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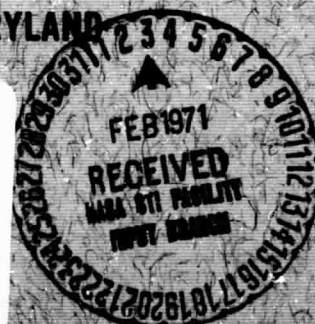


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# SOLAR CYCLE VARIATION OF PLANETARY EXOSPHERIC TEMPERATURES

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The purpose of this note is to show that a simple scaling law, based on an approximate solution of the heat balance equation, can be used to determine the variation of exospheric temperature  $T_{\infty}$  over a solar cycle.  $T_{\infty}$  is defined as the temperature of the isothermal region (on account of heat conduction) extending into the exosphere. Thus,  $T_{\infty}$  also controls the escape of gases from planetary atmospheres.

It is assumed that in the thermosphere, the effective heat input due to EUV radiation,  $Q_{uv}$ , is balanced by the divergence of a conductive heat flux,  $E$  (neglecting IR radiation from atmospheric constituents) according to

$$\text{div } E = Q_{uv} \quad (1)$$

where  $E = -K_n(T) dT/dz$ , with  $K_n(T)$  the thermal conductivity and the  $dT/dz$  the temperature gradient with altitude and  $Q_{uv} = n_j \sigma_a \epsilon_j I_{\infty} e^{-\tau}$  with  $n_j$  the number density of the absorbing constituent,  $\sigma_a$  its average absorption cross section,  $\epsilon_j$  the heating efficiency of solar EUV radiation of intensity  $I_{\infty}$  outside the atmosphere (i.e., in the exosphere) and  $\tau = \int_0^z n_j \sigma_a dz$  the optical depth.

The integrated form of (1) is given by

$$\int_{z_0}^{z_{\infty}} \epsilon_j I_{\infty} (1 - e^{-\tau}) dz = \int_{T_0}^{T_{\infty}} K_n(T) dT \quad (2)$$

where  $T_\infty$  is the exospheric temperature,  $T_0$  is the temperature at the mesopause and  $z_\infty$  and  $z_0$  are the corresponding altitudes.

In the aeronomic literature it is generally assumed that  $K_{nj} = A_j T^{1/2}$  based on a rigid sphere approximation.<sup>1</sup> However, as pointed out recently,<sup>2</sup> such a temperature dependence may not always be applicable. Experimental data<sup>3</sup> and quantum-mechanical calculations<sup>4</sup> show that for atmospheric gases  $K_{nj} = K_0 T^{\nu_j}$ , where  $\nu_j > 0.5$ , so that the temperature dependence in the heat conduction equation will not be identical for all planetary atmospheres as the result of their different composition. ( $K_O = 67 T^{0.71}$ ,  $K_{CO_2} = 1.5 \times T^{1.23}$ ,  $K_{He} = 21 T^{0.75}$  and  $K_H = 16.4 \times T^{0.73}$ )<sup>3,4</sup>

Integration of (2) following some simplifications leads to

$$\frac{\epsilon_j I_\infty}{n(z_\infty) \sigma_a} \simeq K_0 (T_\infty^{\nu+1} - T_0^{\nu+1}) \quad (3)$$

if we assume that  $\tau \rightarrow 0$  at  $z_\infty$ , i.e., in the exosphere, and  $\tau \rightarrow \infty$  at  $z_0$ , i.e., at the base of the thermosphere, and allowing for  $(n(z_\infty) \sigma_a)^{-1} \gg z_\infty - z_0$  (since  $\sigma_a \simeq 10^{-17} \text{ cm}^2$  and exospheric densities are of the order  $10^8 \text{ cm}^{-3}$ ).

For  $n(z_\infty)$  we substitute the density at the exobase  $n_b \simeq (\sigma_c H)^{-1}$  which derives from the exospheric condition that the mean free path is of the order of the local scale height  $H = k T_\infty / mg$ , where  $k$  is the Boltzmann constant,  $m$  is the mean mass and  $g$  is the acceleration of gravity at the exobase and  $\sigma_c$  is the gas kinetic collision cross section.

Thus, we can write (3), neglecting  $T_0$  compared to  $T_\infty$  (which seems permissible, at least for the terrestrial planets) as

$$T_{\infty}^{\nu} \propto \frac{\epsilon_j I_{\infty} k \sigma_c}{K_0 m g \sigma_a} \quad (4a)$$

and thus

$$T_{\infty} \propto I_{\infty}^{1/\nu} \quad (4b)$$

From satellite drag observations over a solar cycle<sup>5</sup> it is known that the terrestrial exospheric (daytime) temperature ranges from  $\sim 800^{\circ}\text{K}$  at solar minimum to  $\sim 2000^{\circ}\text{K}$  at solar maximum leading to a ratio  $T_{\infty} (\text{SMax})/T_{\infty} (\text{SMin}) \simeq 2.5$ . Although an empirical relation exists between solar 10.7 cm radio flux,  $S_{10.7}$  (which is an indicator of solar activity) and exospheric temperature,<sup>5</sup> the exact relationship between  $S_{10.7}$  and the EUV intensity  $I_{\infty}$  has yet to be established. Since for the terrestrial thermosphere,  $T_{\infty} \propto I_{\infty}^{1/0.71}$ , we infer from a temperature ratio of 2.5, that the EUV intensity has changed by a factor  $\sim 1.9$  over the solar cycle. This seems to be consistent with observations<sup>6</sup> and modelling of the terrestrial ionosphere.<sup>7</sup>

For a  $\text{CO}_2$  atmosphere (Mars, Venus) where  $T_{\infty} \propto I_{\infty}^{1/1.23}$  holds, we obtain a temperature ratio over the solar cycle of  $\sim 1.7$ . For Mars, Mariner 4 observations<sup>8</sup> near solar minimum showed an exospheric temperature  $T_{\infty} \approx 300^{\circ}\text{K}$ , while Mariner 6 and 7 measurements<sup>9</sup> near solar maximum showed  $T_{\infty} \approx 500^{\circ}\text{K}$ . Our estimated ratio of  $T_{\infty}$  over a solar cycle, thus agrees well with the observations. For Venus, Mariner 5 measured  $T_{\infty} \approx 650^{\circ}\text{K}$  near the middle of the solar cycle.<sup>10</sup> Thus, according to our estimate, the exospheric temperature of Venus at solar minimum should be  $T_{\infty} \approx 470^{\circ}\text{K}$  and that at solar maximum  $T_{\infty} \approx 800^{\circ}\text{K}$ .

Recently it was suggested<sup>11</sup> that the ratio of thermal energy  $3/2 kT$ , to gravitational energy  $m_H gr$ , of hydrogen is about  $1/2$  for the solar corona as well as for the earth's exosphere ( $r$  is the planetocentric distance). This leads to the condition that the thermal velocity of hydrogen  $v_{th} = (3 kT/m_H)^{1/2}$  is about equal to its circular orbital velocity  $v_0 = \sqrt{gr}$ . (Note that the escape velocity  $v_\infty = \sqrt{2} v_0$ ). From this relationship Matora<sup>11</sup> estimated the "limiting" exospheric temperature of a planet to be given by  $T_\infty \approx m_H gr/3 k$ . Accordingly, he obtained the estimates of  $T_\infty \approx 400^\circ K$  for Mercury,  $800^\circ K$  for Venus,  $2000^\circ$  for Earth,  $300^\circ K$  for Mars and  $25,000^\circ K$  for Jupiter as exospheric temperature for a gravitationally stable atmosphere. This limit is similar to the condition of Öpik<sup>12</sup> according to which atmospheres with an escape parameter  $B \equiv mgr/kT_\infty < 1.5$  are unstable. This condition corresponds to  $v_{th}/v_0 = \sqrt{2}$  or  $v_{th}/v_\infty = 1$ . (It should be noted, however, that with significant escape, the "effective" temperature well above the exobase, will be less than  $T_\infty$  due to the loss of high velocity particles).<sup>13</sup>

Thus, Earth and Venus seem to have exospheres which are stable even at solar maximum although they are subject to escape ( $B < 15$ ), while the exosphere of Mars is stable only near solar minimum, and escape should be excessive near solar maximum.

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