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SOME ASPECTS OF WIND SHEAR IN THE UPPER ATMOSPHERE

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GODDARD SPACE FLIGHT CENTER

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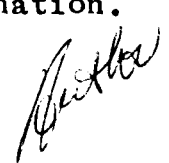
Some Aspects of Wind Shear in the
Upper Atmosphere

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R. G. Roper

ABSTRACT

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The wind motions responsible for the shearing of sodium vapor trails ejected from rockets in the 70 to 140 km region of the upper atmosphere are subjected to an analysis based on generally accepted theories of hydrodynamic turbulence. The region from 80 to 100 km is of particular interest in that here the predictions of shear turbulence theory are well substantiated. The energy spectrum of the height shear is found to follow the $4/3$ power law proposed by Tchen, and is associated with a vertical correlation distance of approximately 6 km. The existence of an isotropic inertial region of maximum scale 3 km, previously indicated by analysis of meteor data, is confirmed. The vertical scale of the turbulent eddies is found to be the atmospheric pressure scale height, a phenomenon which has been observed by others, but which, as yet, has no satisfactory explanation.



1. Turbulence Theory

The complete development of the relationships used in the analysis of wind shear is beyond the scope of this work; the following, used in conjunction with the references quoted, should provide an adequate background for consideration of the subsequent analysis.

1.1 Energy Spectrum Analysis

If there exists in a turbulent flow field a range of scales which receive energy from larger scale motions and pass it on undiminished to smaller scale motions, then, for this so-called inertial (nondissipative) range of scales, the only form of energy spectrum function dimensionally possible is

$$E(k) = \alpha \epsilon^{2/3} k^{-5/3} \quad (\text{Kolmogoroff, 1941})$$

where ϵ is the rate at which the turbulent energy is received by (and leaves) the inertial range of scales, k is the wavenumber vector corresponding to the real space scale r , and α is an absolute constant of order unity.

Batchelor (1953) has shown that, for such an inertial region which also possesses the property of isotropy, the fluid velocity differences measured as a function of the separation r follow the relation

$$[\overline{u(x) - u(x + r)}]^2 = 4.82\alpha(\epsilon r)^{2/3} \dots\dots\dots 1$$

In real space, the energy spectrum function E(r) is defined by

$$E(r) = \overline{[u(x) - u(x + r)]^2} \dots\dots\dots 2$$

Tchen (1954) has considered an otherwise isotropic region subjected to a mean shear, and finds that equation 1 is modified, becoming

$$[u(x) - u(x + r)]^2 = a r^{4/3} \dots\dots\dots 3$$

where a involves α , ϵ , and the mean gradient. Thus, for what may be termed shear turbulence, in real space

$$E(r) \sim r^{4/3} \dots\dots\dots 4$$

1.2 Correlation Analysis

An energy spectrum function equivalent to that based on velocity differences can be formulated from the lateral or longitudinal velocity correlations defined as

$$f(r) = \frac{\overline{u_f(x) u_f(x + r)}}{u_f^2}$$

and

$$g(r) = \frac{\overline{u_n(x) u_n(x+r)}}{u_n^2} \dots\dots\dots 5$$

where $u_f(x)$, $u_f(x+r)$, $u_n(x)$, $u_n(x+r)$ are the turbulent components of the velocity at two points x and $x+r$ respectively, measured parallel (suffix f) and normal (suffix n) to the vector separation r . In isotropic turbulence, $u_f^2 = u_n^2 = u_o^2$, the velocity characteristic of the energy bearing eddies of the Kolmogoroff spectrum.

Introduction to the equation of continuity for incompressible fluids leads to the relation

$$g(r) = f(r) + 1/2 r \frac{\partial f}{\partial r} \dots\dots\dots 6$$

(von Karman and Howarth, 1938)

between the functions f and g , or, if the turbulence is isotropic in two dimensions only,

$$g(r) = f(r) + r \frac{\partial f}{\partial r} \dots\dots\dots 7$$

These functions, f and g , may be called Eulerian space correlation functions, and either may be denoted by $R(r)$. If the range of scales under observation is inertial, then $R(r)$ must depend only on ϵ , and the only form dimensionally possible is

$$u_0^2 [1 - R(r)] \sim \epsilon^{2/3} r^{2/3}$$

i.e. $1 - R(r) \sim r^{2/3}$

(for any given flow field, u_0 and ϵ are constant).

Introduction of the relations 6 and 7 above gives,

with c a constant

$$f = 1 - cr^{2/3} \dots\dots\dots 8$$

and $g = 1 - 4/3 cr^{2/3} \dots\dots\dots 9$

for three dimensional isotropy

or $f = 1 - 5/3 cr^{2/3} \dots\dots\dots 10$

for isotropy in two dimensions only.

Since the variations of the correlation difference function $[1-R(r)]$ with r is the same as that of $E(r)$ in equation 1, we may suppose that the dependence of $[1-R(r)]$ on r will be similarly modified in the presence of a mean shear

i.e. $[1-R(r)] \sim r^{4/3} \dots\dots\dots 11$

If, in fact, equation 11 is pertinent, then the relationships between f , g , and r (equations 8, 9, 10) will now become

$$f = 1 - ar^{4/3}$$

and $g = 1 - 5/3 ar^{4/3}$

or $g = 1 - 7/3 ar^{4/3}$

An indication of the degree of isotropy can be obtained by considering the ratio

$$S = \frac{1 - f}{1 - g}$$

For two dimensional isotropy

S = 0.60 without mean shear

and S = 0.43 with mean shear

For three dimensional isotropy

S = 0.75 without mean shear

and S = 0.60 with mean shear

The relative importance of mean shear will be indicated by the form of either the energy spectrum function $E(r)$, or the correlation difference function $[1 - R(r)]$.

2. The Practical Application of Turbulence Theory

2.1 The Mean Wind Profile

In applying the relations developed above to the wind vectors measured by means of sodium vapor trails ejected as a tracer into the upper atmosphere, the relative importance of the mean motion must not be overlooked. The velocities used in the energy spectrum and correlation analyses of the previous section must be the turbulent velocities, or departures from the mean motion. In normal correlation analysis, the mean of a set of observations is usually taken as the mean value of the measured data, which, when applied to wind height shear data, would be tantamount to the assumption of a mean wind profile which is

constant with height. Such a profile is the exception rather than the rule in meteorological phenomena.

Practically any attempt to prescribe a mean wind profile to the measured winds will be subjective to a certain extent. If one uses a polynomial fit, for example, it must be truncated before it can assume any of the features of the measured profile which are due to the turbulent motions present. From experience based on the measurement of winds in the 75 to 105 km region by means of radio reflections from meteor trails (Elford, 1958, 1964), a quadratic change with height of the mean wind over any given 20 km interval should best describe the contribution of the mean motion without destroying any of the characteristics of the turbulent flow field. In the present analysis, the windspeed/azimuth data are converted to zonal and meridional components, and a polynomial profile of order $Z + 1$, where

$$Z = \frac{\text{Total height range covered by data(km)}}{20} ,$$

is fitted to each. The profiles thus determined are subtracted from the relevant measured profiles to give zonal and meridional turbulent velocities.

2.2 The Determination of E(Δh)

The available sodium trail data lists wind speed and azimuth against height, and usually involves irregularly spaced height intervals. The sampling irregularity exists for two reasons:

- a) there is an occasional difficulty in absolutely identifying the same point on the trail in consecutive photographs;
- b) the wind profile between consecutive observational heights is linear. This is usually obvious from the photographs, and can easily be allowed for in subsequent analysis.

Whereas equal height interval sampling is not absolutely necessary for subsequent reduction, it does simplify the analysis, and so linear interpolation between the listed data points is used to provide a profile with data points spaced 0.2 km in height.

The energy spectrum function is computed as

$$E(\Delta h_i) = \frac{1}{N} \sum_{k=1}^N (u_{k+i} - u_k)^2 \dots\dots\dots 12$$

in which

$$N = 5(H_2 - H_1) - i$$

where $i = 1, 2, 3, \dots, N$

such that $\Delta h_1 = 0.2$ km.

$$\Delta h_2 = 0.4$$
 km

etc.

and H_1 , H_2 (in km) are the lower and upper bounds respectively of the region for which $E(\Delta h)$ is being determined. Such partitioning of the height range is necessary since there is considerable variation in the characteristics of the flow over the total height range sampled (usually some 70 to 200 km).

Energy spectrum functions may be calculated using

- a) zonal turbulent velocities
 - b) meridional turbulent velocities
- and
- c) turbulent windspeed.

The turbulent windspeed w is defined here as

$$w = \text{measured windspeed} - \left((\text{mean zonal wind})^2 + (\text{mean meridional wind})^2 \right)^{1/2}$$

2.4 Correlation Analysis

The two correlation functions f and g defined in 1.2 refer to turbulent velocity components measured parallel and normal to the separation vector. Since vertical velocities in this region of the upper atmosphere are so much less than the associated horizontal components, the magnitude of the vertical component cannot be determined from sodium trail photographs. (Most workers in this field consider the upper limit for mean plus random vertical

motions to be some 10 metres/sec). However, we may redefine f and g in terms of the orthogonal zonal and meridional flow fields, with a view to investigating possible isotropy. This has been done by meteorologists in the past, with at least partial success (see, for example, Hutchings (1955)).

The normalizing factors \bar{u}_f^2 and \bar{u}_n^2 are best estimated by the standard correlation function definition. The correlation functions f and g then becomes

$$f(\Delta h_i) = \frac{\sum_{k=1}^N u_{k+i} u_k}{\left[\sum_{k=1}^N u_{k+1}^2 \sum_{k=1}^N u_k^2 \right]^{1/2}} \dots\dots\dots 14$$

where u are the zonal turbulent wind velocities and

$$g(\Delta h_i) = \frac{\sum_{k=1}^N v_{k+1} v_k}{\left[\sum_{k=1}^N v_{k+i}^2 \sum_{k=1}^N v_k^2 \right]^{1/2}} \dots\dots\dots 15'$$

where v are meridional turbulent wind velocities, and N and i are as previously defined for equation 12.

3. Preliminary Results

The analysis of Section 2 has been applied to data obtained from a sodium trail release over the Eglin Air Force Base, Florida (29.6°N, 86.6°W) at 1910 CST on May 21, 1963 (Edwards et al, 1963). In this experiment, wind speed and azimuth were obtained over a height range of 69 to 140 km.

3.1 The Mean Wind Profile

Since the data covers the height range from 69 to 140 km, polynomials of order 4 are fitted to the zonal and meridional measured profiles. These yield mean zonal and meridional profiles

$$u_{\text{mean}} = 38.6 - 126h - 244h^2 + 174h^3 + 205h^4$$

$$v_{\text{mean}} = -30.4 + 19.8h + 148h^2 - 16.8h^3 - 125h^4$$

where u , v are in metres/sec and h is the normalized height given by

$$h = (2z - z_{\text{min}} - z_{\text{max}}) / (z_{\text{min}} - z_{\text{max}})$$

where z is the height variable

z_{\max} the maximum and

z_{\min} the minimum heights of the available data,

all heights being in kilometers.

Normalization of the height range stabilizes the least squares fitting process, and makes the relative importance of the individual terms of the fitted polynomials more obvious than is the case when the mean velocities are expressed as power series in the actual height z . Results are plotted in Figs.1 and 2.

The meteorological significance of these profiles, in particular the reversal of both the zonal and meridional components above 110 km, cannot be evaluated from consideration of this single firing.

3.2 The Turbulent Wind Profile

As mentioned in Section 2, the 70 to 200 km height range covers a number of characteristically different regions. The wind motions observed below approximately 105 km indicate the presence of small-scale structure, while those above 110 km do not appear to be at all turbulent, even though vertical shear is present. In the results presented here, discussion is confined to the consideration of the region from 80 to 100 km, which has been found to be representative of a turbulent region which can be adequately

described by available statistical theories of hydrodynamic turbulence (Blamont and Jager, 1961; Zimmerman, 1962; Roper, 1962).

The measured zonal and meridional profiles, and the deviations from the mean wind for the height range 80 to 100 km are plotted in Figures 3 to 6. Whereas the immediately obvious wavelike nature of the turbulent profile would suggest a wave theory approach as likely to be the most profitable for consideration of the wind motions in this region, applications of available wave theories (e.g., Hines, 1959) have not produced consistent results. The possibility of generation of turbulence in the 80 to 100 km region by vertically propagating gravity waves has been proposed by Hines (1963). The purpose of the present work, however, is not to determine the source of the turbulent energy, but rather to substantiate the evidence that the observed shears characterize a region of hydrodynamic turbulence.

3.3 The Energy Spectrum Functions

The energy spectrum functions for the zonal and meridional turbulent wind profiles for the 80 to 100 km region have been computed using the methods described in Sections 2.2 and 2.3. As can be seen from Figures 7 and 8, the velocity difference spectrum $E(\Delta h)$ and the correlation difference function $[1-f(\Delta h)]$ for the zonal turbulent velocities are completely equivalent. This is not surprising, since both functions are solutions of the same equation

$$F = a \Delta h^m$$

(see Sections 1.1 and 1.2).

Similarly for the meridional spectra of Figures 9 and 10.

The slope m of the log log plots of both the zonal and meridional $E(\Delta h)$ (or $[1-R(\Delta h)]$) spectrum functions against height difference Δh is constant at $4/3$ for separations Δh up to 3 km, indicative of an isotropic region subject to a mean wind shear (Tchen, 1954).

3.4 Isotropy

Since the energy spectrum functions and the correlation difference functions all follow an established (shear) law at small scales in the height range from 80 to 100 km, it is possible to determine the nature of the isotropy at these scales in this region.

Figure 13, curve A, is a plot of the variation of

$$S = \frac{1-f}{1-g}$$

against Δh .

In the region up to a scale of 3.5 km, S has a value of approximately 0.7, indicating an isotropy lying somewhere between two and three dimensional. The increase in S at scales greater than approximately 3 km is due to the breakdown at this scale of the $4/3$ power law in both zonal and meridional spectrum functions.

The function S has been found to be quite sensitive to variation in the parameters specifying the mean wind profiles. For example, if, instead of the fourth order polynomial fits to the total height range of the zonal and meridional data, linear mean wind profiles are fitted over the height range 80 to 100 km only, subsequent spectrum analysis yields the S function of curve B of Figure 13. This curve would indicate the existence of three dimensional isotropy for scales up to 3 km. It is possible that the 4th order polynomial fits are, in fact, attributing a small fraction of the random wind variations to the mean motion. However, in order to correlate spectra obtained for different height strata (as is done in the next section), a continuous mean profile is required, and has therefore been used throughout the analysis.

3.5 Vertical Scale

The vertical scale associated with the turbulent wind structure is conveniently defined by the Δh corresponding to the maximum value of either $E(\Delta h)$ or $[1-R(\Delta h)]$. As can be seen from Figures 7, 9 and 11, which are plotted using zonal, meridional and windspeed turbulent velocities respectively, the vertical correlation distance as defined above is not the same for each component, i.e., the wind motions are not isotropic at the maximum of $E(\Delta h)$. However, for the purpose of comparison with stratospheric and

lower mesospheric data, where meridional winds are, for the most part, negligible, consideration need only be given to the zonal energy spectrum.

The variation of vertical scale with height in the stratosphere has been determined by Webb (1964) from a series of Robin soundings at Eglin. His results indicate an exponential increase of the vertical scale with height, from approximately 800 metres at 35 km to 2 km at 55 km. He has extrapolated this exponential to a height of 90 km, and finds a vertical scale of 6 km which is the value determined for this height by Greenhow (1959) from a correlation analysis of wind shears determined by means of radio reflections from meteor trails at Jodrell Bank (53°N).

In order to investigate the change with height of the vertical correlation distance in the 80 to 130 km region, the sodium trail data was subjected to a stepwise analysis, starting with the 75 to 95 km height range, and proceeding via the 80 to 100, 85 to 105, etc. ranges to 115 to 135 km. The maxima of the resultant $E(\Delta h)$ curves were then plotted as the vertical scales at the midpoints of the respective height ranges. The results are shown in Figure 14, together with Greenhow's determination, and the 7.8 km at 94 km (October, 1961) determined from radio meteor trail shears at Adelaide (35°S) (Roper, 1962). The dashed line is Webb's extrapolated variation. While there is excellent

agreement with Webb's proposed exponential increase, there is also good agreement between the vertical scales measured at various altitudes for the Eglin firing, and the atmospheric pressure scale height (also shown in Figure 14). This phenomenon has also been observed by others (e.g., Zimmerman (1964)). As yet, no satisfactory explanation for such a dependence has been proposed.

Above 125 km, the magnitudes of the deviations of the measured winds from the mean wind values becomes insignificant. Wind motions at these heights for this particular firing are characteristically nonturbulent.

Conclusions

The techniques of spectrum analysis based on hydrodynamic turbulence theory can be profitably applied to wind data obtained from sodium trail rocket firings, at least in the height range from 80 to 100 km. The analysis of further firings should indicate whether or not the mean profiles determined as a prelude to spectrum analysis have any meteorological significance.

References

General

- Kolmogorff, A., C. R. Akad. Sci. U.R.S.S., 30, 301, 1941.
- Batchelor, G. K., The Theory of Homogeneous Turbulence,
Cambridge University Press, 1953
- Corrsin, S., Jour. Geophys. Res., 64, 2134, 1959
- Tchen, C. M., Phys. Rev., 94, 4, 1954
- Karman, T., von and Howarth, L., Proc. Roy. Soc., London
164, 192, 1938
- Elford, W. G., Plan. and Space Sci., 1, 94, 1958
- Elford, W. G., 1964 (unpublished)
- Hutchings, J. W., Jour. Met., 12, 263, 1955
- Edwards, H. D. et al., Jour. Geophys. Res. 68, 6062, 1963
- Blamont, J. E., and De Jagar, C., Ann. de Geophys., 17, 134, 1961
- Zimmerman, S. P., Ann de Geophys. 18, 116, 1962.
- Roper, R. G., Ph.D. Thesis, University of Adelaide, 1962
- Hines, C. O., Jour. Geophys. Res., 64, 2210, 1959
- Hines, C. O., Quart. Jour. Roy. Met. Soc., 89, 1, 1963
- Webb, W. L., U.S. Army E.R.D.A. - 130 , White Sands, New Mexico.
April 1964
- Greenhow, J. S., and Neufeld, E. L., Jour. Geophys. Res., 64,
2129, 1959
- Zimmerman, S. P., Jour. Geophys. Res., 65, 784, 1964.

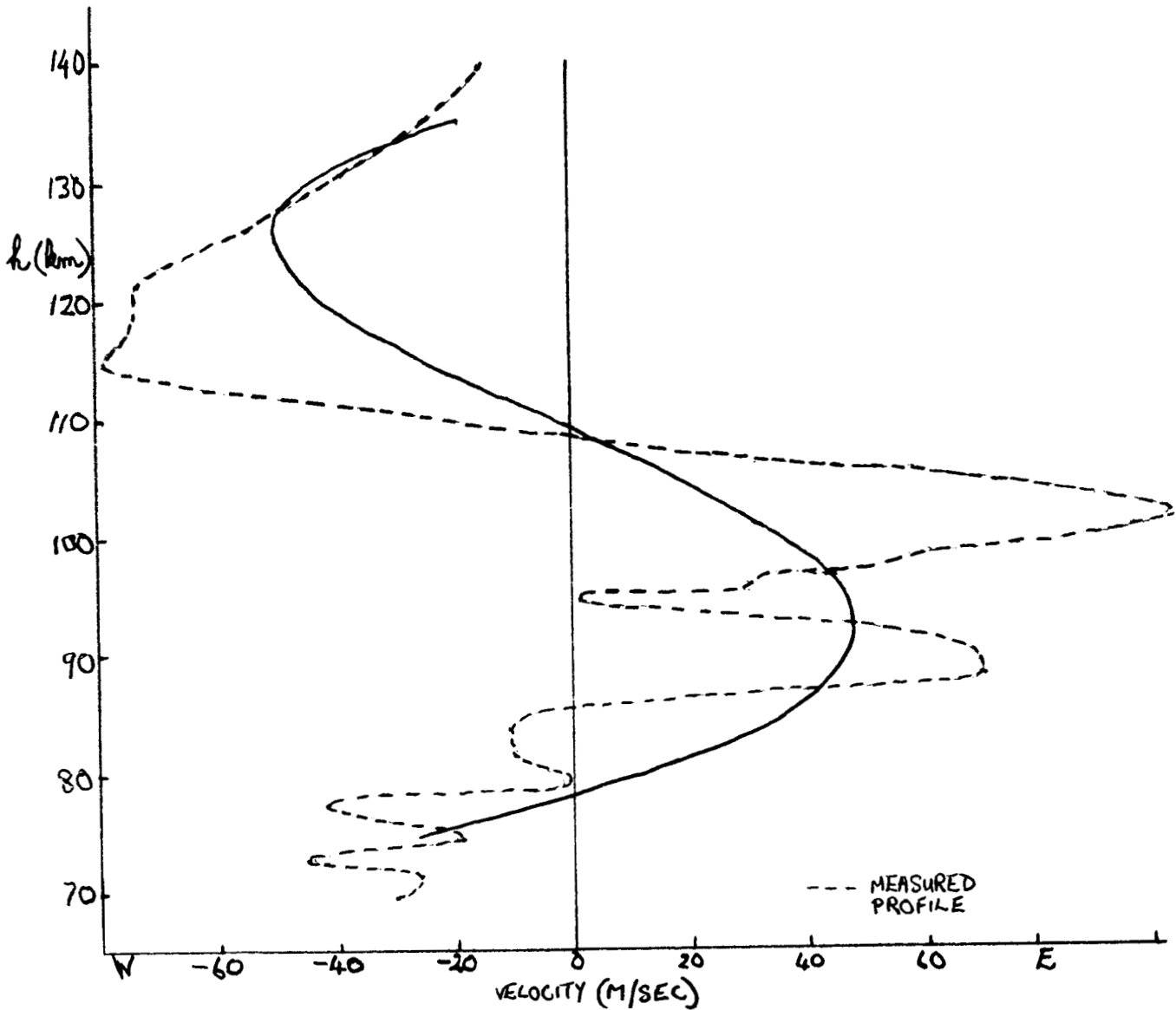


FIG 1. MEAN ZONAL PROFILE

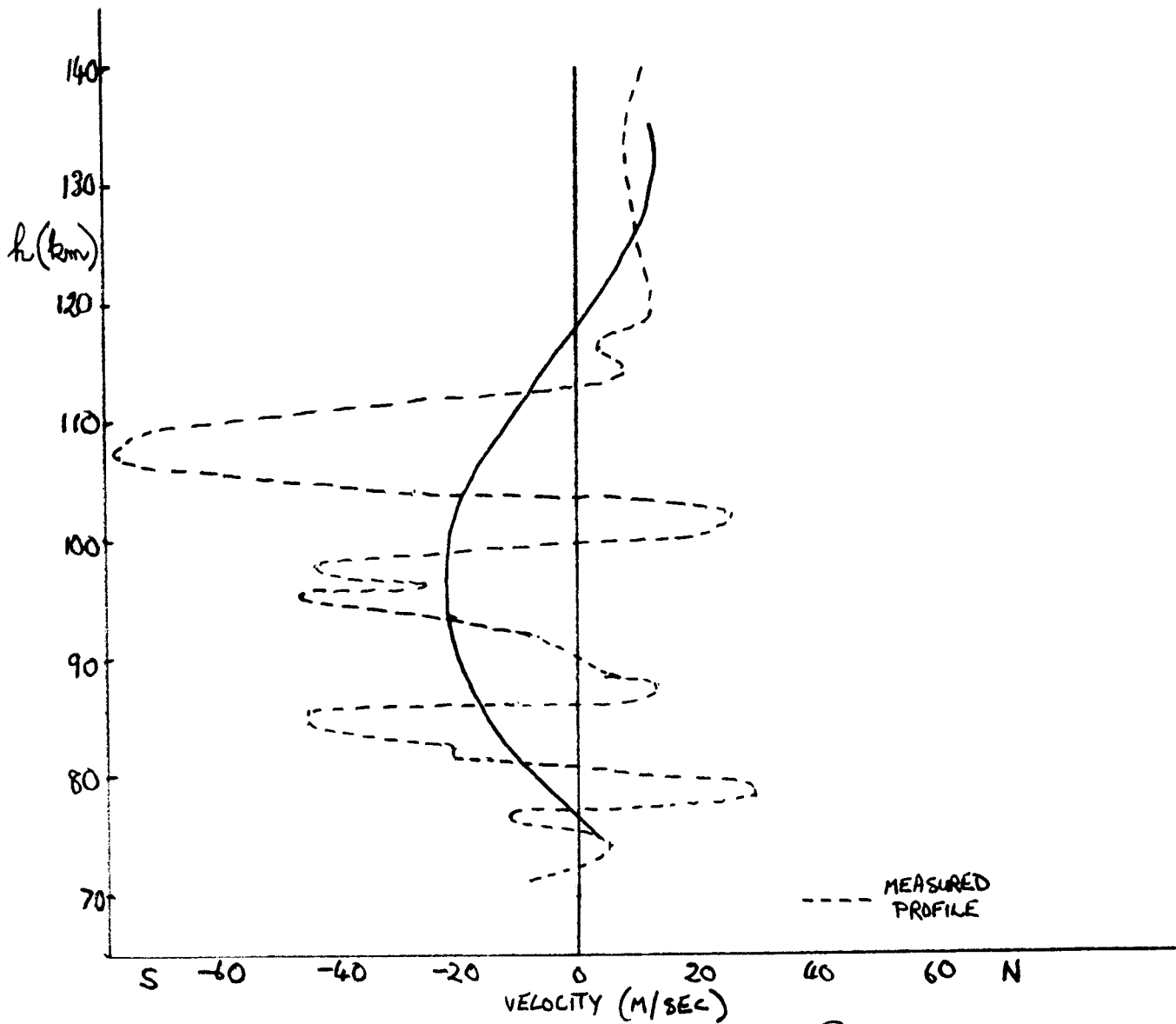


FIG. 2. MEAN MERIDIONAL PROFILE

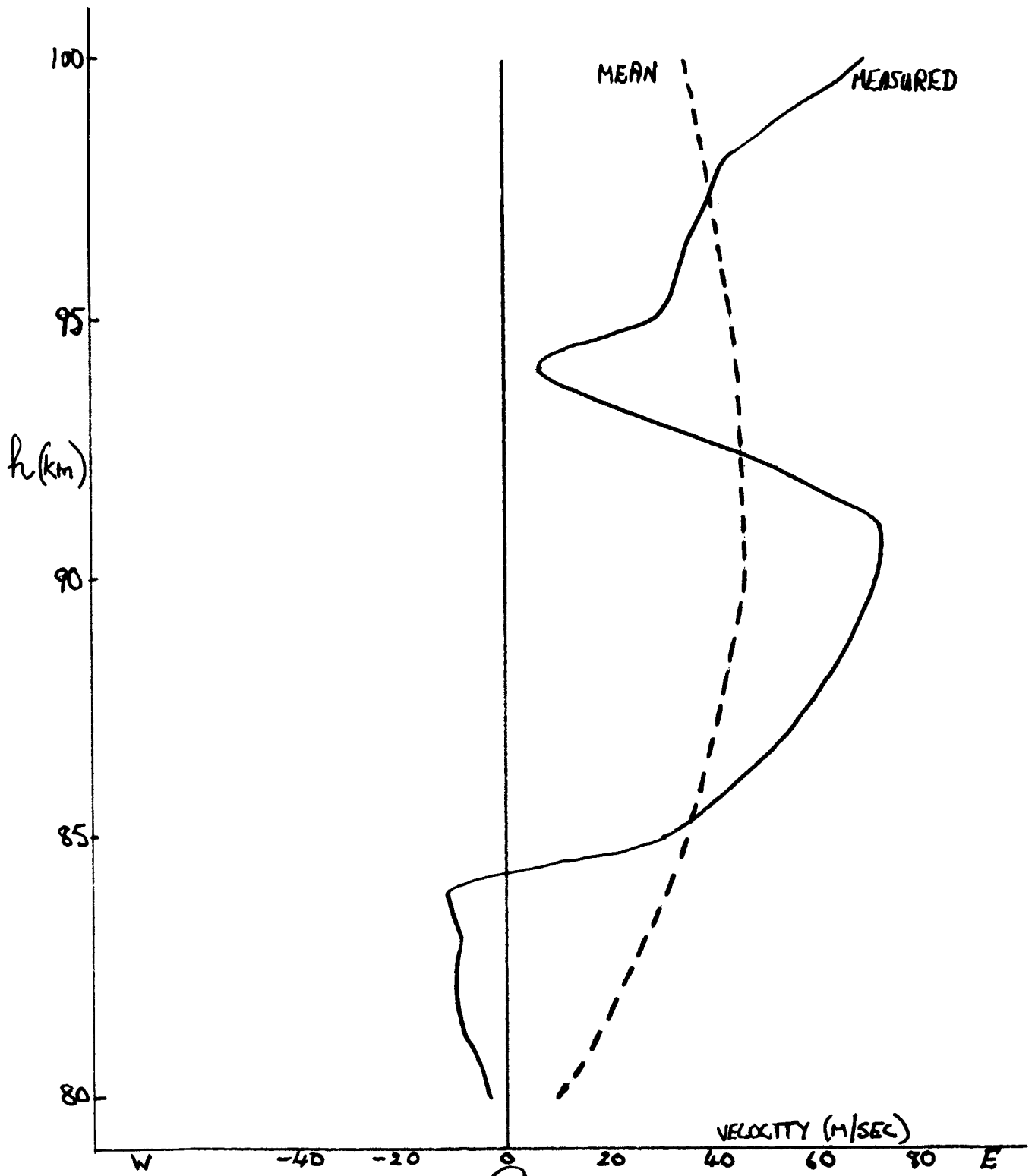


FIG. 3. ZONAL PROFILE, 80-100 KM

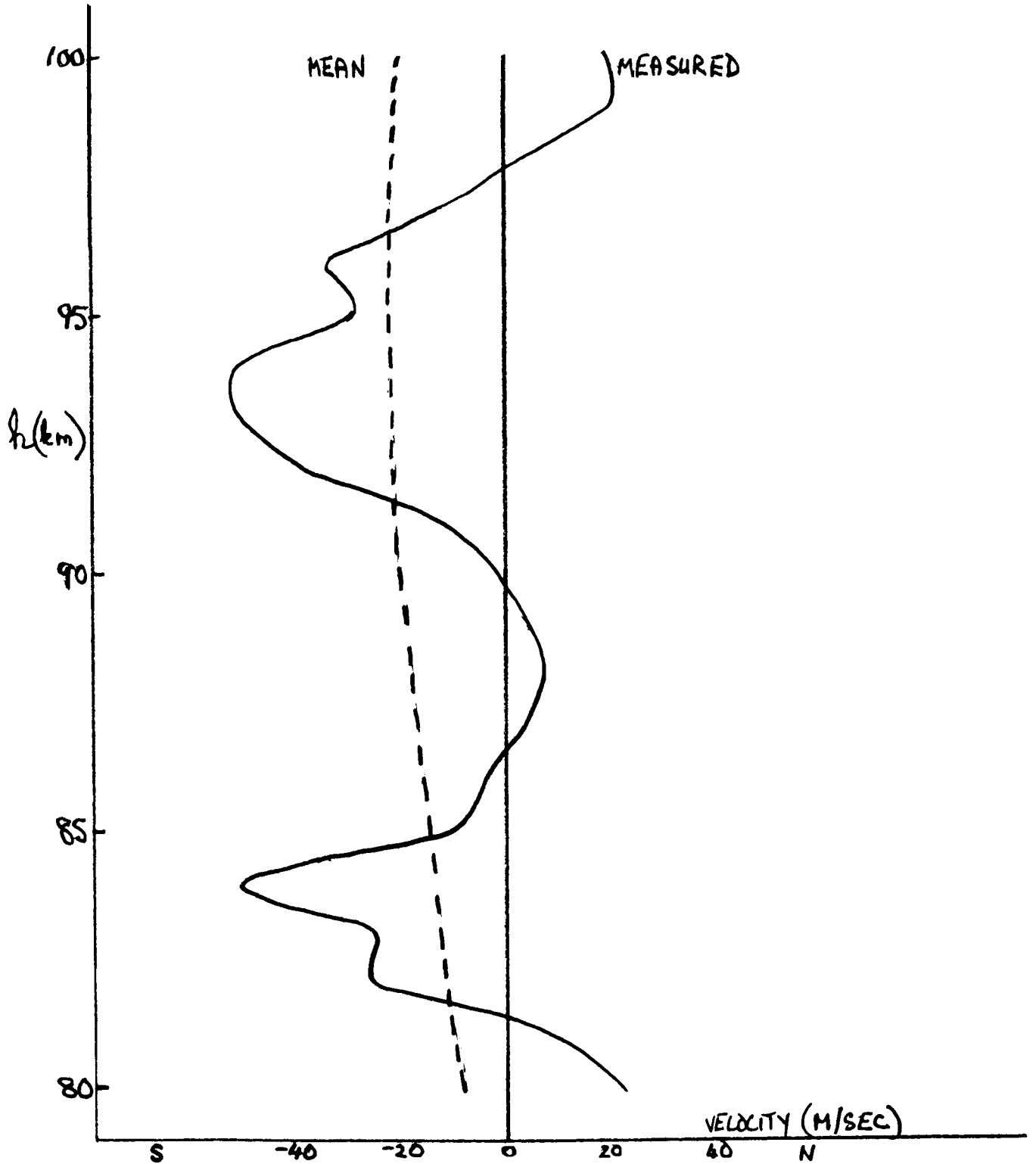


FIG. 4. MERIDIONAL PROFILE, 80-100KM.

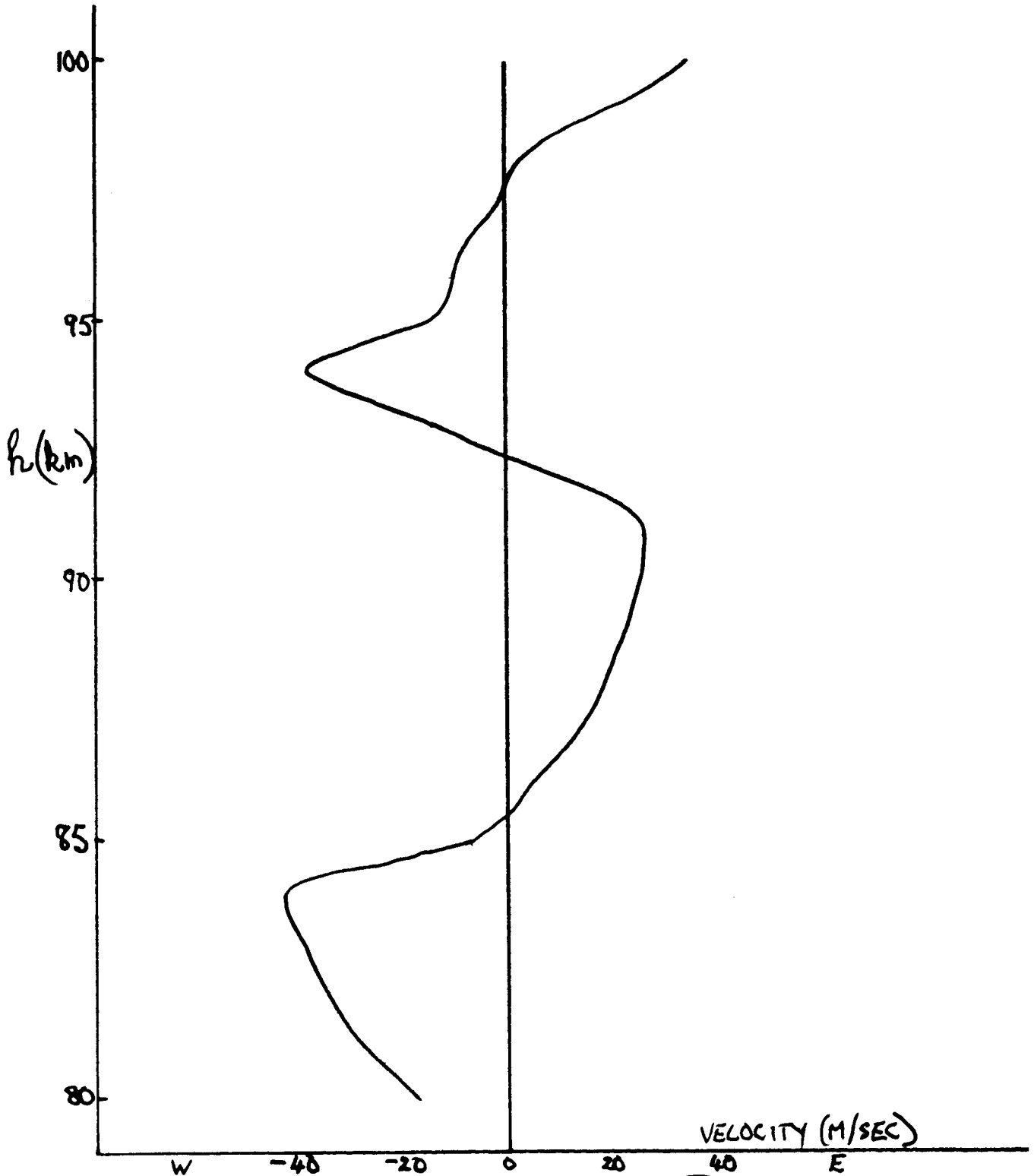


FIG. 5. TURBULENT ZONAL PROFILE

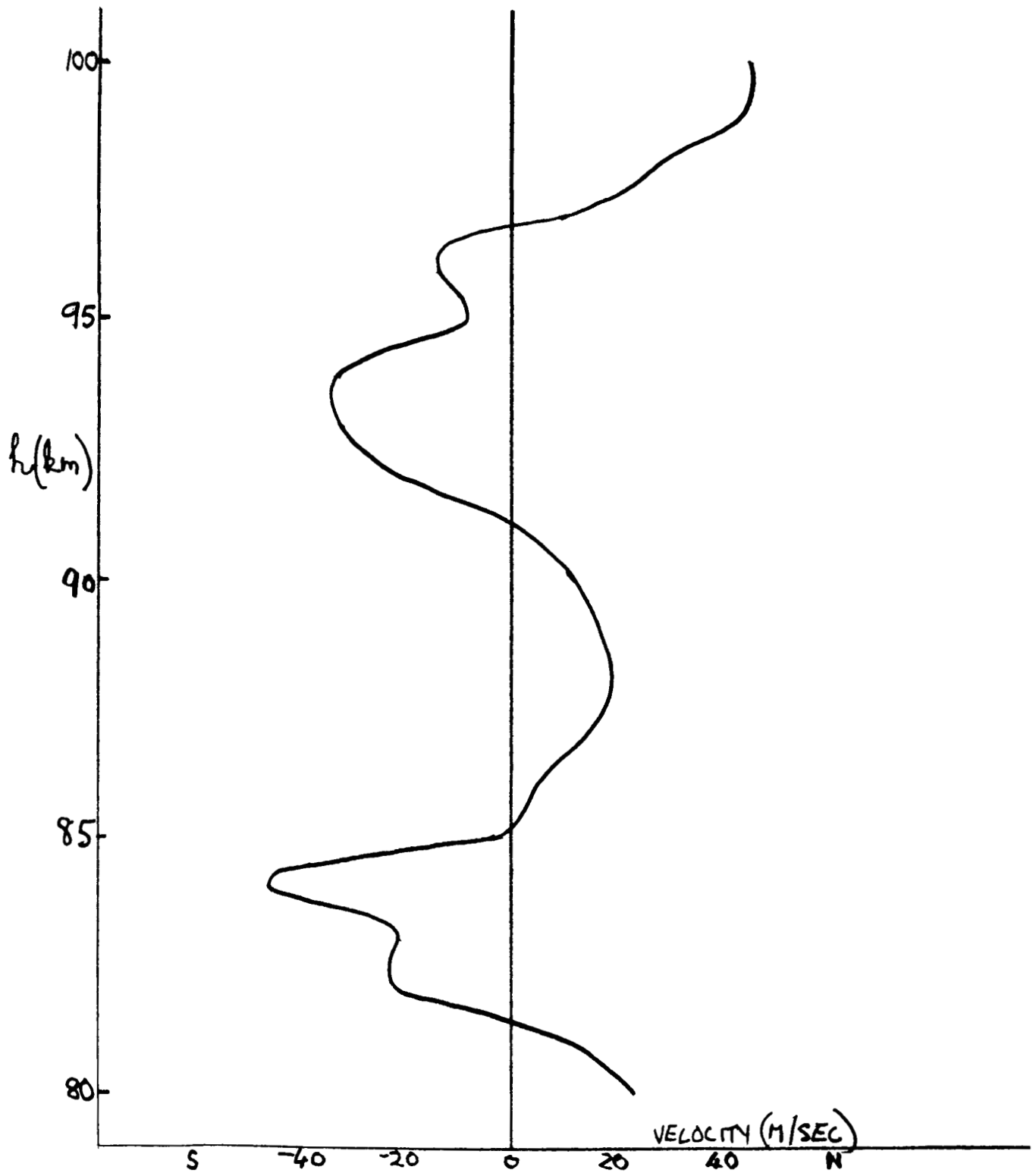


FIG.6. TURBULENT MERIDIONAL PROFILE

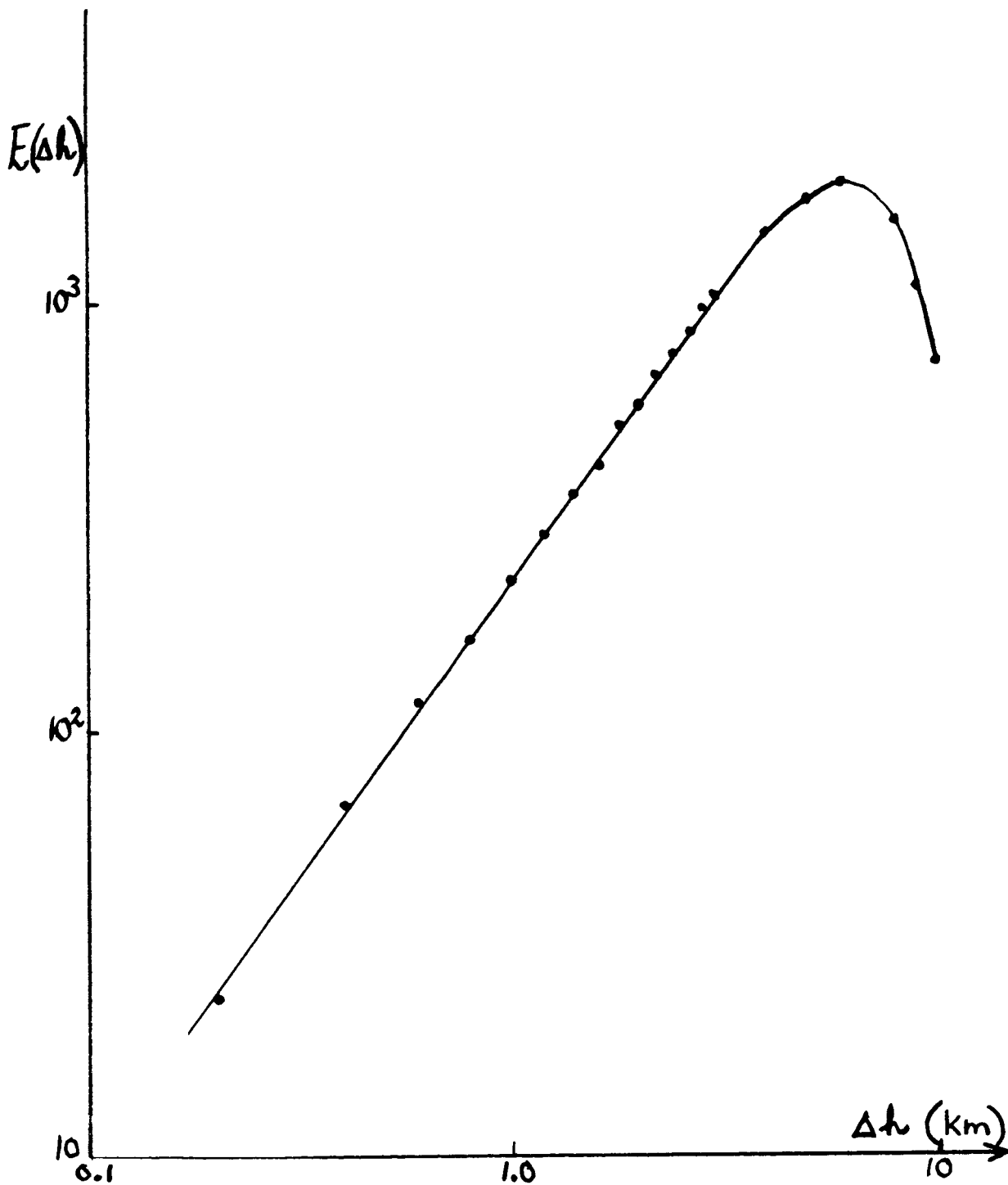


FIG. 7. THE ZONAL ENERGY SPECTRUM FUNCTION

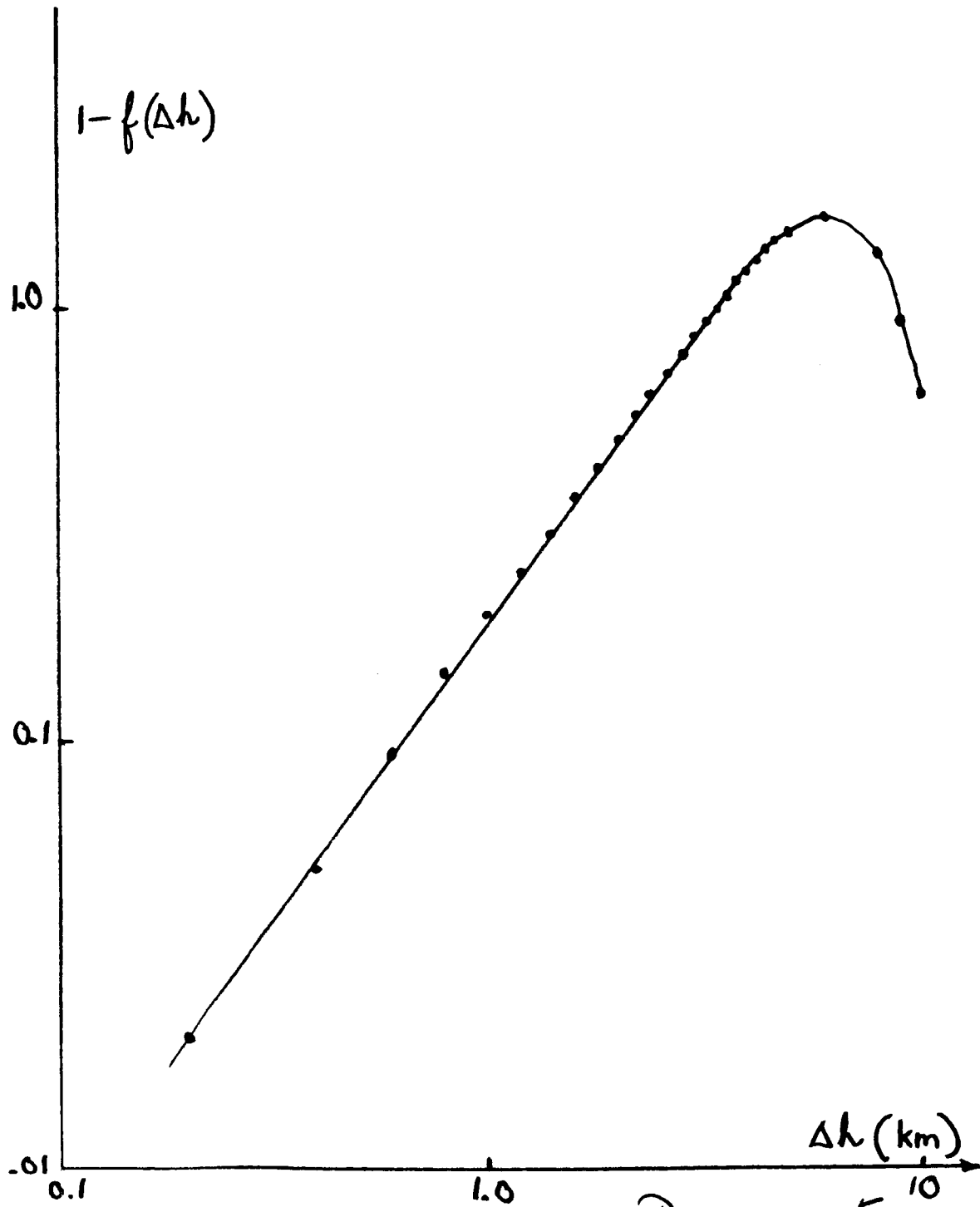


FIG. 8. THE ZONAL CORRELATION DIFFERENCE FUNCTION.

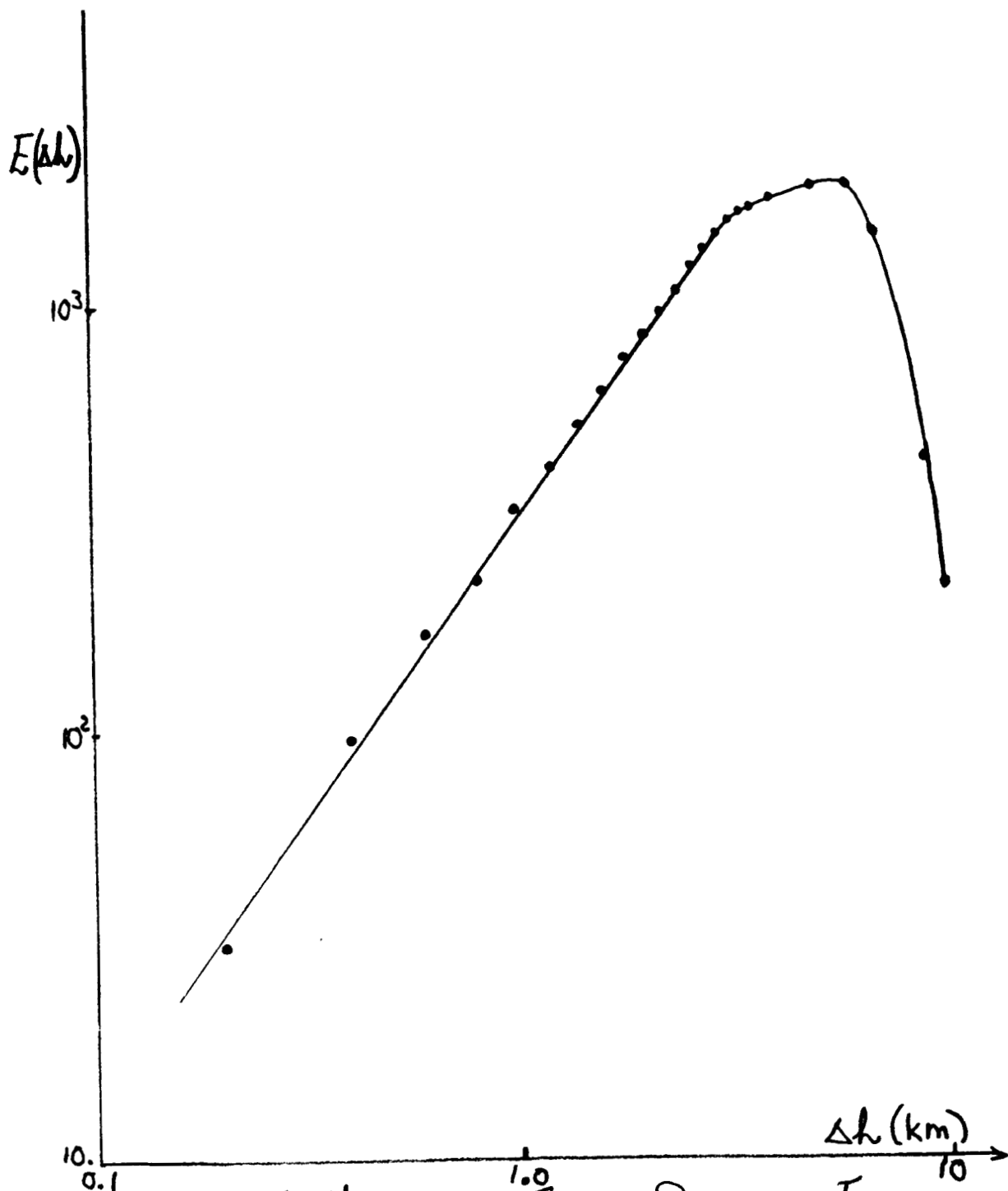


FIG. 9. THE MERIDIONAL ENERGY SPECTRUM FUNCTION.

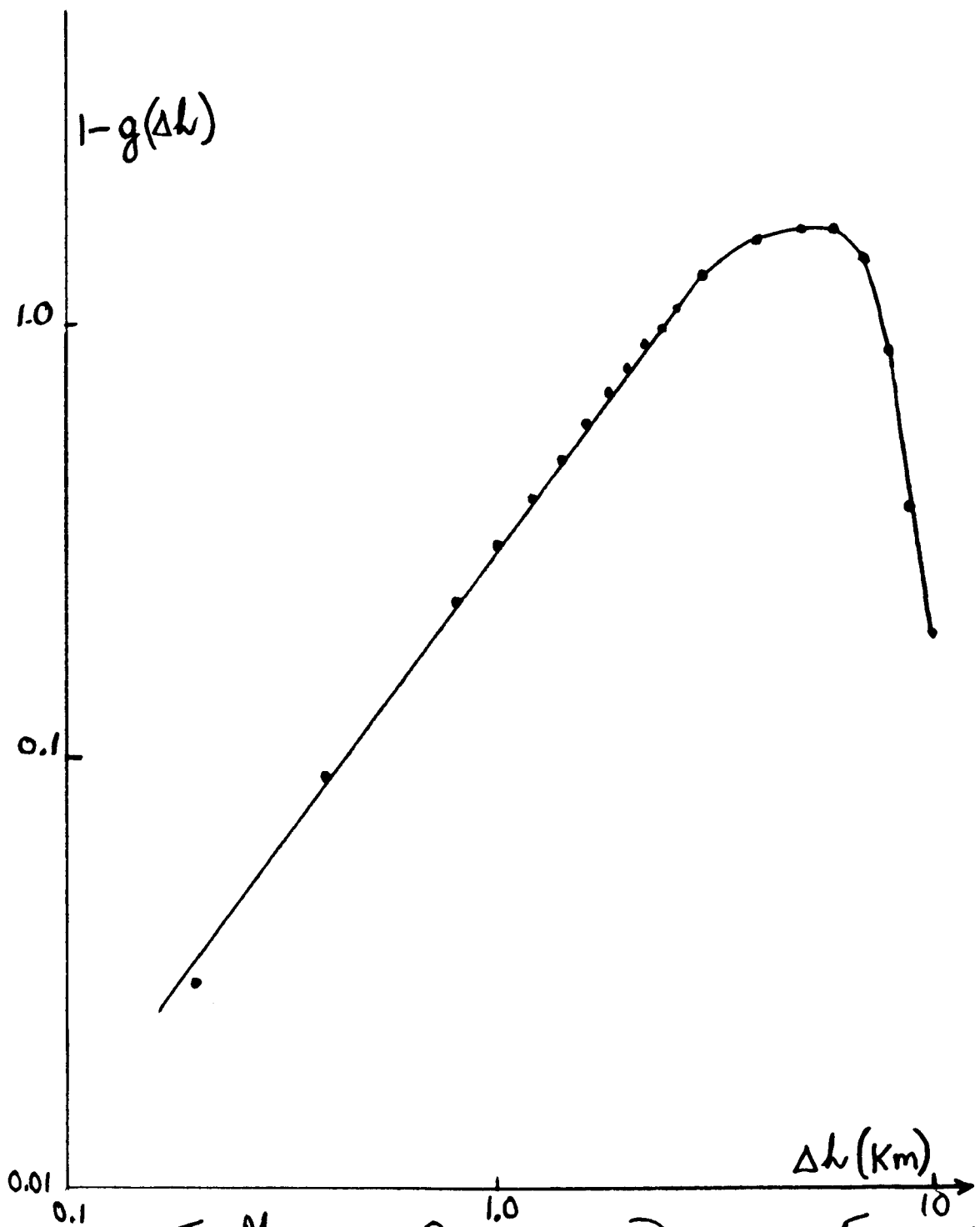


FIG. 10. THE MERIDIONAL CORRELATION DIFFERENCE FUNCTION.

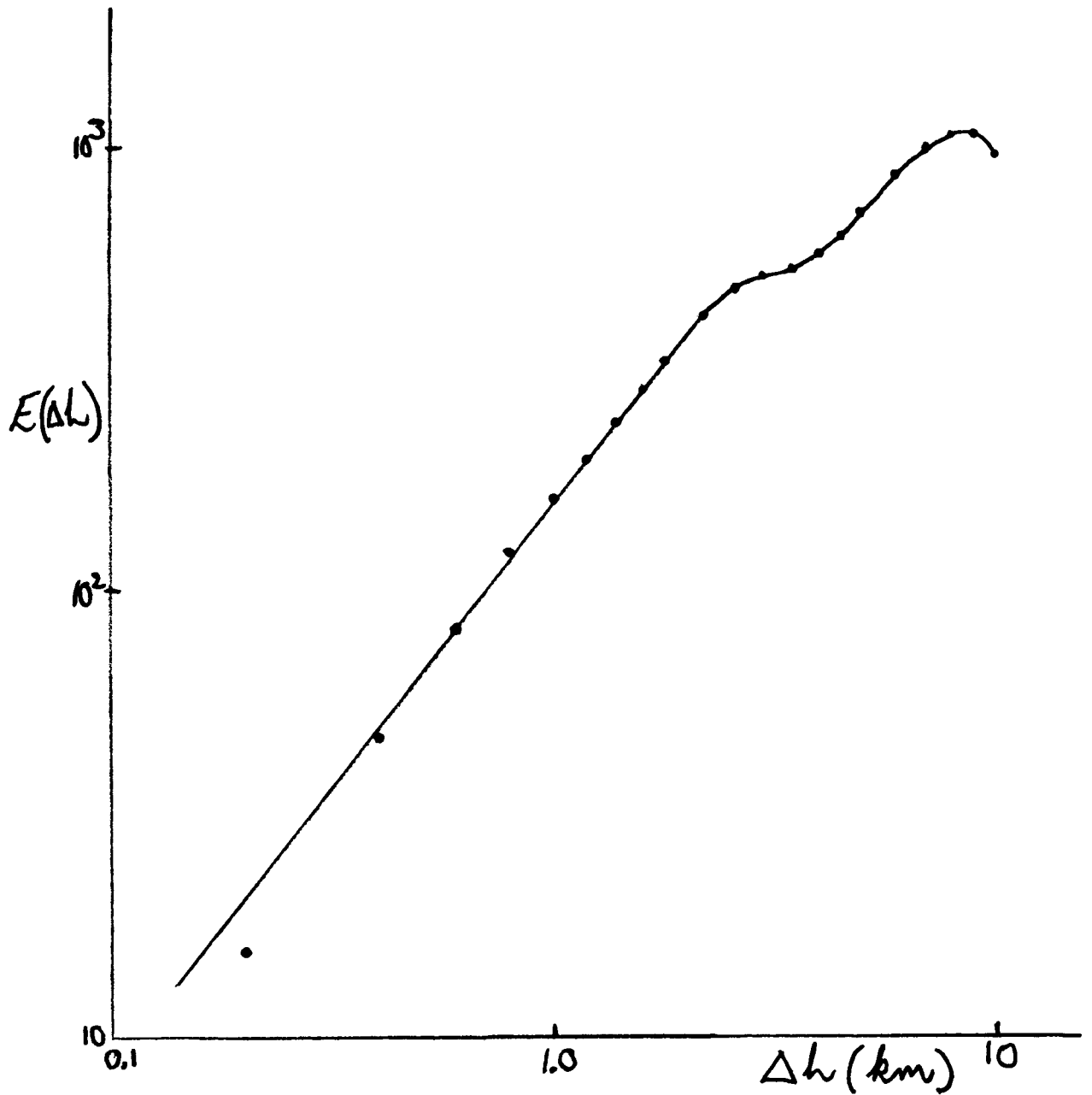


FIG 11. THE WINDSPEED ENERGY SPECTRUM FUNCTION.

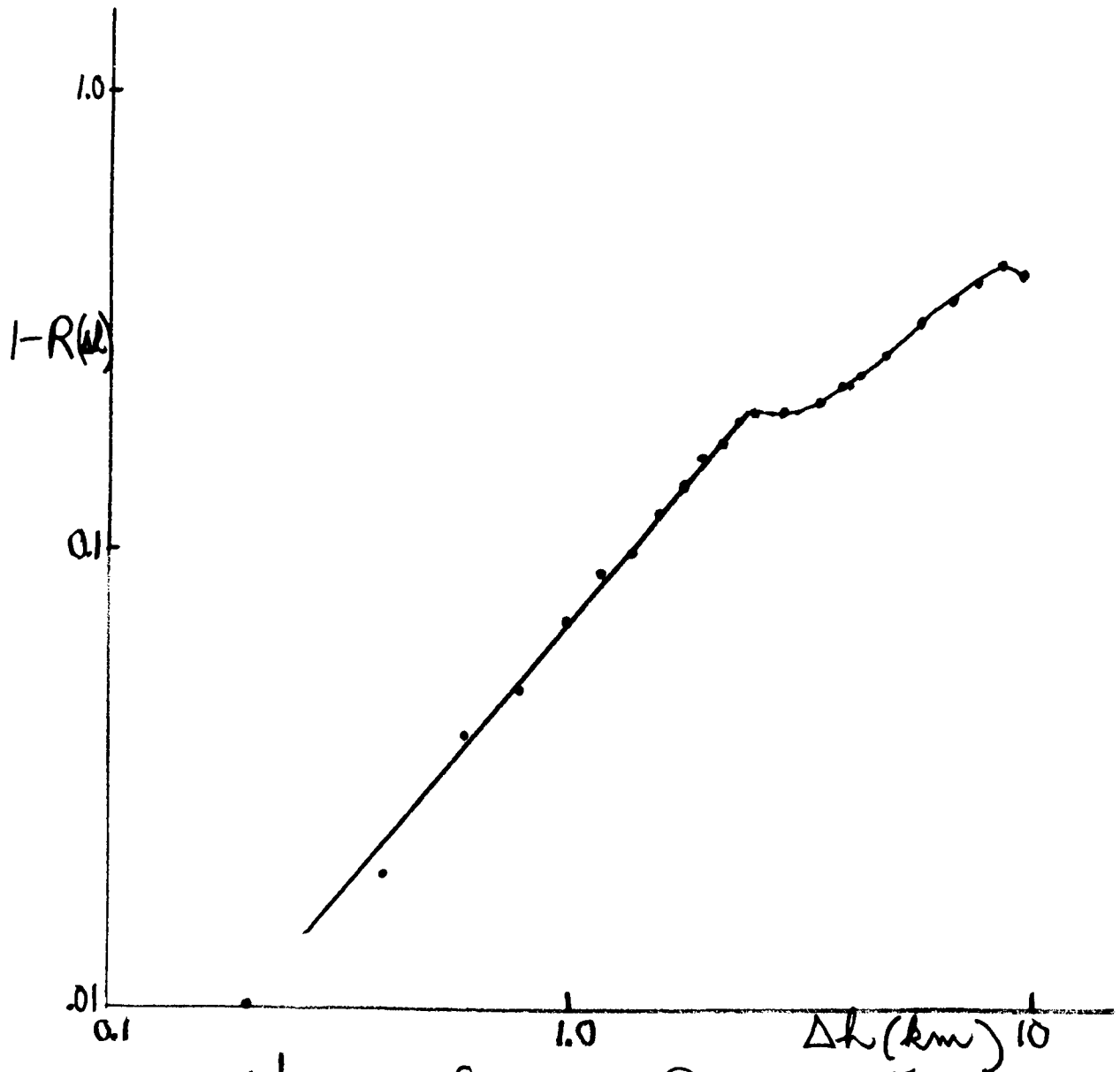


FIG 12. THE WINDSPEED CORRELATION DIFFERENCE FUNCTION

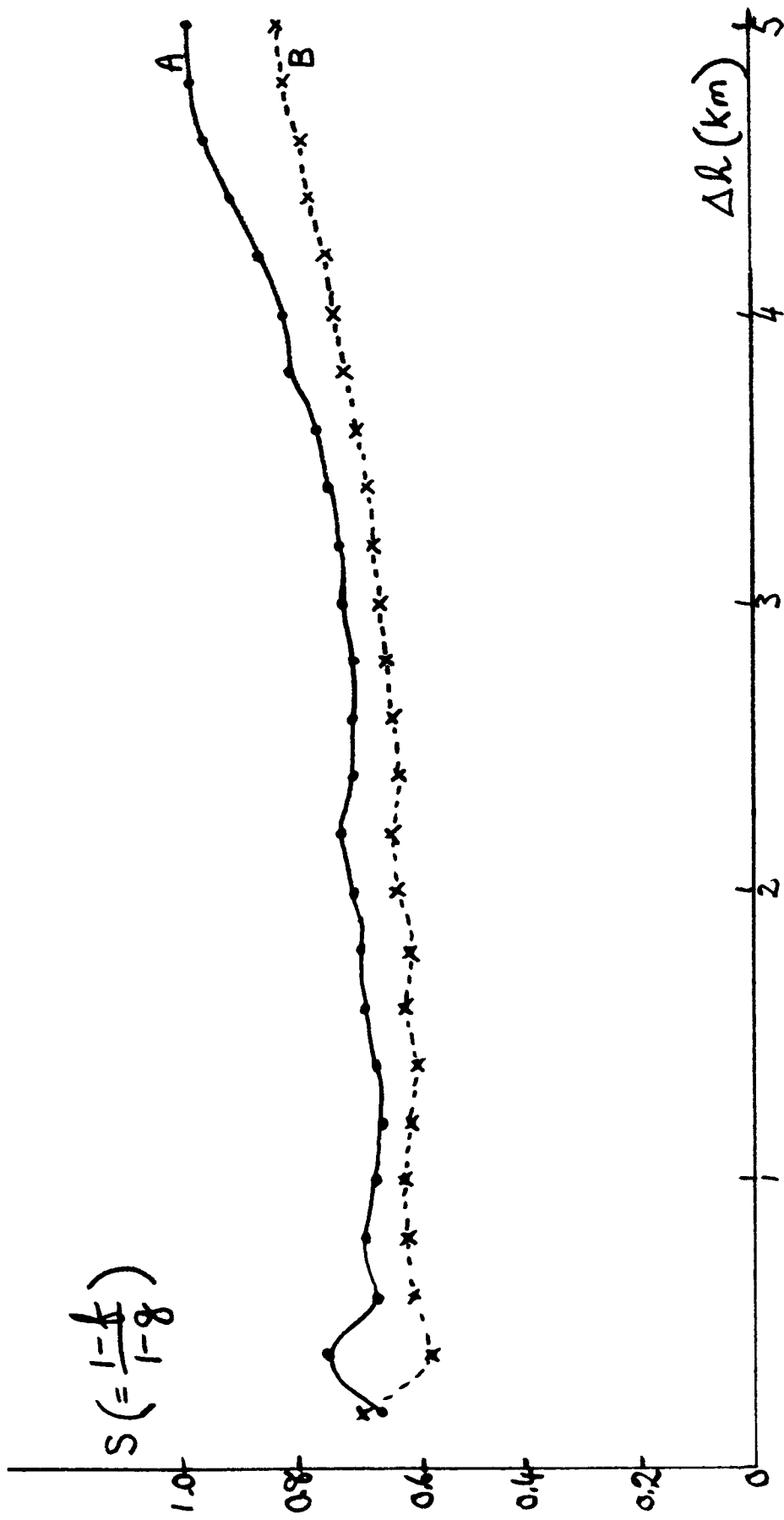


FIG. 13. VARIATION OF THE ISOTROPY PARAMETER "S" WITH SCALE.

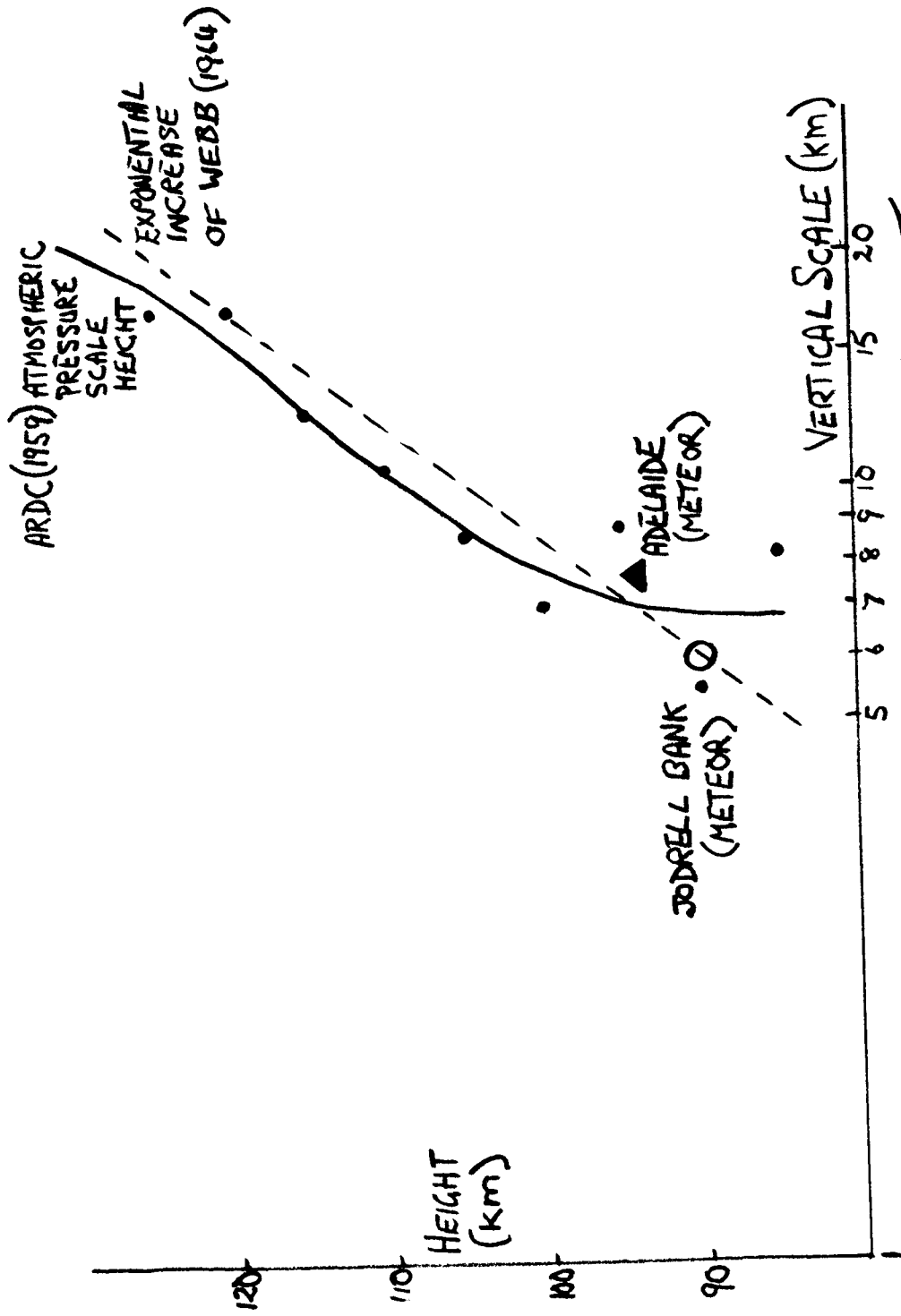


FIG. 14. VARIATION OF THE VERTICAL SCALE OF THE ZONAL TURBULENT WIND COMPONENT WITH HEIGHT

Appendix I

Resume of the Computer Program

The computer program has been written in FORTRAN IV for execution on an IBM 7094, and consists of the following routines:

1. The main program SHEAR which reads and does some processing of the input data (which is punched on cards).

SHEAR calls subroutines

a) FIT, which fits polynomial profiles to the zonal and meridional wind velocity / height profiles by the method of least squares, utilizing

b) MATS, to solve the set of independent linear equations

c) The total wind speed profile is calculated by TRADE.

d) OUTPUT performs the energy spectrum and correlation analyses, and printout of same.

e) PAGE is a utility routine which turns and numbers the output pages, and prints an appropriate heading on each.

Input Data

The input data is on punched cards, and is read as follows:

1. Header Card; the information punched on this card is printed at the top of each page of output. The full 72 columns can be used for any type of alphanumeric heading.

2. The parameters specifying the mean zonal and meridional profiles. These determine the order of the polynomial fitted to

- a) the zonal (punched in columns 2,3) and
- b) the meridional (punched in columns 5,6) measured data to determine the mean profile.

e.g., a 5 punched in column 3 will fit a polynomial of the form

$$a_1 + a_2h + a_3h^2 + a_4h^3 + a_5h^4$$

to the zonal wind data.

The maximum allowable value of the subscript i in the a_i above is 10.

3. The height range (must be integral, and in km) over which the energy spectrum and correlation analyses are to be performed. This can be any fraction of the height interval covered by the input data, but must not be less than 10 km unless an appropriate spectral range is punched in column 13. The minimum of the height range interval is punched in columns 2 to 5, the maximum in columns 6 to 9. If OUTPUT is required to produce spectra over height differences to other than a maximum of 10 km, the required maximum height difference can be punched in columns 10 to 13. This entry must not exceed the height range specified in columns 2 to 9. If this last field is left blank, spectra to 10 km are output.

4. The trial data, which must not exceed a maximum to minimum height range of 150 km. The data is punched as year, month, day, hour (local time, 24 hour clock), height (km), windspeed (metres/sec), and azimuth (degrees) as follows

YR	MO	DAY	HR	HEIGHT	WINDSPEED	AZIMUTH
3	3	9	14	19	24	29

all integer

must be right adjusted if integer;
anywhere in appropriate field if
decimal punch is included.

5. After all data of 4. a blank card to flag "end of data".

6. Further height ranges as for 3 as required.

7. a) If a change in the mean wind profile is required a card with a 2 punched in column 1 may be substituted for any of the cards 6, and followed by a card of format 2 above which specified the new profile. This card is then followed by further height ranges (as for 6.) as required.

7 b) If a completely new set of header plus data cards is to be processed after either stages 5, 6, or 7a, a card with a 1 punched in column 1 will return control to the start of the program, which will then read in the new set of data cards sequenced as from 1 above.

APPENDIX II

Listing of the FORTRAN IV Program

C SODIUM TRAIL HEIGHT SHEAR ANALYSIS PROGRAM
 C
 C CALCULATES THE ENERGY SPECTRUM FOR TOTAL SHEAR, AND FOR ZONAL AND
 C MERIDIONAL SHEAR COMPONENTS
 C
 C READS INPUT DATA AS FOLLOWS
 C A HEADER CARD, PUNCHED WITH DETAILS OF FIRING TIME, ETC.
 C FORMAT 12A6
 C THE PARAMETERS SPECIFYING THE MEAN ZONAL AND MERIDIONAL PROFILES,
 C FORMAT 2I3
 C THE HEIGHT RANGE OVER WHICH ANALYSIS IS TO BE PERFORMED, ZMIN,
 C ZMAX, AND THE MAXIMUM OF THE OUTPUT SPECTRAL RANGE DESIRED.
 C FORMAT 1X2F4.0, I4.
 C SPECTRAL RANGE MUST NOT EXCEED ZMAX-ZMIN.
 C IF NO SPECTRAL RANGE IS SPECIFIED, ZMAX-ZMIN MUST BE GREATER THAN
 C 10KM, AND SPECTRA WILL BE OUTPUT TO 10KM.
 C THE TRAIL DATA, YEAR MONTH DAY LOCAL TIME (HOURS AND MINUTES,
 C 24 HOUR CLOCK) HEIGHT(KM) WIND SPEED(METRES/SEC) WIND AZIMUTH
 C (DEGREES).
 C FORMAT 3I3, I5, F5.1, 2F5.0
 C A BLANK CARD
 C FURTHER ZMIN, ZMAX AS REQUIRED. IF A '2' APPEARS IN COLUMN 1,
 C PROGRAM WILL READ NEXT CARD AS NEW PROFILE SPECIFICATION, FOLLOWED
 C BY FURTHER ZMIN, ZMAX CARDS. IF A '1' APPEARS IN COLUMN 1, THE
 C PROGRAM WILL READ NEXT CARD AS HEADER CARD OF A COMPLETELY NEW SET
 C OF DATA.
 C

```

DIMENSION BZ(10), BM(10)
DIMENSION RESULT(12), ZI(750), WINDI(750), AZRADI(750), ZONAL(750),
MERID(750),          TZONAL(750), TMERID(750)
DIMENSION TSPEED(750), AW(5)
COMMON ZI, WINDI, AZRADI, ZONAL, ERID, TZONAL, TMERID, TSPEED
C=0.05
CS=0.000001
PI=3.1415926
TWOPI=2.0*PI
999 NGO=0
NI=0
I=0
NSUM=0
ZERO=0.0
READ (2,1)RESULT
1 FORMAT(12A6)
17 READ (2,4)NP,NC
4 FORMAT (2I3)
NFIT=-1
2 READ(2,3)NGO,ZMIN,ZMAX,NCIFF
3 FORMAT(I1,2F4.0,I4)
IF(NGC)19,19,18
18 IF(NGC-2)999,17,17
19 CONTINUE
IF(NCIFI)20,20,21
20 NDIFF=10
21 MIN=ZMIN+C
MAX=ZMAX+C
  
```

```

C      IF(NI)5,5,101
      TRAIL DATA INPUT
5 READ      (2,6)MYEAR,MCNTH,JCUR,LTIME,Z,WIND,AZDEG

16 FORMAT(3I3,15,F5.1,2F5.C)
      IF(MYEAR)100,100,7
17 I=I+1
      IF(I-1)11,11,8
18 ZDIFF=Z-ZI(I-1)
      MZ=10.0*Z+C
      LZ=10.0*ZI(I-1)+C
      LDIFF=MZ-LZ
      IF(LDIFF-2)12,11,9
12 WRITE      (3,6)MYEAR,MCNTH,JCUR,LTIME,Z,WIND,AZDEG
      WRITE(3,13)
13 FORMAT(1X/////1X26+DATA CARDS OUT OF SEQUENCE/////1X20HEXECUTICN
      1TERMINATED/////1X)
      PRINT 13
      CALL EXIT
C      INTERPOLATION ROUTINE
      9 NPOINT=LDIFF/2-1
      NSUM=NSUM+NPOINT
      GRAD=(WIND - WINDI(I-1))/ZDIFF
      AZDIFF=AZDEG/57.3-AZRADI(I-1)
      IF(ABS(AZDIFF)-PI)94,94,91
91 IF(AZDIFF)92,94,93
92 AZDIFF=AZDIFF+TWCP1
      GOTO94
93 AZDIFF=AZDIFF-TWCP1
94 AZGRAC=AZDIFF/ZDIFF
      DU10J=1,NPCINT
      ZI(I)=ZI(I-1)+0.2
      WINDI(I)=WINDI(I-1)+0.2*GRAC
      AZRADI(I)=AZRADI(I-1)+0.2*AZGRAC
      IF(AZRADI(I))95,96,96
95 AZRADI(I)=AZRADI(I)+TWCP1
96 IF(AZRADI(I)-TWCP1)98,98,97
97 AZRADI(I)=AZRADI(I)-TWCP1
98 CONTINUE
      I=I+1
10 CONTINUE
11 ZI(I)=Z
      WINDI(I)=WIND
      AZRADI(I)=AZDEG/57.3
      GO TO 5
100 NI=1
C      ALL DATA IN. COMMENCE COMPUTATION
      HDIFF=ZI(NI)-ZI(1)
      HPLUS=ZI(NI)+ZI(1)
C      CALCULATE ZONAL AND MERIDIONAL WIND COMPONENTS
      DU201J=1,NI
      ZONAL(J)=WINDI(J)* SIN(AZRADI(J))
201 ERID(J)=WINDI(J)* CCS(AZRADI(J))
101 NZ=0
      LMIN=10*MIN
      LMAX=10*MAX

```

```

CALL PAGE(RESULT, ZERO)
DU104J=1,NI
LZ=ZI(J)*10.C+C
IF(LZ-LMIN)104,102,102
102 IF(LZ-LMAX)103,103,104
103 NZ=NZ+1
104 CONTINUE
WRITE (3,105)NI,NSUM,ZMIN,ZMAX,NZ,NP,NG
105 FORMAT(1X////1X33FICTAL NUMBER OF INPUT DATA POINTS 18//1X26HINC
1LUDING INTERPOLATION OF 15,7H POINTS ////1X45HNUMBER OF DATA POINT
2S WITHIN THE HEIGHT RANGE F5.0,6H KM TO F5.0,3H KM 18//1X31HEAST
3 WEST PROFILE SPECIFICATION 15 //1X33FNORTH SOUTH PROFILE SPECIFI
4CATION 13 //1X)
START=1
CALL PAGE(RESULT,START)
WRITE (3,106)ZMIN,ZMAX
106 FORMAT(1X50HCALCULATED ZONAL AND MERIDIONAL MEAN WIND PROFILES
110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1H0
2/10X 6HHEIGHT 10X5HZONAL 13X10HMERIDIONAL 10X10HWIND SPEED /1X/
31X,18X,3(4X12HDATA MEAN,4X),/1X/)
NHITE=(ZMAX-ZMIN)/20.C-C
NHITE=NHITE+1
INK=5*NHITE-1
NFIT=NFIT+1
IF(NFIT)202,202,203
C CALCULATE COEFFICIENTS OF MEAN PROFILE POLYNOMIALS
202 CALL NCRNAL(ZCNAL,ERID,ZI,NI,BZ,BM,MP,NG)
203 CONTINUE
C DETERMINE AND PRINT CLT MEAN WIND PROFILES
SUBS=(ZMAX+ZMIN-ZI(1))*5.0+1.0+C
DU108L=MIN,MAX,NHITE
IZ=SUBS-FLCAT(L*5)
J=MAX+MIN-L
H=J
S=(2.0*H-HPLUS)/HDIFF+CS
AZONAL=0.0
AMERID=0.0
DU204M=1,10
AZONAL=AZCNAL+BZ(M)*S**(M-1)
AMERID=AMERID+BM(M)*S**(M-1)
204 CONTINUE
MZONAL=AZCNAL
MMERID=AMERID
MSPEED=SQRT(AZCNAL**2+AMERID**2)
WRITE(3,107)J,ZCNAL(IZ),MZCNAL,ERID(IZ),MMERID,WINDI(IZ),MSPEED
107 FORMAT(1X113,4X,3( F8.0,4X14,4X),/1X)
108 CONTINUE
CALL PAGE(RESULT,START)
WRITE(3,109)ZMIN,ZMAX
109 FORMAT(1X,22HTURBULENT WIND PROFILE,10X12HHEIGHT RANGE F5.0,
1 6HKM TO F5.0,3H KM/1X/1X,10X6HHEIGHT,10X6HTZONAL,10X11HTMERIDICNA
2L,10X6HTSPEED/1X)
C DETERMINE AND PRINT CLT TURBULENT WIND PROFILES
NH=(MAX-MIN)*5+1
K=(ZMAX-ZI(1))*5.0+2.0+C
NK=INK

```



```

RMSZ=0.0
RMSM=0.0
RMST=0.0
SUMSQT=0.0
DO112J=1,NP
TZONAL(J)=0.0
TMERIC(J)=0.0
K=K-1
H=ZMAX-0.2*FLOAT(J-1)
IH=H+C
S=(2.0*H-FPLUS)/FCIFF+CS
NK=NK+1
DU205M=1,10
TZONAL(J)=TZONAL(J)+BZ(M)*S**(M-1)
TMERIC(J)=TMERIC(J)+BM(M)*S**(M-1)
205 CONTINUE
TZONAL(J)=ZONAL(K)-TZONAL(J)
TMERIC(J)=ERIC(K)-TMERIC(J)
CALL TRADE(J,K)
C CALCULATE RMS TURBULENT VELOCITIES FOR THE HEIGHT RANGE ZMIN/ZMAX
RMSZ=TZONAL(J)**2+RMSZ
RMSM=TMERIC(J)**2+RMSM
RMST=TSPEED(J)**2+RMST
SUMSQT=SUMSQT+1.0
IF(NK-INK)112,112,110
110 WRITE(3,111)IH,TZONAL(J),TMERIC(J),TSPEED(J)
111 FORMAT(1X,10X14,12XF5.0,11XF8.0,13XF5.0/1X)
NK=0
112 CONTINUE
CALL PAGE(RESULT,START)
RMSZ=SQRT(RMSZ/SUMSQT)
RMSM=SQRT(RMSM/SUMSQT)
RMST=SQRT(RMST/SUMSQT)
WRITE(3,206)ZMIN,ZMAX,(BZ(I),I=1,NP)

206 FORMAT(1X1CHNCRMALIZED
1 1X12HHEIGHT RANGE10XF7.0,7H KM. TO F7.0,4H KM./1H0/
2 1X42HCOEFFICIENTS GIVING BEST FIT TO ZONAL DATA/1X/10E11.3)
WRITE(3,207)(BM(I),I=1,NC)
207 FORMAT(1X/1X/
11X/1X47HCCEFFICIENTS GIVING BEST FIT TO MERIDIONAL DATA/1X/1X10E11
2.3/1X/)
WRITE(3,208)RMSZ,RMSM,RMST
208 FORMAT(1X/1X
1 1X22HRMS TURBULENT VELOCITY/1X/1X10X5H-ZONALF11.0,11HMETRES/S
2EC./1X/1X10X10HMERIDIONAL F6.0,11HMETRES/SEC./1X/1X10X10HWIND SPEE
3D F6.0,11HMETRES/SEC.)
CALL PAGE(RESULT,START)
WRITE (3,113)ZMIN,ZMAX
113 FORMAT(1X46HENERGY SPECTRUM OF ZONAL WIND HEIGHT VARIATION,
110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)
C CALCULATE AND OUTPUT ENERGY SPECTRUM FUNCTIONS
CALL OUTPUT(TZONAL,NF,NDIFF)
CALL PAGE(RESULT,START)
WRITE (3,114) ZMIN,ZMAX
114 FORMAT(1X51HENERGY SPECTRUM OF MERIDIONAL WIND HEIGHT VARIATION,

```

```
/$ID JO2T          RGR   BLDC 11
$PAUSE
$EXECUTE          IBJOB
IBJOB VERSION 2,  7090 - PR - 929
```

```
1
0
$IBJOB           GC
$IBFTC OUT       M94,XR7
1               CUT
                EXTERNAL FORMULA NUMBER - SOURCE STATEMENT - INTERI
```

```
0
C               SUBROUTINE OUTPUT(TSPEED,NH,NDIFF)
C               PRODUCES HEIGHT SPECTRUM AND PERFORMS CORRELATION ANALYSIS.
C
```

```
                DIMENSION TSPEED(750)
                WRITE(3,1)
1  FORMAT(1X,10X7HDELTA H,5X10HE(DELTA H),
15X11HCORRELATION,3X13HSIGNIF 1-G /1X)
                NEND=5*NDIFF
                LINE=0
                DO125K=1,NEND
                LINE=LINE+1
                DELTAH=FLCAT(K)*C.2
                SUM=0.0
                ENERGY=0.0
                ERGX=0.0
                GJPKXJ=0.0
                SQJ=0.0
                SQJPK=0.0
                DO123J=1,NH
                IF(NH-J-K)124,122,122
122  JPK=J+K
                ENERGY=ENERGY+(TSPEED(JPK)-TSPEED(J))**2
                SUM=SUM+1.0
                GJPKXJ=GJPKXJ+TSPEED(JPK)*TSPEED(J)
                SQJ=SQJ+TSPEED(J)**2
                SQJPK=SQJPK+TSPEED(JPK)**2
123  CONTINUE
124  IF(SUM)117,117,118
117  G=0.0
                SIGNIF=9.99
                GOTO119
118  ENERGY=ENERGY/SUM
                G=GJPKXJ/SQRT(SQJ*SQJPK)
                SIGNIF=(1.0-G**2)/SQRT(SUM-1.0)
119  DIFF=1.0-G
                IF(LINE-50)125,125,116
116  WRITE(3,2)
                2  FORMAT(1H1/1X)
                WRITE(3,1)
                LINE=0
125  WRITE(3,115)DELTAH,ENERGY,      G,SIGNIF,DIFF
115  FORMAT(1X,6XF8.1,  7XF8.1 ,8XF8.3,F11.3,F8.2)
                RETURN
                END
```

SHEAR
EXTERNAL FORMULA NUMBER - SOURCE STATEMENT - INTERNAL 08

```
110X12HEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)
CALL CUTPUT(TMERICID,NF,NCIFF)
CALL PAGE(RESULT,START)
WRITE      (3,115)ZMIN,ZMAX
115 FORMAT(1X46HENERGY SPECTRUM OF WIND SPEED HEIGHT VARIATION,
110X12HEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)
CALL CUTPUT(TSPEED,NF,NCIFF)
GOTO2
END
```

\$IBJOB GC
\$IBFTC NORMAL M94, XR7

NCRMAL
EXTERNAL FORMULA NUMBER - SOURCE STATEMENT - INTE

SUBROUTINE NORMAL (ZCNAL,ERID,ZI,NI,BZ,BM,NP,NQ)

C
C
C
DETERMINES BEST FIT PROFILE TO THE NORMALIZED TOTAL HEIGHT RANG

DIMENSION ZONAL(750),ERID(750),ZI(750),BZ(10),BM(10),ZINORM(750)
DO1J=1,NI

ZINORM(J)=(2.0*ZI(J)-ZI(NI)-ZI(1))/(ZI(NI)-ZI(1))

1 CONTINUE

CALL FIT (ZCNAL,NP,BZ,NI,ZINORM)

CALL FIT (ERID, NQ,BM,NI,ZINORM)

RETURN

END

\$IBJCB CC
\$IBFTC TRADE MS4, XR7

SUBROUTINE TRADE(J,K)

C
C
C

CALCULATES TURBULENT WINDSPEED PROFILE

DIMENSION ZI(750)
DIMENSION WINDI(750),AZRADI(750),ZONAL(750),ERID(750),TZONAL(750),
1TMERID(750),TSPEED(750)
COMMON ZI, WINDI,AZRADI,ZONAL,ERID,TZONAL,TMERID,TSPEED
TSPEED(J)=WINDI(K)-SQRT(((ZONAL(K)-TZONAL(J))**2+(ERID(K)-TMERID(J)
1)**2))
RETURN
END

\$IBFTC FIT M94, XR7

SUBROUTINE FIT (BLOW, NSPEC, A, N, X)

```

C
C FITS A HEIGHT PROFILE SPECIFIED BY NSPEC TO THE MEASURED DATA
C BLOW, DETERMINING THE COEFFICIENTS A BY THE METHOD OF LEAST
C SQUARES. PROFILE FITTED OVER TOTAL HEIGHT RANGE OF INPUT DATA.
C
    DIMENSION BLOW(750), A(10), X(750), S(10,11)
    DO3J=1,10
3  A(J)=0.0
    IF(NSPEC)100,100,4
4  NADD=NSPEC+1
    DO10J=1,NSPEC
    DO10K=1,NADD
10 S(J,K)=0.0
    DO5J=1,NSPEC
    DO5K=1,NSPEC
    DO5I=1,N
5  S(J,K)=S(J,K)+X(I)**(K+J-2)
    DO6J=1,NSPEC
    K=0
    DO6I=1,N
    K=K+1
6  S(J,NADD)=S(J,NADD)+X(I)**(J-1)*BLOW(K)
    CALL MATS(S,A,NSPEC,MISS)
    IF(MISS)100,100,7
7  WRITE(3,8)
8  FORMAT(1X75HERROR IN INPUT DATA HAS RESULTED IN MATRIX S BEING UNS
    SUITABLE FOR INVERSION ////1X20HEXECUTION TERMINATED //////////)
    PRINT 8
    CALL EXIT
100 RETURN
    END

```

SUBROUTINE MATS (S,A,NSPEC,MISS)

C
 C
 C

REDUCES THE AUGMENTED MATRIX S TO PRODUCE THE COEFFICIENTS A

DIMENSION S(10,11), A(10)

MISS=-1

MM=NSPEC+1

N=NSPEC

DO15I=2,N

70 II=I-1

7 DO15J=1,II

8 IF(S(I,J))9,15,9

9 IF(ABS(S(J,J))-ABS(S(I,J)))11,10,10

10 R=S(I,J)/S(J,J)

GOTO 130

11 R=S(J,J)/S(I,J)

DO12K=1,MM

B=S(J,K)

S(J,K)=S(I,K)

12 S(I,K)=B

130 JJ=J+1

13 DO14K=JJ,MM

14 S(I,K)=S(I,K)-R*S(J,K)

15 CONTINUE

IF(ABS(S(N,N))-1.0E-10)16,16,17

16 MISS=+1

GOTO29

17 A(N)=S(N,MM)/S(N,N)

DO28I=2,N

JJ=N-I+1

B=0.0

II=N-I+2

DO25K=II,N

25 B=B+S(JJ,K)*A(K)

IF(ABS(S(JJ,JJ))-1.0E-10)16,16,28

28 A(JJ)=(S(JJ,MM)-B)/S(JJ,JJ)

29 RETURN

END

08/17

\$IBFTC PAGE M94, XR7

SUBROUTINE PAGE(RESET,START)

C
C
C
C

TURNS PAGE, NUMBERS IT, AND WRITES HEADING AS APPEARING
ON RESULT CARD.

DIMENSION RESULT(12)

IF(START)1,1,2

1 NOP=0

2 NOP=NOP+1

WRITE (3,3)RESULT,NOP

3 FORMAT(1H1/1X12A6,30X4HPAGE15//)

RETURN

END

APPENDIX III

SAMPLE DATA AND COMPUTER OUTPUT

63 5 21 1910105.6 140 130

63 5 21 1910105.8 130 128

63 5 21 1910107.0 125 132

63 5 21 1910108.0 92 140

63 5 21 1910109.0 50 160

63 5 21 1910110.0 42 240

63 5 21 1910111.0 56 250

63 5 21 1910113.0 60 262

63 5 21 1910114.0 84 274

63 5 21 1910115.0 68 270

63 5 21 1910117.2 72 282

63 5 21 1910125.0 45 286

63 5 21 1910130.0 30 296

63 5 21 1910135.0 14 310

63 5 21 1910140.0 15 316

← BLANK CARD

7
8

← E O F

PRELIMINARY PAPER EGLIN DATA, MAY 21, 1963. 16/9/64. PROFILE,5,5.

5 5

80 100

63 5 21 1910 69.0 34 255

63 5 21 1910 70.6 31 256

63 5 21 1910 71.0 27 270

63 5 21 1910 72.4 43 272

63 5 21 1910 72.6 20 290

63 5 21 1910 77.0 41 256

63 5 21 1910 80.4 25 358

63 5 21 1910 82.0 08 310

63 5 21 1910 82.4 28 202

63 5 21 1910 83.6 27 200

63 5 21 1910 84.0 54 190

63 5 21 1910 85.0 54 194

63 5 21 1910 85.4 30 108

63 5 21 1910 88.0 60 080

63 5 21 1910 91.6 76 095

63 5 21 1910 95.0 50 178

63 5 21 1910 96.0 40 120

63 5 21 1910 96.6 54 150

63 5 21 1910 97.6 40 115

63 5 21 1910 99.0 42 084

63 5 21 1910 99.6 64 078

63 5 21 1910100.0 58 070

63 5 21 1910100.8 65 069

63 5 21 1910102.0 80 079

63 5 21 1910103.0 93 094

63 5 21 1910104.0 115 104

63 5 21 1910104.2 111 120

63 5 21 1910105.0 129 122

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. 15/9/64. PROFILE,5,5.

TOTAL NUMBER OF INPUT DATA POINTS 356

INCLUDING INTERPOLATION OF 313 POINTS

NUMBER OF DATA POINTS WITHIN THE HEIGHT RANGE 80. KM TO 100. KM 101

EAST WEST PROFILE SPECIFICATION 5

NORTH SOUTH PROFILE SPECIFICATION 5

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE, 5, 5.

CALCULATED ZONAL AND MERIDIONAL MEAN WIND PROFILES

HEIGHT RANGE

HEIGHT	ZONAL		MERIDIONAL		WIND SPEED	
	DATA	MEAN	DATA	MEAN	DATA	MEAN
100	70.	34	20.	-21	72.	40
99	56.	37	21.	-21	60.	43
98	42.	39	4.	-22	42.	45
97	39.	41	-11.	-22	41.	47
96	34.	43	-35.	-22	48.	49
95	31.	45	-28.	-22	42.	50
94	6.	46	-51.	-22	52.	51
93	31.	46	-51.	-21	59.	51
92	55.	47	-38.	-21	67.	51
91	73.	47	-13.	-21	74.	51
90	72.	46	-1.	-20	72.	50
89	67.	45	4.	-19	67.	49
88	62.	43	8.	-18	63.	47
87	55.	41	5.	-18	55.	45
86	44.	39	-4.	-16	44.	43
85	31.	36	-9.	-15	32.	39
84	-12.	33	-53.	-14	54.	36
83	-9.	29	-25.	-13	27.	32
82	-10.	24	-26.	-11	28.	27
81	-8.	19	10.	-10	12.	22
80	-3.	14	23.	-8	23.	16

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE, 5, 5.

TURBULENT WIND PROFILE HEIGHT RANGE 80. KM TO 100. KM

HEIGHT	TZONAL	TMERIDIONAL	TSPEED
100	35.	41.	42.
99	19.	42.	26.
98	2.	26.	16.
97	-3.	11.	31.
96	-10.	-12.	32.
95	-14.	-6.	29.
94	-40.	-29.	3.
93	-16.	-29.	31.
92	8.	-16.	58.
91	26.	8.	62.
90	25.	19.	52.
89	22.	24.	45.
88	18.	27.	38.
87	13.	24.	33.
86	4.	13.	30.
85	-6.	7.	27.
84	-45.	-38.	18.
83	-39.	-12.	-5.
82	-35.	-14.	0.
81	-28.	20.	3.
80	-18.	31.	10.

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. 15/9/64. PROFILE,5,5.
NORMALIZED HEIGHT RANGE 80. KM. TO 100. KM.

COEFFICIENTS GIVING BEST FIT TO ZONAL DATA

0.386E 02 -0.126E 03 -0.244E 03 0.174E 03 0.205E 03

COEFFICIENTS GIVING BEST FIT TO MERIDIONAL DATA

-0.304E 02 0.198E 02 0.148E 03 -0.168E 02 -0.125E 03

RMS TURBULENT VELOCITY

ZONAL 22.METRES/SEC.

MERIDIONAL 23.METRES/SEC.

WIND SPEED 17.METRES/SEC.

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

ENERGY SPECTRUM OF ZONAL WIND HEIGHT VARIATION

HEIGHT RANGE 80.

DELTA H	DELTA V**2	CORRELATION	SIGNIF	1-G
0.2	24.3	0.979	0.004	0.02
0.4	64.9	0.943	0.011	0.06
0.6	113.7	0.900	0.019	0.10
0.8	169.4	0.851	0.028	0.15
1.0	230.7	0.797	0.037	0.20
1.2	297.1	0.739	0.047	0.26
1.4	369.4	0.676	0.056	0.32
1.6	450.4	0.604	0.066	0.40
1.8	534.8	0.525	0.076	0.47
2.0	619.9	0.449	0.084	0.55
2.2	703.2	0.374	0.091	0.63
2.4	789.0	0.297	0.097	0.70
2.6	877.1	0.217	0.102	0.78
2.8	964.9	0.136	0.106	0.86
3.0	1054.0	0.053	0.108	0.95
3.2	1133.4	-0.023	0.109	1.02
3.4	1209.9	-0.099	0.109	1.10
3.6	1285.2	-0.176	0.107	1.18
3.8	1360.8	-0.257	0.104	1.26
4.0	1436.0	-0.342	0.099	1.34
4.2	1507.0	-0.429	0.092	1.43
4.4	1567.1	-0.514	0.083	1.51
4.6	1614.3	-0.597	0.073	1.60
4.8	1681.4	-0.651	0.066	1.65
5.0	1745.7	-0.696	0.060	1.70
5.2	1804.0	-0.735	0.053	1.74
5.4	1852.3	-0.766	0.048	1.77
5.6	1885.5	-0.788	0.045	1.79
5.8	1897.3	-0.800	0.043	1.80
6.0	1880.3	-0.801	0.043	1.80
6.2	1863.6	-0.795	0.044	1.80
6.4	1847.4	-0.783	0.047	1.78
6.6	1827.7	-0.762	0.051	1.76
6.8	1803.8	-0.730	0.057	1.73
7.0	1775.0	-0.690	0.065	1.69
7.2	1740.9	-0.642	0.073	1.64
7.4	1701.2	-0.588	0.082	1.59
7.6	1656.5	-0.528	0.092	1.53
7.8	1608.4	-0.466	0.100	1.47
8.0	1555.6	-0.401	0.108	1.40
8.2	1497.1	-0.333	0.116	1.33
8.4	1432.6	-0.263	0.122	1.26
8.6	1361.9	-0.190	0.128	1.19
8.8	1285.1	-0.116	0.132	1.12
9.0	1202.6	-0.040	0.135	1.04
9.2	1114.7	0.038	0.136	0.96
9.4	1020.5	0.120	0.135	0.88
9.6	920.9	0.207	0.133	0.79
9.8	819.6	0.296	0.128	0.70
10.0	733.6	0.372	0.122	0.63

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

ENERGY SPECTRUM OF MERIDIONAL WIND HEIGHT VARIATION

HEIGHT RANGE

DELTA H	DELTA V**2	CORRELATION	SIGNIF	1-G
0.2	31.3	0.973	0.005	0.03
0.4	94.1	0.917	0.016	0.08
0.6	164.4	0.853	0.028	0.15
0.8	241.7	0.781	0.040	0.22
1.0	326.3	0.701	0.052	0.30
1.2	418.3	0.612	0.064	0.39
1.4	515.2	0.518	0.076	0.48
1.6	616.0	0.420	0.086	0.58
1.8	723.6	0.316	0.094	0.68
2.0	825.3	0.219	0.100	0.78
2.2	935.1	0.116	0.105	0.88
2.4	1052.2	0.008	0.107	0.99
2.6	1175.9	-0.104	0.106	1.10
2.8	1305.7	-0.219	0.103	1.22
3.0	1429.3	-0.325	0.097	1.33
3.2	1534.6	-0.411	0.091	1.41
3.4	1614.7	-0.479	0.085	1.48
3.6	1667.7	-0.535	0.079	1.54
3.8	1711.0	-0.582	0.074	1.58
4.0	1741.1	-0.617	0.069	1.62
4.2	1756.8	-0.641	0.066	1.64
4.4	1758.3	-0.656	0.065	1.66
4.6	1745.6	-0.662	0.064	1.66
4.8	1776.6	-0.680	0.062	1.68
5.0	1836.6	-0.715	0.056	1.72
5.2	1878.5	-0.733	0.054	1.73
5.4	1900.7	-0.735	0.054	1.74
5.6	1902.8	-0.723	0.056	1.72
5.8	1885.6	-0.699	0.061	1.70
6.0	1853.4	-0.666	0.066	1.67
6.2	1809.8	-0.625	0.073	1.63
6.4	1751.9	-0.573	0.081	1.57
6.6	1679.1	-0.510	0.090	1.51
6.8	1592.6	-0.435	0.100	1.44
7.0	1494.3	-0.350	0.109	1.35
7.2	1385.9	-0.257	0.117	1.26
7.4	1269.8	-0.156	0.123	1.16
7.6	1149.0	-0.051	0.127	1.05
7.8	1028.4	0.057	0.128	0.94
8.0	912.5	0.164	0.126	0.84
8.2	802.4	0.266	0.121	0.73
8.4	697.1	0.366	0.114	0.63
8.6	598.6	0.459	0.105	0.54
8.8	508.4	0.544	0.094	0.46
9.0	427.7	0.620	0.083	0.38
9.2	357.9	0.684	0.072	0.32
9.4	302.8	0.735	0.063	0.27
9.6	260.1	0.774	0.056	0.23
9.8	231.1	0.801	0.050	0.20
10.0	219.4	0.812	0.048	0.19

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE, 5,5.

ENERGY SPECTRUM OF WIND SPEED HEIGHT VARIATION

HEIGHT RANGE 80.

DELTA H	DELTA V**2	CORRELATION	SIGNIF	1-G
0.2	15.1	0.993	0.001	0.01
0.4	47.1	0.979	0.004	0.02
0.6	83.2	0.964	0.007	0.04
0.8	122.0	0.947	0.011	0.05
1.0	161.5	0.930	0.014	0.07
1.2	200.6	0.914	0.017	0.09
1.4	240.7	0.898	0.020	0.10
1.6	286.9	0.879	0.024	0.12
1.8	337.0	0.859	0.028	0.14
2.0	382.2	0.841	0.031	0.16
2.2	421.0	0.827	0.034	0.17
2.4	455.3	0.815	0.036	0.19
2.6	483.4	0.805	0.038	0.20
2.8	502.9	0.799	0.039	0.20
3.0	512.5	0.797	0.040	0.20
3.2	517.8	0.796	0.040	0.20
3.4	528.9	0.793	0.041	0.21
3.6	542.8	0.789	0.042	0.21
3.8	561.7	0.783	0.043	0.22
4.0	584.0	0.775	0.045	0.22
4.2	606.6	0.768	0.046	0.23
4.4	628.3	0.761	0.048	0.24
4.6	650.9	0.755	0.049	0.25
4.8	679.5	0.746	0.051	0.25
5.0	712.5	0.734	0.053	0.27
5.2	748.8	0.722	0.056	0.28
5.4	785.1	0.710	0.058	0.29
5.6	818.7	0.700	0.060	0.30
5.8	847.5	0.692	0.062	0.31
6.0	870.6	0.686	0.063	0.31
6.2	894.8	0.680	0.065	0.32
6.4	919.4	0.674	0.066	0.33
6.6	943.4	0.668	0.068	0.33
6.8	968.1	0.661	0.069	0.34
7.0	996.2	0.653	0.071	0.35
7.2	1023.8	0.645	0.073	0.35
7.4	1048.3	0.638	0.075	0.36
7.6	1070.2	0.630	0.077	0.37
7.8	1088.9	0.623	0.078	0.38
8.0	1103.4	0.616	0.080	0.38
8.2	1114.0	0.608	0.082	0.39
8.4	1121.1	0.600	0.084	0.40
8.6	1120.2	0.593	0.086	0.41
8.8	1114.6	0.588	0.087	0.41
9.0	1103.8	0.585	0.089	0.42
9.2	1087.7	0.584	0.090	0.42
9.4	1066.1	0.586	0.090	0.41
9.6	1039.0	0.590	0.090	0.41
9.8	1007.8	0.596	0.090	0.40
10.0	973.4	0.603	0.090	0.40