
330 for atmospheric electricity investigations in the vicinity of planets possessing an internal  
331 magnetic field could be envisaged from an orbiter.

332

333

(FIGURE 1)

334

## 4. OUTER PLANETS DYNAMICS AND EVOLUTION

### 4.1 Giant Planets Composition

The formation of the gaseous giant planets remains a mystery because current theories are incapable of explaining how their cores can form fast enough and accumulate considerable amounts of gas before the protosolar nebula disappears. In fact, the lifetime of the protoplanetary disk seems to be shorter than the time necessary for planetary core formation. Another open question related to the giant planets formation is their migration. Likely, interaction with the disk causes rapid inward migration and planets would reach the inner regions of the Solar System still as sub-Jovian objects, i.e., mostly as solid bodies (e.g., Benz et al. 2000). On the other hand, according to the nebular hypothesis, Uranus and Neptune are currently located where the low density of the protoplanetary disk would have made their formation improbable. They are believed to have formed in orbits near Jupiter and Saturn and migrated outward to their present positions (e.g., Kallenbach et al. 2003). The unknown abundance of volatiles in the protosolar nebula leads to uncertainty on its gravitational and thermodynamic parameters and hampers the development of accurate accretion models (Guillot 2005). Therefore, an accurate assessment of the ice fraction of volatiles in the giant planets is required for providing a better estimate of the protoplanetary disk initial composition and an improved model of the Solar System evolution.

### 4.2 Jupiter and Saturn

The atmospheres of the Jovian planets are mainly composed of hydrogen and helium with minor mole fractions of other constituents, namely ammonia, methane, and water. Although remote sensing or in situ measurements of Jupiter and Saturn atmospheres have been made, the global composition, and water content in particular, remains uncertain. Additionally, a generalization of the atmospheric composition to the entire fluid envelope may be too broad. In the present work, we consider the conductivity profiles computed by Sentman (1990b), Nellis et al. (1996), and Liu (2006). The electrical conductivity of the interiors of Jupiter and Saturn is mainly due to hydrogen; the mean composition is shown in Table 1. Figure 2 shows the conductivity profile as a function of the planet normalized radial distance.

(FIGURE 2)

366 The conductivity saturation (plateau) is due to hydrogen metallization (Nellis et al. 1996) but  
367 TLF wave reflection (inner boundary of the cavity) takes place at lower depths. For the sake of  
368 clarity, we define the interior of the giant planets as the region where the pressure is larger than 1  
369 bar; this reference level also determines the radius of the planet. We consider conductivity  
370 profiles derived by Sentman (1990b) for Jupiter and by Liu (2006) and Liu et al. (2008) for  
371 Jupiter and Saturn.

372

### 373 **4.3 Uranus and Neptune**

374 Models predict that Uranus and Neptune (called Uranian planets in the rest of the paper) have  
375 similar internal structure (e.g., Lewis 1995). Estimations based on physical characteristics such  
376 as mass, gravity, and rotation period, and on thermodynamic properties as well, predict that the  
377 Uranian planets have an internal rocky core (iron, oxygen, magnesium, and silicon – magnesium-  
378 silicate and iron compounds), surrounded by a mixture of rock and ice (water, ammonia,  
379 methane), and an external gaseous envelope (hydrogen and helium permeated by an unknown  
380 fraction of ice). The intermediate envelope is possibly liquid because of high pressure and  
381 temperature. Considering distances normalized to the radius of the planet, the transition between  
382 the gaseous and intermediate envelopes is located at  $\sim 0.8$  and  $0.84$  for Uranus and Neptune,  
383 respectively. In the present study, we are mainly concerned with the properties of the outer layer,  
384 the gaseous envelope, where TLF-ELF waves would propagate. Figure 3 shows the conductivity  
385 profiles of the interior of Uranus and Neptune as functions of the normalized radial distance. The  
386 sharp variation in conductivity coincides with the transition between the outer and intermediate  
387 envelopes. Conductivity may vary significantly, depending on the water ice mixing ratio in the  
388 gaseous envelope. For the same depth, a water mixing ratio of 0.1 might increase the  
389 conductivity by as much as 10 orders of magnitude compared to that of a dry envelope, a fact  
390 that clearly illustrates the extreme sensitivity of TLF-ELF wave propagation conditions to the  
391 gaseous envelope water mixing ratio.

392

393

(FIGURE 3)

394

395 Unlike Jupiter and Saturn, the magnetic fields of the Uranian planets are quite unusual. The  
396 magnetic fields of Uranus and Neptune are tilted by  $59^\circ$  and  $47^\circ$  with respect to the axes of

397 rotation, and are also displaced from the planet's center by  $0.31 R_U$  and  $0.55 R_N$ , respectively.  
398 This atypical magnetic field structure results in highly asymmetric magnetospheres and suggests  
399 that it is generated in the intermediate, possibly liquid, envelope rather than in the core itself as  
400 in the other planets (Ness et al. 1986; Connerney et al. 1991). In addition to a strong quadrupolar  
401 moment contribution, Uranus sideways rotation complicates even further the magnetic field  
402 distribution. The magnetic field distribution is a second order correction for eigenfrequency and  
403 Q-factor assessments because the medium is highly collisional in most of the envelope.  
404 However, the magnetic field correction is fundamental to investigate the cavity leakage. The  
405 equatorial magnetic fields are given in Table 1.

406  
407 (TABLE 1)  
408

## 409 5. NUMERICAL MODEL

410 The cavities of the gaseous giant planets are intricate, and so the standard analytical  
411 approximations used for Earth are unsuitable; thus, numerical modeling is necessary. We use an  
412 approach similar to that employed by Simões et al. (2008b, 2008c) to study TLF-ELF wave  
413 propagation in planetary environments. The numerical model is based on the finite element  
414 method and solves Maxwell equations with specific boundary conditions and medium properties.  
415 The algorithm calculates the eigenfrequencies,  $Q$ -factors, and electromagnetic field distribution  
416 within the cavity. The most important parameters for running the numerical model include: (i)  
417 the conductivity profile of the atmosphere and ionosphere ( $\sigma_{\text{iono}}$ ), (ii) the conductivity profile of  
418 the interior ( $\sigma_{\text{int}}$ ), (iii) the permittivity profile of the interior ( $\epsilon_{\text{int}}$ ), (iv) the depth of the inner  
419 boundary ( $d$ ), and (v) the height of the outer boundary ( $h$ ). The inner and outer boundaries are  
420 located where  $\delta_i \ll d$  and  $\delta_i \ll h$ , respectively, where  $\delta_i \sim 10$  km. The inner boundary coincides  
421 roughly with the interface between the gaseous envelope and the metal (Jupiter, Saturn) or  
422 icy/liquid (Uranus, Neptune) medium (Liu 2006; Liu et al. 2008).

423 The atmospheric conductivity is computed from the electron density and collision frequency  
424 profiles, which are derived from pressure, temperature, and composition data recorded during  
425 several missions, namely Pioneer, Voyager, Galileo, and Cassini. The conductivity profiles of  
426 the planetary interiors are shown in Figures 2 and 3. Since density increases with depth and the  
427 vacuum approximation is no longer valid, we employ the approach of Simões et al. (2008b,

428 2008c) to derive the permittivity of the interior of the giant planets, assuming that: (i) the  
429 refractivity is a linear function of gas density, (ii) the medium response can be extrapolated from  
430 the radiofrequency to the TLF-ELF range, i.e., non-dispersive medium conditions at low  
431 frequency apply, (iii) the contributions other than that of hydrogen are neglected (more elaborate  
432 approaches are considered if the water content ratio exceeds  $\sim 0.1\%$ ), and (iv) the relative  
433 permittivity of liquid hydrogen is  $\sim 1.25$ . A more elaborated analysis of refractivity effects in  
434 ELF wave propagation can be found elsewhere (Simões et al. 2008c). We first employ the  
435 eigenvalue analysis to determine the eigenfrequencies and Q-factors of isotropic cavities (Simões  
436 et al. 2008b). For a qualitative estimation of the cavity leakage, the electric and magnetic fields  
437 are computed with a full wave harmonic propagation algorithm in an anisotropic medium. For  
438 the sake of simplicity, we employ a vertical Hertz dipole to model the electromagnetic sources  
439 (Simões et al. 2009) and consider a dipolar static magnetic field of known magnitude at the  
440 equator (Table 1).

441 In addition to the conductivity profile variability with water content, estimates of the TLF-  
442 ELF wave magnitude resulting from cavity leakage are invaluable for establishing the detection  
443 range and defining instrumentation requirements. We therefore use a full wave harmonic  
444 propagation model to compute the electric and magnetic field amplitudes as function of distance  
445 to the source. The open boundary ( $r \rightarrow \infty$ ) is approximated by a Perfectly Matched Layer (PML)  
446 placed at  $r \sim 10^2 R$ . The PML approach is used to avoid wave reflection on the edge of the  
447 domain. We consider a vertical Hertz dipole radiating in the TLF range, of arbitrary amplitude  
448 and located at  $r=R$ , and compute the electromagnetic field distribution inside and outside the  
449 cavity. A similar approach to that applied by Simões et al. (2009) to the Earth cavity in the VLF  
450 range is employed here to derive the conductivity tensor on the giant planets, i.e., taking into  
451 account the Pedersen and Hall conductivity corrections. The conductivity tensor is derived from  
452 the Appleton-Hartree dispersion relation that describes the refractive index for electromagnetic  
453 wave propagation in cold magnetized plasma.

454 The present numerical model has already been used for estimating eigenfrequencies of  
455 planetary environments and has been validated against Earth cavity data, namely ELF spectra  
456 and atmospheric conductivity. Consistent results are therefore expected as long as the  
457 conductivity profiles are reliable. For the sake of simplicity, we consider a scalar conductivity to  
458 evaluate cavity eigenvalues because anisotropic corrections (Budden 1979; Simões et al. 2009)

459 are small compared to the conductivity profile uncertainty. Nonetheless, the conductivity tensor  
460 is included in the full wave harmonic propagation model to compute the electric and magnetic  
461 field amplitude resulting from cavity leakage, which allows for spacecraft-planet distance versus  
462 instrument sensitivity assessments. We choose 2D axisymmetric approximations whenever  
463 possible and 3D formulations otherwise.

464

## 465 **6. RESULTS**

466 In this work we address wave propagation primarily in the Uranian planets for the following  
467 reasons. First, the major objective is to investigate the suitability of the proposed technique for  
468 estimating the water content in the gaseous envelopes from Schumann resonance measurements.  
469 Second, water content uncertainty in the gaseous envelope of Uranus and Neptune is large, and  
470 therefore the technique proposed here would be more valuable for those environments. Third,  
471 unless significantly different conductivity profiles are conjectured, the eigenfrequencies and Q-  
472 factors of the cavities of Jupiter and Saturn would be similar to those reported previously (cf.  
473 Table 2 and results reported by Simões et al. 2008b). Finally, enhanced parameterizations are  
474 deemed necessary to quantify electromagnetic field leakage through planetary ionospheres,  
475 namely regarding source characteristics such as spatial and temporal variability of lightning;  
476 since the magnetic fields of the Jovian planets are stronger than those of Uranus and Neptune,  
477 anisotropic corrections should be more important there. A more elaborate model is nevertheless  
478 under development to compute wave propagation through the ionosphere, estimate cavity  
479 leakage as a function of lightning, ionospheric, and magnetospheric characteristics that may  
480 provide useful predictions for the Juno (en route to Jupiter) and Cassini (currently operating in  
481 orbit at Saturn) missions or future endeavors.

482 Voyager 2 measured the ionospheric electron density profile (Lindal et al. 1987; Tyler et al.  
483 1989; Lindal 1992) with some discrepancy between ingress and egress, especially in the case of  
484 Uranus. Two conductivity profiles of the atmosphere and ionosphere are derived for Uranus from  
485 the Voyager data sets, based on analogies with Earth aeronomy and modeling; in the case of  
486 Neptune, a single profile is used (Capone et al. 1977; Chandler & Waite 1986). Since the  
487 eigenfrequencies are little affected by atmospheric conductivity uncertainties due to the  
488 dominance of the interior contribution, the present model takes into account deeper variability  
489 only. Supplementary information regarding the calculation of atmospheric and ionospheric



490 conductivity profiles of the giant planets can be found elsewhere (Sentman 1990b; Simões et al.  
491 2008b).

492 In the case of the Jovian planets, where the water content uncertainty appears to be smaller  
493 than for Uranus and Neptune, we consider the conductivities shown in Figure 2 and compute  
494 eigenfrequencies and Q-factors of the mean, maximum, and minimum profiles. Table 2 shows  
495 the results of the eigenfrequencies and Q-factors of the three lowest eigenmodes of Jupiter and  
496 Saturn. Although the conductivity profile uncertainty produces minor variations in  
497 eigenfrequency and Q-factors, Schumann resonances could be used to confirm whether the  
498 hydrogen ionization processes are realistic as function of depth, and to assess impurity mixing  
499 ratios in the envelope as well. For Jupiter, the results for conductivity profiles derived by  
500 Sentman (1990b) and Liu (2006) produce somewhat dissimilar eigenfrequencies and Q-factors  
501 due to differences in the conductivity profile. These results are also important to confirm that  
502 eigenfrequencies and Q-factors are more sensitive to the conductivity profile than to cavity  
503 shape, e.g., equatorial versus polar radius.

504

505 (TABLE 2)

506

507 Figure 3 takes into account the water content uncertainty in the gaseous envelope of the  
508 Uranian planets and shows the consequences for the conductivity profile. Because of the  
509 significant conductivity profile uncertainty, we compute the eigenfrequencies and Q-factors for  
510 various cavity parameterizations. To facilitate the comparisons among various parameters,  
511 namely water mixing ratio, an exponential conductivity profile with two parameters is  
512 considered

513

$$514 \quad \sigma_{int}(r) = \sigma(r_o) e^{(r_o-r)/H_d} = \sigma_o e^{(r_o-r)/H_d}, \quad (5)$$

515

516 where  $r$  is the radial distance,  $r_o < r < R+h$ ,  $r_o = R-d$ , and  $H_d$  is the interior conductivity profile scale  
517 height. A conductivity profile is therefore defined by the ordered pair  $\{H_d, \sigma_o\}$ . Figures 4-5 show  
518 the eigenfrequencies and Q-factors of the cavities of Uranus and Neptune as a function of  
519 conductivity profile parameterization. Table 2 shows the three lowest eigenmodes computed for  
520 the conductivity profiles shown in Figure 3 (water content: 0, 0.01, and 0.1).

521 The plots presented in Figures 4 and 5 for Uranus and Neptune, respectively, correspond to  
522 specific cavity configurations, using Equation (5) and the conductivity constraints shown in  
523 Figure 3. The water content (magenta lines) is derived from the evaluations of Nellis et al. (1996)  
524 and Liu (2006). These estimates are indicative only and the results should be interpreted with  
525 caution. The conductivity profiles may be unrealistic but they are nonetheless representative and  
526 lend themselves to a qualitative discussion that illustrates how, conversely, the Schumann  
527 resonance characterizes the conductivity profile. Figure 4 presents the eigenfrequencies and Q-  
528 factors of the Uranus cavity for the three lowest eigenmodes as functions of interior scale height,  
529  $H_d$ , and interface conductivity,  $\sigma_o$ . Although one eigenmode is usually sufficient to identify the  
530 corresponding exponential conductivity profile ( $\{f_1, Q_1\} \rightarrow \{H_d, \sigma_o\}$ ), we present the  
531 eigenfrequencies and Q-factors of a few eigenmodes for information. The left-hand side plots in  
532 Figure 4 illustrate the importance of characterizing multiple modes with both the  
533 eigenfrequencies and Q-factors. These plots show two yellowish stripes corresponding to similar  
534 frequencies but different Q-factors; the effect is more evident in the lowest eigenmode. This  
535 effect illustrates how multiple peaks in the Schumann resonance spectrum can be used to further  
536 constrain the water mixing ratio. While the convex boundaries in Figures 4 and 5 represent the  
537 dry envelope limit, the concave ones result from multiple constraints, namely minimum  
538 conductivity close to the inner boundary, dry envelope conditions, and monotonic conductivity  
539 profiles. If the conductivity profile scale heights of Figure 3 are realistic, then the most plausible  
540 cavity parameterizations are found along the gray line. Bites in the plots top-right edge are due to  
541 the lack of eigenvalues; wave resonance is hindered because a critical damping is reached caused  
542 by high water content. As expected, a combination of high  $H_d$  and  $\sigma_o$  entail significant cavity  
543 losses and, comparatively, wave propagation conditions seem more favorable in Uranus than in  
544 Neptune, confirming previous simulations (Simões et al. 2008b). The water content affects both  
545 the frequency and Q-factor though more importantly in the latter (see the magenta isolines in  
546 Figures 4-5).

547

548

(FIGURE 4 and FIGURE 5)

549

550 A single scale height is considered for the Uranus and Neptune interiors for the sake of  
551 simplicity, but a realistic conductivity profile is certainly more intricate. On Earth, for example,

552 the atmospheric conductivity profile is better described by two scale heights (Greifinger &  
553 Greifinger 1978; Sentman 1990a). An improved model that addresses a weakness in the two-  
554 scale height model is when the local scale height changes rapidly in the region of maximum  
555 Joule dissipation, and referred to as the “knee model,” has been described by Mushtak and  
556 Williams (2002). In the case of the giant planets, mainly Uranus and Neptune, a multiple scale  
557 height profile would be preferable in order to differentiate interior, interior-atmospheric, and  
558 ionospheric parameterizations. Information from additional eigenmodes can be used to  
559 characterize conductivity profiles with multiple scale heights. In principle, the number of scale  
560 heights that can be constrained in the profile is equal to the number of spectral peaks detected,  
561 provided eigenfrequencies and Q-factors can be both measured accurately. The most  
562 straightforward approach would consist in solving the direct problem iteratively, starting from an  
563 initial guess and employing minimization techniques to obtain the conductivity profile best fit,  
564 which would then yield an estimation of the water content of the gaseous envelope. If possible, it  
565 should also be attempted to derive conductivity profiles directly from eigenmode information,  
566 i.e., ordered pairs  $\{f_n, Q_n\}$ .

567 Figure 6 shows the electric field magnitude of a radiating dipole as a function of the  
568 normalized radius of the planet. Far from the planet, where the signal resulting from cavity  
569 leakage propagates almost in a vacuum, the electromagnetic field variation with distance  
570 approaches the power law ( $E, B \propto r^{-1}$ ) resulting from spherical wave propagating in a lossless  
571 medium. Amplitude asymptotic convergence to a theoretical solution therefore corroborates the  
572 PML approach at large distance. Since absolute comparisons are not viable because lightning  
573 stroke characteristics are unknown, a source of arbitrary amplitude is selected. In addition to the  
574 previous Uranian environments, we now consider a cavity with an ionospheric parameterization  
575 equivalent to that of Earth. Although physically not representative, this comparison determines  
576 whether detecting leakage from Uranus and Neptune cavities is more demanding than at Earth.  
577 The electric field profiles of Figure 6 suggest that wave leaking detection is more favorable than  
578 at Earth because ionospheric attenuation is weaker. Since the electron density peak in Uranus  
579 and Neptune is about 2 and 3 orders of magnitude lower than at Earth (Lindal et al. 1987; Tyler  
580 et al. 1989), integrated Pedersen and Hall conductivities provide less wave attenuation through  
581 the ionosphere. For example, at a distance of 1.1 R, the expected electric field would be two  
582 orders of magnitude higher than on Earth for a similar electromagnetic source. However,

583 interpretation must be made cautiously because cavity leakage is also a function of Q-factor,  
584 aeronomy processes, and lightning stroke power and rate characteristics. Consequently,  
585 subsequent investigations of atmospheric electricity, namely lightning processes, are needed so  
586 that cavity leakage assessments could be improved. In theory, considering ionospheric plasma  
587 density, magnetic field parameterization and assuming similar electromagnetic source  
588 characteristics, anticipated cavity leakage in the Jovian planets is stronger than on Earth but  
589 weaker than on Uranus and Neptune.

590

591

(FIGURE 6)

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## 7. DISCUSSION

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The most accurate way of evaluating the water content profile of the giant planets is employing in situ techniques for measuring the water mixing ratio in the gaseous envelope. This approach was used by the Galileo Probe Mass Spectrometer during the descent through the atmosphere of Jupiter down to ~20 bar (Mahaffy et al. 2000; Atreya et al. 2003). However, only a small fraction of the envelope has been explored. Other solutions involve Earth-orbiting observatories or dedicated spacecraft around the planets, e.g., Cassini at Saturn, employing infrared, optical, or ultraviolet spectrometry to infer atmospheric composition (e.g., Fouchet et al. 2005; Baines et al. 2009). The microwave radiometer part of Juno, a forthcoming mission to Jupiter, may provide accurate water content estimates possibly down to about 200 bar (Matousek 2007). These options are reliable and accurate but allow for estimates of the envelope outer shallow layer only. Since the connection between water content and electric conductivity is well established, in situ measurements of the conductivity profile would provide an indirect method for water content assessments. During the descent in the atmosphere of Titan, the Permittivity, Wave, and Altimetry analyzer onboard the Huygens Probe performed electric conductivity measurements from about 140 km down to the surface (Fulchignoni et al. 2005; Grard et al. 2006; Hamelin et al. 2007; López-Moreno et al. 2009). This type of approach would be applicable in the giant planets down to moderate depths only. Given that a connection among the planetary Schumann resonance frequencies, conductivity profile, and water content exists, TLF-ELF measurements provide a practical method for inferring the water content in the envelope.

614 On the other hand, C/NOFS data show that measurements inside the cavity are not mandatory  
615 and that a remote sensing method is likely to be practical for planets that possess a magnetic  
616 field. Additionally, unlike other solutions that offer local measurements only, Schumann  
617 resonance measurements would provide a global distribution of the conductivity profile and,  
618 consequently, better estimates of the mean water content. As shown in Table 2 and Figures 4-5,  
619 Schumann resonance modes can be used to estimate global water contents up to a few percent  
620 in the Uranus and Neptune gaseous envelopes and, to a lesser extent, to confirm whether the  
621 conductivity models of Jupiter and Saturn are realistic. Detection of terrestrial Schumann  
622 resonance signatures onboard C/NOFS unveils new remote sensing capabilities for investigating  
623 atmospheric electricity and tropospheric-ionospheric coupling mechanisms, not only on Earth  
624 but also other planetary environments that possess a magnetic field. Observation of Schumann  
625 resonances above the ionospheric F-peak was unexpected and requires revisiting analytical and  
626 numerical models, which are not fully consistent with C/NOFS observations. However, although  
627 analytical and numerical modeling requires significant improvements, it is clear that medium  
628 anisotropy plays a key role in cavity leakage.

629 The snow line is an important concept to address the water ice condensation front in  
630 protoplanetary disk accretion models, to investigate convective and radiation phenomena as well  
631 as and chemical processes, and was allegedly located near the orbit of Jupiter when planets  
632 formed. The condensation front would be expanding during the solar nebula coalescence and  
633 subsequent disk accretion processes, and then receding again throughout the cooling phase. For  
634 example, Stevenson & Lunine (1988) argue that the Galilean satellites formed later than the  
635 proto-Jupiter, allowing for late accretion of water into these moons. Estimates of the relative  
636 abundance and variability of the various elements in the Solar System, in particular with respect  
637 to solar average composition, are frequently achieved from isotopic measurements. Information  
638 on the relative enrichment and depletion of the various elements is then used to investigate the  
639 early stages of the Solar System. Measurements made by the Galileo Probe (Mahaffy et al.,  
640 2000) in Jupiter atmosphere found less water than expected. Several explanations have been  
641 proposed, including (i) non representative measurements due to sampling of a dry area of the  
642 atmosphere, (ii) a larger fraction of oxygen is trapped in the core in the form of silicates, (iii) the  
643 water ratio would be lower than expected in the Solar System, (iv) the snow line was located  
644 farther from the Sun, suggesting more water is diffused toward the periphery of the Solar

645 System. Relocating the snow line farther away would imply that the Uranian planets are water-  
646 enriched; in the case of Neptune, the water enrichment could be several hundred times larger  
647 (e.g., Lodders & Fegley 1994). However, there are also theoretical models that may be consistent  
648 with water and oxygen depletion (Fegley & Prinn 1988). Measurement of water mixing ratios in  
649 the giant planets would thus provide useful data for constraining protoplanetary disk accretion  
650 models, offering a better distribution of water throughout the Solar System.

651 Figure 7 illustrates the rationale linking the water mixing ratio, electrical conductivity profile,  
652 remote sensing and in situ measurement techniques, Schumann resonance spectra, and  
653 protoplanetary disk parameters. The water mixing ratio in the gaseous envelope plays a key role  
654 in atmospheric chemistry, which drives the electrical conductivity profile through molecular  
655 reaction rates - e.g., ionization and recombination - and electron and ion mobility. Along with  
656 geometry parameters such as size, the conductivity profile drives the Schumann resonance  
657 spectrum in the cavity. Both TLF-ELF electric and magnetic field measurements can be used to  
658 estimate Schumann resonance signatures. Remote sensing is often more versatile than in situ  
659 measurements. For example, electric field measurements are frequently noisier onboard descent  
660 probes due to shot noise, mainly below 10 Hz. A descent vessel is also more susceptible to  
661 vibrations, which introduce additional artifacts to the spectrograms. As suggested by our  
662 calculations, high water mixing ratios would shift Schumann resonance toward lower frequencies  
663 and produce broader peaks as well as weaker signatures. In the case of Uranus and Neptune, a  
664 water mixing ratio of  $\sim 0.1$  might change the frequencies and Q-factors by a factor of 2 and 15  
665 compared to those related to dry envelopes. For the sake of comparison, variability of  
666 eigenfrequencies and Q-factors on Earth due to lightning and ionospheric dynamics is less than  
667 10% and 50%, respectively. Since a 50% enrichment or depletion of the water mixing ratio in the  
668 gaseous envelope of Jupiter with respect to the solar average has significant implications for  
669 protoplanetary disk models, discrimination between a water mixing ratio of 0.1 and 0.01 in the  
670 Uranian planets would provide key information for a better understanding of the formation and  
671 evolution of the Solar System.

672

673

(Figure 7)

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## 8. CONCLUSION

Limited data of volatiles abundance, namely water, ammonia, and methane in the outer planets prevent the development of accurate models of the protoplanetary disk dynamics, from which the Solar System evolved. Thus, knowledge of the water mixing ratio in the gas giants is crucial to constraining the protosolar nebula composition. Water content estimates have been measured so far with both in situ and remote sensing techniques. These approaches generally yield local atmospheric composition only, down though to pressure levels of tens of bars. However, extrapolating local composition measurements to the whole gaseous envelope might be inappropriate, particularly at large depths.

We propose here a new approach for estimating the global water content of the giant planet envelopes from Schumann resonance measurements. Water has a clear impact on the electrical conductivity and Schumann resonance signatures. Compared to a dry gaseous envelope, the predicted eigenfrequencies of the cavity of Uranus and Neptune show a 3-fold decrease when the water content reaches 10%. The Q-factors are even more sensitive and decrease by as much as a factor of 40. We therefore advocate performing in situ and remote sensing TLF-ELF electric and magnetic field measurements to probe the water global distribution in the gaseous envelopes, at depths of hundreds, possibly thousands, of kilometers. As seen from the C/NOFS satellite ELF spectra, Schumann resonance detection from orbit is feasible, which presents an obvious advantageous compared to in situ observations. Assuming similar lightning characteristics, preliminary models shows that wave leakage in the outer planets would be stronger than on Earth, suggesting detection of Schumann resonance signatures may even be easier there. Identification of multiple peaks from TLF-ELF spectra would further improve the conductivity profile and corresponding water content estimates. Combining both remote sensing and in situ techniques would of course strengthen synergistic analyses of the volatiles composition.

A Schumann resonance spectrum will be excited in the cavity of the gaseous giants if there are sufficiently powerful electrical drivers, such as lightning. Modeling confirms that with plausible conductivity profiles the distinctive resonance spectrum will form, and therefore be usable for probing the conductivity of the shallow interior of the planets. Electric and magnetic antennas could therefore be used not only to study atmospheric electricity and wave propagation but to estimate water content in the gaseous envelopes, to infer the volatile abundance in the protosolar nebula from which the Solar System evolved, and to constrain the water ice

707 condensation front and better locate the snow line in protoplanetary disk accretion models. The  
708 accurate assessment of the water content in the giant planets could also perhaps contribute for  
709 understanding the formation and dynamics of outer Solar System objects, from the Kuiper belt to  
710 the Oort cloud.

711

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716

## 717 **REFERENCES**

718 Anderson, M. M., Siemion, A. P.V., Barott, W. C. et al. 2011, The Allen telescope array search for electrostatic  
719 discharges on Mars, arXiv:1111.0685v1 [astro-ph.EP] 2 Nov 2011

720

721 Atreya, S. K., Mahaffy, P. R., Niemann, H. B., Wong, M. H., & Owen, T. C. 2003, Composition and origin of the  
722 atmosphere of Jupiter - an update, check and implications for the extrasolar giant planets, *Planet Space Sci*, 51, 105

723

724 Baines, K. H., Delitsky, M. L., Momary, T. W. et al. 2009, Storm clouds on Saturn: Lightning-induced chemistry  
725 and associated materials consistent with Cassini/VIMS spectra, *Planet Space Sci*, 57, 1650, doi:  
726 10.1016/j.pss.2009.06.025

727

728 Balser, M., & Wagner, C. A. 1960, Observations of earth-ionosphere cavity resonances, *Nature*, 188, 638,  
729 doi:10.1038/188638a0

730

731 Barth, E.L. & Rafkin, S. C. R. 2010, Convective cloud heights as a diagnostic for methane environment on Titan,  
732 *Icarus*, 206, 467, doi: doi:10.1016/j.icarus.2009.01.032

733

734 Béghin, C., Canu, P., Karkoschka, E. et al. 2009, New insights on Titan's plasma-driven Schumann resonance  
735 inferred from Huygens and Cassini data, *Planet Space Sci*, 57, 1872

736

737 Béghin, C., Hamelin, M., & Sotin, C. 2010, Titan's native ocean revealed beneath some 45 km of ice by a  
738 Schumann-like resonance, *C R Geosci*, 342, 425

739

740 Béghin, C., Simões, F., Krasnoselskikh, V. et al. 2007, A Schumann-like resonance on Titan driven by Saturn's  
741 magnetosphere possibly revealed by the Huygens Probe, *Icarus*, 191, 251, doi:10.1016/j.icarus.2007.04.005



742  
743 Benz, W., Kallenbach, R., & Lugmair, G. 2000, From Dust to Terrestrial Planets, Space Sciences Series of ISSI,  
744 vol. 9, Reprinted from Space Science Reviews, vol. 92/1-2, 432 p., ISBN: 978-0-7923-6467-2  
745  
746 Bliokh, P. V., Nickolaenko, A. P., & Filippov, Yu. F. 1980, Schumann resonances in the Earth-ionosphere cavity,  
747 (Oxford, England: Peter Peregrinus)  
748  
749 Budden, K. G. 1979, Resonance tunnelling of waves in a stratified cold plasma, Royal Society (London),  
750 Philosophical Transactions, Series A, 290, 405  
751  
752 Capone, L. A., Whitten, R. C., Prasad, S. S., & Dubach, J. 1977, The ionospheres of Saturn, Uranus, and Neptune,  
753 ApJ, 215, 977  
754  
755 Chandler, M. O., & Waite, J. H. 1986, The ionosphere of Uranus - A myriad of possibilities, Geophys Res Lett, 13,  
756 6  
757  
758 Connerney, J. E. P., Acuna, M., Ness, H., & Norman F. 1991, The magnetic field of Neptune, J Geophy Res, 96,  
759 19023  
760  
761 Cooper, G., Kimmich, N., Belisle, W. et al., 2001, Carbonaceous meteorites as a source of sugar-related organic  
762 compounds for the early Earth, Nature, 414, 879, doi: 10.1038/414879a  
763  
764 Encrenaz, T. 2008, Water in the Solar System, Annu Rev Astron Astr 46, 57, doi:  
765 10.1146/annurev.astro.46.060407.145229  
766  
767 Fegley, B., & Prinn, R. G. 1988, Chemical constraints on the water and total oxygen abundances in the deep  
768 atmosphere of Jupiter, ApJ, 324, 621, doi: 10.1086/165922  
769  
770 Fischer, G., & Gurnett, D. A. 2011, The search for Titan lightning radio emissions, Geophys Res Lett, 38, L08206,  
771 doi: 10.1029/2011GL047316  
772  
773 Fouchet, T., Bézard, B., & Encrenaz, T. 2005, The Planets and Titan Observed by ISO, Space Sci Rev, 119, 123,  
774 doi: 10.1007/s11214-005-8061-2  
775  
776 Fulchignoni, M., Ferri F., Angrilli F. et al. 2005, In situ measurements of the physical characteristics of Titan's  
777 environment, Nature, 438, 785  
778

779 Galejs, J. 1972, *Terrestrial Propagation of Long Electromagnetic Waves* (New York, NY: Pergamon)  
780

781 Grard, R., Hamelin, M., López-Moreno, J. J. et al. 2006, Electric properties and related physical characteristics of  
782 the atmosphere and surface of Titan, *Planet Space Sci*, 54, 1124  
783

784 Greifinger, C., & Greifinger, P. 1978, Approximate method for determining ELF eigenvalues in the Earth-  
785 ionosphere waveguide, *Radio Sci*, 13, 831  
786

787 Grimalsky, V., Koshevaya, S., Kotsarenko, A., & Enriquez, R. P. 2005, Penetration of the electric and magnetic  
788 field components of Schumann resonances into the ionosphere, *Ann Geophys*, 23, 2559  
789

790 Guillot, T. 2005, The interiors of giant planets: models and outstanding questions, *Annu Rev Earth Pl Sc*, 33, 493,  
791 doi:10.1146/annurev.earth.32.101802.120325  
792

793 Gurnett, D. A., Morgan, D. D., Granroth, L. J. et al. 2010, Non-detection of impulsive radio signals from lightning  
794 in Martian dust storms using the radar receiver on the Mars Express spacecraft, *Geophys Res Lett*, 37, L17802, doi:  
795 10.1029/2010GL044368  
796

797 Hamelin, M., Béghin, C., Grard, R. et al. 2007, Electron conductivity and density profiles derived from the mutual  
798 impedance probe measurements performed during the descent of Huygens through the atmosphere of Titan, *Planet  
799 Space Sci*, 55, 1964  
800

801 Hansell, S. A., Wells, W. K., & Hunten, D. M. 1995, Optical detection of lightning on Venus, *Icarus*, 117, 345,  
802 doi:10.1006/icar.1995.1160  
803

804 Hofe, R. 2005, *Signal Analysis of the Electric and Acoustic Field Measurements by the Huygens Instrument  
805 HASI/PWA*, Diploma Thesis, Graz University of Technology, Graz, Austria  
806

807 Kallenbach, R., Encrenaz, T., Geiss, J. et al. 2003, *Solar System History from Isotopic Signatures of Volatile  
808 Elements*, Space Sciences Series of ISSI, vol. 16, Reprinted from *Space Science Reviews*, vol. 106/1-4, 444 p.,  
809 ISBN: 978-1-4020-1177-1  
810

811 Krasnopol'sky, V. A., 1980, Lightning on Venus according to information obtained by the satellites Venera 9 and  
812 10, *Kosm Issled*, 18, 429  
813

814 Lewis, J. S. 1995, *Physics and Chemistry of the Solar System* (San Diego, CA: Academic Press)  
815

816 Lindal, G. F. 1992, The atmosphere of Neptune – an analysis of radio occultation data acquired with Voyager 2, *AJ*,  
817 103, 967  
818

819 Lindal, G. F., Lyons, J. R., Sweetnam, D. N., Eshleman, V. R., & Hinson, D. P. 1987, The atmosphere of Uranus –  
820 results of radio occultation measurements with Voyager 2, *J Geophys Res*, 92, 14987  
821

822 Liu, J. 2006, Interaction of magnetic field and flow in the outer shells of giant planets. Ph.D. thesis, Caltech,  
823 California  
824

825 Liu, J., Goldreich, P. M., Stevenson, D. J. 2008, Constraints on deep-seated zonal winds inside Jupiter and Saturn,  
826 *Icarus*, 196, 653, doi: 10.1016/j.icarus.2007.11.036  
827

828 Lodders, K. 2004, Jupiter formed with more tar than ice, *ApJ*, 611, 587  
829

830 Lodders, K., & Fegley, B. 1994, The origin of carbon-monoxide in Neptune atmosphere, *Icarus*, 112, 368, doi:  
831 10.1006/icar.1994.1190  
832

833 López-Moreno, J. J., Molina-Cuberos, G. J., Hamelin, M. et al. 2009, Structure of Titan's low altitude ionized layer  
834 from the Relaxation Probe onboard Huygens, *Geophys Res Lett*, 35, L22104, doi: 10.1029/2008GL035338  
835

836 Lunine, J. I., & Stevenson, D. J. 1987, Clathrate and ammonia hydrates at high pressure-application to the origin of  
837 methane on Titan, *Icarus*, 70, 61  
838

839 Madden, T., & Thompson, W. 1965, Low-frequency electromagnetic oscillations of earth-ionosphere cavity, *Rev*  
840 *Geophys*, 3, 211  
841

842 Mahaffy, P. R., Niemann, H. B., Alpert, A. et al. 2000, Noble gas abundance and isotope ratios in the atmosphere of  
843 Jupiter from the Galileo Probe Mass Spectrometer, *J Geophys Res*, 105, 15061  
844

845 Matousek, S. 2007, The Juno New Frontiers mission, *Acta Astronaut*, 61, 932, 10.1016/j.actaastro.2006.12.013  
846

847 Molina-Cuberos, G. J., Morente, J. A., Besser, B. P. et al. 2006, Schumann resonances as a tool to study the lower  
848 ionospheric structure of Mars, *Radio Sci*, 41, RS1003, doi: 10.1007/s11214-008-9340-5  
849

850 Mushtak, V. C., & Williams, E.R. 2002, ELF propagation parameters for uniform models of the Earth-ionosphere  
851 cavity, *J Atmos Solar-Terr Phys*, 64(18), 1989, doi: 10.1016/S1364-6826(02)00222-5  
852

853 Nellis, W. J., Weir, S. T., & Mitchell, A. C. 1996, Metallization and electrical conductivity of fluid hydrogen in  
854 Jupiter, *Science*, 273, 936  
855

856 Ness, N. F., Acuna, M. H., Behannon, K. W. et al. 1986, Magnetic fields at Uranus, *Science*, 233, 85,  
857 doi:10.1126/science.233.4759.85  
858

859 Nickolaenko, A. P., & Hayakawa, M. 2002, Resonances in the Earth-ionosphere cavity (Dordrecht, Netherlands:  
860 Kluwer Academic)  
861

862 Pechony, O., & Price, C. 2004, Schumann resonance parameters calculated with a partially uniform knee model on  
863 Earth, Venus, Mars, and Titan, *Radio Sci*, 39, RS5007, doi:10.1029/2004RS003056  
864

865 Pfaff, R. F. 1996, in *Modern Ionospheric Science*, ed. H. Kohl et al. (Berlin: Bauer), 459: ISBN 3-9804862-1-4  
866

867 Pfaff, R. F., Rowland, D., Freudenreich, H. et al. 2010, Observations of DC electric fields in the low-latitude  
868 ionosphere and their variations with local time, longitude, and plasma density during extreme solar minimum, *J.*  
869 *Geophys. Res. Space*, 115, A12324, doi:10.1029/2010JA016023  
870

871 Ruf, C., Renno, N. O., Kok, J. F. et al. 2009, Emission of non-thermal microwave radiation by a martian dust storm,  
872 *Geophys Res Lett*, 36, L13202, doi:10.1029/2009GL038715  
873

874 Russell, C. T., Strangeway, R. J., Daniels, J. T. M., Zhang, T. L., Wei, H. Y. 2010, Venus lightning: comparison  
875 with terrestrial lightning, *Planet Space Sci*, 59, 965, doi: 10.1016/j.pss.2010.02.010  
876

877 Schumann, W. O. 1952, On the free oscillations of a conducting sphere which is surrounded by an air layer and an  
878 ionosphere shell (in German), *Z Naturforsch A*, 7, 149  
879

880 Sentman, D. D. 1990a, Approximate Schumann resonance parameters for a two scale-height ionosphere, *J Atmos*  
881 *Terr Phys*, 52, 35  
882

883 Sentman, D. D. 1990b, Electrical conductivity of Jupiter's shallow interior and the formation of a resonant planetary-  
884 ionospheric cavity, *Icarus*, 88, 73  
885

886 Sentman, D. D. 1995, in *Handbook of Atmospheric Electrodynamics*, ed. H. Volland, (Boca Raton, Florida: CRC  
887 Press), pp. 267  
888

889 Simões, F. 2007, Theoretical and experimental studies of electromagnetic resonances in the ionospheric cavities of  
890 planets and satellites; instrument and mission perspectives, Ph.D. thesis, 283 pp., Univ. Pierre et Marie Curie, Paris  
891

892 Simões, F., Berthelier, J. J., Godefroy, M., & Yahi, S. 2009, Observation and modeling of the Earth-ionosphere  
893 cavity electromagnetic transverse resonance and variation of the D-region electron density near sunset, *Geophys Res*  
894 *Lett*, 36, L14816, doi:10.1029/2009GL039286  
895

896 Simões, F., Grard R., Hamelin, M. et al. 2007, A new numerical model for the simulation of ELF wave propagation  
897 and the computation of eigenmodes in the atmosphere of Titan: did Huygens observe any Schumann resonance?,  
898 *Planet Space Sci*, 55, 1978  
899

900 Simões, F., Grard, R., Hamelin, M. et al. 2008b, The Schumann resonance: a tool for exploring the atmospheric  
901 environment and the subsurface of the planets and their satellites, *Icarus*, 194, 30  
902

903 Simões, F., Hamelin, M., Grard, R. et al. 2008c, Electromagnetic wave propagation in the surface-ionosphere cavity  
904 of Venus, *J Geophys Res*, 113, E07007, doi:10.1029/2007JE003045  
905

906 Simões, F., Pfaff, R. F., Berthelier, J.-J., & Klenzing, J. 2011b, A review of low frequency electromagnetic wave  
907 phenomena related to tropospheric-ionospheric coupling mechanisms, *Space Sci Rev*, doi: 10.1007/s11214-011-  
908 9854-0 (in press)  
909

910 Simões, F., Pfaff, R. F., & Freudenreich, H., 2011a, Observation of Schumann resonances in the Earth's ionosphere,  
911 *Geophys Res Lett*, 38, L22101, doi:10.1029/2011GL049668  
912

913 Simões, F., Rycroft, M., Renno, N. et al. 2008a, Schumann resonances as a means of investigating the  
914 electromagnetic environment in the Solar System, *Space Sci Rev*, 137, 455  
915

916 Stevenson, D.J., & Lunine, J. I. 1988, Rapid formation of Jupiter by diffusive redistribution of water-vapor in the  
917 solar nebula, *Icarus*, 75, 146, doi: 10.1016/0019-1035(88)90133-9  
918

919 Tobie, G., Grasset, O., Lunine, J. I., Mocquet, A., & Sotin, C. 2005, Titan's internal structure inferred from a  
920 coupled thermal-orbital model, *Icarus*, 175, 496  
921

922 Tsiganis, K., Gomes, R., Morbidelli, A., & Levison, H. F. 2005, Origin of the orbital architecture of the giant planets  
923 of the Solar System, *Nature*, 435, 459, doi:10.1038/nature03539  
924

925 Tyler, G. L., et al. 1989, Voyager radio science observations of Neptune and Triton, *Science* 246, 1466-1473

926

927 Yair, Y., Fischer, G., Simões, F., Renno, N., & Zarka, P. 2008, Updated review of planetary atmospheric electricity,  
928 Space Sci Rev, 137, 29

929

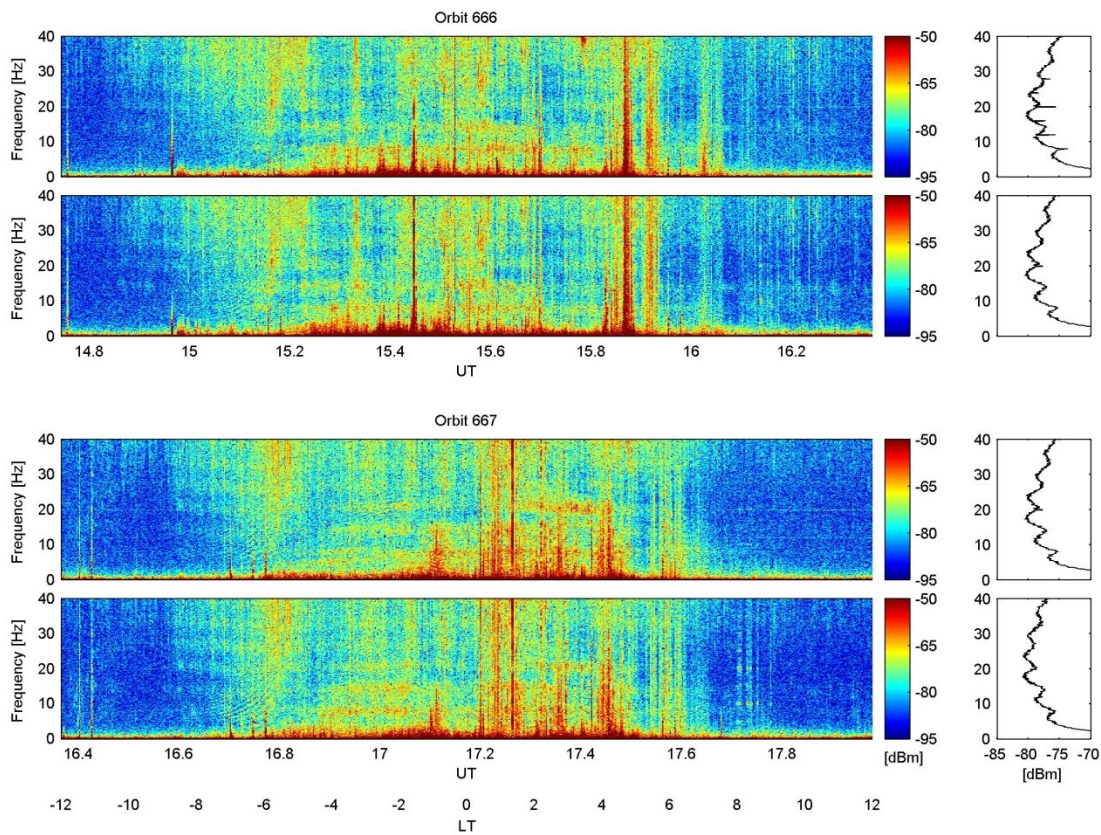
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## Figures

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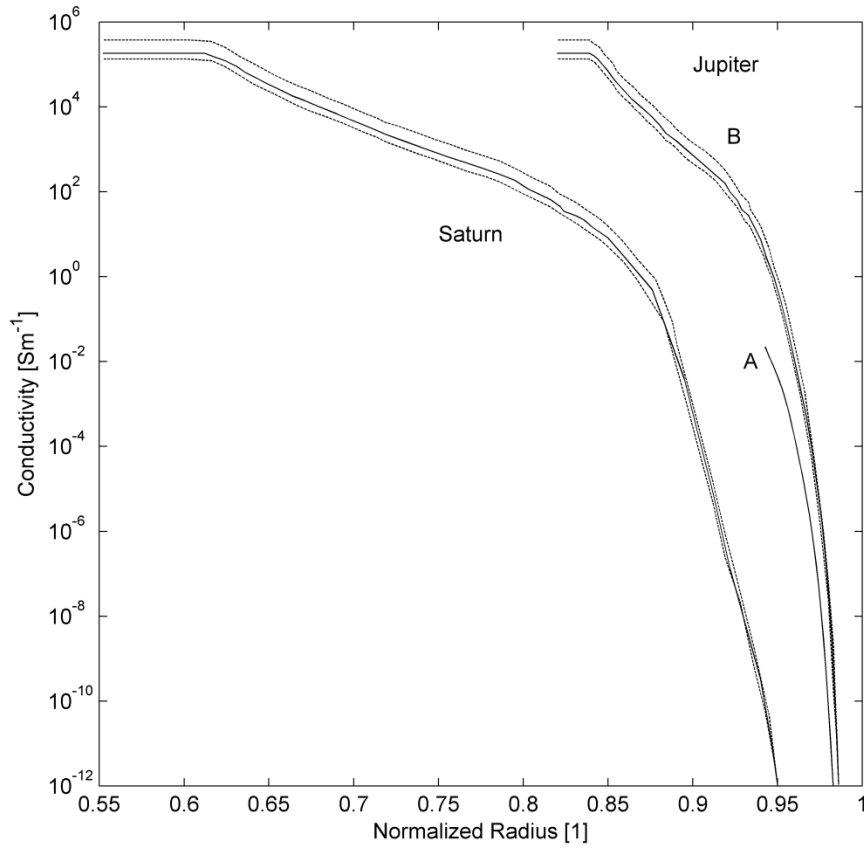


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936 Figure 1: VEFI electric field data recorded on 2008 May 31 during orbits 666 and 667 (top and  
937 bottom panels). (left) Spectrogram and (right) mean spectrum computed all through the orbit.  
938 The top and bottom panel refer to meridional and zonal components, respectively. The fuzzy  
939 horizontal lines seen mostly during nighttime in the left panels and the spectral peaks on the  
940 right-hand-side correspond to Schumann resonance eigenmodes.

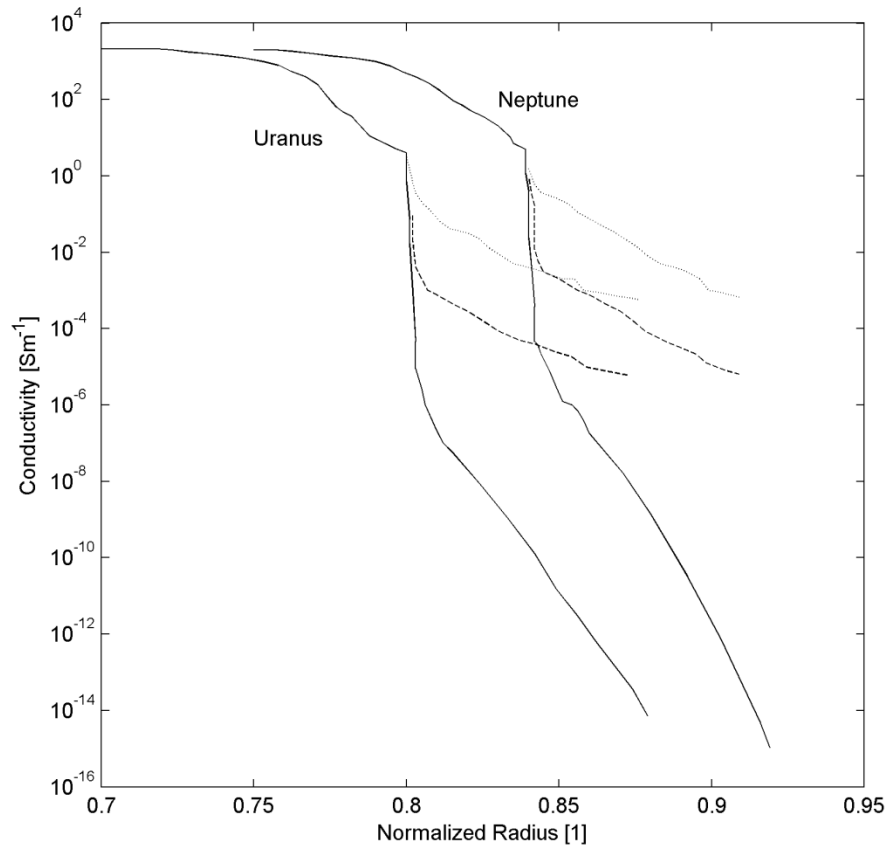
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Figure 2: Conductivity profile of Jupiter and Saturn as a function of normalized radius. The solid and dashed lines represent the mean and uncertainty envelope of the conductivity. The profiles Jupiter-B and Saturn are adapted from Liu (2006). The profile Jupiter-A is taken from Sentman (1990b).

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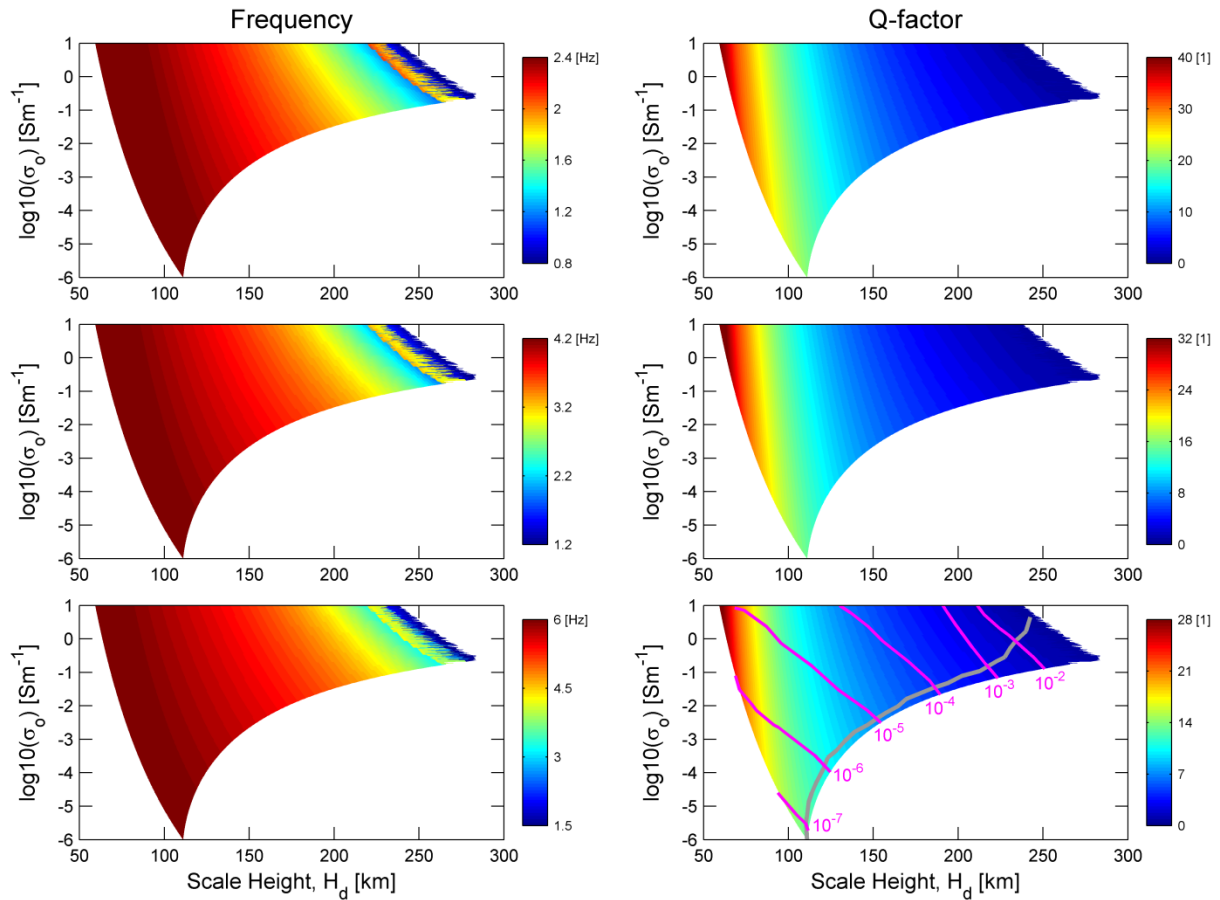


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Figure 3: Conductivity profile of Uranus and Neptune as a function of normalized radius. The solid, dashed, and dotted lines correspond to 0, 0.01, and 0.1 water content, respectively. The conductivity profiles are adapted from Liu (2006) and Liu et al. (2008).



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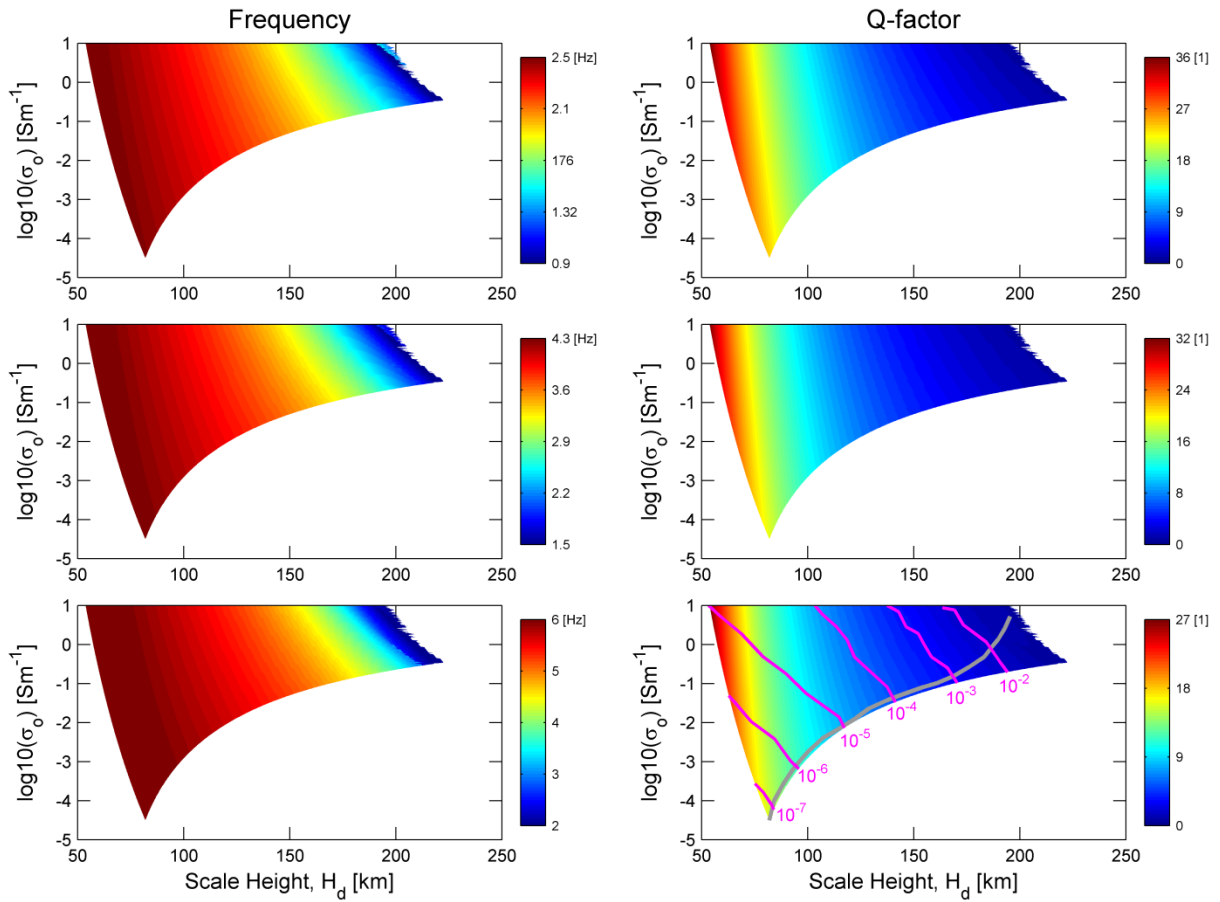


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977 Figure 4: Modeling results of (left) eigenfrequencies and (right) Q-factors of (from top) the three  
978 lowest eigenmodes as function of interface conductivity ( $\sigma_0$ ) and scale height ( $H_d$ ) of the Uranus  
979 cavity. In the bottom-right panel, the magenta curves represent mean water contents in the  
980 gaseous envelope; cavity parameterizations near the gray curve represent the most plausible  
981 conductivity profiles.

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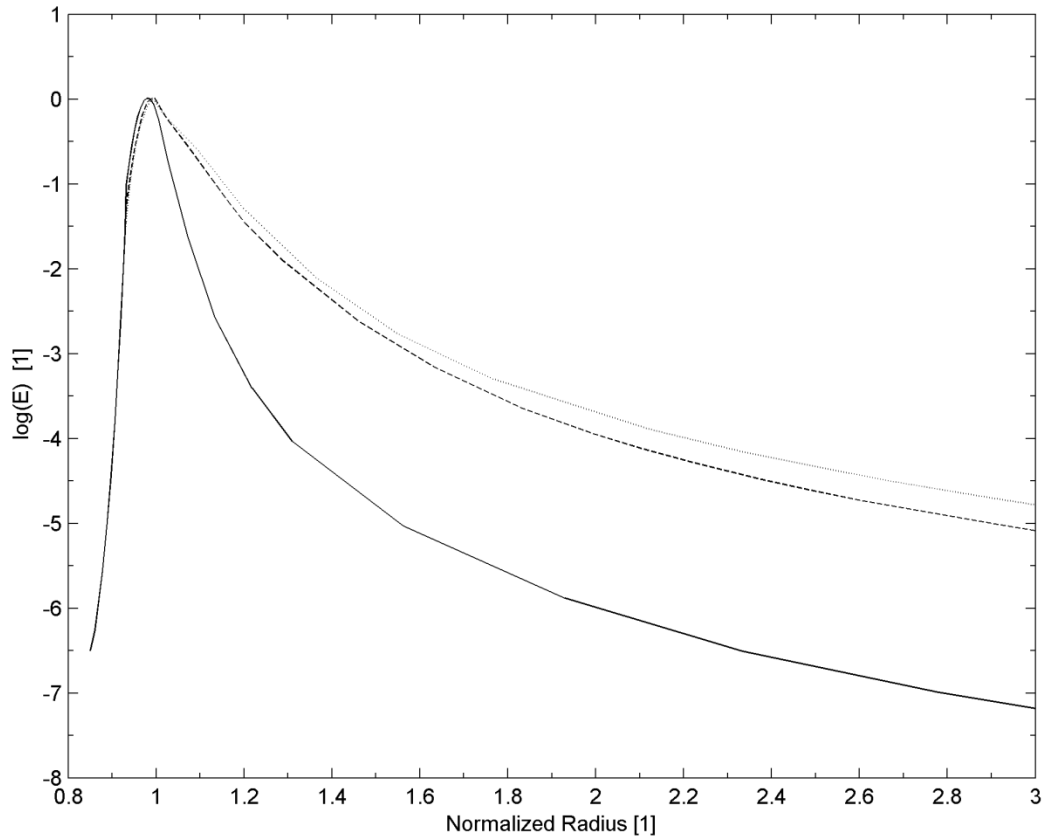
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Figure 5: Same caption as in Figure 4 but for the cavity of Neptune.

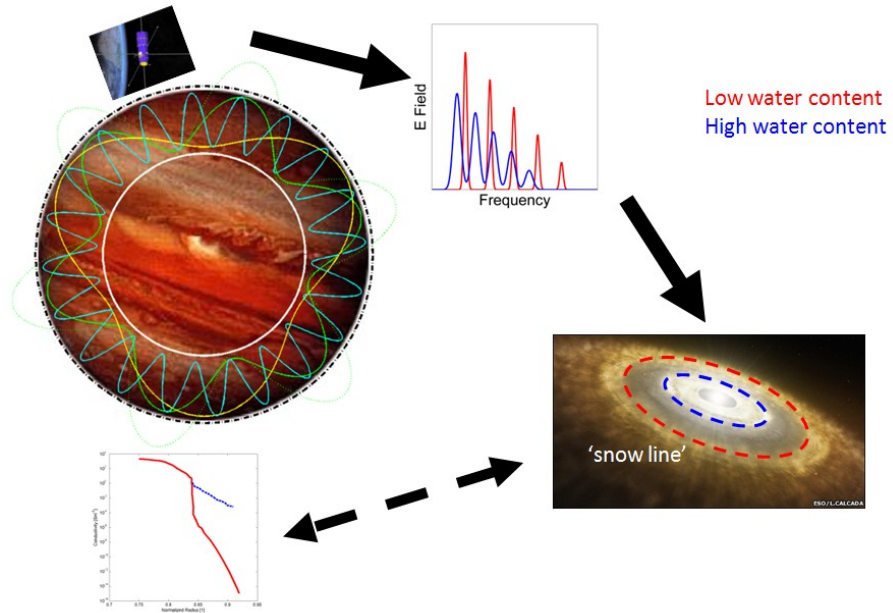
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Figure 6: Normalized electric field amplitude of the lowest Schumann eigenmode as function of distance to planet center for (dash) Uranus, (dot) Neptune, and (solid) Earth ionosphere equivalent configuration.

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Figure 7: Illustration of the connection between the water mixing ratio, conductivity profile, Schumann resonance spectrum, cavity leakage, and water ice condensation front of the Solar System. (--- Artwork to be improved ---)

## Tables

Title: Planetary cavities characteristics

Planet	$R$ [km]	$(R-d)/R$ [1]	$B$ [ $\mu\text{T}$ ]	$h$ [km]	Envelope Composition
Earth	6378	1	32	100	0.78 N <sub>2</sub> ; 0.21 O <sub>2</sub> ; 0.01 Ar (water vapor up to 1%)
Jupiter	71,493	0.84	420	900	0.82 H <sub>2</sub> ; 0.18 He (ice traces)
Saturn	60,268	0.63	20	600	0.94 H <sub>2</sub> ; 0.06 He (ice traces)
Uranus	25,559	0.8	23	600 (?)	0.74 H <sub>2</sub> ; 0.26 He (ice up to 10%)
Neptune	24,764	0.84	14	400 (?)	0.68 H <sub>2</sub> ; 0.32 He (ice up to 10%)

Table 1: Comparison between Earth and giant planets cavity characteristics. The radius and magnetic field refer to equatorial values; since the magnetic field of Uranus and Neptune is irregular and strongly inclined with respect to the rotation axis, values are merely indicative because, locally, they can be 5-fold higher (Ness et al. 1986; Connerney et al. 1991). The third column shows where the gaseous envelope may collapse into a solid or a liquid, and gives an indication to defining the inner boundary of the cavity. The altitude,  $h$ , is evaluated for  $\sigma \sim 10^{-3} \text{ Sm}^{-1}$ , corresponding to skin depths of  $\sim 5$  and 10-15 km for Earth and the giant planets, respectively (cf. Simões et al. 2008b). The composition of the envelope is selected from Lewis (1995) and Liu (2006) and merely indicative due to significant uncertainty. However, further explanation is useful: on Earth, when water vapor is included, composition exceeds one hundred percent – in general, composition refers to a dry atmosphere; on the giant planets, impurities contribution is unknown and composition refers to the expected envelope mean composition rather than atmospheric values.

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1060 Title: Planetary cavities computed eigenfrequencies

Conductivity Profile	Mode	Planet								Reference
		Jupiter		Saturn		Uranus		Neptune		
		$f$ [Hz]	Q [1]	$f$ [Hz]	Q [1]	$f$ [Hz]	Q [1]	$f$ [Hz]	Q [1]	
Minimum	1	0.736	7.022	0.772	3.672	2.429	19.01	2.521	15.46	Liu (2006)
	2	1.308	6.816	1.401	3.705	4.245	17.54	4.416	15.52	
	3	1.885	6.844	2.039	3.707	6.036	16.34	6.293	15.66	
Intermediate	1	0.734	7.896	0.763	3.597	1.025	4.08	1.109	2.04	
	2	1.296	7.913	1.386	3.252	1.992	4.06	2.030	2.11	
	3	1.855	7.717	2.030	3.037	3.037	4.93	2.961	1.75	
Maximum	1	0.752	9.791	0.767	4.062	×	×	×	×	
	2	1.318	10.85	1.381	4.275	×	×	×	×	
	3	1.878	10.82	1.997	4.278	×	×	×	×	
Equatorial radius 71,500 km	1	0.575	5.202	-	-	-	-	-	-	Sentman (1990b)
	2	1.017	5.938	-	-	-	-	-	-	
	3	1.456	6.539	-	-	-	-	-	-	
Mean radius 69,900 km	1	0.584	5.038	-	-	-	-	-	-	
	2	1.040	5.625	-	-	-	-	-	-	
	3	1.495	6.218	-	-	-	-	-	-	
Polar 66,850 km	1	0.616	5.047	-	-	-	-	-	-	
	2	1.100	5.234	-	-	-	-	-	-	
	3	1.588	5.318	-	-	-	-	-	-	

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1062 Table 2: Lowest eigenfrequencies of the giant planets for the conductivity profiles shown in  
1063 Figures 2-3. The maximum conductivity profiles of Uranus and Neptune, corresponding to a  
1064 water mixing ratio 0.1, prevent formation of resonant modes.