

An Analytic Formula for Heating Due to Ozone Absorption

Richard S. Lindzen^{1,2} and Douglas I. Will

University of Chicago

Chicago, IL 60637

(NASA-CR-132778) AN ANALYTIC FORMULA FOR HEATING DUE TO OZONE ABSORPTION (Chicago Univ.) 13 p HC \$3.00 CSCL 04A N73-26382

Unclas G3/13 08475

August, 1972



Now at Harvard University

²Alfred P Sloan Foundation Fellow

From the tropopause to about 80 km, the major source of radiative heating in the atmosphere is the absorption of sunlight by ozone. This absorption occurs in the Hartley (2000 Å - 3000 Å), Huggins (above 3000 Å) and Chappuis (5000 Å - 7000 Å) bands (Craig, 1965, p. 168; also Craig, 1950) and is only slightly temperature and pressure dependent. Consequently, to a high degree of approximation one may write

I_ = I_e

 $Q = \int \kappa \prod_{\nu \in \mathcal{V}} n d\lambda$

where

 ${\tt where}$

and

 $Q = heating in ergs/sec cm^3$

$$u = \int_{\infty}^{z} ndz/cos$$

u = optical depth in cm NTP

 ∂ = zenith angle

n = ozone density in (cm NTP/cm)

 κ_{v} = absorption coefficient (cm NTP)⁻¹ $I_{v\infty}$ = insolation (incident solar intensity) (ergs/cm² sec Å) I_{v} = intensity (ergs/cm² sec Å)

(1)

(2)

z = height (cm)

⁴ This expression is appropriate only for a flat atmosphere and must be modified to take sphericity into account. Distributions of I woo and K are shown in Figs. 1-3. The quantity $Q/n \equiv \eta$ is solely a function of u, and for the choice of spectral data shown in Figs. 1-3 leads to the distribution of η with u shown in Fig. 4. As noted in Lindzen and Goody (1965) there are two regions of the atmosphere where η may be approximated by a constant:

- i) above about 4-5 km where u (2x10⁻³ cm NTP we are in a region which is almost transparent for all ozone bands.
 Hence, ozone band radiation is not significantly attentuated and the heating is largely due to ozone's strongest band, the Hartley band
- ii) below about 30 km where 0.3 cm NTP \rangle u \rangle 0.15 cm NTP, most radiation in the Hartley and Huggins bands has been absorbed but the atmosphere is still almost transparent for the Chappuis which dominates the heating.⁵

The question now arises as to whether a simple expression exists which describes both regions of "constant" η as well as the transition between them. We, moreover, wish a convenient way of relating parameters in the expression to spectral data. Now, in the regions of constant η , η depends only on average values of spectral intensities and absorption coefficients for the Hartley and Chappuis bands. Thus a simple solution might be to use average values for I and κ , 5N. B. Figure 4 is misleading on this matter since it includes values of u which greatly exceed the atmosphere's total ozone content; such values could be important, however, in time dependent calculations where Θ may approach $\pi/2$. appropriate to the Hartley and Chappuis bands, in (1) and (2) modelling both bands as simple Chapman layers (Craig, 1965, pp. 147-150). One then obtains for η the following expression

$$\eta = I_{H} \kappa_{H} \Delta \lambda_{H} e^{-\kappa_{H} u} + I_{c} \kappa_{c} \Delta \lambda_{c} e^{-\kappa_{c} u}$$
(3)

where I κ_{H} and I, κ_{c} are average I, κ_{v} for Hartley (H) and Chappuis (c) bands and Δk_{H} , Δk_{c} are the bandwidths taken for these bands. For the range of u's relevant to the earth's atmosphere, κ_{H} , I_{H} , κ_{c} and I_{c} can be chosen to accurately model the regions of constant n; (3) will then also describe the transition between the two regions. Unfortunately, the accuracy of such an approximation proved poor in the transition zone because of the importance of the Huggins bands in that zone. Modelling the Huggins bands proved reasonably simple. From Fig. 2 we see that κ_{v} is almost a linear function on a semi-log plot. Thus we may write

$$\kappa_{v} = \kappa_{Hu} e^{-M\lambda}$$

(4)

 κ_{Hu} , M are constants of the straight line on the semi-log plot. Taking for I, an average over the Huggins hands, I_{Hu}, we can easily integrate (2) obtaining



where λ long and λ short are the wavelength limits used for the Huggins bands.

Thus an improved approximation to η consists in the sum of (3) and

(5), namely





Expression (6) has, in principle, a substantial number of adjustable parameters: the widths of the Hartley, Huggins and Chappuis bands, I_H , I_c , I_{Hu} , and κ_H and κ_c . If one merely estimates these parameters on the basis of Figs. 1-3, we obtain a uniform 20 percent accuracy in approximating Fig. 4. However, with a little experimentation we were able to choose these parameters so as to obtain the approximation shown in Fig. 4 which, in general, has about a 5 percent accuracy. The parameters chosen are shown in Figs. 1-3. The following formulae should prove useful when dealing with more recent or better data than that shown in Figs. 1-3. Hartley bands:

$$I_{H} = 9 \text{ ergs/cm}^{2} \text{ sec } \text{\AA}$$

$$\kappa_{H} = 260 \text{ cm NTP}^{-1} = 80\% (\kappa_{v} \text{ maximum for Hartley})$$

$$\Delta\lambda_{H} = 375 \text{ \AA}$$

Chappuis bands:

$$I_{c} = 180 \text{ ergs/cm}^{2} \text{ sec } \text{\AA}$$

$$\kappa_{c} = 0.118 \text{ cm NTP}^{-1} = 80\% (\kappa_{v} \text{ maximum for Chappuis})$$
(8)
$$\Delta \lambda_{c} = 1650 \text{ \AA}$$

(7)

Huggins bands:

$$\kappa_{Hu} = 1.99 \times 10^{17}$$

 $M = 0.0126$
 $I_{Hu} = 53 \text{ ergs/cm}^2 \text{ sec } \hat{A}$
 $\lambda_{SHORT} = 2750 \hat{A}$
 $\lambda_{LONG} = 3400 \hat{A}$

Although our formulae for the Huggins bands are not explicitly in a form which would permit adaptation to other data, the method of choosing the parameters, discussed earlier in this note, is obviously adaptable. We chose to model Craig's (1951) somewhat outdated calculation simply because complete spectral information was conveniently available for his calculation. This case is sufficient to demonstrate that ozone heating can be accurately modelled by two Chapman layers plus a modified Chapman layer for the Huggins bands, which is our main result. <u>Acknowledgments.</u> This note was prepared with support from National Science Foundation grant GA-25904 and National Aeronautics and Space Administration contract NASA NGR 14-001-193.

- Craig, R. A., 1950: The observations and photochemistry of atmospheric ozone and their meteorological significance. <u>Metcor. Monogr., 1</u>, No. 2, Boston, Amer. Meteor. Soc., 50 pp.
- ____, 1951: Radiative temperature changes in the ozone layer. <u>Compen-</u> <u>dium of Meteorology</u> (T. F. Malone, Ed.). Boston, Amer. Meteor. Soc., pp. 292-302.

, 1965: The Upper Atmosphere: Meteorology and Physics. New Y York, Academic Press, 509 pp.

Lindzen, R. and Goody, R., 1965: Radiative and photochemical processes in mesospheric dynamics: Part I. Models for radiative and photochemical processes. J. Atmos. Sci., 22, 341-348.

FIGURE LEGENDS

Fig. 1. The solid curve, taken from Craig (1950), shows the spectral distribution of incoming solar radiation. Also shown are the average values of intensity and band limits used in our modelling procedure.

Fig. 2. The solid curve shows absorption by ozone as a function of wavelength for the Hartley and Huggins bands (as taken from Craig, 1950). Also shown (dashed curves) are the approximations to the absorption used in our modelling procedure.

Fig. 3. Same as Fig. 2 but for Chappuis band.

Fig. 4. Specific heating, n, due to ozone as a function of optical depth, u. Solid curve represents Craig's (1951) calculation; dashed curve represents our analytic formula.

Incoming Solar Radiation (ergs cm² sec¹ Å⁻¹) 03 10[°]L 1000 2000 $\overline{\mathbf{O}}$ ō ך מ -= 9.0 Ē 2375 Å 11 53) 3000 1 5000 6000 4000 8000 6000 Wavelength (Å) Hartley Band 2750 Å 3400 Å Huggins Band 5150 Å Band Chappuis Ц С 6800 Ă 7000 11 081 8000

Þ



Absorption Coefficient (cm NTP') 10⁻²/4400 5200 4800 100 \overline{O}_{I} 5150 Å 5600 6400 Wavelength (Å) 6000 Chappuis Band K <u>6800 Å</u> 6800 7200 EB

10

