

DATA FROM EXPLORER 17 ON COMPOSITION OF THE UPPER ATMOSPHERE*

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INTRODUCTION

The Explorer 17 aeronomy satellite carried two double-focusing magnetic mass spectrometers designed [Meadoes, 1960; Hall *et al.*, 1960; Spencer and Reber, 1962] to measure the concentrations of the major neutral-particle constituents of the earth's upper atmosphere. The ambient densities

of helium, oxygen, and nitrogen were obtained between north and south latitudes of 58° and from the perigee altitude of 257 km to an altitude in excess of 700 km. Data for various local times, altitudes, and geographic locations have been combined to provide a broader altitude coverage than could have been experienced at any single

TABLE 1.—*Tabulated Mass Spectrometer Data*

The local sun time, angle of attack (α) geographic latitude and longitude are averaged over the 4-minute pass. The stations involved are: BP, Blossom Point, Md.; COL, College, Alaska; FTM, Fort Myers, Fla.; GF, Grand Forks, Minn.; QUI, Quito, Ecuador; MOJ, Mojave, Calif.; NFL, Newfoundland; OOM, Woomera, Australia; JOB, Johannesburg, South Africa.

Pass and Station	Date	Local Time, hours	α	Geog. Lat.	Geog. Long.
15 BP.....	4/ 4/63	21.15	6°	38.5°	-75.0°
50 COL.....	4/ 6/63	0.65	16°	57.0°	-149.0°
80 COL.....	4/ 8/63	0.99	9°	55.0°	-147.0°
80 FTM.....	4/ 8/63	4.89	63°	18.0°	-92.0°
118 BP.....	4/10/63	18.81	70°	37.0°	-72.0°
120 GF.....	4/11/63	20.32	51°	51.0°	-98.5°
138 BP.....	4/12/63	2.51	12°	37.0°	-84.0°
152 BP.....	4/13/63	2.01	14°	39.5°	-68.5°
167 BP.....	4/14/63	1.65	20°	39.5°	-75.0°
182 BP.....	4/15/63	1.54	25°	37.0°	-78.0°
183 QUI.....	4/15/63	3.26	23°	4.5°	-79.0°
197 BP.....	4/15/63	1.43	27°	34.0°	-81.5°
211 BP.....	4/17/63	0.53	45°	41.5°	-71.5°
226 BP.....	4/18/63	0.48	53°	38.5°	-74.0°
241 BP.....	4/19/63	24.19	62°	38.0°	-79.5°
242 MOJ.....	4/19/63	0.64	54°	31.0°	-121.5°
254 NFL.....	4/20/63	22.75	82°	49.0°	-53.0°
270 BP.....	4/21/63	23.30	80°	41.50°	-71.5°
271 GF.....	4/21/63	22.88	85°	45.0°	-101.0°
708 NFL.....	5/20/63	7.18	39°	49.5°	-49.5°
795 OOM.....	5/26/63	15.81	63°	-34.0°	137.5°
800 JOB.....	5/26/63	15.90	65°	-37.5°	19.0°
888 JOB.....	6/ 1/63	13.24	33°	-27.0°	25.0°

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location. Data are presented for helium, atomic oxygen, and molecular nitrogen; molecular oxygen and atomic nitrogen data will be presented and discussed in a later report along with other data obtained under more varied conditions.

EXPERIMENT

The mass spectrometer measured the number densities of the various species which were in its ionizing, or sampling, region. Before reaching this sampling region, however, some of the neutral particles underwent collisions with various parts of the spectrometer structure; thus the number measured by the spectrometer was generally different from the ambient number density. The relationship between the spectrometer measurements and the ambient number densities was found to be a function of particle species, satellite velocity, and the angle between the spectrometer axis and the satellite velocity vector (angle of attack α). Using satellite velocity and angle of attack, both obtained from the tracking stations and the satellite's optical aspect system, calculations

have been carried out for each species and have been used to convert the measured quantities to the ambient values presented here. The contribution of residual gas has also been taken into account.

The absolute accuracy of the number density data is ± 40 percent, reflecting laboratory vacuum calibration error and the uncertainties in the ambient number density calculations. The relative accuracy of the data obtained for similar angles of attack is about 5 percent; for data obtained when the angles of attack differ by 70° or more, the relative accuracy is about 30 percent. The mean mass data and the concentration ratios are good to ± 10 percent. Since the relative accuracy of the number density varies with the angle of attack, this quantity, the local sun time, and other pertinent information are given in Table 1.

DATA

Figure 1 shows the ambient number densities of the major constituents plotted as a function of

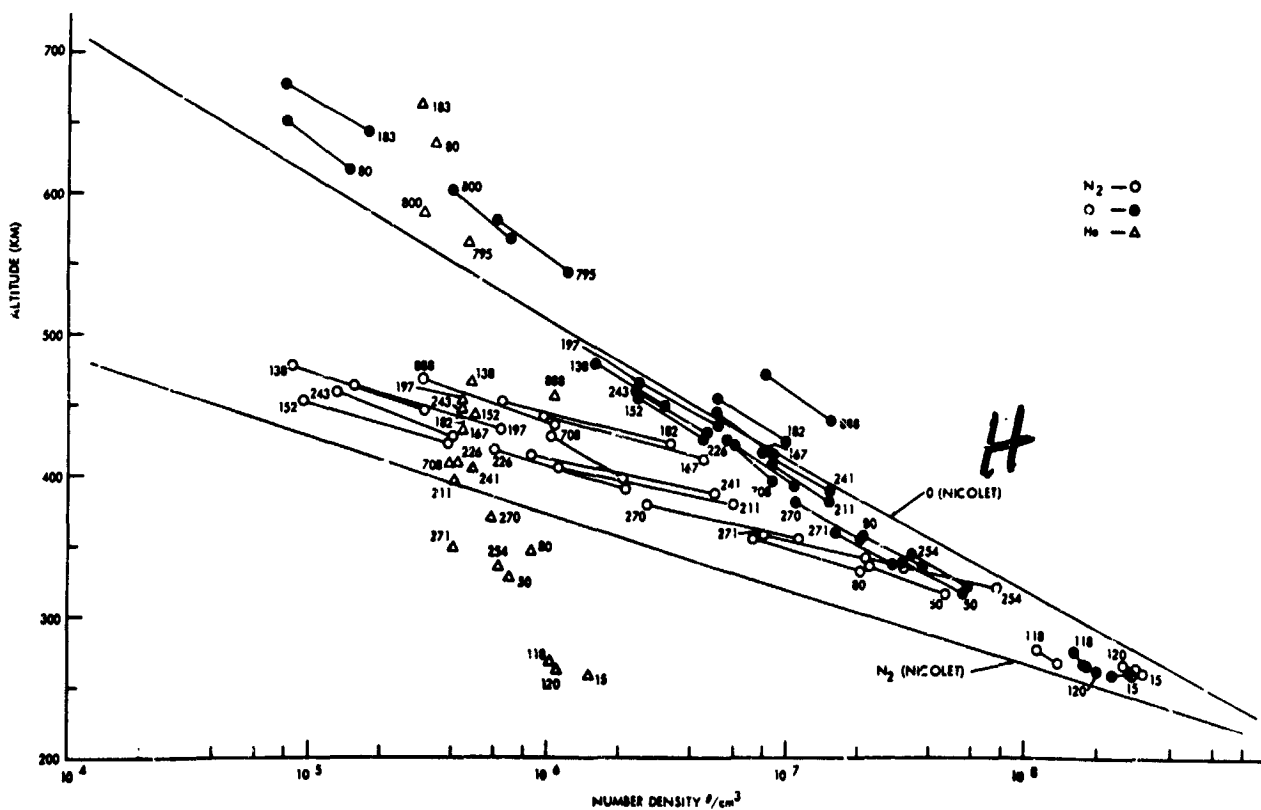


FIGURE 1.—Ambient number densities of the major constituents versus altitude. These data include many local times and geographic locations. The orbit numbers refer to Table 1. Nicolet's model for $T = 700^\circ\text{K}$ is included for reference.

altitude for a number of different local times and locations. A single satellite interrogation (pass) is indicated by two data points joined by a straight line for molecular nitrogen and atomic oxygen and by one point for helium. The Nicolet model (private communication) is included for reference.

For the times concerned, the number densities of helium and atomic oxygen are comparable at about 600 km, with helium predominant at higher

altitudes. Between 300 and about 600 km, atomic oxygen is the major constituent; in the 250- to 300-km region, molecular nitrogen and atomic oxygen have nearly equal concentrations.

There is, at present, insufficient daytime data to determine scale heights for the sunlit atmosphere. However, during the night at higher altitudes the concentration gradients are consistent with the scale heights for temperatures of 700° to 750°K.

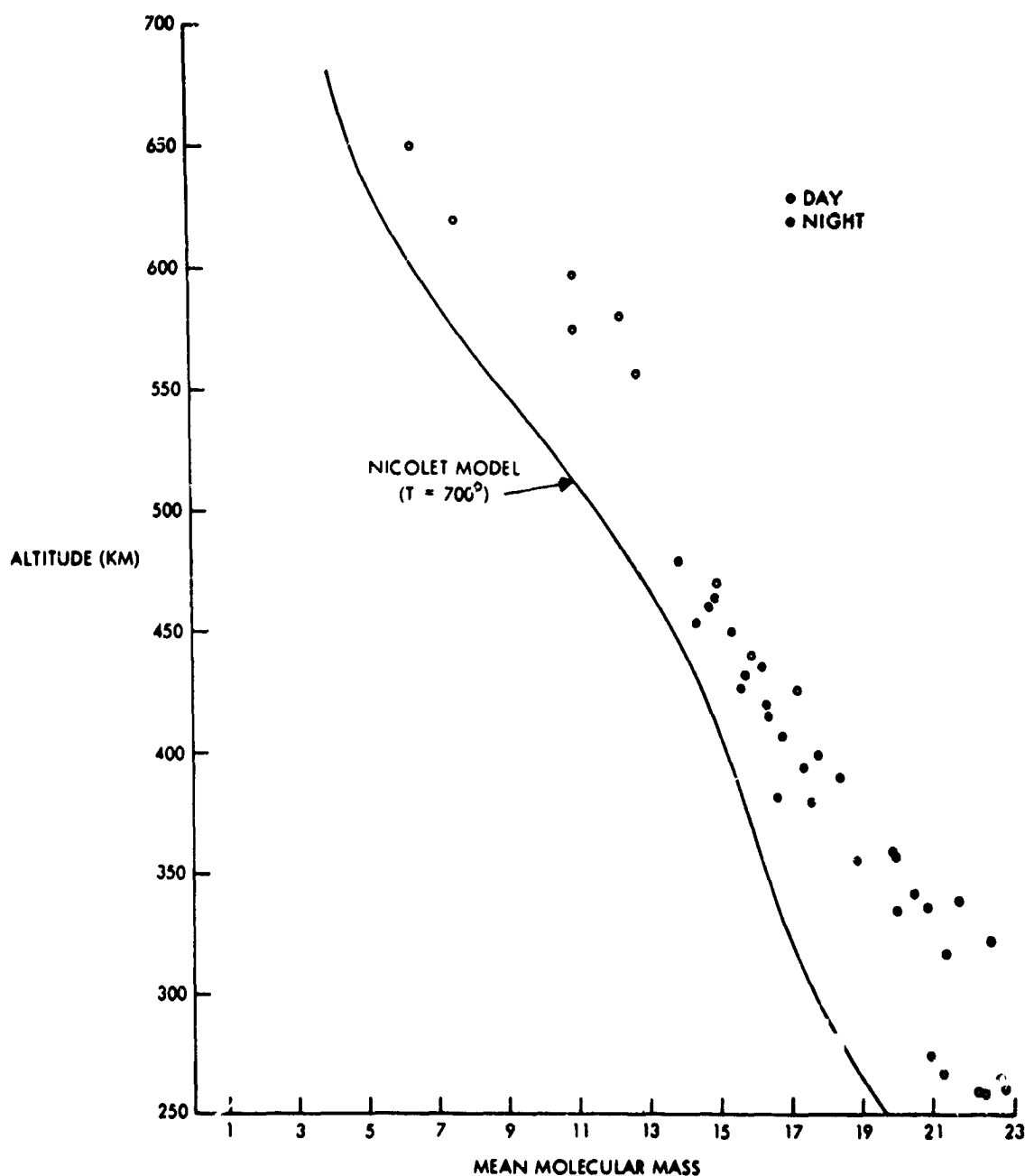


FIGURE 2.—Mean molecular mass versus altitude. The mean mass was calculated using the three major constituents measured, N_2 , O, and He. Nicolet's model is included for reference.

There appears to be a deviation from nighttime diffusive equilibrium in the 300- to 350-km altitude range for N_2 . This may be indicative of longer times required for N_2 to diffuse at these altitudes; however, as noted above, the data shown are representative of many time and geographic locations and may not provide an accurate instantaneous vertical distribution. It is interesting to note that the data from passes 167, 182, and

183, for which the densities are higher than for other nighttime passes, were obtained shortly after a minor magnetic disturbance which occurred on April 14. All data shown are measured values and are not averaged or smoothed in any way.

Figure 2 shows the variation of mean molecular mass with altitude. These data were calculated using the measurements of helium, atomic oxygen,

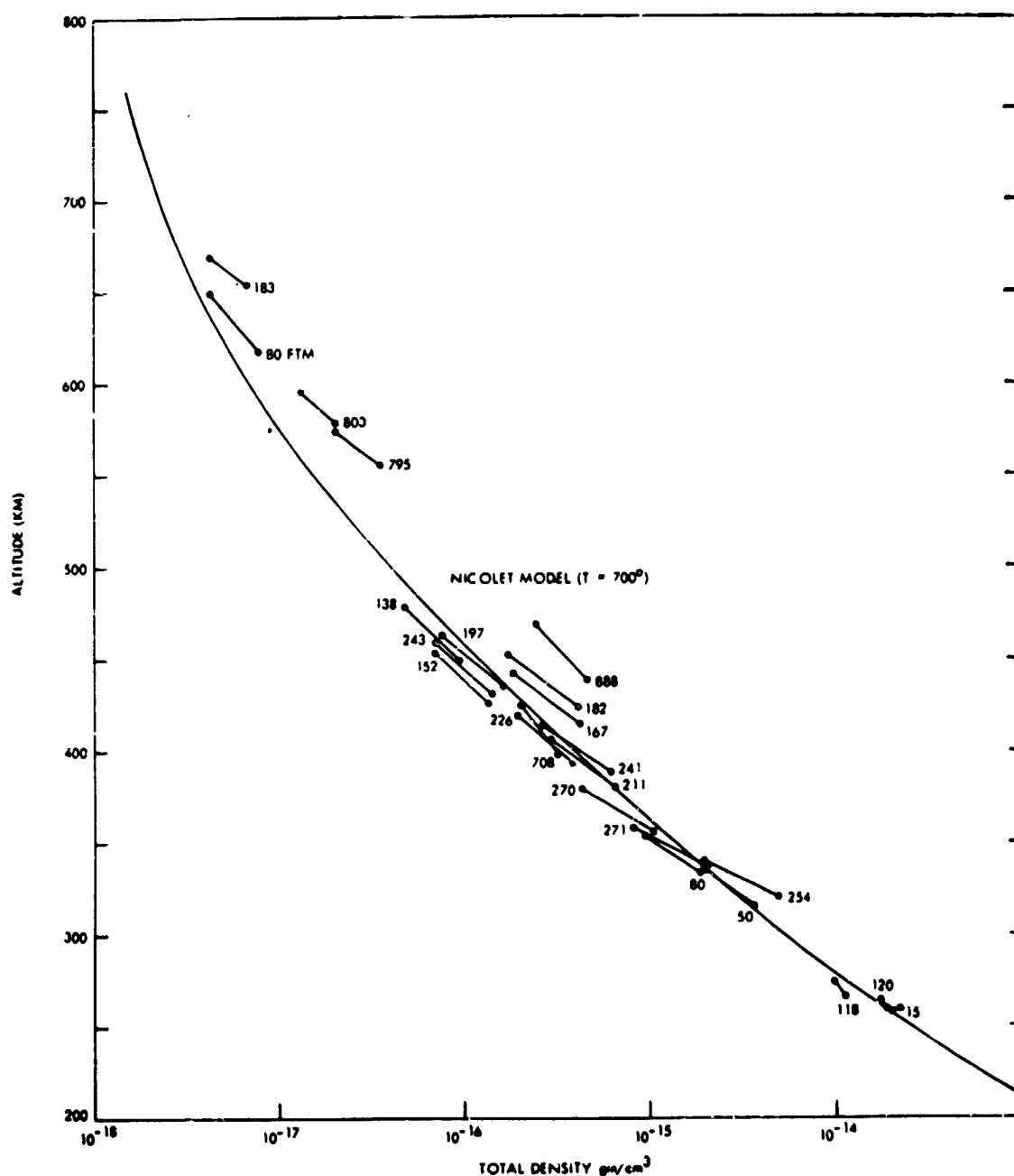


FIGURE 3. Total mass density versus altitude. The total mass density was calculated using the densities for N_2 , O, and He. The orbit numbers refer to Table 1. Nicolet's model is included for reference.

and molecular nitrogen only, other measured gases providing a negligible contribution to the mean mass. The presence of hydrogen, which the instrument was not designed to measure, would reduce the value of the mean mass somewhat at higher altitudes.

The total mass density, shown in Figure 3, was computed by summing the contributions from

molecular nitrogen, atomic oxygen, and helium. The results obtained in this way agree with the densities obtained independently by the pressure gage experiments on Explorer 17. This agreement lends support to the calculations relating the concentrations in the sampling region to the ambient number densities, as the pressure gage kinetics have been studied extensively [Horowitz,

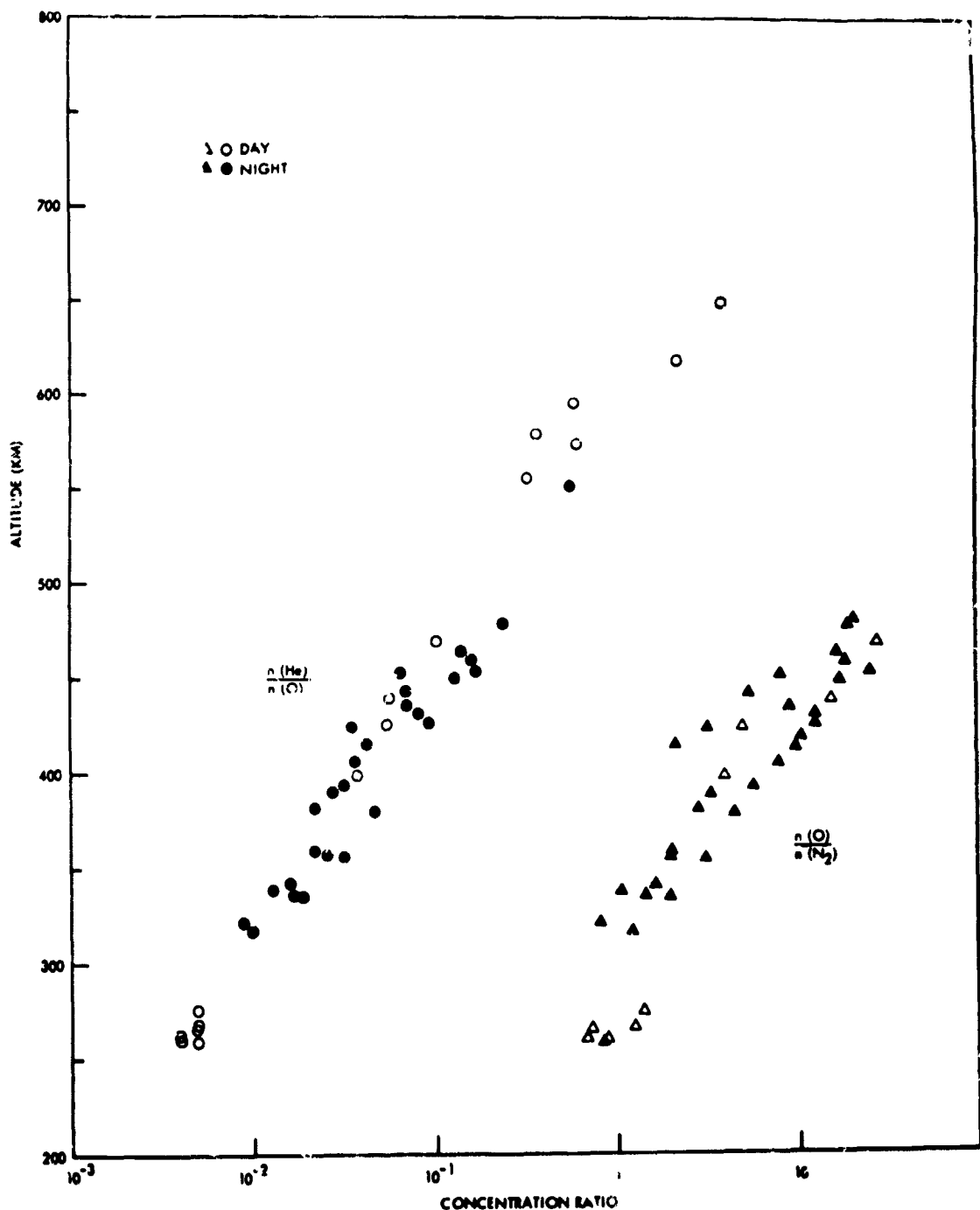


FIGURE 4.—Concentration ratios of the major constituents versus altitude.

1953]. Also, the total density measurements from the mass spectrometers agree in general within a factor of 2 with atmospheric density calculated from satellite drag observations [Bryant, 1964].

Figure 4 shows the variation with altitude of the concentration ratios of helium to atomic oxygen and atomic oxygen to molecular nitrogen.

ACKNOWLEDGMENT

I am indebted to Professor Marcel Nicolet for many hours of stimulating discussions about the data from Explorer 17, particularly with respect to extracting the scientifically interesting from the technically involved.

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164-33630

A NIGHT-TIME MEASUREMENT OF OZONE ABOVE 40 KM*

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The ozone distribution between 40 and 70 km was measured near midnight, May 27, 1960, from Wallops Island, Virginia by means of photometers sensitive to the ultraviolet airglow at wavelengths between 2400 and 2900 Å. Below 60 km, the densities are within a factor of two of the daytime photochemical equilibrium, as represented by Johnson's late afternoon measurement of June 14, 1949. Above 60 km, the ozone density increased with altitude, with its maximum increase, a factor of 6 over the day time value, occurring at 63 km.

INTRODUCTION

At high altitudes, above the principal ozone maximum, ozone concentration as a function of altitude should be governed principally by the presence or absence of sunlight, and vary in a predictable manner from day to night and from season to season [Chapman, 1930]^{a,b}. The first direct measurement of the daytime profile was a result of studies of the sun conducted by the Naval Research Laboratory on a V-2 rocket in 1949. Ozone densities up to 70 km were deduced from solar spectra and were consistent with computed photoequilibrium profiles [Johnson, et al, 1952]. Chapman also suggested that at high altitudes, above the ozone maximum and below the atomic oxygen maximum, ozone would increase at night as a result of a reaction between atomic and molecular oxygen. Several have treated this problem numerically, including Nicolet [1957], Barth [1961], Dütsch [1961], Paetzold [1961], Wallace [1962], and Hunt [1964], but with varying results, depending on the set of reactions, reaction rates, and initial concentrations which were chosen. Of particular difficulty is the computation of the effects of minor constituents such as hydrogen, nitrogen oxides, and the hydroxyl radical.

Ground based measurements of ozone content at these altitudes have not been satisfactory. Measurements of total ozone content are not particularly helpful since variations in the ozone

content below 30 km due to air movements are comparable to the expected night time increase at higher altitudes. However, the discovery of the ultraviolet airglow and a general improvement in the techniques of ultraviolet photometry made a night time measurement of ozone feasible.

In 1957, the Naval Research Laboratory flew an ultraviolet photometer with a response from 2600 to 2900 Å, and did observe an ultraviolet airglow layer centered at 101 km [Tousey, 1958]. This had been predicted from laboratory observations which showed that the Herzberg bands of molecular oxygen, the visible end of which had been observed in the airglow [Chamberlain, 1955], extend to 2563 Å in the ultraviolet [Broida and Gaydon, 1954]. But because an unknown amount of the observed airglow could be due to an OI line at 2972 Å, where the filter transmission is still 10 percent of its maximum; the ozone density could not be determined unambiguously.

In May 1960, Goddard Space Flight Center flew a number of ultraviolet photometers, including some whose filters were centered at 2620 Å and narrow enough so that the absorption cross section of ozone varied by only 50 percent over the bandwidth of the filter. These data, when interpreted with the aid of an airglow spectrum obtained by T. Stecher (of GSFA), provide an ozone density profile between 40 and 70 km.

INSTRUMENTATION

NASA Aerobee 4.05, one of a series of payloads designed for stellar photometry [Boggess, 1961],

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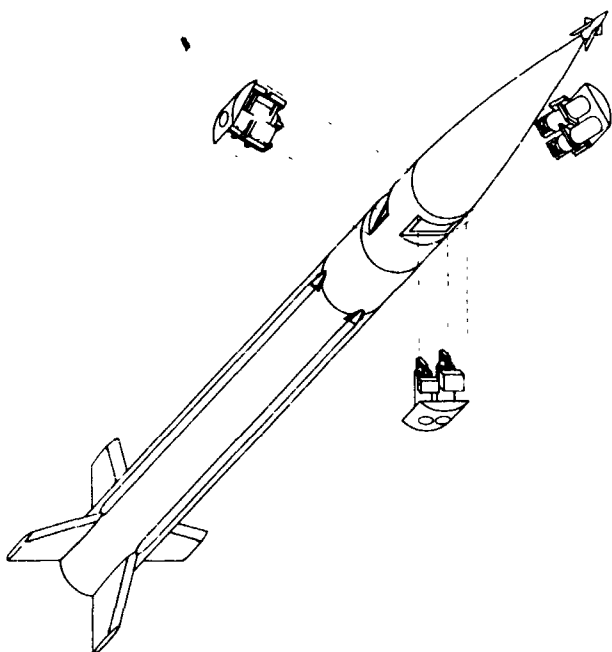


FIGURE 1.—Location of photometers on Aerobee 4.05.

contained three pairs of photoelectric photometers, each pair mounted 120 degrees apart around the rocket axis (See Figure 1) looking out at three different angles to that axis: nominally, 75°, 90°, and 105°. One photometer of each pair was sensitive to light in the spectral region centered near 2620 Å, while the response of the other was centered near 2260 Å, but with their optical axes parallel. Each pair was mounted on a removable door, which, when installed, became an integral part of the rocket skin. Since the principal purpose of this instrumentation was ultraviolet star spectroscopy, 2620 Å photometers were preferred to the 2680 Å ones (to reduce ambiguity in the interpretation of stellar data due to the strong magnesium doublet at 2800 Å). However, to correlate the data from this flight with that of earlier flights, one 2680 Å photometer was included. This was mounted on the door containing the 90° photometers, and also looked at 90° with respect to the rocket axis.

The optical system of each photometer was similar and is shown in Figure 2. Calcium fluoride was used for all lenses in the 2260 Å units; quartz for those in the 2620 and 2680 Å photometers. The field of view had a total width of between 4 to 5°.

Isolation of the 2600 Å region was achieved by combining two millimeters of 0.05 percent lead-

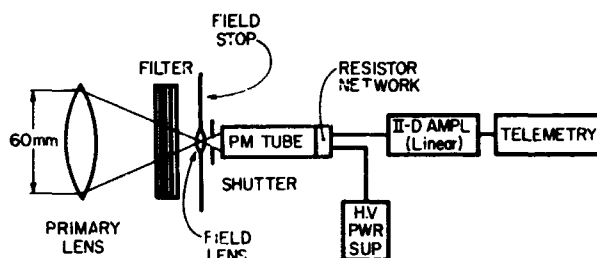


FIGURE 2.—Optical-electrical schematic of a photometer.

doped KC1:KBr (1:1) crystal with Cation-X in thin sheets of polyvinyl alcohol [Childs, 1961]. Three millimeters of nickel sulfate hexahydrate provided a sharp cutoff for longer wavelengths; one Corning No. 7-54 and one Corning 9-54 filter sharpened the shorter wavelength cutoff. A typical filter had a transmittance of 0.18, an effective wavelength of 2620 Å and a 200 Å bandwidth.

The 2700 Å filter consisted of three Corning 7-54 filters, one sheet of Cation-X and 5 millimeters of nickel sulfate hexahydrate. It has an effective wavelength of 2680 Å, a transmittance of 0.16, and a bandwidth of 320 Å.

The relative spectral response of the filters (see Figure 4) was measured by C. Childs, formerly of this laboratory, with a recording spectrophotometer, Cary Model No. 14, with an analytical accuracy of 1/2 of 1 percent for relative spectral transmission, a wavelength calibration of 4 Å, and a resolving power of 1 Å. Over the wavelength regions covered by each filter, the photomultiplier was assumed to have a constant sensitivity and the lenses a constant transmissivity. The rela-

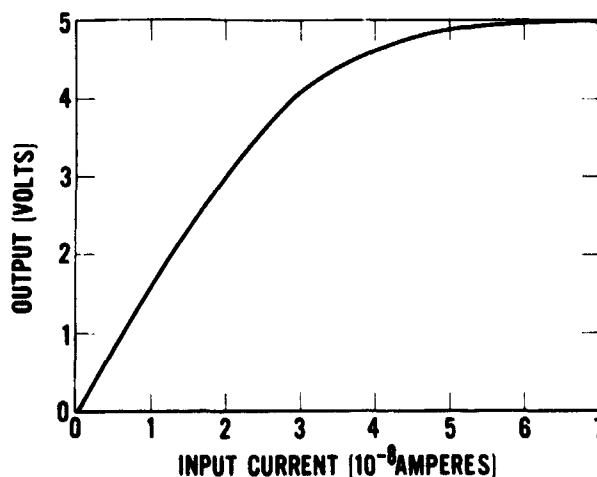


FIGURE 4.—Calibration curve for converter-amplifier.

tive response of the photometers was determined by use of the 2537 Å line, to which all three types of photometers respond with an easily measured signal.

The 2680 Å photometer used RCA's well-known 1P28 as a detector; all others used EMI's No. 6256B, an end-on fused silica window multiplier with cesium antimonide photocathode. A solenoid-operated shutter between the field stop and multiplier gave optical zero signals several times during the flight. The 1P28 and 6256B's were operated at 1000 and 1200 volts respectively; each detector had its own solid state DC to DC inverter power source with a resistor divider network at the base of each multiplier.

Each photomultiplier output was amplified to the zero to five volt range required by the telemetry system with what was essentially an impedance converter. Designed by G. Baker of the laboratory, the converter-amplifier had an electronometer tube (5886) input stage and ended with an emitter follower, with an overall voltage gain of 2.5 for small signals. It was purposely nonlinear in order to extend the dynamic range. A typical calibration curve is shown in Figure 3. The output voltage went directly to a pulse position modulation telemetry transmitter which relayed the data to the ground receiving equipment.

BASIC DATA

Aerobee-Hi NASA 4.05 was launched at 0030 EST on May 27, 1960 from Wallops Island, Virginia (37°50' N, 75°29' W). The vehicle performance was normal: propulsion ended at 52.4

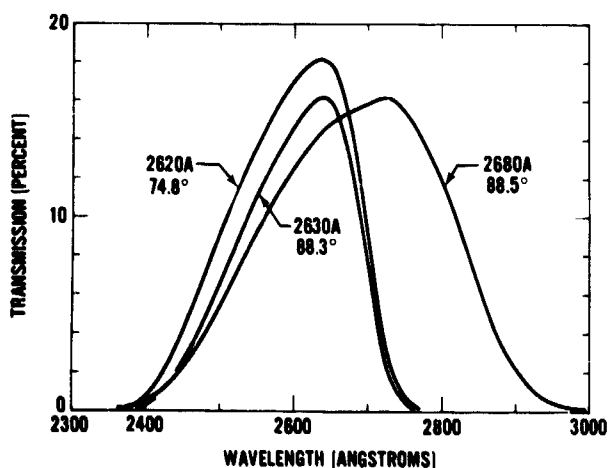


FIGURE 3.—Spectral characteristics of flight filters.

sec after launch at an altitude of 36.6 km with a vertical velocity of 1.84 km/sec and a horizontal velocity of .260 km/sec at an azimuth of 111°. The rocket spun about its longitudinal axis at a rate of 2.16 rps and soon entered a precession cone of 5.7° half angle, whose axis was 15.6° from zenith at an azimuth of 96°, with a period of 75 sec. Aspect during the ascent and free-fall portions of the flight was determined from a combination of data from magnetometers, trajectory information, and horizon and star data from the photometers. Down-leg aspect could not be determined with useful accuracy below the free-fall region. A peak altitude of 215.3 ± 0.2 km was reached 249.4 seconds after launch. Telemetry ceased at 468 sec; no recovery of instrumentation was attempted.

In Figure 5 is a sample of the telemetry record from two of the photometers for a period corresponding to three revolutions of the vehicle, while the airglow layer is still above the rocket. Since these photometers are nearly perpendicular to the rocket axis, the zenith angle of the photometer axis changes from a minimum (76° for the time in Figure 5) as the photometer points skywards, through 90° as it scans the horizon, and to a maximum of 100° as it points earthward. Ozone is relatively opaque around 2600 Å, and the airglow in Figure 5 can be seen only when the zenith angle of the photometer is near a minimum. The ozone is relatively transparent to the light passed by the 2260 Å filter, and the brightening of the airglow at the horizons can be clearly seen. (The light is probably of wavelengths longer than 2700 Å, passed through the long wavelength tail of the 2260 Å filter.) The southern horizon appears wider because it is merged with several bright stars in the Milky Way. As the vehicle increased in altitude, the signal from the 2620 Å photometer resembled that from the 2260 Å photometer, with bright horizons, a less bright sky toward zenith, and a dark earth. Above the airglow layer, the sky was dark (except for the brighter stars) and the earth appeared light. The noise in the record is due partly to the photomultiplier and partly to stars. The records were read at the midpoint between horizons, and analyzed to yield both a distribution of ozone with altitude and the volume emission of the airglow versus altitude.

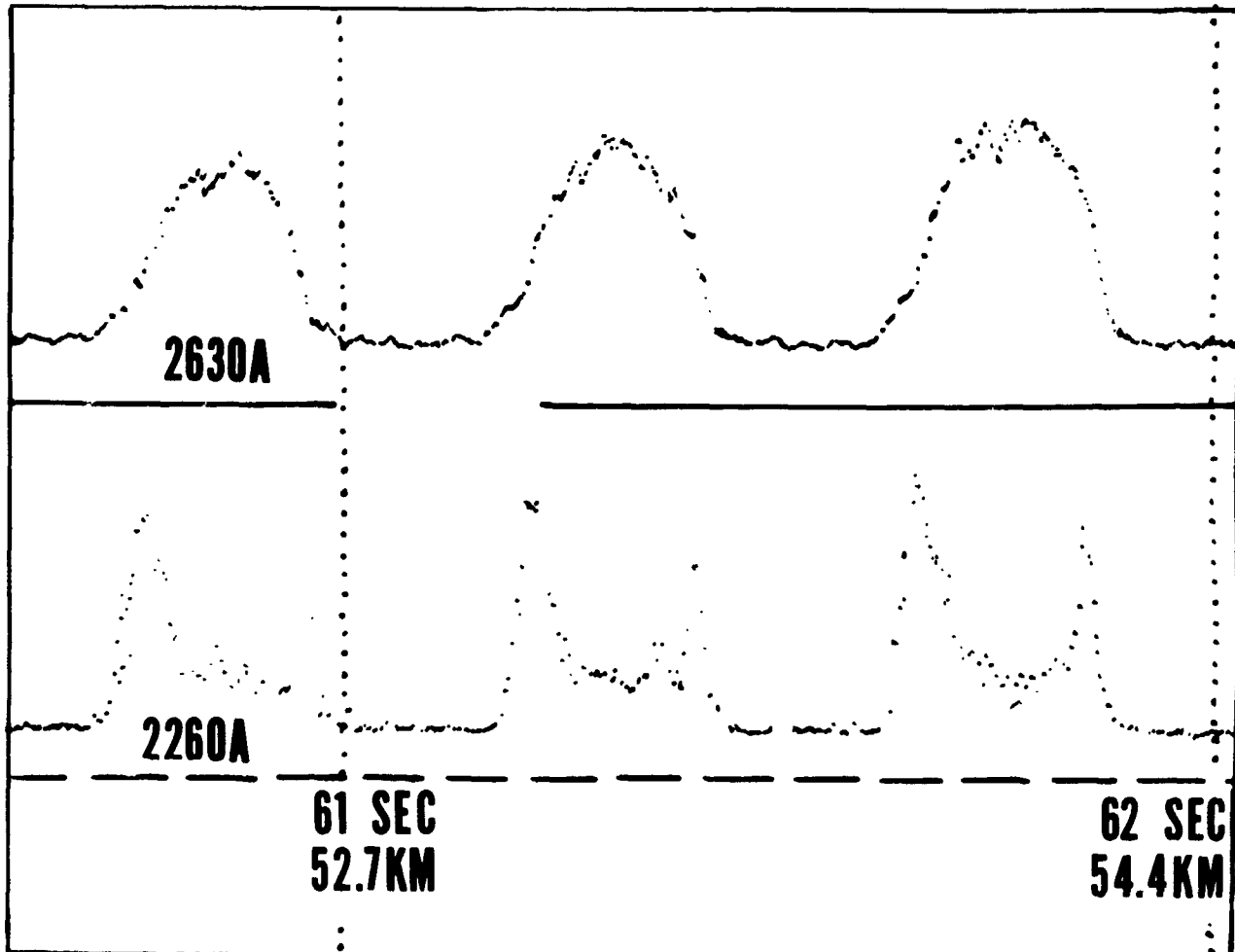


FIGURE 5.—Sample of telemetry record showing roll modulation of the airglow signal. Increasing signal represents increasing light.

The data obtained in the region of interest is shown in Figure 6. Data from 3 of the 7 photometers proved to be useful for ozone measurements. The pass band of the three 2260 Å photometers was too wide to permit an accurate determination of an effective cross section for ozone; the down looking 2620 Å photometer could not see the airglow while the vehicle was below the airglow layer. Data points for the photometers are shown to indicate the scatter in the raw data. As would be expected, the scatter increased rapidly as the signal rose into the non-linear portion of the amplifier response curve. In addition to the data, the angle of the rocket's longitudinal axis with respect to local zenith is given. The rocket took a spiral path with a zenith angle of 8° at thrust termination (52.4 sec) until it entered

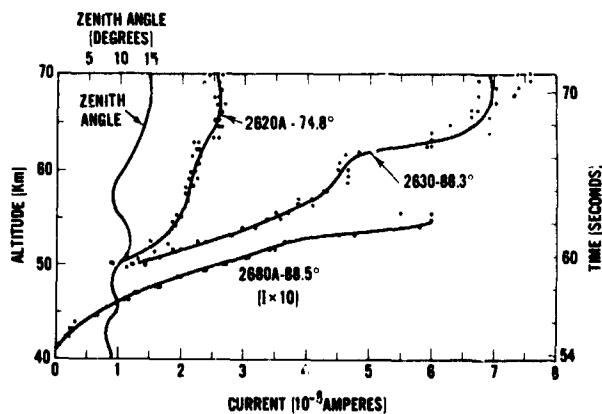


FIGURE 6.—Data from three photometers during the ascent of the rocket. The zenith angle is the angle between the longitudinal axis of the rocket and local zenith; the angle stated for each photometer is the angle between the photometer's optical axis and the rocket's longitudinal axis.

its regular precession cone of motion at about 75 seconds.

OZONE

The ozone content of the atmosphere is obtained from the rate of increase of the airglow signal as the vehicle rose. The energy observed is related to the ozone density in the following manner:

$$n(O_3) = \frac{(\log E_2 - \log E_1)n \cos \gamma}{a(h_2 - h_1)}$$

where $n(O_3)$ is the number of ozone molecules per cm^3

h_1 and h_2 are the lower and upper ends of the altitude interval,

E_1 and E_2 are energies observed at the corresponding altitudes,

n is Loschmidt's number, $2.687 \times 10^{19} \text{ cm}^{-3}$,

γ is the angle between the photometer axis and zenith, and

a is the absorption coefficient, cm^{-1} , base 10.

This was applied to the smooth curves in Figure 6 at one-second intervals.

Since the cross section of ozone varies appreciably over the wavelength interval passed by each filter, some assumption must be made concerning the spectrum of the airglow. For this purpose, the spectrum observed by T. Stecher (private communication) was used, and the absorption coefficient, a , computed for each photometer where

$$a = \frac{\sum E_i R_i a_i}{\sum E_i R_i}$$

where E_i is the energy in a particular wavelength interval, i ,

R_i is the corresponding relative transmission of the filter, and

a_i is the corresponding absorption coefficient.

The absorption coefficients used were from the tabulation of *Inn* and *Tanaka* [1953, 1959] which in the region of interest here are about 10 percent higher than those of *Vigroux* [1953] and 1 or 2

percent lower than those of *DeMore* and *Raper* [1964]. The effective absorption coefficients calculated on this basis are: 2620 Å (88.3°), 114.4 cm^{-1} ; 2620 Å (74.8°), 113.8 cm^{-1} ; and 2630 Å (88.5°), 79.48 cm^{-1} .

The resulting ozone density is shown in Figure 7 as a function of altitude. It is felt that the spectral characteristics of the photometers, the airglow and the ozone absorption coefficient are sufficiently well known such that they contribute no more than a total of about 20 percent uncertainty to the number density of ozone. This would be a systematic error which would not affect the shape of the curve.

The angle of the photometers with respect to zenith could easily contribute an uncertainty of 20 percent to both the absolute and relative values of the ozone densities. It is based upon magnetometer data and the assumption that at the end of thrust the rocket axis was aligned with the velocity vector.

The largest source of error is in the character of the data, which contains noise from the photomultiplier dark current, stray pulses, and stars. It is difficult to separate this from the possible temporal and spatial variations of the airglow itself. A temporal variation could be responsible for some of the shape of the curve, but it is unlikely that the airglow would vary sufficiently in the 16 seconds of time that the ozone curve represents, to be responsible for its major features. Spatial variation is not thought to be a major

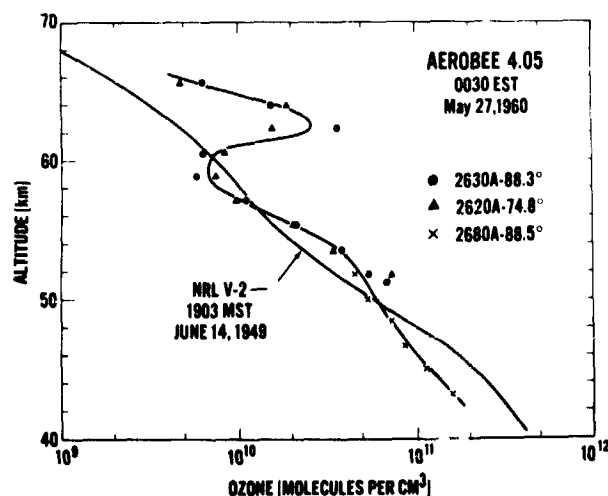


FIGURE 7.—Altitude distribution of ozone.

source of uncertainty in the data since the numbers derived from the two photometers which were looking at different portions of the sky agree reasonably well. The magnitude of the noise-like errors is apparent from the scatter of points about the curve and is on the order of 50 percent.

For comparison the daytime ozone distribution as measured by absorption of the solar spectra between 2500 and 3400 Å [Johnson, *et al*, 1952] is shown. The daytime profile was computed using the ozone absorption coefficients of *Ny* and *Choong* [1933], which in the spectral region of interest are 10 to 15 percent higher than those of *Inn* and *Tanaka* used for the night time profile.

The most important feature of the profile is the factor of 6 increase in ozone density over the daytime profile above 60 km. While this may be in error by 50 percent or more, it is believed that the shape of the curve does indicate an increase of ozone density at night in this region, and that it is on the order of a half of a magnitude.

Techniques are being developed by various workers to use satellites for the measurement of ozone in this region. *Venkateswaran* [1961], observed the sunlight reflected from *Echo I* as it emerged from the earth's shadow, using various wavelength pairs between 4700 and 7000 Å. His results above 55 km are about a factor of 15 higher than the measurements by *Johnson, et al*. However *Venkateswaran* [1963] states that this method probably gives too high values at levels above the principal ozone maximum.

The second type of observation was made at sunrise and sunset by a satellite borne radiometer with a response center at about 2630 Å [*Rawcliffe, et al*, 1963]. At 60 km his data are about 20 percent lower than *Johnson's*, but approach *Johnson's* data, and above 80 km are somewhat higher than the trend of *Johnson's* data.

It is expected that the night time values would be higher. The magnitude of the effect, besides depending on reactions among the various oxygen species and third bodies, depends critically on such things as the initial hydrogen concentrations chosen [*Bates and Nicolet*, 1950, and *Wallace*, 1962] and possible reactions involving atomic nitrogen [*Barth*, 1961]. Perhaps the most recent computation of ozone densities has been done by *B. G. Hunt* [1964] using an atmosphere in which he assumes the only reactive constituent is oxy-

gen. (The effects of atomic nitrogen and hydrogen would be to lower the calculated O and O₃ concentrations.) Between 40 and 50 km, *Hunt's* curve for just before sunrise conditions is as much as 30 percent lower than *Johnson's* daytime profile, crosses it at 53 km, and reaches a maximum value of 5×10^{10} molecules of O₃ per cm at 69 km.

AIRGLOW

The other principal result from the analysis of data from these photometers is information concerning the distribution of the airglow. Volume emission can be deduced from the data obtained as the vehicle passes through the emitting region. The energy calibration of the photometers has been used in preparing these curves, (Figure 8) so that they do rightly represent the relative energy in the portion of the spectrum passed by the different filters. Ten arbitrary units represent on the order of one photon cm⁻² sec⁻¹ per Å. The airglow as measured by the 2680 Å photometer is 2.3 times that sensed by the 2620 Å photometers. The airglow measured by the 2260 Å photometers is 0.15 times that of the 2620 Å photometers; nearly all the energy measured by the 2260 Å photometers has come through a long wavelength tail of the filters. This pattern is completely consistent with a spectrum of the airglow horizon obtained by *T. Stecher* with a spectrograph flown at 0030 local time on July 19, 1963.

The 2680 Å filter was similar in construction and characteristics to those flown in March 1957 [*Tousey*, 1958] and November 1959 [*Friedman*, 1961, *Packer*, 1961]. The altitude of maximum

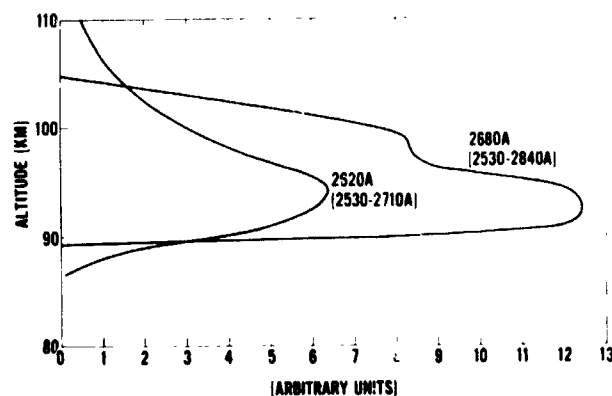


FIGURE 8.—Altitude distribution of the ultraviolet emission.

emission were 101 and 95 km respectively, compared to 92 km for this flight. The zenith intensity for the 1957 flight was 3.4 rayleighs per Angstrom (Dunkelman, private communication) and for the 1959 flight was 1.94 rayleigh per Angstrom. The airglow during the flight of Aerobee 4.05 was somewhat brighter, but within a factor of 10 of these values.

The brighter stars and the Milky Way were readily noted in the records as the photometer scanned across the sky. The signal in the absence of obvious stars indicated that less than 15 percent of the light from extended sources originated above the emission layer indicated in Figure 8.

ACKNOWLEDGMENTS

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SOME ASPECTS OF WIND SHEAR IN THE UPPER ATMOSPHERE*

R. G. ROPER

The wind motions responsible for the shearing of sodium vapor trails ejected from rockets in the 70 to 140 km region of the upper atmosphere are subjected to an analysis based on generally accepted theories of hydrodynamic turbulence. The region from 80 to 100 km is of particular interest in that here the predictions of shear turbulence theory are well substantiated. The energy spectrum of the height shear is found to follow the 4/3 power law proposed by Tchen, and is associated with a vertical correlation distance of approximately 6 km. The existence of an isotropic inertial region of maximum scale 3 km, previously indicated by analysis of meteor data, is confirmed. The vertical scale of the turbulent eddies is found to be the atmospheric pressure scale height, a phenomenon which has been observed by others, but which, as yet, has no satisfactory explanation.

1. TURBULENCE THEORY

The complete development of the relationships used in the analysis of wind shear is beyond the scope of this work; the following, used in conjunction with the references quoted, should provide an adequate background for consideration of the subsequent analysis.

1.1 Energy Spectrum Analysis

If there exists in a turbulent flow field a range of scales which receive energy from larger scale motions and pass it on undiminished to smaller scale motions, then, for this so-called inertial (non-dissipative) range of scales, the only form of energy spectrum function dimensionally possible is

$$E(k) = \alpha \epsilon^{2/3} k^{-5/3} \quad (\text{Kolmogoroff, 1941})$$

where ϵ is the rate at which the turbulent energy is received by (and leaves) the inertial range of scales, k is the wavenumber vector corresponding to the real space scale r , and α is an absolute constant of order unity.

Batchelor (1953) has shown that, for such an inertial region which also possesses the property of isotropy, the fluid velocity differences measured as a function of the separation r follow the relation

$$[u(x) - u(x+r)]^2 = 4.82\alpha(\epsilon r)^{2/3} \quad (1)$$

In real space, the energy spectrum function $E(r)$ is defined by

$$E(r) = [u(x) - u(x+r)]^2 \quad (2)$$

Tchen (1954) has considered an otherwise isotropic region subjected to a mean shear, and finds that equation 1 is modified, becoming

$$[u(x) - u(x+r)]^2 = a r^{4/3} \quad (3)$$

where a involves α , ϵ , and the mean gradient. Thus, for what may be termed shear turbulence, in real space

$$E(r) \sim r^{4/3} \quad (4)$$

1.2 Correlation Analysis

An energy spectrum function equivalent to that based on velocity differences can be formulated from the lateral or longitudinal velocity correlations defined as

$$f(r) = \frac{u_f(x)u_f(x+r)}{u_f^2}$$

and

$$g(r) = \frac{u_n(x)u_n(x+r)}{u_n^2} \quad (5)$$

where $u_f(x)$, $u_f(x+r)$, $u_n(x)$, $u_n(x+r)$ are the turbulent components of the velocity at two points x and $x+r$ respectively, measured parallel (suffix f) and normal (suffix n) to the vector separation r .

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In isotropic turbulence, $u^2 = u_n^2 = u_0^2$, the velocity characteristic of the energy bearing eddies of the Kolmogoroff spectrum.

Introduction to the equation of continuity for incompressible fluids leads to the relation

$$g(r) = f(r) + \frac{1}{2} r \frac{\partial f}{\partial r} \quad (6)$$

(von Karman and Howarth, 1938)

between the functions f and g , or, if the turbulence is isotropic in two dimensions only,

$$g(r) = f(r) + r \frac{\partial f}{\partial r} \quad (7)$$

These functions, f and g , may be called Eulerian space correlation functions, and either may be denoted by $R(r)$. If the range of scales under observation is inertial, then $R(r)$ must depend only on ϵ , and the only form dimensionally possible is

$$u_0^2 [1 - R(r)] \sim \epsilon^{2/3} r^{2/3}$$

$$\text{i.e.} \quad 1 - R(r) \sim r^{2/3}$$

(for any given flow field, u_0 and ϵ are constant).

Introduction of the relations 6 and 7 above gives, with c a constant

$$f = 1 - cr^{2/3} \quad (8)$$

$$\text{and} \quad g = 1 - 4/3 cr^{2/3} \quad (9)$$

for three dimensional isotropy

$$\text{or} \quad g = 1 - 5/3 cr^{2/3} \quad (10)$$

for isotropy in two dimensions only.

Since the variations of the correlation difference function $[1 - R(r)]$ with r is the same as that of $E(r)$ in equation 1, we may suppose that the dependence of $[1 - R(r)]$ on r will be similarly modified in the presence of a mean shear

$$\text{i.e.} \quad [1 - R(r)] \sim r^{4/3} \quad (11)$$

If, in fact, equation 11 is pertinent, then the relationships between f , g , and r (equations 8, 9, 10) will now become

$$f = 1 - ar^{4/3}$$

$$\text{and} \quad g = 1 - 5/3 ar^{4/3}$$

$$\text{or} \quad g = 1 - 7/3 ar^{4/3}$$

An indication of the degree of isotropy can be obtained by considering the ratio

$$S = \frac{1-f}{1-g}$$

For two dimensional isotropy

$$S = 0.60 \text{ without mean shear}$$

$$\text{and} \quad S = 0.43 \text{ with mean shear}$$

For three dimensional isotropy

$$S = 0.75 \text{ without mean shear}$$

$$\text{and} \quad S = 0.60 \text{ with mean shear}$$

The relative importance of mean shear will be indicated by the form of either the energy spectrum function $E(r)$, or the correlation difference function $[1 - R(r)]$.

2. THE PRACTICAL APPLICATION OF TURBULENCE THEORY

2.1 The Mean Wind Profile

In applying the relations developed above to the wind vectors measured by means of sodium vapor trails ejected as a tracer into the upper atmosphere, the relative importance of the mean motion must not be overlooked. The velocities used in the energy spectrum and correlation analyses of the previous section must be the turbulent velocities, or departures from the mean motion. In normal correlation analysis, the mean value of a set of observations is usually taken as the mean value of the measured data, which, when applied to wind height shear data, would be tantamount to the assumption of a mean wind profile which is constant with height. Such a profile is the exception rather than the rule in meteorological phenomena.

Practically any attempt to prescribe a mean wind profile to the measured winds will be subjective to a certain extent. If one uses a polynomial fit, for example, it must be truncated before it can assume any of the features of the measured profile which are due to the turbulent motions present. From experience based on the measurement of winds in the 75 to 105 km region by means of radio reflections from meteor trails (Elford, 1958, 1964), a quadratic change with height of the mean wind over any given 20 km interval should best

describe the contribution of the mean motion without destroying any of the characteristics of the turbulent flow field. In the present analysis, the windspeed/azimuth data are converted to zonal and meridional components, and a polynomial profile of order $Z+1$, where

$$Z = \frac{\text{Total height range covered by data (km)}}{20},$$

is fitted to each. The profiles thus determined are subtracted from the relevant measured profiles to give zonal and meridional turbulent velocities.

2.2 The Determination of $E(\Delta h)$

The available sodium trail data lists wind speed and azimuth against height, and usually involves irregularly spaced height intervals. The sampling irregularity exists for two reasons:

- there is an occasional difficulty in absolutely identifying the same point on the trail in consecutive photographs;
- the wind profile between consecutive observational heights is linear. This is usually obvious from the photographs, and can easily be allowed for in subsequent analysis.

Whereas equal height interval sampling is not absolutely necessary for subsequent reduction, it does simplify the analysis, and so linear interpolation between the listed data points is used to provide a profile with data points spaced 0.2 km in height.

The energy spectrum function is computed as

$$E(\Delta h_i) = \frac{1}{N} \sum_{k=1}^N (u_{k+i} - u_k)^2 \quad (12)$$

in which

$$N = 5(H_2 - H_1) - i$$

where $i = 1, 2, 3, \dots, N$

such that $\Delta h_1 = 0.2$ km.

$$\Delta h_2 = 0.4 \text{ km}$$

etc.

and H_1, H_2 (in km) are the lower and upper bounds respectively of the region for which $E(\Delta h)$ is being determined. Such partitioning of the height range is necessary since there is considerable variation in the characteristics of the flow

over the total height range sampled (usually some 70 to 200 km).

Energy spectrum functions may be calculated using

- zonal turbulent velocities
- meridional turbulent velocities
- turbulent windspeed.

The turbulent windspeed w is defined here as

$$w = \text{measured windspeed} - \\ ((\text{mean zonal wind})^2 + \\ (\text{mean meridional wind})^2)^{1/2}$$

2.3 Correlation Analysis

The two correlation functions f and g defined in 1.2 refer to turbulent velocity components measured parallel and normal to the separation vector. Since vertical velocities in this region of the upper atmosphere are so much less than the associated horizontal components, the magnitude of the vertical component cannot be determined from sodium trail photographs. (Most workers in this field consider the upper limit for mean plus random vertical motions to be some 10 metres/sec). However, we may redefine f and g in terms of the orthogonal zonal and meridional flow fields, with a view to investigating possible isotropy. This has been done by meteorologists in the past, with at least partial success (see, for example, Hutchings (1955)).

The normalizing factors $\bar{u}_f^2$ and $\bar{u}_g^2$ are best estimated by the standard correlation function definition. The correlation functions f and g then become

$$f(\Delta h_i) = \frac{\sum_{k=1}^N u_{k+i} u_k}{\left[\sum_{k=1}^N u_{k+i}^2 \sum_{k=1}^N u_k^2 \right]^{1/2}} \quad (14)$$

where u are the zonal turbulent wind velocities and

$$g(\Delta h_i) = \frac{\sum_{k=1}^N v_{k+i} v_k}{\left[\sum_{k=1}^N v_{k+i}^2 \sum_{k=1}^N v_k^2 \right]^{1/2}} \quad (15)$$

where v are meridional turbulent wind velocities, and N and i are as previously defined for equation 12.

3. PRELIMINARY RESULTS

The analysis of Section 2 has been applied to data obtained from a sodium trail release over the Eglin Air Force Base, Florida (29.6° N, 86.6° W) at 1910 CST on May 21, 1963 (Edwards et al., 1963). In this experiment, wind speed and azimuth were obtained over a height range of 69 to 140 km.

3.1 The Mean Wind Profile

Since the data covers the height range from 69 to 140 km, polynomials of order 4 are fitted to the zonal and meridional measured profiles. These yield mean zonal and meridional profiles

$$u_{\text{mean}} = 38.6 - 126h - 244h^2 + 174h^3 + 205h^4$$

$$v_{\text{mean}} = -30.4 + 19.8h + 148h^2 - 16.8h^3 - 125h^4$$

where u , v are in metres/sec and h is the normalized height given by

$$h = (2z - z_{\min} - z_{\max}) / (z_{\min} - z_{\max})$$

where z is the height variable

$z_{\max}$ the maximum and

$z_{\min}$ the minimum heights of the available data, all heights being in kilometers.

Normalization of the height range stabilizes the least squares fitting process, and makes the relative importance of the individual terms of the fitted polynomials more obvious than is the case when the mean velocities are expressed as power series in the actual height z . Results are plotted in Figs. 1 and 2.

The meteorological significance of these profiles, in particular the reversal of both the zonal and meridional components above 110 km, cannot be evaluated from consideration of this single firing.

3.2 The Turbulent Wind Profile

As mentioned in Section 2, the 70 to 200 km height range covers a number of characteristically different regions. The wind motions observed below approximately 105 km indicate the presence of small-scale structure, while those above 110 km do not appear to be at all turbulent, even though vertical shear is present. In the results presented here, discussion is confined to the consideration of

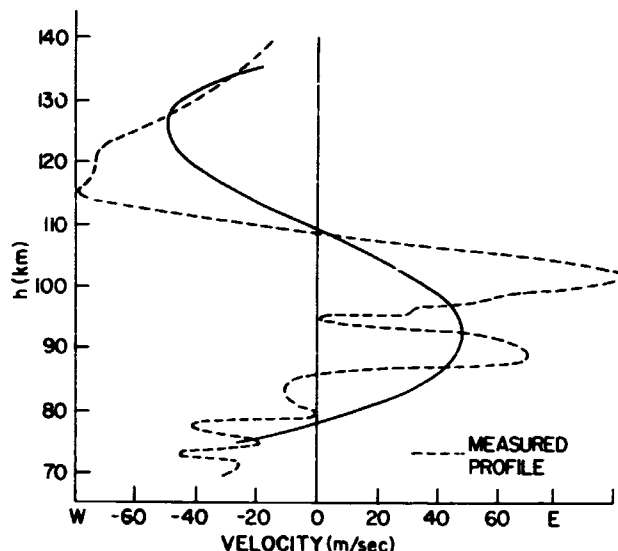


FIGURE 1.—Mean zonal profile.

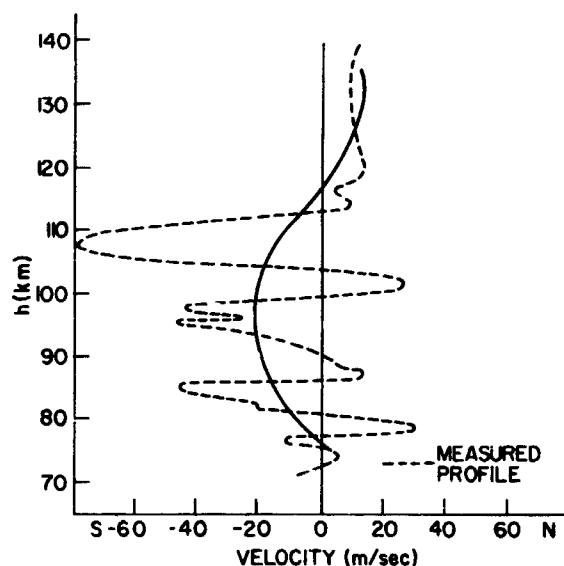


FIGURE 2.—Mean meridional profile.

the region from 80 to 100 km, which has been found to be representative of a turbulent region which can be adequately described by available statistical theories of hydrodynamic turbulence (Elamont and Jager, 1961; Zimmerman, 1962; Roper, 1962).

The measured zonal and meridional profiles, and the deviations from the mean wind for the height range 80 to 100 km are plotted in Figures 3 to 6. Whereas the immediately obvious wave-like nature of the turbulent profile would suggest a wave theory approach as likely to be the most profitable for consideration of the wind motions

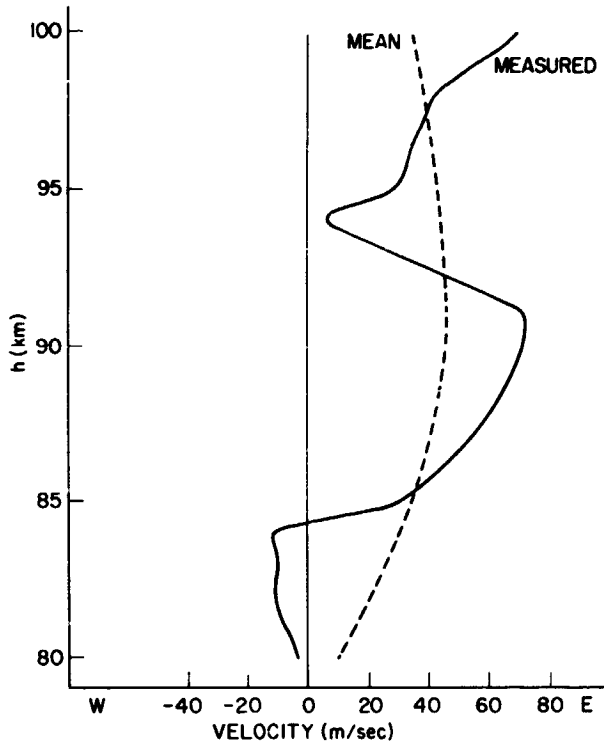


FIGURE 3.—Zonal profile, 80-100 km.

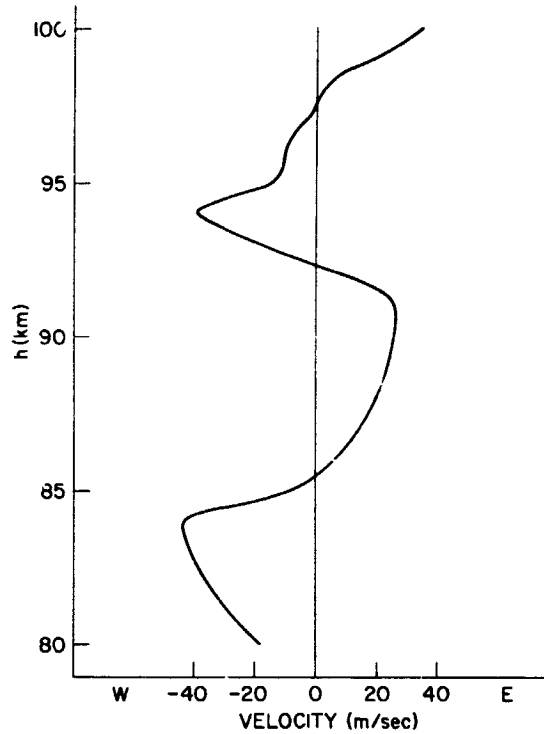


FIGURE 5.—Turbulent zonal profile.

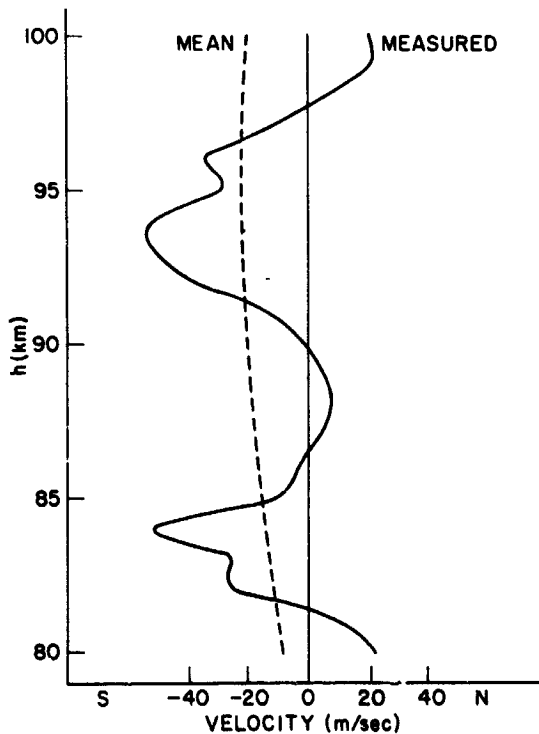


FIGURE 4.—Meridional profile, 80-100 km.

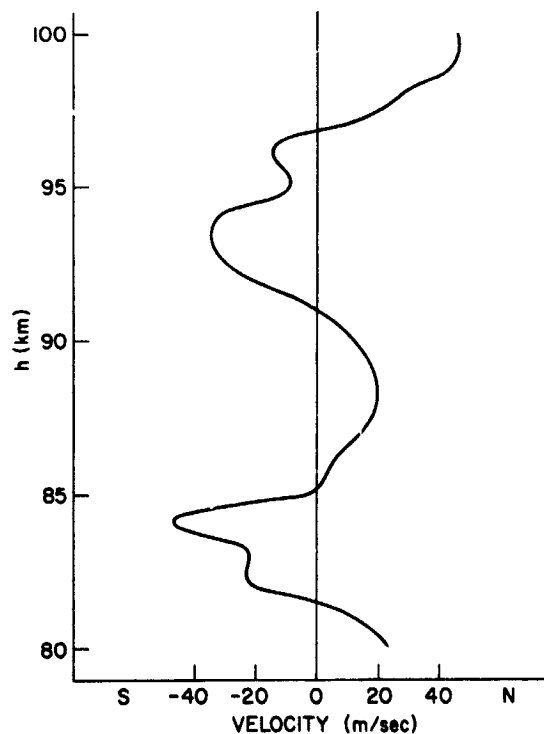


FIGURE 6.—Turbulent meridional profile.

in this region, applications of available wave theories (e.g., Hines, 1959) have not produced consistent results. The possibility of generation of turbulence in the 80 to 100 km region by vertically propagating gravity waves has been proposed by Hines (1963). The purpose of the present work, however, is not to determine the source of the turbulent energy, but rather to substantiate the evidence that the observed shears characterize a region of hydrodynamic turbulence.

3.3 The Energy Spectrum Functions

The energy spectrum functions for the zonal and meridional turbulent wind profiles for the 80 to 100 km region have been computed using the methods described in Sections 2.2 and 2.3. As can be seen from Figures 7 and 8, the velocity difference spectrum $E(\Delta h)$ and the correlation difference function $[1 - f(\Delta h)]$ for the zonal turbulent velocities are completely equivalent. This is not surprising, since both functions are solutions of the same equation

$$F = a\Delta h^m$$

(see Sections 1.1 and 1.2).

Similarly for the meridional spectra of Figures 9 and 10. The slope m of the log log plots of both the zonal and meridional $E(\Delta h)$ (or $[1 - R(\Delta h)]$) spectrum functions against height difference Δh is constant at $4/3$ for separations Δh up to 3 km, indicative of an isotropic region subject to a mean wind shear (Tchen, 1954).

3.4 Isotropy

Since the energy spectrum functions and the correlation difference functions all follow an established (shear) law at small scales in the height range from 80 to 100 km, it is possible to determine the nature of the isotropy at these scales in this region.

Figure 13, curve A, is a plot of the variation of

$$S = \frac{1-f}{1-g}$$

against Δh .

In the region up to a scale of 3.5 km, S has a value of approximately 0.7, indicating an isotropy lying somewhere between two and three dimensional. The increase in S at scales greater than approximately 3 km is due to the breakdown

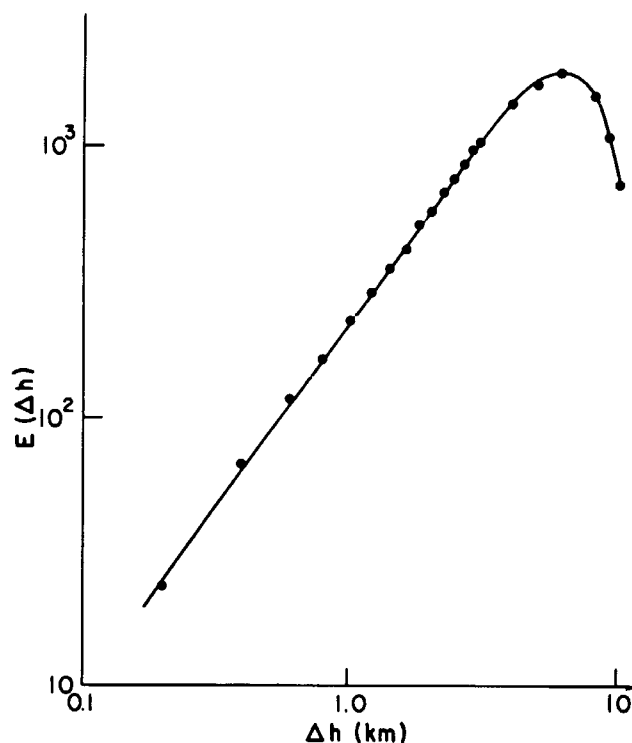


FIGURE 7.—The zonal energy spectrum function.

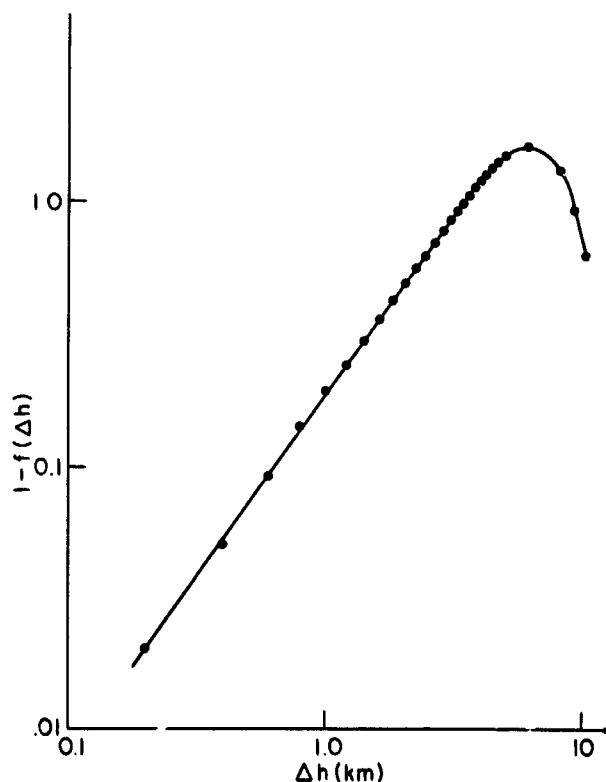


FIGURE 8.—The zonal correlation difference function.

at this scale of the 4/3 power law in both zonal and meridional spectrum functions.

The function S has been found to be quite sensitive to variation in the parameters specifying the mean wind profiles. For example, if, instead of the fourth order polynomial fits to the total height range of the zonal and meridional data, linear mean wind profiles are fitted over the height range 80 to 100 km only, subsequent spectrum analysis yields the S function of curve B of Figure 13. This curve would indicate the existence of three dimensional isotropy for scales up to 3 km. It is possible that the 4th order polynomial fits are, in fact, attributing a small fraction of the random wind variations to the mean motion. However, in order to correlate spectra obtained for different height strata (as is done in the next section), a continuous mean profile is required, and has therefore been used throughout the analysis.

3.5 Vertical Scale

The vertical scale associated with the turbulent wind structure is conveniently defined by the Δh corresponding to the maximum value of either $E(\Delta h)$ or $[1-R(\Delta h)]$. As can be seen from

Figures 7, 9 and 11, which are plotted using zonal, meridional and windspeed turbulent velocities respectively, the vertical correlation distance as defined above is not the same for each component, i.e., the wind motions are not isotropic at the maximum of $E(\Delta h)$. However, for the purpose of comparison with stratospheric and lower mesospheric data, where meridional winds are, for the most part, negligible, consideration need only be given to the zonal energy spectrum.

The variation of vertical scale with height in the stratosphere has been determined by Webb (1964) from a series of Robin soundings at Eglin. His results indicate an exponential increase of the vertical scale with height, from approximately 800 metres at 35 km to 2 km at 55 km. He has extrapolated this exponential to a height of 90 km, and finds a vertical scale of 6 km which is the value determined for this height by Greenhow (1959) from a correlation analysis of wind shears determined by means of radio reflections from meteor trails at Jodrell Bank (53°N).

In order to investigate the change with height of the vertical correlation distance in the 80 to 130 km region, the sodium trail data was sub-

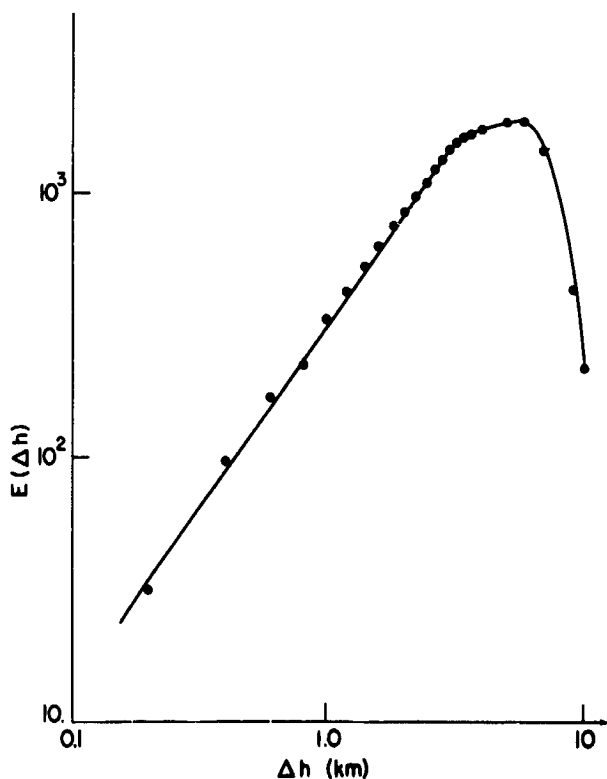


FIGURE 9.—The meridional energy spectrum function.

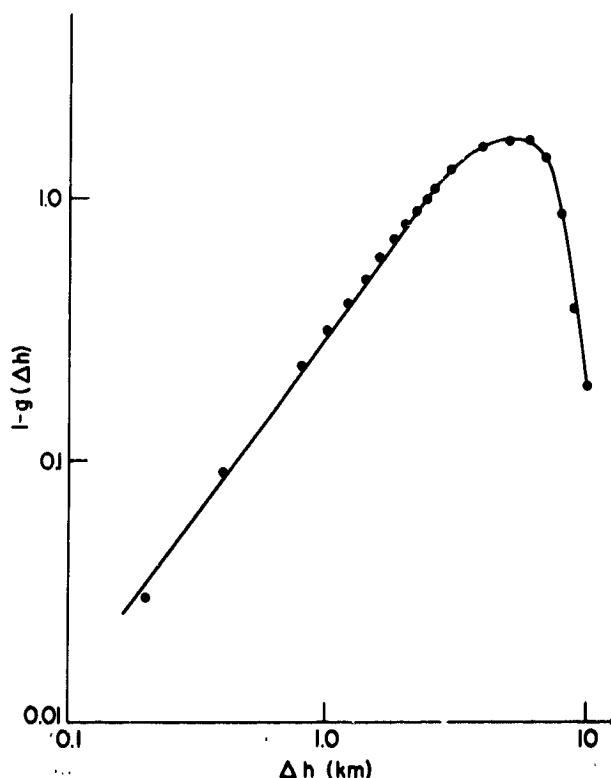


FIGURE 10.—The meridional correlation difference function.

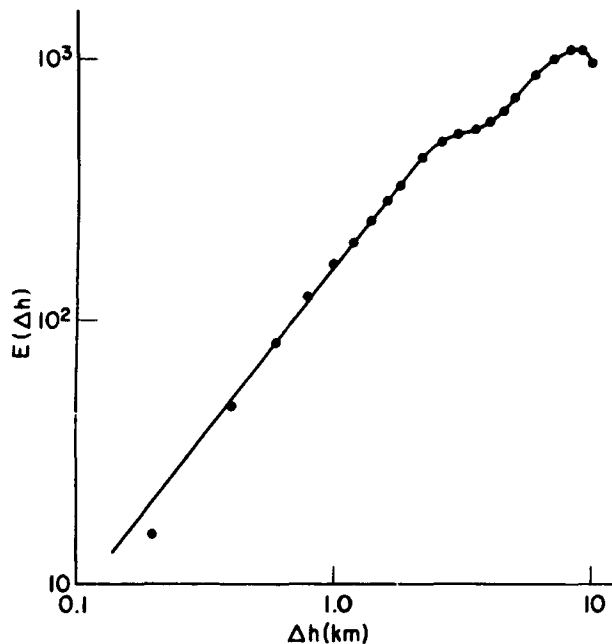


FIGURE 11.—The windspeed energy spectrum function.

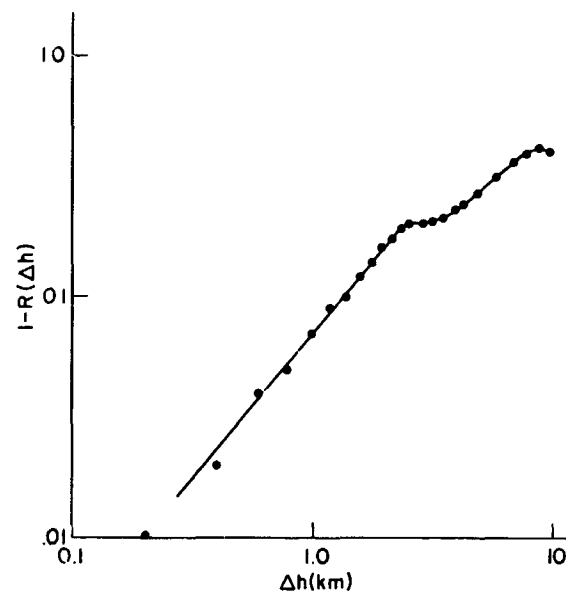


FIGURE 12.—The windspeed correlation difference function

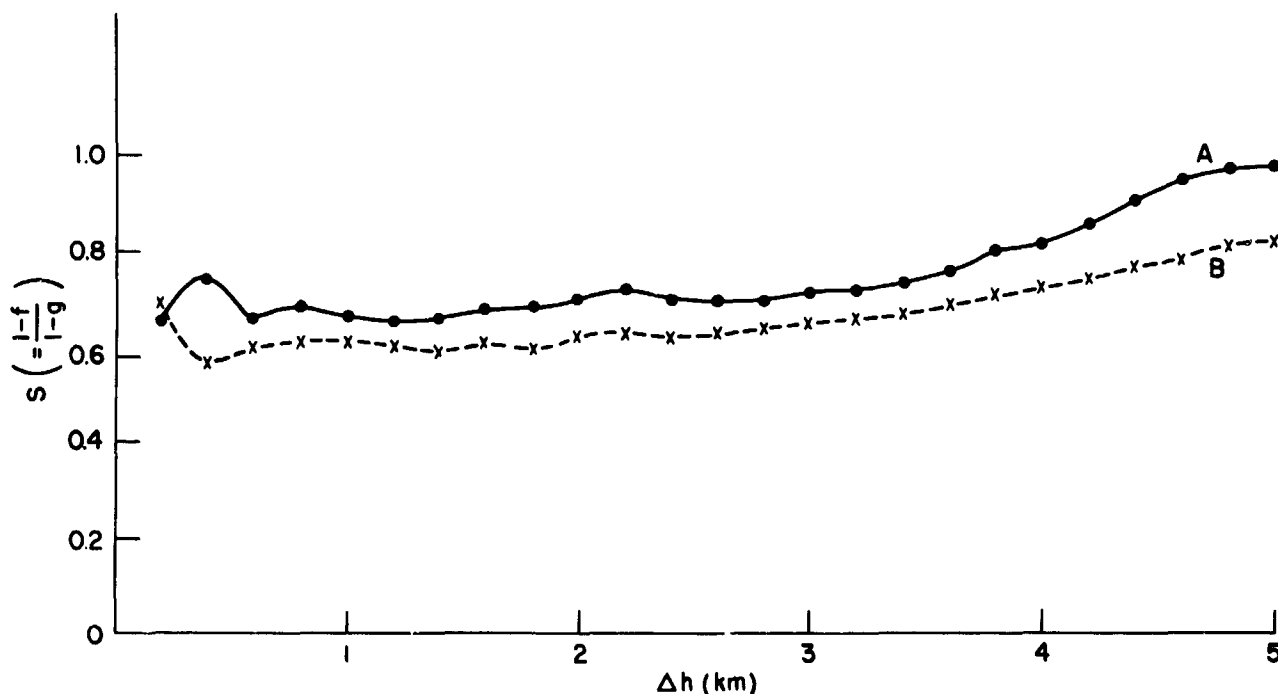


FIGURE 13.—Variation of the isotropy parameter "S" with scale.

jected to a stepwise analysis, starting with the 75 to 95 km height range, and proceeding via the 80 to 100, 85 to 105, etc. ranges to 115 to 135 km. The maxima of the resultant $E(\Delta h)$ curves were then plotted as the vertical scales at the midpoints of the respective height ranges. The results are shown in Figure 14, together with Green-

how's determination, and the 7.8 km at 94 km (October, 1961) determined from radio meteor trail shears at Adelaide (35°S) (Roper, 1962). The dashed line is Webb's extrapolated variation. While there is excellent agreement with Webb's proposed exponential increase, there is also good agreement between the vertical scales measured

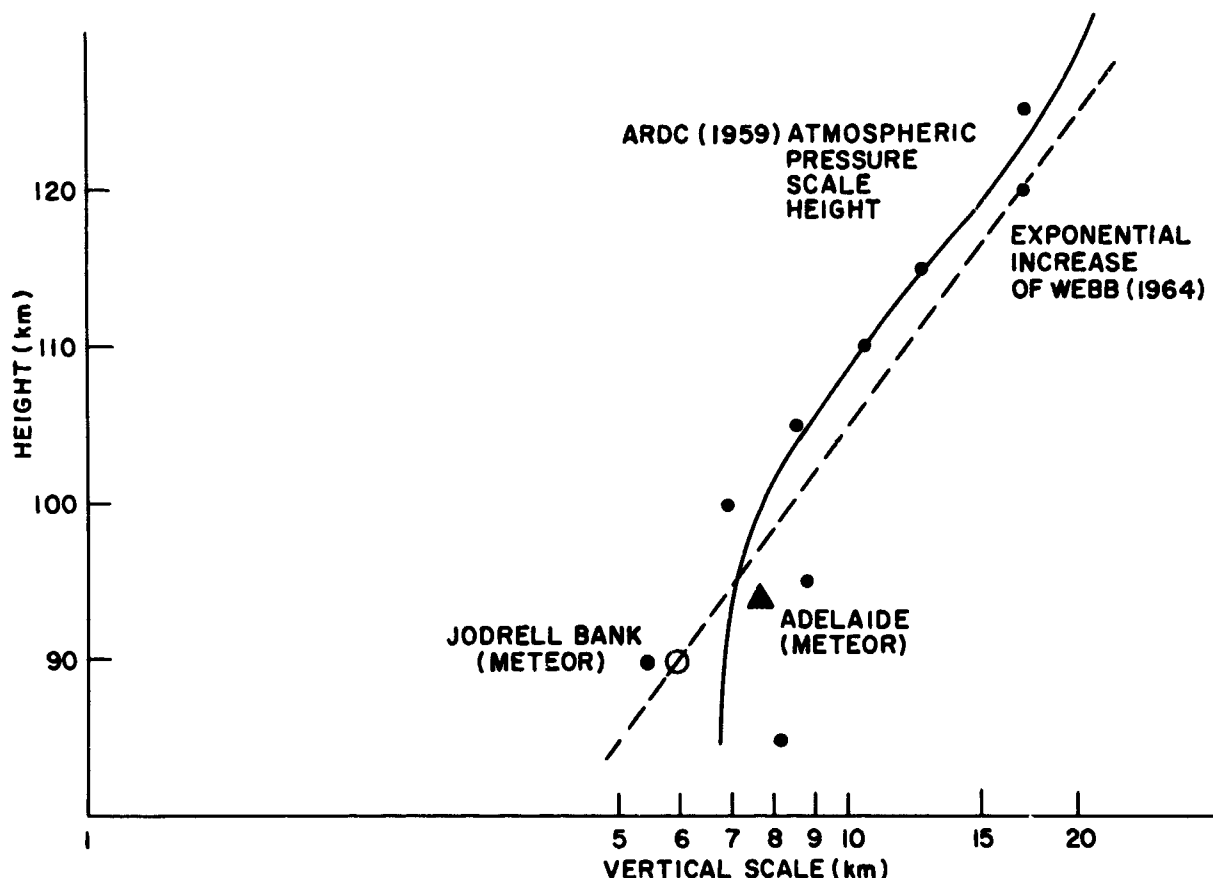


FIGURE 14.—Variation of the vertical scale of the zonal turbulent wind component with height.

at various altitudes for the Eglin firing, and the atmospheric pressure scale height (also shown in Figure 14). This phenomenon has also been observed by others (e.g., Zimmerman (1964)). As yet, no satisfactory explanation for such a dependence has been proposed.

Above 125 km, the magnitudes of the deviations of the measured winds from the mean wind values becomes insignificant. Wind motions at these heights for this particular firing are characteristically nonturbulent.

CONCLUSIONS

The techniques of spectrum analysis based on hydrodynamic turbulence theory can be profitably applied to wind data obtained from sodium trail rocket firings, at least in the height range from 80 to 100 km. The analysis of further firings should indicate whether or not the mean profiles determined as a prelude to spectrum analysis have any meteorological significance.

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APPENDIX I

RESUME OF THE COMPUTER PROGRAM

The computer program has been written in FORTRAN IV for execution on an IBM 7094, and consists of the following routines:

1. The main program SHEAR which reads and does some processing of the input data (which is punched on cards).

SHEAR calls subroutines

a) FIT, which fits polynomial profiles to the zonal and meridional wind velocity/height profiles by the method of least squares, utilizing

b) MATS, to solve the set of independent linear equations

c) The total wind speed profile is calculated by TRADE.

d) OUTPUT performs the energy spectrum and correlation analyses, and printout of same.

e) PAGE is a utility routine which turns and numbers the output pages, and prints an appropriate heading on each.

Input Data

The input data is on punched cards, and is read as follows:

1. Header Card; the information punched on this card is printed at the top of each page of output. The full 72 columns can be used for any type of alphanumeric heading.

2. The parameters specifying the mean zonal and meridional profiles. These determine the order of the polynomial fitted to

a) the zonal (punched in columns 2, 3) and

b) the meridional (punched in columns 5, 6) measured data to determine the mean profile. e.g., a 5 punched in column 3 will fit a polynomial of the form

$$a_1 + a_2h + a_3h^2 + a_4h^3 + a_5h^4$$

to the zonal wind data.

The maximum allowable value of the subscript i in the a_i above is 10.

3. The height range (must be integral, and in km) over which the energy spectrum and correlation analyses are to be performed. This can be any fraction of the height interval covered by the input data, but must not be less than 10 km unless an appropriate spectral range is punched in column 13. The minimum of the height range interval is punched in columns 2 to 5, the maximum in columns 6 to 9. If OUTPUT is required to produce spectra over height differences to other than a maximum of 10 km, the required maximum height difference can be punched in columns 10 to 13. This entry must not exceed the height range specified in columns 2 to 9. If this last field is left blank, spectra to 10 km are output.

4. The trial data, which must not exceed a maximum to minimum height range of 150 km. The data is punched as year, month, day, hour (local time, 24 hour clock), height (km), wind-speed (metres/sec), and azimuth (degrees) as follows

YR	MO	DAY	HOURL	HEIGHT	WINDSPEED	AZIMUTH
3	3	9	14	19	24	29

all integer

must be right adjusted if integer; anywhere in appropriate field if decimal punch is included.

5. After all data of 4. a blank card to flag "end of data."

6. Further height ranges as for 3 as required.

7. a) If a change in the mean wind profile is required a card with a 2 punched in column 1 may be substituted for any of the cards 6, and followed by a card of format 2 above which specified the

new profile. This card is then followed by further height ranges (as for 6.) as required.

7. b) If a completely new set of header plus data cards is to be processed after either stages 5, 6, or 7a, a card with a 1 punched in column 1 will return control to the start of the program, which will then read in the new set of data cards sequenced as from 1 above.

APPENDIX II

Listing of the FORTRAN IV Program

08/

SHEAR
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

```

C   SODIUM TRAIL HEIGHT SHEAR ANALYSIS PROGRAM
C
C   CALCULATES THE ENERGY SPECTRUM FOR TOTAL SHEAR, AND FOR ZONAL AND
C   MERIDIONAL SHEAR COMPONENTS
C
C   READS INPUT DATA AS FOLLOWS
C   A HEADER CARD, PUNCHED WITH DETAILS OF FIRING TIME, ETC.
C   FORMAT 12A6
C   THE PARAMETERS SPECIFYING THE MEAN ZONAL AND MERIDIONAL PROFILES,
C   FORMAT 213
C   THE HEIGHT RANGE OVER WHICH ANALYSIS IS TO BE PERFORMED, ZMIN,
C   ZMAX, AND THE MAXIMUM OF THE OUTPUT SPECTRAL RANGE DESIRED.
C   FORMAT 1X2F4.0, 14.
C   SPECTRAL RANGE MUST NOT EXCEED ZMAX-ZMIN.
C   IF NO SPECTRAL RANGE IS SPECIFIED, ZMAX-ZMIN MUST BE GREATER THAN
C   10KM, AND SPECTRA WILL BE OUTPUT TO 10KM.
C   THE TRAIL DATA, YEAR MONTH DAY LOCAL TIME (HOURS AND MINUTES,
C   24 HOUR CLOCK) HEIGHT(KM) WIND SPEED(METRES/SEC) WIND AZIMUTH
C   (DEGREES).
C   FORMAT 313, 15, F5.1, 2F5.0
C   A BLANK CARD
C   FURTHER ZMIN, ZMAX AS REQUIRED. IF A 2 APPEARS IN COLUMN 1,
C   PROGRAM WILL READ NEXT CARD AS NEW PROFILE SPECIFICATION, FOLLOWED
C   BY FURTHER ZMIN, ZMAX CARDS. IF A 1 APPEARS IN COLUMN 1, THE
C   PROGRAM WILL READ NEXT CARD AS HEADER CARD OF A COMPLETELY NEW SET
C   OF DATA.
C
C   DIMENSION BZ(10), BM(10)
C   DIMENSION RESULT(12), ZI(750), WINDI(750), AZRADI(750), ZONAL(750),
1ERID(750),          TZONAL(750), TMERID(750)
C   DIMENSION TSPEED(750), AW(5)
C   COMMON ZI, WINDI, AZRADI, ZONAL, ERID, TZONAL, TMERID, TSPEED
C   C=0.05
C   CS=0.000001
C   PI=3.1415926
C   TWOPI=2.0*PI
999 NGO=0
C   NI=0
C   I=0
C   NSUM=0
C   ZERO=0.0
C   READ (2,1)RESULT
1  FORMAT(12A6)
17 READ (2,4)NP,NQ
4  FORMAT (213)
C   NFIT=-1
2  READ(2,3)NGO, ZMIN, ZMAX, NDIFF
3  FORMAT(1I,2F4.0,14)
C   IF(NGO)19, 19, 18
18 IF(NGO-2)999, 17, 17
19 CONTINUE
C   IF(NDIFF)20, 20, 21
20 NDIFF=10
21 MIN=ZMIN+C
C   MAX=ZMAX+C

```

SHEAR
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

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```

C      IF(NI)5, 5, 101
      TRAIL DATA INPUT
5      READ      (2,6)MYEAR, MONTH, JOUR, LTIME, Z, WIND, AZDEG

6      FORMAT(3I3, I5, F5.1, 2F5.C)
      IF(MYEAR)100,100, 7
7      I=I+1
      IF(I-1)11, 11, 8
8      ZDIFF=Z-ZI(I-1)
      MZ=10.0 * Z+C
      LZ=10.0 * ZI(I-1)+C
      LDIFF=MZ-LZ
      IF(LDIFF-2)12, 11, 9
12     WRITE      (3,6)MYEAR, MONTH, JOUR, LTIME, Z, WIND, AZDEG
      WRITE(3,13)
13     FORMAT(1X/////1X26HDATA CARDS OUT OF SEQUENCE/////1X20HEXECUTION
14     TERMINATED/////1X)
      PRINT 13
      CALL EXIT
C      INTERPOLATION ROUTINE
9      NPOINT=LDIFF/2-1
      NSUM=NSUM+NPOINT
      GRAD=(WIND - WINDI(I-1))/ZDIFF
      AZDIFF=AZDEG/57.3-AZRADI(I-1)
      IF(ABS(AZDIFF)-PI)94,94,91
91     IF(AZDIFF)92,94,93
92     AZDIFF=AZDIFF+TWOPI
      GOTO94
93     AZDIFF=AZDIFF-TWOPI
94     AZGRAD=AZDIFF/ZDIFF
      DO10J=1,NPOINT
      ZI(I)=ZI(I-1)+0.2
      WINDI(I)=WINDI(I-1)+C.2*GRAD
      AZRADI(I)=AZRADI(I-1)+C.2*AZGRAD
      IF(AZRADI(I))95,96,96
95     AZRADI(I)=AZRADI(I)+TWOPI
96     IF(AZRADI(I)-TWOPI)98,98,97
97     AZRADI(I)=AZRADI(I)-TWOPI
98     CONTINUE
      I=I+1
10     CONTINUE
11     ZI(I)=Z
      WINDI(I)=WIND
      AZRADI(I)=AZDEG/57.3
      GO TO 5
100    NI=I
C      ALL DATA IN. COMMENCE COMPUTATION
      HDIFF=ZI(NI)-ZI(1)
      HPLUS=ZI(NI)+ZI(1)
C      CALCULATE ZONAL AND MERIDIONAL WIND COMPONENTS
      DO201J=1,NI
      ZONAL(J)=WINDI(J)*SIN(AZRADI(J))
201    ERID(J)=WINDI(J)*COS(AZRADI(J))
101    NZ=0
      LMIN=10*MIN
      LMAX=10*MAX

```

SHEAR

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EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

```

CALL PAGE (RESULT, ZERO)
DO104J=1,NI
LZ=ZI(J)*10.0+C
IF(LZ-LMIN)104,102,102
102 IF(LZ-LMAX)103,103,104
103 NZ=NZ+1
104 CONTINUE
WRITE (3,105)NI, NSUM, ZMIN, ZMAX, NZ, NP, NQ
105 FORMAT(1X////1X33HTOTAL NUMBER OF INPUT DATA POINTS I8//1X26HINC
1LUDING INTERPOLATION OF 15,7H POINTS ///1X45HNUMBER OF DATA POINT
2S WITHIN THE HEIGHT RANGE F5.0,6H KM TO,F5.0,3H KM I8///1X31HEAST
3 WEST PROFILE SPECIFICATION I5 ///1X33HNORTH SOUTH PROFILE SPECIFI
4CATION I3 ///1X)
START=1
CALL PAGE (RESULT, START)
WRITE (3,106)ZMIN, ZMAX
106 FORMAT(1X50HCALCULATED ZONAL AND MERIDIONAL MEAN WIND PROFILES
110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1HO
2/10X 6HHEIGHT 10X5HZONAL 12X10HMERIDIONAL 10X10HWIND SPEED /1X/
31X,18X,3(4X12HDATA MEAN,4X) /1X/)
NHITE=(ZMAX-ZMIN)/200-C
NHITE=NHITE+1
INK=5*NHITE-1
NFIT=NFIT+1
IF(NFIT)202,202,203
C CALCULATE COEFFICIENTS OF MEAN PROFILE POLYNOMIALS
202 CALL NORMAL (ZONAL, ERID, ZI, NI, BZ, BM, NP, NQ)
203 CONTINUE
C DETERMINE AND PRINT CUT MEAN WIND PROFILES
SUBS=(ZMAX+ZMIN-ZI(1))*5.0+1.0+C
DO108L=MIN, MAX, NHITE
IZ=SUBS-FLOAT(L*5)
J=MAX+MIN-L
H=J
S=(2.0*H-HPLUS)/HDIFF+CS
AZONAL=0.0
AMERID=0.0
DO204M=1,10
AZONAL=AZONAL+BZ(M)*S***(M-1)
AMERID=AMERID+BM(M)*S***(M-1)
204 CONTINUE
MZONAL=AZONAL
MMERID=AMERID
MSPEED=SQRT(AZONAL**2+AMERID**2)
WRITE(3,107)J, ZONAL(IZ), MZONAL, ERID(IZ), MMERID, WINDI(IZ), MSPEED
107 FORMAT(1X113,4X,3( F8.0,4X114,4X),/1X)
108 CONTINUE
CALL PAGE (RESULT, START)
WRITE(3,109)ZMIN, ZMAX
109 FORMAT(1X,22HTURBULENT WIND PROFILE, 10X12HHEIGHT RANGE F5.0,
1 6HKM TO F5.0, 3H KM/1X/1X, 10X6HHEIGHT, 10X6HTZONAL, 10X11HTMERIDIONA
2L, 10X6HTSPEED/1X)
C DETERMINE AND PRINT CUT TURBULENT WIND PROFILES
NH=(MAX-MIN)*5+1
K=(ZMAX-ZI(1))*5.0+2.0+C
NK=INK

```

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US/

EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

RMSZ=0.0

RMSM=0.0

RMST=0.0

SUMSQT=0.0

DO112J=1,NP

TZONAL(J)=0.0

TMERID(J)=0.0

K=K-1

H=ZMAX-0.2*FLOAT(J-1)

IH=H+0

S=(^ 0 * H-HPLUS)/HDIFF+CS

NK=NK+1

DO205M=1,10

TZONAL(J)=TZONAL(J)+BZ(M)*S*(M-1)

TMERID(J)=TMERID(J)+bM(M)*S*(M-1)

205 CONTINUE

TZONAL(J)=ZONAL(K)-TZONAL(J)

TMERID(J)=ERID(K)-TMERID(J)

CALL TRADE(J,K)

C CALCULATE RMS TURBULENT VELOCITIES FOR THE HEIGHT RANGE ZMIN/ZMAX

RMSZ=^TZONAL(J)*^2+RMSZ

RMSM=TMERID(J)*^2+RMSM

RMST=TSPEED(J)*^2+RMST

SUMSQT=SUMSQT+1.0

IF(NK-INK)112,112,110

110 WRITE(3,111)IH, TZONAL(J), TMERID(J), TSPEED(J)

111 FORMAT(1X, 10X14, 12XF5.0, 11XF8.0, 13XF5.0/1X)

NK=0

112 CONTINUE

CALL PAGE(RESULT, START)

RMSZ=SQRT(RMSZ/SUMSQT)

RMSM=SQRT(RMSM/SUMSQT)

RMST=SQRT(RMST/SUMSQT)

WRITE (3,206)ZMIN, ZMAX, (BZ(I), I=1, NP)

206 FORMAT(1X10HNORMALIZED

1 1X12HHEIGHT RANGE 10XF7.0,7H KM. TO F7.0, 4H KM./1H0/

2 1X42HCOEFFICIENTS GIVING BEST FIT TO ZONAL DATA/1X/10E11.3)

WRITE(3,207)(BM(I),I=1,NQ)

207 FORMAT(1X/1X/

11X/1X47HCOEFFICIENTS GIVING BEST FIT TO MERIDIONAL DATA/1X/1X10E11
2.3/1X/)

WRITE(3,208)RMSZ, RMSM, RMST

208 FORMAT(1X/1X

1 1X22HRMS TURBULENT VELOCITY/1X/1X10X5HZONALF11.0,11HMETRES/S

2E0./1X/1X10X10HMERIDIONAL F6.0,11HMETRES/SEC./1X/1X10X10HWIND SPEE

3D F6.0,11HMETRES/SEC.)

CALL PAGE(RESULT, START)

WRITE (3,113)ZMIN,ZMAX

113 FORMAT(1X46HENERGY SPECTRUM OF ZONAL WIND HEIGHT VARIATION,

110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)

C CALCULATE AND OUTPUT ENERGY SPECTRUM FUNCTIONS

CALL OUTPUT(TZONAL, NH, NDIFF)

CALL PAGE (RESULT, START)

WRITE (3,114) ZMIN, ZMAX

114 FORMAT(1X51HENERGY SPECTRUM OF MERIDIONAL WIND HEIGHT VARIATION,

SHEAR

08/

EXTERNAL FORMULA NUMBER -- SOURCE STATEMENT -- INTERNAL

```

CALL PAGE(RESET, ZERO)
DO104J=1,NI
LZ=ZI(J)*10.0+C
IF(LZ-LMIN)104,102,102
102 IF(LZ-LMAX)103,103,104
103 NZ=NZ+1
104 CONTINUE
WRITE (3,105)NI, NSUM, ZMIN, ZMAX, NZ, NP, NQ
105 FORMAT(1X////1X33HTOTAL NUMBER OF INPUT DATA POINTS I8//1X26HINC
1LUDING INTERPOLATION OF 15,7H POINTS ///1X45HNUMBER OF DATA POINT
2S WITHIN THE HEIGHT RANGE F5.0,6H KM TO,F5.0,3H KM I8///1X31HEAST
3 WEST PROFILE SPECIFICATION I5 ///1X33HNORTH SOUTH PROFILE SPECIFI
4CATION I3 ///1X)
START=1
CALL PAGE(RESET,START)
WRITE (3,106)ZMIN, ZMAX
106 FORMAT(1X50HCALCULATED ZONAL AND MERIDIONAL MEAN WIND PROFILES
110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1HO
2/10X 6HHEIGHT 10X5HZONAL 13X10HMERIDIONAL 10X10HWIND SPEED /1X/
31X,18X,3(4X12HDATA MEAN,4X),/1X/)
NHITE=(ZMAX-ZMIN)/20.0-C
NHITE=NHITE+1
INK=5*NHITE-1
NFIT=NFIT+1
IF(NFIT)202,202,203
C CALCULATE COEFFICIENTS OF MEAN PROFILE POLYNOMIALS
202 CALL NORMAL (ZONAL, ERID, ZI, NI, BZ, BM, NP, NQ)
203 CONTINUE
C DETERMINE AND PRINT CUT MEAN WIND PROFILES
SUBS=(ZMAX+ZMIN-ZI(1))*5.0+1.0+C
DO108L=MIN, MAX, NHITE
IZ=SUBS-FLOAT(L*5)
J=MAX+MIN-L
H=J
S=(2.0*H-HPLUS)/HDIFF+CS
AZONAL=0.0
AMERID=0.0
DO204M=1,10
AZONAL=AZONAL+BZ(M)*S***(M-1)
AMERID=AMERID+BM(M)*S***(M-1)
204 CONTINUE
MZONAL=AZONAL
MMERID=AMERID
MSPEED=SQRT(AZONAL**2+AMERID**2)
WRITE(3,107)J, ZONAL(IZ), MZONAL, ERID(IZ), MMERID, WINDI(IZ),MSPEED
107 FORMAT(1X113,4X,3( F8.0,4X14,4X),/1X)
108 CONTINUE
CALL PAGE (RESET, START)
WRITE(3,109)ZMIN, ZMAX
109 FORMAT(1X,22HTURBULENT WIND PROFILE, 10X12HHEIGHT RANGE F5.0,
1 6HKM TO F5.0,3H KM/1X/1X,10X6HHEIGHT, 10X6HTZONAL, 10X11HTMERIDIONA
2L,10X6HTSPEED/1X)
C DETERMINE AND PRINT CUT TURBULENT WIND PROFILES
NH=(MAX-MIN)*5+1
K=(ZMAX-ZI(1))*5.0+2.0+C
NK=INK

```

08/

SHEAR
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

```

RMSZ=0.0
RMSM=0.0
RMST=0.0
SUMSQT=0.0
DO112J=1,NH
TZONAL(J)=0.0
TMERID(J)=0.0
K=K-1
H=ZMAX-0.2*FLOAT(J-1)
IH=H+C
S=(2.0*H-HPLUS)/HDIFF+CS
NK=NK+1
DO205M=1,10
TZONAL(J)=TZONAL(J)+BZ(M)*S*(M-1)
TMERID(J)=TMERID(J)+BM(M)*S*(M-1)
205 CONTINUE
TZONAL(J)=ZONAL(K)-TZONAL(J)
TMERID(J)=ERID(K)-TMERID(J)
CALL TRADE(J,K)
C CALCULATE RMS TURBULENT VELOCITIES FOR THE HEIGHT RANGE ZMIN/ZMAX
RMSZ=TZONAL(J)**2+RMSZ
RMSM=TMERID(J)**2+RMSM
RMST=TSPEED(J)**2+RMST
SUMSQT=SUMSQT+1.0
IF(NK-INK)112,112,110
110 WRITE(3,111)IH,TZONAL(J),TMERID(J),TSPEED(J)
111 FORMAT(1X,10XI4,12XF5.0,11XF8.0,13XF5.0/1X)
NK=0
112 CONTINUE
CALL PAGE (RESULT, START)
RMSZ=SQRT(RMSZ/SUMSQT)
RMSM=SQRT(RMSM/SUMSQT)
RMST=SQRT(RMST/SUMSQT)
WRITE(3,206)ZMIN,ZMAX,(BZ(I),I=1,NP)

206 FORMAT(1X10HNORMALIZED:
1      1X12HHEIGHT RANGE 10XF7.0,7H KM. TO F7.0,4H KM./1H0/
2      1X42HCOEFFICIENTS GIVING BEST FIT TO ZONAL DATA/1X/10E11.3)
WRITE(3,207)(BM(I),I=1,NQ)
207 FORMAT(1X/1X/
11X/1X47HCOEFFICIENTS GIVING BEST FIT TO MERIDIONAL DATA/1X/1X10E11
2.3/1X/)
WRITE(3,208)RMSZ, RMSM, RMST
208 FORMAT(1X/1X
1      1X22HRMS TURBULENT VELOCITY/1X/1X10X5HZONALF11.0,11HMETRES/S
2EC./1X/1X10X10HMERIDIONAL F6.0,11HMETRES/SEC./1X/1X10X10HWIND SPEE
3D F6.0,11HMETRES/SEC.)
CALL PAGE (RESULT, START)
WRITE (3,113)ZMIN, ZMAX
113 FORMAT(1X46HENERGY SPECTRUM OF ZONAL WIND HEIGHT VARIATION,
110X12HHEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)
C CALCULATE AND OUTPUT ENERGY SPECTRUM FUNCTIONS
CALL OUTPUT(TZONAL, NH, NDIFF)
CALL PAGE (RESULT, START)
WRITE (3,114) ZMIN, ZMAX
114 FORMAT(1X51HENERGY SPECTRUM OF MERIDIONAL WIND HEIGHT VARIATION,

```

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EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

110X12HEIGHT RANGE F5.0,6H KM TO F5.0, 3H KM/1X)
CALL OUTPUT(TMERID, NH, NDIFF)
CALL PAGE (RESULT, START)
WRITE (3,115)ZMIN,ZMAX
115 FORMAT(1X46HENERGY SPECTRUM OF WIND SPEED HEIGHT VARIATION,
110X12HEIGHT RANGE F5.0,6H KM TO F5.0,3H KM/1X)
CALL OUTPUT(TSPEED, NH, NDIFF)
GOTO2
END

```

/$ID JO2T      RGR   BLDG 11
$PAUSE
$EXECUTE      IBJOB
IBJOB VERSION 2, 7090 - PP - 929
1
0
$IBJOB      GO
$IBFTC OUT   M94,XR7
1           OUT
           EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERN
0
SUBROUTINE OUTPUT(TSPEED,NH,NDIFF)
C
C   PRODUCES HEIGHT SPECTRUM AND PERFORMS CORRELATION ANALYSIS.
C
    DIMENSION TSPEED(750)
    WRITE(3,1)
1   FORMAT(1X,10X7HDELTA H,5X10HE(DELTA H),
15X11HCORRELATION,3X13HSIGNIF   1-G /1X)
    NEND=5 * NDIFF
    LINE=0
    DO125K=1,NEND
    LINE=LINE+1
    DELTAH=FLOAT(K) * 0.2
    SUM=0.0
    ENERGY=0.0
    ERGX=0.0
    GJPKXJ=0.0
    SQJ=0.0
    SQJPK=0.0
    DO123J=1,NH
    IF(NH-J-K)124,122,122
122  JPK=J+K
    ENERGY=ENERGY+(TSPEED(JPK)-TSPEED(J)) * * 2
    SUM=SUM+1.0
    GJPKXJ=GJPKXJ+TSPEED(JPK) * TSPEED(J)
    SQJ=SQJ+TSPEED(J) * * 2
    SQJPK=SQJPK+TSPEED(JPK) * * 2
123  CONTINUE
124  IF(SUM)117,117,118
117  G=0.0
    SIGNIF=9.99
    GOTO119
118  ENERGY=ENERGY/SUM
    G=GJPKXJ/SQRT(SQJ * SQJPK)
    SIGNIF=(1.0-G * * 2)/SQRT(SUM-1.0)
119  DIFF=1.0-G
    IF(LINE-50)125,125,116
116  WRITE(3,2)
    2   FORMAT(1H1/1X)
    WRITE(3,1)
    LINE=0
125  WRITE(3,115)DELTAH,ENERGY,      G,SIGNIF,DIFF
115  FORMAT(1X,6XF8.1, 7XF8.1 ,8XF8.3,F11.3,F8.2)
    RETURN
    END

```

\$IBJOB GO
\$IBFTC NORMAL M94,XR7

NORMAL
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTE

SUBROUTINE NORMAL (ZONAL,ERID,ZI,NI,BZ,BM,NP,NQ)
C
C DETERMINES BEST FIT PROFILE TO THE NORMALIZED TOTAL HEIGHT RANG
C
DIMENSION ZONAL(750),ERID(750),ZI(750),BZ(1C),BM(1C),ZINORM(750)
DO1J=1,NI
ZINORM(J)=(2.0 * ZI(J)-ZI(NI)-ZI(1))/(ZI(NI)-ZI(1))
1 CONTINUE
CALL FIT (ZONAL,NP,BZ,NI,ZINORM)
CALL FIT (ERID, NQ,BM,NI,ZINORM)
RETURN
END

\$IBJOB GO
\$IBFTC TRADE M94,XR7

TRADE 08/
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL

SUBROUTINE TRADE(J,K)
C
C CALCULATES TURBULENT WINDSPEED PROFILE
C
DIMENSION ZI(750)
DIMENSION WINDI(750),AZRADI(750),ZONAL(750),ERID(750),TZONAL(750),
1TMERID(750),TSPEED(750)
COMMON ZI, WINDI,AZRADI,ZONAL,ERID,TZONAL,TMERID,TSPEED
TSPEED(J)=WINDI(K)-SQRT((ZONAL(K)-TZONAL(J)) * * 2+(ERID(K)-TMERID(J)
1) * * 2)
RETURN
END

\$IBFTC FIT M94,XR7

08/17

C
C
C
C
C

C
C
C
C

MATS
EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL FO 08/17

SUBROUTINE MATS (S,A,NSPEC,MISS)

C
C REDUCES THE AUGMENTED MATRIX S TO PRODUCE THE COEFFICIENTS A
C

```
DIMENSION S(10,11), A(10)
MISS = -1
MM = NSPEC + 1
N = NSPEC
DO15I = 2,N
70 II = I - 1
7 DO15J = 1,II
8 IF(S(I,J))9,15,9
9 IF(ABS(S(J,J))-ABS(S(I,J)))11,10,10
10 R = S(I,J)/S(J,J)
GOTO 130
11 R = S(J,J)/S(I,J)
DO12K = 1,MM
B = S(J,K)
S(J,K) = S(I,K)
12 S(I,K) = B
130 JJ = J + 1
13 DO14K = JJ,MM
14 S(I,K) = S(I,K) - R * S(J,K)
15 CONTINUE
IF(ABS(S(N,N))-1.0E-10)16,16,17
16 MISS = +1
GOTO29
17 A(N) = S(N,MM)/S(N,N)
DO28I = 2,N
JJ = N - I + 1
B = 0.0
II = N - I + 2
DO25K = II,N
25 B = B + S(JJ,K) * A(K)
IF(ABS(S(JJ,JJ))-1.0E-10)16,16,28
28 A(JJ) = (S(JJ,MM)-B)/S(JJ,JJ)
29 RETURN
END
```

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08/17

\$IBFTC MATS M94,XR7

08/17

\$IBFTC PAGE M94,XR7

PAGE

08/17

EXTERNAL FORMULA NUMBER — SOURCE STATEMENT — INTERNAL FO

SUBROUTINE PAGE (RESULT, START)

C
C
C
C

Turns page, numbers it, and writes heading as appearing
on result card.

DIMENSION RESULT (12)

IF(START)1,1,2

1 NOP=0

2 NOP=NOP+1

WRITE (3,3) RESULT, NOP

3 FORMAT (1H1/1X12A6,30X4HPAGE15//)

RETURN

END

APPENDIX III
Sample Data And Computer Output

PLANETARY ATMOSPHERES

1097

PRELIMINARY PAPER EGLIN DATA, MAY 21, 1963. 16/9/64.

PROFILE, 5, 5.

5	5						
80	100						
63	5	21	1910	69.0	34	255	
63	5	21	1910	70.6	31	256	
63	5	21	1910	71.0	27	270	
63	5	21	1910	72.4	43	272	
63	5	21	1910	72.6	20	290	
63	5	21	1910	77.0	41	256	
63	5	21	1910	80.4	25	358	
63	5	21	1910	82.0	08	310	
63	5	21	1910	82.4	28	202	
63	5	21	1910	83.6	27	200	
63	5	21	1910	84.0	54	190	
63	5	21	1910	85.0	54	194	
63	5	21	1910	85.4	30	108	
63	5	21	1910	88.0	60	080	
63	5	21	1910	91.6	76	095	
63	5	21	1910	95.0	50	178	
63	5	21	1910	96.0	40	120	
63	5	21	1910	96.6	54	150	
63	5	21	1910	97.6	40	115	
63	5	21	1910	99.0	42	084	
63	5	21	1910	99.6	64	078	
63	5	21	1910	100.0	58	070	
63	5	21	1910	100.8	65	069	
63	5	21	1910	102.0	80	079	
63	5	21	1910	103.0	93	094	

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PUBLICATIONS OF GSFC, 1964: I. SPACE SCIENCES

63	5	21	1910	104.0	115	104
63	5	21	1910	104.2	111	120
63	5	21	1910	105.0	129	122
63	5	21	1910	105.6	140	130
63	5	21	1910	105.8	130	128
63	5	21	1910	107.0	125	132
63	5	21	1910	108.0	92	140
63	5	21	1910	109.0	50	160
63	5	21	1910	110.0	42	240
63	5	21	1910	111.0	56	250
63	5	21	1910	113.0	60	262
63	5	21	1910	114.0	84	274
63	5	21	1910	115.0	68	270
63	5	21	1910	117.2	72	282
63	5	21	1910	125.0	45	286
63	5	21	1910	130.0	30	296
63	5	21	1910	135.0	14	310
63	5	21	1910	140.0	15	316

← BLANK CARD

7
8

← EOF

PLANETARY ATMOSPHERES

1099

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. 15/9/64.

PROFILE,5 5.

TOTAL NUMBER OF INPUT DATA POINTS 356

INCLUDING INTERPOLATION OF 313 POINTS

NUMBER OF DATA POINTS WITHIN THE HEIGHT RANGE 80. KM TO 100. KM

101

EAST WEST PROFILE SPECIFICATION 5

NORTH SOUTH PROFILE SPECIFICATION 5

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

CALCULATED ZONAL AND MERIDIONAL MEAN WIND PROFILES

HEIGHT RANGE

HEIGHT	ZONAL		MERIDIONAL		WIND SPEED	
	DATA	MEAN	DATA	MEAN	DATA	MEAN
100	70.	34	20.	-21	72.	40
99	56.	37	21.	-21	60.	43
98	42.	39	4.	-22	42.	45
97	39.	41	-11.	-22	41.	47
96	34.	43	-35.	-22	48.	49
95	31.	45	-28.	-22	42.	50
94	6	46	-51.	-22	52.	51
93	31.	46	-51.	-21	59.	51
92	55.	47	-38.	-21	67.	51
91	73.	47	-13.	-21	74.	51
90	72.	46	-1.	-20	72.	50
89	67.	45	4.	-19	67.	49
88	62.	43	8.	-18	63.	47
87	55.	41	5.	-18	55.	45
86	44.	39	-4.	-16	44.	43
85	31.	36	-9.	-15	32.	39
84	-12.	33	-53.	-14	54.	36
83	-9.	29	-25.	-13	27.	32
82	-10.	24	-26.	-11	28.	27
81	-8.	19	10.	-10.	12.	22
80	-3.	14	23.	-8	23.	16

PLANETARY ATMOSPHERES

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TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

TURBULENT WIND PROFILE

HEIGHT RANGE 80. KM TO 100. KM

HEIGHT	TZONAL	TMERIDIONAL	TSPEED
100	35.	41.	42.
99	19.	42.	26.
98	2.	26.	16.
97	-3.	11.	31.
96	-10.	-12.	32.
95	-14.	-6.	29.
94	-40.	-29.	8.
93	-16.	-29.	31.
92	8.	-16.	58.
91	26.	8.	62.
90	25.	19.	52.
89	22.	24.	45.
88	18.	27.	38.
87	13.	24.	33.
86	4.	13.	30.
85	-6.	7.	27.
84	-45.	-38.	18.
83	-39.	-12.	-5.
82	-35.	-14.	0.
81	-28.	20.	3.
80	-18.	31.	10.

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. 15/9/64.

PROFILE,5,5.

NORMALIZED HEIGHT RANGE

80. KM. TO 100. KM.

COEFFICIENTS GIVING BEST FIT TO ZONAL DATA

0.386E 02 -0.126E 03 -0.244E 03 0.174E 03 0.205E 03

COEFFICIENTS GIVING BEST FIT TO MERIDIONAL DATA

-0.304E 02 0.198E 02 0.148E 03 -0.168E 02 -0.125E 03

RMS TURBULENT VELOCITY

ZONAL 22.METRES/SEC.

MERIDIONAL 23.METRES/SEC.

WIND SPEED 17.METRES/SEC.

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

ENERGY SPECTRUM OF ZONAL WIND HEIGHT VARIATION HEIGHT RANGE 80.

DELTA H	DELTA V * * 2	CORRELATION	SICNIF	I-G
0.2	24.3	0.979	0.004	0.02
0.4	64.9	0.943	0.011	0.06
0.6	113.7	0.906	0.019	0.10
0.8	169.4	0.851	0.028	0.15
1.0	230.7	0.797	p 037	0.20
1.2	297.1	0.739	0.047	0.26
1.4	369.4	0.676	0.056	0.32
1.6	450.4	0.604	0.066	0.40
1.8	534.8	0.525	0.076	0.47
2.0	619.9	0.449	0.084	0.55
2.2	703.2	0.374	0.091	0.63
2.4	789.0	0.297	0.097	0.70
2.6	877.1	0.217	0.102	0.78
2.8	964.9	0.136	0.106	0.86
3.0	1054.0	0.053	0.108	0.95
3.2	1133.4	-0.023	0.109	1.02
3.4	1209.9	-0.099	0.109	1.10
3.6	1285.2	-0.176	0.107	1.18
3.8	1360.8	-0.257	0.104	1.26
4.0	1436.0	-0.342	0.099	1.34
4.2	1507.0	-0.429	0.092	1.43
4.4	1567.1	-0.514	0.083	1.51
4.6	1614.3	-0.597	0.073	1.60
4.8	1681.4	-0.651	0.066	1.65
5.0	1745.7	-0.696	0.060	1.70
5.2	1804.0	-0.735	0.053	1.74
5.4	1852.3	-0.766	0.048	1.77
5.6	1885.5	-0.788	0.045	1.79
5.8	1897.3	-0.800	0.043	1.80
6.0	1880.3	-0.801	0.043	1.80
6.2	1863.6	-0.795	0.044	1.80
6.4	1847.4	-0.783	0.047	1.78
6.6	1827.7	-0.762	0.051	1.76
6.8	1803.8	-0.730	0.057	1.73
7.0	1775.0	-0.690	0.065	1.69
7.2	1740.9	-0.642	0.073	1.64
7.4	1701.2	-0.588	0.082	1.59
7.6	1656.5	-0.528	0.092	1.53
7.8	1608.4	-0.466	0.100	1.47
8.0	1555.6	-0.401	0.108	1.40
8.2	1497.1	-0.333	0.116	1.33
8.4	1432.6	-0.263	0.122	1.26
8.6	1361.9	-0.190	0.128	1.19
8.8	1285.1	-0.116	1.132	1.12
9.0	1202.6	-0.040	0.135	1.04
9.2	1114.7	0.038	0.136	0.96
9.4	1020.5	0.120	0.135	0.88
9.6	920.9	0.207	0.133	0.79
9.8	819.6	0.296	0.128	0.70
10.0	733.6	0.372	0.122	0.63

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

ENERGY SPECTRUM OF MERIDIONAL WIND HEIGHT VARIATION				HEIGHT RANGE
DELTA H	DELTA V * * 2	CORRELATION	SIGNIF	I-G
0.2	31.3	0.973	0.005	0.03
0.4	94.1	0.917	0.916	0.08
0.6	164.4	0.853	0.028	0.15
0.8	241.7	0.781	0.040	0.22
1.0	326.3	0.701	0.052	0.30
1.2	418.3	0.612	0.064	0.39
1.4	515.2	0.518	0.076	0.48
1.6	616.0	0.420	0.086	0.58
1.8	723.6	0.316	0.094	0.68
2.0	825.3	0.219	0.100	0.78
2.2	935.1	0.116	0.105	0.88
2.4	1052.2	0.008	0.107	0.99
2.6	1175.9	-0.104	0.106	1.10
2.8	1305.7	-0.219	0.103	1.22
3.0	1429.3	-0.325	0.097	1.33
3.2	1534.6	-0.411	0.091	1.41
3.4	1614.7	-0.479	0.085	1.48
3.6	1667.7	-0.535	0.079	1.54
3.8	1711.0	-0.582	0.074	1.58
4.0	1741.1	-0.617	0.069	1.62
4.2	1756.8	-0.641	0.066	1.64
4.4	1758.3	-0.656	0.065	1.66
4.6	1745.6	-0.662	0.064	1.66
4.8	1776.6	-0.680	0.062	1.68
5.0	1836.6	-0.715	0.056	1.72
5.2	1878.5	-0.733	0.054	1.73
5.4	1900.7	-0.735	0.054	1.74
5.6	1902.8	-0.723	0.056	1.72
5.8	1885.6	-0.699	0.061	1.70
6.0	1853.4	-0.666	0.066	1.67
6.2	1809.8	-0.625	0.073	1.63
6.4	1751.9	-0.573	0.081	1.57
6.6	1679.1	-0.510	0.090	1.51
6.8	1592.6	-0.435	0.100	1.44
7.0	1494.3	-0.350	0.109	1.35
7.2	1385.9	-0.257	0.117	1.26
7.4	1269.8	-0.156	0.123	1.16
7.6	1149.0	-0.051	0.127	1.05
7.8	1028.4	0.057	0.128	0.94
8.0	912.5	0.164	0.126	0.84
8.2	802.4	0.266	0.121	0.73
8.4	697.1	0.366	0.114	0.63
8.6	598.6	0.459	0.105	0.54
8.8	508.4	0.544	0.094	0.46
9.0	427.7	0.620	0.083	0.38
9.2	357.9	0.684	0.072	0.32
9.4	302.8	0.735	0.063	0.27
9.6	260.1	0.774	0.056	0.23
9.8	231.1	0.801	0.050	0.20
10.0	219.4	0.812	0.048	0.19

TRIAL ANALYSIS. EGLIN DATA, MAY 21, 1963. CONTINUOUS PROFILE,5,5.

ENERGY SPECTRUM OF WIND SPEED HEIGHT VARIATION			HEIGHT RANGE 80.	
DELTA H	DELTA V * * 2	CORRELATION	SIGNIF	I-G
0.2	15.1	0.993	0.001	0.01
0.4	47.1	0.979	0.004	0.02
0.6	83.2	0.964	0.007	0.04
0.8	122.0	0.947	0.011	0.05
1.0	161.5	0.930	0.014	0.07
1.2	200.6	0.914	0.017	0.09
1.4	240.7	0.898	0.020	0.10
1.6	286.9	0.879	0.024	0.12
1.8	337.0	0.859	0.028	0.14
2.0	382.2	0.841	0.031	0.16
2.2	421.0	0.827	0.034	0.17
2.4	455.3	0.815	0.036	0.19
2.6	483.4	0.805	0.038	0.20
2.8	502.9	0.799	0.039	0.20
3.0	512.5	0.797	0.040	0.20
3.2	517.8	0.796	0.040	0.20
3.4	528.9	0.793	0.041	0.21
3.6	542.8	0.789	0.042	0.21
3.8	561.7	0.783	0.043	0.22
4.0	584.0	0.775	0.045	0.22
4.2	606.6	0.768	0.046	0.23
4.4	628.3	0.761	0.048	0.24
4.6	650.9	0.755	0.049	0.25
4.8	679.5	0.746	0.051	0.25
5.0	712.5	0.734	0.053	0.27
5.2	748.8	0.722	0.056	0.28
5.4	785.1	0.710	0.058	0.29
5.6	818.7	0.700	0.060	0.30
5.8	847.5	0.692	0.062	0.31
6.0	870.6	0.686	0.063	0.31
6.2	894.8	0.680	0.065	0.32
6.4	919.4	0.674	0.066	0.33
6.6	943.4	0.668	0.068	0.33
6.8	968.1	0.661	0.069	0.34
7.0	996.2	0.653	0.071	0.35
7.2	1023.8	0.645	0.073	0.35
7.4	1048.3	0.638	0.075	0.36
7.6	1070.2	0.630	0.077	0.37
7.8	1088.9	0.623	0.078	0.38
8.0	1103.4	0.616	0.080	0.38
8.2	1114.0	0.608	0.082	0.39
8.4	1121.1	0.600	0.084	0.40
8.6	1120.2	0.593	0.086	0.41
8.8	1114.6	0.588	0.087	0.41
9.0	1103.8	0.585	0.089	0.42
9.2	1087.7	0.584	0.090	0.42
9.4	1066.1	0.586	0.090	0.41
9.6	1039.0	0.590	0.090	0.41
9.8	1007.8	0.596	0.090	0.40
10.0	973.4	0.603	0.090	0.40

N64-31451

TEMPERATURE, PRESSURE, DENSITY, AND WIND MEASUREMENTS WITH THE ROCKET GRENADE EXPERIMENT, 1960-1963*

W. SMITH, L. KATCHEN, P. SACHER, P. SWARTZ, AND J. THEON

Complete data from 28 rocket grenade experiments at Wallops Island, Virginia, and Fort Churchill, Canada, are presented. Pressures, temperatures, densities, and winds have been derived directly from the recorded times of explosions and sound arrivals; but no attempt has been made to analyze the meteorological significance of these measurements. Error analyses on 16 of the Wallops experiments are also included.

INTRODUCTION

During the period 1960-1963 a total of 28 rocket grenade experiments were carried out by Goddard Space Flight Center at Wallops Island, Virginia, and Fort Churchill, Canada. It is the purpose of this report to present the complete data from these experiments. The data have been reduced but not analyzed, i.e., the physical measurements (pressure, density, temperature and winds) have been derived directly from recordings of the times of grenade explosions and sound arrivals, and from the rocket trajectories; but no attempt has been made to analyze the measurements for their meteorological significance. Rather, this report should be considered as a record of the unsmoothed, raw measurements which may serve as the basis of further investigation and interpretation of the structure of the atmosphere. This is the first such report since the program started in 1960. As the program continues, future results will be documented in similar fashion.

The grenade experiment is one of the several techniques employed in NASA's Meteorological Rocket Sounding Program, the objective of which is to obtain as representative as possible a sample of the synoptic structure of the mesosphere and lower ionosphere. In order to extend the findings over the widest possible geographic area, and perhaps eventually to attain a global network

of soundings, the NASA launchings are coordinated whenever possible with soundings in other parts of the world. The instrumental details and methods of data reduction have been thoroughly described elsewhere.†

EXPERIMENTAL METHOD

One and two pound explosive charges (grenades) are carried aloft in the nosecone of a Nike-Cajun sounding rocket. The grenades are ejected and exploded at 4-6 km intervals. Either a precise radar such as the FPS-16 or a doppler tracking system, or both, are used to determine the rocket position and hence the exact position of the explosion. The time of the explosion is detected by small rocket borne infrared photocells and telemetered to the ground. A ground based array of hot wire microphones with frequency response peaked at 4 cycles per second is used to detect and record the sound waves from each exploding grenade. The measured experimental parameters are the times of the grenade explosions, the positions of the grenade explosions, and the times of arrival of the sound waves at the ground based microphones.

Elevation and azimuth angles are computed for each arriving sound wave front by applying a least-squares fit to the arrival times at the various microphones. Each wave is then analytically traced up through the atmosphere by means of Snell's law. Data from radiosonde ascents or

*Published as NASA Technical Report R-211, October 1964.

†Nordberg, W., and Smith W., "The Rocket Grenade Experiment," NASA Technical Note D-2107, March 1964.

from small meteorological rockets obtained at the time of the grenade soundings are used for this tracing up to the first explosion; above this, the results of the experiment itself are used for each successive explosion. This *apparent* position of the *wave* is then compared with the *known* position of the *wave source*, the grenade explosion. The amount by which the sound wave has been displaced horizontally from one explosion to the next is a measure of the average wind velocity vector in the layer between any two adjacent explosions. The average speed of the sound, and hence average temperature between two adjacent explosions, may also be determined. This analysis yields an altitude profile of temperatures and winds as a direct measurement. The temperature profile consists of discrete points, each representing an average value for an altitude layer between grenade explosions. It may be used to derive density and pressure profiles, if the pressure or density is

known at a given level at the bottom of the temperature profile; and the latter data are available from the accompanying radiosonde. Pressure is calculated by use of the hydrostatic equation, and the ideal gas law; an integration is performed over the temperature profile starting at the known pressure. The density at any level is then computed from the pressure, temperature, and universal gas constant.

During the period between July 8, 1960 and December 7, 1963, 23 successful rocket grenade soundings were conducted from Wallops Island, Virginia (37° 50'N, 75° 29'W). In addition to offering excellent launch and tracking support, the geographical location of Wallops Island makes it a good middle-latitude observation site. Five soundings were conducted from a sub-arctic range, Fort Churchill, Canada, (58° 47'N, 94° 17'W) in the period between December 1962 and March 1963. Originally, it was planned to distribute

Table 1
Dates, Times, and Locations of GSFC Grenade Experiments 1960-1963

Date	Local Time	Location
8 July 1960	2259	Wallops Island
14 February 1961	1850	Wallops Island
16 February 1961	2126	Wallops Island
5 April 1961	0757	Wallops Island
5 May 1961	1800	Wallops Island
5 May 1961	2354	Wallops Island
13 July 1961	1707	Wallops Island
14 July 1961	1102	Wallops Island
20 July 1961	0530	Wallops Island
16 September 1961	1855	Wallops Island
1 March 1962	1901	Wallops Island
2 March 1962	0615	Wallops Island
23 March 1962	1854	Wallops Island
27 March 1962	1904	Wallops Island
17 April 1962	0428	Wallops Island
6 June 1962	2005	Wallops Island
7 June 1962	2053	Wallops Island
1 December 1962	1625	Wallops Island
4 December 1962	0105	Fort Churchill
6 December 1962	0032	Wallops Island
5 December 1962	2343	Fort Churchill
20 February 1963	1734	Fort Churchill
20 February 1963	1847	Wallops Island
28 February 1963	1547	Fort Churchill
28 February 1963	1711	Wallops Island
8 March 1963	1901	Wallops Island
8 March 1963	1801	Fort Churchill
7 December 1963	0812	Wallops Island

the number of firings about equally between the two sites, but a fire early in 1961 destroyed part of the Fort Churchill launch facility, and it was not operational again until November 1962. Table 1 lists the soundings conducted. Four of the recent firings from each site were conducted nearly simultaneously.

The results of the 28 soundings are presented in Figures 1 through 43. In Figures 1 through 28, the *directly measured* parameters of temperatures and winds are tabulated (actual computer print-outs) in the columns on the left side of the page

immediately above the graph. Tables of pressure and density *derived* from the temperature profiles are also shown. The temperatures were linearly interpolated between each of the measured points.

Winds and temperatures below 30 km obtained from radiosondes, and below 50 km from smaller meteorological sounding rockets (Arcas* and Hasp† whenever available) launched nearly

*"Data Report of the Meteorological Rocket Network, Spring 1961 Firings," U.S. Army Signal Missile Agency, White Sands Missile Range, New Mexico, September 1961.

†Private Communication from M. J. Parker, Naval Ordnance Laboratory, White Oak, Maryland.

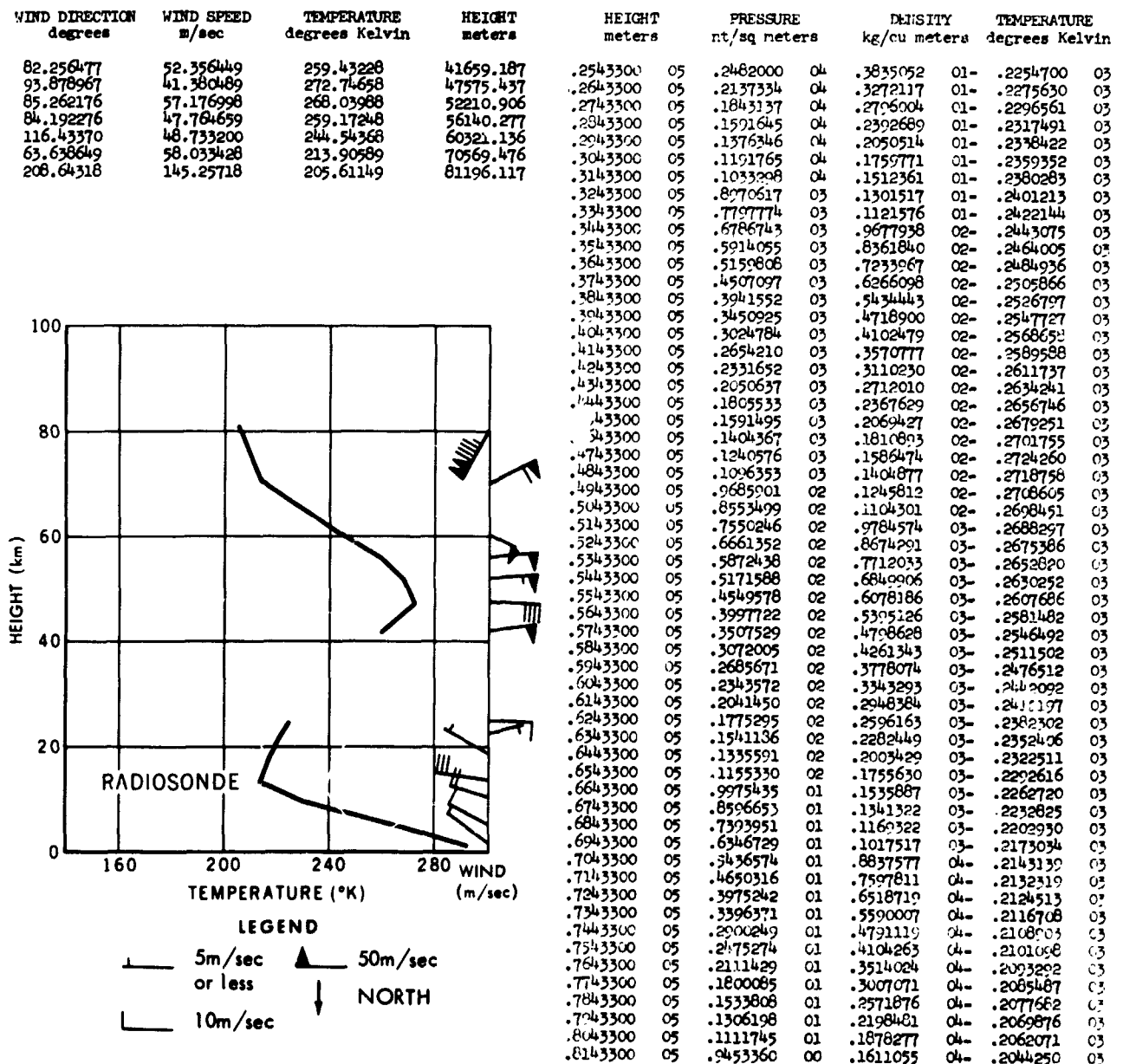


FIGURE 1

simultaneously with the grenade soundings are shown graphically. Winds and temperatures above 40 km from the grenade soundings are shown on the graph.

RESULTS

In the results from these soundings all data points, regardless of our degree of confidence, are presented in tabular form. Some individual data points may be subject to later revision, as there may be occasional errors in some of the records which have not yet been detected. In the preparation of these graphs some data points, where such errors seem obvious, were omitted and the graphs note the omissions. For the most part,

these occur at the upper altitude limit of the experiment where the signal-to-noise ratio of detected acoustic waves is poor. Random errors contained in either the measurements or in the data reductions can usually be recognized by comparing two adjacent data points. Since these errors will generally be contained in the parameters (time and space coordinates of the explosion) associated with one individual grenade explosion, the temperature data points both below and above this explosion will contain errors of approximately equal magnitude but opposite sign. Therefore, if large excursions of this nature occur between two adjacent data points, it is very likely that the recordings from the explosion between these two data points were in error.

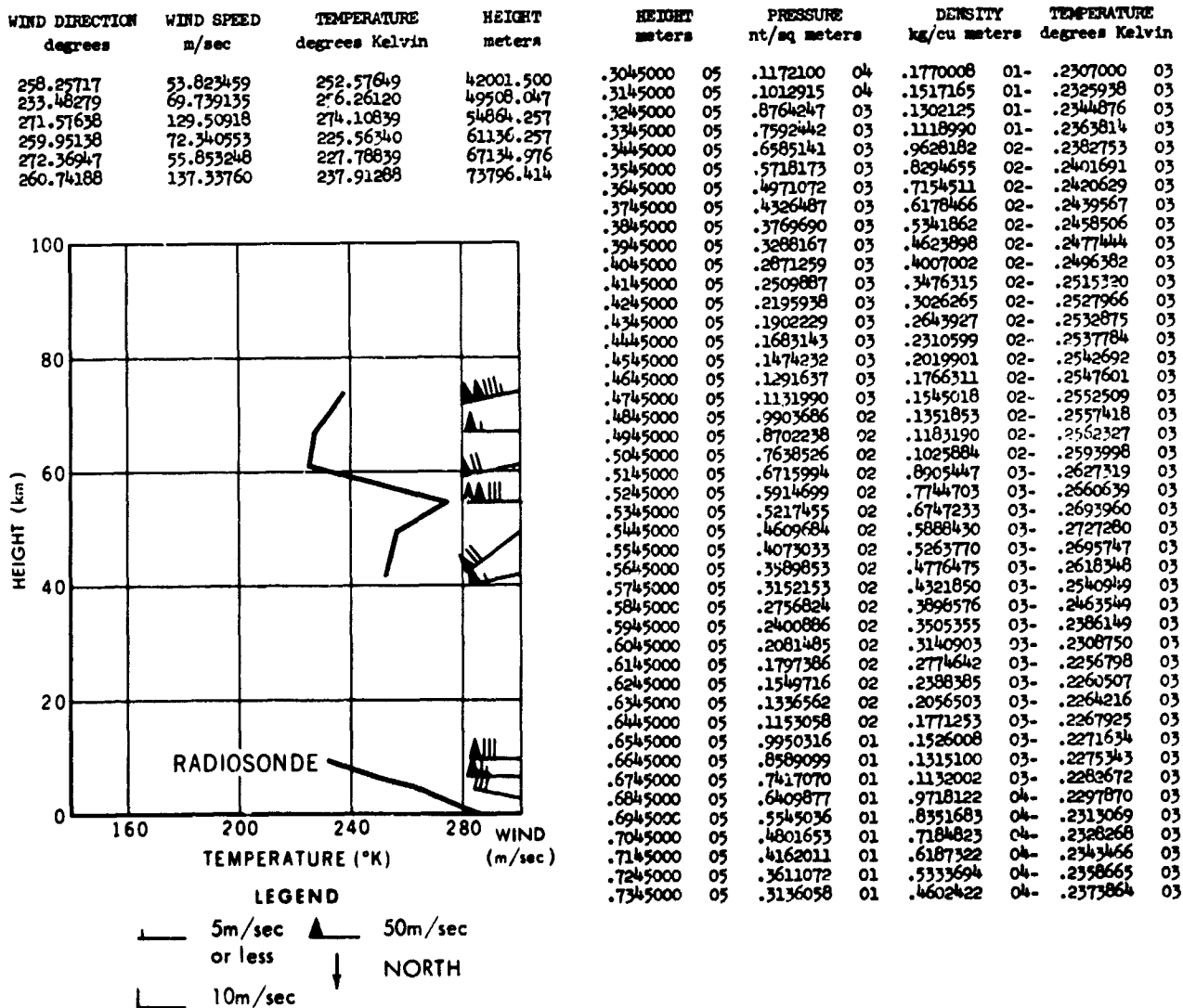


FIGURE 2

ERROR ANALYSIS

An error analysis was run with the aid of a digital computer on 16 of the 28 experiments reported. The firings considered were the first 16 conducted from Wallops Island. The following nomenclature is used in the error analysis:

Δw = Difference in west coordinate between successive grenade explosions.

Δn = Difference in north coordinate between successive grenade explosions.

h = Difference in altitude between successive grenade explosions.

ω = Wind direction in the layer.

W = Wind speed in the layer.

c = Speed of sound in the layer = $k(T)^{1/2}$, where k is a proportionality constant depending on molecular mass, the ratio of specific heats, and the universal gas constant.

ϕ_u = Azimuth angle of sound wave for upper explosion.

τ = Travel time of sound wave in layer.

$K = c_0 \sec \theta_u + W_0 \cos(\phi_u - \omega_0) - W \cos(\phi_u - \omega)$, where c_0 is the mean speed of sound over

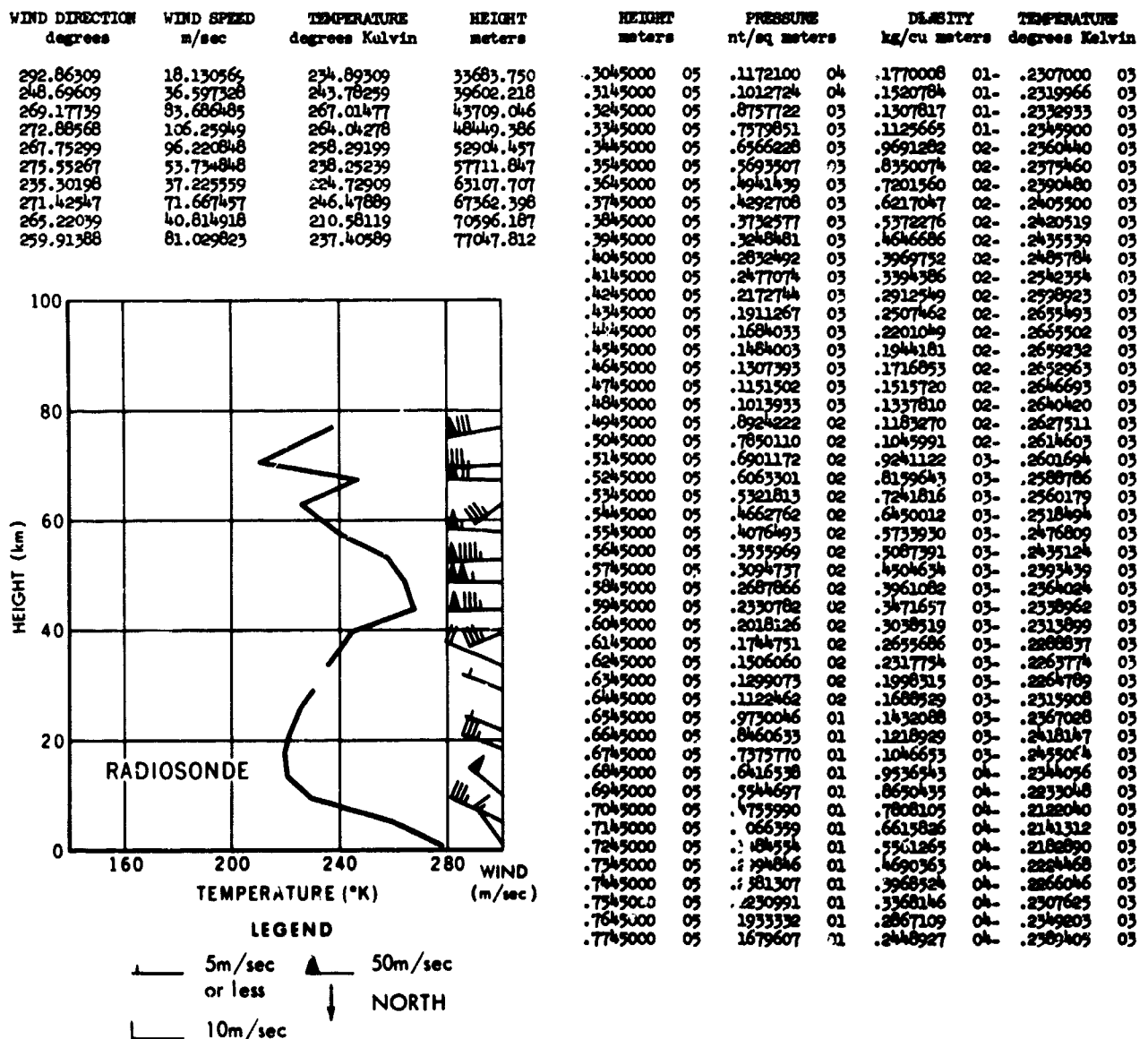


FIGURE 3

the microphone array; θ_u is the elevation angle of sound wave for upper explosion; W_0 is the wind speed over the microphone array; and ω_0 is the wind direction over the microphone array.

t = Travel time of sound wave from grenade explosion to ground.

T = Mean temperature in layer.

Δt = Time difference between microphones introduced by changing the time of arrival at one microphone.

ϕ'_u = Elevation angle from ground to apparent position of upper grenade explosion as obtained by ray tracing.

In order to determine the error function, deliberate errors were introduced in the following experimental parameters:

A. Position of grenade explosions

1. North: 200 meters
2. West: 200 meters
3. Up: 50 meters

B. Travel time of sound from explosions to ground: 0.3 second.

C. Time of arrival at one of the 6 microphones relative to the other five microphones: 0.02 second.

WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees kelvin
274.04568	37.738597	258.56097	44334.699	.2877900	05 .1516800	04 .2323785	01- .2274000
267.40359	40.801197	267.20568	51244.250	.2977900	05 .1307583	04 .1986374	01- .2294032
335.20639	41.921638	269.77447	56158.988	.3077900	05 .1109419	04 .1700349	01- .2314063
177.28808	39.535289	258.58029	61859.648	.3177900	05 .9765115	03 .1457528	01- .2334095
160.35488	66.775573	233.18949	69577.523	.3277900	05 .8453935	03 .1251036	01- .2354107
94.937965	25.101959	49.960929	75558.476	.3377900	05 .7328085	03 .1075323	01- .2374159
				.3477900	05 .6360084	03 .925 .697	02- .2394191
				.3577900	05 .5526712	03 .7975310	02- .2414223
				.3677900	05 .4808323	03 .6881542	02- .2434255
				.3777900	05 .4188070	03 .5945215	02- .2454287
				.3877900	05 .3652430	03 .5142620	02- .2474318
				.3977900	05 .3168796	03 .4445376	02- .2494350
				.4077900	05 .2787158	03 .3861788	02- .2514382
				.4177900	05 .2438811	03 .3352423	02- .2534414
				.4277900	05 .2136335	03 .2913607	02- .2554446
				.4377900	05 .1873388	03 .2535110	02- .2574478
				.4477900	05 .1644407	03 .2210924	02- .2594510
				.4577900	05 .1444530	03 .1932845	02- .2614542
				.4677900	05 .1268788	03 .1681058	02- .2634574
				.4777900	05 .1116025	03 .1480057	02- .2654606
				.4877900	05 .9830851	02 .1261710	02- .2674638
				.4977900	05 .8658475	02 .1036504	02- .2694670
				.5077900	05 .7630761	02 .9471741	02- .2714702
				.5177900	05 .6728685	02 .8673650	02- .2734734
				.5277900	05 .5935381	02 .7971521	02- .2754766
				.5377900	05 .5237114	02 .7370438	02- .2774798
				.5477900	05 .4622301	02 .6865007	02- .2794830
				.5577900	05 .4080211	02 .6435611	02- .2814862
				.5677900	05 .3600563	02 .6070670	02- .2834894
				.5777900	05 .3170350	02 .5764530	02- .2854926
				.5877900	05 .2784159	02 .5513810	02- .2874958
				.5977900	05 .2438266	02 .5308800	02- .2894990
				.6077900	05 .2109195	02 .5134191	02- .2915022
				.6177900	05 .1853705	02 .4980970	02- .2935054
				.6277900	05 .1610078	02 .4848614	02- .2955086
				.6377900	05 .1399185	02 .4732689	02- .2975118
				.6477900	05 .1214860	02 .4630078	02- .2995150
				.6577900	05 .1054423	02 .4540523	02- .3015182
				.6677900	05 .9148293	01 .4462893	02- .3035214
				.6777900	05 .7934166	01 .4395009	02- .3055246
				.6877900	05 .6878553	01 .4335200	02- .3075278
				.6977900	05 .5943888	01 .4280780	02- .3095310
				.7077900	05 .5080049	01 .4230076	02- .3115342
				.7177900	05 .4222664	01 .4183656	02- .3135374
				.7277900	05 .3379725	01 .4141418	02- .3155406
				.7377900	05 .2554000	01 .4102674	02- .3175438
				.7477900	05 .1750806	01 .4067555	02- .3195470
				.7577900	05 .1060058	01 .4035040	02- .3215502

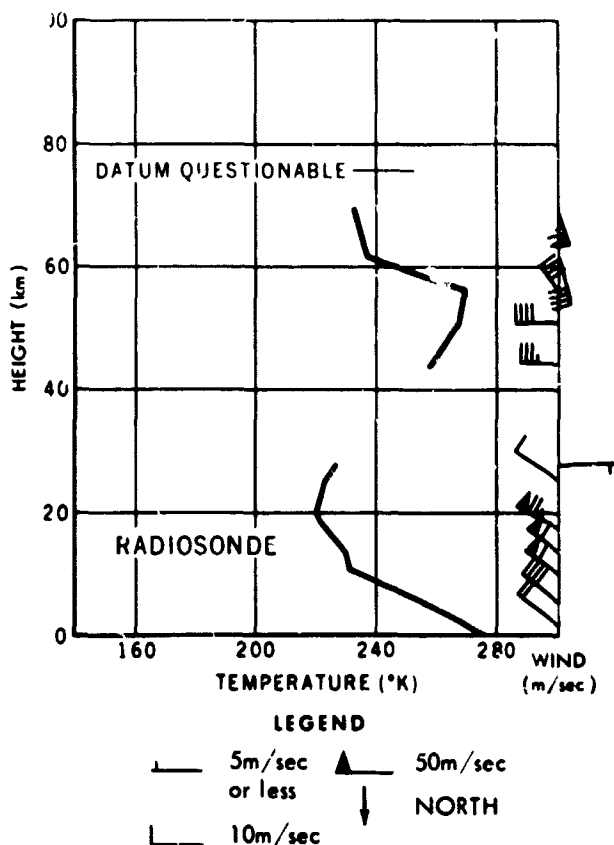


FIGURE 4

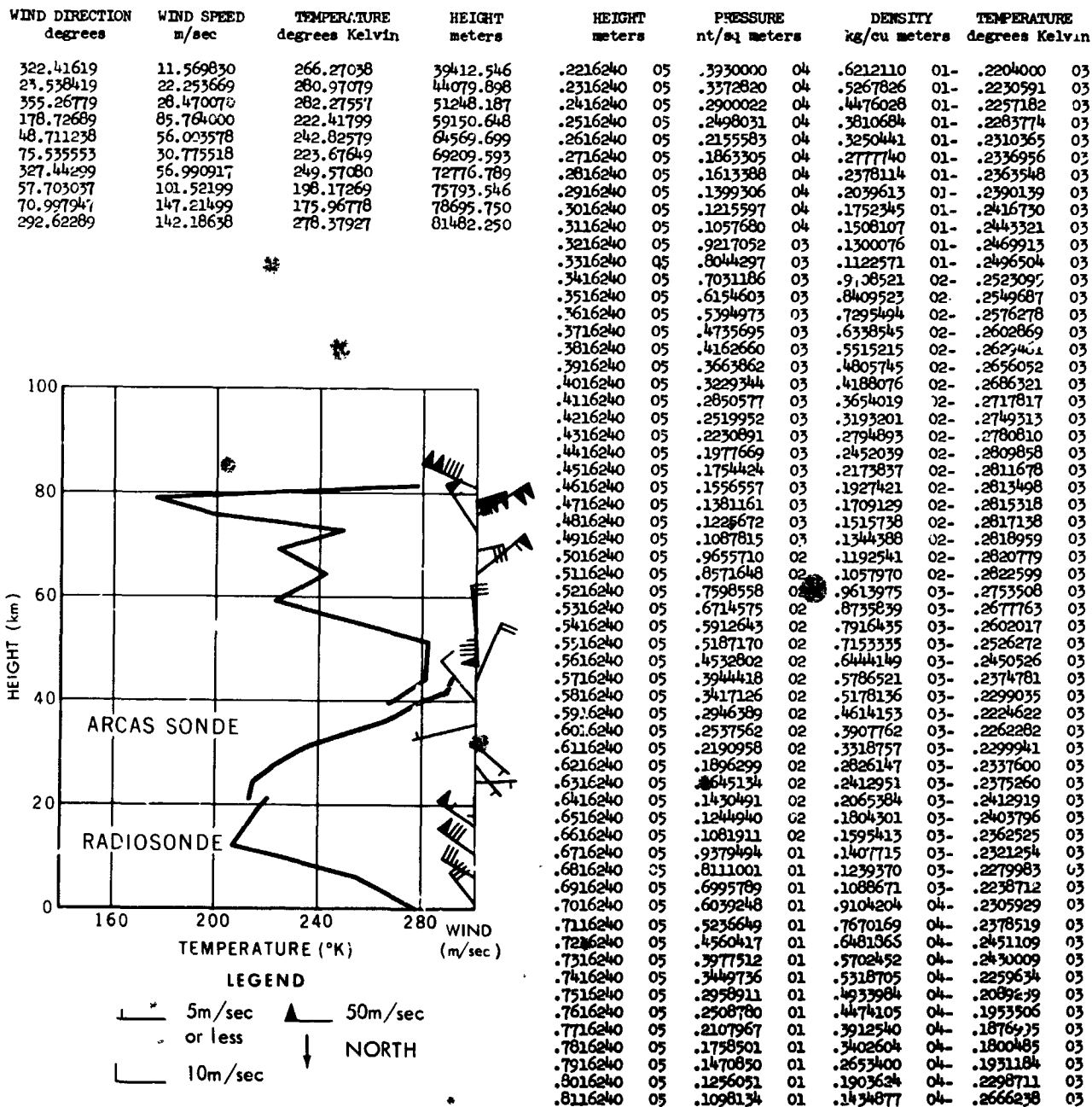


FIGURE 5

It was determined that an error in one of the foregoing parameters for any given grenade will cause wind and temperature errors in the two layers adjacent to the grenade. The temperature errors, in general, will be of opposite sign and approximately equal in magnitude. There are slight errors introduced in the remaining layers above the grenade; however, these are negligibly small. The errors in the temperatures and winds due to errors in the coordinates of the explosions are given by:

$$\frac{\partial W}{\partial \Delta w} = \frac{\sin \omega}{\tau}; \quad (1)$$

$$\frac{\partial W}{\partial \Delta n} = \frac{\cos \omega}{\tau}; \quad (2)$$

$$\frac{\partial W}{\partial h} = \frac{\cot \theta'_u}{\tau} \cos(\omega - \phi_u); \quad (3)$$

$$\frac{\partial c}{\partial \Delta w} = \frac{\sin \phi_u}{c\tau} \left[\frac{h^2 K}{\tau^2 (K^2 - 2c^2)} - \frac{c^2}{K} \right]; \quad (4)$$

$$\frac{\partial c}{\partial \Delta n} = \frac{\cos \phi_u}{c\tau} \left[\frac{h^2 K}{\tau^2 (K^2 - 2c^2)} - \frac{c^2}{K} \right]; \quad (5)$$

$$\frac{\partial c}{\partial h} = \frac{1}{c\tau} \left[\frac{h^2 K \cot \theta'_u - h K^2 \tau}{\tau^2 (K^2 - 2c^2)} + \frac{c^2 \cot \theta'_u}{K} \right]. \quad (6)$$

In the error analysis, the north and west coordinates of alternate grenades were changed by -200m , and the up coordinates were changed by -50m . This means that $d\Delta w$ and $d\Delta n$ for adjacent layers were alternately $+200$ and -200 , and dh for adjacent layers were alternately $+50$ and -50 . The azimuth angle changes slowly with altitude. Generally, $\sin \phi_u$ and $\cos \phi_u$ do not change sign during an experiment; thus the signs of $(\partial c/\partial \Delta w)d\Delta w$ and $(\partial c/\partial \Delta n)d\Delta n$ alternate regularly from layer to layer as the coordinates of alternate grenade explosions are changed. This accounts for the regular alternation of the temperature error, since all other signs in the equation for $\partial c/\partial \Delta w$, $\partial c/\partial \Delta n$ remain constant during an experiment. Similarly, it can be seen from the expression for $\partial c/\partial h$ that all the quantities retain

the same sign throughout an experiment. Thus, the same regular alternation for the signs of the $(\partial c/\partial h)dh$ will occur for successive grenades.

The situation is different, however, for $\partial W/\partial \Delta w$, $\partial W/\partial \Delta n$, $\partial W/\partial h$. The sign of $\partial W/\partial \Delta w$ depends on the sign of $\sin \omega$ which in turn depends on the value of ω . Since ω may vary from 0 to 360° , $\sin \omega$ may be either positive or negative. The same is true for the sign of $\partial W/\partial \Delta n$, which depends on $\cos \omega$. Hence, we expect no regular alternation from layer to layer in the sign of the wind error. The sign of $\partial W/\partial h$ depends on the sign of $\cos(\omega - \phi_u)$ which in turn depends on the value of $\omega - \phi_u$. Although ϕ_u changes slowly and in general remains in the same quadrant, ω can vary from 0 to 360° , and hence $\omega - \phi_u$ can be positive or negative. Thus, $(\partial W/\partial h)dh$ exhibits no regular alternation in sign. In summary, we can see that the signs of $\partial c/\partial \Delta w$, $\partial c/\partial \Delta n$, $\partial c/\partial h$ are constant and hence an error in the coordinates of one individual grenade explosion will cause a positive error in the measured temperatures in the layer below the explosion if a negative error is caused in the layer above the explosion, and vice versa. In general, the magnitudes of the temperature errors caused by an error of the explosion coordinates between two adjacent layers are approximately equal.

However, this does not hold for wind errors. The error functions for the wind speed are proportional to $\sin \omega$, $\cos \omega$ and $\cos(\omega - \phi_u)$, which may vary considerably from layer to layer. The error functions of c , however, are proportional to ϕ_u and θ'_u which do not vary greatly from layer to layer.

The two other parameters of importance to the errors in addition to the grenade explosion coordinates were investigated in a similar manner. These parameters were: the time differences, Δt , between arrival of the sound wave at any two microphones within the array of six; and the travel time, t , of the sound from the explosion to the ground. The parameter Δt is important because the direction of the arriving sound wave front is derived from those differences, while t enters primarily in the calculation of the speed of sound in the layer. An error of .02 second was introduced into the arrival time at one microphone without altering the arrival times at the remaining five microphones, while an error of

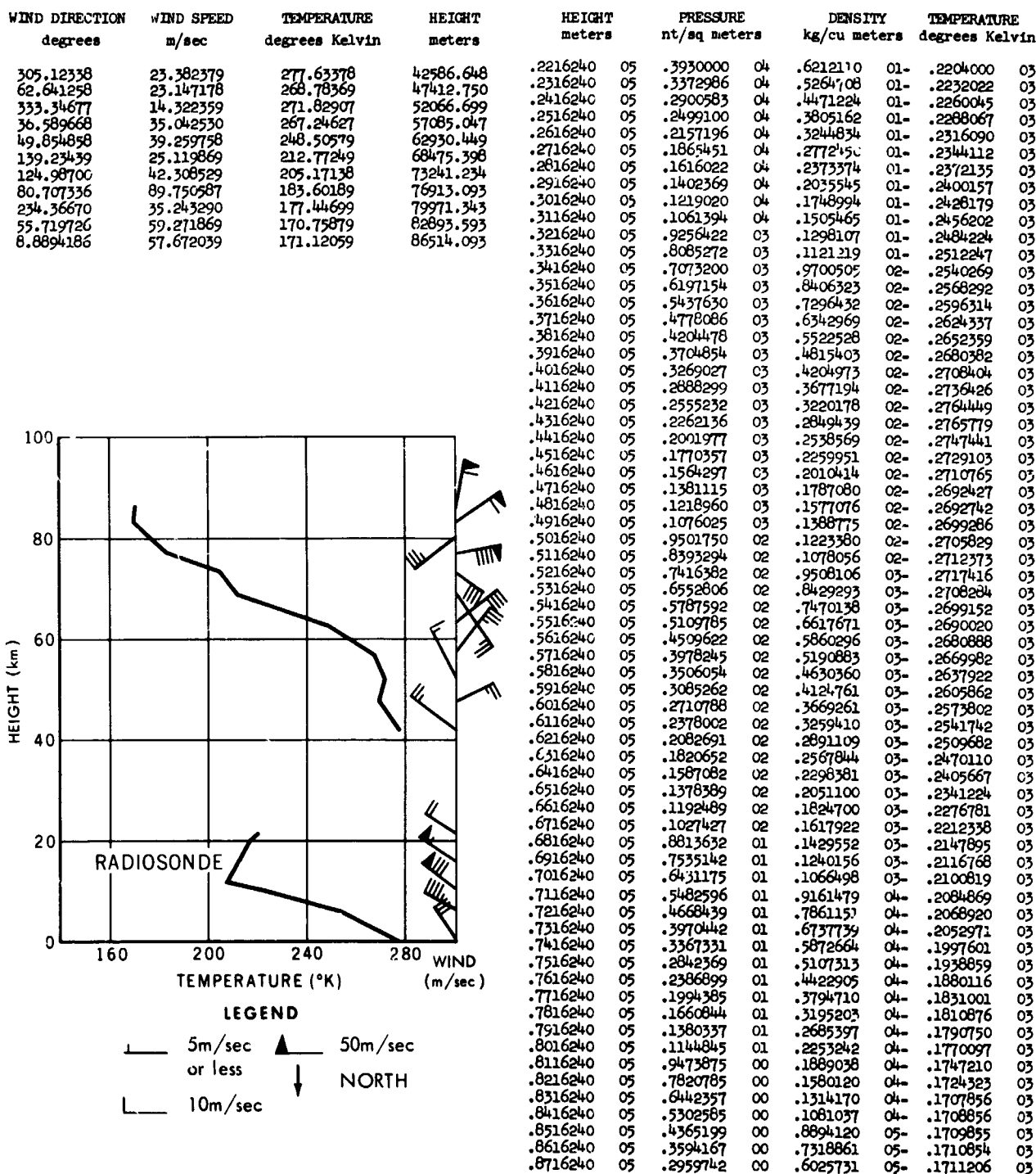


FIGURE 6

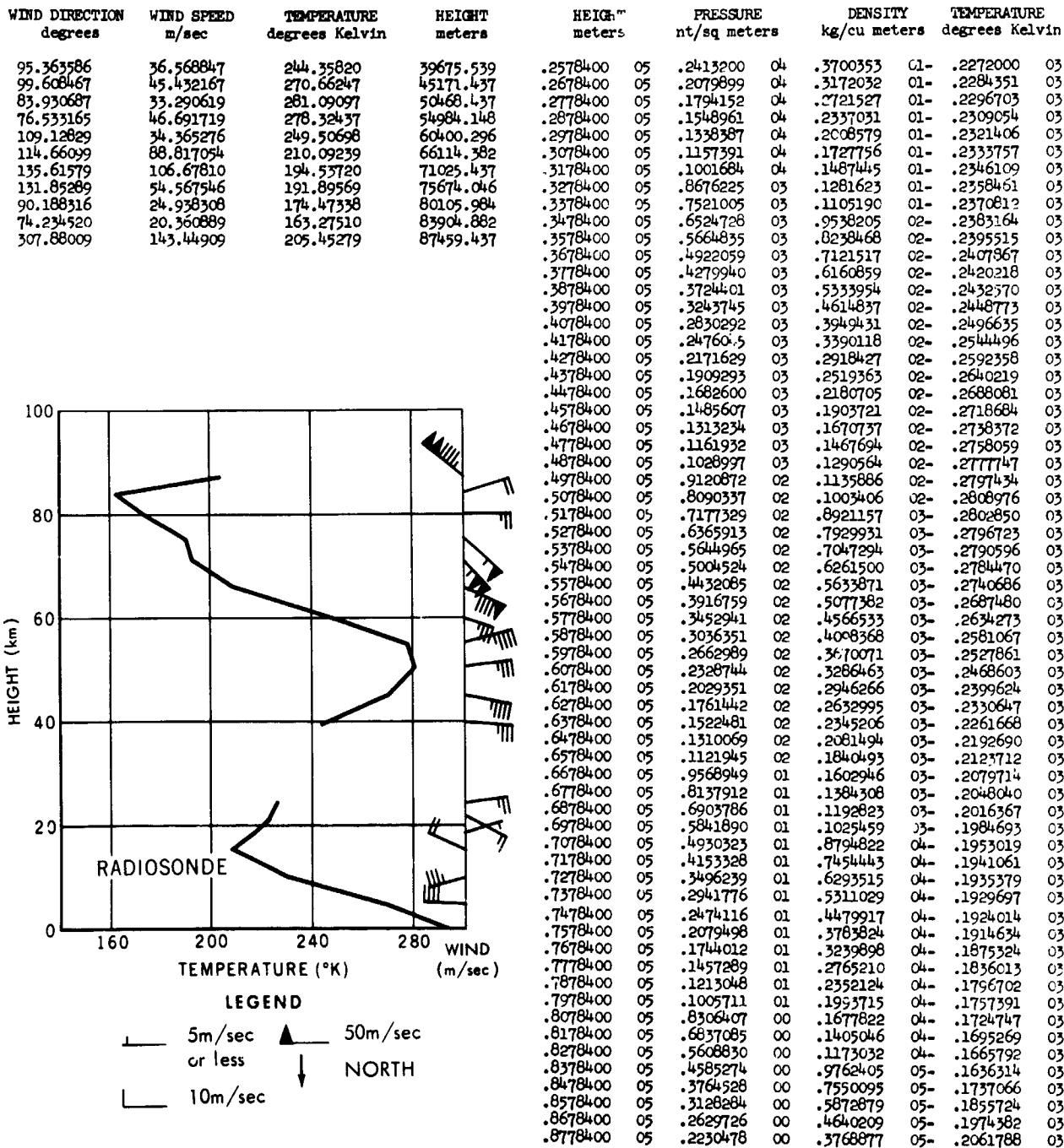


FIGURE 7

.3 second was introduced into t . As in the case of the coordinates of the grenade explosions, these errors were introduced only for alternate grenades.

A typical result of the error analysis is presented in Figures 29 through 38. From a study of these figures, the following conclusions can be reached:

1. Relative errors in the North and West coordinates of 200 meters will significantly affect only winds by about 10 m/sec but not the temperatures.
2. Relative errors in the up coordinates of 50 meters will cause significant errors of about 5°K in temperature, but not in winds.

3. An error of .3 second in t will give rise to large temperature errors—about 10°K —but in general only very small wind errors.
4. An error of .02 in the relative arrival times between microphones will cause the largest errors in both winds and temperatures: 10–40 m/sec and $10\text{--}25^{\circ}\text{K}$. The error analysis clearly shows that an error introduced into the input parameters of a grenade significantly affects only the layers adjacent to the particular grenade. No significant error is introduced into the calculated winds and temperature for suc-

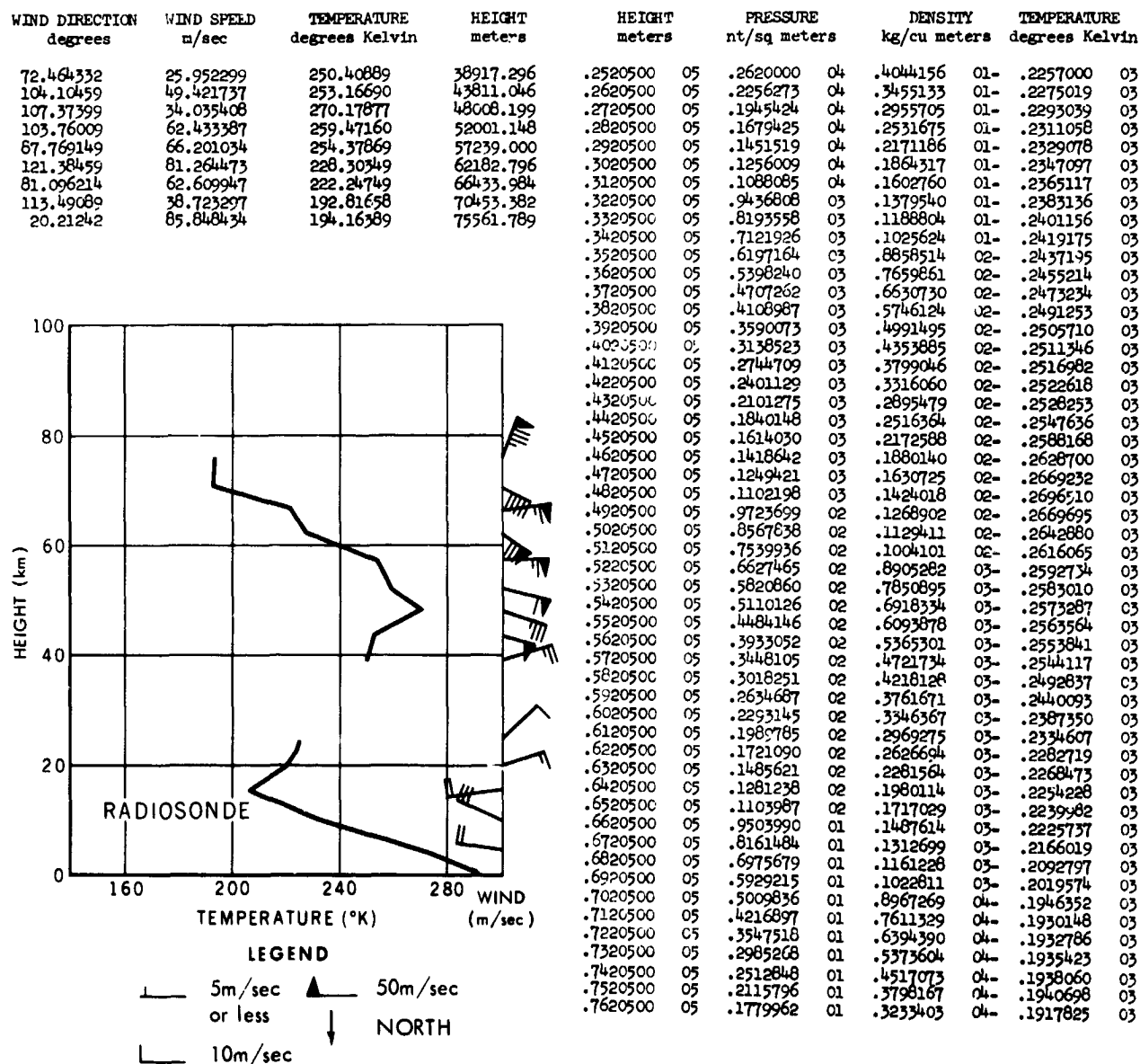


FIGURE 8

ceeding grenades, and hence only the layers adjacent to the grenade in question need be considered.

Table 2 gives the average error functions for the 16 firings tabulated by layers. Figures 39 and 40 are graphs of the two functions which are most strongly altitude dependent, $\partial T/\partial \Delta t$ and $\partial W/\partial \Delta t$, showing quite clearly the increase of the error function with altitude.

$\partial W/\partial \Delta h$, etc., by the value of the actual errors in Δw , Δn , h , Δt , and t . The maximum error is calculated by summing all the partial errors:

$$\left(\frac{\partial W}{\partial \Delta w} d\Delta w + \frac{\partial W}{\partial \Delta n} d\Delta n + \frac{\partial W}{\partial \Delta h} d\Delta h + \dots \text{etc.} \right)$$

The actual errors in coordinates and times are estimated in the following manner:

The position of the grenade explosion is determined by two totally independent tracking meth-

Table 2
Absolute Average Error Functions

Layer	t + .3 sec		North-200m		West-200m		Up-50m		$\Delta t + .02$ sec	
	Temp. (°K)	Wind (m/sec)	Temp. (°K)	Wind (m/sec)	Temp. (°K)	Wind (m/sec)	Temp. (°K)	Wind (m/sec)	Temp. (°K)	Wind (m/sec)
1	10.1	.5	.9	3.9	.7	9.2	5.4	.8	11.7	12.2
2	10.2	.7	.9	5.3	.7	10.6	5.6	.6	15.2	18.5
3	13.0	1.0	.9	5.2	.8	11.9	6.7	.8	16.3	22.0
4	10.5	1.1	.6	4.6	.6	10.7	5.8	.6	17.8	22.3
5	8.4	.9	.4	2.5	.4	9.5	4.6	.5	12.0	16.9
6	7.7	1.0	.4	3.7	.5	8.7	4.5	.7	15.1	26.4
7	10.0	1.3	.4	6.0	.5	10.6	5.6	.8	16.3	25.3
8	9.2	1.2	.4	6.3	.7	10.3	5.6	.7	20.9	31.9
9	11.7	2.3	.5	5.8	.8	12.1	6.4	.8	21.4	30.5
10	8.3	2.4	.5	8.2	.8	9.3	5.4	.7	22.5	41.1
11	11.5	5.2	.4 ^b	9.8	.7	11.4	6.4	.5	22.4	35.2

Figure 41 shows the dependence of $\partial T/\partial \Delta t$ upon zenith angle. As the zenith angle increases, the error function increases nearly linearly. Thus, to make the experiments as accurate as possible, it is desirable that the sound propagation be as nearly vertical as possible. A zenith angle of 0° would, of course, be ideal; however, range safety considerations make this impossible, and at Wallops Island the zenith angles usually lie in the range from 15° — 30° . The zenith angles at Fort Churchill are generally less than 10° , and the probable error in the Fort Churchill data is therefore correspondingly less.

The magnitudes of the errors assumed in this analysis were chosen arbitrarily large for convenience in calculation. The actual resulting errors will be much smaller, as will be shown below. The magnitude of the actual errors in the rocket grenade experiment is determined by multiplying the error functions $\partial W/\partial \Delta w$, $\partial W/\partial \Delta n$,

ods, DOVAP and radar. The agreement of these two systems with regard to the coordinates Δn , Δw and h in general is better than 20 meters. The result is a maximum temperature error of 1°K and maximum wind error of 1–2 meters/sec, according to the error functions. The time of the grenade explosion, determined primarily by rocket borne infrared photocells, is telemetered to the ground equipment. The time of grenade explosions may also be determined by ground based flash detectors or radar signal strength records. It has been shown that this time can generally be determined within $\pm .001$ second. The errors in determining the arrival time of the sound at the ground range from $\pm .001$ to $\pm .010$ second. Thus, the maximum error expected in determining the travel time (t) of the sound between explosion and ground is approximately .01 second, which results in a temperature error of 0.3°K and virtually no errors in winds. By far the largest contribution

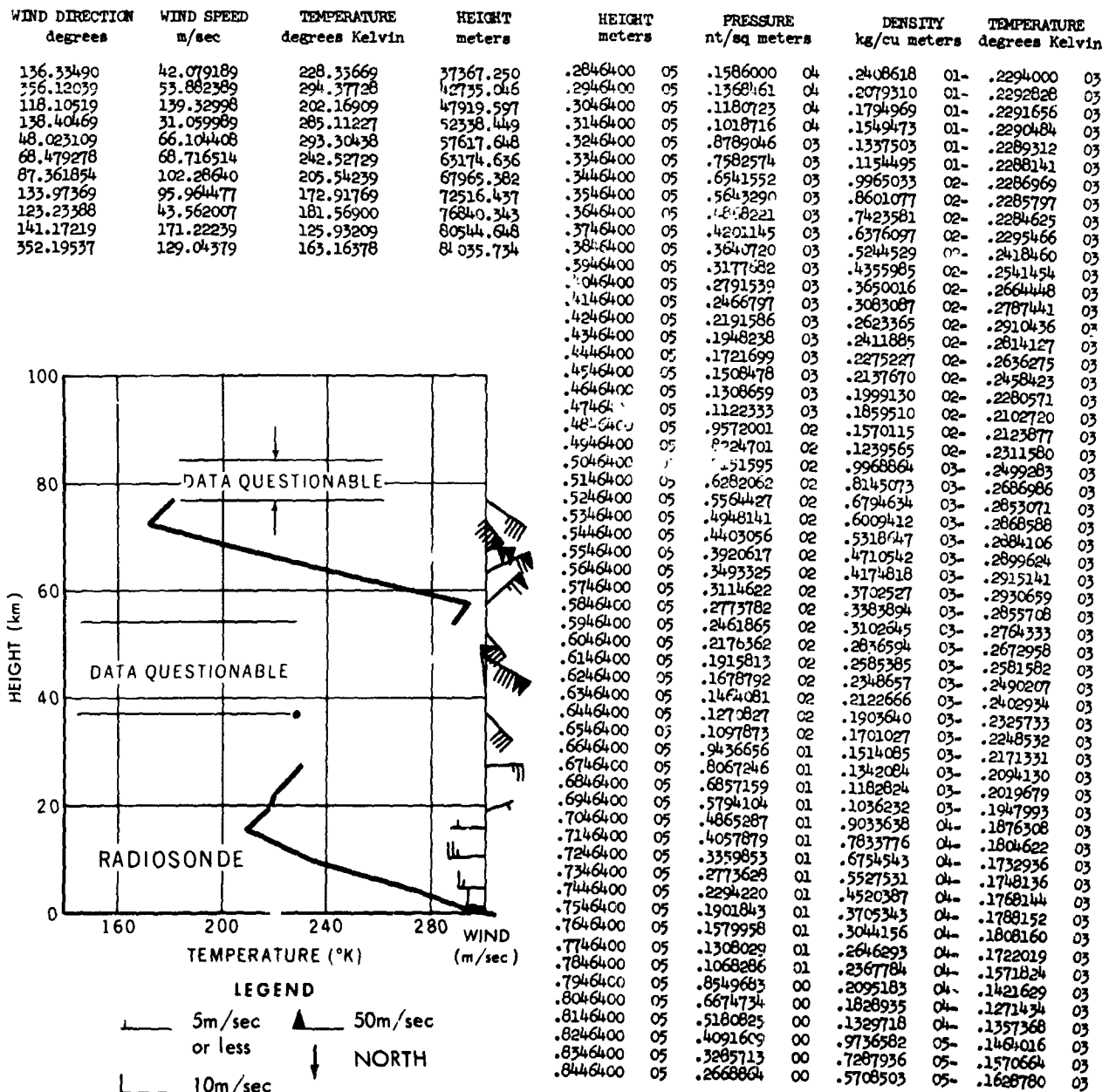


FIGURE 9

to the total error comes from errors in Δt , which also range from .001 to .010 second, depending upon the background noise level and the altitude of the explosion. The error in Δt is obtained by determining typical discrepancies among a number of independent Δt readings for each explosion. These errors result in temperature errors ranging from 1°K at 40 km to 15°K at 90 km, and wind errors from 1 to 16 m/sec over the same altitude range.

The maximum errors in the 16 soundings considered in this error analysis are presented in Figures 42 and 43, which clearly show the increase of the maximum errors with altitude. This increase is to be expected, since the amplitude of the sound waves will decrease rapidly with increas-

ing altitudes, making it difficult to distinguish the grenade explosion from the background noise level. Since the determination of the relative arrival times between microphones is by far the most critical measured parameter, the largest part of the maximum error is due to the attenuation of the sound wave with altitude.

ACKNOWLEDGMENTS

The authors greatly appreciate the contributions of: Superior Engineering Company in designing and constructing the payload instrumentation for the grenade experiment; Texas Western College in conducting the sound ranging portion of this experiment; and New Mexico State University in operating the DOVAP tracking system.

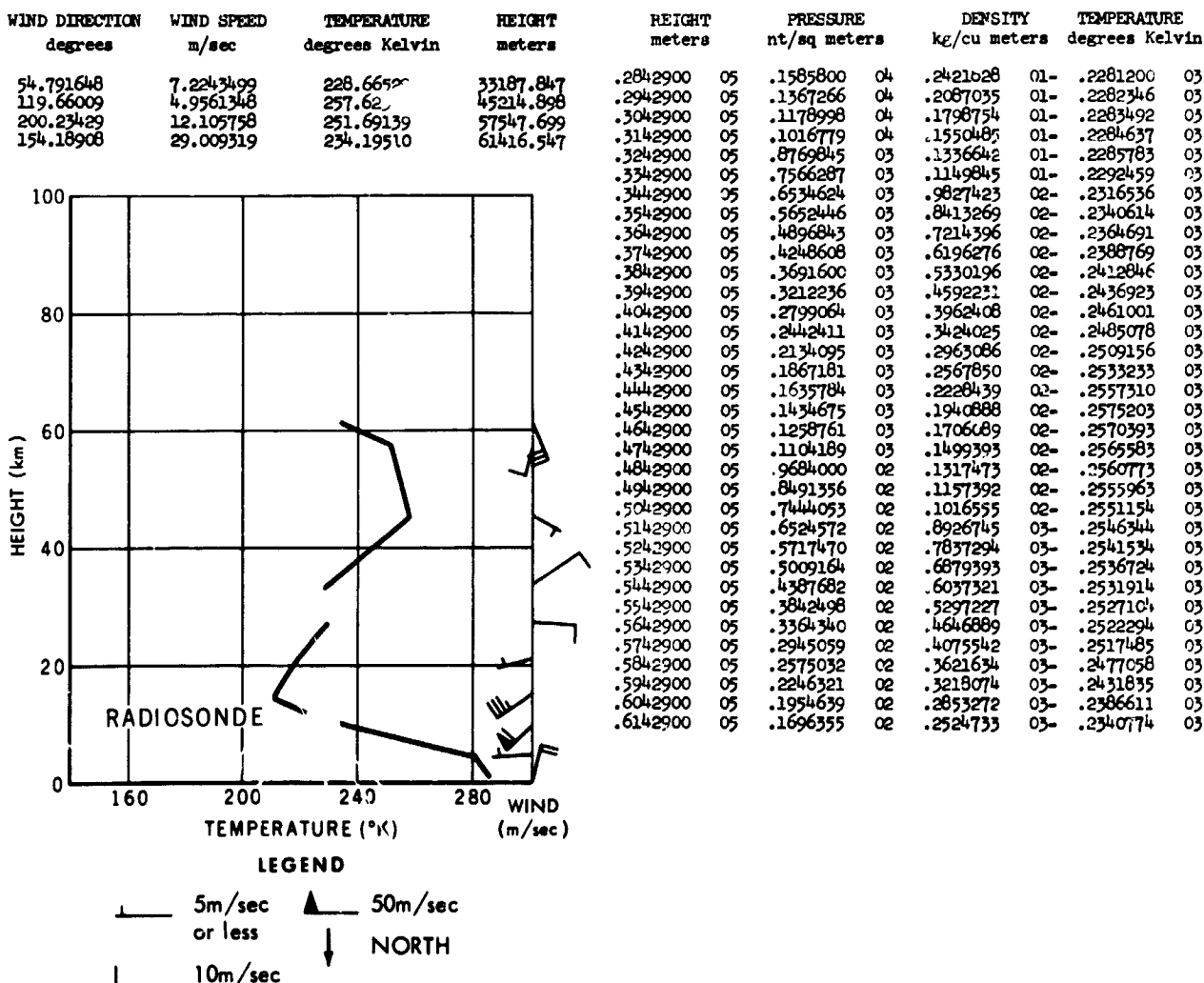


FIGURE 10

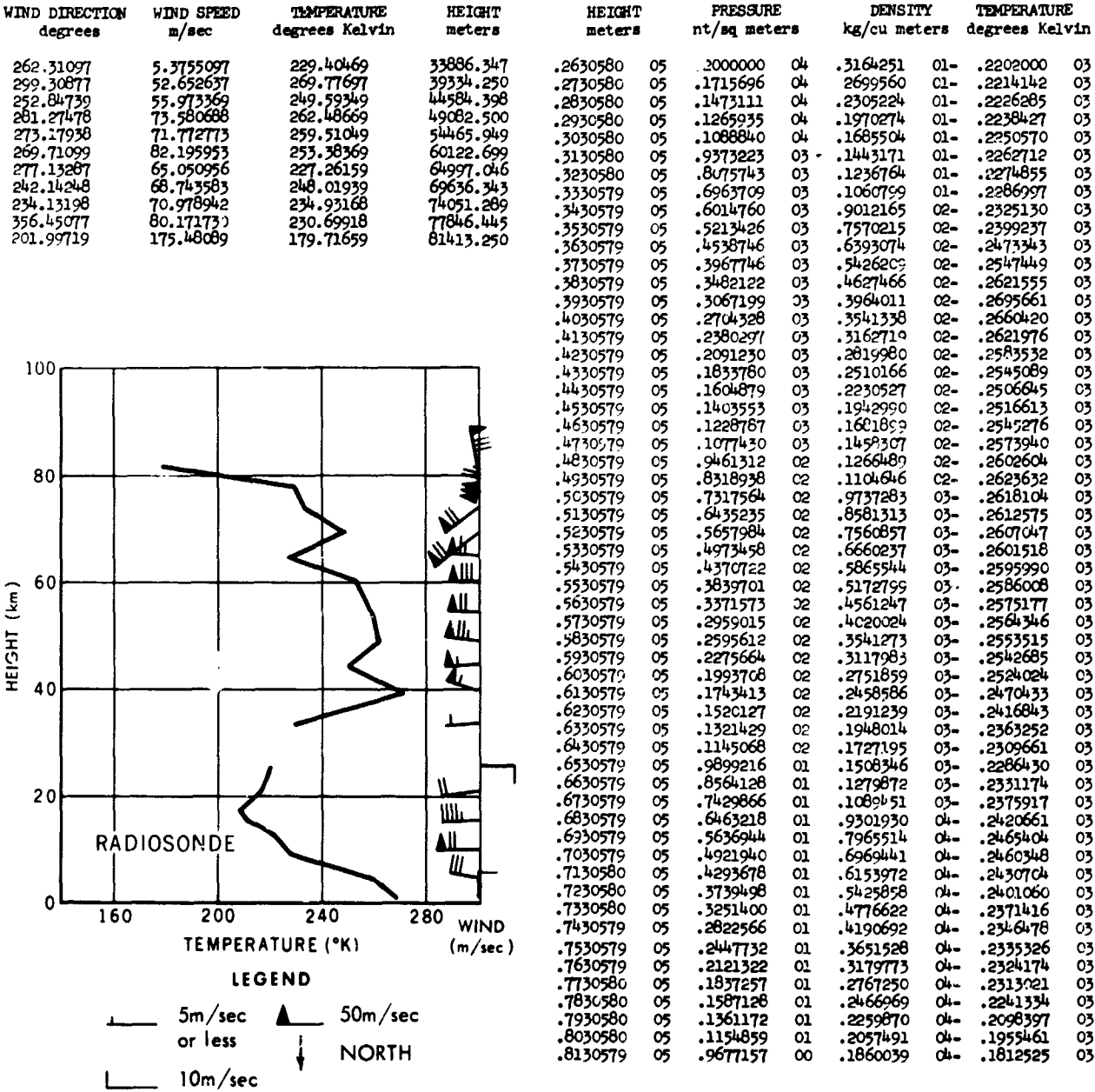
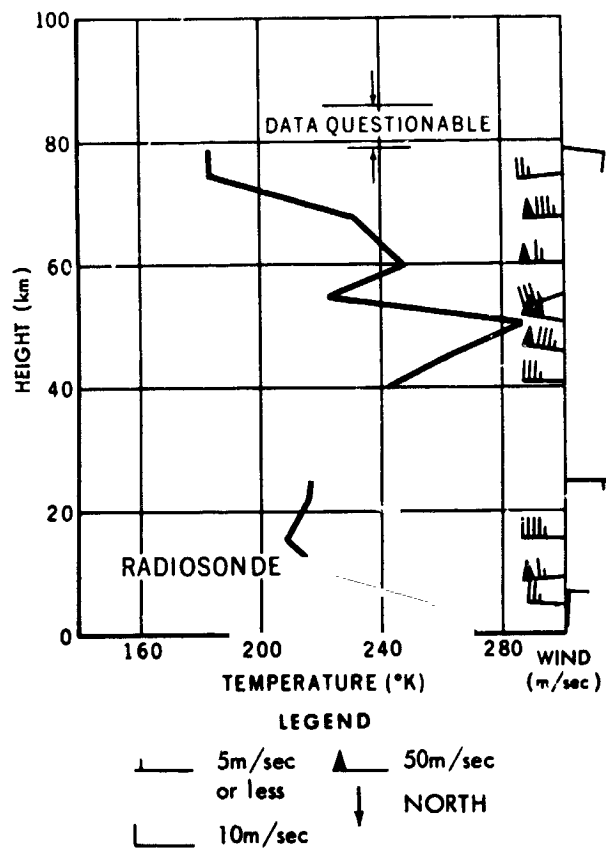


FIGURE 11

WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees Kelvin
274.23077	34.899269	242.51269	40513.449	.2478430	05 .2600000	04 .4192929	01- .2160300
286.33287	85.033699	263.57437	45697.347	.2578430	05 .2224041	04 .3558897	01- .2177136
281.50439	101.26220	285.79770	50676.898	.2678430	05 .1904829	04 .3024704	01- .2193973
243.78489	33.179687	222.30879	54934.796	.2778430	05 .1633445	04 .2574016	01- .2210810
268.60199	62.581386	247.68548	60029.648	.2878430	05 .1402426	04 .2193269	01- .2227647
260.13058	85.637725	231.54458	67638.132	.2978430	05 .1205521	04 .1871185	01- .2244483
259.93988	24.920852	183.24919	74377.484	.3078430	05 .1037483	04 .1598369	01- .2261320
106.51109	10.605539	183.39129	78539.037	.3178430	05 .8939038	03 .1366990	01- .2278157
296.23410	234.90919	391.75579	82104.632	.3278429	05 .7710758	03 .1170506	01- .2294993
150.42019	103.22889	152.16539	85453.297	.3378429	05 .6658746	03 .1003448	01- .2311830



.2478430	05	.2600000	04	.4192929	01-	.2160300	03
.2578430	05	.2224041	04	.3558897	01-	.2177136	03
.2678430	05	.1904829	04	.3024704	01-	.2193973	03
.2778430	05	.1633445	04	.2574016	01-	.2210810	03
.2878430	05	.1402426	04	.2193269	01-	.2227647	03
.2978430	05	.1205521	04	.1871185	01-	.2244483	03
.3078430	05	.1037483	04	.1598369	01-	.2261320	03
.3178430	05	.8939038	03	.1366990	01-	.2278157	03
.3278429	05	.7710758	03	.1170506	01-	.2294993	03
.3378429	05	.6658746	03	.1003448	01-	.2311830	03
.3478429	05	.5756651	03	.8612332	02-	.2328667	03
.3578429	05	.4982214	03	.7400217	02-	.2345503	03
.3678429	05	.4316613	03	.6365886	02-	.2362340	03
.3778429	05	.3743909	03	.5482224	02-	.2379177	03
.3878429	05	.3250594	03	.4726412	02-	.2396014	03
.3978429	05	.2825198	03	.4079215	02-	.2412850	03
.4078429	05	.2458427	03	.3515725	02-	.2436131	03
.4178429	05	.2143231	03	.3014694	02-	.2476760	03
.4278429	05	.1872701	03	.2591650	02-	.2517389	03
.4378429	05	.1639926	03	.2233464	02-	.2558018	03
.4478429	05	.1437150	03	.1929377	02-	.2598647	03
.4578429	05	.1265579	03	.1670341	02-	.2639624	03
.4678429	05	.1115287	03	.1447509	02-	.2684253	03
.4778429	05	.9849314	02	.1257417	02-	.2728882	03
.4878429	05	.8716010	02	.1094828	02-	.2773511	03
.4978429	05	.7720464	02	.9554075	03-	.2818141	03
.5078429	05	.6863049	02	.8413115	03-	.2841963	03
.5178429	05	.6078233	02	.7863622	03-	.2892854	03
.5278429	05	.5346258	02	.7322078	03-	.2943746	03
.5378429	05	.4666310	02	.6788785	03-	.2994637	03
.5478429	05	.4037547	02	.6264078	03-	.3045529	03
.5578429	05	.3479388	02	.5755767	03-	.3096420	03
.5678429	05	.3005305	02	.5222272	03-	.3147312	03
.5778429	05	.2604042	02	.4835940	03-	.3198203	03
.5878429	05	.2263203	02	.4450095	03-	.3249094	03
.5978429	05	.1972702	02	.4074777	03-	.3300000	03
.6078429	05	.1721774	02	.3715531	03-	.3350906	03
.6178429	05	.1501787	02	.3364581	03-	.3401812	03
.6278429	05	.1306400	02	.3021810	03-	.3452718	03
.6378429	05	.1133582	02	.2684604	03-	.3503624	03
.6478429	05	.9896233	01	.2351055	03-	.3554530	03
.6578429	05	.8591097	01	.2021035	03-	.3605436	03
.6678429	05	.7448874	01	.1712064	03-	.3656342	03
.6778429	05	.6448914	01	.1417186	04-	.3707248	03
.6878430	05	.5565639	01	.11362091	04-	.3758154	03
.6978430	05	.4780546	01	.87704613	04-	.3809060	03
.7078429	05	.4085400	01	.6810036	04-	.3859966	03
.7178429	05	.3472421	01	.5093764	04-	.3910872	03
.7278429	05	.2934271	01	.35251311	04-	.3961778	03
.7378429	05	.2464043	01	.21457308	04-	.4012684	03
.7478430	05	.2058228	01	.13912701	04-	.4063590	03
.7578430	05	.1715362	01	.9261254	04-	.4114496	03
.7678430	05	.1430568	01	.6718503	04-	.4165402	03
.7778429	05	.1192818	01	.4266286	04-	.4216308	03
.7878429	05	.1001379	01	.21764406	04-	.4267214	03
.7978429	05	.8639777	00	.1175026	04-	.4318120	03
.8078430	05	.7685141	00	.8510464	05-	.4369026	03
.8178430	05	.6974343	00	.6513443	05-	.4419932	03
.8278430	05	.6355071	00	.4652434	05-	.4470838	03
.8378430	05	.5700333	00	.312438	05-	.4521744	03
.8478430	05	.4945087	00	.18612610	05-	.4572650	03

FIGURE 12

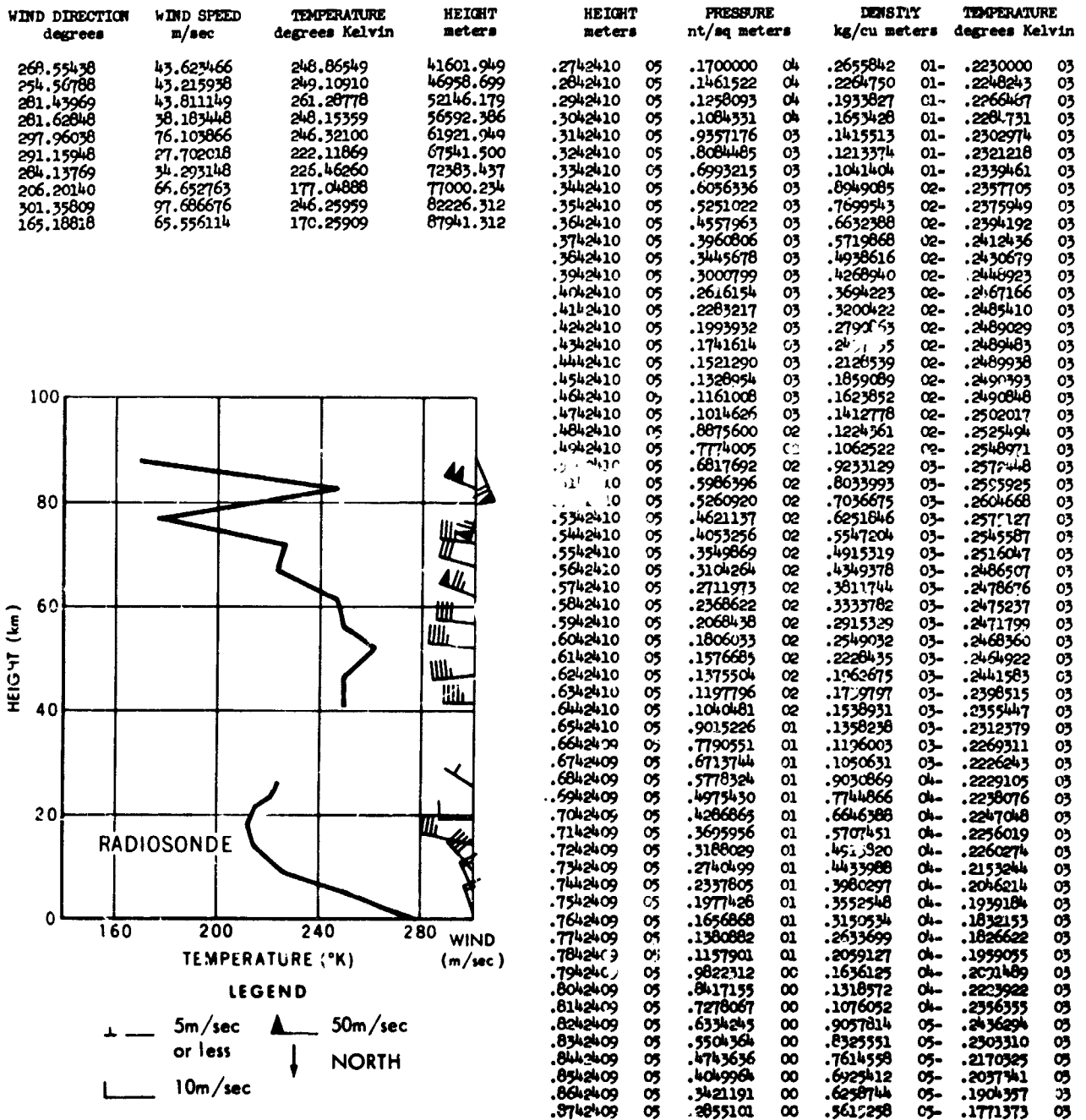


FIGURE 13

WIND DIRECTION	WIND SPEED	TEMPERATURE	HEIGHT	HEIGHT	PRESSURE	DENSITY	TEMPERATURE				
degrees	m/sec	degrees Kelvin	meters	meters	nt/sq meters	kg/cu meters	degrees Kelvin				
262.26339	37.318786	261.81979	42956.867	.2692060	05	.1800000	04	.2833670	01-	.2213000	03
258.63888	59.576107	258.50137	48370.109	.2792060	05	.1545038	04	.2406303	01-	.2238267	03
264.73590	57.798050	262.25830	53607.328	.2892060	05	.1330240	04	.2047392	01-	.2263535	03
264.37527	55.014717	256.49829	58100.378	.2992060	05	.1146528	04	.1745157	01-	.2288803	03
222.00199	1.9646849	228.56479	63469.269	.3092060	05	.098465	03	.1453217	01-	.2314070	03
267.01077	53.375399	225.95019	69130.968	.3192060	05	.0859803	03	.1274762	01-	.2339338	03
266.98429	27.363178	227.11749	74019.578	.3292060	05	.07414078	03	.1092337	01-	.2364605	03
259.86737	37.858348	213.65579	78665.953	.3392060	05	.06431811	03	.9375978	02-	.2389873	03
291.07809	93.457389	213.82579	83089.789	.3492060	05	.5588272	03	.8061081	02-	.2415140	03
115.27018	117.01439	151.20039	87297.484	.3592060	05	.4862684	03	.6941795	02-	.2440408	03
355.37219	95.003875	205.23109	91614.125	.3692060	05	.4237558	03	.5987394	02-	.2465676	03
				.3792060	05	.3698137	03	.5172223	02-	.2490943	03
				.3892060	05	.3231954	03	.4474827	02-	.2516211	03
				.3992060	05	.2828461	03	.3877233	02-	.2541478	03
				.4092060	05	.2478712	03	.3364351	02-	.2567446	03
				.4192060	05	.2175109	03	.2923492	02-	.2593014	03
				.4292060	05	.1911191	03	.2543969	02-	.2617281	03
				.4392060	05	.1680204	03	.2240778	02-	.2642290	03
				.4492060	05	.1476788	03	.1974128	02-	.2667159	03
				.4592060	05	.1297656	03	.1738760	02-	.2692029	03
				.4692060	05	.1139952	03	.1531058	02-	.2716899	03
				.4792060	05	.1001146	03	.1347814	02-	.2741769	03
				.4892060	05	.8791714	02	.1183062	02-	.2766632	03
				.4992060	05	.7722527	02	.1036311	02-	.2791536	03
				.5092060	05	.6786043	02	.9081320	03-	.2816403	03
				.5192060	05	.5965486	02	.7961284	03-	.2841273	03
				.5292060	05	.5246212	02	.6982185	03-	.2866143	03
				.5392060	05	.4614768	02	.6139661	03-	.2891013	03
				.5492060	05	.4058291	02	.5425867	03-	.2915883	03
				.5592060	05	.3566801	02	.4792330	03-	.2940753	03
				.5692060	05	.3132953	02	.4230331	03-	.2965623	03
				.5792060	05	.2750211	02	.3732070	03-	.2990493	03
				.5892060	05	.2410731	02	.3213449	03-	.3015363	03
				.5992060	05	.2107799	02	.2946640	03-	.3040233	03
				.6092060	05	.1837712	02	.2753440	03-	.3065103	03
				.6192060	05	.1597495	02	.2585715	03-	.3090000	03
				.6292060	05	.1384386	02	.2427121	03-	.3114900	03
				.6392060	05	.1196664	02	.2288628	03-	.3139800	03
				.6492059	05	.1033289	02	.2150078	03-	.3164700	03
				.6592059	05	.8920137	01	.2000000	03-	.3189600	03
				.6692059	05	.7698772	01	.1982112	03-	.3214500	03
				.6792059	05	.6643095	01	.1821939	03-	.3239400	03
				.6892059	05	.5730858	01	.1682686	04-	.3264300	03
				.6992059	05	.4943632	01	.1561604	04-	.3289200	03
				.7092059	05	.4265180	01	.1456905	04-	.3314100	03
				.7192059	05	.3680600	01	.1365826	04-	.3339000	03
				.7292059	05	.3176766	01	.1287857	04-	.3363900	03
				.7392059	05	.2742458	01	.12207197	04-	.3388800	03
				.7492059	05	.2366110	01	.1163335	04-	.3413700	03
				.7592059	05	.2038362	01	.1112640	04-	.3438600	03
				.7692059	05	.1753177	01	.1071657	04-	.3463500	03
				.7792059	05	.1505403	01	.10311364	04-	.3488400	03
				.7892059	05	.1290703	01	.10016102	04-	.3513300	03
				.7992059	05	.1105743	01	.10790597	04-	.3538200	03
				.8092059	05	.9470517	00	.1056573	04-	.3563100	03
				.8192059	05	.8109328	00	.10318262	04-	.3588000	03
				.8292059	05	.6942042	00	.1130039	04-	.3612900	03
				.8392059	05	.5914250	00	.9303110	05-	.3637800	03
				.8492059	05	.4982278	00	.8396304	05-	.3662700	03
				.8592059	05	.4137934	00	.7508170	05-	.3687600	03
				.8692059	05	.3379475	00	.6999732	05-	.3712500	03
				.8792059	05	.2738231	00	.6548786	05-	.3737400	03
				.8892059	05	.2239461	00	.6168878	05-	.3762300	03
				.8992059	05	.1857787	00	.5865654	05-	.3787200	03

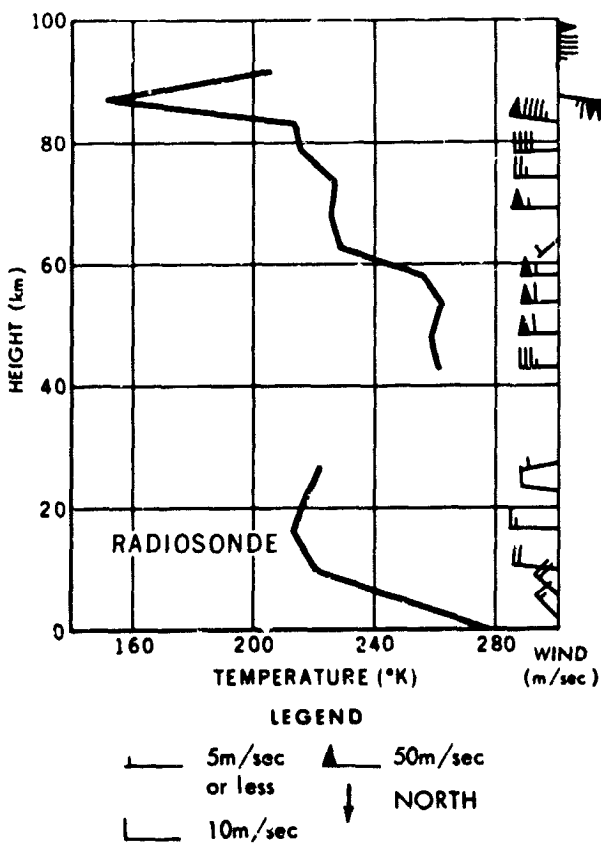


FIGURE 14

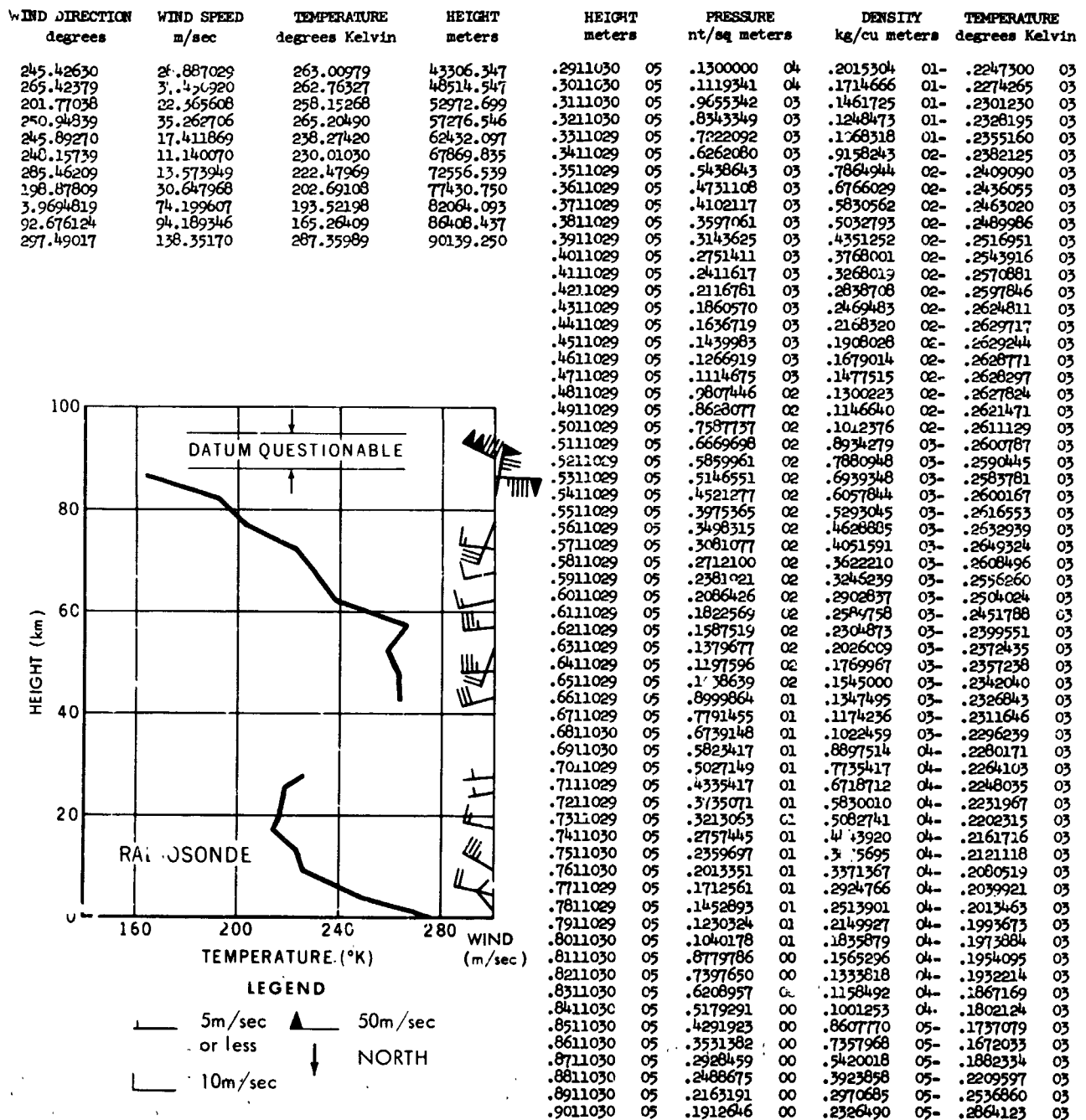
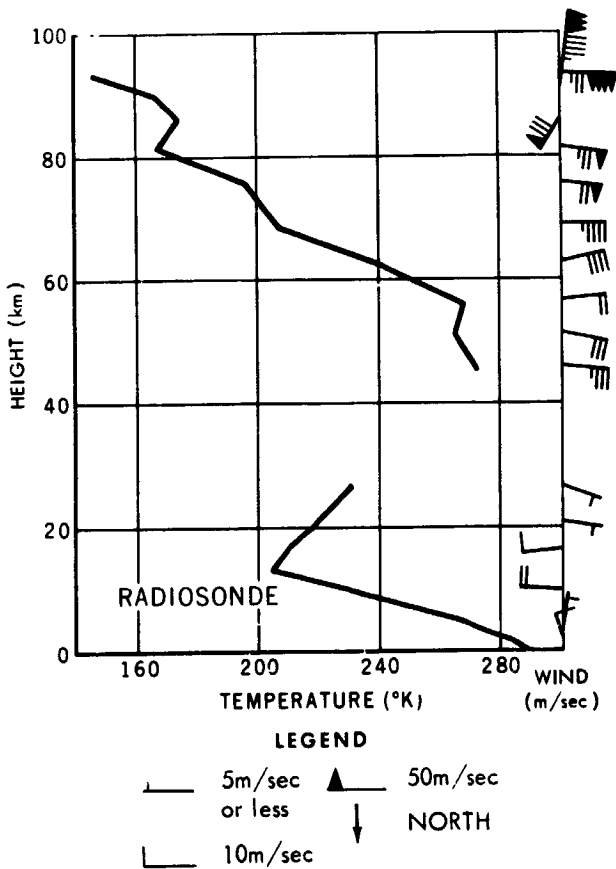


FIGURE 15

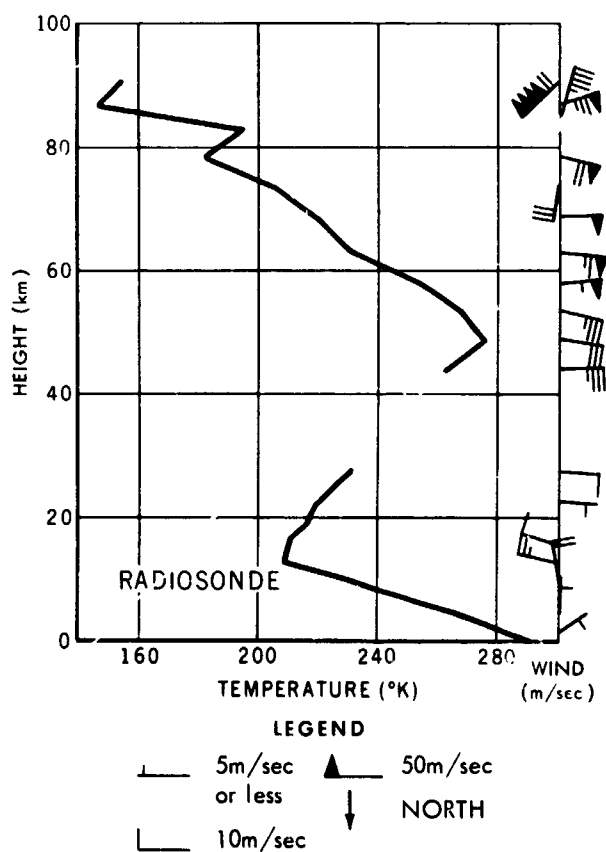
WIND DIRECTION	WIND SPEED	TEMPERATURE	HEIGHT	HEIGHT	PRESSURE	DENSITY	TEMPERATURE				
degrees	m/sec	degrees Kelvin	meters	meters	nt/sq meters	kg/cu meters	degrees Kelvin				
100.23349	35.039177	271.12729	45967.500	.2782370	05	.1700000	04	.2561204	01-	.2312400	03
112.12339	31.985488	265.44738	51428.398	.2882370	05	.1469582	04	.2193208	01-	.2334384	03
85.151390	20.540390	268.92038	56624.398	.2982370	05	.1272189	04	.1880904	01-	.2356366	03
76.249366	37.992378	241.96479	62624.597	.3082370	05	.1102834	04	.1615445	01-	.2376352	03
86.465759	47.108570	208.31379	69435.039	.3182370	05	.0973240	03	.1389457	01-	.2400336	03
102.19909	72.065656	197.25399	75832.382	.3282370	05	.0832120	03	.1196778	01-	.2422320	03
100.33849	77.084068	168.90759	81279.049	.3382370	05	.0724249	03	.1032254	01-	.2444304	03
219.76649	81.912193	174.49778	85927.898	.3482370	05	.06311593	03	.0915658	02-	.2466288	03
5.6612997	193.16529	167.00509	89940.898	.3582370	05	.05507353	03	.07710869	02-	.2488272	03
97.556236	223.62478	145.78129	93665.695	.3682370	05	.04811559	03	.06677687	02-	.2510256	03
				.3782370	05	.04008920	03	.05709443	02-	.2530240	03



.3782370	05	.4208802	03	.5790443	02-	.2532240	03
.3882370	05	.3685967	03	.5027484	02-	.2554224	03
.3982370	05	.3231886	03	.4370523	02-	.2576207	03
.4082370	05	.2837027	03	.3804088	02-	.2598191	03
.4182370	05	.2493247	03	.3315073	02-	.2620175	03
.4282370	05	.2193578	03	.2892359	02-	.2642159	03
.4382370	05	.1932052	03	.2526500	02-	.2664143	03
.4482370	05	.1703549	03	.2209460	02-	.2686127	03
.4582370	05	.1503671	03	.1934393	02-	.2708111	03
.4682370	05	.1327794	03	.1711766	02-	.2720267	03
.4782370	05	.1172096	03	.1516882	02-	.2691966	03
.4882370	05	.1034197	03	.1343609	02-	.2681565	03
.4982370	05	.9121133	02	.1189615	02-	.2671164	03
.5082370	05	.8040792	02	.1052812	02-	.2660763	03
.5182370	05	.7086315	02	.9291121	03-	.2657116	03
.5282370	05	.6245859	02	.8168621	03-	.2663800	03
.5382370	05	.5507040	02	.7184333	03-	.2670483	03
.5482370	05	.4857337	02	.6320927	03-	.2677168	03
.5582370	05	.4285786	02	.5563270	03-	.2683852	03
.5682370	05	.3781911	02	.4915637	03-	.2680339	03
.5782370	05	.3333633	02	.4406094	03-	.2635859	03
.5882370	05	.2932304	02	.3942177	03-	.2591379	03
.5982370	05	.2573672	02	.3520461	03-	.2546900	03
.6082370	05	.2253809	02	.3137727	03-	.2502420	03
.6182370	05	.1969090	02	.2750953	03-	.2457940	03
.6282370	05	.1716138	02	.2478020	03-	.2412713	03
.6382370	05	.1491656	02	.2199319	03-	.2362863	03
.6482370	05	.1292725	02	.1947090	03-	.2313013	03
.6582369	05	.1116886	02	.1719297	03-	.2263163	03
.6682369	05	.9618718	01	.1514023	03-	.2213313	03
.6782369	05	.8255966	01	.1329464	03-	.2163463	03
.6882369	05	.7061436	01	.1163927	03-	.2113613	03
.6982369	05	.6020476	01	.1010123	03-	.2070419	03
.7082369	05	.5122460	01	.8666698	04-	.2059131	03
.7182369	05	.4352695	01	.7426672	04-	.2041842	03
.7282369	05	.3693669	01	.6356042	04-	.2024554	03
.7382369	05	.3130174	01	.5432774	04-	.2007266	03
.7482369	05	.2648987	01	.4637561	04-	.1989978	03
.7582369	05	.2238624	01	.3953489	04-	.1972690	03
.7682369	05	.1886279	01	.3420963	04-	.1920948	03
.7782369	05	.1582039	01	.2949091	04-	.1868905	03
.7882369	05	.1320369	01	.2531813	04-	.1816861	03
.7982369	05	.1096266	01	.2164086	04-	.1764818	03
.8082369	05	.9052001	00	.1841207	04-	.1712774	03
.8182369	05	.7445886	00	.1529835	04-	.1695625	03
.8282369	05	.6123329	00	.1249243	04-	.1707650	03
.8382369	05	.5042907	00	.1021628	04-	.1719675	03
.8482369	05	.4158986	00	.8367063	05-	.1731700	03
.8582369	05	.3434779	00	.6862448	05-	.1743725	03
.8682369	05	.2836304	00	.5717470	05-	.1728252	03
.8782369	05	.2337198	00	.4764040	05-	.1709581	03
.8882369	05	.1922939	00	.3961897	05-	.1690910	03
.8982369	05	.1578366	00	.3288277	05-	.1672239	03
.9082369	05	.1290005	00	.2774782	05-	.1619749	03
.9182369	05	.1047229	00	.2334559	05-	.1562769	03
.9282369	05	.8435857	01-	.1951745	05-	.1505790	03
.9382369	05	.6706524	01-	.1684508	05-	.1387020	03

FIGURE 16

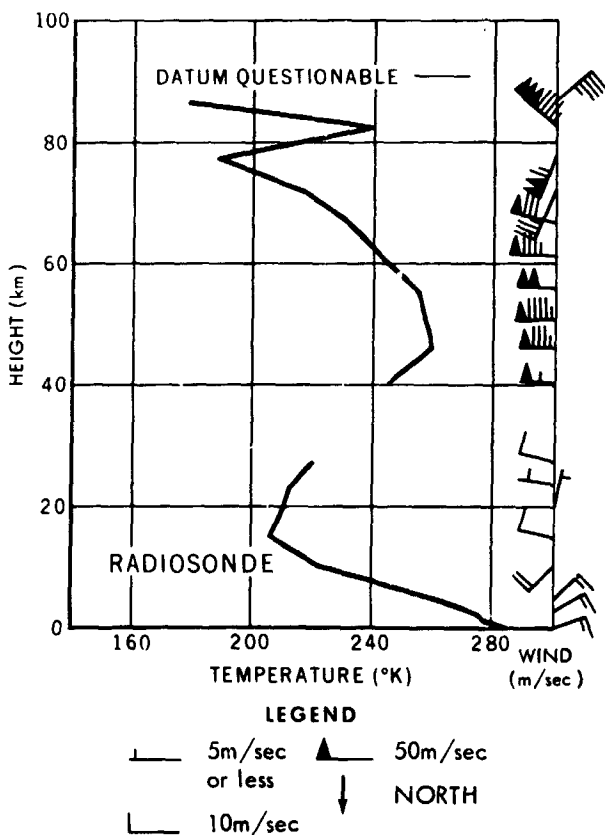
WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees Kelvin
88.135818	35.642078	262.44598	43731.648	.3128600	05 .1000000	04 .1498425	01- .2325000
118.24260	29.022520	275.65139	49039.000	.3228600	05 .8653305	03 .1283352	01- .2349061
121.08769	36.761146	266.55227	53575.898	.3328600	05 .7499351	03 .1100935	01- .2373123
80.334419	54.959678	254.40489	57950.200	.3428600	05 .6508961	03 .9459506	02- .2397184
96.692169	64.717727	231.56219	63171.789	.3528600	05 .5657608	03 .8140521	02- .2421245
80.894264	51.982868	220.42698	68672.593	.3628600	05 .4924644	03 .7016162	02- .2445307
192.69558	28.052829	206.80468	73406.648	.3728600	05 .4290652	03 .6056169	02- .2469368
121.72178	70.112663	182.82728	78311.234	.3828600	05 .3746909	03 .5235211	02- .2493430
31.833500	44.285537	194.84425	82984.882	.3928600	05 .3274960	03 .4532067	02- .2517491
70.778137	72.619316	147.24949	86953.234	.4028600	05 .2866244	03 .3928911	02- .2541552
224.92109	217.80018	154.34379	90628.484	.4128600	05 .2511792	03 .3410755	02- .2565614



.3128600	05	.1000000	04	.1498425	01-	.2325000	03
.3228600	05	.8653305	03	.1283352	01-	.2349061	03
.3328600	05	.7499351	03	.1100935	01-	.2373123	03
.3428600	05	.6508961	03	.9459506	02-	.2397184	03
.3528600	05	.5657608	03	.8140521	02-	.2421245	03
.3628600	05	.4924644	03	.7016162	02-	.2445307	03
.3728600	05	.4290652	03	.6056169	02-	.2469368	03
.3828600	05	.3746909	03	.5235211	02-	.2493430	03
.3928600	05	.3274960	03	.4532067	02-	.2517491	03
.4028600	05	.2866244	03	.3928911	02-	.2541552	03
.4128600	05	.2511792	03	.3410755	02-	.2565614	03
.4228600	05	.2203976	03	.2964967	02-	.2589675	03
.4328600	05	.1936301	03	.2580888	02-	.2613737	03
.4428600	05	.1703242	03	.2249149	02-	.2638252	03
.4528600	05	.1500085	03	.1962371	02-	.2663134	03
.4628600	05	.1322774	03	.1714400	02-	.2688015	03
.4728600	05	.1167819	03	.1499686	02-	.2712697	03
.4828600	05	.1032229	03	.1313517	02-	.2737778	03
.4928600	05	.9132106	02	.1156246	02-	.2751560	03
.5028600	05	.8078318	02	.1030332	02-	.2731504	03
.5128600	05	.7139951	02	.9173859	03-	.2711448	03
.5228600	05	.6305047	02	.8161490	03-	.2691392	03
.5328600	05	.5562814	02	.7254777	03-	.2671336	03
.5428600	05	.4902878	02	.6455825	03-	.2645803	03
.5528600	05	.4315882	02	.5743182	03-	.2618033	03
.5628600	05	.3794151	02	.5103039	03-	.2590063	03
.5728600	05	.3330998	02	.4528661	03-	.2562493	03
.5828600	05	.2919949	02	.4021824	03-	.2529359	03
.5928600	05	.2554564	02	.3580484	03-	.2485612	03
.6028600	05	.2229696	02	.3181136	03-	.2441865	03
.6128600	05	.1941446	02	.2820414	03-	.2398119	03
.6228600	05	.1686231	02	.2495171	03-	.2354372	03
.6328600	05	.1460886	02	.2200091	03-	.2313310	03
.6428600	05	.1263302	02	.1919326	03-	.2293067	03
.6528600	05	.1091084	02	.1672441	03-	.2272824	03
.6628600	05	.9411522	01	.1455585	03-	.2252531	03
.6728600	05	.8107781	01	.1265319	03-	.2232338	03
.6828600	05	.6975484	01	.1098572	03-	.2212095	03
.6928600	05	.5992222	01	.9547134	04-	.2186610	03
.7028600	05	.5138047	01	.8295381	04-	.2157844	03
.7128600	05	.4396759	01	.7194509	04-	.2129069	03
.7228600	05	.3754634	01	.6227959	04-	.2100293	03
.7328600	05	.3199472	01	.5380811	04-	.2071518	03
.7428600	05	.2718522	01	.4676853	04-	.2023057	03
.7528600	05	.2301037	01	.4056556	04-	.1976169	03
.7628600	05	.1939649	01	.3506196	04-	.1907282	03
.7728600	05	.1627946	01	.3019336	04-	.1878394	03
.7828600	05	.1360111	01	.2589994	04-	.1809506	03
.7928600	05	.1135026	01	.2133584	04-	.1853336	03
.8028600	05	.9495193	00	.1760451	04-	.1879048	03
.8128600	05	.7963040	00	.1456454	04-	.1904761	03
.8228600	05	.6694253	00	.1208083	04-	.1930473	03
.8328600	05	.5629857	00	.1025636	04-	.1912328	03
.8428600	05	.4704215	00	.9143498	05-	.1792392	03
.8528600	05	.3882380	00	.8087262	05-	.1672456	03
.8628600	05	.3158847	00	.7088422	05-	.1552520	03
.8728600	05	.2536854	00	.5975985	05-	.1478918	03
.8828600	05	.2029426	00	.4719060	05-	.1498221	03
.8928600	05	.1628251	00	.3738041	05-	.1517524	03
.9028600	05	.1310111	00	.2969896	05-	.1536827	03

FIGURE 17

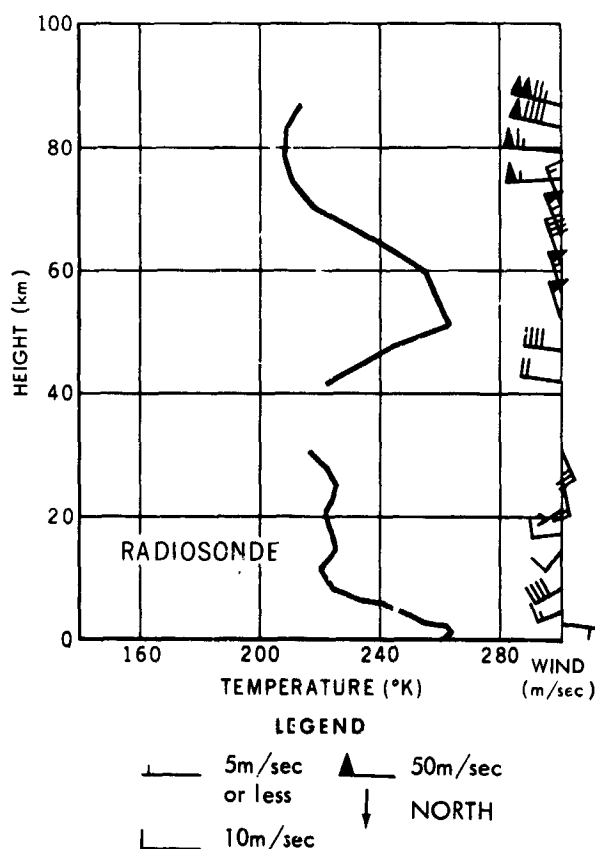
WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees Kelvin				
273.83029	53.767238	245.67549	40618.347	.2889260	05	.1300000	04	.2061350	01-	.2197100	03
267.66479	83.497436	259.56347	46206.437	.2989260	05	.1115347	04	.1750977	01-	.2219244	03
265.29300	94.068039	256.97729	50987.546	.3089260	05	.9584243	03	.1489701	01-	.2241388	03
263.99978	99.636138	255.52859	55623.097	.3189260	05	.8248469	03	.1269536	01-	.2263532	03
276.75988	85.044586	241.85888	61199.898	.3289260	05	.7109576	03	.1083645	01-	.2285676	03
295.78497	80.256599	230.72470	67064.695	.3389260	05	.6136999	03	.9264295	02-	.2307820	03
216.34069	36.389019	216.65219	72110.546	.3489260	05	.5305159	03	.7932451	02-	.2329964	03
203.29139	115.02049	188.27578	77416.687	.3589260	05	.4592600	03	.6802359	02-	.2352108	03
329.80078	196.25229	239.28830	82479.836	.3689260	05	.3981302	03	.5841933	02-	.2374252	03
46.157417	45.832988	179.11309	86794.796	.3789260	05	.3456103	03	.5024425	02-	.2396396	03
346.41769	389.29647	271.31540	90863.945	.3889260	05	.3004224	03	.4327502	02-	.2418540	03



.2889260	05	.1300000	04	.2061350	01-	.2197100	03
.2989260	05	.1115347	04	.175097	01-	.2219244	03
.3089260	05	.9584243	03	.1489701	01-	.2241388	03
.3189260	05	.8248469	03	.1269536	01-	.2263532	03
.3289260	05	.7109576	03	.1083645	01-	.2285676	03
.3389260	05	.6136999	03	.9264295	02-	.2307820	03
.3489260	05	.5305159	03	.7932451	02-	.2329964	03
.3589260	05	.4592600	03	.6802359	02-	.2352108	03
.3689260	05	.3981302	03	.5841933	02-	.2374252	03
.3789260	05	.3456103	03	.5024425	02-	.2396396	03
.3889260	05	.3004224	03	.4327502	02-	.2418540	03
.3989260	05	.2614078	03	.3732486	02-	.2440684	03
.4089260	05	.2278092	03	.3222819	02-	.2463571	03
.4189260	05	.1988968	03	.2784593	02-	.2488423	03
.4289260	05	.1738276	03	.2409554	02-	.2513276	03
.4389260	05	.1521260	03	.2088084	02-	.2538129	03
.4489260	05	.1333123	03	.1812104	02-	.2562982	03
.4589260	05	.1169791	03	.1574816	02-	.2587835	03
.4689260	05	.1027264	03	.1380759	02-	.2591923	03
.4789260	05	.9021051	02	.1215068	02-	.2586514	03
.4889260	05	.7920117	02	.1069016	02-	.2581105	03
.4989260	05	.6951930	02	.9403056	03	.2575695	03
.5089260	05	.6100673	02	.8269028	03	.2570286	03
.5189260	05	.5352674	02	.7264614	03	.2566944	03
.5289260	05	.4695803	02	.6380881	03	.2563819	03
.5389260	05	.4119050	02	.5603992	03	.2560694	03
.5489260	05	.3612704	02	.4921112	03	.2557569	03
.5589260	05	.3167762	02	.4330075	03	.2548680	03
.5689260	05	.2775331	02	.3830493	03	.254168	03
.5789260	05	.2428479	02	.3384638	03	.2499656	03
.5889260	05	.2122270	02	.2987159	03	.2475145	03
.5989260	05	.1852264	02	.2633194	03	.2450633	03
.6089260	05	.1614159	02	.2318330	03	.2426121	03
.6189260	05	.1405455	02	.2035546	03	.2405438	03
.6289260	05	.1222150	02	.1784144	03	.2386453	03
.6389260	05	.1061614	02	.1562215	03	.2367468	03
.6489260	05	.9211609	01	.1366489	03	.2348484	03
.6589260	05	.7984050	01	.1154040	03	.2329499	03
.6689260	05	.6912293	01	.1042249	03	.2310514	03
.6789260	05	.5976166	01	.9114561	04	.2284157	03
.6889260	05	.5158077	01	.7964442	04	.2256268	03
.6989260	05	.4444041	01	.6947799	04	.2228379	03
.7089260	05	.3821849	01	.6050795	04	.2200490	03
.7189260	05	.3280607	01	.5260566	04	.2172600	03
.7289260	05	.2808563	01	.4605162	04	.2144699	03
.7389260	05	.2395546	01	.4029362	04	.2071221	03
.7489260	05	.2034884	01	.3513436	04	.2017742	03
.7589260	05	.1721051	01	.3052474	04	.1964264	03
.7689260	05	.1448981	01	.2641856	04	.1910785	03
.7789260	05	.1218167	01	.2198107	04	.1933070	03
.7889260	05	.1029564	01	.1765646	04	.2031459	03
.7989260	05	.8773210	00	.1433463	04	.2132211	03
.8089260	05	.7531704	00	.1175086	04	.2232964	03
.8189260	05	.6505872	00	.9718126	05	.2333717	03
.8289260	05	.5645094	00	.8421372	05	.2335321	03
.8389260	05	.4873984	00	.7732802	05	.2195864	03
.8489260	05	.4168033	00	.7061230	05	.2056407	03
.8589260	05	.3525550	00	.6407291	05	.1916949	03
.8689260	05	.2949986	00	.5667752	05	.1813290	03
.8789260	05	.2482211	00	.4239283	05	.2039870	03
.8889260	05	.2127082	00	.3265587	05	.2266468	03
.8989260	05	.1849879	00	.2585053	05	.2493057	03
.9089260	05	.1628287	00	.2090840	05	.2713116	03

FIGURE 18

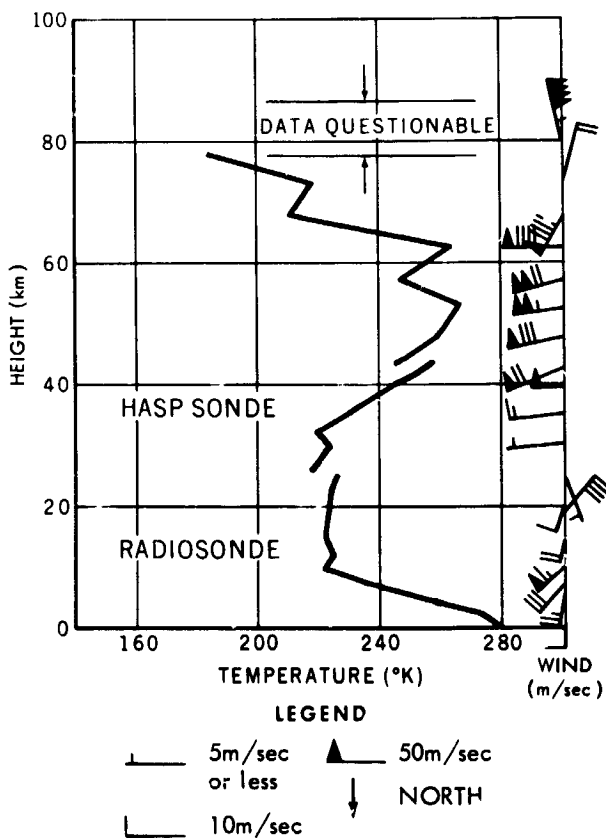
WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees Kelvin
290.36468	20.766498	221.75138	41866.148	.3064000 05	.1000000 04	.1603978 01-	.2172000 03
288.81070	40.355209	242.00459	47108.296	.3164000 05	.8557217 03	.1370001 01-	.2176054 03
330.36560	48.105667	262.88278	51592.347	.3264000 05	.7325066 03	.1170551 01-	.2180108 03
334.51089	54.224269	258.55648	55932.187	.3364000 05	.6272446 03	.1000484 01-	.2184163 03
339.43869	39.816047	254.45228	60114.148	.3464000 05	.5372898 03	.8554141 02-	.2188217 03
328.10589	55.908588	238.15919	65105.547	.3564000 05	.4503895 03	.7316263 02-	.2192271 03
333.09668	13.910879	218.05169	70373.898	.3664000 05	.3946274 03	.6259631 02-	.2196325 03
263.82138	54.381446	210.59829	74878.593	.3764000 05	.3383709 03	.5357393 02-	.2200380 03
279.66458	64.996559	208.21719	79168.898	.3864000 05	.2902301 03	.4586734 02-	.2204434 03
294.59338	91.706909	208.93289	83243.593	.3964000 05	.2490202 03	.3928237 02-	.2208488 03
299.68167	125.85369	212.67819	86702.132	.4064000 05	.2137321 03	.3365396 02-	.2212543 03



.4164000 05	.1835045 03	.2884153 02-	.2216597 03
.4264000 05	.1577466 03	.2445319 02-	.2247412 03
.4364000 05	.1359248 03	.2071437 02-	.2286047 03
.4464000 05	.1174197 03	.1750688 02-	.2324682 03
.4564000 05	.1016835 03	.1498048 02-	.2363318 03
.4664000 05	.8826577 02	.1280224 02-	.2401953 03
.4764000 05	.7680436 02	.1094461 02-	.2441482 03
.4864000 05	.6700262 02	.9369423 03-	.2481363 03
.4964000 05	.5860218 02	.8044393 03-	.2521324 03
.5064000 05	.5138202 02	.6926204 03-	.2561485 03
.5164000 05	.4515614 02	.5985374 03-	.2601835 03
.5264000 05	.3971945 02	.5284795 03-	.2642384 03
.5364000 05	.3492161 02	.4664186 03-	.2683145 03
.5464000 05	.3068940 02	.4114640 03-	.2724146 03
.5564000 05	.2695780 02	.3628259 03-	.2765317 03
.5664000 05	.2366912 02	.3197815 03-	.2806618 03
.5764000 05	.2077214 02	.2817140 03-	.2848040 03
.5864000 05	.1822136 02	.2480678 03-	.2889581 03
.5964000 05	.1597643 02	.2183424 03-	.2931242 03
.6064000 05	.1399718 02	.1929444 03-	.2973023 03
.6164000 05	.1224605 02	.1710147 03-	.3014924 03
.6264000 05	.1069561 02	.1513431 03-	.3056945 03
.6364000 05	.9324993 01	.1337218 03-	.3099086 03
.6464000 05	.8115300 01	.1179595 03-	.3141347 03
.6564000 05	.7048782 01	.1040017 03-	.3183728 03
.6664000 05	.6109125 01	.9161846 04-	.3226229 03
.6764000 05	.5282436 01	.8054392 04-	.3268850 03
.6864000 05	.4556648 01	.7065775 04-	.3311591 03
.6964000 05	.3920623 01	.6184899 04-	.3354442 03
.7064000 05	.3365729 01	.5388349 04-	.3397403 03
.7164000 05	.2884390 01	.4653130 04-	.3440474 03
.7264000 05	.2469075 01	.4013393 04-	.3483655 03
.7364000 05	.2111117 01	.3458677 04-	.3526946 03
.7464000 05	.1802536 01	.2976941 04-	.3570347 03
.7564000 05	.1538304 01	.2550018 04-	.3613858 03
.7664000 05	.1312049 01	.2180589 04-	.3657479 03
.7764000 05	.1118589 01	.1864000 04-	.3701200 03
.7864000 05	.9532087 00	.1592791 04-	.3745021 03
.7964000 05	.8122319 00	.1358467 04-	.3788942 03
.8064000 05	.6920720 00	.1156523 04-	.3832963 03
.8164000 05	.5877942 00	.9847768 05-	.3877084 03
.8264000 05	.5027252 00	.8386632 05-	.3921305 03
.8364000 05	.4286470 00	.7132757 05-	.3965626 03
.8464000 05	.3657261 00	.6054450 05-	.4010047 03
.8564000 05	.3123110 00	.5143724 05-	.4054568 03
.8664000 05	.2669252 00	.4373834 05-	.4100000 03

FIGURE 19

WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meters	TEMPERATURE degrees kelvin
248.97470	71.465309	244.66299	43146.937	.2604670 05	.2000000 04	.3237469 01-	.2152200 03
259.58279	82.276358	259.48867	48478.148	.2704670 05	.1709929 04	.2745953 01-	.2160418 03
257.88967	103.08649	265.55178	53019.250	.2804670 05	.1463812 04	.2332207 01-	.2186635 03
250.40109	122.64289	247.70709	57394.699	.2904670 05	.1254708 04	.1983436 01-	.2203853 03
271.84619	91.785987	262.91629	62650.347	.3004670 05	.1076817 04	.1689032 01-	.2221071 03
212.86879	87.717720	210.57238	68162.734	.3104670 05	.9252830 03	.1440175 01-	.2238289 03
18.724060	20.152278	217.48979	72900.132	.3204670 05	.7760349 03	.1220550 01-	.2255507 03
348.64187	202.72158	184.07159	77840.836	.3304670 05	.6856570 03	.1051037 01-	.2272715 03
132.55758	339.43707	93.948997	82529.546	.3404670 05	.5912771 03	.8995484 02-	.2289943 03
272.95788	385.39788	327.72747	86532.296	.3504670 05	.5104778 03	.7708274 02-	.2307161 03



.2604670 05	.2000000 04	.3237469 01-	.2152200 03
.2704670 05	.1709929 04	.2745953 01-	.2160418 03
.2804670 05	.1463812 04	.2332207 01-	.2186635 03
.2904670 05	.1254708 04	.1983436 01-	.2203853 03
.3004670 05	.1076817 04	.1689032 01-	.2221071 03
.3104670 05	.9252830 03	.1440175 01-	.2238289 03
.3204670 05	.7760349 03	.1220550 01-	.2255507 03
.3304670 05	.6856570 03	.1051037 01-	.2272715 03
.3404670 05	.5912771 03	.8995484 02-	.2289943 03
.3504670 05	.5104778 03	.7708274 02-	.2307161 03
.3604670 05	.4412218 03	.6613148 02-	.2324379 03
.3704670 05	.3617895 03	.5680284 02-	.2341597 03
.3804670 05	.3307277 03	.4884667 02-	.2358814 03
.3904670 05	.2868072 03	.4205280 02-	.2376032 03
.4004670 05	.2489863 03	.3624479 02-	.2393250 03
.4104670 05	.2163815 03	.3127353 02-	.2410468 03
.4204670 05	.1882424 03	.2701365 02-	.2427686 03
.4304670 05	.1639311 03	.2335919 02-	.2444904 03
.4404670 05	.1429421 03	.2014796 02-	.2471651 03
.4504670 05	.1248333 03	.1739972 02-	.2499461 03
.4604670 05	.1091867 03	.1505138 02-	.2527270 03
.4704670 05	.9564529 02	.1304119 02-	.2555079 03
.4804670 05	.8390687 02	.1131749 02-	.2582888 03
.4904670 05	.7370027 02	.9865980 03-	.2602477 03
.5004670 05	.6479087 02	.8629041 03-	.2615829 03
.5104670 05	.5699818 02	.7552638 03-	.2629181 03
.5204670 05	.5017729 02	.6615231 03-	.2642532 03
.5304670 05	.4420119 02	.5801310 03-	.2654998 03
.5404670 05	.3891117 02	.5186698 03-	.2613614 03
.5504670 05	.3418705 02	.4629228 03-	.2572831 03
.5604670 05	.2997563 02	.4124342 03-	.2532047 03
.5704670 05	.2622805 02	.3667791 03-	.2491263 03
.5804670 05	.2292788 02	.3200282 03-	.2459539 03
.5904670 05	.2006179 02	.2768137 03-	.2524877 03
.6004670 05	.1758143 02	.2398406 03-	.2553816 03
.6104670 05	.1543129 02	.2081503 03-	.2582755 03
.6204670 05	.1356434 02	.1809400 03-	.2611694 03
.6304670 05	.1192634 02	.1603281 03-	.2591526 03
.6404670 05	.1045592 02	.1459073 03-	.2496570 03
.6504670 05	.9100528 01	.1323047 03-	.2401613 03
.6604669 05	.7912255 01	.1195027 03-	.2306656 03
.6704669 05	.6823545 01	.1074836 03-	.2211699 03
.6804669 05	.5846780 01	.9622920 04-	.2116742 03
.6904669 05	.4993262 01	.8210832 04-	.2118631 03
.7004669 05	.4267132 01	.6968768 04-	.2133233 03
.7104669 05	.3650684 01	.5921498 04-	.2147834 03
.7204669 05	.3126745 01	.5037411 04-	.2162436 03
.7304669 05	.2679761 01	.4312238 04-	.2164984 03
.7404669 05	.2291380 01	.3806145 04-	.2097346 03
.7504669 05	.1949340 01	.3345894 04-	.2029707 03
.7604669 05	.1649379 01	.2928630 04-	.1962069 03
.7704669 05	.1387493 01	.2551587 04-	.1894430 03
.7804669 05	.1158476 01	.2240764 04-	.1801147 03
.7904669 05	.9526619 00	.2062806 04-	.1608935 03
.8004669 05	.7641867 00	.1870198 04-	.1416723 03
.8104669 05	.5936266 00	.1688918 04-	.1224512 03
.8204669 05	.4416710 00	.1490566 04-	.1032300 03
.8304669 05	.3293313 00	.9241337 05-	.1241528 03
.8404669 05	.2644035 00	.5045758 05-	.1825572 03
.8504669 05	.2257515 00	.3263930 05-	.2409617 03
.8604669 05	.1995179 00	.2321867 05-	.2993661 03

FIGURE 20

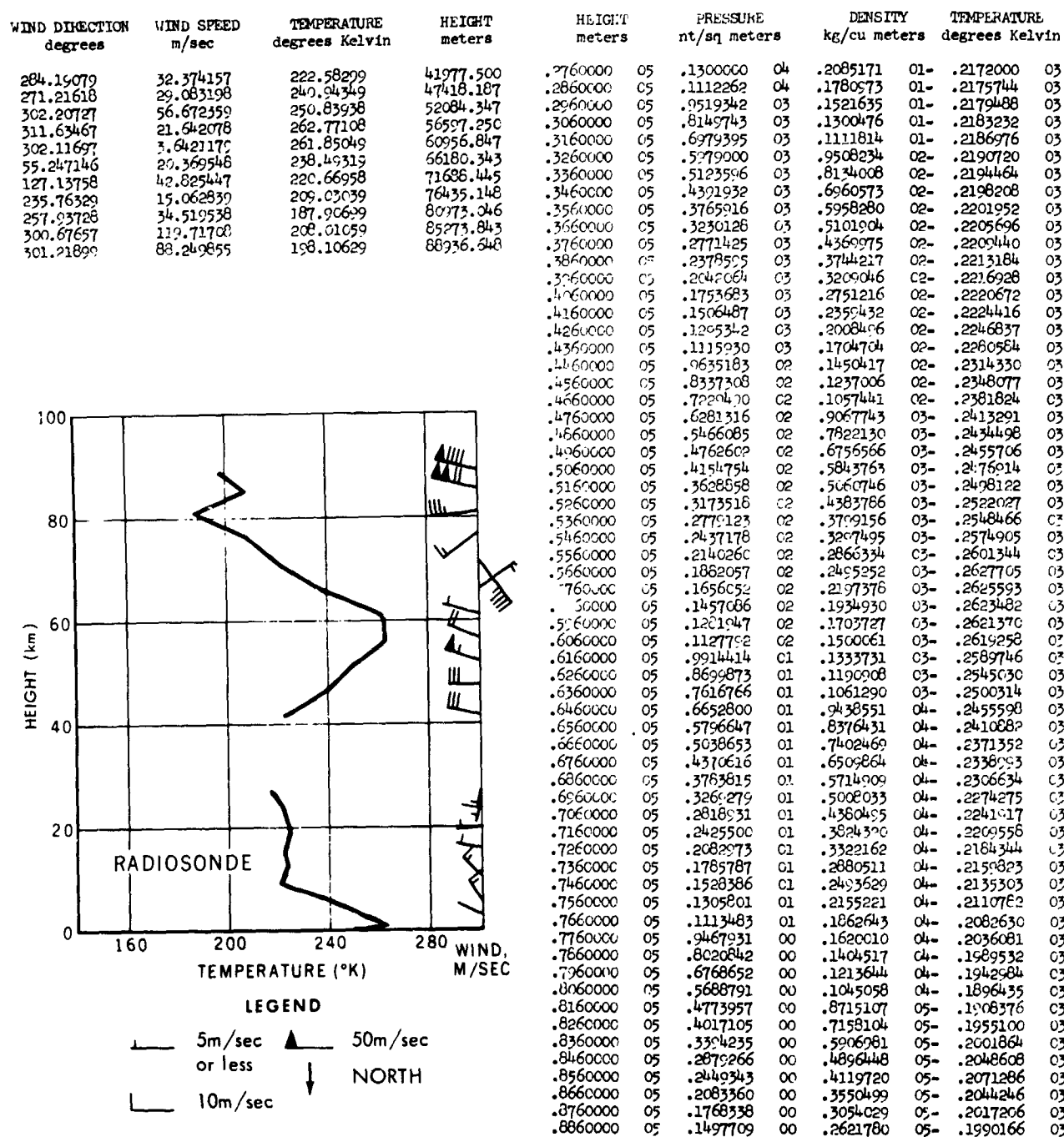


FIGURE 21

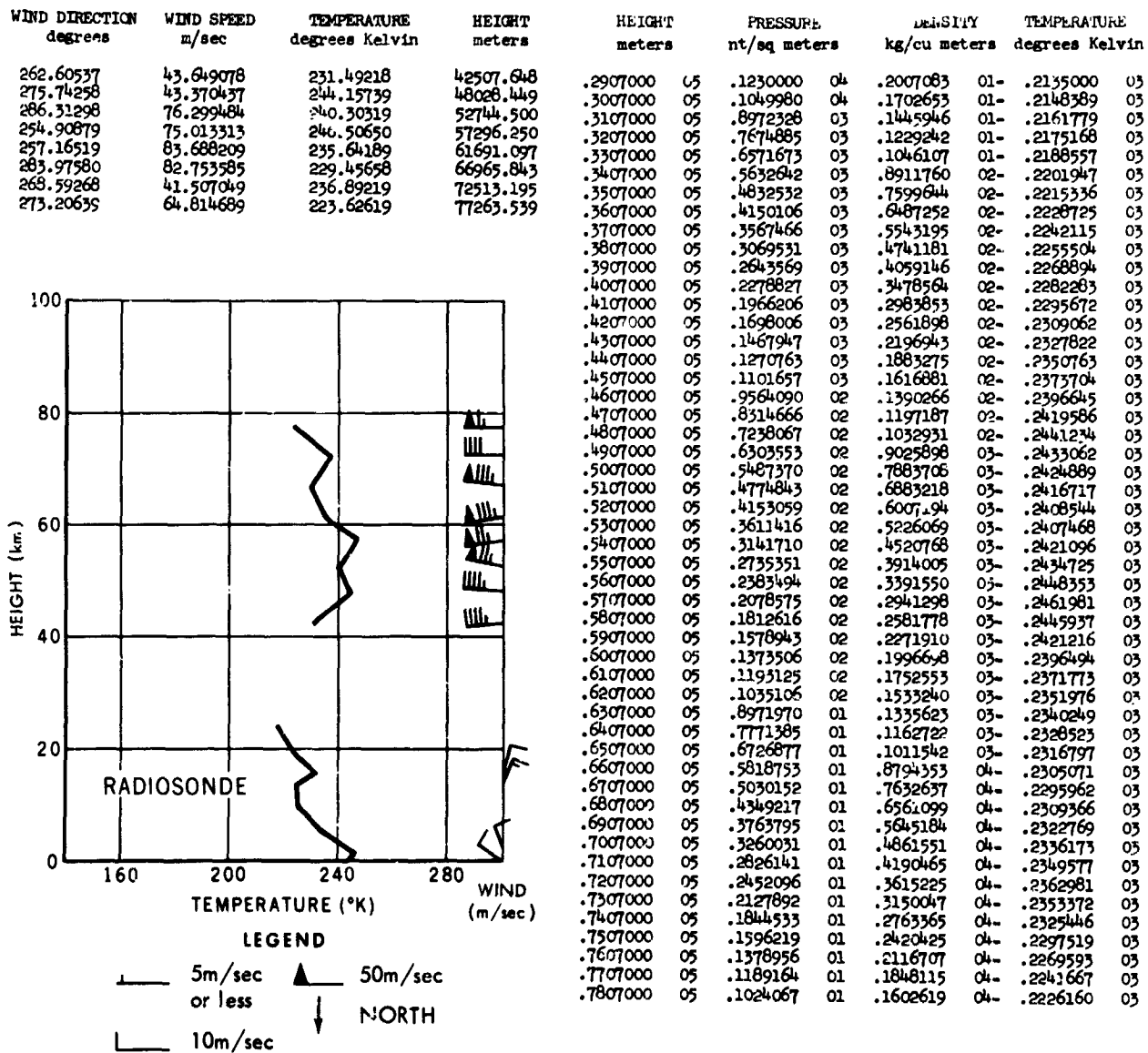


FIGURE 22

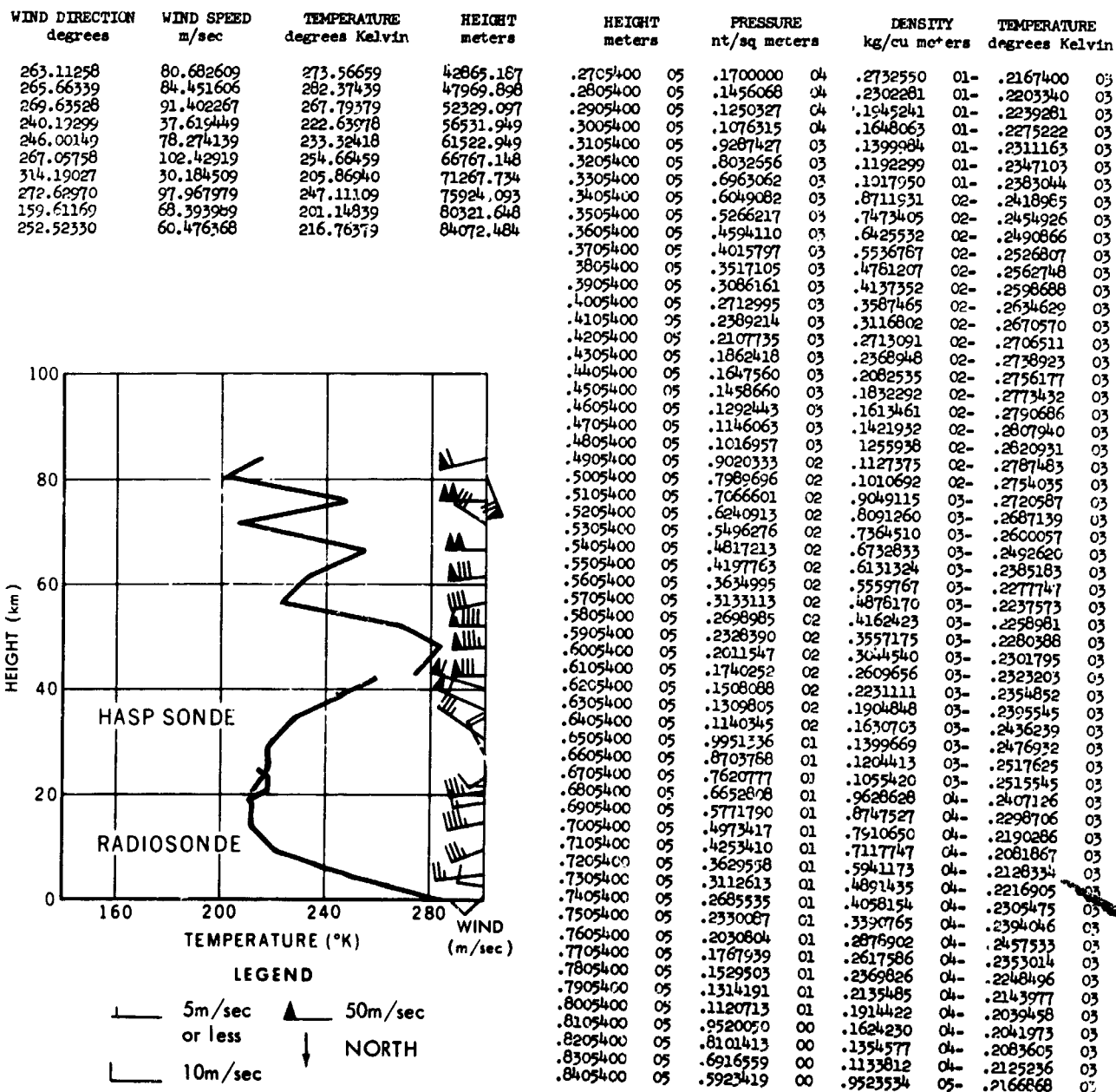


FIGURE 23

WIND DIRECTION degrees	WIND SPEED m/sec	TEMPERATURE degrees Kelvin	HEIGHT meters	HEIGHT meters	PRESSURE nt/sq meters	DENSITY kg/cu meter	TEMPERATURE degrees Kelvin
262.00698	43.808818	241.95039	42457.347	.3336000	05 .6200000	03 .9972211	02 .2166000
265.68618	54.826538	238.57479	48063.000	.3436000	05 .5308386	03 .8429634	02 .2193875
270.27807	148.54609	264.38278	57444.046	.3536000	05 .4554130	03 .7141151	02 .2221751
227.51558	100.64650	243.20858	67609.046	.3636000	05 .3914702	03 .6062425	02 .2240626
296.37479	103.47199	222.15139	73349.937	.3736000	05 .3371486	03 .5157281	02 .2277502
119.05209	51.014370	212.26008	78287.843	.3836000	05 .2909064	03 .4306118	02 .2305376
196.22400	58.051452	218.47559	83003.48	.3936000	05 .2514635	03 .3754665	02 .2333253
91.258529	172.84658	170.39270	87487.71	.4036000	05 .2177547	03 .3212065	02 .2361129

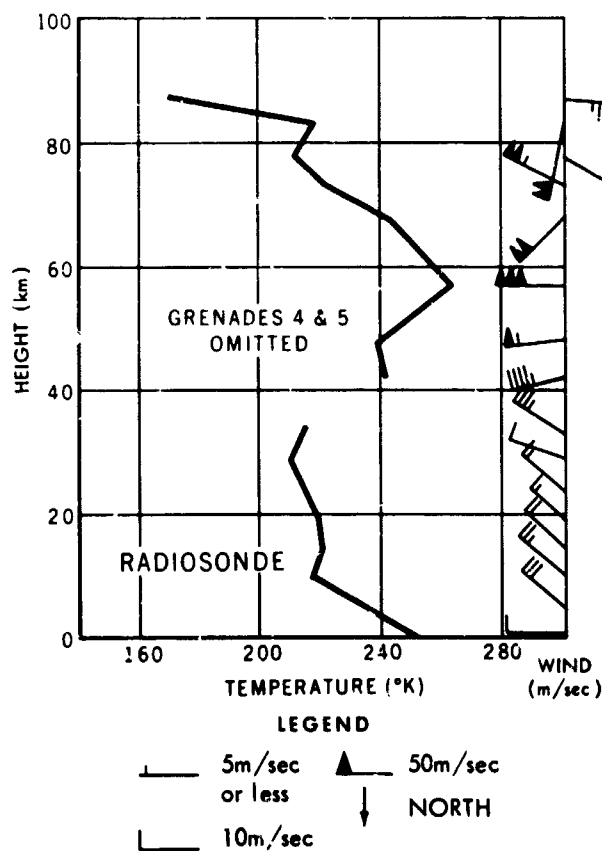


FIGURE 24

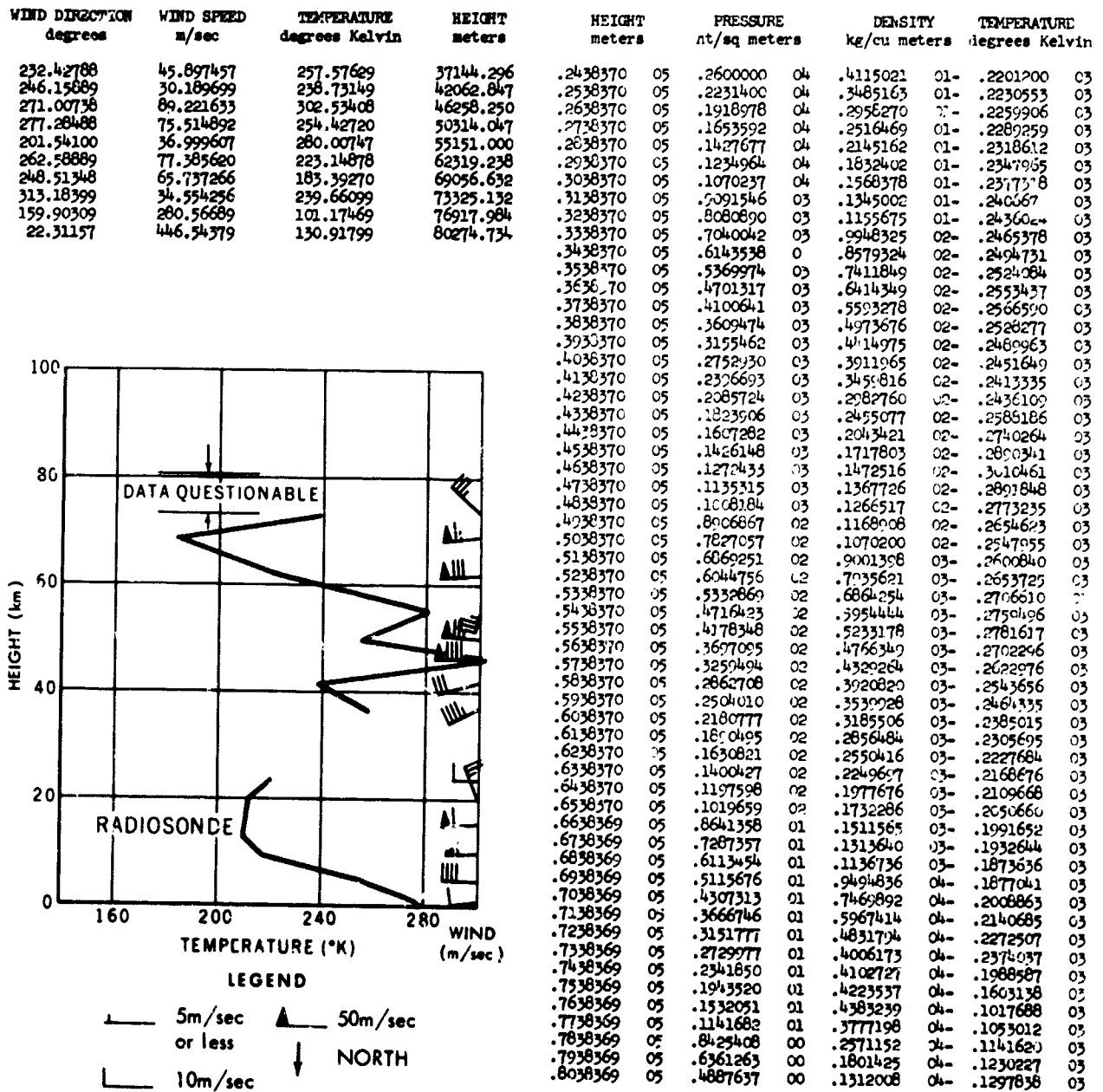


FIGURE 25

PLANETARY ATMOSPHERES

1135

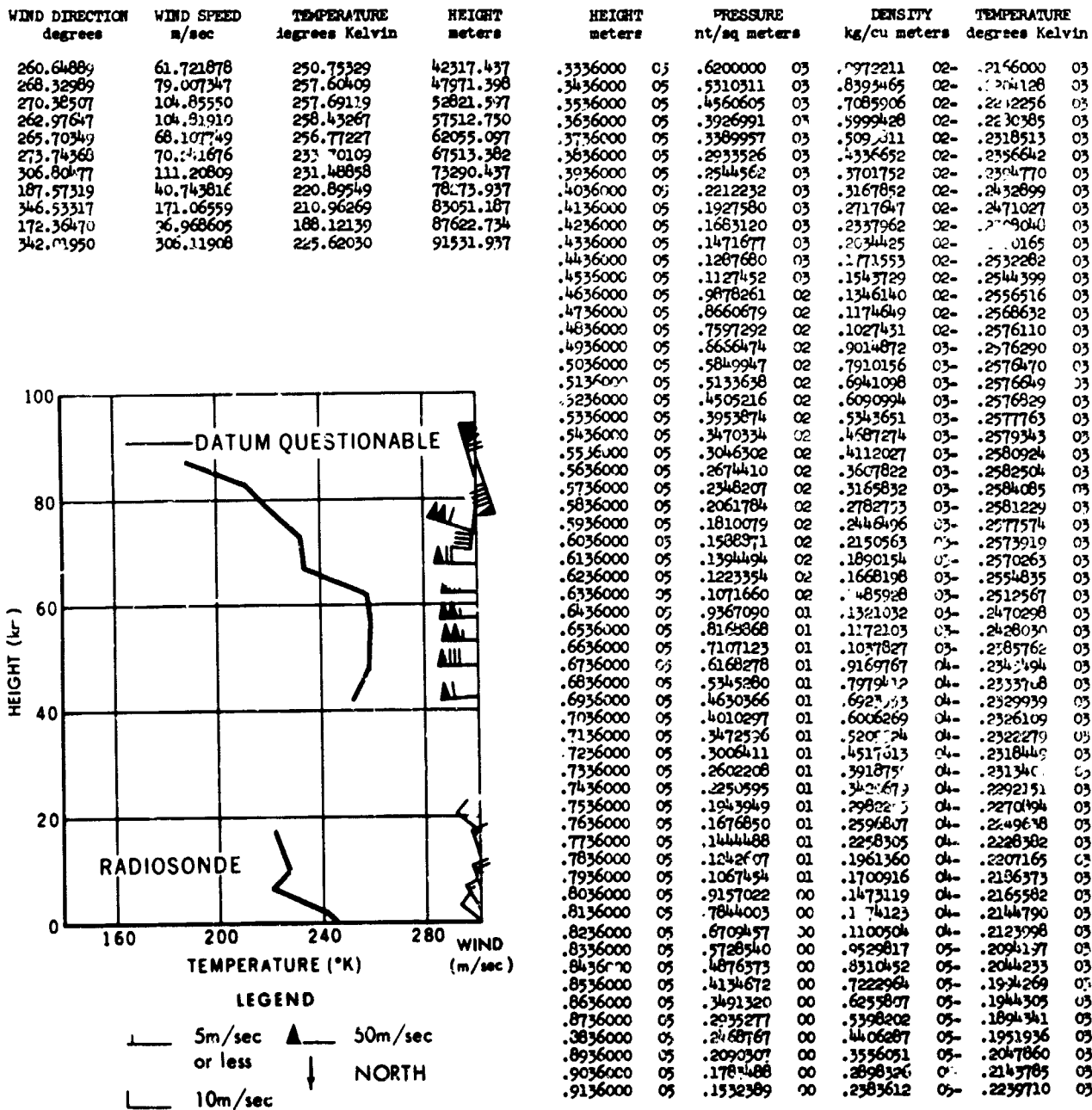


FIGURE 26

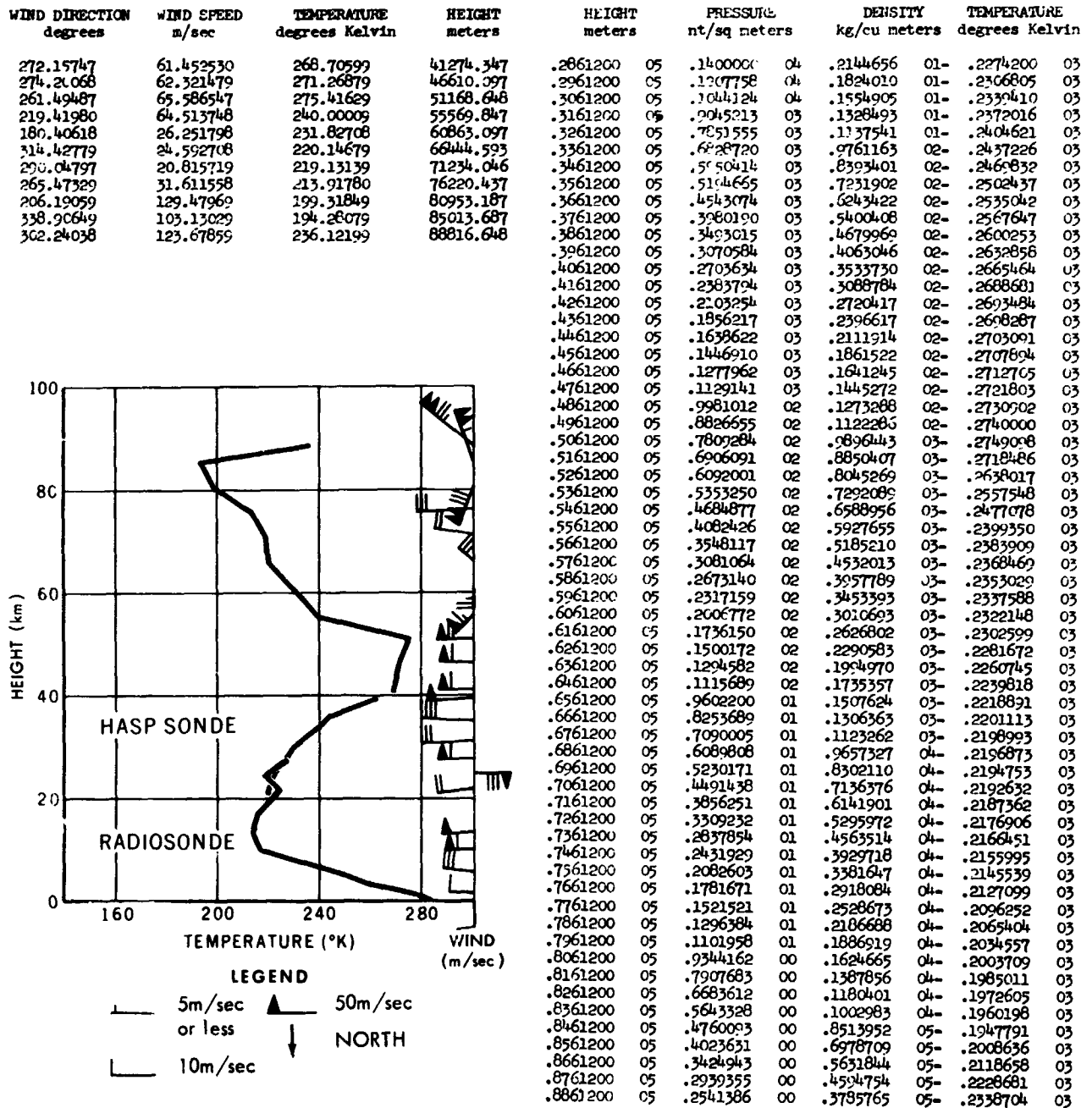


FIGURE 27

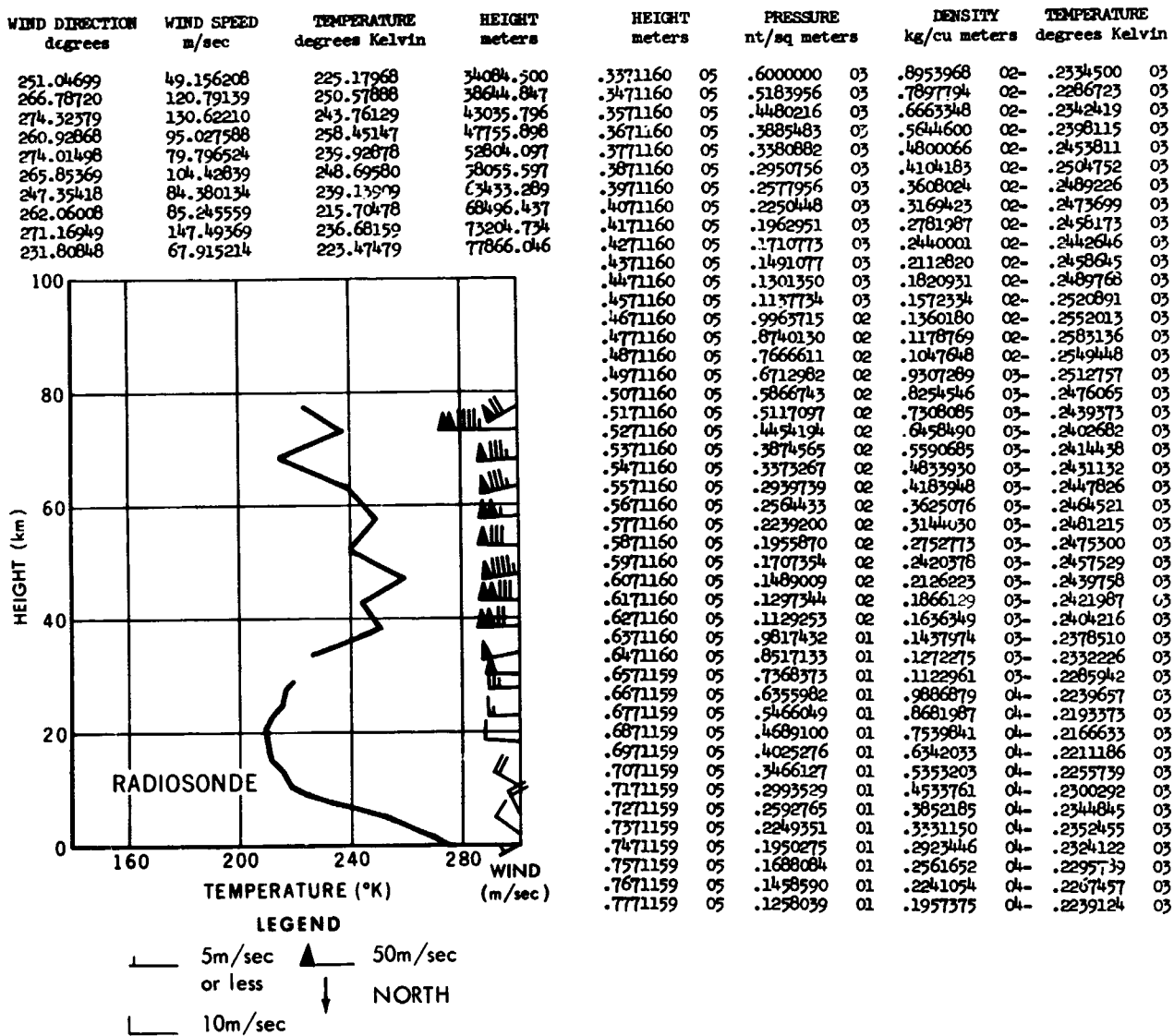


FIGURE 28

ERROR ANALYSIS 7/JUNE/62 NORTH COÖRD. - 200m

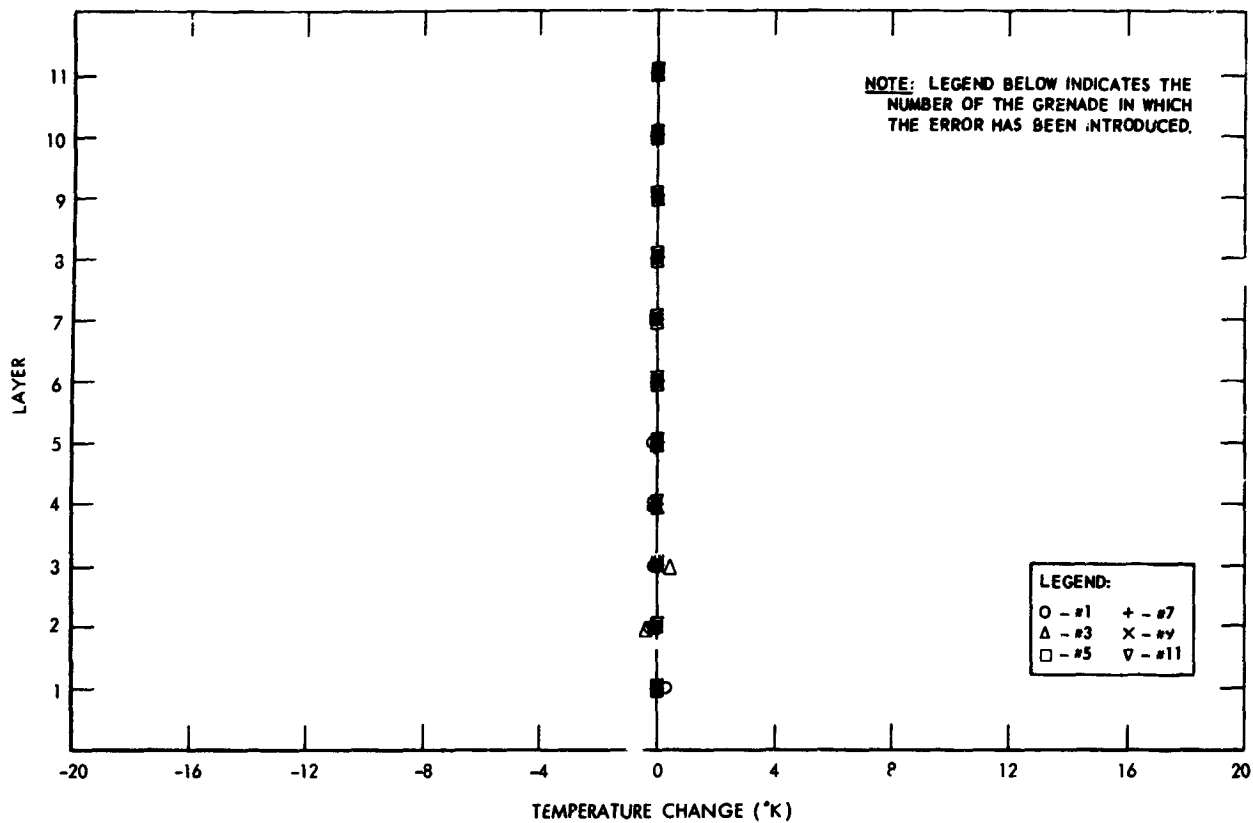


FIGURE 29

ERROR ANALYSIS 7/JUNE/62 NORTH COÖRD. - 200m

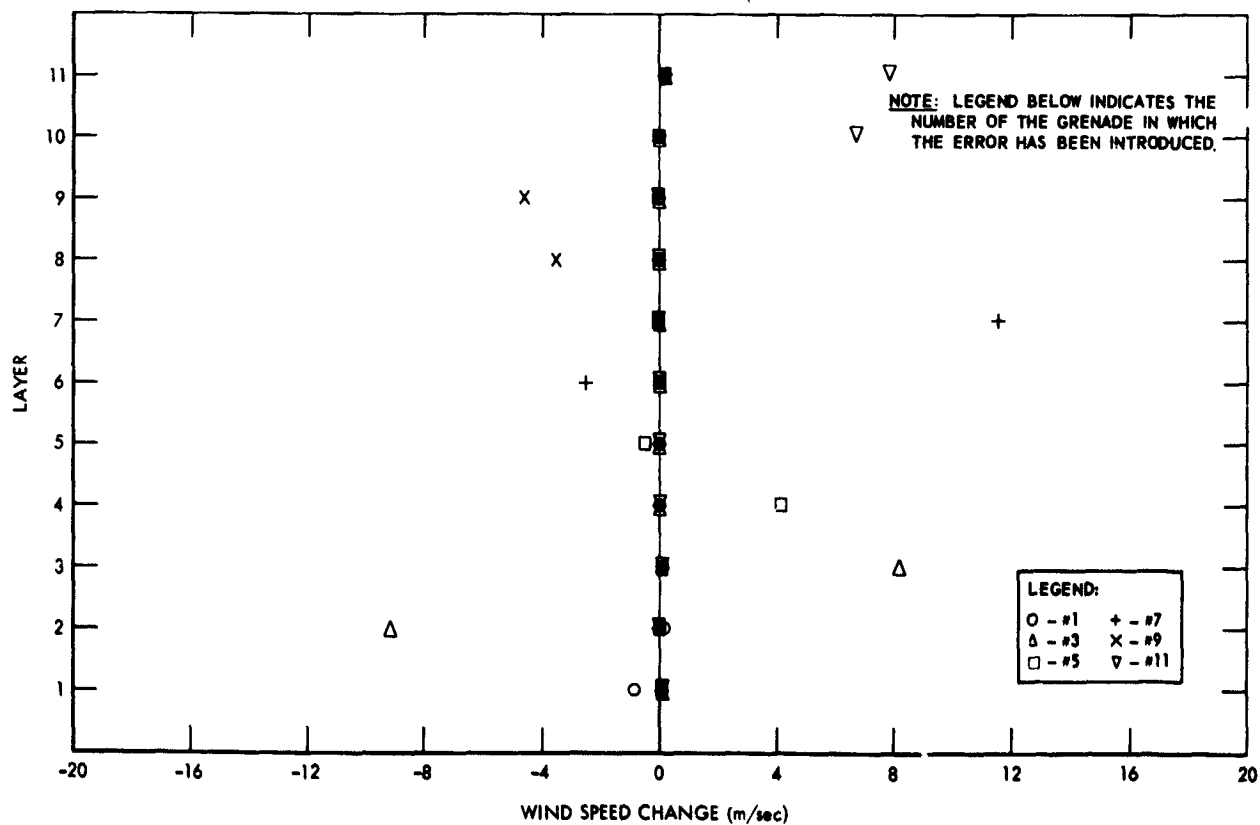


FIGURE 30

PLANETARY ATMOSPHERES

1139

ERROR ANALYSIS 7/JUNE/62 WEST COÖRD. - 200m

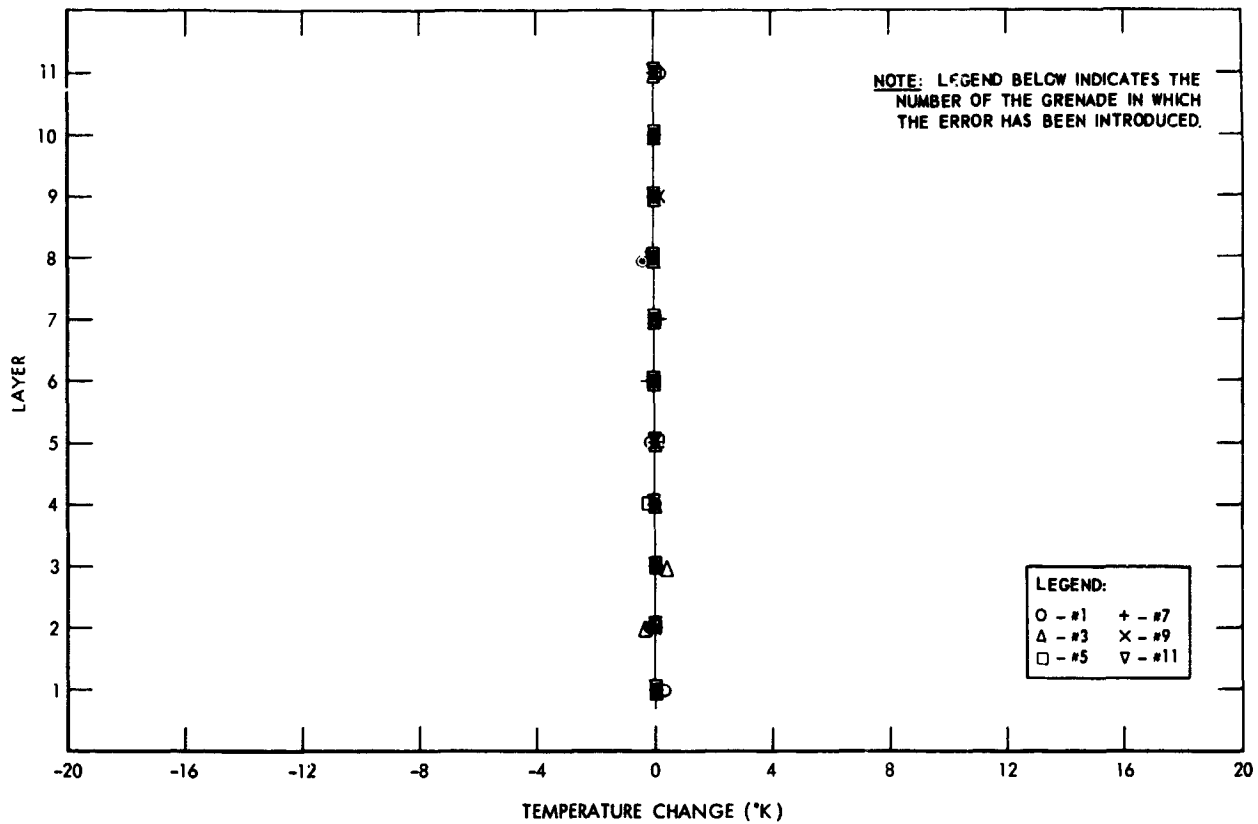


FIGURE 31

ERROR ANALYSIS 7/JUNE/62 WEST COORD. - 200m

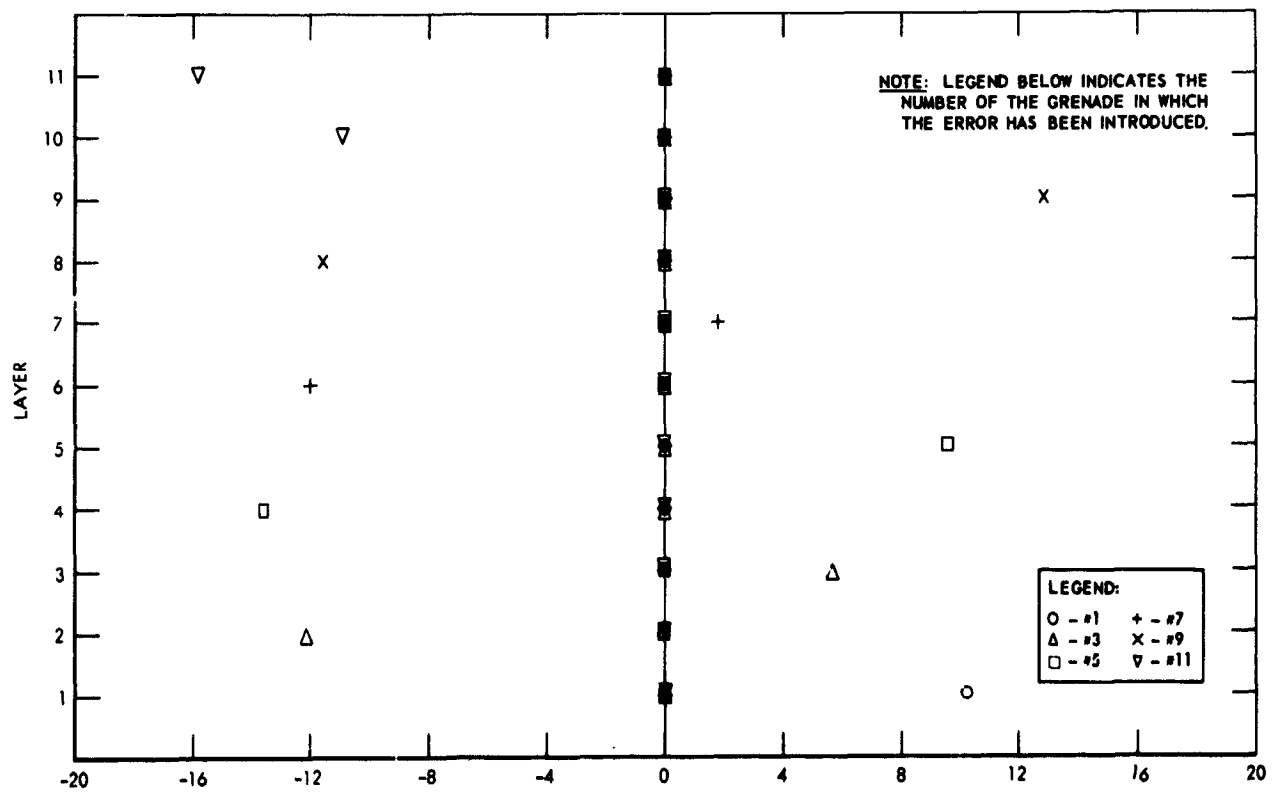


FIGURE 32

ERROR ANALYSIS 7/JUNE/62 UP COÖRD. - 50m

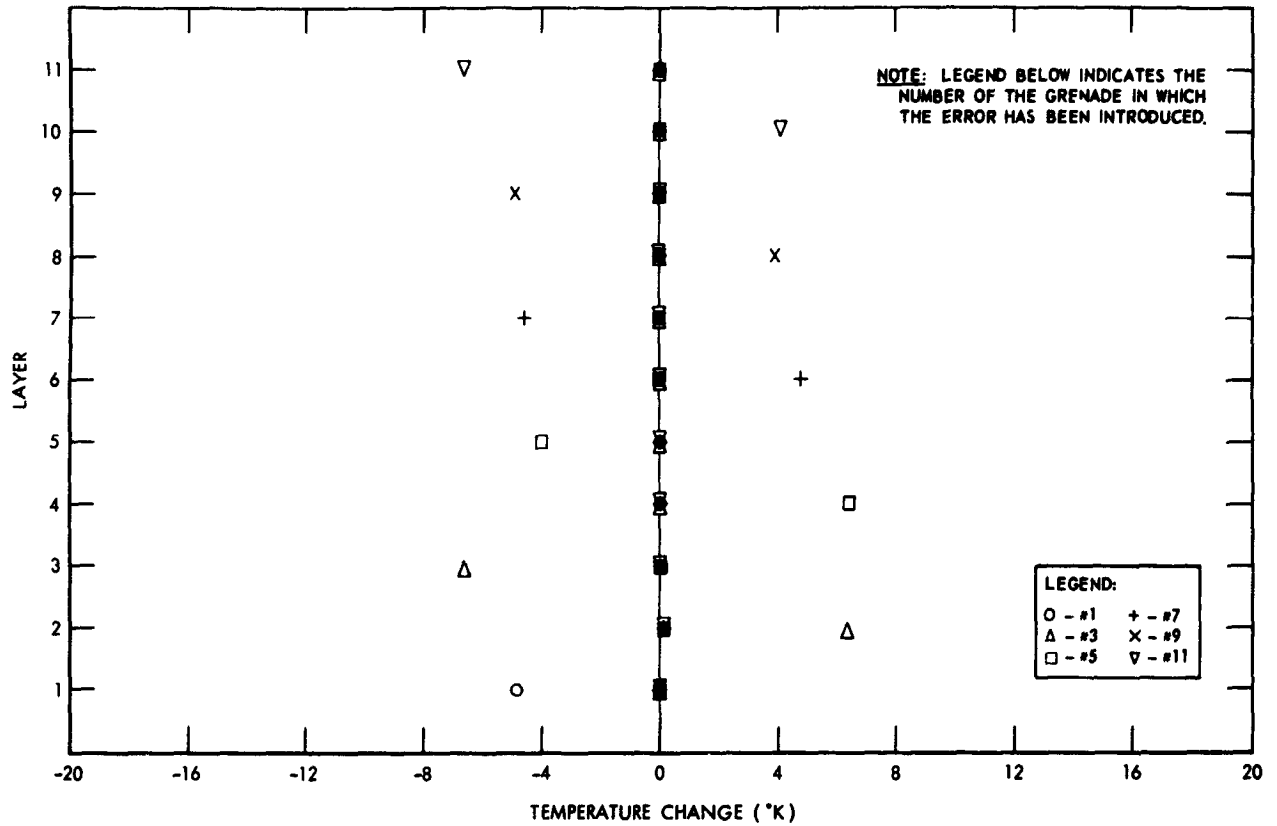


FIGURE 33

ERROR ANALYSIS 7/JUNE/62 UP COÖRD. - 50m

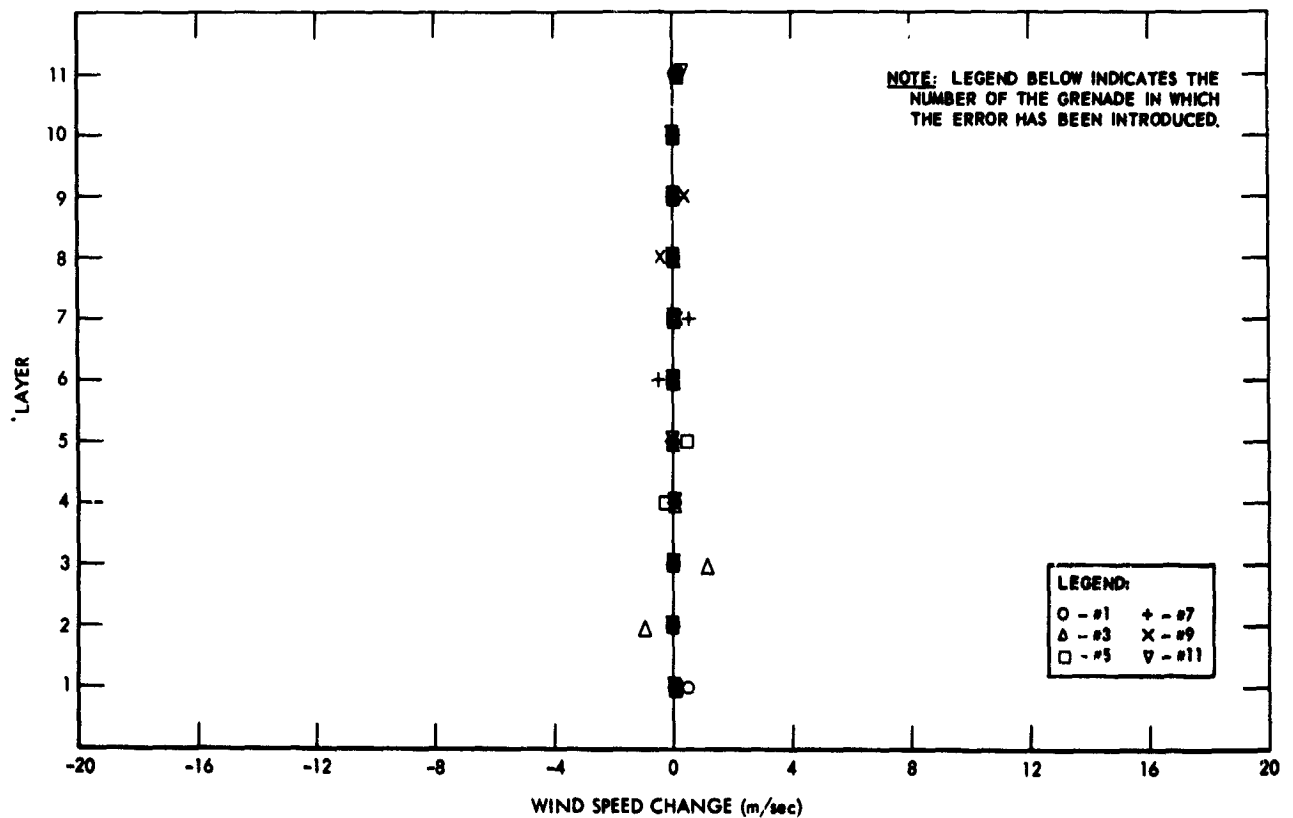


FIGURE 34

PLANETARY ATMOSPHERES

1141

ERROR ANALYSIS 7 JUNE '62 t + .3 sec

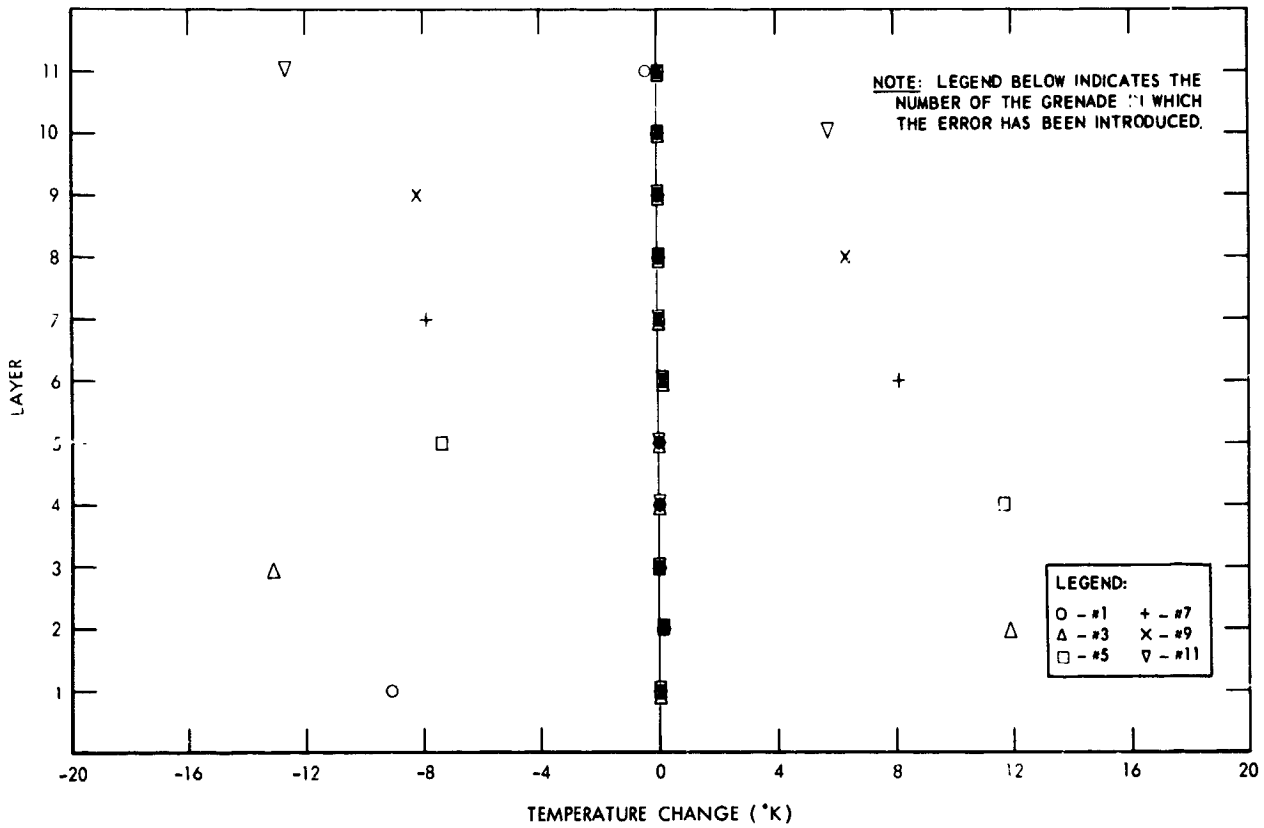


FIGURE 35

ERROR ANALYSIS 7/JUNE/62 t + .3 sec

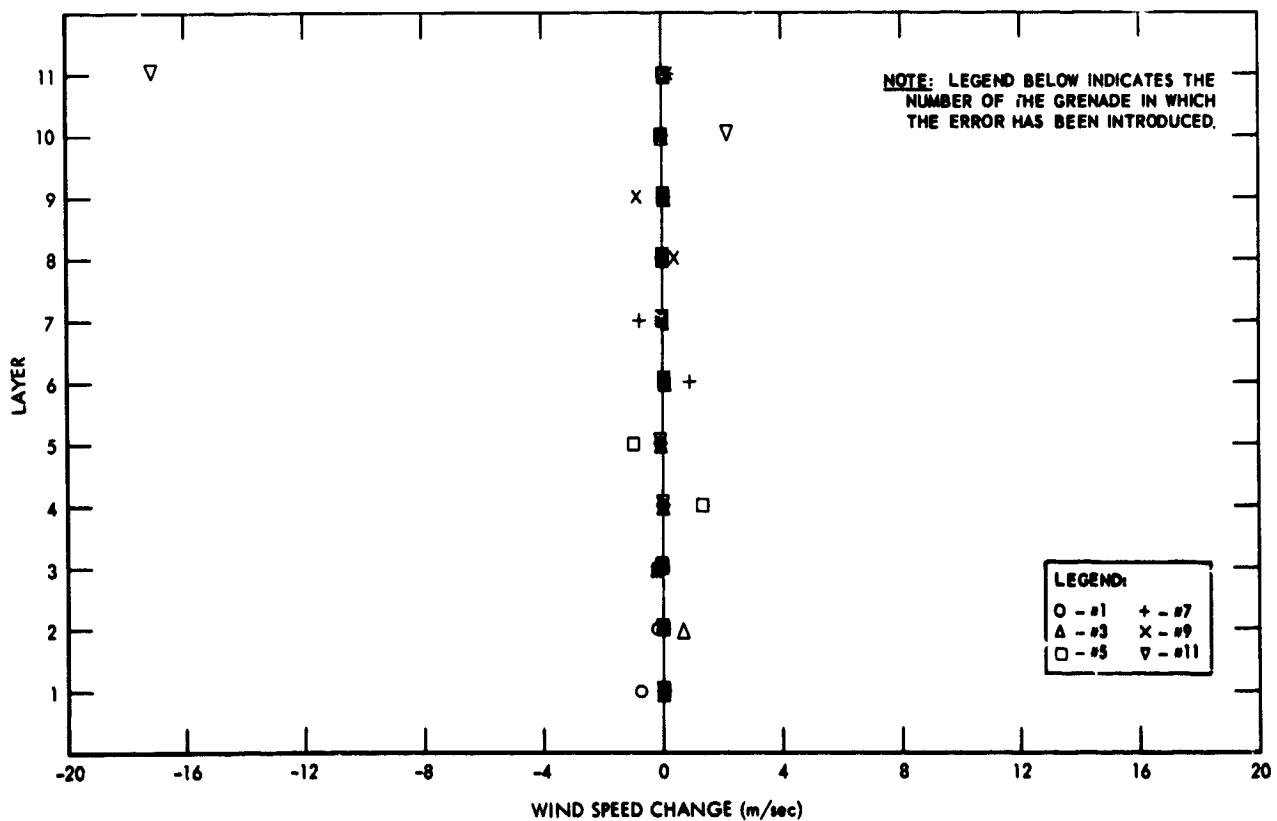


FIGURE 36

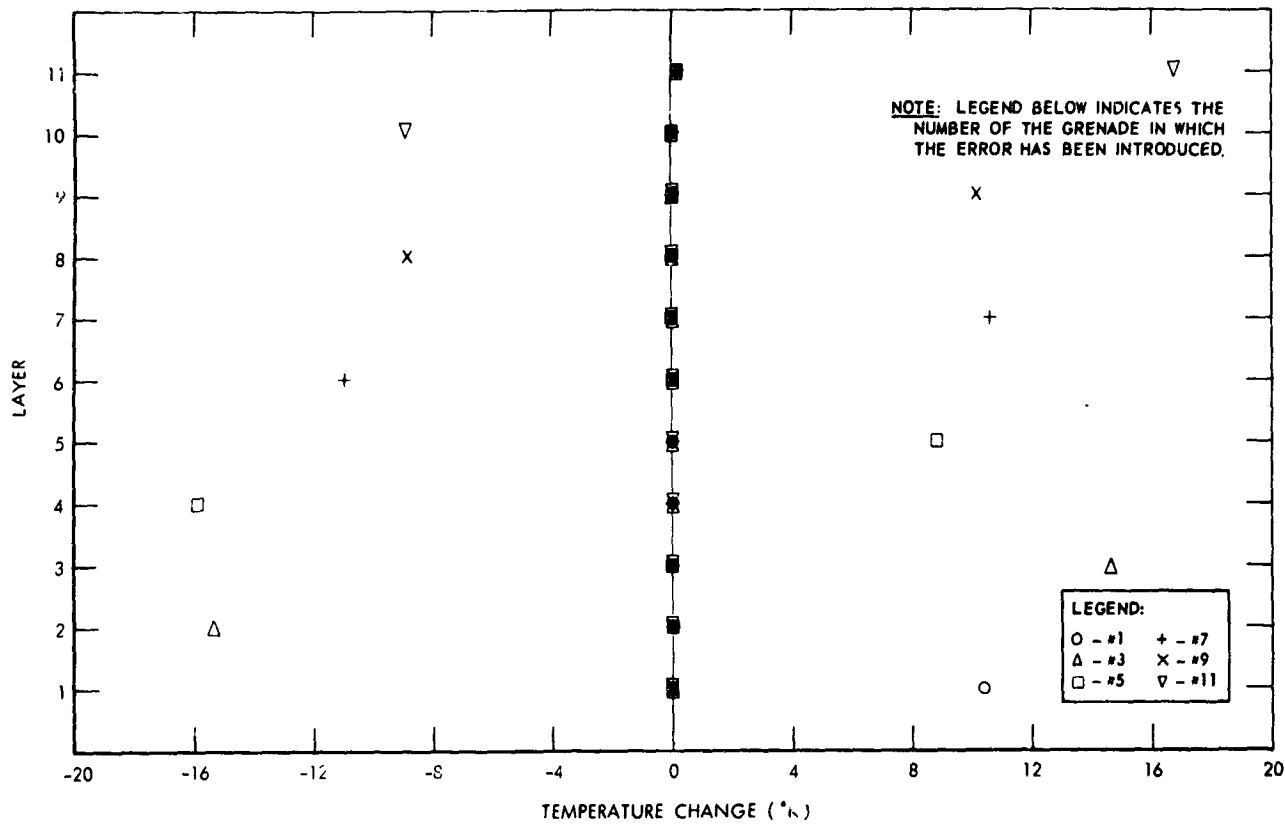
ERROR ANALYSIS 7 JUNE 62 $\Delta t = .02 \text{ sec}$ 

FIGURE 37

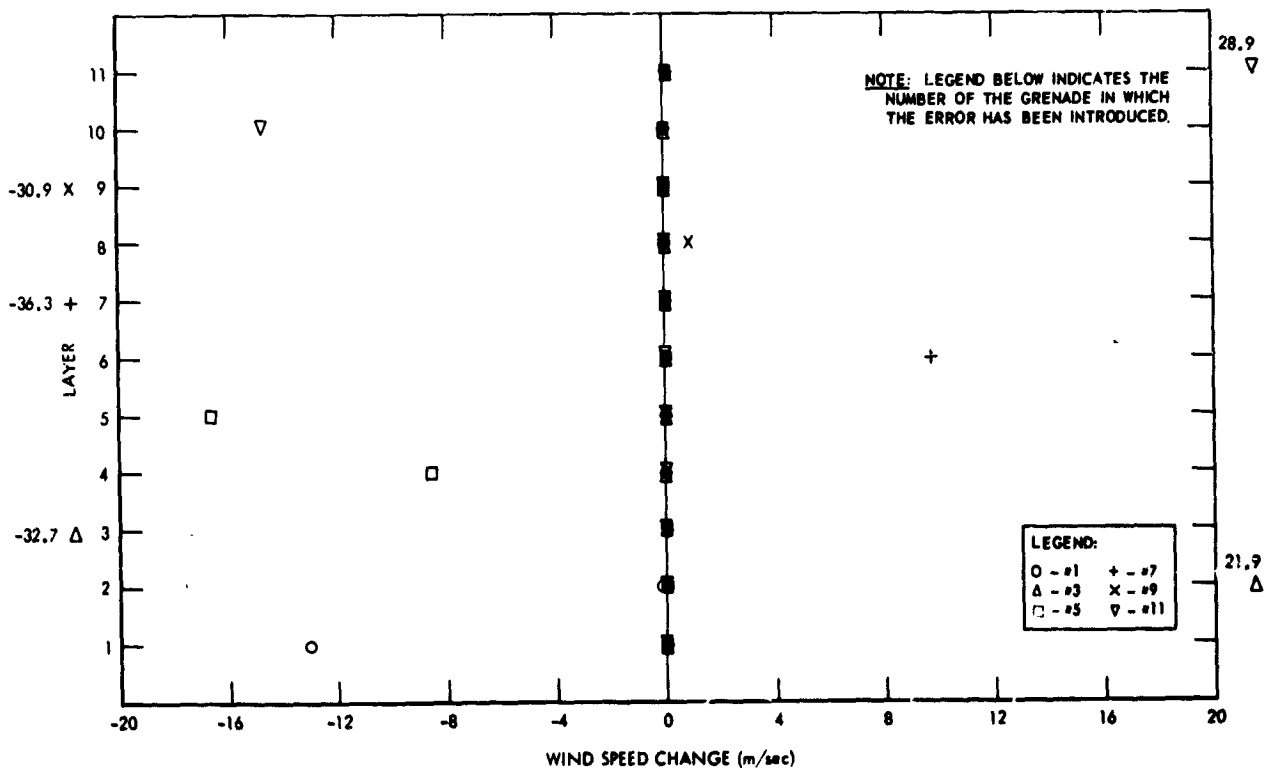
ERROR ANALYSIS 7/JUNE/62 $\Delta t = .02 \text{ sec}$ 

FIGURE 38

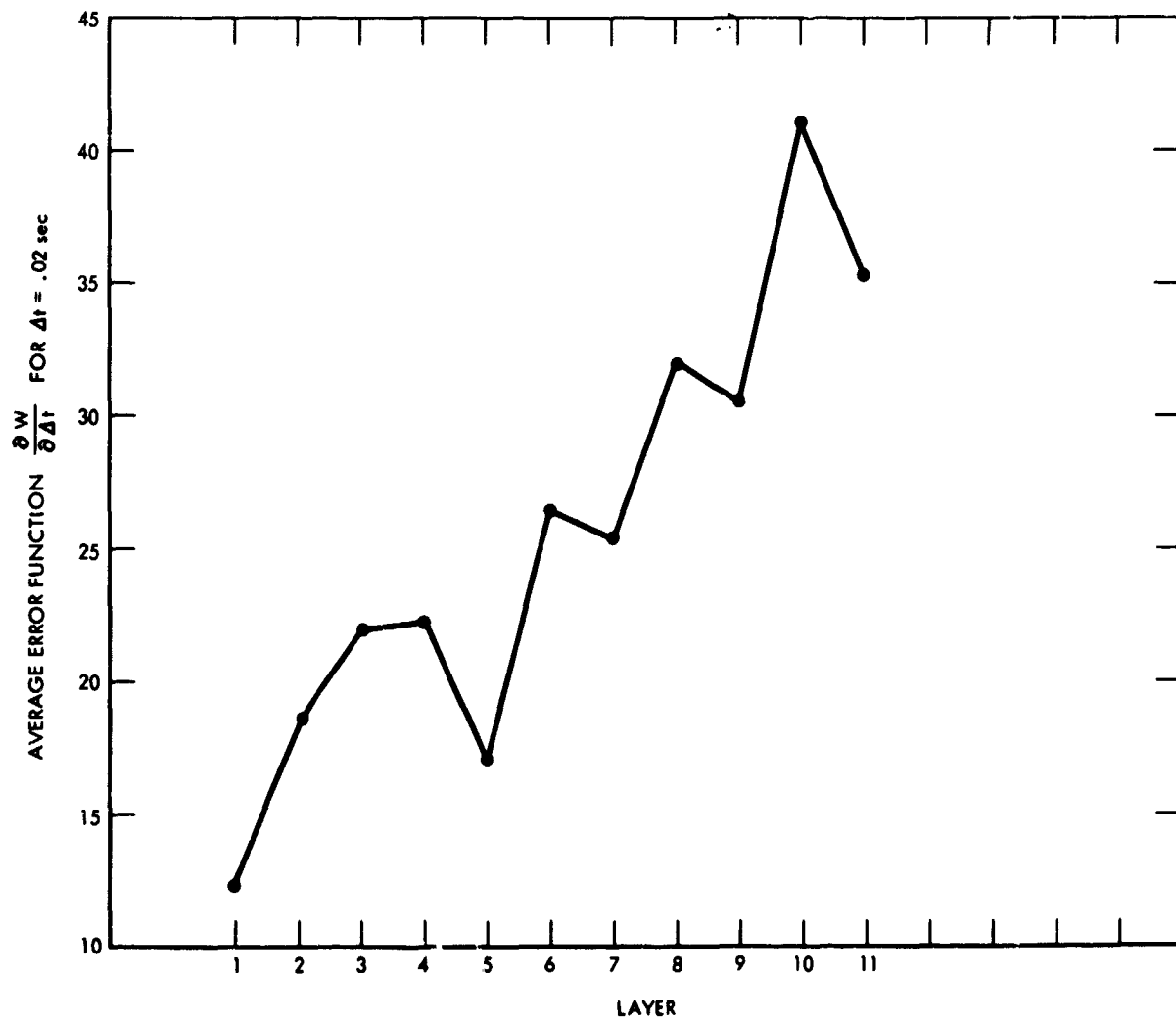


FIGURE 39.—Graph of $\frac{\partial T}{\partial \Delta t}$ vs. layer where $\Delta t = .02 \text{ sec}$. (average of 16 firings).

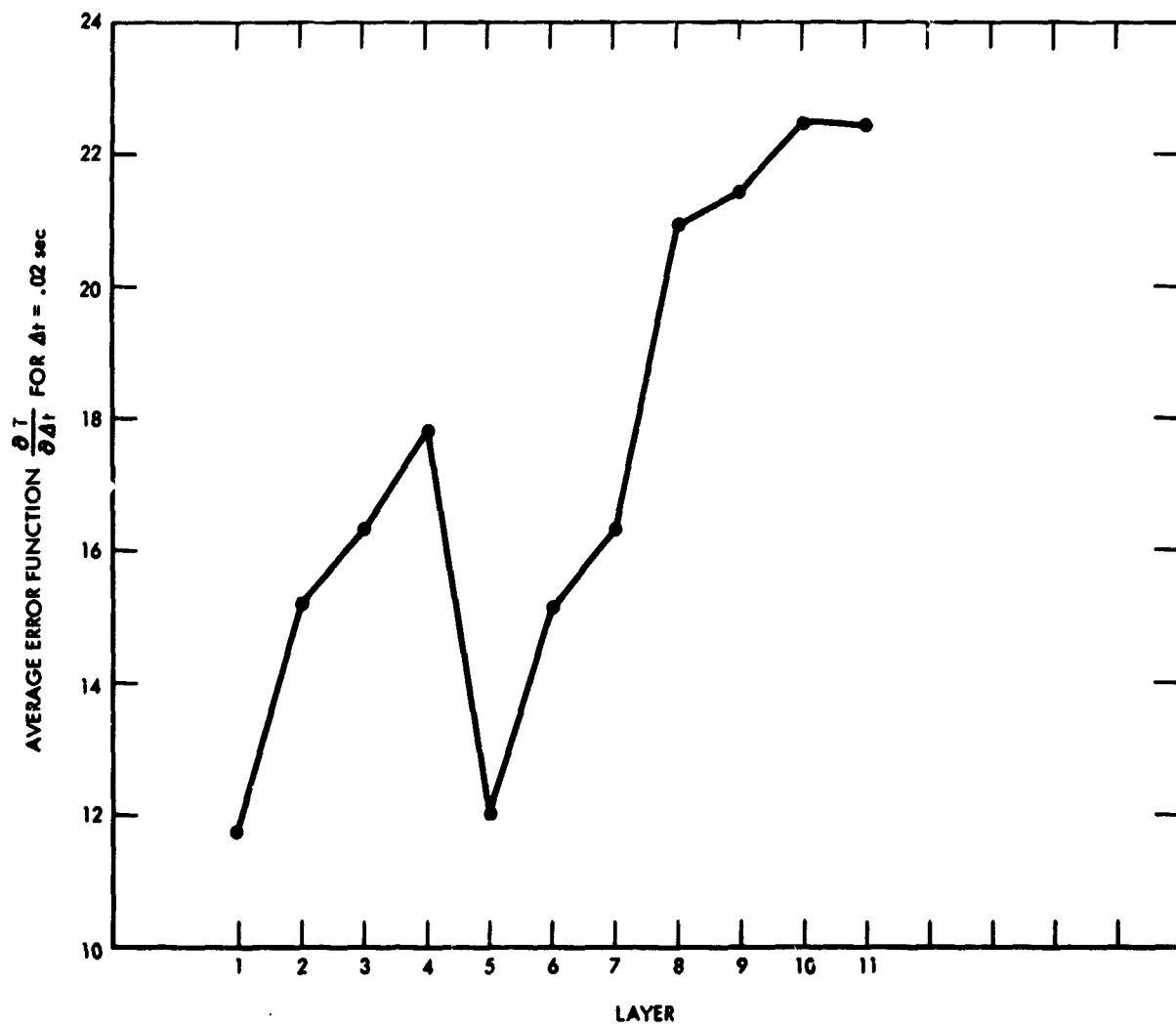


FIGURE 40.—Graph of $\frac{\partial W}{\partial \Delta t}$ vs. layer where $\Delta t = .02 \text{ sec}$. (average of 16 firings).

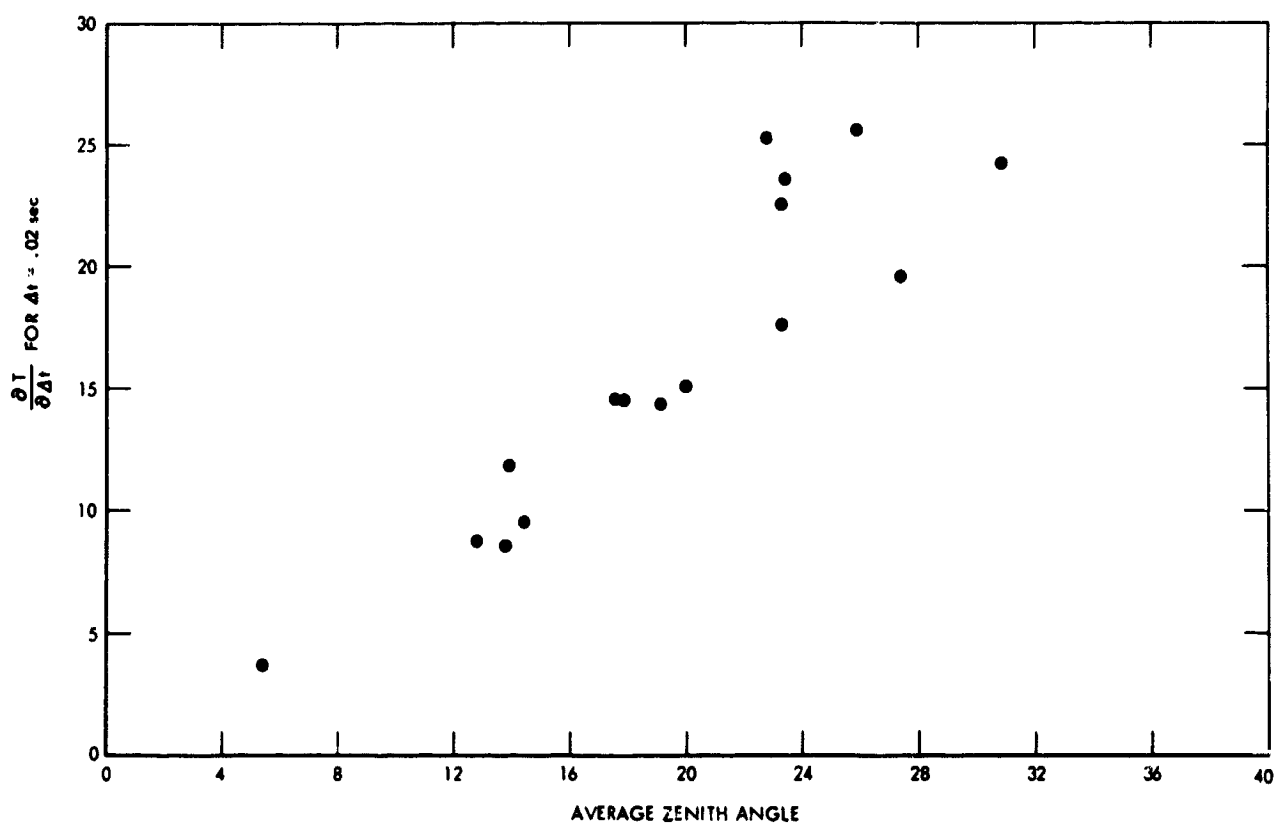


FIGURE 41.—Dependence of $\frac{\partial T}{\partial \Delta t}$ upon zenith angle, where $\Delta t = .02$ sec. for the 16 firings 1960–1963.

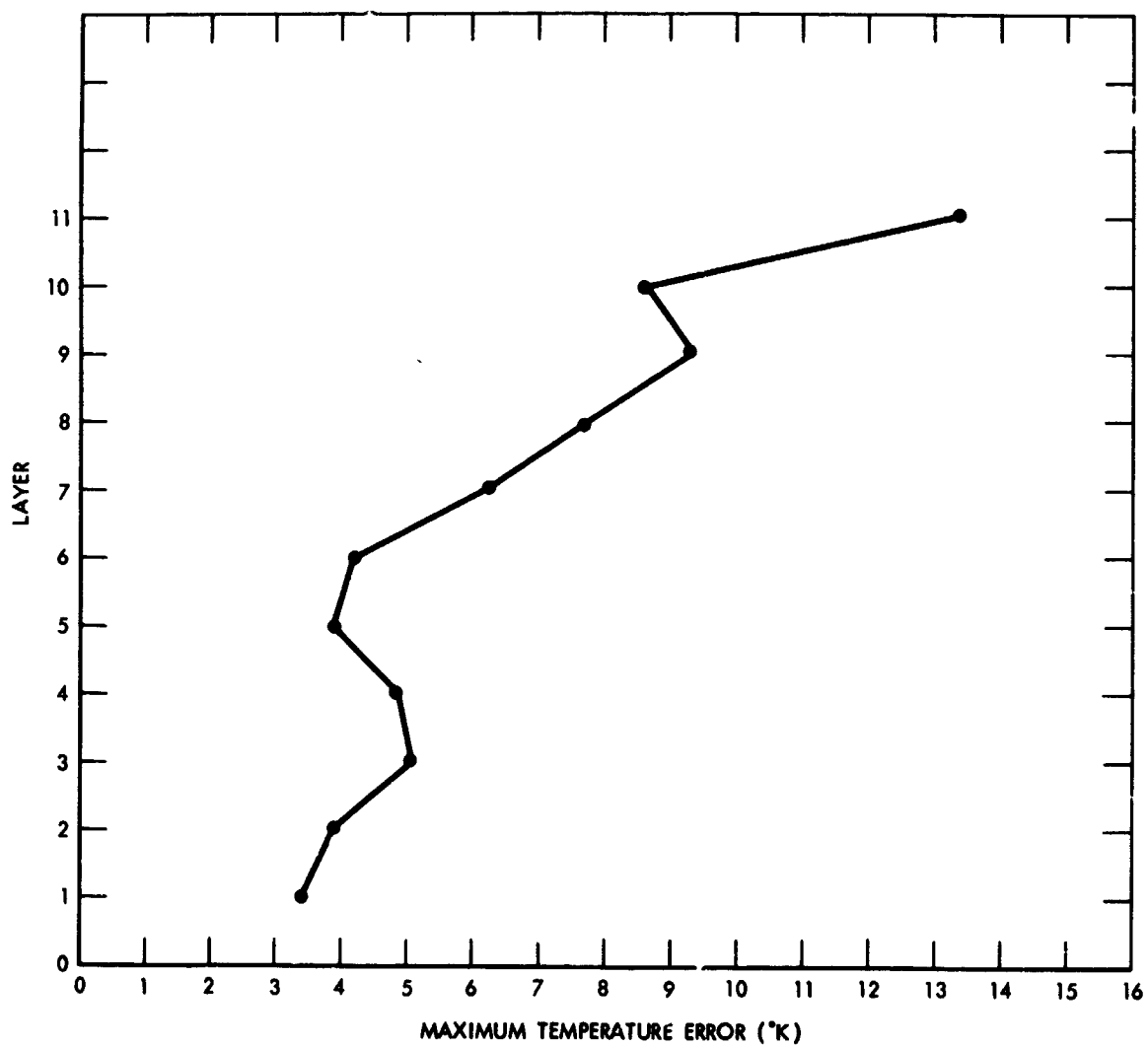


FIGURE 42.—Maximum temperature errors for the 16 firings 1960–1963.

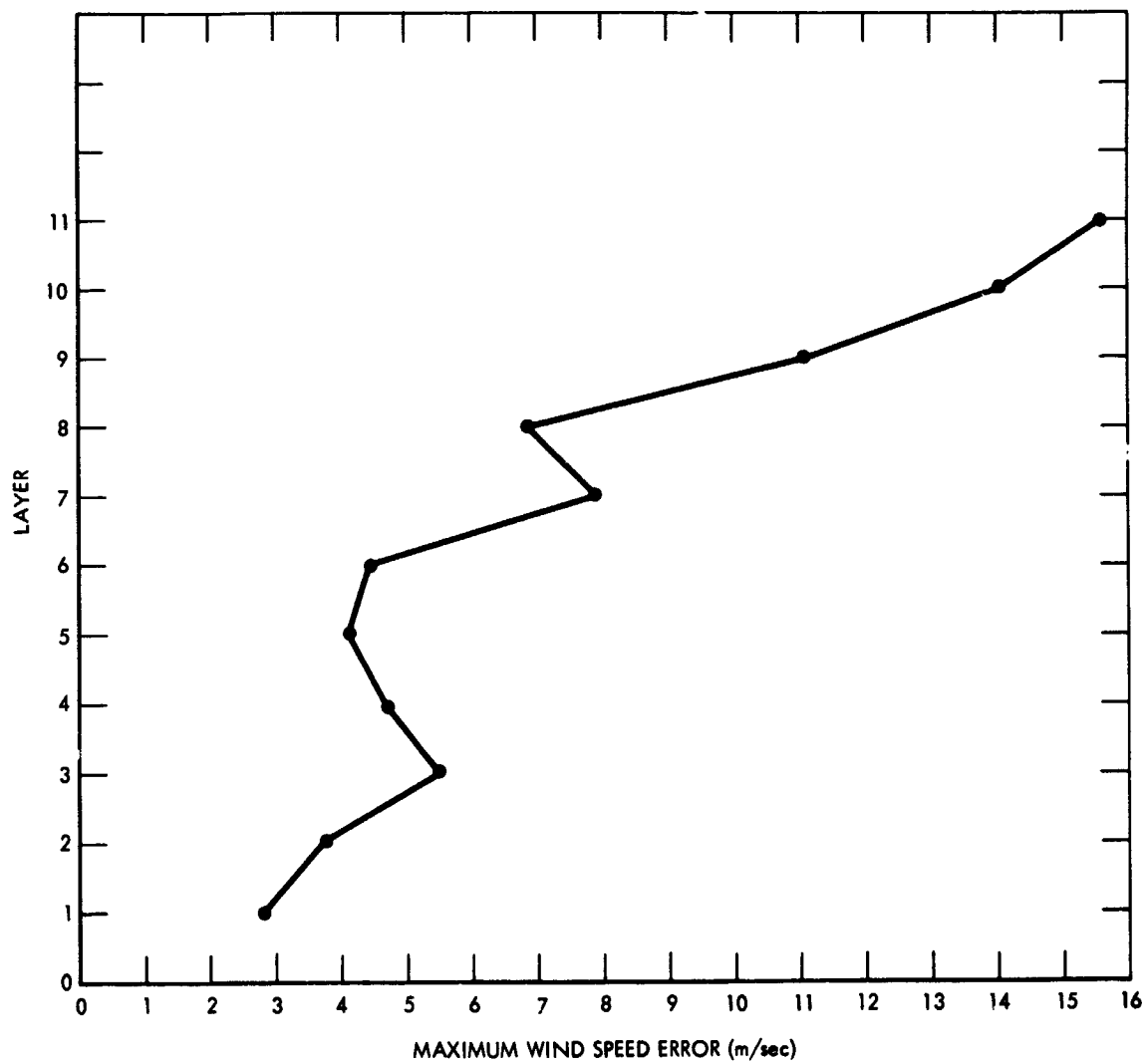


FIGURE 43.—Maximum wind errors for the 16 firings 1960–1963.

NO 5-10370

NEW KNOWLEDGE OF THE EARTH'S ATMOSPHERE FROM THE AERONOMY SATELLITE (EXPLORER XVII)*

N. W. SPENCER, G. P. NEWTON, C. A. REBER, L. H. BRACE AND R. HOROWITZ

The Explorer XVII satellite performed direct, very localized measurements of the total neutral particle density, the concentration of neutral particle masses 4, 14, 16, 28, and 32 and the temperature and concentration of thermal electrons, between the altitudes of 258 km and 920 km over those regions of the earth where the satellite was accessible to the Minitrack network, and in particular between $\pm 58^\circ$ latitude. Pressure gages on the satellite showed that the total density at 280 km was about 50 percent lower than is given by the appropriate atmospheric models based on satellite drag measurements. Daily variations in total density are more strongly dependent on a_p (the magnetic index) than had been believed previously. Neutral mass spectrometers showed that He is the predominant neutral constituent above 600 km, O is predominant between 250 km and 600 km, and N₂ is predominant below 250 km. The scale heights of the various constituents agree in general with the corresponding model atmospheric scale heights. Langmuir probe results confirmed the global extent of thermal non-equilibrium ($T_e > T_g$) and provided high resolution of the diurnal variation of electron temperature and density at several stations. For example, the electron temperatures near the F₂ maximum over Blossom Point show a nighttime value of about 1100° K, followed by a mid-morning maximum of 2800° K and an afternoon plateau of 2200° K. A consistent and strong latitude effect, evident particularly at Blossom Point, caused a significant positive gradient in electron temperature (the order of 25° K/degree of latitude) and an inverse gradient in electron density in a manner approximately in accord with recent theories of Hanson and Dalgarno.

INTRODUCTION

The Explorer XVII satellite, Figure 1, was designed to provide direct measurements of aeronomic parameters as a basis for new studies of the physics of the Earth's upper atmosphere. Thus, instruments were selected for the satellite which would provide both total and relative concentration of the neutral particles, and high-resolution measurements of the electron temperature and density; all of considerable significance in studies of the physical processes controlling the upper atmosphere. These data would help also to (a) clarify and define the structural properties of the atmosphere, previously established primarily through inferences from satellite drag measurements, and (b) investigate the variability and dependence of the atmosphere on solar conditions.

The technological advance of measurement techniques was also an objective of the project, as

part of a continuing effort to improve experimental capability. The application of laboratory-developed techniques required engineering as well as measurement technique refinement and adaptation to new environments. The atmospheric data to be obtained, if it were to be of maximum benefit consistent with its timeliness, required computer usage for processing the large quantity of data (the order of 2×10^9 bits of information).

EXPERIMENTS

The choice of experiments to obtain the desired data was based mainly upon laboratory vacuum technique experience; thus both thermionic and cold-cathode pressure gages, derived from proven laboratory sensors, were selected for the measurements of the total neutral particle density. Double-focusing magnetic mass spectrometers were chosen for the measurements of neutral particle concentrations. Two Langmuir probe experiments were employed for measurements of the electron temperature and ion density.

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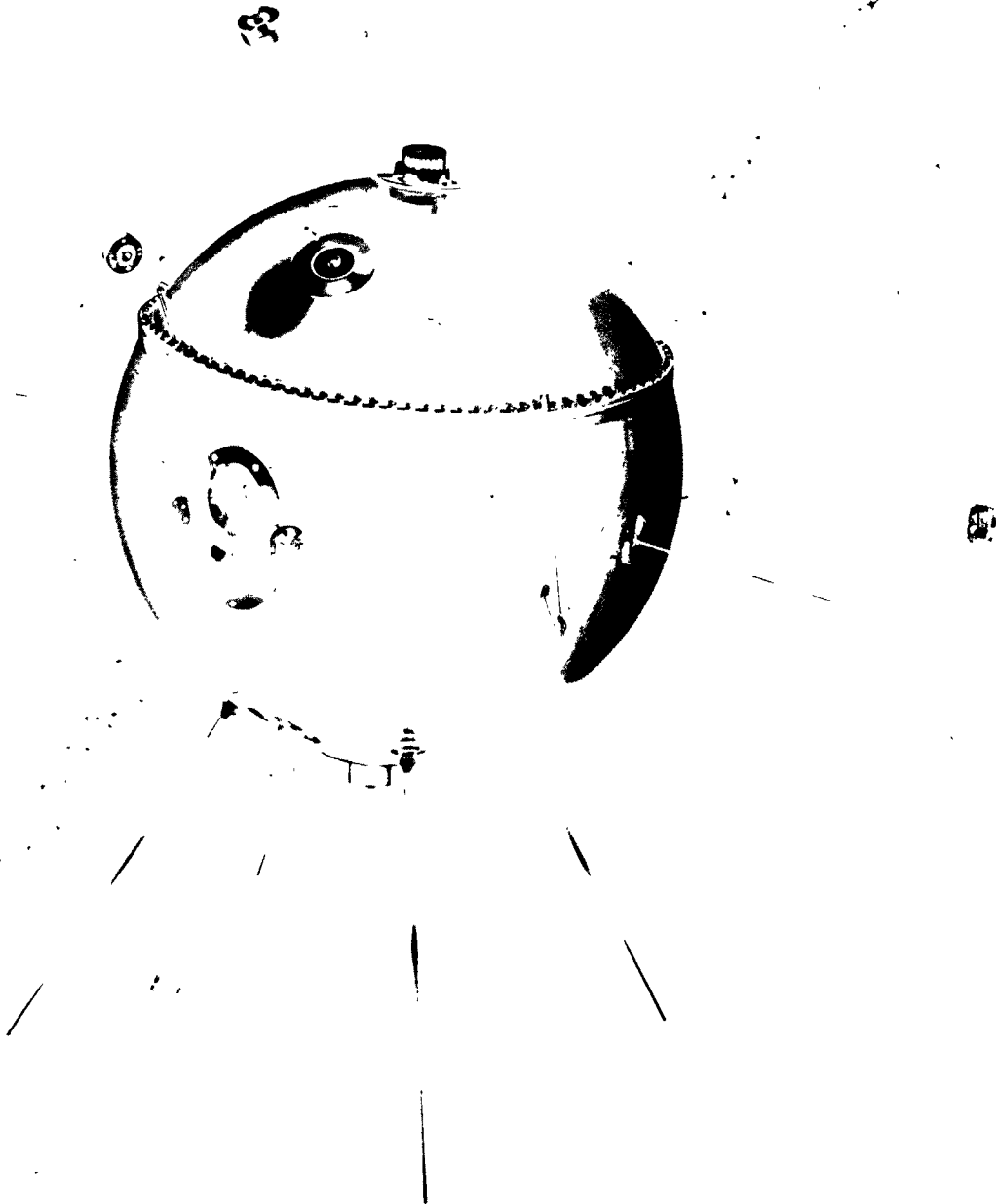


FIGURE 1.—Artist's conception of Explorer XVII in orbit.

It is of special significance from the point of view of determining local values of the ambient concentrations of constituents that proven laboratory vacuum system techniques were employed in construction of the satellite. For example, to minimize potential sources of contaminating gases, nearly all external surfaces were constructed of stainless steel, and all satellite joints were either welded or utilized copper shear gaskets for vacuum sealing. A significant effort was thus devoted to making the satellite-sensor system a true inside-out ultra-high vacuum system so that the experimenter could be confident that the atmospheric samples measured were not contaminated by materials carried aloft by the satellite.

Pressure Gage Experiments

The density measurements made by the Explorer XVII ionization gages extended to satellite usage the technique employed in rocket manometer experiments (Schultz, Spencer, Reifman, 1948) (Horowitz, Kleitman, 1953) to measure the total atmospheric density. The description of the instrumentation, the calibration technique, and the flight response of the Explorer XVII sensors have been reported (Newton, Pelz, Miller, Horowitz, 1963).

Four independent pressure gage experiments were employed: two Bayard-Alpert type (thermionic cathode) ionization gages and two Redhead type (cold-cathode, magnetic) gages. Each sensor was equipped with a special vacuum-sealed orifice that could be opened after the satellite was in orbit. Thus it was possible for the sensor to be properly cleaned, calibrated, sealed under vacuum, and opened on command to the space environment. This procedure, whose validity was previously established through rocket experience, assured the necessary high degree of vacuum cleanliness for the sensors.

The use of both cold and thermionic cathode gages was considered essential because of uncertainties in (a) the response of a hot-cathode gage in a sometimes predominately atomic oxygen environment (not subject to adequate laboratory calibration) and, (b) the general applicability of ionization gages to the high-velocity satellite environment. At the same time, a desirable redundancy was accomplished and valuable studies of the usefulness of the two fundamentally differ-

ent sensors were made possible. Each pressure gage was provided with an appropriate electrometer amplifier and other electronic support devices which enabled conversion of the sensor output current to a voltage suitable for telemetry. The electronic systems also included provision for in-orbit current calibration of the amplifiers once during each operation of the gages.

Neutral Particle Mass Spectrometers

Two identical double focusing magnetic mass spectrometers were employed for the determination of the local concentrations of atmospheric helium (mass 4), atomic nitrogen (mass 14), atomic oxygen (mass 16), molecular nitrogen (mass 28) and molecular oxygen (mass 32). Although detailed descriptions of the spectrometers are provided elsewhere (Meadows, 1960) (Hall, Howden, Iwasaki, 1960), (Spencer, Reber, 1962), (Reber, Hall, in prep.), the significant features are summarized here. The external ion source, designed to reduce the interactions of the sampled particles with the sensor, was followed by an electrostatic ion lens which focused the relatively high energy ions on the entrance slit of the analyzer. In this way, it was possible for the instrument to accept particles from a 2π steradian solid angle, and up to 12 *ev* kinetic energy. In the magnetic analyzer the beam of ions was separated according to mass, a given mass falling on the appropriate collector electrode in spectrograph fashion.

Each of the spectrometers was provided with a sensitive electrometer and logarithmic amplifier for conversion of the collected ion current to the proper telemetry voltages. Electronic logic circuitry accomplished the required changes in sensitivity and mass-number selection. As with the pressure gages, a vacuum-seal arrangement was adopted which permitted exposure of the spectrometer ion source to the atmosphere after orbit was attained.

The primary data analysis task was to determine the relationship of the measured ion currents to the ambient neutral particle densities. Particles could enter the region of ionization in one of three ways: (1) directly with no collisions; (2) after suffering one or more collisions with surfaces in the source region; (3) by entering the spectrometer analyzer, becoming thermalized, and

subsequently being re-emitted. The relationships between the measured ion currents and the ambient atmosphere were computed on the basis of these three mechanisms and the laboratory gas calibrations. The validity of these calculations is demonstrated by the fact that the total mass-density measured by the spectrometer is in satisfactory agreement with that obtained independently by the companion pressure gage experiments described above.

Langmuir Probes

Two independent Langmuir probe systems, based on established techniques and previous rocket usage (Spencer, Brace, Carignan, 1962), (Brace, Spencer, Carignan, 1963), (Nagy, Brace, Carignan, Kanai, 1963) were employed to provide measurements of the ion concentration (N_i), and the electron temperature (T_e) of the ionosphere. Each probe system used a cylindrical electrode (projecting into the plasma) whose potential was varied with respect to the satellite shell. The resulting current to the probe was converted to a voltage suitable for telemetry.

Using the following equation, the temperature was derived from the electron current to the probe as it was swept from the satellite potential to the plasma potential:

$$\frac{d \log_e i_e}{dV} = \frac{e}{kT_e} \quad (1)$$

To localize the T_e measurement, the electron temperature probe was swept at a rate of 10 sweeps per second; and to maximize the resolution, the voltage was swept in two ranges, 0 to $\frac{3}{4}$ V and 0 to $1\frac{1}{2}$ V, respectively. As a result, each temperature measurement was completed in less than 400 meters of the satellite path, and to that extent represents a point measurement. The telemetry sampling rate was sufficiently high to permit determination of temperature values as low as 400°K (although the lowest temperatures actually recorded were about 900°K).

The ion density probe was swept from -3 to +2 volts in a 2-second period, which was long compared with the satellite spin period of 0.7 seconds. N_i was derived from the ion current maxima which occurred each time the probe axis was perpendicular to the velocity vector.

Sensor Location

As noted above, the capability to provide measurements of the constituents of the space environment required close adherence to established laboratory vacuum techniques. Thus, in addition to providing a sealed housing which would not itself contaminate the local atmosphere, the eight sensors were located so as to provide maximum separation from each other.

Consideration of sensor orientation with respect to the direction of motion required that the mass spectrometers be located at the two ends of the spin axis to minimize changes in orientation during each mass-sampling sequence. The four pressure gages were distributed uniformly about the spherical satellite to insure that at least one gage would always experience pressure variations due to satellite spin, regardless of the spin axis orientation. One of each type of pressure gage was located on the satellite equator to enable a comparison of the gage responses under identical conditions. The two Langmuir probes were also located on the equator at points nearly diametrically opposite each other.

SUPPORTING SYSTEM

Interpretation of the data from the various sensors required detailed knowledge of the instantaneous angle between any sensor and the direction of motion of the satellite. This information was provided through the use of a multiple optical sensor arrangement, which enabled sensing the direction of the sun and/or moon and the instants of passage, during spin, of the Earth's horizons.

Direct measurement sensors like those employed in Explorer XVII provide time rates-of-change of data requiring high telemetry sampling rates. For example, the spectrometers and pressure gages required 60 samples per second and the "high speed" Langmuir probe required 180 samples per second. To meet these needs a pulse code modulation (PCM) telemetry system capable of 1000 samples per second was selected. This system had the additional advantage of providing a digital format which facilitated computer data processing.

The satellite was powered exclusively by silver-zinc cells, since solar cells presented the possibility of local contamination of the atmosphere. The

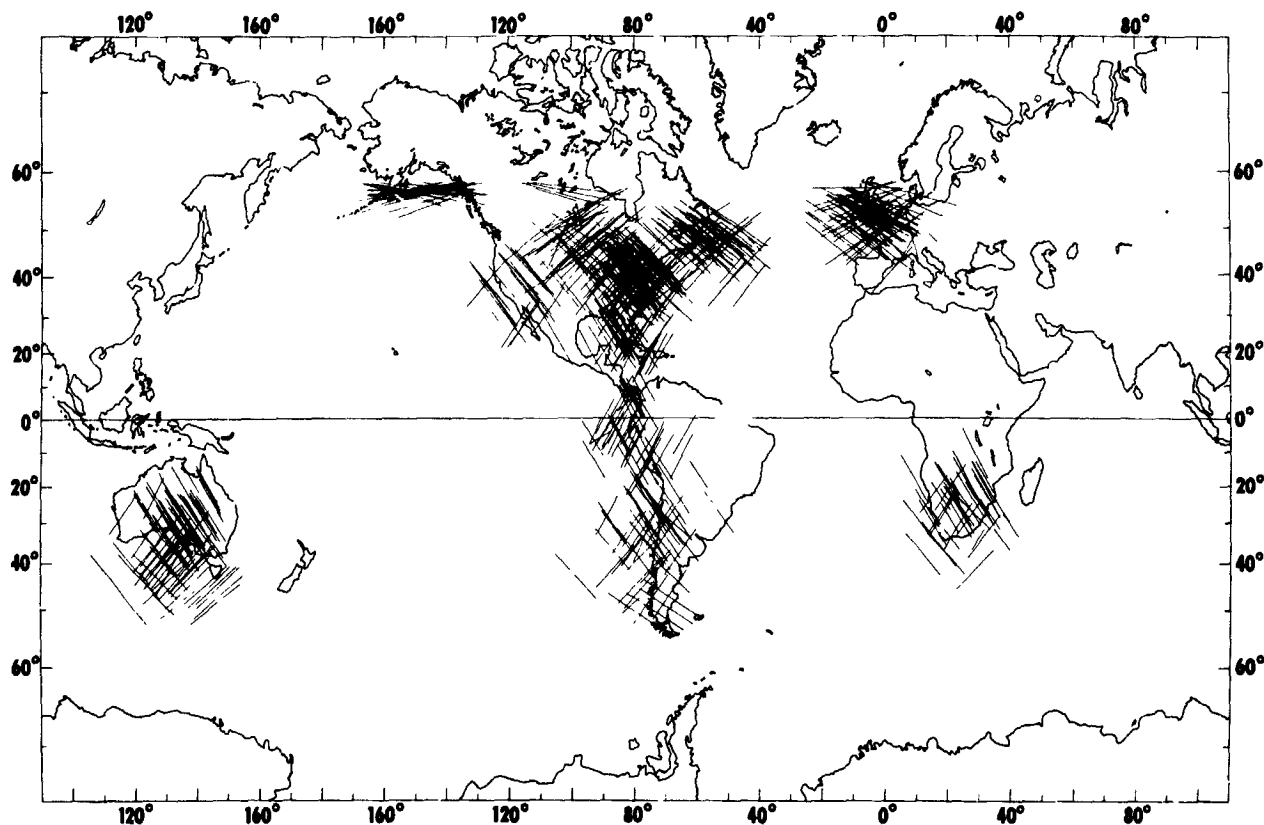


FIGURE 2.—Explorer XVII geographic coverage.

150 pounds of cells which were employed provided adequate energy to operate the entire satellite system for a total of 75 hours. A command-control system permitted the experiments to be turned on for four-minute periods, each of which was terminated by an internal programmer. Because a tape recorder was not employed, responses were confined to geographic regions of approximately 4000 kilometers diameter about each minitrack command station. Figure 2 illustrates the geographic coverage attained by showing the path of the satellite during each data-producing response. Table I summarizes some of the pertinent statistics of the satellite, and other information.

EXPERIMENTAL RESULTS

Pressure Gage Results

As noted above, four independent gage systems were employed on the satellite, and all gages operated in orbit. Three of the gages gave useful and meaningful data during the active lifetime of the satellite. The fourth gage, however, experienced an apparent decrease in sensitivity after

TABLE I.—Explorer XVII Statistics

Launch Date.....	April 3, 1963
Inclination.....	58°
Perigee.....	258 km
Apogee Range.....	920-870 km
Useful lifetime.....	100 days
Perigee motion.....	+39° to +58° to -18°
Data responses.....	650 on command
Telemetry.....	PCM—8640 bits/sec
Spin rate.....	90 RPM
Power supply.....	Chemical
Size and shape.....	1 meter sphere
Weight.....	410 pounds

opening in orbit and no atmospheric structure data is available at present from this gage. Some of the results from the other three gages for five northern mid-latitude stations, representing approximately 25 percent of the total available data are presented here.

Figure 3 shows the measured atmospheric density over Grand Forks, Minnesota, during a pass at 2000 hours local time and demonstrates typical resolution of the density gage data. It may be

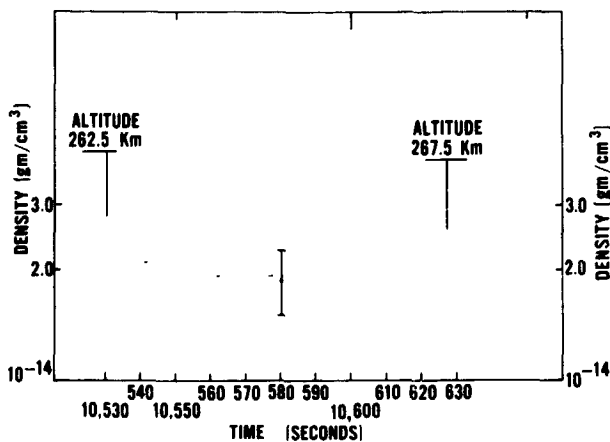


FIGURE 3.—Density data derived from a single pass over Grand Forks, Minn., demonstrating the resolution of the pressure gauge data.

seen that the total altitude change was only 5 km for this pass, and at the average altitude of 265 km, the atmospheric density was 2×10^{-14} gm/cm³. Each point shown is the average of the three independent density measurements. The resolution of the density data from each gage is such that the density was measured every 700 milliseconds or once every six kilometers along the satellite orbit. The observed small scatter of the density data makes quantitative density scale-height determination possible for passes possessing significant altitude changes. The magnitude of the error in the absolute value of the density is ± 25 percent for this pass and can be attributed primarily to uncertainty in the absolute calibration of the sensors in the laboratory. At an altitude of 600 km, the error in the absolute value of the measured density is ± 55 percent for some passes. Generally, the precision of the density measurements is better than ± 20 percent.

Figure 4 (Newton, Horowitz, Priester, 1964) provides a comparison of the atmospheric densities directly measured by the gages in the altitude range 258 to 300 kilometers, to the density obtained from drag observations of Explorer XVII (Bryant, 1964) and Injun 3 (Jacchia, Slowey, 1964). All data are normalized to a height of 280 kilometers for comparison purposes, by using density gradients obtained from the Harris and Priester model with $S=90$.

The Injun 3 data, indicated by squares, were selected for quiet geomagnetic conditions ($A_p=2$), for the time interval of 18 February through

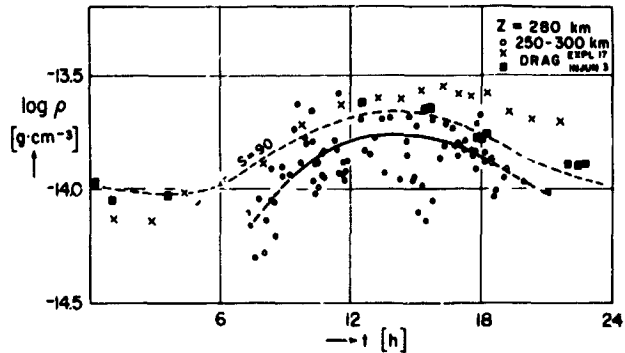


FIGURE 4.—Comparison of directly-measured and drag-derived atmospheric densities.

30 June 1963. During this period, the latitude of the satellite perigee varied between -40° and $+60^\circ$. The drag data from Explorer XVII, indicated by crosses, correspond to the time interval from 3 April to 6 July 1963, when the latitude of the satellite perigee was between $+58^\circ$ and -20° . All Explorer XVII data in the figure have been reduced to quiet conditions ($A_p=0$) by using the preliminary linear reduction relation:

$$\log \rho = 0.006 A_p$$

Further analysis of the gage data indicates, however, that the correlation between density and geomagnetic activity should be a steeper, non-linear relation with the steepest portion applicable for low A_p (Newton, Horowitz, Priester, 1964). Application of the modified relation is expected to (a) remove some of the scatter from the gage data (which reflects real atmospheric variations), and (b) lower the average value of the directly measured densities by a small amount.

It is observed that the densities determined from drag are systematically 40 percent to 50 percent greater than the normalized densities measured by the gages, and that this separation is just outside the combined, stated uncertainties of the two sets of data. This difference is significant but at this time is not considered serious, since it could be accounted for by modest changes in the altitude to which the drag data are assigned, the drag coefficient, or the gage calibration constants.

Figure 5 shows measured atmospheric density-versus-altitude for the altitude range 258 to 600 km. These data result from approximately 60 passes for an A_p between 0 and 10, $F_{10.7}$ between

70 and 100, and most local times. It is seen that considerable variation in the atmospheric density occurs, resulting primarily from the differences in local time, a factor of 5 diurnal variation at 360 km being observed. The Harris and Priester model densities ($S=90$) for 0400 and 1400 hours are shown for comparison purposes.

Continued analysis of the Explorer XVII data is currently underway to further define:

- (1) The quiet atmosphere and its variation with local time.
- (2) The variations from the quiet atmosphere resulting from solar and geomagnetic disturbances.
- (3) Other effects not now apparent.

Mass Spectrometer Experiment Data

Concentrations as a function of altitude are shown in Figure 6 (Reber and Nicolet, 1964) where Nicolet's model (private communication) is included for reference. The data points shown are measured values, converted to ambient number densities. Measurements of each mass taken on the same satellite interrogation are made approximately one minute apart and are shown in the figure joined by straight lines. The numbers refer to orbit numbers. Table II lists pertinent information for the passes shown in the figure.

Of particular interest are the helium concentration vs. altitude, and the altitude regions where helium, atomic oxygen and molecular nitrogen are predominant. It is seen that helium is the major constituent above 600 km, molecular nitrogen is predominant below 250 km, and atomic

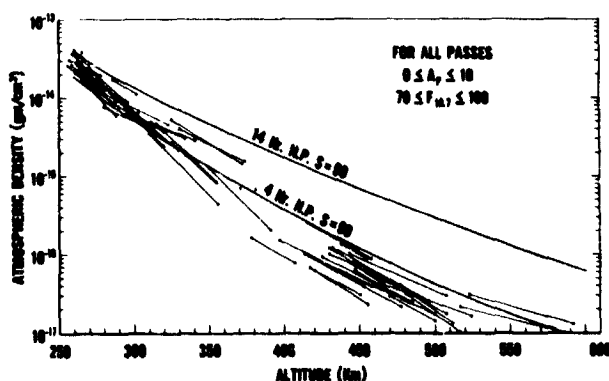


FIGURE 5.—Atmospheric density versus altitude measured by the pressure gauge experiment.

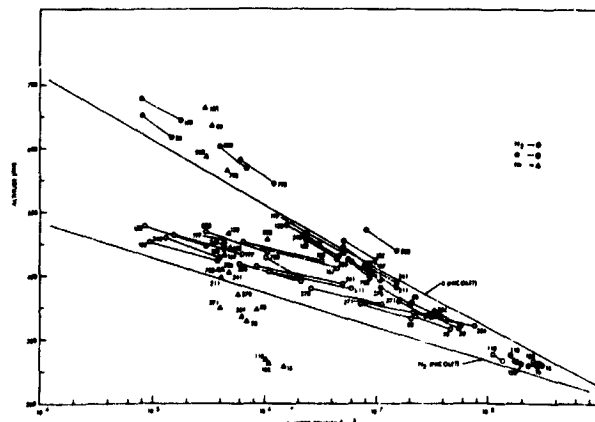


FIGURE 6.—Average daytime and nighttime concentrations of He, O, and N₂ from Explorer XVII mass spectrometer experiment.

oxygen is predominant between these levels. It should be noted that the scale heights of the constituents at higher altitudes correspond to temperatures of about 700° at night.

A possible deviation of the nighttime N₂ distribution from a diffusive equilibrium condition at altitudes less than 400 km is also suggested by this figure. One possible explanation is that the diffusion time in this altitude region is the same order of magnitude as the diurnal variation period. It should also be kept in mind, however, that these data represent a variety of times and geographic locations and thus do not accurately present an instantaneous vertical profile.

Figure 7 shows the variation of mean mass with altitude. The mean mass was computed using the major constituents, N₂, O and He. Hydrogen, which the instrument was not designed to measure, could significantly reduce the value of the mean mass at higher altitudes.

The variation of concentration ratios of helium-to-oxygen and atomic oxygen-to-molecular nitrogen with altitude is shown in Figure 8. The solid lines drawn through the points are averages; as in all the data presented, the points are measured values and are not averaged, nor do they reflect any smoothing. It can be seen here again that helium is the dominant component above 600 km and that molecular nitrogen is predominant below about 250 km.

Langmuir Probe Experiment Results

Figure 9 shows the detailed variation of T_e and

TABLE II.—*Tabulated Mass Spectrometer Data*

Pass & Station	Date	Local Time	α	Geo. Lat.	Geo. Long.
# 15 BP.....	4/4/63	21.15 hrs.	6°	38.5°	-75.0°
# 50 COL.....	4/6/63	0.65	16°	57.0°	-149.0°
# 80 COL.....	4/8/63	0.99	9°	55.0°	-147.0°
# 80 FTM.....	4/8/63	4.89	63°	18.0°	-92.0°
#118 BP.....	4/10/63	18.81	70°	37.0°	-72.0°
#120 GF.....	4/11/63	20.32	51°	51.0°	-98.5°
#138 BP.....	4/12/63	2.51	12°	37.0°	-84.0°
#152 BP.....	4/13/63	2.01	14°	39.5°	-68.5°
#167 BP.....	4/14/63	1.65	20°	39.5°	-75.0°
#182 BP.....	4/15/63	1.54	25°	37.0°	-78.0°
#183 QUI.....	4/15/63	3.26	23°	4.5°	-79.0°
#197 BP.....	4/16/63	1.43	27°	34.0°	-81.5°
#211 BP.....	4/17/63	0.53	45°	41.5°	-71.5°
#226 BP.....	4/18/63	0.48	53°	38.5°	-74.0°
#241 BP.....	4/19/63	24.19	62°	38.0°	-79.5°
#242 MOJ.....	4/19/63	0.64	54°	31.0°	-121.5°
#254 NFL.....	4/20/63	22.75	82°	49.0°	-53.0°
#270 BP.....	4/21/63	23.30	80°	41.5°	-71.5°
#271 GF.....	4/21/63	22.88	85°	45.0°	-101.0°
#708 NFL.....	5/20/63	7.18	39°	49.5°	-49.5°
#795 OOM.....	5/26/63	15.81	63°	-34.0°	137.5°
#800 JOB.....	5/26/63	15.90	65°	-37.5°	19.0°
#888 JOB.....	6/1/63	13.24	33°	-27.0°	25.0°

The local sun time, angle of attack (α), geographic latitude and longitude are averaged over the four minute pass. The stations involved are: BP—Blossom Point, Md.; COL—College, Alaska; FTM—Fort Myers, Fla.; GF—Grand Forks, Minn.; QUI—Quito, Ecuador; MOJ—Mojave, Calif.; NFL—Newfoundland; OOM—Woomera, Australia; JOB—Johannesburg, South Africa.

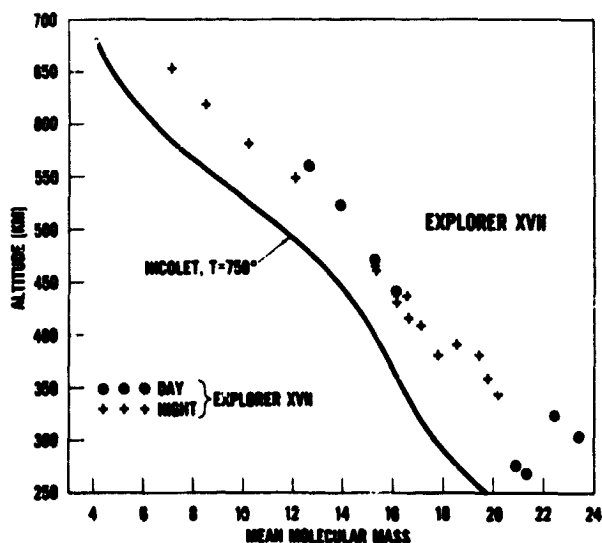


FIGURE 7.—Mean molecular mass versus altitude from mass spectrometer.

N_2 , measured during a series of near perigee passes near the F_2 maximum over Blossom Point, Maryland. These data were obtained over the three

month satellite lifetime during which the rotation of the orbit plane caused a complete diurnal variation to occur. The data are shown as points or pairs of points, the latter corresponding to the measured values at the beginning and at the end of individual satellite passes.

At first glance one is struck by the lack of correlation between the gross diurnal variations (smoothed curves) and the changes during the individual passes, especially in the afternoon. Since perigee passes such as these can exhibit very little altitude change, this in-pass variation must be largely latitude dependent (10° change in latitude within average pass). The average in-pass change in T_e implies a latitude gradient near Blossom Point of approximately 25° K per degree of latitude, corresponding to about a 10 percent change in T_e within a pass. With few exceptions the changes in T_e within a pass are accompanied by an inverse change in N_2 which is even greater than 10 percent.

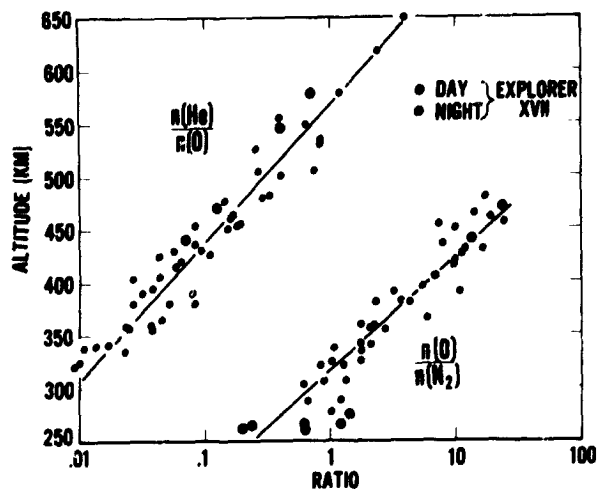


FIGURE 8.—Ratios of $n(\text{He})/n(\text{O})n(\text{N}_2)$ versus altitude from mass spectrometer experiment.

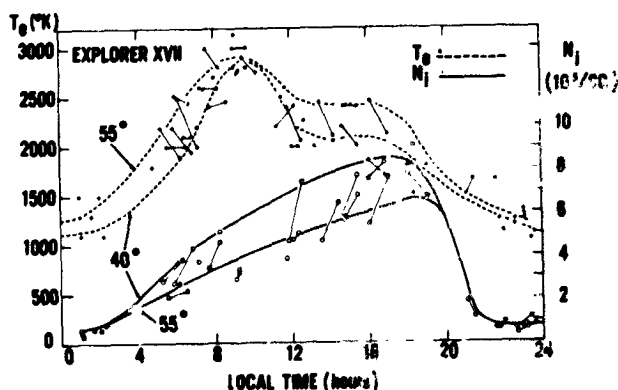


FIGURE 9.—Diurnal variation of T_e and N_i for magnetic latitudes 40–50°N and altitudes 258 to 350 km, from the Langmuir probes.

Plots similar to Figure 9 have been prepared from T_e and N_i measurements at two other latitudes (10° N at Quito, 60° N at College), and the resulting gross diurnal variation curves at all three latitudes are shown in Figure 10. These data also correspond to the region of the F_2 maximum (below 400 km).

The T_e variation at the F_2 maximum at all latitudes shown is characterized by a steep morning rise, a midmorning maximum, and afternoon plateau, and a gradual decrease near sunset. The nighttime values of T_e are somewhat variable but are always significantly above the neutral particle temperature (Harris and Priester, 1962), particularly at College, Alaska where the summer night at F_2 altitudes is short or non-existent.

The values of N_i rose gradually throughout the

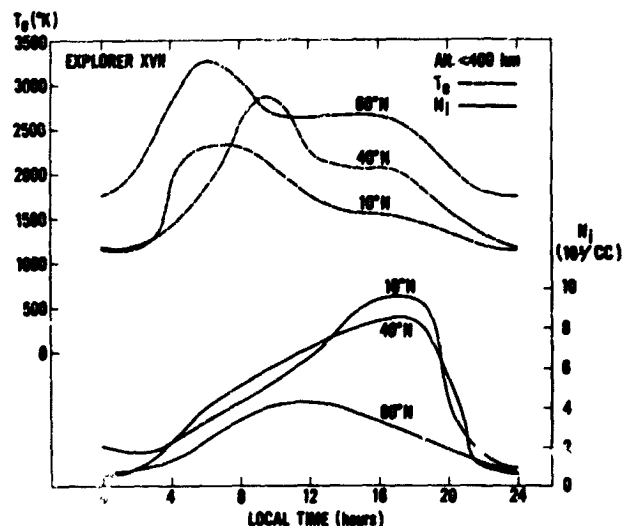


FIGURE 10.—Averaged T_e and N_i showing the diurnal variation above three selected stations; Quito (10°N), Blossom Point (40°N) and College (60°N).

day, reaching a maximum density in the late afternoon, except at College where the maximum occurred in the early afternoon.

It should be noted that the curves in Figure 10 represent direct "in situ" measurements above specific geographic locations during the late spring and summer of 1963, and therefore should not be considered models of the diurnal variation at other altitudes, longitudes, and seasons. However, the analysis of higher altitude data from these sites, as well as data from other sites is now in progress, and it is hoped that these data will help provide a better understanding of the global structure of the ionosphere.

In conclusion, it is clear from the simultaneous measurements that T_e and N_i are related in a generally inverse manner which agrees reasonably well with the inverse square relationship predicted by Hanson (1962) and Dalgarno, et al (1963). The elevated nighttime values of T_e show that there is a heat source for the electrons at night, although its energy content is only a few percent of that required to cause the electron temperatures observed in the daytime (Brace, Spencer, Dalgarno, 1964).

ACKNOWLEDGMENTS

The success of a satellite laboratory such as discussed above results, clearly, from the efforts of many individuals. The authors gratefully and

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