TECHNICAL
MEMORANDUM

LUNAR CONDUCTIVITY MODELS
FROM THE APOLLO 12
MAGNETOMETER EXPERIMENT

Bellcomm

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ABSTRACT

A number of conductivity models have been investigated
for compatibility with the Apollo 12 magnetometer data. Except
at the highest frequencies, a simple core-crust model is compatible
with the observed dayside transfer function, which is expressed
as the ratio of the lunar surface field spectrum to the inter-
planetary magnetic field spectrum. All conductivity profiles
exhibit a peak near 1500 km, when the models are constrained to
conform to the observed flat response at the higher frequencies.
However, at frequencies above .01 Hz the long wavelength limitation
of the theoretical model is no longer valid. A combination of
dayside and nightside models and data indicate that a conductivity
profile with a peak (.003 mho/m) near 1500 km radius and a core
conductivity of about .01 mho/m at 1000 km is compatible with
the observations, as is a monotonic conductivity profile with
.0005 mho/m at 1600 km and a core conductivity of .01/mho/m at
1000 km radius.

A plausible explanation for the difference between the
north-south and east-west transfer functions is that it is due
to a time varying compression of the remanent (dc) field at the
Apollo 12 site by fluctuations in the solar wind plasma. Provid-
ing that the spectrum of these compressive fluctuations is not
strongly frequency dependent, the result of removing this effect
will be to decrease slightly the estimated conductivity.

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TECHNICAL MEMORANDUM

INTRODUCTION

The Apollo 12 Lunar Surface Magnetometer Experiment
has produced some unexpected results, such as the observation
of a large remanent (dc) magnetic field, an inductive response
to time fluctuations in the interplanetary magnetic field that
is a function of the polarization of the fluctuations, and
apparent peaks in the observed response (transfer function)
to the time fluctuations [Dyal and Parkin, 1971; Sonett et al.,
1971 a, 1971 b]. On the interpretational side, the rapid rise
and then leveling off of the transfer function, as determined
from the lunar dayside data, has been taken as indicating a
thin, highly conducting shell at a depth of about 250 km
[Sonett et al., 1971 a, 1971 b], while the nightside data has
been interpreted in terms of a monotonically increasing con-
ductivity structure. In this paper I intend to pursue some
of the implications of these observations and determine their
effects on the interpretation of the data in terms of the
conductivity models.

Theoretical Model

Figure 1 shows the geometry on the dayside lunar
equator and in this coordinate system the field components
due to the interplanetary field and the induced fields are
[Sill and Blank, 1970]

$$B_x = B_{ox} \tag{1}$$

$$B_y = H_p B_{oy} - H_t (V_x * B_{oy}) + H_t (V_y * B_{ox}) \tag{2}$$

$$B_z = H_p B_{oz} - H_t (V_x * B_{oz}) \tag{3}$$

where all quantities are to be considered as functions of frequency, H_p and H_t are the Poloidal and Toroidal transfer functions, respectively, and the solar wind components are (neglecting aberation and fluctuations out of the ecliptic)

$$V_x = -V_{sw} \cos \phi \tag{4}$$
$$V_y = V_{sw} \sin \phi$$

The constraints imposed in the above derivation are [Blank and Sill, 1969; Sill and Blank, 1970], (1) the wavelength of the source field fluctuation is much greater than the radius of the moon, which, for spatial irregularities convected with the solar wind, requires that the frequency of the associated time variations be much less than 0.1 Hz, (2) the lunar conductivity profile is radially symmetric, (3) a thin current sheet at the lunar surface confines the induced fields to the lunar interior, (4) the day-night asymmetry in the solar wind plasma interaction with the moon, which results in the plasma void behind the moon, produces small effects on the induced fields on the sunlit hemisphere.

Model studies indicate that H_p is an increasing function of frequency in the frequency range where induction occurs in the more conductive portions of the lunar interior, while H_t is independent of frequency until induction occurs in the least conductive regions near the lunar surface, at which point H_t decreases with frequency [Sill and Blank, 1970].

Discussion of the Observations

Figure 2 shows the ratio of the power spectra ($A_i = |B_i| / |B_{oi}|$), i = x, y, z, (average of 14 spectra representing 21 hours of data) as determined from measurements made on the lunar surface by the Apollo 12 magnetometer and in the solar wind by Explorer 35 [Sonett et al., 1971 b]. The observed vertical transfer function (A_x) is near unity in accord with (1), but it shows a tendency to decrease at the higher frequencies, perhaps indicating that the approximation of an infinitely thin confining current sheet is no longer valid. The observed horizontal transfer functions (A_y, A_z) increase with frequency indicating a dominant contribution from the poloidal response. However, A_z is consistently

larger than A_y and this has been suggested as an indication of a contribution from the toroidal response [Sonett et al., 1971 a, 1971 b]. From equations 2 and 3, assuming that V_x, V_y have a dc component only and that H_t is real, the ratios of the power spectra are,

$$A_z^2 = |H_p|^2 - 2V_x H_t \text{ Real } [H_p] + (V_x H_t)^2 \tag{6}$$

$$A_y^2 = A_z^2 + (V_y H_t)^2 + 2V_y H_t \text{ Real } [(H_p - V_x H_t)$$

$$B_{ox}^* B_{oy} / B_{oy} B_{oy}^*] \tag{7}$$

Therefore $A_z > A_y$, if the sum of the last two terms in (7) is negative. Since H_t, H_p and $-V_x$ are positive, a necessary condition is that the product V_y Real $[B_{ox}^* B_{oy}]$ be negative. From (5) and Figure 1 we see that V_y is negative before local noon and positive after local noon. Observations of the x and y components of the interplanetary magnetic field show both positive and negative correlations, e.g., Figures 14a, 14b of Dyal et al., [1970]. Therefore, if we require $A_z > A_y$ by the above mechanism, those spectra averaged in Figure 2 which represent measurements made before local noon ($V_y < 0$) must be dominated by positively correlated events in the cross-spectrum (Real $[B_{ox}^* B_{oy}] > 0$) and those which represent measurements made after noon ($V_y > 0$) by negatively correlated events. These requirements seem like a rather restrictive set of conditions.

The last condition, that the magnitude of the last term in (7) be greater than the second, poses no serious problem, if the poloidal response is greater than the toroidal response and if the spectral ratio term is not too small. For H_p in the range from 1 to 3, and ($B_{ox}^* B_{oy} / B_{oy} B_{oy}^*$) about 1/2, values for $V_y H_t$ in the range from 1/2 to 2 are compatible with the observed differences between A_z and A_y. A test for the contribution from the toroidal interaction can be made by the analysis of data collected at local noon, for then A_z should equal A_y.

Alternatively, if the toroidal response is negligible, then the difference between the horizontal transfer functions could be attributed to a more complicated conductivity structure including variations in latitude and longitude, or to other asymmetric aspects of the interaction such as the day-night asymmetry.

Another possibility is a contribution to the lunar surface spectra from a variable compression of the local remanent (dc) field. The remanent field should interact with the incoming solar wind plasma in a fashion similar to that of the induced fields [Blank and Sill, 1969]. The gist of this interaction is that the deflection of the particles by the magnetic field gives rise to a current which excludes the field from the bulk of the plasma and confines it to a region near the lunar surface. The thickness of the current sheath is of the order of c/ω_p or about 5 to 10 km. Perturbations in the solar wind plasma parameters should then cause fluctuations in the interaction current and thereby give rise to a fluctuating magnetic field which is proportional to the remanent field.

Dyal and Parkin [1971] observed fluctuations in the 3 hour averages of the field components that are proportional to the remanent field and the energy density of the solar wind. They suggest that these are due to the compression of the remanent field by the solar wind. Assuming a mechanism of this type and neglecting the toroidal interaction, the field at the lunar surface would be

$$B_y = H_p B_{oy} + K B(dc)_y \tag{8}$$

$$B_z = H_p B_{oz} + K B(dc)_z \tag{9}$$

where K is the spectrum of the compression factor, which is a function of the solar wind parameters, and $B(dc)_y = 13\gamma$, $B(dc)_z = 25.6\gamma$ are the horizontal components of the remanent field at the Apollo 12 site. Since the remanent z component is about twice as large as the y component, this mechanism will contribute proportionately more power to the z spectra, thereby providing an explanation for the observation that A_z is greater than A_y. In order for the remanent field fluctuations to be effective, their amplitude should be of the same order as the inductive fluctuations or about several gammas.

Fluctuations of this order are observed in the three hour averages of the difference between the lunar surface field and the interplanetary magnetic field [Dyal and Parkin, 1971]. For the case where the confining current layer thickness, remanent field scale size and distance from the source are all of the same order (10 km), a ten percent fluctuation in the solar wind plasma density will give rise to a fluctuation of similar magnitude in the confined remanent field or a fluctuation of several gammas. Therefore, such a mechanism would appear quite plausible.

Inspection of the individual spectra, as measured in the solar wind and on the lunar surface, [Sonett et al., 1971 b] indicates that for the surface fields $|B_z|^2 > |B_y|^2$ over the whole frequency range, while in the solar wind $|B_{oz}|^2 < |B_{oy}|^2$ for $f < 10^{-2}$ Hz and $|B_{oz}|^2 \approx |B_{oy}|^2$ for $f > 10^{-2}$ Hz. Therefore, the relative contribution from fluctuating compression is greater at the lower frequencies. In any case, A_y is less contaminated by these effects simply because the y component of the remanent field is less than the z component.

Finally, some explanation must be offered for the apparent peaks in the observed transfer functions as they are not consistent with the proposed theoretical model. One possible explanation is that they are the result of the excitation of higher order modes by the shorter wavelengths of the high frequencies. As was pointed out previously, the approximation of a uniform source field limits the application of the theoretical model to frequencies much less than 0.1 Hz. For the convection of spatial fluctuations in the interplanetary field, the phase difference across the moon is as least 15° at 10^{-2} Hz and it increases to at least 45° at 3×10^{-2} Hz. Certainly the approximation of a uniform source field is not applicable at these high frequencies. On the other hand, there is no evidence that the excitation of higher order modes will lead to a peaked response. Consideration of a "resonant" cavity effect leads to improbably large length scales or extremely low electromagnetic wave velocities.

With respect to these peaks in the observed transfer function, it is important to note that the individual spectra, in the solar wind as well as on the lunar surface, exhibit peaks near and above 10^{-2} Hz [Sonett et al., 1971 b] and that there is even a suggestion that the peaks are harmonically related. Peaks in the spectra of the interplanetary magnetic field at these frequencies have also been observed by earth

orbiting spacecraft [Fairfield, 1969; Scarf et al., 1970]
The fluctuations responsible for these peaks are observed
upstream from the earth's bowshock when the field line from
the spacecraft intersects the shock. It has been proposed
that these fluctuations are waves generated through a plasma
instability by high velocity protons traveling upstream from
the shock. Correlations between the magnetic field fluctuations
and suprathermal protons as well as other solar wind parameters
have been observed [Scarf et al., 1970].

Such fluctuations in the solar wind parameters could
lead to peaks in the lunar surface spectrum, through a variable
compression of the remanent field, if the spectrum of the com-
pression factor in (8) and (9) is peaked. Peaks in the spectrum
of the solar wind velocity at frequencies near 10^{-2} Hz have
been reported [Coleman, 1966]. As noted previously, modest
changes in the solar wind parameters ($\sim 10\%$) could lead to com-
pressional fluctuations of a few gammas.

A priori, it is also possible that the toroidal
interaction, which is dependent on the convolution of $B_o(\omega)$
and $V_{sw}(\omega)$, could contribute to the peaks in the spectra. In
this case the convolution of the spectra would give rise to
peaks at the sum and difference frequencies for fluctuations
in B_o and V_{sw} which are correlated. Thus, this mechanism is
capable of introducing additional power into different portions
of the surface spectra and this characteristic could be used
as a diagnostic feature of this interaction. Measurements
of the solar wind spectrum [Coleman, 1968] show that typical
fluctuations in the frequency range from 10^{-3} to 10^{-2} are of
the order of 10 km/sec. Since the toroidal contribution from
the steady (dc) component of the wind is at least an order of
magnitude greater than the fluctuations at these frequencies,
it seems less likely that the above mechanism could produce
an important effect.

Conductivity Models

With the above discussion of the data and the
theoretical model in mind, let us examine several conduc-
tivity models which purport to fit the observed transfer
function. Figure 3 shows a comparison of three poloidal
response models to the observed RMS transfer function,
$\langle A \rangle = [.5(A_y^2 + A_z^2)]^{1/2}$. The simple core-crust model (2)
provides an adequate fit to the data except at the highest
frequencies where the response is still increasing with
frequency, while the observed response has flattened out.

As has been mentioned several times previously, this high frequency region is where the assumption of a uniform source field breaks down and perhaps because of this (and the peaks), we should use caution in constraining the models to fit the data in this region, at least until we have a better understanding of the nature of the interaction in this frequency range. Sonett et al., [1971 b] attempted to fit a core-crust model (core radius = 1560 km, core conductivity = 7.4 x 10^{-4} mho/m) to the data, but the calculated response did not show very good agreement. Apparently, some of the problem is attributable to a computation technique which leads to an overestimate of the high frequency response. This effect can also be seen by comparing models 3 and Sonett's best fit model (1) in Figure 3. Here we note that the response of model 3, even with its slightly higher conductivity exterior to the peak, is lower and flatter than Sonett's model at the higher frequencies.

In Figure 4, models 4 and 5 show some of the effects of changing the conductivity profile near the peak. Model 4 has a broader and less conductive shell near 1500 km as compared to model 3. The response of model 4, compared to 3, is slightly greater at frequencies above 10^{-2} Hz, which is partly attributable to the higher conductivity in the region between 1525 km and 1550 km. This is partly in contradiction of the statement in Sonett et al., [1971 b], that a better fit of the model at the higher frequencies will probably lead to both a higher value of the maximum conductivity and a steeper slope of the conductivity profile at the outer edge. Model 5 shows that the transfer function is not very sensitive to the structure just below the peak; here the increase in the conductivity between 1200 km and 1450 km is compensated by a slight thinning of the highly conducting shell. Figure 5 pursues this effect. In these models the thinning of the shell is compensated by an increase in the shell and core conductivity.

These models illustrate something of the permissible range of models that will fit the data. If we require the response to be flat at the higher frequencies, all the models exhibit a peak in the conductivity near 1500 km. However, there is some spread in the models and the conductivity profile just below the peak is not strongly limited. If we relax the constraint on the flatness of the high frequency response, a simple core-crust model will fit the low frequency data.

So far the models have been fit to the observed RMS transfer function, but in the previous discussion of why $A_z > A_y$, two of the mechanisms (toroidal interaction and compression of B(dc)) were based on effects which would channel

more extraneous power into the z component spectra at the lunar surface. In both these cases the A_y transfer function should be closer to the pure poloidal transfer function. Figure 6 is a fit of a model to A_y and A_z - .4 where .4 is the average difference between A_z and A_y. Model 6 is similar to models 3 and 5 with a slightly thinner and less conductive shell and a less conductive core. The core-crust model 7 is slightly less conductive than that in Figure 3.

If the difference between A_z and A_y is due to contamination of the spectra from either a weak toroidal interaction or a variable compression of the remanent field, then the appropriate models will be much like those in Figures 3-6 with slightly greater depths of the layers and slightly reduced conductivities. However, the picture could be altered significantly if the proposed contamination mechanisms are strongly frequency dependent, as seems more likely for the compression of the remanent field. Some of these questions about the mechanism could be answered by a careful examination of the data, especially at low frequencies, and also by measurements that will be made at future Apollo landing sites.

The analysis of the nighttime data, when the magnetometer is exposed to the plasma void behind the moon, has proceeded by matching the observed time domain response to models of the moon represented by a sphere in a vacuum [Dyal et al., 1970, Dyal and Parkin, 1971]. The response of a sphere in a vacuum to a step function change in the source field is characterized by a decay from the initial to the final value for the radial component and by an overshoot and then a decay to the final value for the tangential components. Qualitatively such a response is observed in the nighttime data; however, deviations from the theoretical response of a sphere in a vacuum are observed. Some of these effects are illustrated in Figure 7 [Dyal et al., 1970]. Note that the step change in the tangential field as observed at Explorer 35 results in an overshoot at the lunar surface which is more than twice the magnitude of the source field change. Also, the radial component (x) at Explorer 35 shows no pronounced low frequency change at the time of the step change in the tangential components, but at the lunar surface there is a large transient in the radial component.

A model consisting of a sphere with confining currents distributed along a tangential cylinder, representing the boundary of the plasma void, would probably be a more appropriate model for the nighttime response. Such a model would account for the amplification greater than 1.5 and the crosstalk between components by the compression and distortion of the induced dipole-like field.

The absence of the diamagnetic solar wind plasma on the night side could also contribute to an observed amplification greater than 1.5. In this case, suppose that the source field in the void is about 30% greater than in the solar wind because of the plasma diamagnetism, the initial overshoot could be nearly twice as great as the magnitude of the step as observed in the solar wind. Another effect would be that the decay would asymptotically approach a value some 30% greater than that seen in the solar wind. Since the field is rarely steady for long enough periods, it is difficult to test this effect. In any case the crosstalk between components and the suggestion that the amplification may be even larger than 2 at the higher frequencies favors partial compression and distortion as a dominant mechanism.

Figure 8 shows the calculated time response of layered sphere models in a vacuum to a simulation of an observed transient [Dyal et al., 1970] in the radial component and Figure 9 shows the frequency domain response. In the time domain, the calculated response of the core-crust model suggested by Dyal and Parkin [1971] is larger than the observed response for all times shown. The 3-layer model of Dyal and Parkin [1971] has a highly conductive core and its response is larger than the observed one for the first 200 sec, but at longer time it matches the observed surface response. Model 3, after Sonett's conductive shell model, fits the first 200 sec of the observed response rather well, but it is larger than the observed for later times. A model (3c) composed of model 3 with a highly conductive core as in the 3-layer model of Dyal and Parkin seems to give a very good fit to the surface response over the complete time span shown, as does model 2c which is like the model 2 with the addition of a conductive core. The fact that the calculated response is below the input response at 300 sec indicates that induced currents are still flowing in the conductive interior and Figure 9 shows that these models have the largest response at the low frequencies.

For comparison Figure 10 shows the observed and calculated transients in the z component which occurred at the same time as the x component transient shown in Figure 8. Here, none of the models provides an adequate fit to the observation. The obvious discrepancy is that the magnitude of the observed surface response is much greater than the calculated response. The initial value of the observed overshoot ($\sim14\gamma$) is about twice as large as the step change in the solar wind field ($\sim7\gamma$), while the calculated overshoot is about what would be expected from Figures 9 and 10, (i.e. $1.3 \times 7\gamma \cong 9\gamma$). This discrepancy is probably due to partial confinement effects.

As a final example Figure 11 shows another observed transient in the radial component and the response of the same five models. During the first 300 sec all the models give a response that is greater than the observed one. The models with the least conductive interiors (core-crust and 3) have as expected the largest response at 300 sec, however, the decay of the conductive shell model (3) after 300 sec is much slower than the core-crust model. The 3-layer model has a very rapid decay after 300 sec when compared to the core-crust model because the currents induced in the innermost core during the first 300 sec very nearly cancel the current induced by the drop in the field at 300 sec. Here, again, the conductive shell model with a highly conductive core (3c) and model 2c provide a better, but not wholely adequate, fit to the observed response.

Summary

For a completely dominant poloidal interaction, a simple core-crust model provides an adequate fit to the observed dayside RMS transfer function except at the highest frequencies. The requirement that the calculated poloidal response match the relatively flat observed response in the high frequency region leads to models which have a peak in the conductivity profile near 1500 km. However, examination of several models indicates that the conductivity structure below this depth is not strongly limited by the present dayside data. Extension of the frequency domain analysis to lower frequencies will be required to determine the deeper conductivity structure and will aid in the clarification of other questions such as the importance of the toroidal interaction.

It seems plausible that the difference between the observed horizontal transfer functions (i.e., $A_z > A_y$) is due to a time variable compression of the remanent field driven by fluctuations of the solar wind plasma. In such a case the amplitude of the fluctuation is proportional to the strength of the remanent field and, therefore, the observed transfer function A_y is closer to the true inductive transfer function simply because the magnitude of the y component of the remanent field is smaller than the z component. The models that are fit to the y transfer function are slightly less conductive than those that are fit to the RMS transfer function. Provided that the spectrum of the compressive fluctuations is not a strong function of frequency, we anticipate that the complete removal of the compressive effects will lead to models slightly less conductive than those that have been previously presented. If the compressive fluctuations of the remanent field are strongly frequency dependent, then the resulting models could be signif-icantly different. (Preliminary results from the Apollo 15

magnetometer indicate a very small remanent field, <5γ [Dyal, 1971, personal communication]. Contamination of the induction field spectra by compression of the remanent field should be minimum at this location.)

During the lunar night the data is less likely to be contaminated by compression of the remanent field because of the absence of the solar wind plasma at the lunar surface. However, the partial confinement of the induction field within the plasma void limits the utility of the sphere-in-a-vacuum model for the nightside response. Within the limitation of this model, it is found that either Sonett's conductive shell model with the addition of a highly conductive core (model 13c) or a monotonic conductivity model (2c) provide the best (but not completely adequate) fit to the nightside radial component transient data. Inclusion of the conductive core in either of these models does not appreciably alter the dayside transfer function above 10^{-3} Hz, so that these models are compatible with the present dayside data as well.

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Attachments

REFERENCES

Blank, J. L., and W. R. Sill, Response of the Moon to the time-varying interplanetary magnetic field, J. Geophys. Res., 74, 736, 1969.

Coleman, P. J., Jr., Variations in the interplanetary magnetic field: Mariner 2, J. Geophys. Res., 71, 5509, 1966.

Coleman, P. J., Jr., Turbulence, viscosity and dissipation in the solar wind plasma, Astrophys. J., 153, 371, 1968.

Dyal, P., C. W. Parkin, C. P. Sonett, and D. S. Colburn, Electrical conductivity and temperature of the lunar interior from magnetic transient measurements, NASA TM X-62, 012, 1970.

Dyal, P., and C. W. Parkin, The Apollo 12 magnetometer experiment: Internal lunar properties from transient and steady magnetic field measurements, Proceedings of the Apollo 12 Lunar Science Conference, (in press) The M.I.T. Press, 1971.

Fairfield, D. H., Bow shock associated waves observed in the far upstream interplanetary medium, J. Geophys. Res., 74, 3541, 1969.

Scarf, F. L., R. W. Fredricks, L. A. Frank, C. T. Russel, P. J. Coleman, Jr., and M. Neugebauer, Direct correlations of large amplitude waves with suprathermal protons in the upstream solar wind, J. Geophys. Res., 75, 7316, 1970.

Sill, W. R., and J. L. Blank, Method for estimating the electrical conductivity of the lunar interior, J. Geophys. Res., 75, 201, 1970.

Sonnett, C. P., D. S. Colburn, P. Dyal, B. F. Smith, G. Schubert, and K. Schwartz, Lunar electrical conductivity profile, Nature, 230, 359, 1971a.

Sonett, C. P., G. Schubert, B. F. Smith, K. Schwartz, and D. S. Colburn, Lunar electrical conductivity from Apollo 12 magnetometer measurements: Compositional and thermal inferences, Proceedings of the Apollo 12 Lunar Science Conference, (in press) The M.I.T. Press, 1971b.
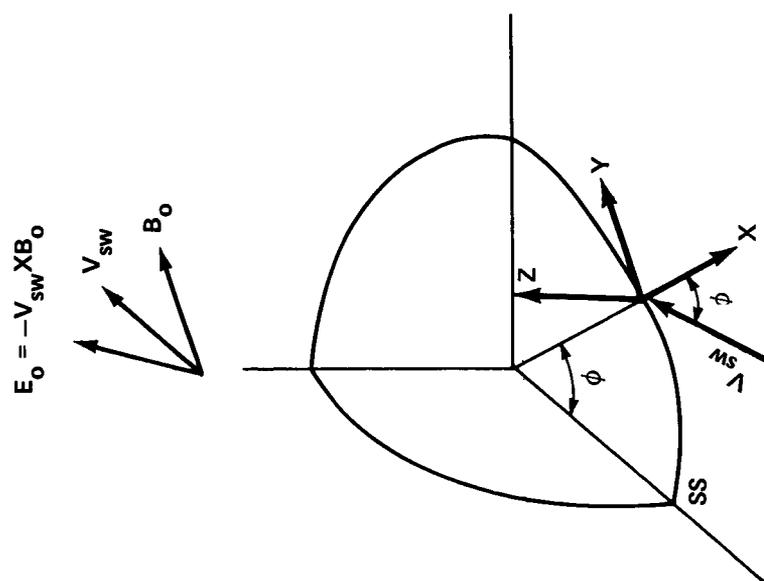
FIGURE 1 - QUADRANT OF THE MOON SHOWING THE INTERPLANETARY SOURCE FIELDS B_0, E_0, V_{sw} AND A LOCAL COORDINATE SYSTEM ON THE LUNAR SURFACE AT THE EQUATOR, X(VERTICAL), Y (EAST), Z (NORTH). SS IS THE SUBSOLAR POINT AND ϕ IS THE ANGLE BE-TWEEN THE DIRECTION TO THE SUN AND THE VERTICAL.
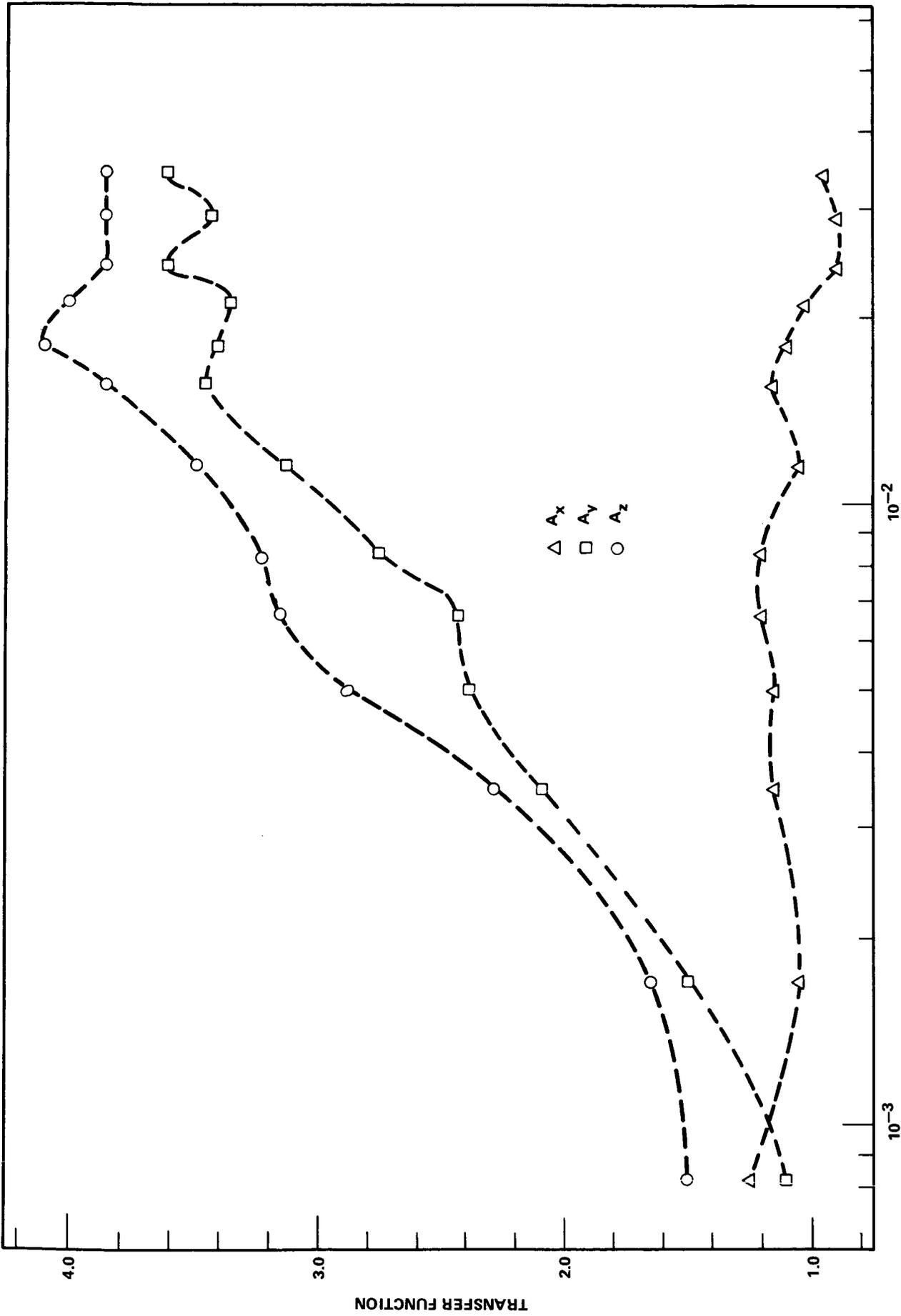
FIGURE 2 - THE EXPERIMENTAL TRANSFER FUNCTIONS A_x (VERTICAL), A_y (EAST) AND A_z (NORTH) [SONETT ET AL., 1971b].
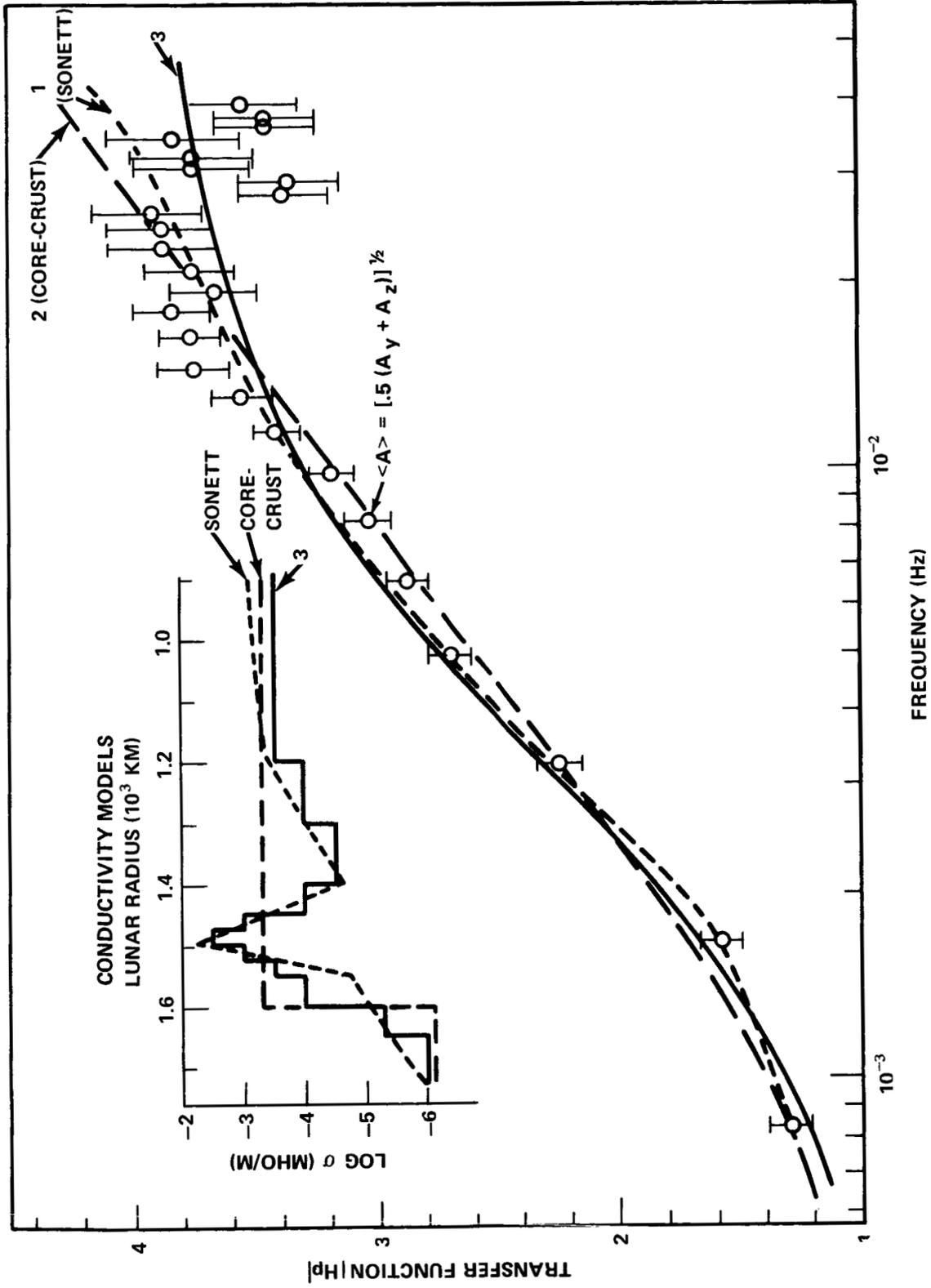
FIGURE 3 - THE EXPERIMENTAL RMS TANGENTIAL TRANSFER FUNCTION SHOWN AS CIRCLES WITH ERROR BARS FROM SONETT ET AL. [1971b]. ALSO SHOWN ARE THE CALCULATED POLOIDAL TRANSFER FUNCTIONS FOR THE CONDUCTIVITY MODELS SHOWN IN THE INSET.
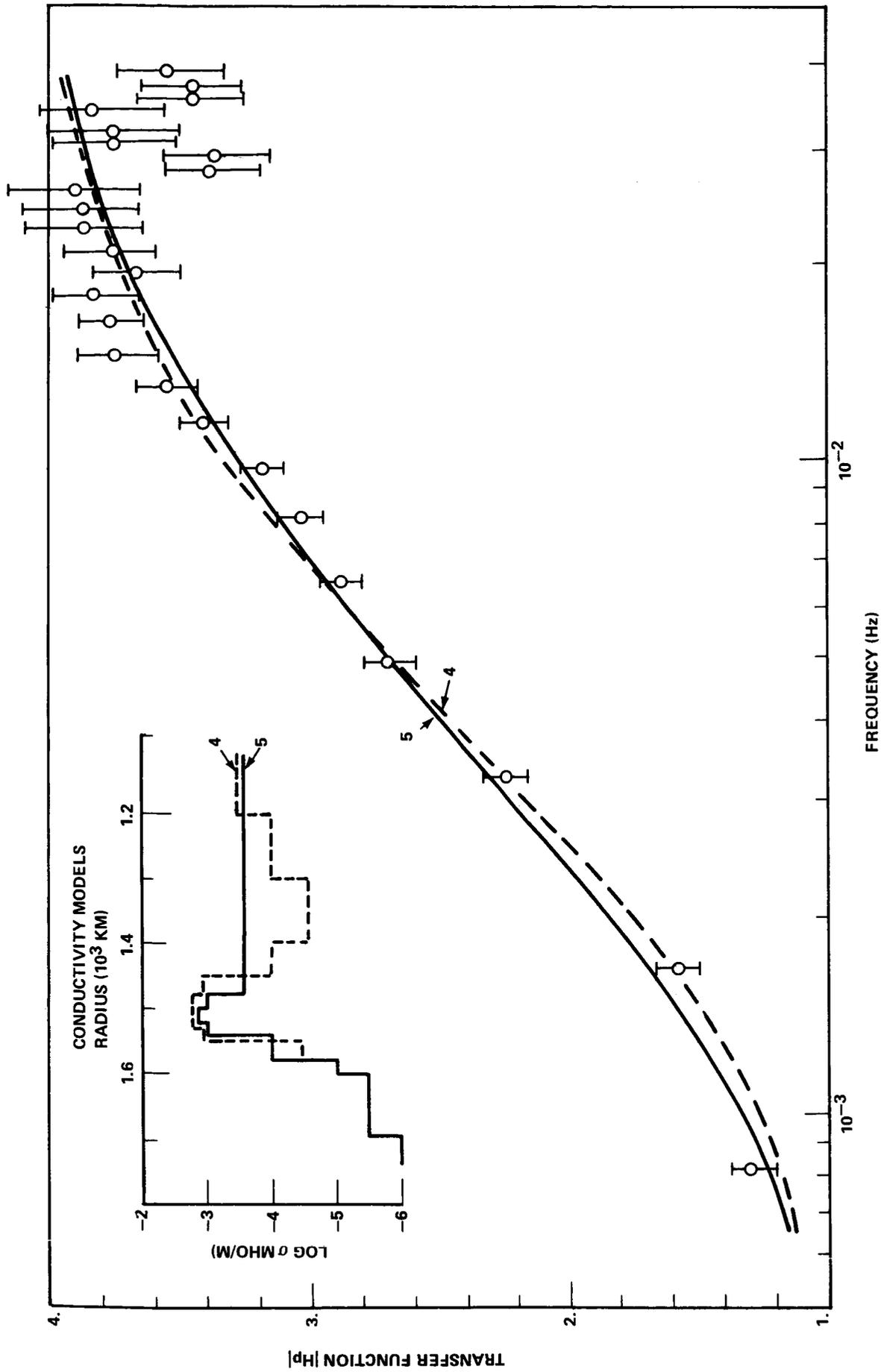
FIGURE 4 - THE EXPERIMENTAL RMS TANGENTIAL TRANSFER FUNCTION SHOWN AS CIRCLES WITH ERROR BARS AND THE CALCULATED POLOIDAL TRANSFER FUNCTIONS FOR THE CONDUCTIVITY MODELS SHOWN IN THE INSET.
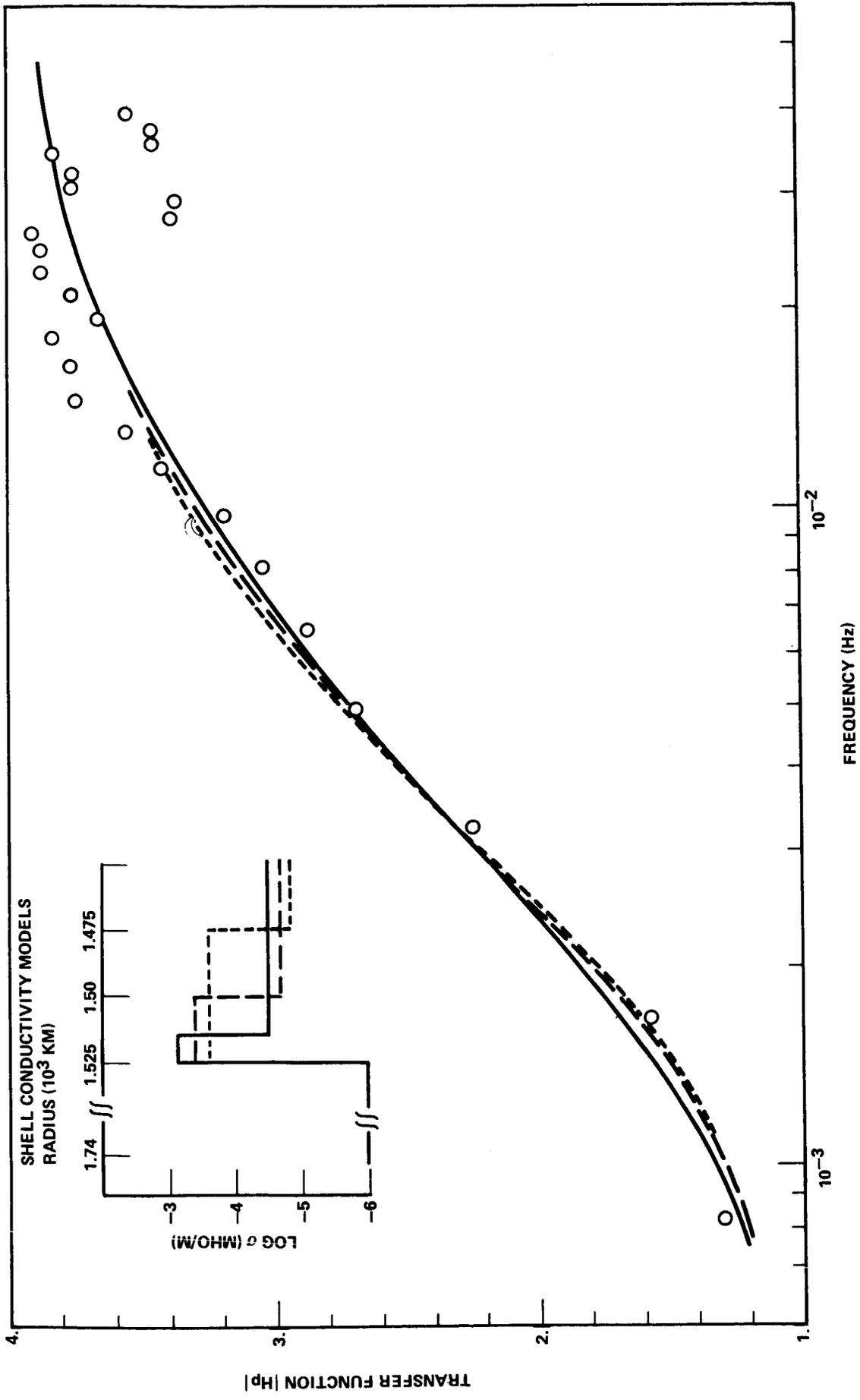
FIGURE 5 - THE EXPERIMENTAL RMS TANGENTIAL TRANSFER FUNCTION SHOWN AS CIRCLES AND THE CALCULATED POLOIDAL TRANSFER FUNCTIONS FOR THE CONDUCTIVITY MODELS SHOWN IN THE INSET.
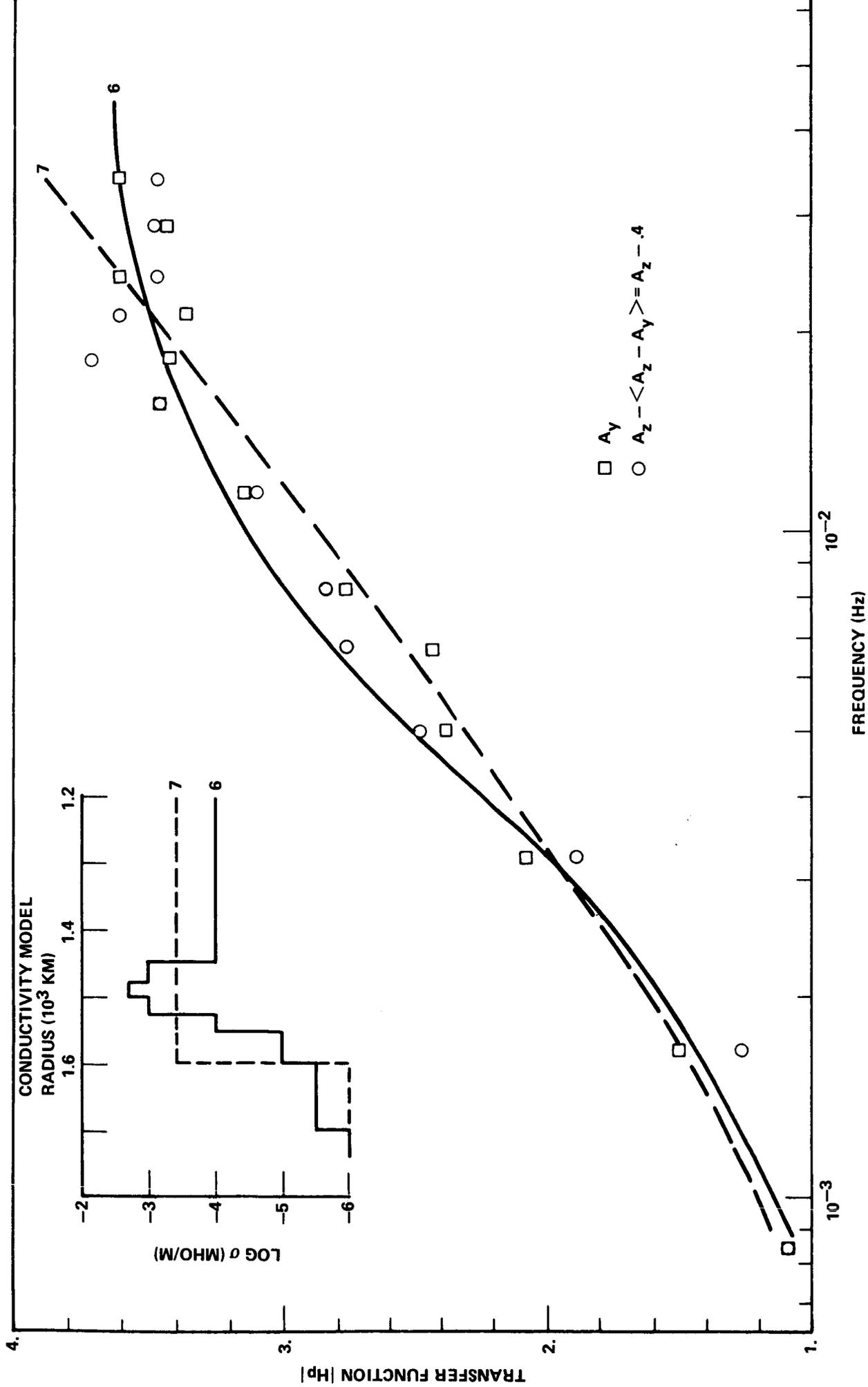
FIGURE 6 - THE Y COMPONENT OF THE EXPERIMENTAL TRANSFER FUNCTION AND THE Z COMPONENT MINUS THE AVERAGE DIFFERENCE (.4) BETWEEN THE Z AND Y COMPONENT TRANSFER FUNCTIONS. ALSO SHOWN ARE THE CALCULATED POLOIDAL TRANSFER FUNCTIONS FOR THE CONDUCTIVITY MODELS SHOWN IN THE INSET.

☐ A_y

○ $A_z - \langle A_z - A_y \rangle = A_z - .4$
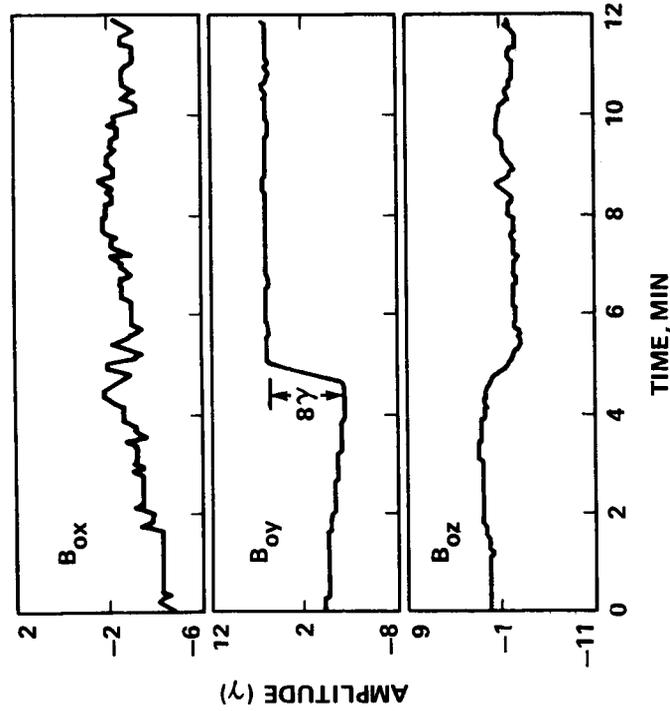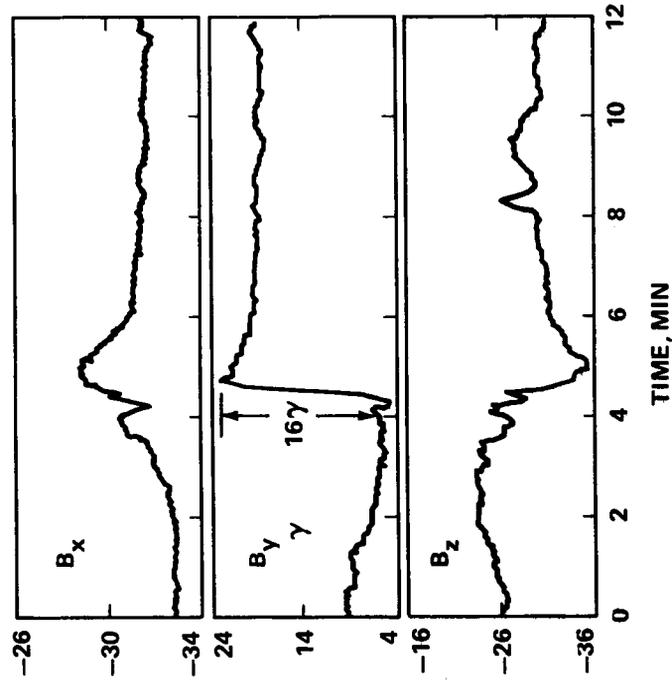
FIGURE 7 - THE LUNAR NIGHTSIDE TRANSIENT RESPONSE. AT THIS TIME EXPLORER 35 MEASURED THE FIELD IN THE SOLAR WIND WHILE APOLLO 12 WAS EXPOSED TO THE PLASMA VOID BEHIND THE MOON.
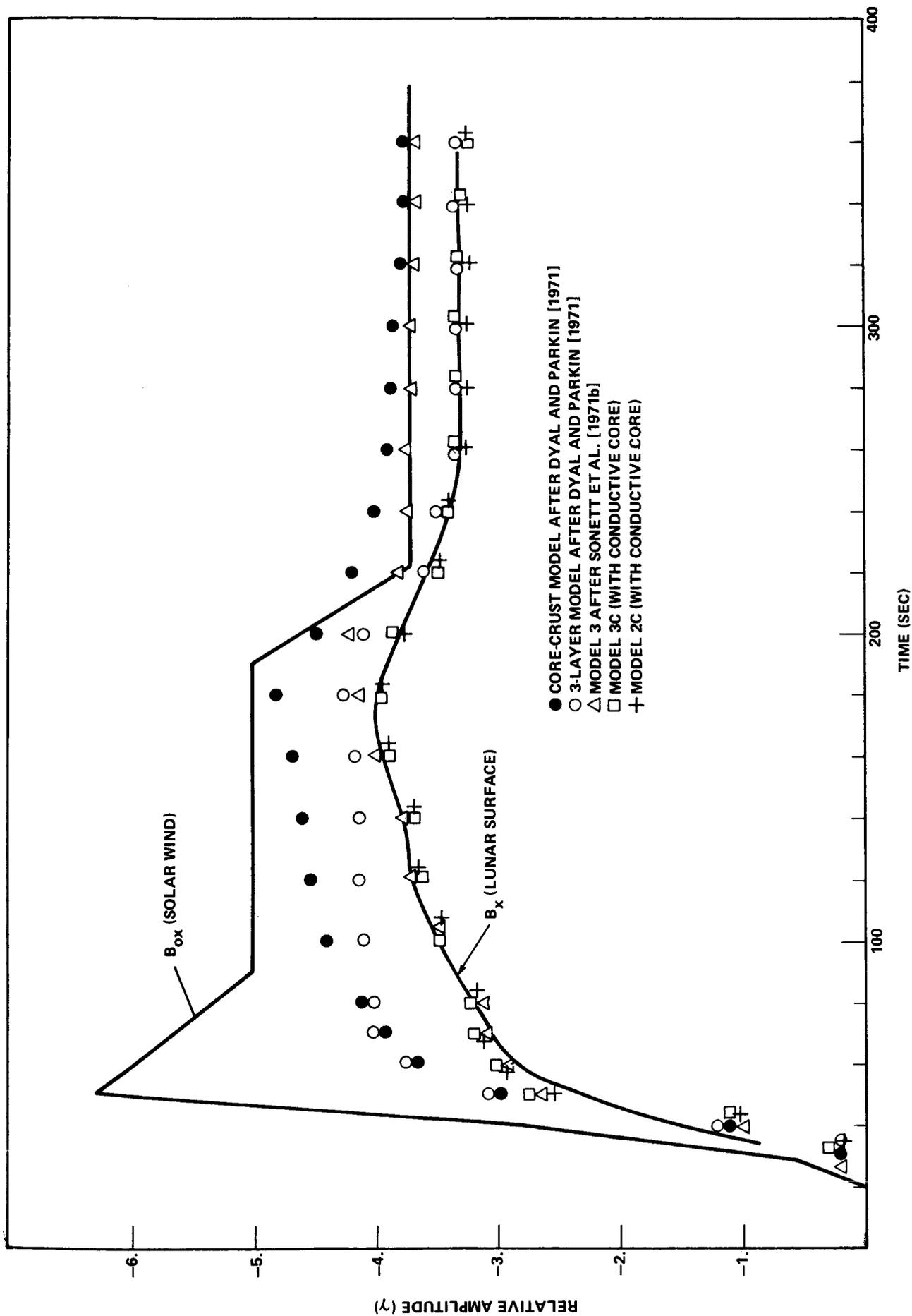
FIGURE 8 · THE LUNAR NIGHTSIDE TRANSIENT RESPONSE AND THE CALCULATED TIME DOMAIN RESPONSE FOR THE CONDUCTIVITY MODELS SHOWN IN FIGURE 9. B_{ox} IS A SIMULATION OF THE SOURCE FIELD AS OBSERVED IN THE SOLAR WIND AND B_x IS THE OBSERVED LUNAR SURFACE FIELD [DYAL ET AL., 1970].
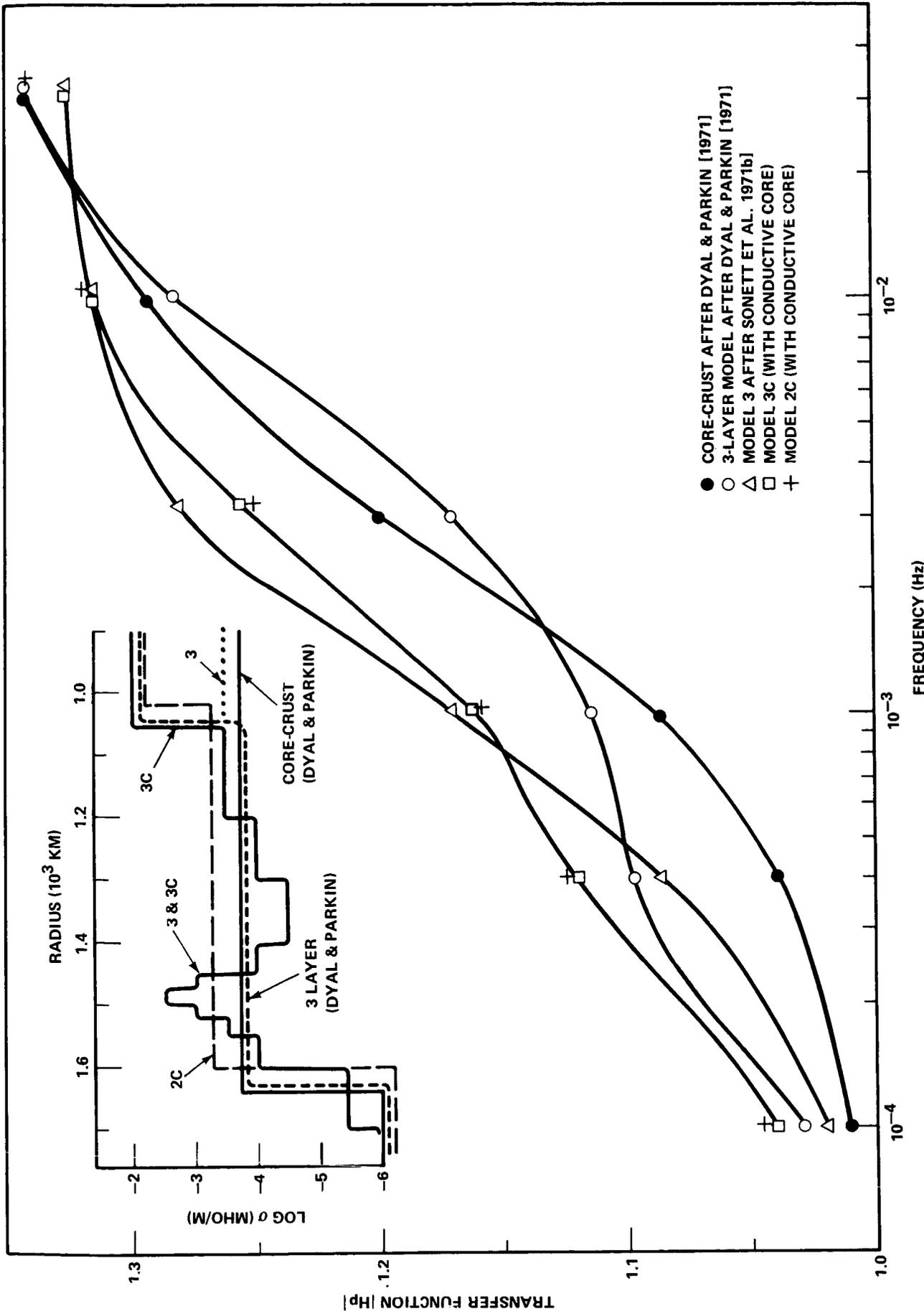
FIGURE 9 - THE CALCULATED FREQUENCY DOMAIN RESPONSE FOR A SPHERE IN A VACUUM. THE CONDUCTIVITY MODELS ARE SHOWN IN THE INSET.

Legend:
● CORE-CRUST AFTER DYAL & PARKIN [1971]
○ 3-LAYER MODEL AFTER DYAL & PARKIN [1971]
◁ MODEL 3 AFTER SONETT ET AL. 1971b]
□ MODEL 3C (WITH CONDUCTIVE CORE)
+ MODEL 2C (WITH CONDUCTIVE CORE)

FREQUENCY (Hz)

TRANSFER FUNCTION |Hp|

Inset axes:
RADIUS (10³ KM)
LOG σ (MHO/M)

Inset labels: 3, 3C, 3 & 3C, 2C, CORE-CRUST (DYAL & PARKIN), 3 LAYER (DYAL & PARKIN)
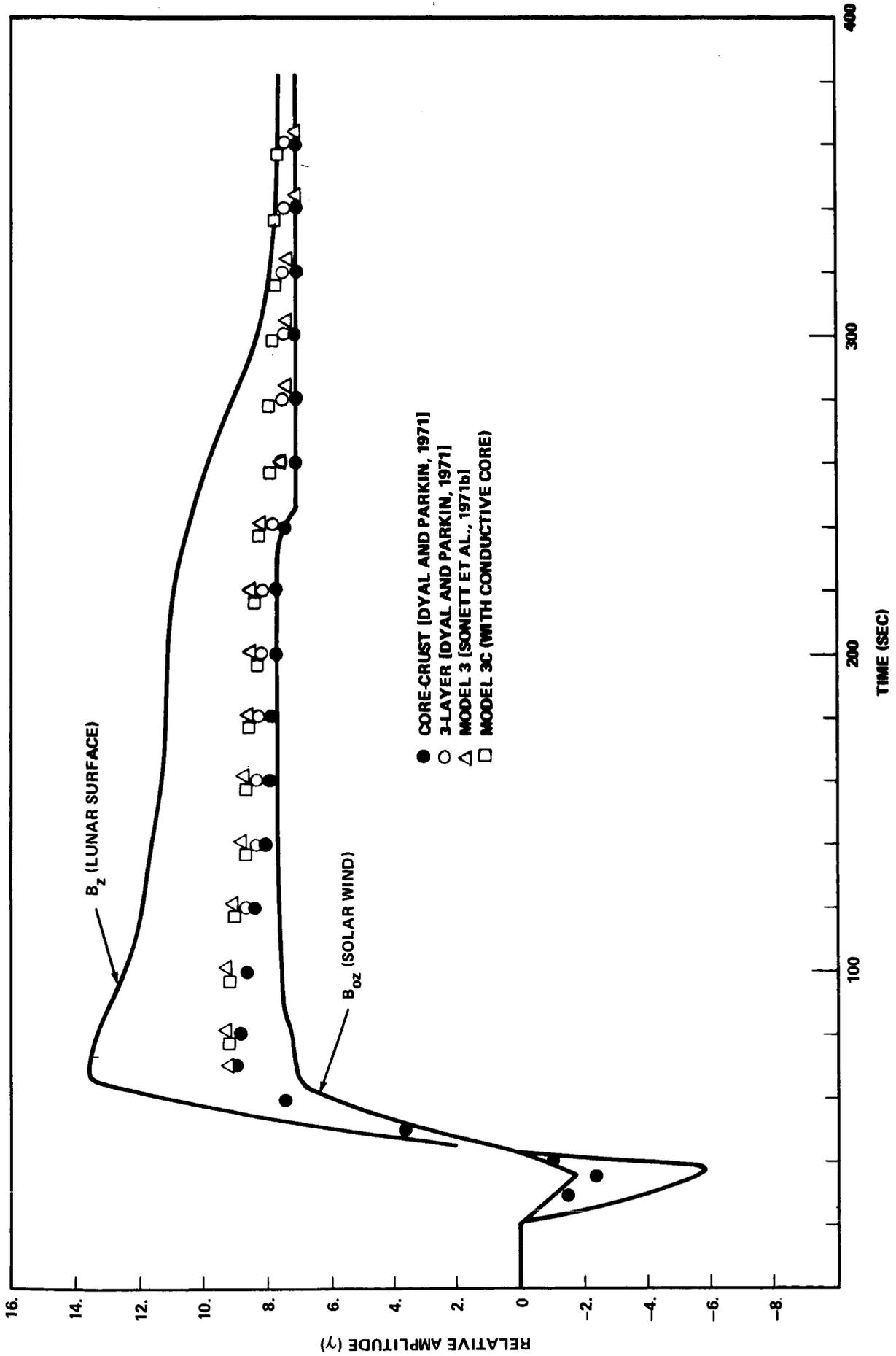
FIGURE 10 - THE LUNAR NIGHTSIDE TRANSIENT RESPONSE AND THE CALCULATED TIME DOMAIN
RESPONSE. THIS EVENT OCCURRED AT THE SAME TIME AS THE EVENT SHOWN IN
FIGURE 8.
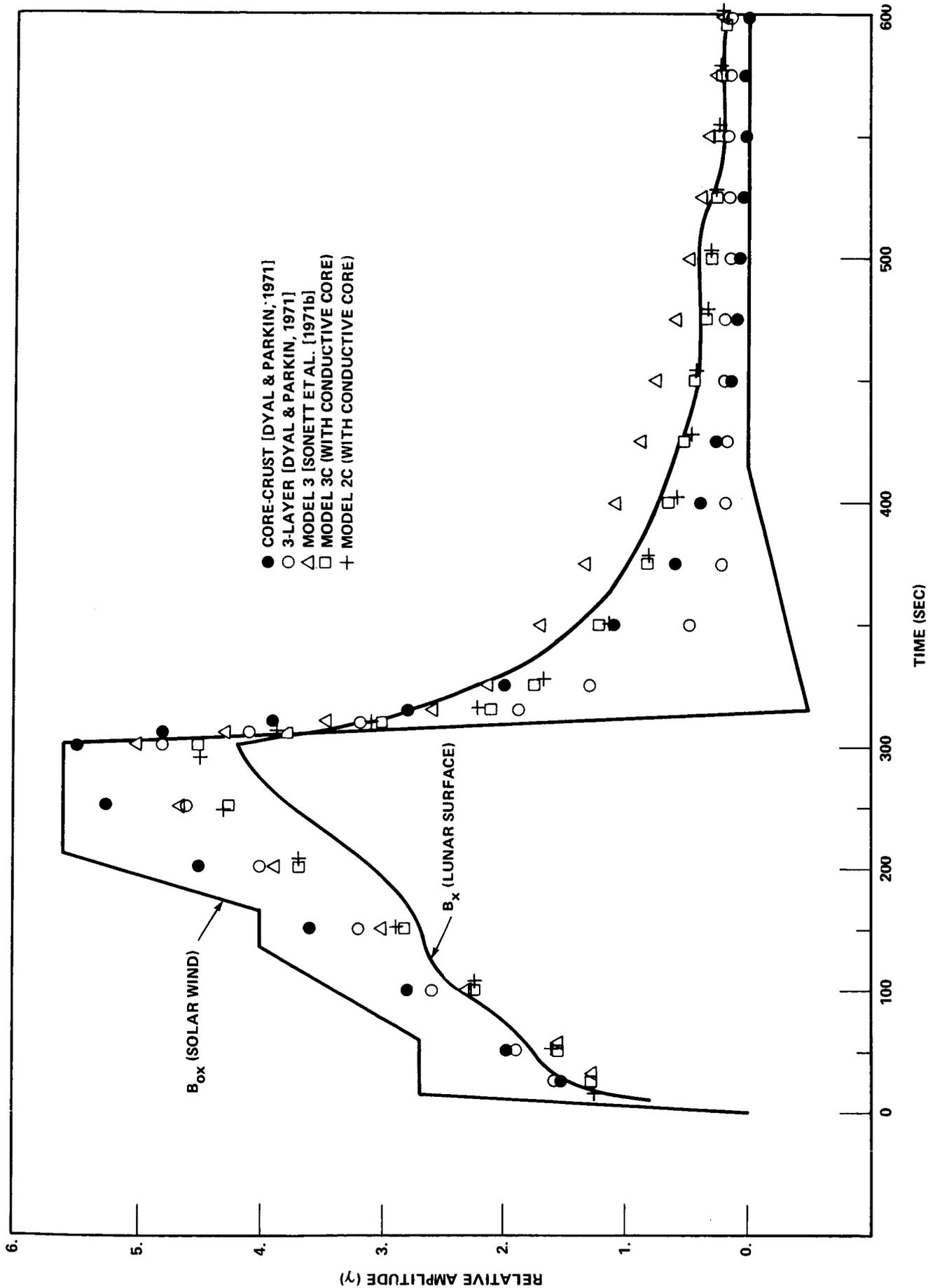
FIGURE 11 - THE OBSERVED NIGHTSIDE TRANSIENT RESPONSE [DYAL AND PARKIN, 1971] AND THE CALCULATED TIME DOMAIN RESPONSES OF THE MODELS IN FIGURE 9.