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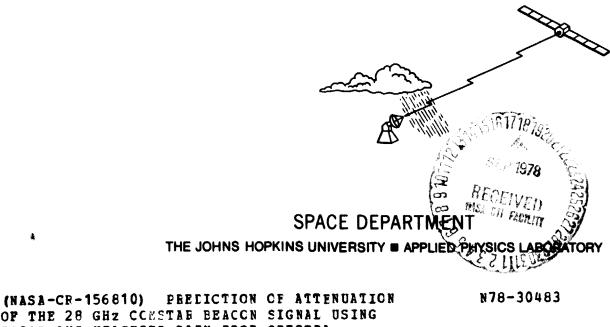
APL/JHU S1R78U-009 **FEBRUARY 1977** 



# PREDICTION OF ATTENUATION OF THE 28 GHz COMSTAR BEACON SIGNAL USING RADAR AND ( MEASURED RAIN DROP SPECTRA

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by JULIUS GOLDHIRSH



RADAR AND MEASUFED RAIN DRCP SPECTRA (Applied Physics Lab.) 46 p HC A03/MF A01 Unclas CSCL 20N G3/32 30521 APL/JHU S1R78U-009 FEBRUARY 1977

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# PREDICTION OF ATTENUATION OF THE 28 GHz COMSTAR BEACON SIGNAL USING RADAR AND MEASURED RAIN DROP SPECTRA

by JULIUS GOLDHIRSH

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SPACE DEPARTMENT

THE JOHNS HOPKINS UNIVERSITY APPLIED PHYSICS LABORATORY Johns Hopkins Road, Laurel, Maryland 20810 Operating under Contract N00017-72-C-4401 with the Department of the Navy

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# PREDICTION OF ATTENUATION OF THE 28 GHz COMSTAR BEACON SIGNAL USING RADAR AND MEASURED RAIN DROP SPECTRA

#### Abstract

A program to measure the rain attenuation of the COMSTAR beacon signal at 28.56 GHz has been in continuous operation since March of 1977 at Wallops Island, Virginia. During the summer of 1977, simultaneous radar and disdrometer measurements were also made and used for predicting path attenuation. The best fit values of a and b of the relation  $k = a Z^b$ were deduced for each rain period from the raindrop size measurements; where k is the attenuation coefficient [dB/km] and Z is the relfectivity factor  $[mm^6/m^3]$ . The measured k-Z relations and the simultaneous radar reflectivity measurements along the beacon path were injected into a computer program for estimating the path attenuation. Predicted attenuations when compared with the directly measured ones showed generally good correlation on a case by case basis and very good agreement statistically.

The results demonstrate the utility of using radar in conjunction with disdrometer measurements for predicting fade events and long term fade distributions associated with earth-satellite telecommunications. It also suggests its use for arriving at sampled fade distributions from which prediction criteria associated with variable path angle, frequency, and site diversity can be tested.

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# PREDICTION OF ATTENUATION OF THE 28 GHz COMSTAR BEACON SIGNAL USING RADAR AND MEASURED RAIN DROP SPECTRA

#### 1.0 INTRODUCTION

The ability to use radar as a predictor of rain attenuation has been previously tested against actual direct measurements by several investigators. McCormick [1] and Strickland [2] correlated their 3 GHz radar data with path attenuations at 15.3 GHz; the former investigator using a beacon on an aircraft, and the latter using transmissions from the ATS-5 satellite. Goldhirsh [3] correlated his 3 GHz radar measurements with 13 and 18 GHz measured attenuations using the ATS-6 in an uplink mode. Although some statistical aspects were considered by the above investigators, none examined how well cumulative fade distributions as derived from radar compared with those obtained from direct measurements. More recently, Hodge and Austin [4] compared radar measurements at 3 GHz using the McGill radar with radiometer-predicted attenuation at 13 GHz and were able to arrive at predictive parameters as well as a radar calibration by matching radar and radiometer-derived cumulative fade distributions.

We describe here an experiment performed at Wallops Island, Va., using a disdrometer and radar system for estimating the rain attenuation of the COMSTAR beacon signal at 28.56 GHz. Drop size distributions measured during each rain period enabled the calculation to be made of empirical parameters which were injected into the radar prediction scheme. In this way, the dominant drop size distribution for each rain period was used. Results using the measured drop size spectra are compared with those using the Marshall-Palmer distribution [5]. Individual case histories and overall cumulative fade distribution comparisons are presented. The data base corresponds to five rain days during the summer of 1977 (June through September) and represents 13 attenuation events and 304 minutes of simultaneous radar data.

## 2.0 EXPERIMENTAL CONFIGURATION

#### 2.1 General

The experimental configuration which is depicted in Fig. 1 consists of a phase locked loop receiving system operating at 28.56 GHz, an

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S-band radar (f = 2.84 GHz) located approximately 30 m away, and a system of three raingages and two disdrometers for measuring rain drop size distributions located in the immediate vicinity of the receiving antenna. Both antennas are fixed and point in the direction of the COMSTAR geosynchronous satellite [95°W longitude  $\pm$  0.1°] with elevation and azimuth angles of 41.6° and 210°, respectively.

In Fig. 2 are shown the geographical location of Wallops Island (WI), Va., and the azimuth pointing direction. The circled region about WI represents the radar observation area.

The satellite receiver as well as the disdrometer-raingage system are functioning continuously for arriving at long term statistics (to be described in a future report). During selected periods of rain, the radar monitors the reflectivity along a parallel path for purposes of testing radar methods for predicting the path attenuation at 28.56 GHz. The disdrometer data are used as input in these prediction methods.

## 2.2 Receiver System

The pertinent parameters for the COMSTAR receiving system are given in Table 1. The system (Fig. 3) consists of an rf front end mounted at the antenna location. The IF output of 1.05 GHz feeds a phase lock loop receiver located in the radar building. The output voltage ranges from 0 to 5 VDC and covers the input dynamic range of 32 dB. This output is continuously recorded on a calibration strip chart. In addition, the output signal is coded and recorded on one channel of a four channel tape recorder. The tape recorder is automatically actuated whenever the attenuation exceeds a defined threshold, a raingage bucket tip occurs, or a drop impacts on the disdrometer sensors.

Calibration of the receiver system (Fig. 4) is performed by injecting a stable 1.05 GHz signal at various levels into the front end where it is upconverted and routed (switch in up position) through the mixer preamplifier downconverter. The 1.05 GHz IF output passes through the phase locked loop system. Calibrations are performed daily and recorded on tape. On several occasions, the satellite beacon signal was used as a source with

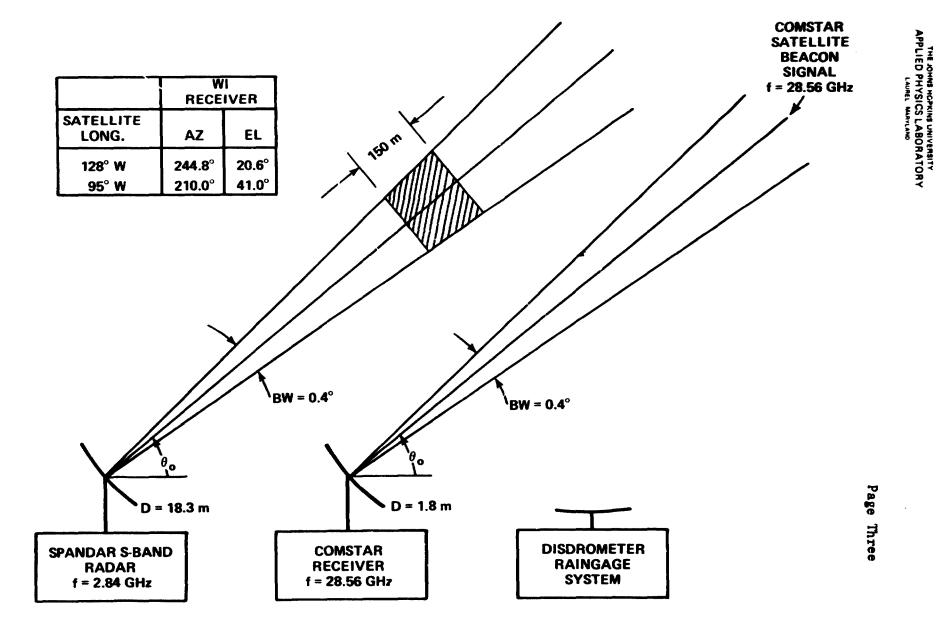
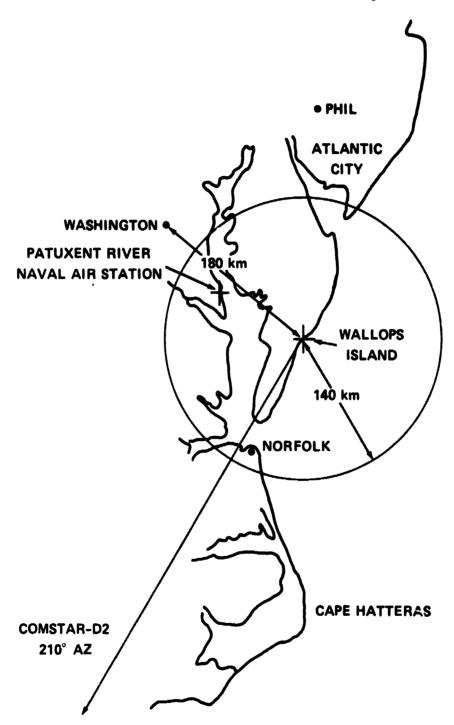
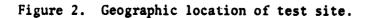


Figure 1. Experimental configuration at Wallops Island, Virginia. Antennas print in the direction of the Satellite at 95°W longitude.







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# Table 1

## PERTINENT COMSTAR RECEIVING SYSTEM PARAMETERS

Antenna Gain		52.7 JB	
Beamwidth	0.4°		
Antenna Diameter		1.83 m (6 ft)	
Free Space Power Received		-106 dBm	
ERP		55.8 dBm	
Path Loss		-213.4 dB	
Line Loss	11	-0.5 dB	
Sky Loss		-0.3 dB	
Phaselock Hold-In Threshold		-138 dBm	
Phaselock Acquisition Threshold		-133 dBm	
Phaselock Acquisition Threshold-133 dBmDynamic Range (Minimum)32 dB		32 dB	
Predetection 3 dB Bandwidth		50-100 Hz	
Post Detection 1 dB Bandwidth		10-20 Hz	
Amplitude Measurement Error		<u>+</u> 1 dB	
anna Diameter1.83 m (6 ft)e Space Power Received-106 dBmERP55.8 dBmPath Loss-213.4 dBLine Loss11Sky Loss-0.5 dBSky Loss-0.3 dBselock Hold-In Threshold-138 dBmselock Acquisition Threshold-133 dBmamic Range (Minimum)32 dBdetection 3 dB Bandwidth50-100 Hzt Detection 1 dB Bandwidth10-20 Hz			

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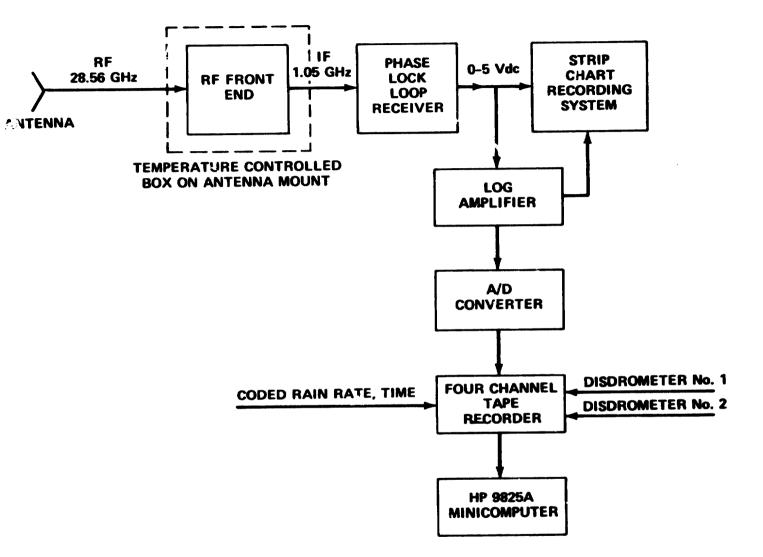


Figure 3. Elements of satellite receiving configuration.

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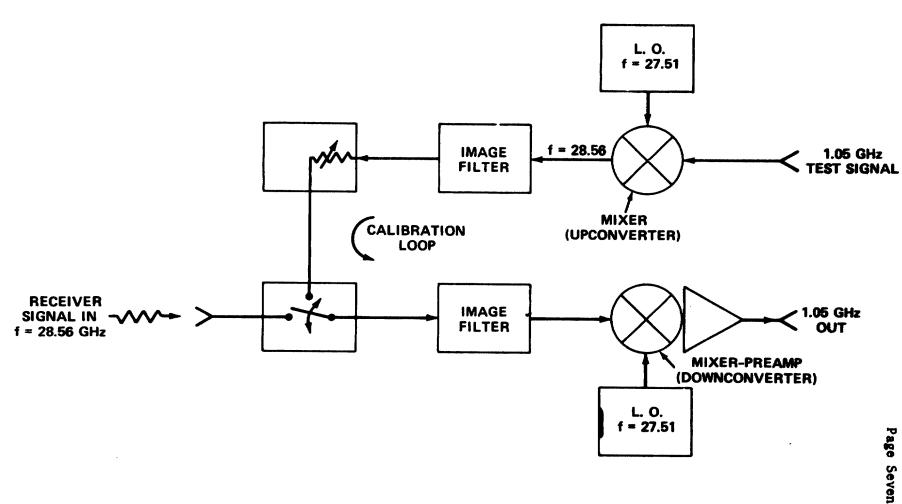


Figure 4. Calibration configuration for the COMSTAR receiver.

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the attenuation placed at the rf front end input. These latter calibrations are within + 0.5 dB of the loop calibrations.

Since the satellite position is maintained fixed through stationkeeping to within  $\pm 0.1^{\circ}$ , the receiver antenna is also kept fixed. However, an approximate weekly adjustment of the antenna pointing is made whenever the satellite assumes a new nodal position. The motion of the satellite within its  $\pm 0.1^{\circ}$  spatial box results in a free space diurnal variation of less than 2 dB which is pindictable to within  $\pm 0.5$  dB.

The reflection losses due to water on the feed were measured by controlled water spray tests and were observed to be on the average of 2 dB within an uncertainty of + 1 dB using a hydrophobic agent (Fusidox M)\*.

2.3 Radar System

The operating parameters of the radar (SPANDAR) are listed in Table 2.

During selected periods of rain, the radar reflectivity is monitored at adjacent range bins along the path where each bin has resolution of 150 m (1  $\mu$ sec pulsewidth). The resulting reflectivity represents the average of 128 independent samples obtained in 0.4 seconds (prf - 320 Hz) using a frequency diversity transmission scheme where each pulse in the prf is shifted by an amount greater than the inverse pulsewidth [6]. The integrated data are stored on IBM compatible tape for later analysis on the IBM 360/91 computer.

The electrical parameters of the radar have been directly calibrated [7] and are routinely checked by tracking 6-inch spheres on ballons [8] and by monitoring the solar flux from the sum at 2.84 GHz [9]. Relative standard gain horn measurements are also periodically performed. The accuracy of the absolute calibration of SPANDAR based on on-going measurements as well as previous component calibrations [7] is well within  $\pm$  1.5 dB.

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# Table 2

## OPERATING PARAMETERS OF SPANDAR

Radar	FPS/18
Gain	50.6 dB
Peak Power	1.0 MW
Beamwidth	0.4°
Diameter	18.3 m (60 ft)
Frequency Diversity:	
Center Frequency	2.840 GHz
Frequency Excursion, $\Delta f$	<u>+</u> 13.5 MHz
Pulse Length	l µsec
Pulse Range Resolution	150 m
prf	320 pps
Number of Pulses Integrated	128
Averaging Time	0.4 sec
Number of Adjacent Range Gates	871
Calibration Error (rms)	< 2 dB

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#### 2.4 Disdrometer-Raingage System

The raingages used are of the tipping bucket variety and the disdrometers which were developed at APL [10,11] are of the impact type. In Fig. 5 is shown a cross section view of the sensor itself. When a drop impacts at terminal velocity on the plexiglass aperture, a voltage is induced at the piezoelectric transducer and this voltage is proportional to the drop momentum. The transducer output voltage is passed through a preamplifier, shaping circuitry, a biphase encoder, and ultimately to one channel of the four channel tape recorder. Calibrations of the disdrometers are performed by dropping drops of known diameters through capillaries at sufficient heights so that drop terminal velocities are achieved.

The basic elements of the disdrometer-raingage system are given in the block diagram in Fig. 6. The coded outputs of two disdrometers are each fed into one channel of a four channel tape recorder. Outputs of three raingages are multiplexed with a time code (IRIG B Clock) and fed into another channel of the tape recorder. Simultaneously the raingage and disdrometer rain rates obtained using hard wire circuitry are recorded side by side on a chart recording system. A real time comparison of these rain rates represents an on-going check on the performance of the disdrometer.

#### 3.0 ANALYTICAL ASPECTS

#### 3.1 Radar Near Field Correction

Since reflectivities are measured at ranges as close as 600 m away from SPANDAR, and the far zone distance,  $r_f$ , is

$$\mathbf{r}_{f} = \frac{2D^2}{\lambda} = 6.3 \text{ km}$$
(3.1)

it is necessary to account for near zone effects. In this section we describe a method which enabled the measurement of reflectivities to be made as close in as 600 m by correcting the range-dependent gain and beamwidth in the near field region.

For a tapered aperture illumination, the gain reduces and the beamwidth increases for axial distances smaller than  $r_f$ . At the shorter

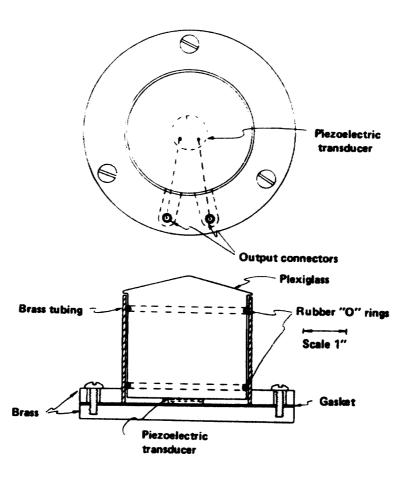
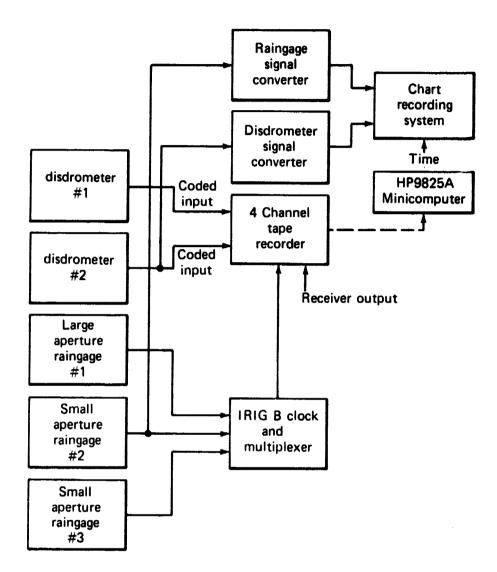
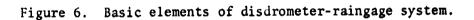


Figure 5. Cross sectional view of the disdrometer sensor.

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ranges there results a monotonic increase in power density (at a rate smaller than  $1/r^2$ ) down to 0.1  $(2D^2/\lambda)$ . At closer ranges oscillations occur with a diminishing envelope [12].

It is apparent from the above that the radar equation must be appropriately modified so as to account for the near field effects. In Appendix A we have modified the radar equation using the near field formulation of Hansen [12] to be applicable down to 0.1  $(2D^2/\lambda) \approx 600$  m. At closer ranges the clutter of SPANDAR does not allow any further adjustment and we therefore assume the reflectivity to be uniform for the first 600 m of range and use the value measured at 600 m. Using the results of Appendix A, the near zone power, P(r), is obtained by **digi**ding the far zone power formulation, P<sub>o</sub>(r), by the appropriate range dependent correction factor, C(r). That is,

$$P(\mathbf{r}) = \frac{P_{o}(\mathbf{r})}{C(\mathbf{r})}$$
(3.2)

where,

$$\left[C(\mathbf{r})\right]^{-1} = \left(\frac{\mathbf{r}}{\mathbf{r}_{f}}\right)^{2} \beta(\mathbf{r})$$
(3.3)

and where,

$$r \leq r_f = \frac{2D^2}{\lambda}$$
 (3.4)

$$\beta(\mathbf{r}) = \frac{256}{\pi^2} \left\{ 1 - \frac{16X}{\pi} \sin\left(\frac{\pi}{8X}\right) + \frac{128}{\pi^2} X^2 \left[ 1 - \cos\left(\frac{\pi}{8X}\right) \right] \right\}$$
(3.5)

$$X = \frac{r}{r_f}$$
(3.6)

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The reflectivity factor in the near zone, Z [dBZ] is related to the reflectivity factor using the far zone formulation by the relation,

$$Z[dBZ] = Z_{o}[dBZ] + 10 \ Log_{10} \ C(r)$$
(3.7)

The correction factor C(r) may be shown to be the ratio of the far field gain to the near field gain. That is,

$$C(r)_{dB} = 10 \ \log_{10}\left\{\frac{G_{o}}{G(r)}\right\}$$
 (3.8)

In Fig. 7,  $C(r)_{dB}$  is plotted as a function of range from SPANDAR. We may deduce from this curve, for example, that at 0.1  $(2D^2/\lambda) = 600$  m from SPANDAR, a correction of 4.3 dB must be added to the reflectivity factor calculated using the far zone formulation. The correction factor given by (3.3) was used in the radar program to arrive at the corrected reflectivity factors in the near region as described above.

## 3.2 Radar Prediction Formulation

We here describe the formulations used in arriving at the radar and disdrometer predicted attenuations and subsequent probability levels.

The radar power recorded in each of the adjacent range bins is converted into reflectivity factor levels,  $Z \ [mm^6/m^3]$  and these are injected in the formulation for path attenuation given by,

$$A_{R}(t) = \sum_{i=1}^{N} k_{i} \Delta r \qquad [dB] \qquad (3.9)$$

where

$$k_{i} = a Z_{i}^{b} [dB/km]$$
 (3.10)

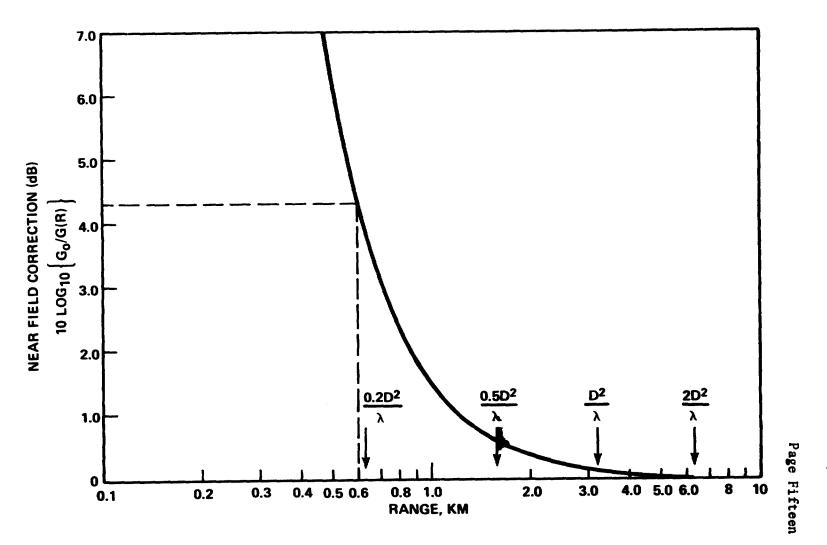


Figure 7. Near field correction factor as a function of range from SPANDAR.

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and where

A<sub>R</sub>(t) = the radar predicted attenuation in time, t [dB]
r<sub>i</sub> = the attenuation coefficient pertaining to i<sup>th</sup> range bin [dB/km]
Z(r<sub>i</sub>) = the reflectivity factor at i<sup>th</sup> range bin [mm<sup>6</sup>/m<sup>3</sup>]
Ar = range resolution interval (150 m)

N = number of range bins used in summation (see text)

The limit, N, in (3.9) is obtained in the following way. Through the observation of photographs of RHI's taken at 1/2 hour to 1 hour intervals, it is noted as to whether the rains are either convective or nonconvective; the latter showing a bright band and the former showing cellular activity rising to elevations above the zero degree isotherm height. The summation is taken up to a range corresponding to a zero degree isotherm height when a bright band appears since the dominant presence of ice above this height causes negligible attenuation. On the other hand, the summation is taken along the entire path when the rain is convective; implying the dominant presence of water above the zero degree isotherm height. This procedure has been found to give good results both in the previous work [3] as well as in the present experiment.

The values of a and b were arrived at by sampling continuous 30 second drop size distributions (hereafter referred to as DSD) with the APL disdrometer during each rain period and calculating

$$k = \int_{D_{\min}}^{D_{\max}} N(D) C_{ext}(D) dD \quad [dB/km] \qquad (3.11)$$

$$Z = \int_{D_{\min}}^{D_{\max}} N(D) D^{6} dD \quad [mm^{6}/m^{3}]$$
(3.12)

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where

k	= theoretical attenuation coefficient [dB/km]
N (D) dD	= number of drops per unit volume between drop sizes
	D and D+dD (DSD)
C <sub>ext</sub> (D)	= extinction factor $(dB/km) \times cm^3$
Z	= theoretical reflectivity factor $[mm^6/m^3]$
D <sub>min</sub> , D <sub>max</sub>	= the minimum and maximum diameter measured,
man musi	respectively.

Best fit values of a and b were generated for each rain period; the rain period representing a continuous rain event for the day in question. The extinction factors for 28.56 GHz,  $C_{ext}(D)$ , were arrived at by interpolating the results of Medhurst [13].

## 4.0 RESULTS

As mentioned previously, the results presented here correspond to the summer of 1977 during which approximately five hours of radar data of rain were obtained simultaneous with path attenuation measurements. This period encompassed five rain days and 13 attenuation events. The disdrometer results presented were taken from data acquired over a period which flanked the radar measurements.

In Table 3 are summarized the sampling days and times for both the radar and disdrometer measurements. Also listed are the measured best fit parameters a and b (5th and 6th columns) pertaining to the relation (3.10) for the indicated rain events. In the 7th column are given values of the coefficient of correlation,  $r^2$ , which establishes a measure of the goodness of fit; 1 being perfect. In establishing the number of samples and to ensure a statistically adequate drop size number, only those spectra were selected in which the rain rate exceeded 2.5 mm/hr.

#### 4.1 Empirical Relations from Drop Size Spectra

Examples showing scatter plots of log k versus log Z are given in Figs. 8 through 11 for days 157 (June 6), 236 (August 24, afternoon and evening rain periods), and 257 (September 14). Each point is representative

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# Table 3

## SUMMARY OF SAMPLING DAYS AND TIMES AND CORRESPONDING DISDROMETER RESULTS

Date (Julian)	Event No.	Sampling Time Intervals Disdrometer Best Fit Para		Fit Parame	ters		
(Julian) Day 1977		Radar	Disdrometer	a	b	r <sup>2</sup>	No. of Samples
June 6	1	21:26-21:55	21:26-22:45	$1.17 \times 10^{-3}$	0.825	0.93	123
(157)	2	21:57-22:26					
	3	22:36-22:48					
June 9	4	12:44-13:13	12:44-14:49	7.36x10 <sup>-3</sup>	0.617	0.75	180
(160)	5	13:24-13:53					
	6	14:41-14:50					
Aug 24	7	14:56-15:24	14:20-16:30	$1.23 \times 10^{-2}$	0.492	0.60	102
(236)	8	15:38-16:07					
	9	16:09-16:25					
Aug 24-25 (236-237)	10	23:35-23:59	23:30-00:52	$1.45 \times 10^{-3}$	0.806	0.94	107
(230-237)	11	00:05-00:34					
Sept 14 (257)	12	20:03-20:32	20:03-20:46	$3.47 \times 10^{-3}$	0.733	0.95	53
(237)	13	20:33-20:48					
Overall Data Base		304 min	458 min	1.87x10 <sup>-3</sup>	0.775	0.89	565
Marshall-Palmer				2.01x10-3	0.773		

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of a 30-second drop size distribution where the corresponding calculations are made using (3.11) and (3.12). The solid line shown in each of the figures represent the best fit linear regression line used for that rain period.

In Fig. 12 are shown the best-fit k-Z curves plotted side by side for the five rain days. Also shown is the Marshall-Palmer (hereafter referred to as M-P) best-fit curve. We note that the distributions for three of the five rain periods (June 6, August 24-25, and September 14) are close to the M-P line. It may be noted that when all the drop size distributions are considered together, the values of a and b (Table 1) are almost identical with those corresponding to the M-P distribution (i.e.,  $a = 2.10 \times 10^{-3}$ , b = 0.773). This is attributed to the fact the M-P distribution is representative of a large data base for the geographic region of Washington, D. C. (mid-Atlantic east coast). On an individual rain period basis the values of a and b may be markedly different from the M-P distribution as is indicated for the June 9 and August 24 periods.

Other interesting features of the above results are that the drop size distributions for two distinct rain periods occurring for the same day may be markedly different as evidenced by the results of August 24. It also may be noted from Table 1 that the smaller the coefficient of correlation, the greater the deviation from the M-P, and where  $r^2$  exceeds 0.9, the fit is close to M-P.

## 4.2 Measured and Predicted Attenuation Events

In Figs. 13 through 17 are given examples of attenuation versus time plots of the predicted and measured attenuations corresponding to the best-fit regression parameters a and b obtained from the scatter plots of Figs. 8 through 11. The ordinate and abscissa in each of these curves represent the attenuation in dB and time in GMT, respectively. The solid curve represents the measured level, the dashed curve the predicted level using (3.9) and the measured empirical parameters, a and b. Also plotted is a dot-dashed curve which represents the attenuation obtained using (3.9) and a and b corresponding to the M-P distribution. The directly mecsured

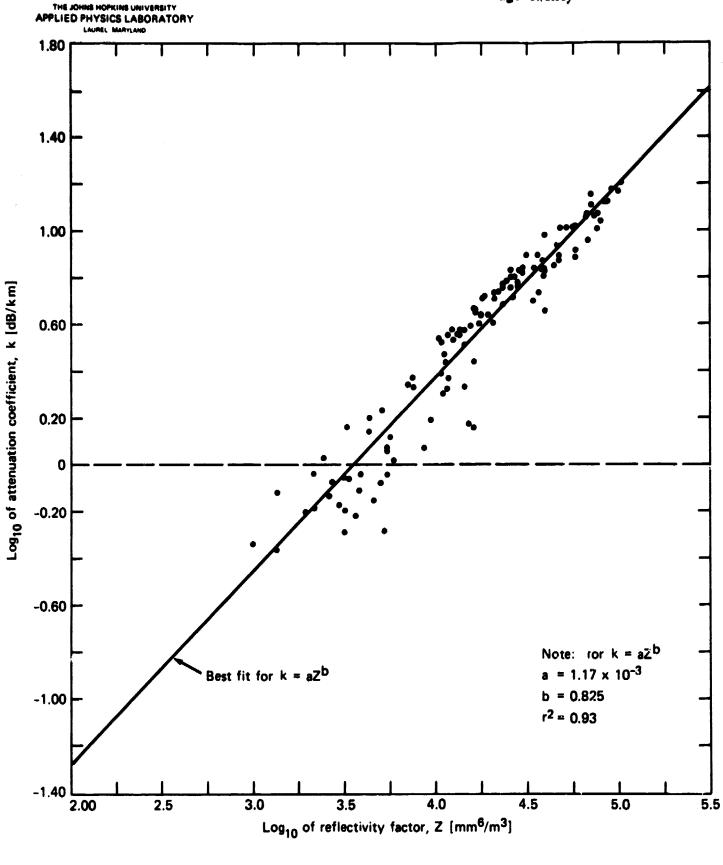


Figure 8. Scatter plot of log k vs log Z for June 6, 1977 (day 157) from 21 hr, 26 min to 22 hr, 45 min GMT. Solid line is best fit for  $k = a Z^b$  relation.

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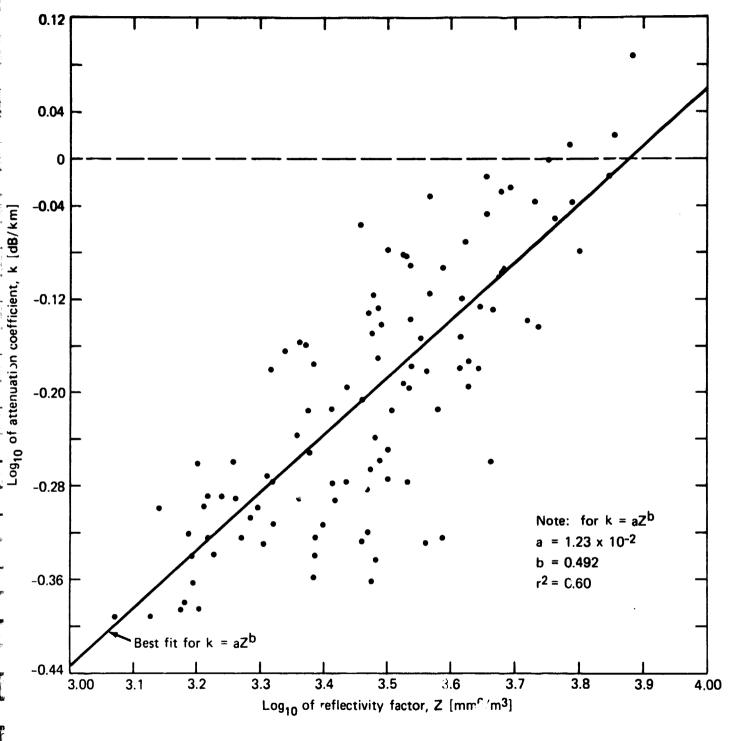
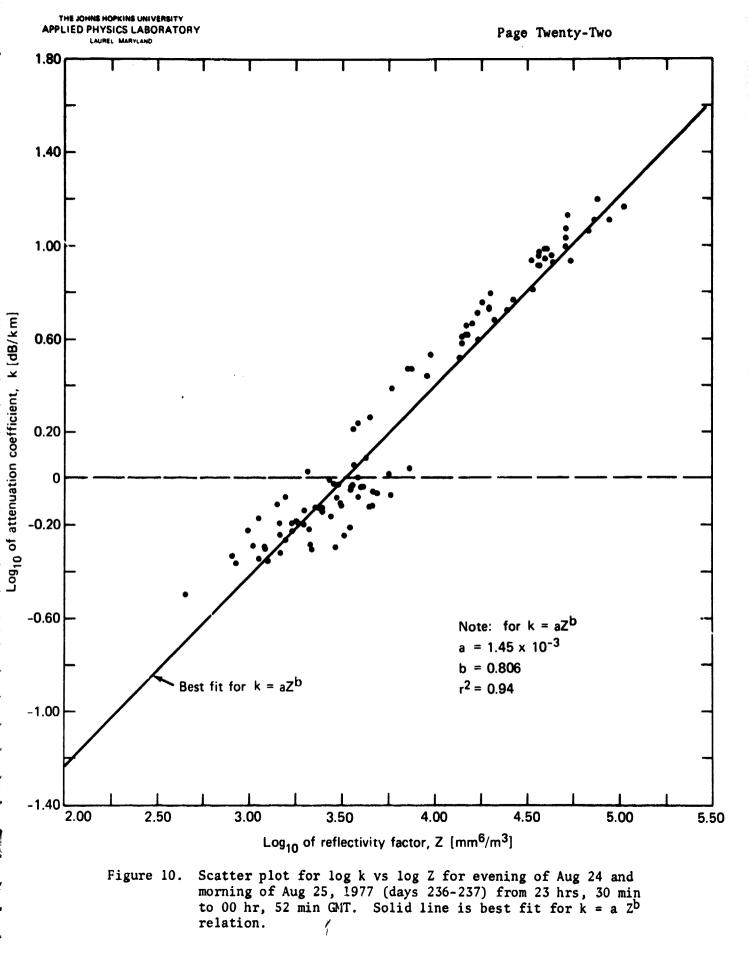


Figure 9. Scatter plot of log k vs log Z for early afternoon of Aug 24, 1977 (day 236) from 14 hr, 21 min to 15 hr, 30 min GMT. Solid line is best fit for k = a  $Z^b$  relation.



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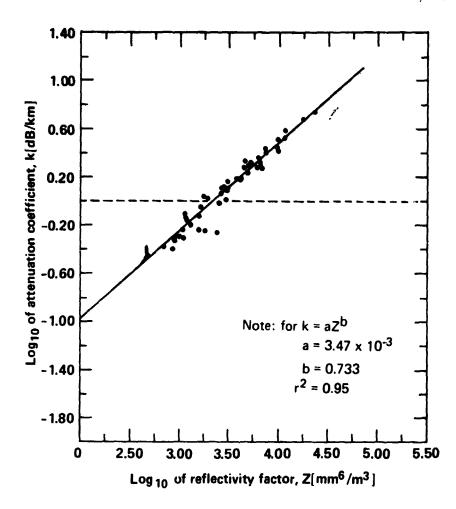
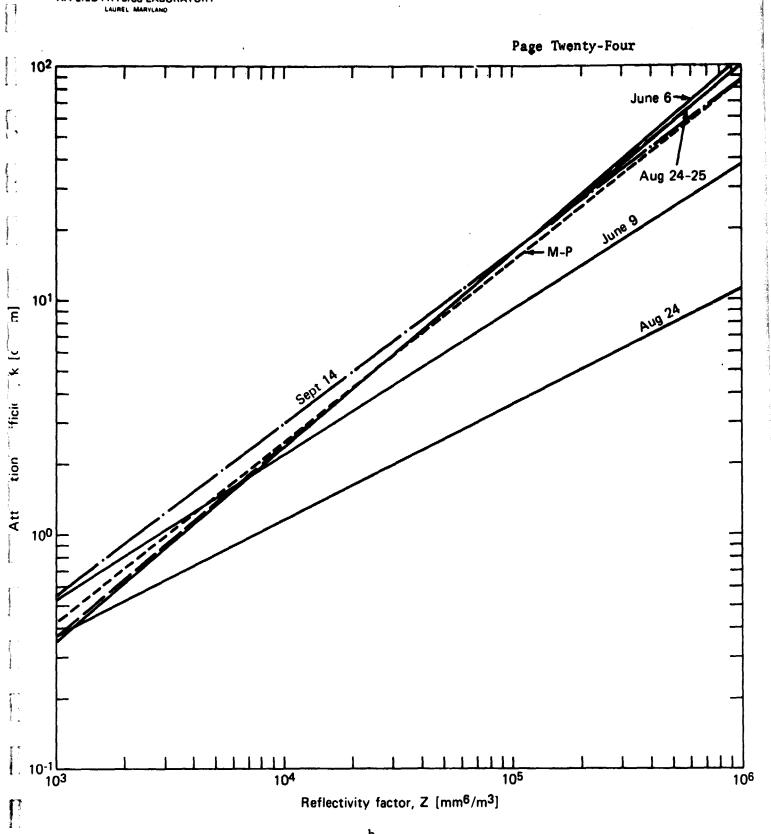
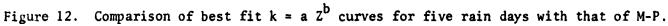


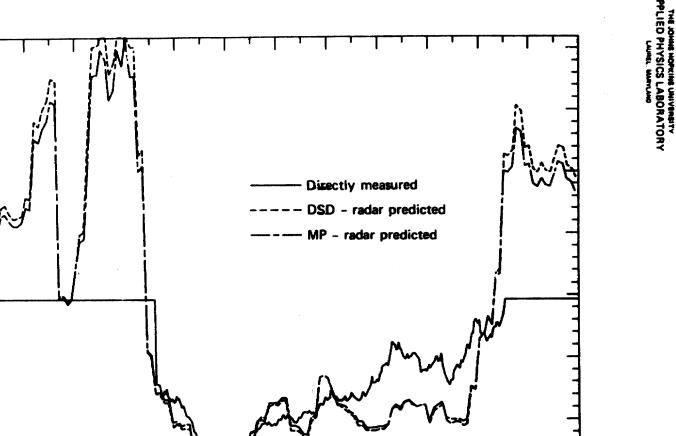
Figure 11. Scatter plot of log k vs log Z for Sept 14, 1977 (day 257) from 20 hr, 03 min to 20 hr, 46 min GMT. Solid line is best fit for  $k = a Z^b$  relation. THE JOHNS HOPKINS UNIVERSITY APPLIED PHYSICS LABORATORY LAUREL MARYLAND

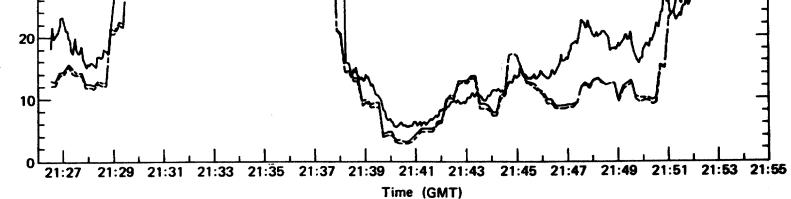
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Attenuation, dB

Figure 13. Comparison of directly measured fade event with DSD-radar predicted and M-P-predicted for June 6, 1977 (day 157) from 21 hr, 26 min to 21 hr, 55 min GMT.



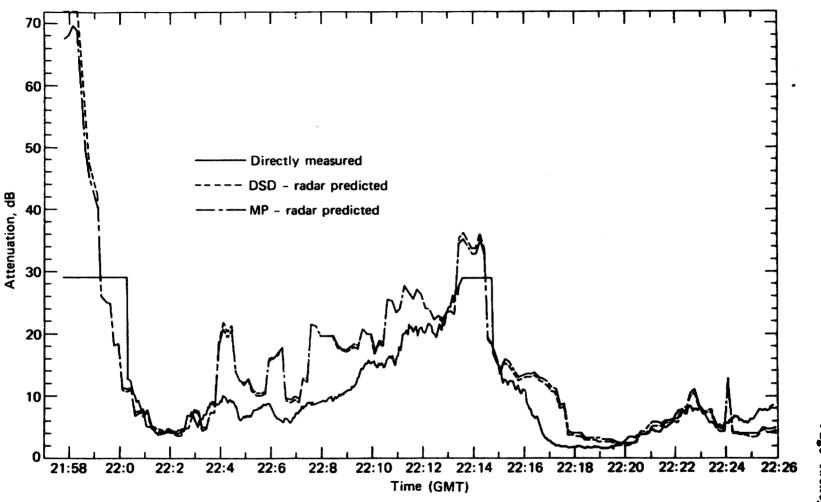


Figure 14. Comparison of directly measured fade event with DSD-radar predicted and M-P-radar predicted for June 6, 1977 (day 157) from 21 hr, 57 min to 22 hr, 25 min GMT.

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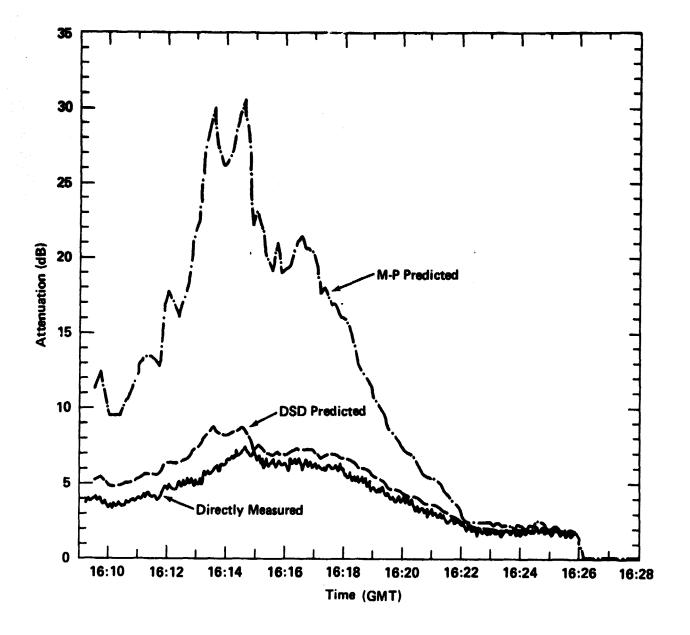


Figure 15. Comparison of directly measured fade event with DSD-radar predicted and M-P-radar predicted for early afternoon of Aug 24, 1977 (day 236) from 16 hr, 09 min to 16 hr, 26 min GMT.

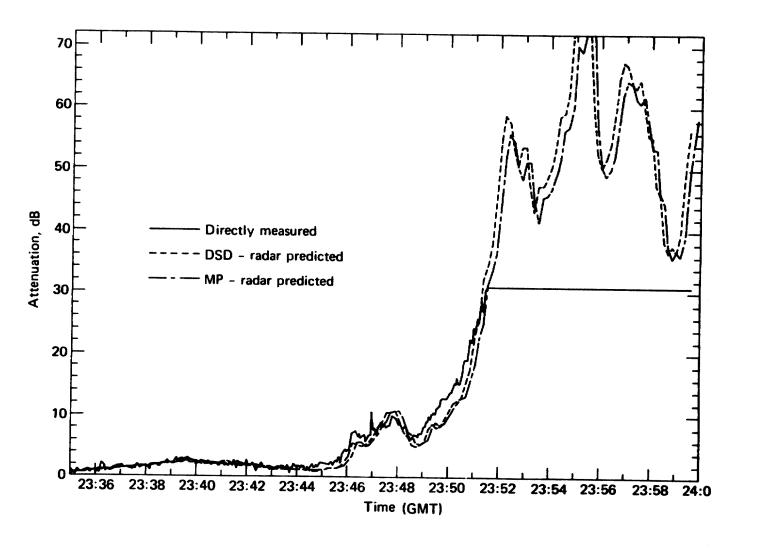


Figure 16. Comparison of directly measured fade event with DSD-radar predicted and M-P-radar predicted for evening of Aug 24, 1977 (day 236) from 23 hr, 35 min to 23 hr, 59 min GMT.

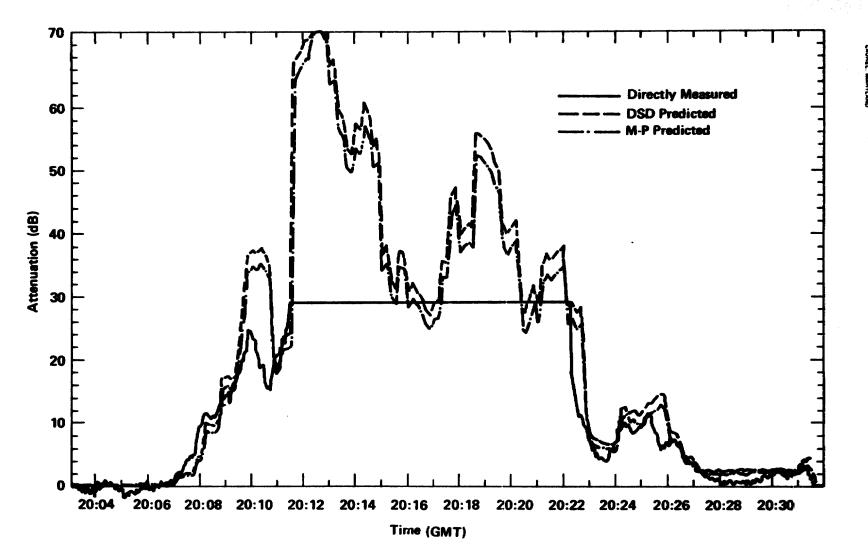


Figure 17. Comparison of directly measured fade event with DSD-radar predicted and M-P-radar predicted for Sept 14, 1977 (day 257) from 20 hr, 03 min to 20 hr, 32 min GMT.

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curves are plotted from receiver data sampled at 3 second intervals and the radar predicted curves are plotted from data samples at 10 second intervals.

The following interesting features may be noted from the above comparisons: (1) The receiver dynamic range limits the maximum attenuation excursion due to rain to approximately 30 dB beyond which a loss of lock condition results. This accounts for the flattened sections of the curves. The radar predicted levels do not suffer from this condition and show attenuations sometimes in excess of 72 dB. (2) The directly measured and predicted curves using DSD and radar measurements show overall good agreement although slots of time do exist exhibiting large deviations. (3) On those days in which the measured drop size spectra differed from M-P, the radar predicted attenuation using the M-P distribution showed poor agreement with the directly measured attenuation. An example of such case is depicted in Fig. 15. On the other hand, the DSD predicted results show rather close agreement. (4) For June 6 (Figs. 13 and 14), evening of August 24 (Fig. 16), and September 14 (Fig. 17), the predicted attenuations using the measured distributions are practically coincident with those of M-P. This, of course, should be expected, as the measured values of a and b are similar to those of M-P.

#### 4.3 Measured and Predicted Probability Distributions

We here compare the cumulative fade distributions of the directly measured and radar predicted cases. Such a comparison is given in Fig. 18. The vertical scale represents the probability the attenuation exceeds the abscissa; the abscissa corresponding to various attenuation levels. The solid points, circles, and square points represent, respectively, the directly measured, the DSD-radar predicted, and the M-P-radar predicted distributions.

The table in Fig. 18 summarizes quantitatively how well the radar predicted fade distributions agree with the measured ones. We note the rms and average dB deviations for the radar-DSD case are 1.2 and 1.0 and the average ratio of probabilities associated with the predicted and measured points is 1.1. For the M-P case, the respective values are 5.6, 5.0, and 1.37.

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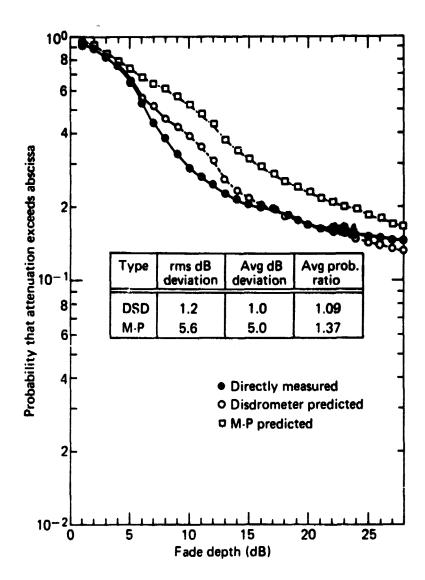


Figure 18. Comparison of cumulative distributions for directly measured, DSD-radar predicted, and M-P-radar predicted cases covering five rain days, 304 minutes of simultaneous radar sampling during summer of 1977.

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The average probability ratio of the two curves was obtained by averaging the ratio of the larger probability to the smaller one for each dB level up to 28 dB. This gives a measure of the percentage uncertainty of the predicted curve relative to the measured one. Hence, on an absolute basis, the disdrometer curve shows very good agreement with the directly measured whereas the M-P curve shows poorer agreement. Whether this would tend to obviate the use of the M-P spectra for such radar predictions is too early to say, as an increased data base might bring the M-P radar predicted distribution in closer agreement with the directly measured one.

#### 4.4 Empirical Calibration Adjustment

An adjustment in the radar calibration was deemed necessary in arriving at the reflectivity factors, Z; otherwise, the radar predicted attenuation levels consistently underestimate the directly measured values. When the reflectivity values were increased by multiplying by an empirically determined fixed factor of 3.5 (5.4 dB), both the case-by-case attenuation events as well as the cumulative distribution showed good agreement. The reason for this required adjustment is not at present understood although it might in part be due to a phenomenon described by Rogers [14]. Rogers demonstrated that for the case in which the backscatter is measured through a logarithmic receiver and there are simultaneously large gradients of reflectivity in the radar pulse volume, estimated reflectivity levels deviate from the true averages by an amount depending upon the gradient value of Z in the pulse volume. For example, according to one model, if there is a logarithmic variation of Z in the pulse volume with a ratio of maximum to minimum Z of 6, this could account for a 6 dB underestimation of the reflectivity factor.

Another possible reason for underestimations of the peak attenuation values may be due to the comparative spatial resolutions of the receiving and radar systems. The attenuation of the received beacon signal takes place essentially through the first Fresnel spheroidal region whose diameter is given by

$$D(FZ) = 2(r_0 \lambda)^{1/2} = 0.205 r_0^{1/2}$$
 [m]

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where  $r_0$  [m] is the axial distance from the receiving antenna and  $\lambda$  [m] is the beacon wavelength. The diameter of the cross section of the radar beamwidth ( $\theta = 0.4^\circ$ ) is given by,

 $D_{R} = 6.98 \times 10^{-3} r_{o}$  [m]

and the ratio of these diameters is,

$$\frac{D_R}{D(FZ)} = 3.4 \times 10^{-2} r_0^{1/2} [m]$$

We note that at a distance of 6 km the diameter associated with the radar beamwidth is approximately 2.5 times the Fresnel zone diameter and at 1 km they are equivalent. The received beacon signal is therefore expected to have greater attenuation peaks than that estimated by the radar because of the differences in averaging that takes place across the beams.

#### 5.0 SUMMARY, DISCUSSION, AND CONCLUSIONS

Although the overall data base was limited, the DSD and radar results showed the following salient features:

- (1) The radar predicted attenuations using the measured rain drop spectra showed generally good agreement on a case by case basis with the directly measured levels; although slots of time exist exhibiting deviations. On the other hand, extended rain periods exist where the M-P radar predicted curves show overall poor agreement.
- (2) The DSD-radar predicted cumulative fade distributions show very good agreement with the directly measured ones whereas the M-P results showed poorer agreement.
- (3) When the coefficient of correlation was high (i.e.,  $r^2 > 0.9$ ), the measured disdrometer results were in close agreement with the M-P spectra; and conversely.

Two distinct advantages of using radar results appear in the curves shown in Figs. 13, 14, 16, and 17. First, no limitation for examining severe

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attenuation levels exists in the radar predicted cases as does exist in a fixed dynamic range receiver. Secondly, the receiver used could not reacquire lock until it was 7 to 8 dB below the loss of lock condition and this reacquisition might take place only after a frequency search which could take several minutes. The radar measurements do not suffer from these intrinsic receiver characteristics. On the other hand, the radar predicted measurements do suffer from other features and these are: (1) An absolute calibration of the radar is necessary. (2) The drop size distribution may be both spatially and temporally variable, (3) The attenuation for the received satellite signal occurs along a narrow Fresnel ellipsoid whose maximum cross section diameter is 20 m whereas the radar prediction measures the average reflectivity of the radar beamwidth which may have approximately 3 times the cross section diameter. (4) The radar reflectivities are averaged over a pulse volume defined by the antenna beamwidth and the range resolution interval of 150 m within which the reflectivities may be highly non-uniform. (5) The beam axis of the radar and the receiver system are not coincident: being displaced by 30 m. (6) Near field correction formulations applied to the radar antenna system are based on idealized assumptions. and (7) The pulse volumes may contain snow or ice near the zero degree isotherm height.

In spite of the many reasons for non-correlation, the radar results do correlate remarkable well for the above example. In fact, the above results demonstrate the utility of using radar coupled with disdrometers measurements for predicting individual fade events as well as long term fade distributions associated with satellite communications through rain. It also suggests its use for arriving at sampled fade distributions from which prediction criteria associated with variable frequencies, path angles, and site diversity can be tested [15].

In the absence of DSD spectra, the M-P distribution might represent a viable alternative for use with the radar measurements; at least for the mid-Atlantic coast geographic region. An insufficient data base, however, exists here to be definitive about this assessment. THE JOHNS HOPKINS UNIVERSITY APPLIED PHYSICS LABORATORY LAUREL MARYLAND

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## 6.0 ACKNOWLEDGEMENTS

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## Appendix A

## DERIVATIONS OF THE RADAR EQUATION IN THE NEAR FIELD OF SPANDAR

The antenna of SPANDAR may be approximated by a circular aperture with a tapered illumination having the cross section distribution given by  $\left[1-\left(\frac{\rho}{D/2}\right)^2\right]$  where  $\rho$  is the cross section radial distance and D/2 is the radius. Using the results of Hansen [12] the gain, G(r), and beamwidth,  $\theta(r)$ , in the near region of this aperture (i.e.,  $r \leq 2 D^2/\lambda$ ) are range dependent and are given by,

$$\theta(\mathbf{r}) = \theta_{0} \gamma(\mathbf{r}) \tag{A.1}$$

$$G(\mathbf{r}) = \frac{G_0}{\gamma^2}$$
(A.2)

where

$$\gamma(\mathbf{r}) = \frac{\pi}{16X} \left\{ 1 - \frac{16X}{\pi} \operatorname{Sin}\left(\frac{\pi}{8X}\right) + \frac{128X^2}{\pi^2} \left[ 1 - \operatorname{Cos}\left(\frac{\pi}{8X}\right) \right] \right\}^{-1/2}$$
(A.3)

where

$$X = \frac{r}{r_f}; r_f = \frac{2D^2}{\lambda}$$
(A.4)

and where  $G_0$  and  $\theta_0$  are the far field gain and beamwidth, respectively.

We assume that the radar equation in the near region to have the same form as that in the far region with the exception that the beamwidths and gains are range dependent. That is,

$$P_{\mathbf{r}} = \left(\frac{c}{1024 \pi^2 \ln 2}\right) \left(P_{\mathbf{t}} \tau \lambda^2 \eta\right) \frac{G^2(\mathbf{r}) \theta_{\mathbf{e}}(\mathbf{r}) \theta_{\mathbf{h}}(\mathbf{r})}{2}$$
(A.5)

Page Two-A

where

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$$\lambda$$
 = the wavelength (m)

= the range (m) r

Dividing (A.5) by the far field formulation, we obtain

$$\frac{P_{r}}{P_{ro}} = \frac{G^{2}(r) \theta_{e}(r) \theta_{h}(r)}{G_{o}^{2} \theta_{eo} \theta_{ho}}$$
(A.6)

where the subscript o denotes the far field values.

Substituting (A.1) and (A.2) into (A.6)

$$\frac{P_{\mathbf{r}}}{P_{\mathbf{r}o}} = \frac{1}{\gamma^{2}(\mathbf{r})} = \frac{G(\mathbf{r})}{G_{o}} = \left(\frac{\mathbf{r}}{\mathbf{r}_{f}}\right)^{2} \beta(\mathbf{r})$$
(A.7)

where from (A.3),  $\beta(r)$  is given by,

$$\beta(\mathbf{r}) = \frac{256}{\pi^2} \left\{ 1 - \frac{16X}{\pi} \sin\left(\frac{\pi}{8X}\right) + \frac{128X^2}{\pi^2} \left[ 1 - \cos\left(\frac{\pi}{8X}\right) \right] \right\}$$
(A.8)

We note from (A.7) that the far zone power formulation,  $P_{ro}$ , is modified by a correction factor equal to the ratio of the near to the far field gains. In Fig. 7, 10  $\log_{10}$  of  $\gamma^2(r)$  is plotted down to 600 m from SPANDAR. We note that this parameter monotonically increases down to this range reaching a value of 4.3 dB. The far zone power formulation must therefore be reduced by the indicated levels in the near zone region.