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Documentation of the Fourth Order Band Model

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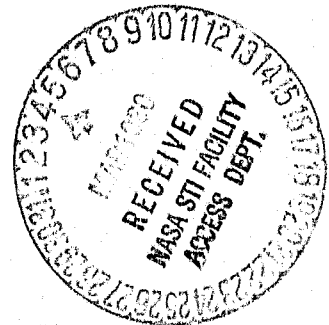
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Greenbelt, Maryland 20771



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Fourth Order Band Model

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November 1979

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Preface

We have decided to compile a preliminary documentation of the new GLAS Fourth Order General Circulation Model. The present documentation has not been subjected to a careful editing process; we hope that its possible usefulness will compensate for some of its defects. The model dynamics (COMP0, COMP1 and COMP2) is still undergoing minor improvements, especially in the time differencing scheme which we hope will improve its efficiency. The "physics" routine (COMP3) has not been documented because it is being thoroughly revised. The present version of COMP3, similar to the one used in the GLAS/GISS models (see the documentation by Tsang and Karn), with modifications introduced by Y. Sud (1979) is included in the code. Criticisms and suggestions for improvement will be greatly appreciated, since a final documentation will be prepared in 1980.

We are very grateful to all the people that have helped us generously. In particular, Dr. N. Rushfield had a major impact in the process of making the model operational. W. Connelly, D. Edelmann, D. Han, S. Breining, P. Anolick and M. Almeida were very helpful in the development of the model. The documentation was expertly and cheerfully typed by S. Mathis; D. Edelmann and D. Rosen have also cooperated in its compilation. We want to express our special gratitude toward Dr. Y. Sud, who offered us generously both his advice and his time in the development of the "physics" routine, and most especially to Dr. M. Halem without whose many useful suggestions, constant encouragement, and long patience, this work would not have been finished.

Eugenia Kalnay-Rivas

November 1979

I. Introduction

The band fourth order model is a GCM which uses quadratically conservative, fourth order horizontal space differences on an unstaggered grid and second order vertical space differences with a Matsuno (forward-backward) or a smooth leapfrog time scheme to solve the primitive equations of motion.

This program numerically solves these equations one latitude band at a time which greatly reduces the amount of computer core storage needed to run the program. It also uses the same variable names, order of computations, I/O, post-processing as the standard second order GCM. Appropriate modifications have been made for the fourth order differences and leapfrog scheme. (See the 1978 Goddard Modeling and Simulation Research Review for an overview of the fourth order band model.)

The main feature of this model is that fourth order approximations are used for all the horizontal derivatives. The derivative $\frac{\partial q}{\partial x}$ is approximated by

$$\frac{4}{3} \left(\frac{q(x+\Delta x) - q(x-\Delta x)}{2\Delta x} \right) - \frac{1}{3} \left(\frac{q(x+2\Delta x) - q(x-2\Delta x)}{4\Delta x} \right)$$

and the derivative $\frac{\partial}{\partial x} (qT)$ by

$$\frac{4}{3} \left[\frac{(T(x)+T(x+\Delta x))(q(x)+q(x+\Delta x)) - (T(x)+T(x-\Delta x))(q(x)+q(x-\Delta x))}{4\Delta x} \right] \\ - \frac{1}{3} \left[\frac{(T(x)+T(x+2\Delta x))(q(x)+q(x+2\Delta x)) - (T(x)+T(x-2\Delta x))(q(x)+q(x-2\Delta x))}{8\Delta x} \right]$$

The second approximation is derived by averaging the flux qT to yield a conservative form of the dynamic equations. Note that if T is equal to 1 the second equation reduces to the first.

The primary variables are the horizontal components of the wind velocity, $W=(u,v)$, the temperature, T , the specific humidity, q , and the shifted surface pressure, π , ($\pi=P_s-P_{top}$, $P_{top}=10$ mb).

The secondary variables are the geopotential, ϕ , the vertical wind velocity, $\dot{\sigma}$, and the pressure, p .

The following pages give the differential equations of motion for the GCM model with the initial and boundary conditions. This is followed by the equations with the corresponding fourth order approximations which use the same notation as the current second order model. A complete description of the primitive equations with the σ coordinate system is found in the Arakawa UCLA notes (1976).

II. Primitive Equations of Motion

1.&2. Horizontal momentum equations

$$\begin{aligned} \mathbf{V} \frac{d\pi}{dt} + \pi \frac{d\mathbf{V}}{dt} &= \frac{\partial \pi \mathbf{V}}{\partial t} + \nabla \cdot (\pi \mathbf{V} \mathbf{V}) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma} \mathbf{V}) \\ &= -\pi \nabla \phi - \pi \sigma \frac{RT}{p} \nabla \pi - (f + u \frac{\tan \phi}{a}) k \times \pi \mathbf{V} + \pi \mathbf{F} \end{aligned}$$

3. Continuity equation

$$(3.1) \quad \frac{\partial \pi}{\partial t} + \nabla \cdot (\pi \mathbf{V}) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma}) = 0, \quad \text{or}$$

$$(3.2) \quad \frac{\partial \pi}{\partial t} = - \int_0^1 \nabla \cdot (\pi \mathbf{V}) d\sigma = - \nabla \cdot \int_0^1 \pi \mathbf{V} d\sigma$$

4. Equation of state

$$\alpha = \frac{RT}{p}$$

5. First law of thermodynamics

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + \frac{\partial}{\partial \sigma} (\pi \dot{\sigma} T) = \frac{\pi \omega \alpha}{C_p} + \frac{\pi Q}{C_p} \quad (\omega = \frac{dp}{dt})$$

From $\theta = T/p^k$, $p = p_T + \sigma \pi$, $\omega = \dot{\sigma} \pi + \dot{\pi} \sigma$, $\dot{\pi} = \frac{\partial \pi}{\partial t} + \mathbf{V} \cdot \nabla \pi$, $k = R/C_p$ we get

$$\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} = p^k \frac{\partial \pi \dot{\sigma} \theta}{\partial \sigma} + \pi \frac{\dot{\sigma} \alpha \pi}{C_p} \quad \text{Replacing in 5,}$$

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + p^k \frac{\partial \pi \dot{\sigma} \theta}{\partial \sigma} = \frac{\pi \sigma k T}{p} \left(\frac{\partial \pi}{\partial t} + \mathbf{V} \cdot \nabla \pi \right) + \frac{\pi Q}{C_p}$$

6. Humidity equation

$$\frac{\partial \pi q}{\partial t} + \nabla \cdot (\pi \mathbf{V} q) = 0$$

7. Hydrostatic equation

$$\frac{\partial \phi}{\partial p} = -C_p \theta \quad \left(\text{from } \frac{\partial p}{\partial \phi} = -\rho = -\frac{1}{\alpha} \right)$$

Of the variables $\pi, u, v, T, q, \phi, \alpha, \dot{\sigma}$ we update the 5 primary variables $\pi, u, v, T,$ and q using equations 1, 2, 3.2, 5 and 6. From equations 3.1, 4, and 7 we can obtain $\phi, \alpha,$ and $\dot{\sigma}$ which are our secondary variables. Note that $p = \sigma\pi + p_{\text{top}}$.

Sea level pressure (used only in the smoothing routine SMSHAP)

$$\text{Hydrostatic eq. } \Rightarrow \frac{\partial p}{\partial \phi} = -p = -\frac{1}{\alpha} = -\frac{p}{RT} \therefore \log\left(\frac{\text{SLP}}{p}\right) = -\int_{\phi_s}^0 \frac{d\phi}{RT}$$

$$\therefore \text{SLP} = p(\sigma=1) \exp\left(\frac{\phi_s}{RT}\right)$$

DERIVATION OF THE EQUATIONS AT THE POLES

Consider the continuity equation

$$\frac{\partial \pi}{\partial t} + \nabla \cdot (\pi \mathbf{V}) + \frac{\partial \pi \dot{\sigma}}{\partial \sigma} = 0$$

coupled with a conservation equation $\frac{dT}{dt} = S$ which can be expanded into

$$\frac{\partial \pi T}{\partial t} + \nabla \cdot (\pi \mathbf{V} T) + \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} = \pi S$$

If we integrate this equation over a polar cap of radius $\Delta\phi$

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi T}{\partial t} a^2 \cos\phi d\phi d\lambda = - \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \nabla \cdot (\pi \mathbf{V} T) a^2 \cos\phi d\phi d\lambda$$

$$- \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} a^2 \cos\phi d\phi d\lambda - \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \pi S a^2 \cos\phi d\phi d\lambda$$

and we assume the value of $\frac{\partial \pi T}{\partial t}$ to be approximately constant over the polar cap

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \frac{\partial \pi T}{\partial t} a^2 \cos\phi d\phi d\lambda \approx \left(\frac{\partial \pi T}{\partial t} \right) \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} a^2 \cos\phi d\phi d\lambda$$

$$= 2\pi a^2 (1 - \cos\Delta\phi) \left(\frac{\partial \pi T}{\partial t} \right)_{NP}$$

The first term in the rhs is, using Gauss' theorem

$$- \int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta\phi} \nabla \cdot (\pi \mathbf{V} T) a^2 \cos\phi d\phi d\lambda = - \int_0^{2\pi} \pi v T a \sin\Delta\phi d\lambda$$

This can be approximated as

$$-\frac{2\pi}{IM} \sum_{i=1}^{IM} \pi_i v_i T_i a \sin \Delta \phi$$

The third term on the rhs is

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2 - \Delta \phi} \frac{\partial \pi \dot{\sigma} T}{\partial \sigma} a^2 \cos \phi d\phi d\lambda \approx 2\pi a^2 (1 - \cos \Delta \phi) \left(\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} \right)_{NP}$$

and similarly with the source term.

From these equations we obtain

$$\left(\frac{\partial \pi T}{\partial t} \right)_{\text{Pole}} = -(-1)^m \frac{\cot \Delta \phi / 2}{aIM} \sum_{i=1}^{IM} (\pi_i v_i T_i) \frac{\pi}{2} - \Delta \phi - \left(\frac{\partial \pi \dot{\sigma} T}{\partial \sigma} \right)_{\text{Pole}} + S$$

$$\text{since } \frac{2\pi a \sin \Delta \phi / 2}{IM} \frac{1}{2\pi a^2 (1 - \cos \Delta \phi)} = \frac{2 \sin \Delta \phi / 2 \cos \Delta \phi / 2}{IM a 2 \sin^2 \frac{\Delta \phi}{2}} = \frac{\cot \frac{\Delta \phi}{2}}{aIM}$$

This formulation is used in the continuity, momentum temperature and moisture equations. Note that the first term changes sign in the South Pole ($m=1$). In the momentum equations we make use of the transformation

$$U_i = -\sin \lambda u_i - (-1)^m \cos \lambda v_i$$

$$V_i = (-1)^m \cos \lambda u_i - \sin \lambda v_i$$

where $m=1$ for the South Pole and $m=2$ for the North Pole. U_i, V_i are the "cartesian" velocities at longitude λ_i , and u_i, v_i the corresponding spherical velocities.

The pressure gradient terms are computed making use of Green's

$$\text{Theorem: } \iint \frac{\partial Q}{\partial x} dx dy = \oint Q dy; \iint \frac{\partial P}{\partial y} dx dy = - \oint P dx$$

For example

$$\int_0^{2\pi} \int_{\pi/2}^{\pi/2-\Delta\phi} \frac{\partial \phi}{\partial x} a^2 \cos \phi d\phi d\lambda = -\phi \phi a \Delta\phi (-\cos \lambda d\lambda)$$

$$\phi = \frac{\pi}{2} - \Delta\phi$$

$$\approx + \frac{a\Delta\phi 2\pi}{IM} \sum_{i=1}^{IM} \phi_i \cos \lambda_i$$

In the U-momentum equation we have then the following pressure terms

$$\left(\frac{\partial \pi U}{\partial t}\right)_P = - \frac{\Delta\phi}{a(1-\cos\phi) IM} \sum_{i=1}^{IM} (\phi_i + \left(\frac{\sigma RT}{\rho}\right)_P \pi_i) \cos \lambda_i$$

and similarly for πV .

In the model we have approximated $\Delta\phi \approx 2\sin\frac{\Delta\phi}{2}$

$$\text{Then } \frac{\Delta\phi}{a(1-\cos\phi) IM} \sim \frac{1}{a \sin \frac{\Delta\phi}{2} IM}$$

Based on this formulation we construct the fourth order scheme at the Poles by taking $\frac{4}{3}$ of the differences evaluated at $\Delta\phi$ from the Poles (as expanded here), minus $\frac{1}{3}$ of the differences at $2\Delta\phi$ from the Poles.

This formulation is not conservative at the Poles. However we have found that this has had no noticeable effect in the conservation of mass or energy in the model. In our shallow water experiments we studied a set of equations that were quadratically conservative, but inconsistent at the Pole, and another scheme analogous to the GFDL scheme, which is both quadratically conservative and consistent

at the Poles, but suffers from a serious truncation error near the Poles in the pressure gradient term. The scheme that we chose gave better results than the other two (Kalnay-Rivas, 1976).

Computation of the horizontal pressure gradient as suggested by N. A. Phillips

(1) Let $\theta = \bar{\theta} + \theta'$, $\bar{\theta} = 280^\circ\text{K}/1000^k$ i.e. constant

(2) $\phi = \bar{\phi} + \phi'$, $k = R/C_p = .286$

$$\frac{\partial \bar{\phi}}{\partial p^k} = -C_p \bar{\theta} \quad \& \quad \bar{\phi} = \phi_0 - C_p \bar{\theta} p^k \quad \text{with } \phi_0 = C_p \bar{\theta} 1000^k$$

(3) $T = \bar{T}(p) + T'$, $\bar{T}(p) = \bar{\theta} p^k$

Thus our new dependent variables are

$$\phi' = \phi + C_p \bar{\theta} (p^k - 1000^k)$$

$$T' = T - \bar{\theta} p^k$$

In this way $\pi(\nabla\phi + \frac{\sigma RT}{p} \nabla\pi)$, the pressure gradient in the momentum equations gets transformed into

$$\begin{aligned} (4) \quad & \pi(\nabla_{\phi'} \bar{\phi} + \nabla_{\phi'} \phi' + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi + \frac{\sigma RT'}{p} \nabla\pi) \\ & = \pi(\nabla_{\phi'} \phi' + \frac{\sigma RT'}{p} \nabla\pi) + \pi(\nabla_{\phi'} \bar{\phi} + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi) \end{aligned}$$

But the second parenthesis is zero:

$$\nabla_{\phi'} \bar{\phi} + \sigma R \bar{\theta} \frac{p^k}{p} \nabla\pi = -C_p \bar{\theta} \nabla p^k + R \bar{\theta} \frac{p^k}{p} \nabla p = 0$$

In regions of steep orography, the second parenthesis in (4) is much larger than the first. When the horizontal pressure gradient terms are computed in their original form, the near cancellation of the two terms introduces large truncation errors. The procedure suggested by Phillips greatly reduces this truncation error. We have chosen a simpler definition of ϕ_0 than the one suggested by Phillips.

III. Finite Difference Variables and Grid

The notation used in fourth band model is the same as the standard GLAS(GISS) second order GCM except that we use a non-staggered horizontal grid. A complete description of the variables can be found in the TSANG-KARN documentation of the GISS 9 level model.

Let $u_{ijk}, v_{ijk}, T_{ijk}, q_{ijk}, \pi_{ij}$ be the finite difference approximations to the primary variables $u, v, T, q,$ and π at the mesh point $(i\Delta\lambda, j\Delta\phi, (k-1/2)\Delta\sigma)$. Also the scaled variables $\pi u, \pi v, \dots$ are approximated by $\pi u_{ijk}, \pi v_{ijk}, \dots$.

The finite difference equations also use the following geometric arrays:

$$DXP(j) = m_j = a \cos \phi_j \Delta\lambda, \quad DYU(j) = n_j = a \Delta\phi_j,$$

We use a factor of 12 in DXYP to make our scaled fourth order differences simpler.

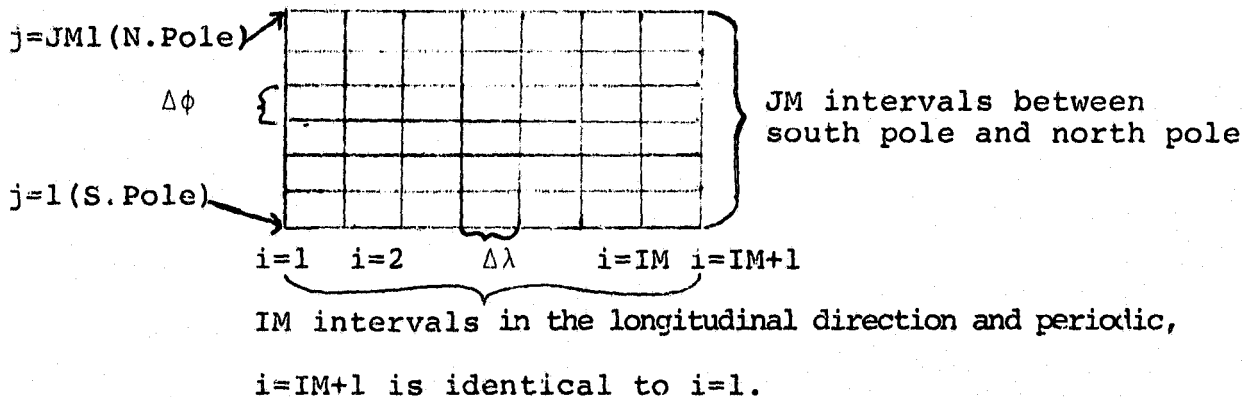
$$DXYP(j) = 12 \cdot n_j \cdot n_j, \quad \Pi_{ij} = DXYP(j) \cdot \pi_{ij}, \quad U_{ijk} = n_j \pi_{ij} U_{ijk}$$

$$\dot{S}_{ijk} = \Pi_{ij} \cdot \dot{\sigma}_{ijk}, \quad ADLDP = 12 \cdot a \cdot \Delta\lambda \Delta\phi, \quad V_{ijk}^* = m_j \pi_{ij} V_{ijk}$$

$$F_{ijk} = DXYP(j) f_j + ADLDP \cdot \sin \phi_j U_{i\&j}$$

Horizontal Grid

The fourth order band model uses an unstaggered grid in the horizontal direction.



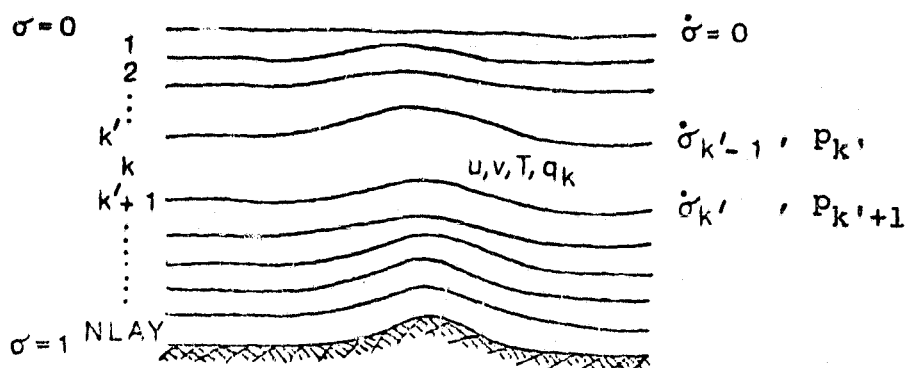
$$\Delta\lambda = \text{DLON} = \frac{\pi}{JM} \quad \Delta\phi = \text{DLAT} = \frac{\pi}{JM}$$

$$\lambda_i = (i-1)\Delta\lambda - \pi = (i-1)5^\circ - 180^\circ \quad \phi_j = (j-1)\Delta\phi - \frac{\pi}{2} = (j-1)4^\circ - 90^\circ$$

$$\text{JSP} = 1 \quad \text{JNP} = \text{JM}1 = \text{JM}+1$$

Vertical Grid

The vertical grid is staggered; the values of all the variables u, v, T, q, π, \dots except $\dot{\sigma}$ are computed at the center of each layer. The values of $\dot{\sigma}$ are computed at the edges of the layers.



In the case of uniform vertical resolution and $\text{NLAY}=9$ vertical layers,

$$\sigma_k = \text{SIG}(K) = \frac{k-1}{9} (0, 1/9, \dots, 1)$$

$$\sigma_{k'} = \text{SIGE}(K) = \frac{2k-1}{18} = \left(\frac{1}{18}, \dots, \frac{17}{18}\right)$$

$$\text{and } \Delta\sigma = 1/9.$$

$\dot{\sigma}_{ijk}$ and its scaled version \dot{S}_{ijk} are defined at the eight interior edges, i.e. for $\sigma_{k'}$ with $k=1$ to 8 since $\dot{\sigma}(0)=\dot{\sigma}(1)=0$ from the boundary conditions. The pressure $p_{k'} = \pi\sigma_{k'} + p_{\text{TOP}}$ is defined at the same level as $\dot{\sigma}_{k'-1}$. Note for level k we need \dot{S}_{ijk} and \dot{S}_{ijk-1} to form the second order vertical differences:

$$\dot{S}_{ijk} (v_{ijk} + v_{ijk+1}) - \dot{S}_{ijk-1} (v_{ijk} + v_{ijk-1})$$

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III-2 Periodic Filtering of Short Waves

An integral part of the numerical scheme is the periodic application (every ISMTH steps, generally 2 hours) of a 16th order Shapiro filter. This has the effect of removing waves shorter than $4\Delta x$, which are not resolved in the model, while waves longer than $4\Delta x$, which are accurately computed by the difference scheme, are not affected by the filter (Kalnay-Rivas and Hoitsma, 1979).

The filter is applied to an array q_j in the following way:

$$\text{Let } d_+ (q_j) = q_{j+1} - q_j, \quad d_- (q_j) = q_j - q_{j-1}$$

Then a Shapiro filter of order $2N$ is given by

$$\bar{q}_j = q_j - (-1)^N (d_+ d_-)^N q_j$$

The response of the filter applied to a wave of the form

$$q_j = Q \exp(i \frac{2\pi}{L} \Delta x j) \text{ is}$$

$$\bar{q}_j = (1 - \sin^{2N} \frac{\pi \Delta x}{L}) q_j$$

The 2-dimensional filter is applied as a product of 1-dimensional filters (first in longitude, then in latitude). In latitude we filter the fields on great circles formed by meridians of longitude λ and $\lambda + \pi$, where $0 \leq \lambda < \pi$. We are presently filtering only potential temperature and sea level pressure. These fields were chosen because they are not very affected by orography. Winds are not currently filtered, because the adjustment between mass and velocity fields does not allow the development of short waves in the winds alone. However, in the tropics, where the adjustment of winds to the mass field is minimal, the winds are somewhat noisy, and we may opt to filter them too.

IV. Boundary Conditions

Periodicity in the Zonal (East-West) direction

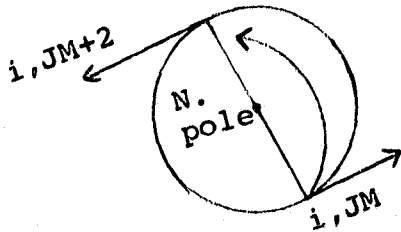
$$\pi_{IM+mj} = \pi_{mj} \quad m=1,2,3,\dots, \quad j=2,\dots,JM$$

$$Q_{IM+mjk} = Q_{mjk} \quad m=1,2,3,\dots, \quad j=2,\dots,JM, \quad k=1,\dots,NLAY$$

for $Q=u,v,T,q,\phi,\delta,\pi u,\pi v,\dots$

Boundary conditions at the north and south poles. Define the array

INDEX as follows:



$$INDEX(i) = i + \frac{IM}{2} \quad i=1,2,\dots,\frac{IM}{2}$$

$$INDEX(i + \frac{IM}{2}) = i \quad i=\frac{IM}{2}+1,\dots,IM$$

i.e., INDEX: $\frac{IM}{2}+1, \frac{IM}{2}+2, \dots, IM, 1, 2, \dots, \frac{IM}{2}$

Then we can define

$$\pi_{iJM+2} = \pi_{INDEX(i)JM} \quad \text{for the points needed "beyond" the North Pole,}$$

and $\pi_{i0} = \pi_{INDEX(i)2}$ for the points "beyond" the South Pole

and similarly for T,q,ϕ .

For the horizontal velocity $V=(u,v)$ we have

$$V_{iJM+2k} = -V_{iJMk}$$

$$V_{i0k} = -V_{i2k}$$

V. Finite Difference Equations

THE ZONAL (U) MOMENTUM EQUATION

$$\frac{\partial \pi u}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial (\pi u \cdot u)}{\partial \lambda} + \frac{\partial (\pi v \cos \phi \cdot u)}{\partial \phi} \right] - \frac{\partial (\pi \dot{\sigma} u)}{\partial \sigma}$$

$$- \frac{\pi}{a \cos \phi} \left[\frac{\partial \phi}{\partial \lambda} + \frac{\sigma RT}{p} \frac{\partial \pi}{\partial \lambda} \right] + \left(f + \frac{u \tan \phi}{a} \right) \pi v + \pi F_x$$

$$\frac{\partial \pi U_{ijk}}{\partial t} = \left\{ \begin{array}{l} \text{COMP1: HA: } \overbrace{PU1_{i-1/2j}}^{(U_{ijk}^* + U_{i-1jk}^*)} (U_{ijk} + U_{i-1jk}) - \overbrace{PU1_{i+1/2j}}^{(U_{ijk}^* + U_{i+1jk}^*)} (U_{ijk} + U_{i+1jk}) \\ - .5 \left[\overbrace{PU2_{i-1}}^{(U_{ijk}^* + U_{i-2jk}^*)} (U_{ijk} + U_{i-2jk}) - \overbrace{PU2_{i+1}}^{(U_{ijk}^* + U_{i+2jk}^*)} (U_{ijk} + U_{i+2jk}) \right] \\ + 4 * \left[\overbrace{PV1}^{(V_{ijk}^* + V_{ij-1k}^*)} (U_{ijk} + U_{ij-1k}) - \overbrace{(V_{ijk}^* + V_{ij+1k}^*)}^{PV1} (U_{ijk} + U_{ij+1k}) \right] \\ - .5 * \left[\overbrace{(V_{ijk}^* + V_{ij-2k}^*)}^{PV2} (U_{ijk} + U_{ij-2k}) - \overbrace{(V_{ijk}^* + V_{ij+2k}^*)}^{PV2} (U_{ijk} + U_{ij+2k}) \right] \\ \text{VA: } \\ + .5 * \left[\dot{S}_{ijk-1} (U_{ijk} + U_{ijk-1}) - \dot{S}_{ijk} (U_{ijk} + U_{ijk+1}) \right] / \Delta \sigma_k \\ \text{COMP2: PG: } \\ + (\pi_{ij} * n_j * \left\{ 8 * [\phi'_{i-1jk} - \phi'_{i+1jk} + \frac{\sigma_k RT'_{ijk}}{p_k} (\pi_{i-1j} - \pi_{i+1j})] + \right. \\ \left. [\phi'_{i+2jk} - \phi'_{i-2jk} + \frac{\sigma_k RT'_{ijk}}{p_k} (\pi_{i+2j} - \pi_{i-2j})] \right\}) \\ \text{C: } \\ + F_{ijk} * \pi_{ij} * V_{ijk} \end{array} \right\} + (\pi F_x)$$

Note: COMP1, COMP2, COMP3 are the names of the three subroutines where the different terms are computed. HA: Horizontal advection terms. VA: Vertical advection terms. PG: Pressure gradient terms. C: Coriolis term.

Also note that PV2 is set equal to zero for $j=0$ and $j=JM$, i.e., there is no transport of mass across the poles.

THE MERIDIONAL (V) MOMENTUM EQUATION

$$\frac{\partial \pi V}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial (\pi u \cdot v)}{\partial \lambda} + \frac{\partial (\pi v \cos \phi v)}{\partial \phi} \right] - \frac{\partial \pi \dot{\sigma} v}{\partial \sigma}$$

$$- \frac{\pi}{a} \left[\frac{\partial \phi}{\partial \phi} + \frac{\sigma RT}{p} \frac{\partial \pi}{\partial \phi} \right] - \left(f + \frac{u \tan \phi}{a} \right) \pi u + \pi F_y$$

COMP1: HA:

$$\frac{\partial \pi V_{ijk}}{\partial t} = \left\{ 4. * [(U_{ijk}^* + U_{i-1jk}^*) (V_{ijk} + V_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (V_{ijk} + V_{i+1jk})] \right.$$

$$- .5 [(U_{ijk}^* + U_{i-2jk}^*) (V_{ijk} + V_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (V_{ijk} + V_{i+2jk})]$$

$$+ 4. * [(V_{ijk}^* + V_{ij-1k}^*) (V_{ijk} + V_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (V_{ijk} + V_{ij+1k})]$$

$$- .5 * [(V_{ijk}^* + V_{ij+2k}^*) (V_{ijk} + V_{ij-2k}) - (V_{ijk}^* + V_{ij+2k}^*) (V_{ijk} + V_{ij+2k})]$$

VA:

$$+ .5 * [\dot{S}_{ijk-1} (V_{ijk} + V_{ijk-1}) - \dot{S}_{ijk} (V_{ijk} + V_{ijk+1})] / \Delta \sigma_k \}$$

COMP2: P:

$$+ (\pi_{ij} * m_j * \{ 8. * [\phi_{ij-1k}^* - \phi_{ij+1k}^* + \frac{\sigma_k RT_{ijk}^*}{p_k} (\pi_{ij-1} - \pi_{ij+1})]$$

$$+ [\phi_{ij+2k}^* - \phi_{ij-2k}^* + \frac{\sigma_k RT_{ijk}^*}{p_k} (\pi_{ij+2} - \pi_{ij-2})] \}$$

C:

COMP3:

$$- F_{ijk} * \pi_{ij} * U_{ijk} + (\pi F_y)$$

THE THERMODYNAMIC ENERGY EQUATION

$$\frac{\partial \pi T}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial \pi U T}{\partial \lambda} + \frac{\partial \pi V \cos \phi \cdot T}{\partial \phi} \right] - p^k \frac{\partial \pi \dot{\sigma} T / p^k}{\partial \sigma}$$

$$+ \frac{\pi \sigma K T}{p} \left(\frac{\partial \pi}{\partial t} + \frac{U}{a \cos \phi} \frac{\partial \pi}{\partial \lambda} + \frac{V}{a} \frac{\partial \pi}{\partial \phi} \right) + \frac{\pi Q}{C_p}$$

COMP1: HA:

$$\frac{\partial \pi T_{ijk}}{\partial t} = \{ 4 * [(U_{ijk}^* + U_{i-1jk}^*) (T_{ijk} + T_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (T_{ijk} + T_{i+1jk})]$$

$$- .5 * [(U_{ijk}^* + U_{i-2jk}^*) (T_{ijk} + T_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (T_{ijk} + T_{i+2jk})]$$

$$+ 4 * [(V_{ijk}^* + V_{ij-1k}^*) (T_{ijk} + T_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (T_{ijk} + T_{ij+1k})]$$

$$- .5 * [(V_{ijk}^* + V_{ij-2k}^*) (T_{ijk} + T_{ij+2k}) - (V_{ijk}^* + V_{ij+2k}^*) (T_{ijk} + T_{ij+2k})] \}$$

VA:

$$+ .5 p_{ijk}^k \left[\dot{s}_{ijk-1} \left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk-1}}{p_{ijk-1}^k} \right) - \dot{s}_{ijk} \left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk+1}}{p_{ijk+1}^k} \right) \right] / \Delta \sigma_k \}$$

COMP2:

$$\left(+ \frac{\sigma_k K T_{ijk}}{p_{ijk}} \left[\pi_{ij} \frac{\partial \pi}{\partial t} + U_{ijk}^* \{ 8 * [\pi_{i+1j} - \pi_{i-1j}] + \pi_{i-2j} - \pi_{i+2j} \} \right. \right.$$

$$\left. \left. + V_{ijk}^* \{ 8 * [\pi_{ij+1} - \pi_{ij-1}] + \pi_{ij-2} - \pi_{ij+2} \} \right] \right)$$

COMP3:

$$+ \left(\frac{\pi Q_{ijk}}{C_p} \right)$$

THE MOISTURE EQUATION

$$\frac{\partial \pi q}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial \pi U \cdot q}{\partial \lambda} + \frac{\partial \pi V \cos \phi \cdot q}{\partial \phi} \right] - \frac{\partial \pi \delta q}{\pi \sigma} + \pi (E - C)$$

COMP1: HA

$$\begin{aligned} \frac{\partial \Pi q_{ijk}}{\partial t} = & \{ 4. * [(U_{ijk}^* + U_{i-1jk}^*) (q_{ijk} + q_{i-1jk}) - (U_{ijk}^* + U_{i+1jk}^*) (q_{ijk} + q_{i+1jk})] \\ & - .5 * [(U_{ijk}^* + U_{i+2jk}^*) (q_{ijk} + q_{i-2jk}) - (U_{ijk}^* + U_{i+2jk}^*) (q_{ijk} + q_{i+2jk})] \\ & + 4. * [(V_{ijk}^* + V_{ij-1k}^*) (q_{ijk} + q_{ij-1k}) - (V_{ijk}^* + V_{ij+1k}^*) (q_{ijk} + q_{ij+1k})] \\ & - .5 * [(V_{ijk}^* + V_{ij-2k}^*) (q_{ijk} + q_{ij-2k}) - (V_{ijk}^* + V_{ij+2k}^*) (q_{ijk} + q_{ij+2k})] \\ & + .5 * [\dot{S}_{ijk-1} (q_{ijk} + q_{ijk-1}) - \dot{S}_{ijk} (q_{ijk} + q_{ijk+1})] / \Delta \sigma_k \} \end{aligned}$$

COMP3:

$$+ \Pi (E_{ijk} - C_{ijk})$$

Note: The current transport scheme for the moisture field is being modified.

THE PRESSURE-TENDENCY EQUATION

$$\frac{\partial \pi}{\partial t} = - \sum_{\ell=1}^L \frac{(\Delta \sigma)_{\ell}}{a \cos \phi} \left[\frac{\partial \pi u}{\partial \lambda} + \frac{\partial \pi v \cos \phi}{\partial \phi} \right]$$

COMPl:

$$\begin{aligned} \text{CONV}_{i\ell j} = & \Delta \sigma_{\ell} \{ 8. * (U_{i-1j\ell}^* - U_{i+1j\ell}^*) + U_{i+2j\ell}^* - U_{i-2j\ell}^* \\ & + 8. * (V_{ij-1\ell}^* - V_{ij+1\ell}^*) + V_{ij+2\ell}^* - V_{ij-2\ell}^* \} \end{aligned}$$

COMPl:

$$\frac{\partial \pi_{ij}}{\partial t} = \left(\sum_{\ell=1}^L \text{CONV}_{i\ell j} \right)$$

THE VERTICAL VELOCITY EQUATION

$$\frac{\partial \pi}{\partial t} = - \frac{1}{a \cos \phi} \left[\frac{\partial \pi u}{\partial \lambda} + \frac{\partial \pi v \cos \phi}{\partial \phi} \right] - \frac{\partial}{\partial \sigma} (\pi \dot{\sigma})$$

Thus

$$\Delta \sigma_k \frac{\partial \pi}{\partial t} = \text{CONV}_{ikj} - \Delta \dot{S}_{ijk}$$

giving:

COMPl:

$$\dot{S}_{ijk} = \{ \dot{S}_{ijk-1} + \text{CONV}_{ikj} - \Delta \sigma_k \frac{\partial \pi}{\partial T} \}$$

VI. The Forecast Equations at the Poles

The fourth order band model uses a spherical cap at the poles, and the finite difference approximations to the equations of motion must be derived for this spherical region. Stereographic projection is used to give us a well-defined velocity vector at the poles.

From trigonometry, (Fig. VIa) if the vector (U, V) in X - Y coordinates is represented by (U', V') in X' - Y' coordinates where the prime coordinate axes are rotated by an angle λ then,

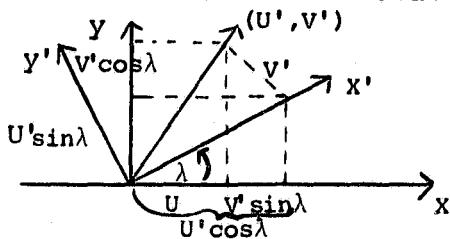


Fig. VIa

From Fig. VIb, which shows a unit vector in both coordinate systems, we see that, at the North Pole, spherical coordinates are rotated by an angle $(\frac{\pi}{2} + \lambda)$ with respect to the polar stereographic coordinates. Therefore,

$$U_{NP} = U \cos(\lambda + \frac{\pi}{2}) - V \sin(\lambda + \frac{\pi}{2})$$

$$V_{NP} = U \sin(\lambda + \frac{\pi}{2}) + V \cos(\lambda + \frac{\pi}{2})$$

or

$$\begin{aligned} (3) \quad U_{NP} &= -U \sin \lambda - V \cos \lambda \\ V_{NP} &= U \cos \lambda - V \sin \lambda \end{aligned}$$

Similarly, for the South Pole:

$$\begin{aligned} (4) \quad U_{SP} &= -U \sin \lambda + V \cos \lambda \\ V_{SP} &= -U \cos \lambda - V \sin \lambda \end{aligned}$$

also,

$$\begin{aligned} (5) \quad U &= -U_{NP} \sin \lambda + V_{NP} \cos \lambda \\ V &= -U_{NP} \cos \lambda - V_{NP} \sin \lambda \end{aligned}$$

$$\begin{aligned} (6) \quad U &= -U_{SP} \sin \lambda - V_{SP} \cos \lambda \\ V &= U_{SP} \cos \lambda - V_{SP} \sin \lambda \end{aligned}$$

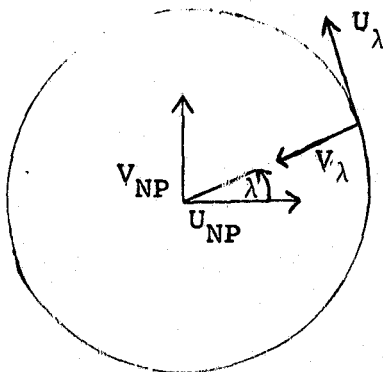


Fig. VIb

The initial values for U_{NPK} , V_{NPK} , U_{SPK} , V_{SPK} are obtained from equations 3 and 4 by averaging in the zonal direction on the line of latitude $j=45$ for the north pole and $j=2$ for the south pole. In the forecast stereographic velocities are advected and transformed back into spherical velocities after each time step by equations 5 and 6.

In the following equation we will denote the polar velocities by $VP_{k,m}$, $VP_{k,m}$ and the temperature and other variables in a similar way, $TP_{k,m}$, $\pi P_{k,m}$, ... where $m=1$ for the south pole and $m=2$ for the north pole. The following constants are used,

$$COEF1 = (-1)^m \quad COEF2 = -COEF1 \quad RADIM = 3 \cdot a \cdot IM$$

$$CON1 = 4COT(.5\Delta\phi) / RADIM$$

$$CON2 = -COT(\Delta\phi) / RADIM$$

$$CON3 = 4DT / (RADIM * SIN(.5\Delta\phi))$$

$$CON4 = DT / (RADIM * SIN(\Delta\phi))$$

$$JPOL(K,M) = \begin{pmatrix} 2 & JM \\ 3 & JM-1 \end{pmatrix}$$

$r = JPOL(1,m)$ = first interior value of j (2 for the S. Pole, JM for the N. Pole)

$s = JPOL(2,m)$ = second interior value of j (3 for the S. Pole, $JM-1$ for the N. Pole).

ZONAL (U) MOMENTUM EQUATION (POLES)

$$\begin{aligned}
 \frac{\partial \pi UP_{k,m}}{\partial t} = & \text{COMP1:} \\
 & \text{IM} \\
 & \{ \text{COEF1} * [\text{CON1} * \sum_{i=1}^{\text{IM}} -\pi_{ir} V_{irk} (U_{irk} \text{Sin} \lambda_i + \text{COEF1} * V_{irk} \text{COS} \lambda_i) \\
 & + \text{CON2} * \sum_{i=1}^{\text{IM}} -\pi_{is} V_{isk} (U_{isk} \text{Sin} \lambda_i + \text{COEF1} * V_{isk} \text{COS} \lambda_i)] \\
 & + .5 [\dot{S}P_{k-1} (UP_k + UP_{k-1}) - \dot{S}P_k (UP_k + UP_{k+1})] / \Delta \sigma_k \} \\
 & \text{COMP2:} \\
 & + \pi P_m \cdot [-\frac{\text{CON3}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{irk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{ir}) \text{COS} \lambda_i \\
 & + \frac{\text{CON4}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{isk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{is}) \text{COS} \lambda_i] \\
 & \text{COMP3:} \\
 & + f_m \pi P_m \cdot VP_{k,m} + (\pi P_m \cdot FP_{k,m}
 \end{aligned}$$

MERIDIONAL (V) MOMENTUM EQUATION (POLES)

$$\begin{aligned}
 \frac{\partial \pi VP_{k,m}}{\partial t} = & \text{COMP1:} \\
 & \text{IM} \\
 & \{ \text{COEF1} * [\text{CON1} * \sum_{i=1}^{\text{IM}} -\pi_{ir} V_{irk} (\text{COEF2} * U_{irk} * \text{COS} \lambda_i + V_{irk} \text{Sin} \lambda_i) \\
 & + \text{CON2} * \sum_{i=1}^{\text{IM}} -\pi_{is} V_{isk} (\text{COEF2} * U_{isk} * \text{COS} \lambda_i + V_{isk} \text{Sin} \lambda_i)] \\
 & + .5 [\dot{S}P_{k-1} (VP_k + VP_{k-1}) - \dot{S}P_k (VP_k + VP_{k+1})] / \Delta \sigma_k \} \\
 & \text{COMP2:} \\
 & + \pi P_m \cdot [-\frac{\text{CON3}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{irk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{ir}) \text{Sin} \lambda_i \\
 & + \frac{\text{CON4}}{\text{DT}} * \sum_{i=1}^{\text{IM}} (\phi'_{isk} + \frac{\sigma_k \text{RTP}'_k}{pP_k} \pi_{is}) \text{Sin} \lambda_i] \\
 & \text{COMP3:} \\
 & - f_m \pi P_m \cdot UP_{k,m} + (\pi P_m \cdot SP_{k,m})
 \end{aligned}$$

THE THERMODYNAMICS ENERGY EQUATION (POLES)

COMP1:

$$\frac{\partial \pi TP_{k,m}}{\partial t} = \{ \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} T_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk} T_{isk})$$

$$+ .5 p_{k,m}^k * \{ \dot{S}_{k-1,m} \left(\frac{TP_{k,m}}{p_{k,m}^k} + \frac{TP_{k-1,m}}{p_{k-1,m}^k} \right)$$

$$- \dot{S}_{k,m} \left(\frac{TP_{k,m}}{p_{k,m}^k} + \frac{TP_{k+1,m}}{p_{k+1,m}^k} \right) \} / \Delta \sigma_k \}$$

COMP2:

$$+ \left(\frac{\pi_{m,k}^p \sigma_k K_{TP_{k,m}}}{p_{k,m}^k} \right) \left\{ \frac{\partial \pi_{m,k}}{\partial t} + \text{CON5} * \sum_{i=1}^{\text{IM}} [(\text{COEF1} * U_{P_{k,m}} \cos \lambda_i + V_{P_{k,m}} \sin \lambda_i) * \right.$$

COMP3:

$$\left. \left(\frac{\pi_{m,k}^p}{C_p} Q_{P_k} \right) \right\} (8\pi_{ir} - \pi_{is}) \}$$

THE MOISTURE BALANCE EQUATION (POLES)

COMP1:

$$\frac{\partial \pi q_{k,m}}{\partial t} = \{ \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} q_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk} q_{isk})$$

$$+ .5 [\dot{S}_{k-1,m} (q_{k,m}^p + q_{k-1,m}^p) - \dot{S}_{k,m} (q_{k,m}^p + q_{k+1,m}^p)] / \Delta \sigma_k \}$$

COMP3:

$$+ (\pi_{m,k}^p * EP_k)$$

THE PRESSURE TENDENCY EQUATION (POLES)

$$\text{CONVPL}_{k,m} = \text{COEF1} * (\text{CON1} * \sum_{i=1}^{\text{IM}} \pi_{ir} V_{irk} + \text{CON2} * \sum_{i=1}^{\text{IM}} \pi_{is} V_{isk}) \Delta\sigma_k$$

$$\frac{\partial \pi P_m}{\partial t} = \sum_{k=1}^{\text{NLAY}} \text{CONVPL}_{k,m}$$

THE VERTICAL VELOCITY EQUATION (POLES)

$$\dot{S}P_{k,m} = \dot{S}P_{k-1,m} + \text{CONVPL}_k - \Delta\sigma_k \frac{\partial \pi P_m}{\partial t}$$

VII. Diagnostic Equations ($\phi, \dot{\sigma}, p$)

Once the updated values of πU , πV , πT , πq are found we unscale:

$$U_{ijk}^{n+1} = \pi U_{ijk}^{n+1} / \pi_{ij}^{n+1} \quad \text{for all } i, j, k.$$

Similarly for V, T, q . We also filter the fields near the poles to prevent linear instability (see subroutine AVRX). $\dot{\sigma}$ is obtained from \dot{S} by unscaling also.

We determine from π_{ij}^{n+1} and σ_k , p_{ijk}^{n+1}

$$p_{ijk}^{n+1} = \sigma_k \pi_{ij}^{n+1} + p_{TOP} \quad p_{TOP} = \text{constant.}$$

ϕ_S is the surface geopotential (a function only of latitude and longitude)

For the Phillips geopotential we define $(p)^k$ at the center of the layer in the following way:

$$(p)_{ijk}^k = \frac{p_{ijk+1}^{k+1} - p_{ijk}^{k+1}}{(k+1)(p_{ijk+1}^{k+1} - p_{ijk}^{k+1})} \quad \left(p^k = \frac{1}{k+1} \frac{\partial p^{k+1}}{\partial p} \right)$$

$k=1, 2, \dots, N_{LAY}$

where

$$p_{ijk}^{k+1} = (\text{SIGE}(k) * \pi_{ij} + p_{TOP})^{k+1} \quad (\text{SIGE}(k) = \sigma_k'), \text{ i.e.}$$

p^{k+1} is obtained by exponentiation and $(p)^k$ by differences.

The following equations represent the geopotential calculations used in the old fourth order model.

Let

$$C_{ijk} = \frac{\pi_{ij} \sigma_k R \Delta \sigma_k}{p_{ijk}} - \frac{C_p}{2} \left[\sigma_k \left(\frac{p_{ijk+1}^k}{p_{ijk}^k} - 1 \right) + \sigma_{k-1} \left(1 - \frac{p_{ijk-1}^k}{p_{ijk}^k} \right) \right]$$

for $k=1, \dots, NLAY$

with

$$p_{ij0}^k = p_{ij1}^k \text{ and } p_{ijNLAY+1}^k = p_{ijNLAY}^k$$

An optimized version of C_{ijk} is:

$$C_{ijk} = \frac{\pi_{ij} \sigma_k R \Delta \sigma_k}{p_{ijk}} - \frac{.5 C_p}{p_{ijk}^k} \left\{ \Delta \sigma_k (p_{ijk+1}^k - p_{ijk}^k) + \sigma_{k-1} (p_{ijk+1}^k - p_{ijk-1}^k) \right\}$$

Rather than compute ϕ and then subtract $\bar{\phi}$ we do everything at once:

$$(1) \quad \phi'_{ijNLAY} = \phi'_S - CPTH * (PSKAPA - p_{ijNLAY}^k) + \sum_{\ell=1}^{NLAY} C_{ij\ell} T_{ij\ell}$$

$$(2) \quad \phi'_{ij\ell} = \phi'_{ij\ell+1} + \frac{C_p}{2} (p_{ij\ell+1}^k - p_{ij\ell}^k) \left(\left(\frac{T_{ijk}}{p_{ijk}^k} + \frac{T_{ijk+1}}{p_{ijk+1}^k} \right) - 2\bar{\theta} \right)$$

where

$$CPTH = C_p \cdot \bar{\theta} \quad PSKAPA = 1000^k = p_S^k$$

VIII. The Time Differencing Scheme

The fourth order band model has the option of using the Matsuno time scheme or the smooth leapfrog scheme (see MWR-Vol 100 (487-490) R. Asselin).

Let Q^n represent a typical variable that is to be updated to time $n+1$, and let $D(Q^n)$ represent the nonlinear space differences. The Matsuno (Euler-backward) scheme is as follows:

$$\begin{aligned}\tilde{Q} &= Q^n + \Delta t D(Q^n) \\ Q^{n+1} &= Q^n + \Delta t D(\tilde{Q})\end{aligned}$$

The standard leapfrog scheme is given by

$$(1) \quad \frac{Q^{n+1} - Q^{n-1}}{2\Delta t} = D(Q^n)$$

For the smooth leapfrog scheme we replace Q^{n-1} by \bar{Q}^{n-1}

$$(2) \quad Q^{n+1} = \bar{Q}^{n-1} + 2\Delta t D(Q^n)$$

with

$$(3) \quad \bar{Q}^n = (1-\nu)Q^n + .5\nu(\bar{Q}^{n-1} + Q^{n+1})$$

Equation (3) represents a simple time filter except \bar{Q}^{n-1} is used instead of Q^{n-1} in order to save core storage. The above equations (2) and (3) represent the order in which the smooth leapfrog scheme is evaluated. For $n=1$ we define $\bar{Q}^0 = Q^0$ then we update in equation (2) followed by the filtering in equation (3) which is needed for the next time step.

The smoothing step introduces dissipation with respect to time, controlled explicitly by the parameter ν , as compared to the implicit dissipation in the Matsuno scheme. The amplification factor can be found in the paper by Asselin.

A further modification must be made to the smooth leapfrog scheme when source terms are included. Essentially the idea is that we must include the source term effect over two steps rather than one. If we do not do this, then the source effects (COMP3) are included only in every other step which will introduce large discretization errors. (For details see the attached report, Appendix B.)

If the source terms are called every NCOMP3 steps, then for step $n = \text{NCOMP3}$

$$Q_*^n = \bar{Q}^{n-2} + 2\Delta t D(Q^{n-1})$$

$$\bar{Q}_*^{n-1} = (1-\nu) Q^{n-1} + .5\nu(\bar{Q}^{n-2} + Q^n)$$

Then compute the source terms S^n and include in both steps n and $n-1$,

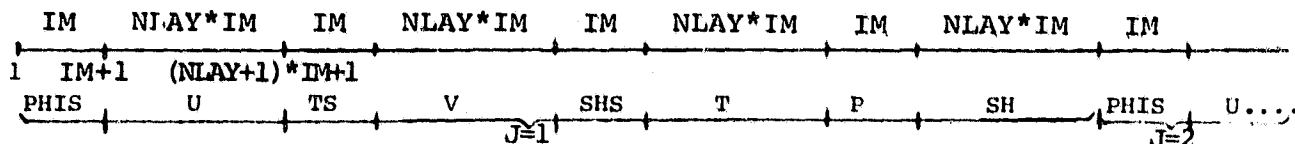
$$Q^n = Q_*^n + \text{NCOMP3} \cdot \Delta t S^n$$

$$\bar{Q}^{n-1} = \bar{Q}_*^{n-1} + \text{NCOMP3} \cdot \Delta t S^n$$

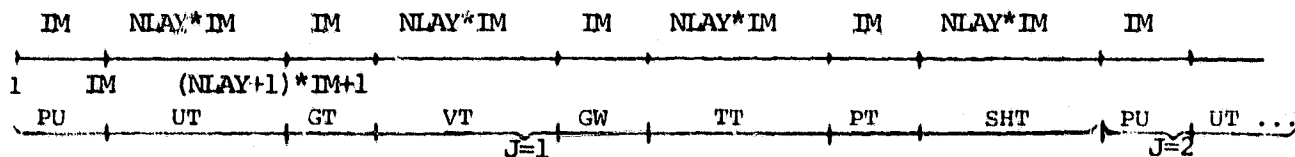
The actual code is complicated by the fact that we actually use scaled and unscaled variables, but the generalization is straightforward.

IX. Documentation of the Code (Preliminary)

The band fourth model uses special equivalences and nonstandard dimensions in order to have the variables P,U,V,T,SH stored in contiguous locations for each line of latitude. Thus, we desire to have the variables stored as follows.



The scaled variables PT, UT, VT, TT, SHT are to be stored as follows:



The above storage designation is accomplished by dimensioning PHIS(4*(NLAY+1)*IM,JNP) instead of PHIS(IM,JNP) (for a fine grid we have PHIS(2880,46), for ultrafine we have PHIS(4800,72)). Then we equivalence NLAY*IM locations of the first line of latitude of U with PHIS(IM+1,1) to PHIS((NLAY+1)*IM,1) (the first IM locations of PHIS are used for the quantity PHIS itself).

Similarly, we equivalence TS(1,1) to TS(IM,1) with PHIS((NLAY+1)*IM+1,1) to PHIS((NLAY+2)*IM,1) and so forth. (See the enclosed computer codes for the exact values in the fine and ultrafine versions.)

In order to have the successive lines of latitude arranged properly in storage we dimension our variables to achieve this purpose. Thus, we have $U(IM, 4*(NLAY+1), 1)$ instead of $U(IM, NLAY, JM)$, similarly for V, T, SH, UT, VT, TT , and SHT . For P we have $P(4*(NLAY+1)*IM, 1)$ instead of $P(IM, JNP)$ and similarly for TS, SHS, GT, GW , and PT . Note that the variables U, UT, \dots are only computed at $IM*NLAY*JNP$ points, the special dimensions are needed to properly align the variables in storage. It is important to recall that we are using two properties of the FORTRAN Compiler:

First, by equivalencing two elements of two different arrays we implicitly equivalence the other elements of the arrays. Second, that last array dimension can be left as 1 as long as it is dimensioned properly in the calling routine or it is equivalenced to a properly dimensioned array. Note we could have dimensioned U as $U(IM, 4*(NLAY+1), JNP)$ we would use the same amount of storage. It is crucial to have the $IM*(NLAY+1)$ dimension because it would cause the computations on U for the successive lines of latitude to be shifted $IM*(NLAY+1)$ locations where the next line of latitude of U are stored. For clarity and simplicity in programming we use the standard equivalence of U, V, T, SH with $Q(I, L, N, J)$ with U equivalent to $Q(I, L, 1, J)$ and so forth. Since we want P, U, V, T , and SH in contiguous storage locations and the equivalence of Q with U, V, T , and SH , then we must dimension Q as $Q(IM, NLAY+1, 4, 1)$ instead of $Q(IM, NLAY, 4, 1)$. We fill in the extra locations by including $PHIS, TS, SH$, and P .

COMPØ Description

This subroutine contains the time schemes and controls the calling sequences of the routines COMP1, COMP2, COMP3, and the polar filtering routine AVRX.

The logic of this subroutine involves three considerations. First, it permits one to choose either the Matsuno or the smooth leapfrog time scheme. Second, the latter scheme involves the use of storage arrays PSM and QSM. Third, the unscaling, smoothing of the updated variables QT and the call of subroutine COMP3 occur at the value $JS2=J-2$ when J is the value of the current line of latitude which is being computed. The routine COMP1 and COMP2 require the values of U,V,... at JS2 in order to compute the updated values at J because fourth order differences are used in the meridional (J) direction. Only after the COMP1 and COMP2 are called for value J can we unscale, smooth, and finish processing the variables at JS2.

The code for the poles is identical in format with the code for the other J values except the variables are scaled by the pressure PPOl only.

The main program contains two calls to COMPØ which cannot be treated independently because of the calling sequence: COMPØ(Q,QT), COMPØ(QT,Q). If the first call is a leapfrog step (LF), the second one must be also LF. If the first call is Matsuno predictor (MP), the second one must be Matsuno corrector (MC). This is represented symbolically by $LF \rightarrow LF$, or $MP \rightarrow MC$. The second call can be followed by either MP or LF. Each of these combinations requires different transfers.

Description of the Time Step Sequence Parameters

NSTEP: Counts the time steps. Starts and restarts both begin with **NSTEP=0**.

NSEQ: The number of steps (combined matsuno and leapfrog) in each (repeated) sequence of time steps.

MLF(I): **MLF(I)=0** or **1** according to whether the **I**th step in the sequence is Leapfrog or Matsuno, respectively. First step is always Matsuno (**MLF(1)=1**).

ISMTH,) Smoothing routine (**SMSHAP**) is called **MOD(NSTEP-NSM1,**
NSM1:) **ISMTH)**; if **ISMTH=0**, there is no smoothing.

NCOMP3,) Physics routine (**COMP3**) is called **MOD (NSTEP-NCM1,**
NCM1:) **NCOMP3)**.

Sample Runs

Matsuno only:

NSEQ=1, MLF(1)=1, MATSUN=1, DT=750., NSM1=0, NCM1=0
BCINO3=4, ISMTH=8

Leapfrog only:

NSEQ=1, MLF(1)=1, MATSUN=0, DT=600., NSM1=0, NCM1=0
NCOMP3=5, ISMTH=10

1 Matsuno, 4 Leap-Frog:

NSEQ=5, MLF=(1,0,0,0,0), MATSUN=not needed, NSM1=0,
NCM1=0, DT=600., NCOMP3=5, ISMTH=10

COMPl Description

The COMPl subroutine contains the horizontal and vertical advection differences. The DO loops over I are arranged to make use of the periodicity of the variables in the zonal (I) direction. For example, suppose we are to compute $D(I)=Q(I+1)-Q(I)$ for $I=1,\dots,IM$. Then the corresponding code is

```

I=IM
DO 10 IP1=1,IM
D(I)=Q(IP1)-Q(I)
I=IP1
10 CONTINUE
    
```

Where we used the periodicity $Q(IM+1)=Q(1)$.

In the meridional (J) direction we compute our difference approximations in stages in order to make maximum use of each line of latitude of a typical variable when it is in core. The fourth order difference approximation to $\frac{\partial Q}{\partial \phi}(i\ell j\Delta\phi)=D(i,\ell,j)$

$$(1) \quad \frac{4}{3} \left(\frac{Q_{i\ell j+1} - Q_{i\ell j-1}}{2\Delta\phi} \right) - \frac{1}{3} \left(\frac{Q_{i\ell j+2} - Q_{i\ell j-2}}{4\Delta\phi} \right)$$

Thus we see that $Q_{i\ell j}$ will be needed in the difference approximations to $\frac{\partial Q}{\partial \phi}$ for $j-2$, $j-1$, $j+1$, and $j+2$. The corresponding code is

```

DO 20 I = 1,IM
QFLUX1= 4*(Q(I,L,JP1)+Q(I,L,J))
QFLUX2=-.5*(Q(I,L,JP2)+Q(I,L,J))
D(I,L,JP2)= D(I,L,JP2)-QFLUX2
D(I,L,JP1)= D(I,L,JP1)-QFLUX1
D(I,L,J) = D(I,L,J)+QFLUX1+QFLUX2
20 CONTINUE
    
```

For simplicity the array D is initialized to zero and Q is scaled so that (1) contains no divisions.

At the poles ($j=1$ or $J=JNP=46$) we use the values given in section IV on boundary conditions. We have special code for these cases denoted $J=2$ or $J=JM$ corrections.

COMP2 Description

This routine contains the Coriolis force term, the geopotential calculation (which should be made into a separate subroutine), and the pressure gradient and energy term calculations.

The geopotential PHI is dimensioned PHI(72,9,5) since we only need at most five storage locations for any computation. We use modular arithmetic (MOD5) to compute the indices JMOD, JP1MOD, JP2MOD which correspond to the standard index values of J, J+1, and J+2. Thus PHI(I,L,6) is stored in PHI (I,L,1), and we avoid shifting array values by using the JMOD index as a pointer.

For the south pole calculation we need geopotential values at J=2 and 3, thus, for J=1 we compute PHI for J=1,2, and 3. For successive values of J we need only compute PHI at JP2 which is needed in the pressure gradient calculation at J. Therefore, the calculation of PHI and the associated array PK are coded for calculation at JP2. Except for the first J value we are only computing the geopotential at one latitude value for each pressure gradient calculation.

X: Flow Charts

COMPO (Q,QT)

```

JS1=1
JS2=1

((MAIN LOOP))
DO 10 J=1,JM

((COMPUTE ALL J-PARAMETERS))

IF(JP2>JM)                                GO TO 25
IF(J>1)                                    GO TO 18

JP2=JP2MOD=2
((SAVE QTPOL(M) IN QSMPOL(M),M=1,2))

18 ((SAVE QJ(JP2) IN QSM(JP2MOD)))

IF(JP2>2)                                GO TO 25
((INCREMENT JP2,JP2MOD))                  GO TO 18

25 CALL COMP1(Q,QT,J)
CALL COMP2(Q,QT,J)

((ELIMINATE NEGATIVE HUMIDITIES))

IF(PT<(400,1100.))                        STOP
IF(J<3)                                    GO TO 200
IF(J=3)                                    GO TO 70

29 ((UNSCALE QT(JS2)))
CALL AVRX(QT,JS2)

((LEAP FROG, MATSUNO AS RELATED TO THE
SEQUENCE OF 2 CALLS TO COMPO:
1st CALL ⇒ LF ⇒ LF or MP ⇒ MC,
2nd CALL ⇒ LF ⇒ LF or LF ⇒ MP or MC ⇒ MP))

IF(LF→LF)                                GO TO 45
IF(LF→MP.OR.MC→MP)                       GO TO 58

((STATUS: MP→MC OR MC→LF))
((P(I,JS2)+PSM(I,JS2MOD),Q+QSM))

IF(MC→LF)                                GO TO 63

(STATUS: MP→MC))

```

(contin.)

COMPJ (Q,QT)

```

58      ((P(I,JS2)+PT(I,JS2),Q-QT*DXYP*PT)) GO TO 64
45      ((P(I,JS2)+ *P(I,JS2)+α*(PSM+PT),
        Q-β*DXYP*P*Q+α*(QSM+QT*DXYP*PT))
63      CONTINUE
        ((SOURCE TERM CORRECTION FOR LEAP FROG))
      IF (NOT COMP3 CALL) GO TO 67
        ((Q-Q-QT*DXYP*PT))
64      IF (NOT COMP3 CALL) GO TO 67
      IF (MATSUNO PREDICTOR STEP) GO TO 67
        CALL COMP3 (QT,JS2)
        ((COMPLETE SOURCE TERM CORRECTION,
        Q-Q+QT*DXYP*PT))
67      IF (J<JM) GO TO 200
        ((INCREMENT JS2,JS2MOD))
      IF (JS2<JM) GO TO 29
70      ((POLES))
200     JS2=JS1
        JS1=J
10      CONTINUE
        ((END OF J LOOP))
      RETURN

```

COMP1 (Q,QT,J)

```

        ((COMPUTE JP2-JS2MOD))
      IF (J>2) GO TO 2150
        JS1=JS2-JS1MOD=JS2MOD=1

```

(contin.)

COMPL (Q,QT,J)

```

2150 IF (J=JM)                                GO TO 2158
      ((COMPUTE PV1, PV2 = Vj* + Vj+2*))
      ----- ● -----
IF (J=1)                                      GO TO 2225

2158      ((COMPUTE PU1, PU2 = U1* + U1+2*))
      ((COMPUTE HORIZ.ADVEC.-IN LONG. DIREC.))
IF (J = JM)                                  GO TO 2237

2225      ((COMPUTE HORIZ.ADV.-IN LAT. DIREC.))
IF (J>1)                                      GO TO 2290
      ((J=2 CORREC. TO HORIZ. ADVEC & CONV.))
      ((CONV. CALC. FOR CONT. EQ.))

2290 IF (J<JM)                                GO TO 2405

2237      ((J=JM CORREC. TO HORIZ. ADVEC. & CONV.))
      ----- ● -----

2405      ((CONV + CONV * DSIG(L)))
IF (L<NLAY)                                  GO TO 2150
IF (J=1)                                      GO TO 2600
      ((COMPUTE SIGDOT,PT))
      ((COMPUTE VERTICAL ADVEC.))

IF (J<JM)                                    RETURN

2600      ((POLES, M=1 or 2))
      ((HORIZ. ADVEC.))
      ((CONT. E.Q.))
      ((SIGDOT PL, PTPOL))
      ((VERT. ADVEC.))
      RETURN

```

END OF COMPL

Alternative code from -----●----- to -----●-----

```

IF (J.EQ.1)                                  GO TO 2222
      ((COMPUTE PU1, PU2))
      ((COMPUTE HA(I) ))
      GO TO 2235

2222      ((J=2 CORREC ))

2235 IF (J<JM)                                GO TO 2225
      ((J=JM CORREC))
      GO TO 2405

2225      ((HORIZ ADVEC(J) ))
      ((CONV.CALC.FORCONT.EQ.))

2405      ((CONV CONV*DSIG(L)))
      -----●-----

```

COMP2 (Q,QT,J)

```

((COMPUTE JP2=... ))
IF(J>1) ((JP2=JP2MOD=JPKP2=1)) GO TO 3001
:3005
3001 ((CORIOLIS ))
IF(J>JM) GO TO 3032
((FIRST MAIN LOOP IN L))
3005 ((DO 3030 LX=1,NLAY ))
((COMPUTE ENERGY TERM IN
THERMODYNAMIC EQN ))
((L=NLAY+1-LX ))
IF(L<NLAY) GO TO 3055
3007 ((COMPUTE PK(I,LL,JPKP2)FORLL=1,NLAY))
((COMPUTE PHI'(I,NLAY,JP2MOD)) :3065
3055 ((COMPUTE PHI' AT JP2MOD FOR L<NLAY))
3065 IF(JP2=1 or JP2=JNP) GO TO 3030
((COMPUTE W(JP2MOD)))
3030 ((INCREMENT LX ))
((IF J=1 GO TO 3005 AND COMPUTE PHI(2),W(2)))
(( 2nd MAIN LOOP IN L))
3032 ((DO 3031 LX=1, NLAY))
IF(J=1) GO TO 3111 ( POLES)
IF (J > 2) GO TO 3085
((J=2: VT,TT CORREC.))
3085 (((P) FOR U EQ.))
IF (J=2orJM) GO TO 3135
(((P) FOR V EQ.))
3135 IF (J < JM) GO TO 3031
(( CORREC. TO (P) J=JM))
3111 ((POLES))
3031 ((CONTINUE))
((RETURN))
END

```

A FOURTH-ORDER FORECASTING MODEL*

(E. Kalnay-Rivas, D. Hoitsma, and P. Anolick)

The GISS fourth-order model (Kalnay-Rivas, et al., 1977), which is a fourth-order, energy-conserving GCM on an unstaggered grid, had shown promising capabilities. It produced forecasts that showed an improvement over the second-order GISS forecasts with the same fine grid ($4^{\circ} \times 5^{\circ}$) resolution, but that were somewhat inferior to the "ultrafine" forecasts. However, the first version of the model required excessive amounts of computer memory and time for execution.

The model has been reprogrammed into the "fourth-order band model." The new program solves the primitive equations one latitude band at a time. The arrays are stored in an interlaced way, with all arrays being updated at the same latitude stored contiguously, and similarly for all arrays used in the computation of the time derivatives. This design of the program makes effective use of the virtual memory capability of the Modeling and Simulation Facility's IBM 370/165 or Amdahl computers. The virtual memory facility permits the execution of programs whose core size is larger than the one available, by placing the excess on disk and reading in those pages of information needed in the current calculations. The band fourth-order data structure and computations were constructed to optimize this virtual I/O process; a possible improvement may be to interlace also the arrays being updated with those used to compute the time derivatives. The use of the virtual memory facility avoids the explicit I/O used in the current Kern model, and yields a simpler program. (The band structure was suggested by G. Russell.)

The band fourth-order model computations have been optimized so that each time step is computed in the Amdahl in half the time required by the old model. The array structure has also been designed to reduce the amount of overall storage and high-speed memory by a factor of two (see Table 1).

Table 1. Comparison of Fourth-Order Model Computing Requirements.

	Original 4th-Order Model	Band 4th-Order Model
Core (bytes)	3500K	1500K
CPU time per step (COMP1, COMP2)	34 sec	17 sec

*Reprint from Atmospheric and Oceanographic Research Review-1978, G-AS, NASA Technical Memorandum 80253.

The model has been programmed so that it can use both the Matsuno and the leapfrog time schemes, the latter with the Robert time smoother. If Q^n and $D(Q^n)$ denote the fields at time $n\Delta t$ and the corresponding time derivative computed from the space differences, the smoothed leapfrog scheme is

$$Q^{n+1} = \bar{Q}^{n-1} + 2\Delta t D(Q^n)$$

$$\bar{Q}^n = Q^n + .5v(Q^{n-1} + Q^{n+1} - 2Q^n)$$

with $\bar{Q}^0 = Q^0$. The use of a smoothing coefficient requires a slightly smaller time step. For example, with $v = .1$, the model is marginally stable with $\Delta t = 288$ sec., compared to $\Delta t = 300$ sec. for the Matsuno and leapfrog schemes.

The smoothed leapfrog scheme is further modified to include the subgrid "physics" terms, and scaling and spatial smoothing procedures. Since the "physics" is called every few time steps, unless the leapfrog scheme is restarted after every call to the "physics," only one of the two consecutive fields will be affected. The restarting procedure is time-consuming, so it has been replaced by the following algorithm

$$Q^{n-1} = Q^{n-1} - Q^n$$

$$\tilde{Q}^n = Q^n + S^n$$

$$Q^{n-1} = Q^{n-1} + \tilde{Q}^n$$

where S^n corresponds to the "physics" terms. This procedure, which ensures that the physics is applied to two consecutive time steps, has been tested with good results.

The model can be extended into an "ultrafine" version in a straightforward way.

Preliminary Results. The first numerical integrations performed with the new band fourth-order model show dramatic improvements over forecasts made with the GLAS model with the same resolution ("fine grid" or 4° latitude by 5° longitude). The quality of the forecasts is now comparable with those produced with the "ultrafine" (2.5° latitude by 3° longitude) version of the model. Results from a 3-day numerical integration are presented in Figure 1. During an extended 8-day integration of the new fourth-order model, the atmospheric systems remained remarkably smooth, exhibiting a realistic behavior both with respect to position and intensities.

References

Kalnay-Rivas, E., A. Bayliss, and J. Storch, 1977: The fourth-order GISS model of the global atmosphere. Contributions to Atmos. Phys., 20, pp. 299-311.

The 4th Order GISS Model of the Global Atmosphere*

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Abstract: The new GISS 4th order model of the global atmosphere is described. It is based on 4th order quadratically conserving differences with the periodic application of a 16th order filter on the sea level pressure and potential temperature equations, a combination which is approximately enstrophy conserving. Several short range forecasts indicate a significant improvement over 2nd order forecasts with the same resolution (~ 400 km). However the 4th order forecasts are somewhat inferior to 2nd order forecasts with double resolution. This is probably due to the presence of short waves in the range between 1000 km and 2000 km, which are computed more accurately by the 2nd order high resolution model. An operation count of the schemes indicates that with similar code optimization, the 4th order model will require approximately the same amount of computer time as the 2nd order model with the same resolution.

It is estimated that the 4th order model with a grid size of 200 km provides enough accuracy to make horizontal truncation errors negligible over a period of a week for all synoptic scales (waves longer than 1000 km).

1. Introduction

It is generally accepted that the use of 4th order finite differences is more efficient in reducing space truncation errors than the use of higher resolution on a 2nd order model (KREISS and OLIGER, 1972). Linear analyses and shallow water type of experiments give *an upper limit* of the improvement that can be expected from the reduction of errors in the horizontal differences. For example, Table 1 corresponds to a linear wave equation with phase speed $c = 11 \text{ ms}^{-1}$, typical of atmospheric motions. It provides a measure of the "computational predictability period" after which the errors introduced by horizontal truncation alone become very serious. The table suggests that for waves longer than 2000 km a 4th order -- 400 km grid model is preferable to a 2nd order -- 200 km grid model. Waves shorter than 2000 km are forecast more accurately by a 2nd order -- 200 km grid model. In order to insure that horizontal truncation errors are small in the 1--3 week period of atmospheric predictability for all synoptic scale waves it is necessary to use either 4th order differences with a grid resolution of the order of 200 km or 2nd order differences with a grid of the order of 100 km. Numerical experiments with simple nonlinear models (WILLIAMSON and BROWNING, 1973, KALNAY-RIVAS 1976a, from now on I) also indicate dramatic reduction of errors by the use of 4th order differences.

All these studies indicate that a considerable improvement in forecasting skill is to be expected from the use of 4th order differences *if a substantial portion of the forecasting errors is due to horizontal truncation errors*. In actual numerical forecasts there are several other important sources of errors: vertical truncation errors, errors in the initial conditions, and poor "physics" (i.e. parameterization of physical processes like radiation, dissipation, subgrid transports, cumulus convection, boundary layers, etc.)

* (Paper presented during the DMG-AMS Meeting, Hamburg 1976; see Preface to Issue 1 - 2/1977)

■ **Table 1.** Flapsed time T after which a wave of wavelength L moving with a phase speed $c = 11 \text{ ms}^{-1}$ lags by more than 100 km due to space truncation errors.

	$\Delta x = 400 \text{ km}$	$\Delta x = 200 \text{ km}$
2nd order differences	L = 2000 km, T = 0.4 days L = 4000 km, T = 1.6 days	L = 2000 km, T = 1.6 days L = 4000 km, T = 6.4 days
4th order differences	L = 2000 km, T = 1.5 days L = 4000 km, T = 21 days	L = 2000 km, T = 21 days L = 4000 km, T = 328 days

At the Goddard Institute for Space Studies (GISS), New York, we have developed a 4th order general circulation model (GCM) with the expectation that it will yield more accurate short range forecasts, and probably more realistic climate simulations, but with the improvement limited by the other sources of errors. The model is described in Section 2. The results of several experimental short range forecasts are discussed in Section 3. A sample 36-hour forecast is presented and compared with forecasts by the standard 2nd order GISS model, which has the same resolution as the 4th order model, and by the "ultrafine" 2nd order GISS model, which has twice the horizontal resolution. Section 4 contains the conclusions and a discussion of future work.

2. Description of the model

The 4th order GISS global atmospheric model is a primitive equation model using longitude (λ), latitude (ϕ) and sigma (σ) coordinates. The basic equations and the parameterizations of physical processes are the same as those of the standard 2nd order GISS global model (Somerville et al, 1974). The finite difference scheme is quite different and is described in the following subsections.

2.1. Finite-difference scheme

The horizontal grid is uniform (constant $\Delta\lambda$ and $\Delta\phi$) and non-staggered. There are two ways to define such a grid, depending on whether variables are defined at the poles, as done by WILLIAMSON and BROWNING (1973), or half a grid size away from the poles (HOLLOWAY et al, 1973). We chose the former method because in the absence of smoothing near the poles it allows a time step twice as large as the latter. The singularity that spherical coordinates have at the poles is explicitly avoided by the use of a polar cell where "stereographic" velocities are used.

In the vertical direction we use a staggered grid with σ , the vertical velocity, defined at the boundaries of each layer, and all other variables in the center of the layer. The vertical grid can be non uniform although experiments have been made so far with constant $\Delta\sigma$. Experiments reported in I indicated that no significant improvement in accuracy was obtained when a 4th order staggered conservative scheme was used in the vertical direction. Therefore the model has 2nd order vertical differences. The computation of the geopotential Φ is performed as indicated by ARAKAWA (1972), with a modification suggested by PHILLIPS (1974).

a) Forecast equation away from the poles

Several systems of horizontal difference were tested as reported in I. The scheme chosen for the basis of these experiments consists of the simplest possible quadratically conservative 4th order differences.

We define the finite difference horizontal divergence operator at the vertical level k as

$$D_k(\mathbf{g}) = \frac{1}{a \cos \phi} \left[\frac{4}{3} \delta_\lambda (\Pi u^\lambda \mathbf{g}^\lambda) - \frac{1}{3} \delta_{2\lambda} (\Pi u^{2\lambda} \mathbf{g}^{2\lambda}) + \frac{4}{3} \delta_\phi (\Pi v \cos \phi^\phi \mathbf{g}^\phi) - \frac{1}{3} \delta_{2\phi} (\Pi v \cos \phi^{2\phi} \mathbf{g}^{2\phi}) \right]_k \quad (2.1)$$

and the finite difference gradient operator as

$$\nabla \mathbf{g} = \frac{\mathbf{i}}{a \cos \phi} \left[\frac{4}{3} \delta_\lambda \mathbf{g}^\lambda - \frac{1}{3} \delta_{2\lambda} \mathbf{g}^{2\lambda} \right] + \frac{\mathbf{j}}{a} \left[\frac{4}{3} \delta_\phi \mathbf{g}^\phi - \frac{1}{3} \delta_{2\phi} \mathbf{g}^{2\phi} \right] \quad (2.2)$$

Here $\Pi = p_S - p_T$ is the difference between the pressure at the surface and the constant pressure at the top of the model, a is the radius of the earth, $\mathbf{w} = u\mathbf{i} + v\mathbf{j}$ is the horizontal velocity vector in spherical coordinates, and we use the finite difference notation

$$\delta_{n\lambda} \mathbf{g} = [\mathbf{g}(\lambda + n\Delta\lambda/2, \phi, \sigma, t) - \mathbf{g}(\lambda - n\Delta\lambda/2, \phi, \sigma, t)] / (n\Delta\lambda), \quad (2.3)$$

$$\mathbf{g}^{n\lambda} = [\mathbf{g}(\lambda + n\Delta\lambda/2, \phi, \sigma, t) + \mathbf{g}(\lambda - n\Delta\lambda/2, \phi, \sigma, t)] / 2,$$

and similar formulas for the other independent variables. With this notation, the continuity equation away from the poles is

$$\frac{\partial \Pi}{\partial t} = -D_k(\mathbf{w}) - \Pi \delta_\sigma \dot{\sigma}_k \quad (2.4a)$$

or, if we integrate it in the vertical and make use of the boundary condition

$$\dot{\sigma} = 0 \text{ at } \sigma = 0, 1,$$

$$\frac{\partial \Pi}{\partial t} = - \sum_{k=1}^K D_k(\mathbf{w}) \Delta \sigma_k \quad (2.4b)$$

where K is the number of vertical layers. Equation (2.4b) is the forecast equation for Π . The momentum equation is

$$\left(\frac{\partial \mathbf{w}}{\partial t} \right)_k = -D_k(\mathbf{w}) - \Pi \delta_\sigma (\dot{\sigma} \mathbf{w}^\sigma)_k + \left(f + \frac{u_k \tan \phi}{a} \right) \Pi \mathbf{k} \times (\Pi \mathbf{w})_k - \Pi \left[\nabla \Phi' + \frac{\sigma RT'}{p} \nabla \Pi \right]_k + \mathbf{F}_k. \quad (2.5)$$

The first law of thermodynamics is

$$\left(\frac{\partial \Pi T}{\partial t} \right)_k = -D_k(T) - \{ p^k \Pi \delta_\sigma [\dot{\sigma} (T/p^k)^\sigma] \}_k + \left\{ \frac{\Pi \sigma k T}{p} \left[\frac{\partial \Pi}{\partial t} + \mathbf{w} \cdot \nabla \Pi \right] \right\}_k + \frac{\Pi}{C_p} Q. \quad (2.6)$$

Here T is absolute temperature, \mathbf{F} is the frictional force, $\kappa = R/C_p$, Q is the diabatic heating per unit mass, p is the pressure and p^h is computed as indicated in Subsection 2.1c. The moisture equation for the water vapor mixing ratio q is

$$\left(\frac{\partial \Pi q}{\partial t}\right)_k = -D_k(q) - \Pi \delta_{jk} (\dot{q}^o)_k + \Pi (E - C). \quad (2.7)$$

E and C are the rates of evaporation and condensation.

In the momentum equation (2.5) we follow a device suggested by PHILLIPS (1974) to alleviate the difficulties of the computation of the pressure gradient in σ -coordinates in the vicinity of orography. We define $T' = T - T(p)$, $\Phi' = \Phi - \Phi(p)$, where $T = \theta p^h$, $\theta = 280 \text{ K}/(1000 \text{ mb})^h$, $\Phi = \Phi_0 \cdot C_p \theta p^h$ and $\Phi_0 = [(1000 \text{ mb})^h \cdot p^h]^{-1}$. We have simplified PHILLIPS' expression for Φ_0 since its precise value is not important. With this procedure, in the pressure gradient term there is an exact cancellation of the terms $\nabla \Phi + \frac{\sigma R T}{p} \nabla \Pi$, and this implies a significant reduction of truncation errors in regions with steep orography.

b) Forecast equations at the Poles

We have followed the method used by WILLIAMSON and BROWNING (1973) and define a polar cap of radius $a \Delta \phi$ on which we use "stereographic" (or rather "cartesian") velocity components defined by the transformation

$$\begin{aligned} U_{1p} &= -u \sin \lambda + v \cos \lambda \\ V_{1p} &= +u \cos \lambda + v \sin \lambda \end{aligned} \quad (2.8)$$

with inverse

$$\begin{aligned} u &= -U_{1p} \sin \lambda + V_{1p} \cos \lambda \\ v &= +U_{1p} \cos \lambda + V_{1p} \sin \lambda \end{aligned} \quad (2.9)$$

where the top and bottom signs correspond to the north and south poles respectively. The positive x -axis of the cartesian coordinates coincides with the meridian of longitude $\lambda = 0$. The difference in signs at the south pole is due to the choice of a right handed system of coordinates with the vertical unit vector pointing outwards. Formulas (2.8) and (2.9) are used to define "spherical" or "cartesian" velocity components wherever they are required in the finite difference equations.

The finite difference horizontal divergence operator at the poles is defined by an average over all longitudes λ_i :

$$\begin{aligned} D_{1p, k}(g) &= + \sum_{i=3}^{l+2} \left\{ \frac{4}{3} A_1 (\text{Hvg})_{\lambda_i} + \left(\frac{\pi}{2} - \Delta \phi\right) \right. \\ &\quad \left. - \frac{1}{3} A_2 (\text{Hvg})_{\lambda_i} + \left(\frac{\pi}{2} - 2 \Delta \phi\right) \right\}_k \end{aligned} \quad (2.10)$$

where $\lambda_i = (i-3) \Delta \lambda$, $\Delta \lambda = 2\pi/l$, and $A_n = \Delta \lambda \sin(n \Delta \phi) / [2\pi a^2 (1 - \cos(n \Delta \phi))]$. The finite difference gradient operator at the poles is also defined by an average over all longitudes:

$$\nabla_{1p} g = \frac{2}{l a \Delta \phi} \sum_{i=3}^{l+2} \left\{ \frac{4}{3} g_{\lambda_i} + \left(\frac{\pi}{2} - \Delta \phi\right) - \frac{1}{3} g_{\lambda_i} + \left(\frac{\pi}{2} - 2 \Delta \phi\right) \right\} \cdot \{\cos \lambda_i \mathbf{i} + \sin \lambda_i \mathbf{j}\} \quad (2.11)$$

Then the continuity equation at the poles is

$$\frac{\partial \Pi_{\pm p}}{\partial t} = \mp \sum_{k=1}^K D_{\pm p, k}(1) \Delta \sigma_k, \quad (2.12)$$

the momentum equation is

$$\begin{aligned} \left(\frac{\partial \Pi V}{\partial t} \right)_k &= \mp D_{\pm p, k}(V) - \Pi_{\pm p} \delta_\sigma \left(\dot{\sigma}_{\pm p} \bar{V}_{\pm p}^\sigma \right)_k + f_{\pm p} \mathbf{k} \times (\Pi V)_{\pm p, k} \\ &- \Pi_{\pm p} \left[\nabla_{\pm p} \Phi' + \frac{\sigma RT'}{p} \nabla_{\pm p} \Pi \right]_k + F_{\pm p, k} \end{aligned} \quad (2.13)$$

where $V_{\pm p} = U_{\pm p} \mathbf{i} + V_{\pm p} \mathbf{j}$. The thermodynamic equation is

$$\begin{aligned} \left(\frac{\partial \Pi T}{\partial t} \right)_{\pm p, k} &= \mp D_{\pm p, k}(T) - \{ p_{\pm p}^k \Pi_{\pm p} \delta_\sigma \left[\dot{\sigma}_{\pm p} \overline{\left(T_{\pm p} / p_{\pm p}^k \right)} \right] \}_k \\ &+ \left\{ \frac{\Pi_{\pm p} \sigma \kappa T_{\pm p}}{p_{\pm p}} \left[\frac{\partial \Pi_{\pm p}}{\partial t} + V_{\pm p} \cdot \nabla_{\pm p} \Pi \right] \right\}_k + \frac{\Pi_{\pm p}}{C_p} Q_{\pm p} \end{aligned} \quad (2.14)$$

and the moisture equation is similar.

c) Diagnostic equations

In σ -coordinates, the pressure is defined by $p = \sigma \Pi + p_T$. Following a suggestion by PHILLIPS (1974) we define at the center of a layer $p_k^k = 1/(\kappa + 1) \delta_\sigma p_k^{k+1} / \delta_\sigma p_k$, instead of $p_k^k = (\bar{p}_k^\sigma)^k$ as in ARAKAWA (1972). This formula was derived under the assumption that the potential temperature varies in a σ -layer much less than either the temperature or the pressure. PHILLIPS indicated that a more accurate relationship between temperature and geopotential can be expected from this formula.

The geopotential Φ is determined following ARAKAWA (1972): If Φ_S is the surface geopotential, then the geopotential at the center of the lowest layer is

$$\Phi_K = \Phi_S + \sum_{k=1}^K C_k T_k \quad (2.15)$$

and at other levels

$$\Phi_k = \Phi_{k+1} + C_p \Delta \sigma_k + \frac{1}{2} \delta_\sigma \left(p_k^k + \frac{1}{2} \right) \left(\frac{T_k + 1/2 \sigma}{p_k^k + 1/2} \right) \quad (2.16)$$

The coefficients C_k are determined from

$$C_k = \frac{\Pi \sigma_k R \Delta \sigma_k}{p_k} - \frac{C_p}{p_k^k} \left(\sigma \Delta \sigma \delta_\sigma p^k \right)_k \sigma \quad (2.17)$$

The vertical velocity in σ -coordinate is determined from (2.4a) and (2.4b), and the boundary condition $\dot{\sigma} = 0$ at $\sigma = 0, 1$:

$$\delta_\sigma \dot{\sigma}_k = \frac{1}{\Pi} \left\{ \sum_{k=1}^K D_k(1) \Delta \sigma_k - D_K(1) \right\} \quad (2.18)$$

d) **Boundary conditions**

In the east-west direction we use periodicity: $g(\lambda + 2\pi, \phi) = g(\lambda, \phi)$ for all variables. When the value of a variable is needed "beyond" the poles, we define it by continuation along the same meridian:

$$w\left(\lambda, \pm\left(\frac{\pi}{2} + \Delta\phi\right)\right) = -w\left(\pi + \lambda \pm\left(\frac{\pi}{2} - \Delta\phi\right)\right); g\left(\lambda, \pm\left(\frac{\pi}{2} + \Delta\phi\right)\right) = g\left(\pi + \lambda, \pm\left(\frac{\pi}{2} - \Delta\phi\right)\right),$$

where g represents all variables other than the two horizontal velocity components. In the vertical the top and bottom are material surfaces through which there is no flux ($\dot{\sigma} = 0$ at $\sigma = 0, 1$) except for subgrid boundary layer fluxes of momentum and heat included in F and Q .

2.2. Filtering near the poles and high order filtering

The CFL computational stability condition requires a very small time step unless linearly unstable short waves are filtered out near the poles. For this purpose several alternative procedures have been tried but so far the method found to give most satisfactory results is the fourier filtering of the prognostic variables u, v, T and *the indirect smoothing of Π through the filtering of the sea level pressure (SLP) field.*

The fourier filtering is performed polewards of 66° latitude. The amplitudes of the fourier components of zonal wavenumber n are multiplied by a transfer function which is 1 for $n \leq N$ and decreases linearly to zero between $n = N$ and $n = N + 5$. The number of retained modes is defined by $N(\phi) = \text{integer part of } (90 \cos \phi)$. We have tried filtering the stereographic velocities U, V instead of u, v but no improvement was obtained.

Since both the surface geopotential and the surface pressure fields contain large amplitude short wave components due to the presence of orography, an artificial smoothing of these fields represents a distortion of the real geometry of the boundary. On the other hand, the SLP is an intrinsically smooth field, and its high wavenumber components more closely represent *atmospheric waves*. Therefore in our model *the SLP is filtered near the poles and Π is recovered by solving the transcendental equation used to relate them.*

It was found that the 4th order model forecasts were less smooth than those of the 2nd order GISS model, especially in regions with steep orography. Based on the same considerations we have introduced a periodic application of a high (16th) order filter (SHAPIRO, 1970) on the SLP and potential temperature fields, which are not very affected by orography.

The filter is of the form $\bar{g} = \{1 - (F_\phi^2)^8\} \{1 - (F_\lambda^2)^8\} g$, where $F_\lambda^2(g_{ij}) = (g_{i+ij} - 2g_{ij} + g_{i-ij})/4$ and has a response $F_\lambda^2[\exp(ik\lambda)] = -\sin^2(k\Delta\lambda/2) \exp(ik\lambda)$. This filter eliminates waves shorter than $4\Delta\lambda$ and even after hundreds of applications has negligible damping effect on waves longer than $4\Delta\lambda$.

In the meridional direction there are three simple alternative forms for F_ϕ^2 :

$$F_{\phi I}^2(g_{ij}) = (g_{ij+1} - 2g_{ij} + g_{ij-1})/4$$

$$F_{\phi II}^2(g_{ij}) = [(g_{ij+1} - g_{ij}) \cos \phi_{j+1/2} - (g_{ij} - g_{ij-1}) \cos \phi_{j-1/2}] / (4 \cos \phi_j)$$

and

$$F_{\phi III}^2(g_{ij}) = [F_{\phi I}^2(g_{ij} \cos \phi_j)] / \cos \phi_j$$

FRANCIS (1975) used $F_{\phi II}^2$, which is the only one that conserves the area weighted average of g . However, we have found that the three meridional filters produced virtually identical results. This could

be expected because the waves affected by the filter are too short to be strongly influenced by the convergence of the meridians. Since $F_{\phi I}^2$ and $F_{\phi II}^2$ can be programmed as efficiently as F_{λ}^2 we find them preferable. In the model we use $F_{\phi I}^2$, with g continued "beyond" the poles as indicated in Subsection 2.1d. At the poles $\bar{g}_{\pm p}$ is defined as the average over all longitudes λ_i of the filtered values $\bar{g}_{i, \pm p}$. At the present we are applying the high order filter to the SLP and potential temperature fields once every two hours.

2.3. Enstrophy constraint

BAYLISS and ISAACSON (1975) have developed a simple method to make any finite difference scheme conservative with respect to any quantity. The method has been tested in our model by forcing conservation of enstrophy on the dry adiabatic version of the model, although in such a model it is potential enstrophy that is conserved. The procedure is the following: Let the vectors U and V represent the values of the velocity components discretized over the grid and let the functional $G(U, V)$ denote the consistent 4th order approximation of the mean square vorticity. At each time step a correction U', V' is added to the predicted values \bar{U}, \bar{V} such that $G(\bar{U} + U', \bar{V} + V') = G(U_0, V_0)$ where U_0, V_0 denote the velocities at time $t = 0$. Since this equation cannot be solved explicitly, it is linearized about the predicted values \bar{U}, \bar{V} :

$$\left(\frac{\partial G}{\partial U}\right)_{\bar{U}, \bar{V}} \cdot U' + \left(\frac{\partial G}{\partial V}\right)_{\bar{U}, \bar{V}} \cdot V' = G(U_0, V_0) - G(\bar{U}, \bar{V}) \quad (2.19)$$

If we assume

$$(U', V') = \alpha \left(\frac{\partial G}{\partial U}, \frac{\partial G}{\partial V}\right)_{\bar{U}, \bar{V}} \quad (2.20)$$

then α can be determined from (2.19):

$$\alpha = \frac{G(U_0, V_0) - G(\bar{U}, \bar{V})}{\left|\frac{\partial G}{\partial U}\right|^2 + \left|\frac{\partial G}{\partial V}\right|^2} \quad (2.21)$$

It may be easily shown that the choice (2.20) minimizes $\|U'\| + \|V'\|$, the norm of the correction vector. It has been found that this procedure improves the forecasting skill of the model. However, when it is used in combination with the periodic application of the high order filter the improvement is marginal. Therefore this option is not included in the standard version of our model.

2.4. Summary of the properties of the model

The finite differences of the model have the following properties:

- a) Horizontal differences are performed on a nonstaggered grid and have 4th order truncation errors.
- b) Vertical differences are performed on a staggered grid and have 2nd order truncation errors.
- c) The differences are conservative in the sense that in the absence of diabatic and dissipation, mass and energy are conserved except for marginal terms at the poles. Non-conservative differences similar to those used by WILLIAMSON and BROWNING (1973), *starting from real data*, proved to be nonlinearly unstable after less than one day of integration.
- d) Unlike the "box method", the horizontal differences remain 4th order near the poles (KALNAY-RIVAS, 1976b)

- e) The model contains *no horizontal diffusion* except for the possible use of a dissipative time scheme and for the periodic application of the high order filter. It has been shown in I that the combination of a 4th order quadratically conservative scheme with the *periodic* application of a high order filter *replaces successfully the use of an enstrophy-conserving scheme*. This is because waves shorter than four times the grid size are the ones subject to aliasing and to large truncation errors, and they are removed by the filter *before they attain finite amplitude*. Waves longer than four times the grid size are accurately computed by the 4th order scheme and are not affected by the filter.

3. Numerical forecasts

3.1. Analytical initial conditions

The model was tested using 3 dimensional extensions of both the nontrivial steady state solution and the Rossby-Haurwitz wave initial conditions used in I. The forecasts remained smooth during the several days of interaction. Errors in the steady state case were an order of magnitude larger than in I because of the use of single precision arithmetic on the IBM/360-95 computer. However, it was found necessary to use double precision for the longitudinally averaged pressure gradient at the poles in order to avoid excessive local error growth.

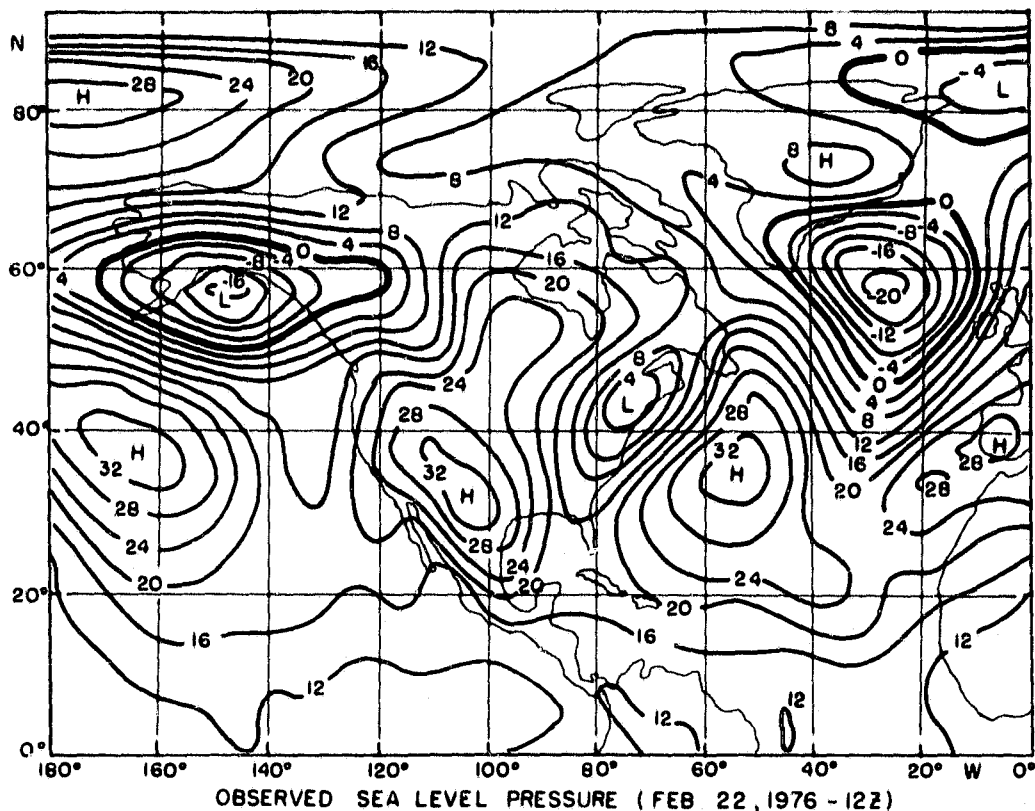
3.1. Forecasts with real initial data

As mentioned before, the parameterization of physical processes included in the terms F , Q , C and E of Equations (2.5) -- (2.7) has been adapted from the standard 2nd order GISS model (SOMERVILLE et al, 1974). The only change that has been introduced is the reduction by a factor of 2 of the surface drag coefficient. This was found to be necessary in order to avoid an excessive damping of the pressure systems after one or two day forecasts. We believe that the fact that in the staggered grid the surface winds are obtained by a horizontal average reduces the relative effect of friction in the standard model. At the present time (November 1976), we have performed 5 experimental short range forecasts. Since they have been made with slightly different versions of the 4th order model, we don't yet have a reliable measure of the model's forecasting skill. The initial data so far has been available in the staggered grid used by the standard model so that winds have been linearly interpolated to the nonstaggered grid. This is a source of error which may have adversely affected the 4th order forecasts.

We have performed a single 4-day forecast which indicated that the model remains very stable and that synoptic systems show no tendency to become unrealistically weak.

All the numerical 4th order forecasts show a significant improvement over the 2nd order forecasts performed with the same resolution. This improvement appears as a better estimation of the changes in position and intensity of several pressure systems. This is an encouraging result, especially in view of the study made by BAUMHEFNER and DOWNEY (1976) which indicates that the standard 2nd order GISS model forecasts compare favorably with those made by the NMC and NCAR models with similar resolution.

On the other hand, and contrary to our expectations, the forecasts made with the 2nd order GISS model with double horizontal resolution were found to be either similar or superior to the 4th order forecasts. This is more true in the sea level pressure than in the 500 mb forecasts. We consider that there are several possible reasons for this result. One is the extra errors introduced in the 4th order model initial

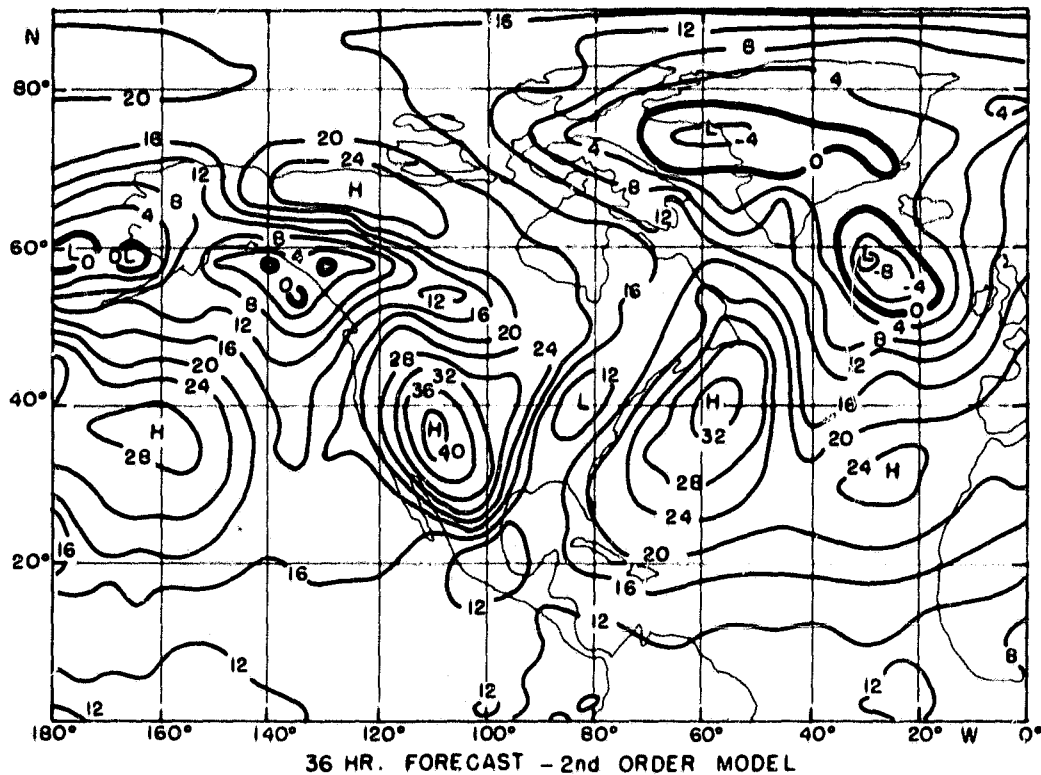


● Figure 1. Observed 36 hr. sea level pressure at 0000 GMT, 22 February 1976

data by the averaging of the winds. A second reason (and probably the most important one) is the existence of waves with significant amplitudes and wavelengths between 1000 and 2000 km. These short waves are better detected in the initial data and more accurately forecast by the 2nd order model with double resolution. The nonlinear interaction of these short waves with longer synoptic waves which according to linear theory are better forecasted by the 4th order model is clearly very important. A third important reason is the fact that errors introduced by the parameterization of subgrid processes become less important as the grid size is reduced.

We present an example of a 36 hour forecast of the sea level pressure corresponding to February 22, 1976, 12Z. Figure 1 is the verification sea level pressure map. Figure 2 is the forecast computed with the standard 2nd order GISS model, which has a resolution $\Delta\lambda = 5^\circ$, $\Delta\phi = 4^\circ$, $\Delta\sigma = 1/9$. Figure 3 is the 4th order forecast computed with the same resolution, and Figure 4 is the forecast obtained using the "ultrafine" 2nd order GISS model which has a resolution $\Delta\lambda = 2.5^\circ$, $\Delta\phi = 2^\circ$, $\Delta\sigma = 1/9$.

The most striking improvement showed by the 4th order forecast is in the position and intensity of the low in the eastern portion of North America. In fact, this low and the high over the Atlantic Ocean have been forecasted better by the 4th order model than by the 2nd order model with double resolution. On the other hand the developing cold high in central Canada and the intense low pressure center south of Greenland have been predicted better by the double resolution 2nd order model. As usual in this

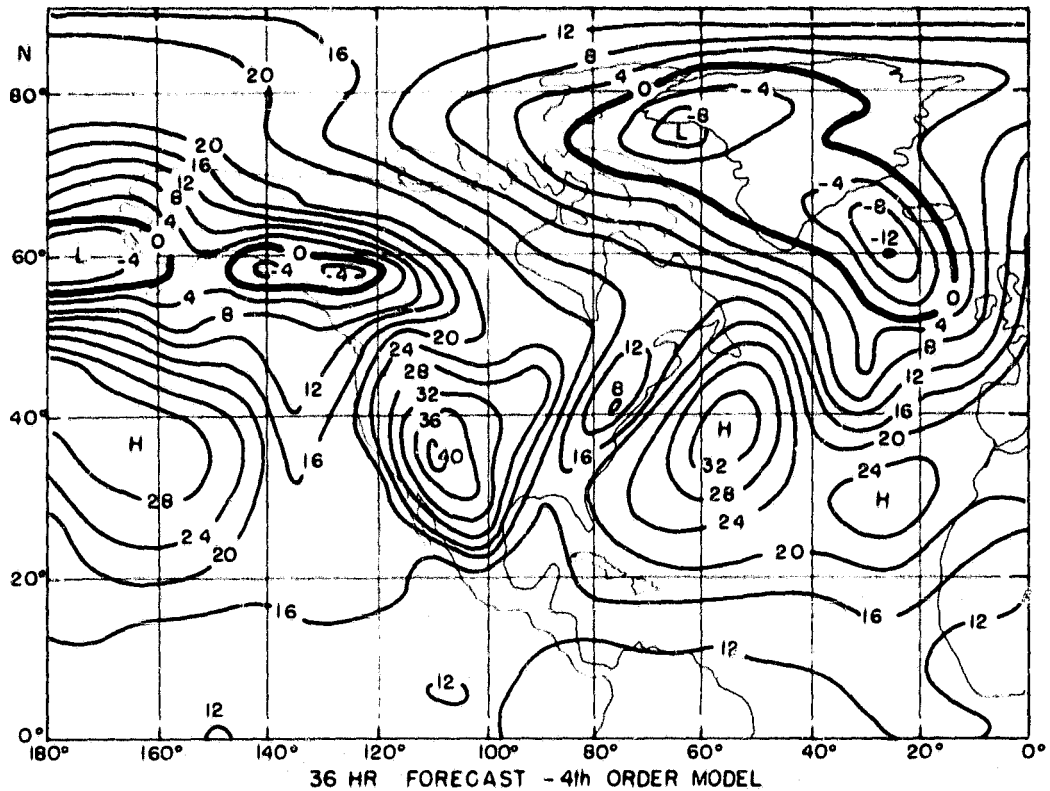


● Figure 2. 36 hr. forecast using the standard 2nd order GISS model

type of comparison, the three forecasts share several important deficiencies. For example the intensity of the high over southwest USA has been overpredicted by the three models, probably because of difficulties introduced by the orography. The low south of Alaska has been erroneously split by the three models. This splitting, which is slightly less pronounced in the 4th order forecast, may be due to both errors in the initial data and orographic and coastal problems.

4. Summary and conclusions

We have described the characteristics of the 4th order GISS model of the global atmosphere. It is based on a quadratically conservative scheme with the periodic application of a 16th order filter on the sea level pressure and potential temperature fields. As shown in I this combination is approximately enstrophy-conserving. An operation count of the numerical schemes indicates that with similar code optimization the 4th order model will require approximately the same amount of computer time as the 2nd order enstrophy-conserving GISS model with the same resolution. We also plan to introduce a simplified semi-implicit scheme, and to study the possibility of using the combination of a nonstaggered vertical grid with the Kreiss 4th order method described in I.

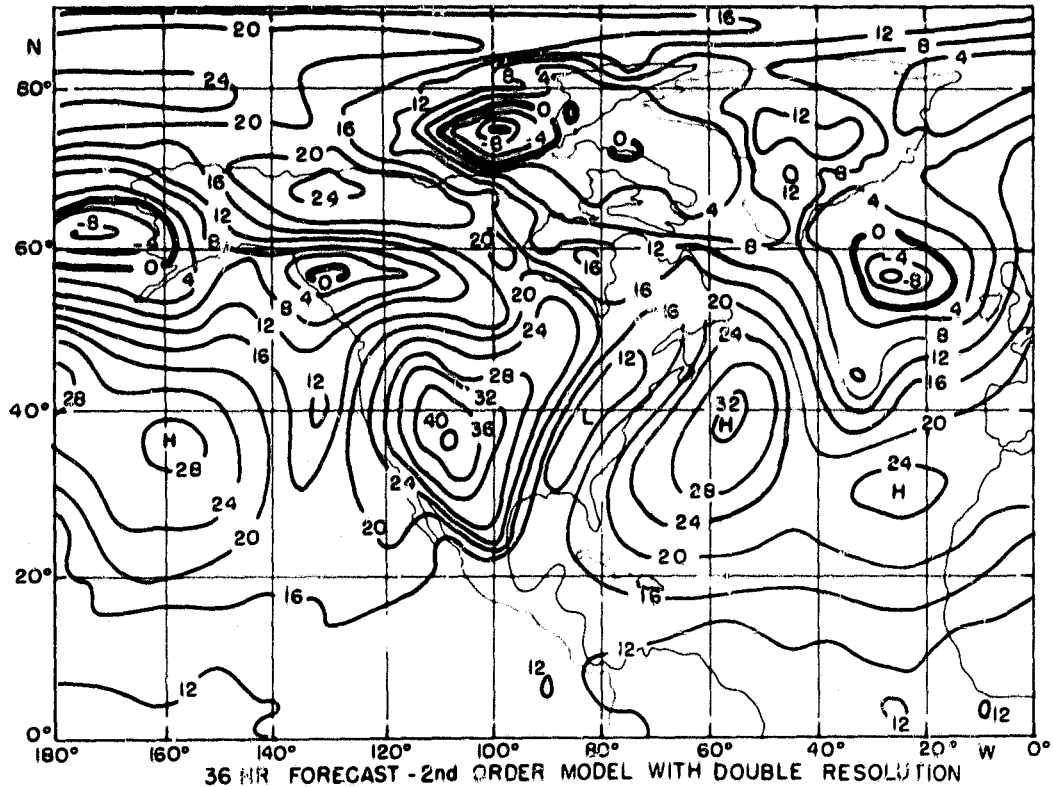


● Figure 3. 36 hr. forecast using the 4th order GISS model with the same resolution as in Figure 2

The results of several short range forecasts indicate a significant improvement over the 2nd order forecast with the same resolution. This improvement is shown in the estimations of changes in position and intensity of several pressure systems. We plan to study the impact that the greater accuracy of 4th order differences has on the forecasting skill of variables of more practical importance, such as temperature and precipitation.

It has been found that the 4th order forecasts are somewhat inferior to the forecasts made with a 2nd order model with double horizontal resolution. We consider that one of the most important reasons for this result are the presence of waves in the range between 1000 km and 2000 km which are computed more accurately by the high resolution 2nd order model than by the 4th order model. Another important reason is that errors introduced by the parameterization of subgrid processes become smaller as the size is reduced.

It is estimated that the 4th order model with a grid size of 200 km provides enough accuracy to make horizontal truncation errors negligible over a period of a week for all synoptic scales (waves longer than 1000 km).



● Figure 4. 36 hr. using the "ultrafine" 2nd order GISS model with double horizontal resolution

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THE EFFECT OF ACCURACY, CONSERVATION AND FILTERING
ON NUMERICAL WEATHER FORECASTING

ORIGINAL PAGE IS
OF POOR QUALITY

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1. INTRODUCTION

The design of a numerical model for atmospheric simulation is not a straight-forward procedure. Both in the areas of mathematical and numerical analysis, and in the parameterization of physical processes not explicitly resolved, the modeler faces several difficult choices between equally reasonable methods, and sometimes between similarly unsatisfactory methods.

In this paper we discuss the considerations leading to the numerical design of the GLAS Fourth-Order Global Atmospheric Model. This model, which was briefly described in Kalnay-Rivas et al., [1977], has been restructured, and several minor changes were introduced. *The computation time and memory requirements for the 4th order model are now similar to those of the present second order GLAS model with the same 4° latitude, 5° longitude and 9 vertical-level resolution [Somerville et al., 1974]. However, the fourth-order model forecast skill is significantly better than that of the current GLAS model, and after 3 days it is comparable or better than that obtained with the 2.5° by 3° version of the GLAS model.*

A discussion of several of the basic characteristics of the model design is contained in section 2. For each of them we present some of the possible alternatives, their advantages and disadvantages, and the reason for our choice. In section 3, we discuss the effect on numerical forecasts of changes in the accuracy, resolution and conservation properties of the models. ⁽¹⁾ Section 4 contains some final remarks.

(1) As of this writing, several of the numerical experiments are not complete.

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2. DISCUSSION OF THE DESIGN OF THE MODEL

Different numerical analysts and atmospheric modelers often take different approaches in the design of a numerical model for weather prediction. Sometimes they even have different basic philosophies. For example Arakawa [1966, 1972], has been a pioneer in the development and use of numerical schemes that reproduce as closely as possible the conservation properties of the continuous equations of fluid dynamics that determine the motion of the atmosphere. On the other extreme, Kreiss and Olinger [1972, 1973] have advocated the use of more accurate schemes even when they don't formally satisfy any conservation properties.

The current GLAS model is based on Arakawa's [1972] second order scheme using a staggered grid B. The scheme is energy conserving and approximately enstrophy conserving for non-divergent flow. In this section we isolate the areas in which the Fourth-Order model is different from the GLAS model, discuss some of the alternatives, and the justification of our choice.

2.1 Accuracy

There is a consensus among modelers that for finite difference models with second-order accuracy, horizontal resolution of about 400 km, and about 10 vertical levels, horizontal truncation errors are the most important source of errors. The truncation errors can be reduced by any of the following methods:

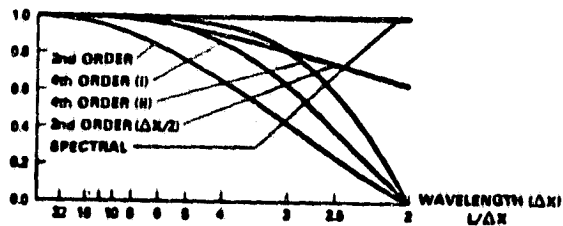
a. Increased horizontal resolution, retaining second-order differences.

Advantages: If the horizontal grid size is reduced by a factor of 2, truncation errors are reduced by a factor of 4. A comparison of the effect of truncation errors on the computational phase speed and group velocity of a linear wave is presented in Fig. 1. Another advantage is that increasing

the resolution allows smaller but possibly important scales to be explicitly included in the model.

Disadvantages: The reduction of error is slow, and the computation time is increased by a factor of 8.

COMPUTATIONAL PHASE SPEED c/c



COMPUTATIONAL GROUP VELOCITY c_g/c

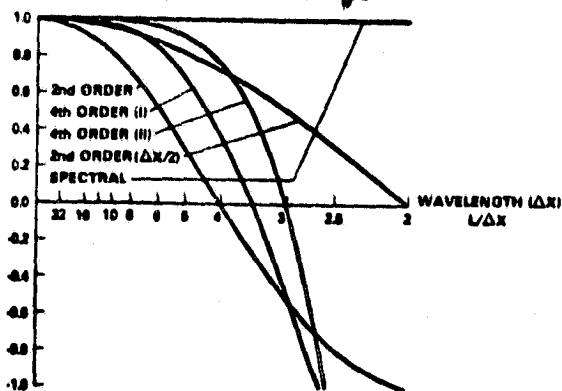


Fig. 1 Computational phase speed and group velocity for a linear wave of constant phase speed c , for different numerical schemes as a function of wavelength L .

b. Fourth order schemes of the first kind. We include in this group the common 5-point explicit fourth order schemes [Kreiss and Olliger, 1972].

Advantages: Phase speed errors are much smaller, especially for long waves (Fig. 1). Waves longer than $5\Delta x$ have smaller errors in their phase speed with fourth order differences than with second order differences with twice the resolution. If we consider the computational group velocity, the crossover occurs at about $7\Delta x$. The computational time is only increased by a factor between 2 and 3. **Disadvantages:** Waves in the range of $2\Delta x$ to $4\Delta x$ are still grossly misrepresented.

c. Fourth order schemes of the second kind. In this group we include linear finite elements, cubic splines and the "compact" or Pade type 3-point 4th order differences introduced by Kreiss [Orszag & Israeli, 1972], all of which have the same fourth order accuracy.

Advantages: These fourth-order differences are considerably more accurate than the fourth-order differences of the first kind (Fig. 1). The crossover with double-resolution second-order differences in the phase speed error occurs at $3\Delta x$, and in the group velocity error at about $4\Delta x$.

Disadvantages: In the simplest case, Kreiss' fourth order differences require the solution of tri-diagonal matrices. Finite element schemes require the solution of at least block tri-diagonal matrices. Even though there are efficient methods to perform these inversions, they are still computationally expensive. The extra accuracy can be compensated with 4th order differences of the first kind by increasing the resolution, which has other advantages, as we mentioned earlier. These schemes, as well as in higher order finite difference schemes, waves in the range of $2\Delta x$ to $4\Delta x$ are still grossly misrepresented.

d. Spectral schemes.

Advantages: If the basis of the spectral expansion are the eigenfunctions of the wave equation, spectral schemes have no phase speed errors. Because of this they require less resolution than finite difference schemes. **Disadvantages:** Because of the large number of computations required for the nonlinear terms, spectral schemes are competitive with finite difference schemes only in combination with the use of less resolution and semi-implicit time schemes.

Our choice: We chose to use fourth order finite differences of the first kind because they are computationally efficient and have small truncation errors except in the range of waves with wavelengths between $2\Delta x$ and $4\Delta x$.

2.2 Type of Grid

Both staggered and unstaggered grids have been widely used by atmospheric modelers.

a. Unstaggered grid: The advantage of this grid is its simplicity. Higher order schemes are easily developed with this grid. Its disadvantage is that all centered differences have to be computed over a distance of $2\Delta x$.

b. Staggered grids: Several staggered grid configurations are possible as reviewed by Arakawa [1972]. The one he called scheme C, which is the most commonly used, has the pressure defined at the center of a grid cell, and the velocity components u and v defined at their corresponding normal walls. This grid has the advantage that the pressure gradient and velocity divergence terms are computed over a distance of only $1\Delta x$, so that inertia gravity waves are computed with double resolution. Therefore, as pointed out by Arakawa, geostrophic adjustment is best represented in this grid. On the other hand, advection terms are computed with no more accuracy than in the unstaggered grid, and in the Coriolis' acceleration term, it is necessary to take horizontal averages of the velocities. The

higher resolution of inertia gravity waves reduce the maximum time step for explicit time schemes by a factor of two. Full fourth order schemes can be developed with staggered schemes but they are very involved [Kalnay-Rivas, 1976]. To date, modelers using staggered grids have introduced fourth order differences *only* in the advection terms.

Our choice: During extended range forecasts, second order errors in the non-advective terms may become important. For this reason we chose to use an unstaggered grid that allows the use of a simple, full fourth-order scheme.

2.3 Conservation Properties and the Use of Horizontal Diffusion

With respect to conservation properties, there are basically 3 types of finite difference schemes for the primitive equations:

a) Nonconservative schemes, the simplest of which is the one based on the advective form of the equations; b) quadratically or energy conserving schemes; and c) enstrophy conserving schemes. Advective and quadratically conservative schemes can be easily developed using staggered or unstaggered grids [Lilly, 1965; Bryan, 1966]. Enstrophy conserving schemes for the primitive equations have been developed on a grid C by Grammelvedt [1969] and Arakawa and Mintz [1974]. Sadourny [1965a, b] constructed a potential enstrophy conserving scheme on grid C, and Arakawa [1978] has recently developed a potential enstrophy and energy conserving scheme also on grid C.

Nonconservative schemes require a procedure to damp waves shorter than $4\Delta x$, which otherwise grow spuriously causing catastrophic nonlinear instability [Phillips, 1959]. This has usually been done by means of linear or nonlinear horizontal diffusion, or by using dissipative numerical schemes such as the Lax-Wendroff, or schemes that contain explicit horizontal averaging.

Quadratically conservative schemes avoid the unbounded growth of the solutions associated with catastrophic nonlinear instability. However, as Arakawa [1972] and Sadourny [1975a, b] have pointed out, in the course of long integrations, there is still a spurious build up of energy in the shortest waves, which appears as an *unbounded growth of the total enstrophy*. In the absence of horizontal diffusion, this type of *slow nonlinear instability* will completely distort the solution. Enstrophy conserving schemes, on the other hand, impose a stronger constraint on the growth of the smallest scales present in the model. For this reason, the UCLA and the GLAS models, which use enstrophy conserving schemes, *do not need to include horizontal diffusion*.

It should be emphasized that conservation of enstrophy does not necessarily imply a more accurate or realistic simulation. In the real atmosphere, the constraint of quasi-geostrophic motion implies that very little of the energy generated in the baroclinically unstable scales can reach the smallest scales

and be eventually dissipated [Charney, 1972]. In a numerical model, the finite resolution imposes an artificial "wall" at the short end of the spectrum, inducing an excessive accumulation of energy in the shortest waves. This problem is worst in nonconservative schemes, but it appears even in alias-free, energy- and enstrophy-conserving spectral models. This justifies using some parameterization of the unresolved subgrid eddies to withdraw energy from the smallest resolved scales.

For this purpose, Leith [1972] suggested the use of nonlinear horizontal diffusion in which the eddy diffusion coefficient is proportional to the local gradient of vorticity. This formulation is consistent with the transfer of energy to subgrid scales in two-dimensional turbulence. In another widely used formulation, suggested by Smagorinsky [1963] and based on a three-dimensional turbulent cascade theory, the diffusion coefficient is proportional to the deformation tensor. These formulations are better than the use of linear diffusion, but they both share the following problems: a) the diffusion coefficient is computed inaccurately for the shortest waves, and, more importantly, b) when the diffusion coefficient is large enough to avoid the spurious growth of the smallest scales, it produces excessive damping of the larger scales [Merileea, 1975; Williamson, 1978]. Furthermore, at the short end of the spectrum (scales of the order of 100 km), neither quasi-geostrophic nor 3-dimensional isotropic turbulence theories are really justified. Williamson [1978] generalized a higher order diffusion of the form $\nabla^2 \kappa \nabla^2$, suggested by Kreiss and Oliger [1972], still using the deformation type of diffusion coefficient. This formulation has the advantage that, because it is more scale dependent, there is less damping of the longer waves.

In our model, we have taken an approach closer to Fourier filtering the shortest waves, as suggested by Phillips [1959]. Our "subgrid parameterization" is based on the following argument: The 4th order scheme is adequate accurate for waves longer than $4\Delta x$, but grossly inaccurate for waves between $2\Delta x$ and $3\Delta x$. Since the shortest waves *cannot provide any useful information in a finite difference scheme*, we filter them out of the system while their amplitude is still small. Even though they are filtered out, the shortest waves still play an important role in the model: they act as a buffer or "sponge layer" in the spectrum domain, allowing energy to trickle down from longer waves and avoiding the accumulation of energy that would otherwise occur at the short wave cutoff.

The elimination of the short waves is performed in the model with the periodic use of a 16th order Shapiro [1970] filter, first introduced in GCMs by Francis [1975]. Figure 2 indicates the response of Shapiro filters of order 4, 8 and 16. It can be observed that in the case of a 16th order filter waves longer than 4 are scarcely affected at all

even after 128 applications, which in our model correspond to a 10 day integration. Waves shorter than 4 are virtually eliminated. Lower order filters introduce too much decay at long scales.

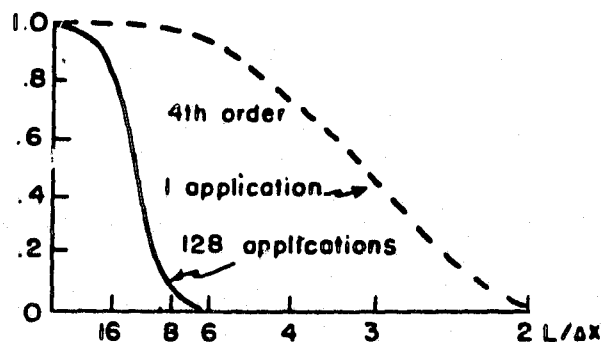
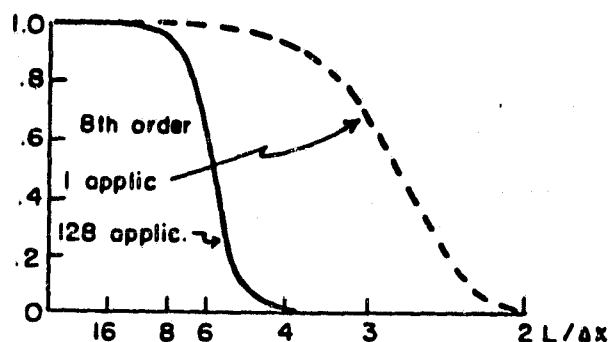
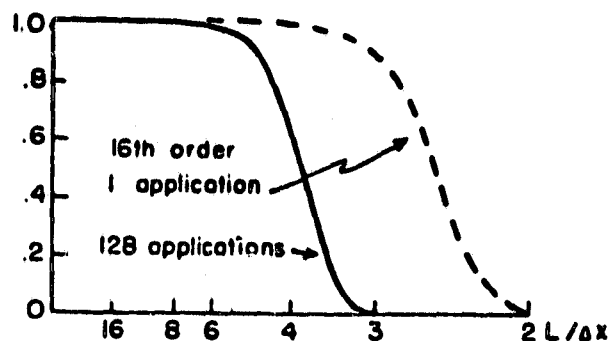


Fig. 2 Response of Shapiro filters after 1 and after 128 applications.

In order to explore the effect of filtering short waves we performed a series of 20 day forecasts with a simple shallow water equation model [Kalnay-Rivas, 1976], using different combinations of schemes, smoothing operators and frequency of application as indicated on Table 1.

A1: 4th order, non-conservative, no smoothing.
 A2: " , " , linear diffusion, $\nu = 10^3 \text{ m}^2/\text{s}$
 A3: " , " , 16th order filter/4 hrs.
 A4: " , " , " /time step.

B1: " , quadrat. conserv., no smoothing.
 B2: " , " , 16th order filter/4 hrs.
 B3: " , " , 8th " " "
 B4: " , " , 4th " " "

C: 2nd order, " , no smoothing

TABLE 1: Characteristics of the different runs in Fig. 3.

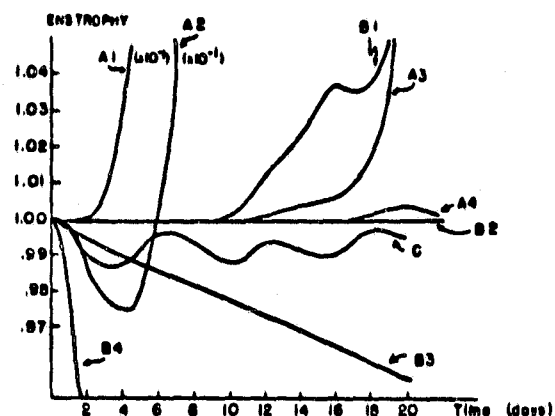
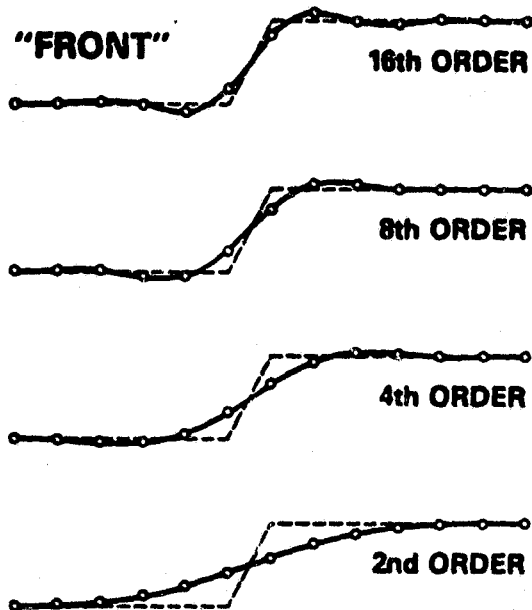


Fig. 3 Variations of total enstrophy during numerical integrations as indicated in Table 1. The scale corresponding to experiments A1 and A2 is multiplied by 10^{-1} .

We found that all stable runs conserved total energy with a high degree of accuracy. Fig. 3 shows the variation in time of the total potential enstrophy, which is conserved exactly in the continuous equations. The results indicate that when the 16th order filter is applied every time step (10 minutes) even a formally nonconservative scheme conserves both total energy and total potential enstrophy during a long integration (Run A4). The quadratically conservative scheme controls better the amount of energy going into the smallest scales, so that in Run B2 it was enough to apply the filter every 4 hours to conserve potential enstrophy within 0.05%, even though such conservation is not formally guaranteed in the scheme.

It may be questioned whether the application of a high order filter eliminates small scale features like frontal zones or the effects of orographic or cumulus convection forcing on small scales. Fig. 4 indicates that this is true for lower order filters. However a 16th order filter eliminates only those components which are not resolved anyway, and still allows for the formation of sharp gradient zones and strong local maxima.

10 PASSES (ONE DAY) OF SHAPIRO FILTER



10 PASSES (ONE DAY) OF SHAPIRO FILTER

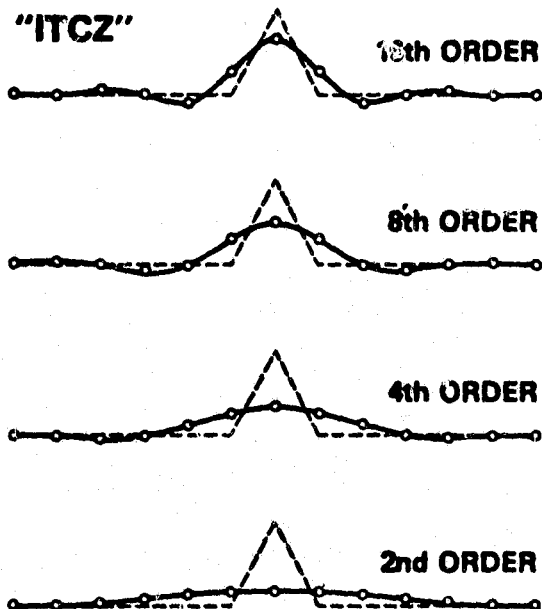


Fig. 4 Effect of 10 applications of a 16th order Shapiro filter on a) a step function b) a spike.

Our choice: We chose to use a quadratically (energy) conserving scheme because it is only slightly less efficient than an advective scheme and it requires the application of a high-order filter only every few hours to avoid the effect of slow nonlinear instability associated with a spurious growth of enstrophy. We apply the filter to the sea level pressure and potential temperature fields in order to compensate for the effect of topography. (3)

3. EXPERIMENTAL FORECASTS: PRELIMINARY RESULTS

We plan to perform an extensive series of forecasts to study the effect of using different schemes and varying resolution on the quality of actual weather forecasts. In this section we present some preliminary results.

We tested the Fourth Order model by making 3-day forecasts from several initial conditions. In every case the model performed much better than the GLAS model with the same 4° by 5° resolution. After 3 days the forecasts were comparable to the 2.5° by 3° version of the GLAS model in the sea level pressure maps, and had slightly less phase errors in the 500 mb maps.

Fig. 5a shows the verification sea level pressure map corresponding to a 3-day forecast with February 19, 1976, 0Z (a case that has been studied in detail by Atlas et al., 1979). Fig. 5b is the 3-day forecast with the Fourth Order 4° by 5° model verifying on the same date. The excessive gradients, especially at high latitudes are due to a faulty computation of the ground temperature by the radiation routine used in that run. Figs. 5c and 5d are the forecasts generated by the GLAS models with 4° by 5° and 2.5° by 3° resolution respectively. Fig. 6 displays the corresponding 500 mb maps.

In Fig. 7a, we present the verification sea level pressure map corresponding to a 3-day forecast with February 1, 1976, 0Z initial conditions. The results of six different 3-day forecasts are shown in Figs. 7b to 7g. The forecast in Fig. 7b was computed with the new full Fourth-Order model, using the same resolution and parameterization of physical processes (except for the long-wave radiation routine) as in the 4° by 5° GLAS model (Fig. 7c). In Fig. 7d, the model was the same as in 7b, but full Second-Order accuracy was used. If we compare these three made with the same resolution, we see that the fourth order model did considerably better, especially in forecasting the development and motion of the cyclone southwest of Greenland. The two second-order forecasts are close to each other.

(3) The idea of filtering these fields is due to Dr. A. Bayliss.

Fig. 7 shows the forecast made with the 2.5° by 3° GLAS model starting from the NASA initial conditions with assimilation of satellite data [Atlas et al., 1979]. All other forecasts were made from NMC's Global Analysis initial conditions. The cyclogenesis forecast was poorer than with the fourth order model, but over continental North America the forecast was better. This may be due to the higher resolution or, possibly, improved initial conditions.

Fig. 7f presents the forecast made with NMC's 6-layer, 380 km resolution, which has resolution comparable to our 4° by 5° grid. For these initial conditions, NMC's forecast was better than the GLAS second order forecast, although it shared some of its errors (such as a spurious anticyclogenesis over the Great Lakes).

Fig. 7g shows a forecast made with a slightly different version of the fourth order model. The differences were as follows: In the 7g forecast we used a Matsuno time step, a surface drag that was .75 of that of the standard model, the Shapiro was applied every two hours, and a slightly different scheme for the vertical advection of moisture was used. In the 7b forecast we used a combined Matsuno-leap-frog scheme, the surface drag was the same as in the standard model and the filter was applied every hour. The two forecasts are quite similar, but from other experiments it seems that the positions of the oceanic lows west of Spain and south of Alaska were somewhat affected by the change in frequency of the Shapiro filter.

4. FINAL REMARKS

Although these are preliminary results, the new Fourth-Order model seems to have very good forecasting skill. Similar excellent results with fourth order schemes were reported by Campana [1978, 1979], and by Williamson (1978). Campana found that most of the improvement over the second order model was obtained just by introducing fourth order accuracy on the horizontal averages performed on the Shuman-Hovermale model. Campana also obtained that for 2-day forecasts, fourth order differences were important in the advection terms but not in the pressure and continuity terms. We made shallow water equation experiments that showed some deterioration of the solution after about 8 days when only the advection terms were written with fourth order accuracy. Since baroclinic models are much more unstable than the SWE, Campana's results may not hold during extended forecasts. We are performing experiments to study this possibility in our model. We are also repeating some of the forecasts using a vector invariant form of the momentum equations. The NMC 6-layer PE model used in Fig. 7f is based on such a formulation, and we want to determine if it affects the solution.

We are performing integrations with a 2.5° by 3° version of the Fourth-Order model. From linear theory, this should make horizontal truncation errors negligible for most synoptic scales for periods of a week or longer. Still, narrow atmospheric features like those due to sharp orographic forcing on the ITCZ will remain poorly resolved.

We have found the fourth order model to be extremely sensitive to the parameterization of physical processes. For example, it became unstable through excessive cooling near the surface, apparently due to a flaw in the radiation routine. None of the second order models was sensitive to this problem.

In this paper we have discussed the use of formally nonconservative schemes coupled with periodic filtering of the shortest waves as an alternative to the use of conservative schemes. It is clear from the experimental results that formal enstrophy conservation has had no beneficial impact on the quality of the forecast (Figs. 7c and 7d). Conservation of potential enstrophy [Arakawa, 1978] might have a more positive effect.

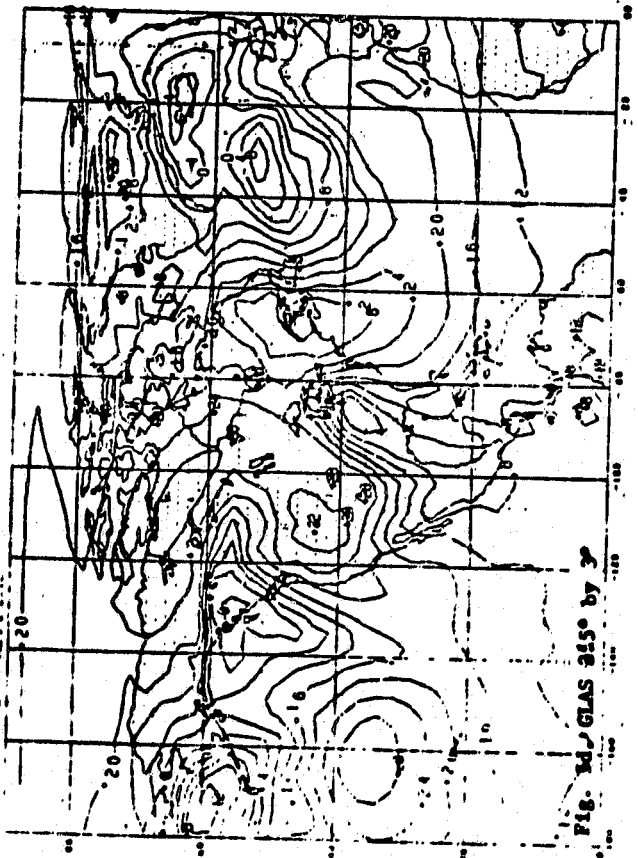
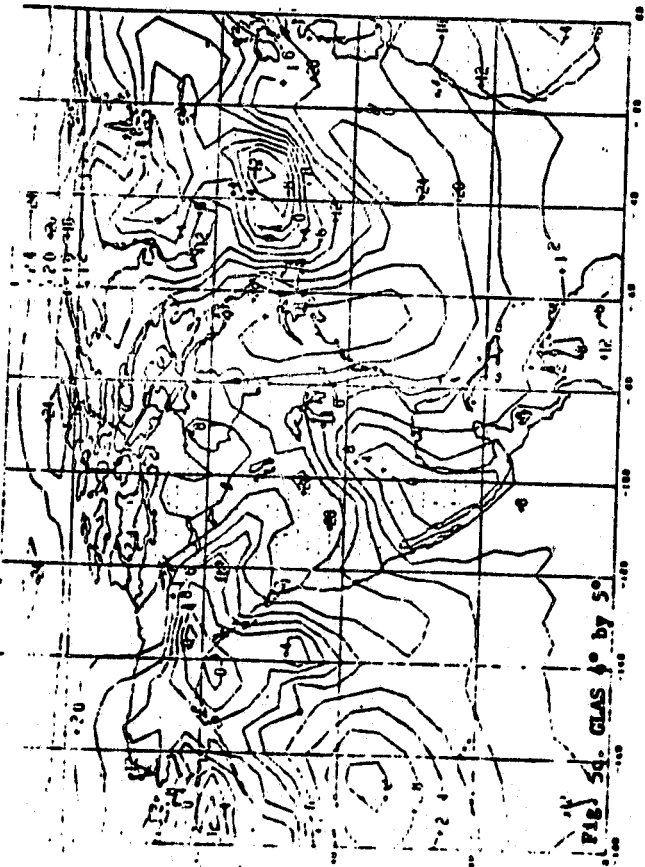
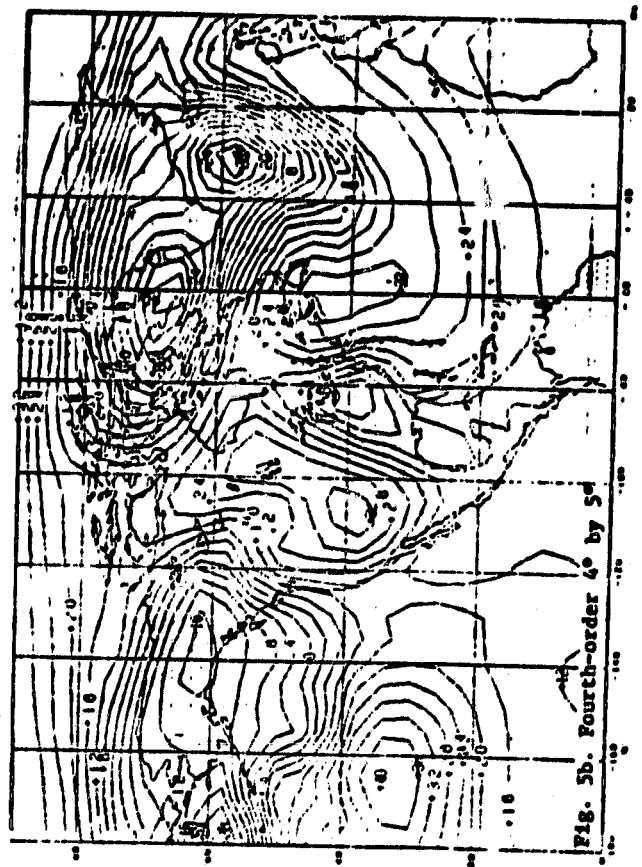
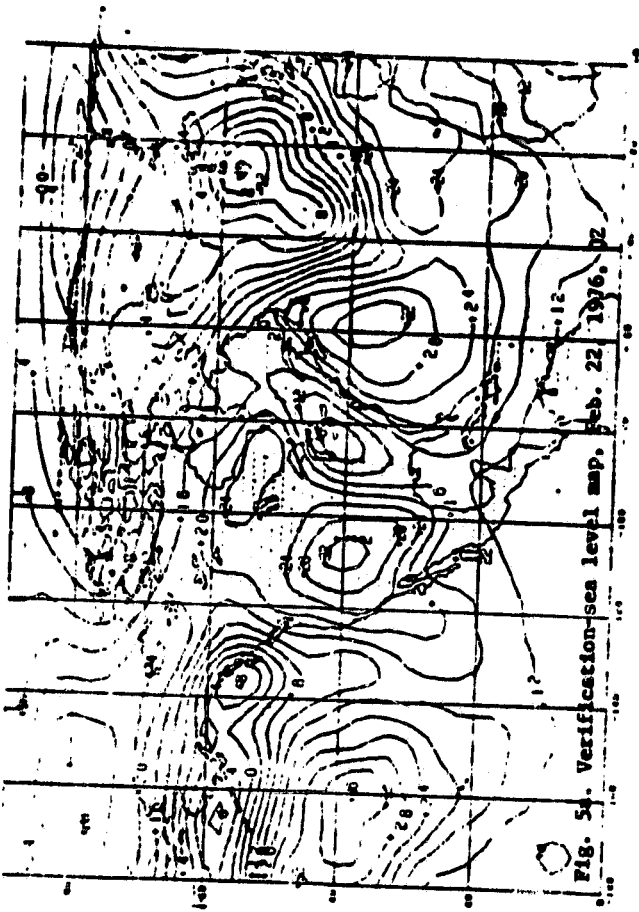
We think that the use of nonconservative schemes with high order periodic filtering is also justified in climate simulations as long as the rates of energy or enstrophy loss due to filtering remain much smaller than the observed rates of generation and dissipation.

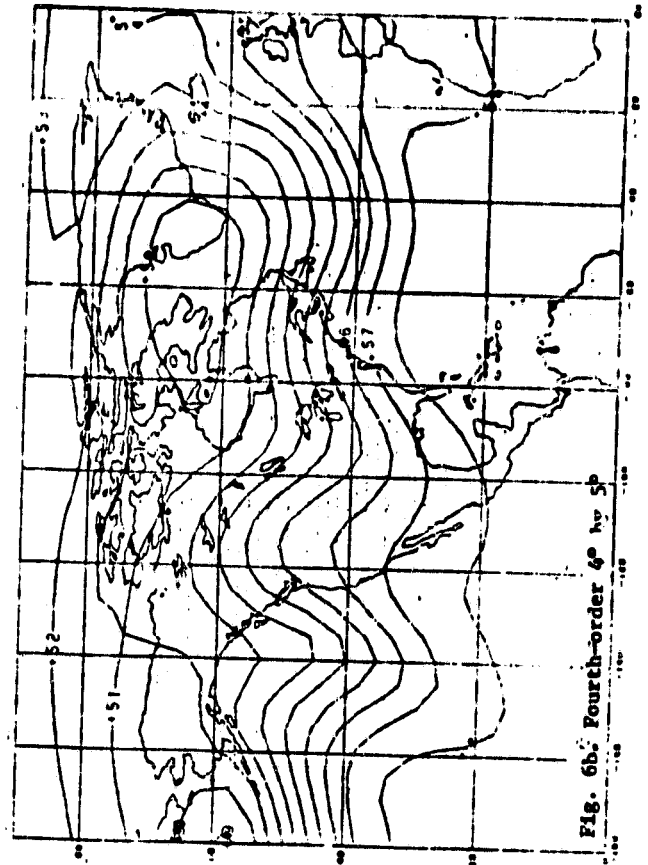
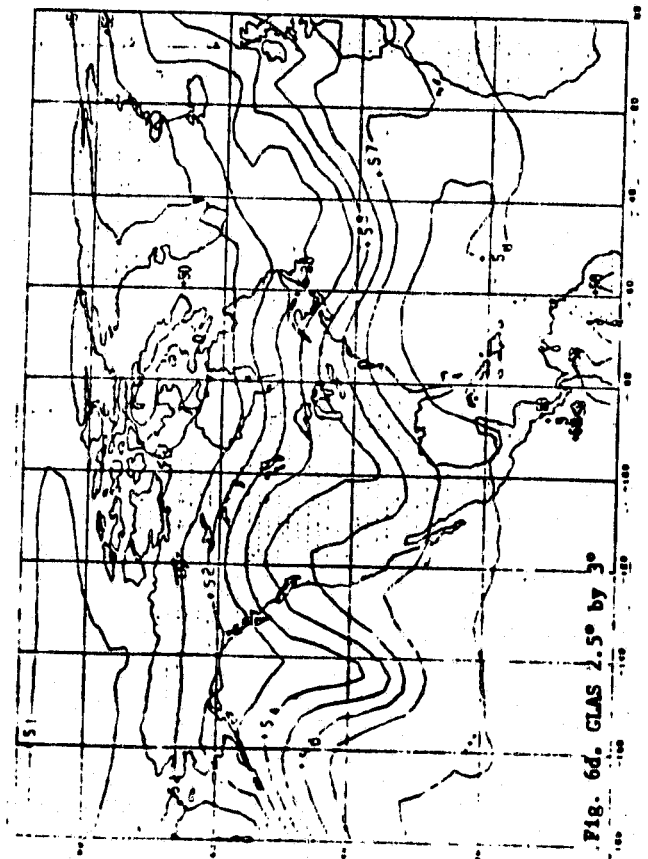
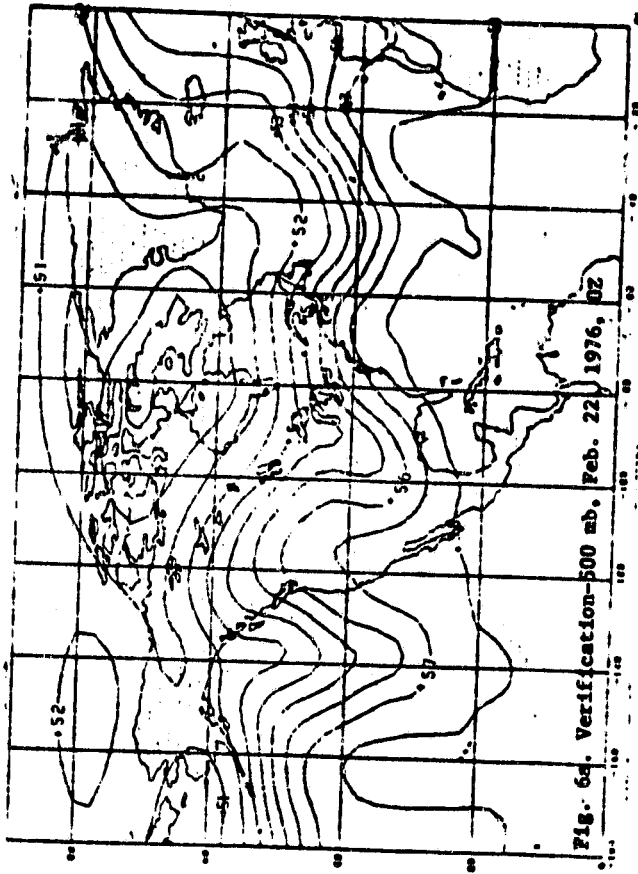
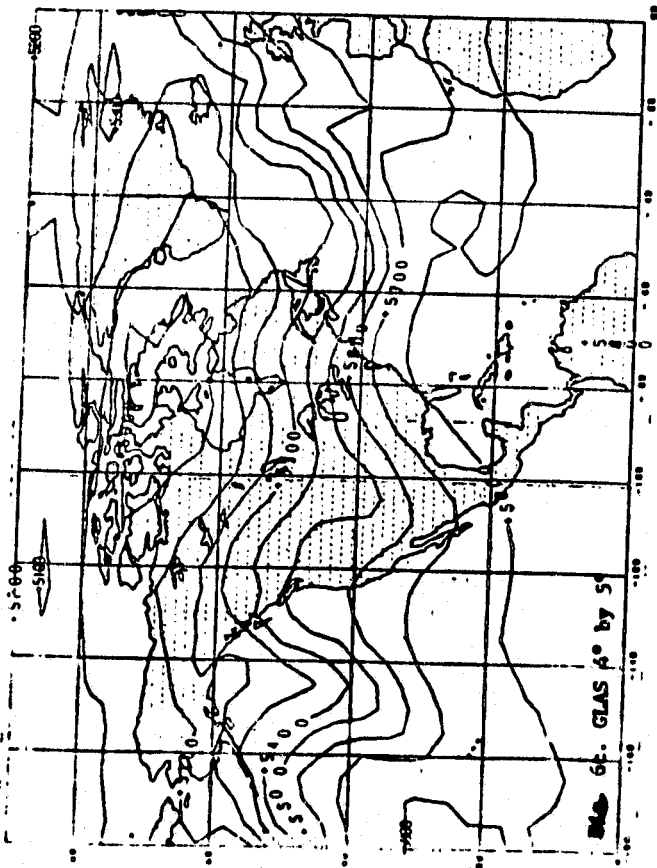
Acknowledgements: Dr. A. Bayliss' contribution was crucial in the development of the model. The authors have benefitted from stimulating discussions with Profs. M. Cane, E. Isaacson and D. Randall. The authors are grateful to Dr. M. Halem for his patience, encouragement and many useful suggestions. Dr. N. Rushfield and Mr. W. Connelly were extremely helpful in the development of the program. Dr. M. Wu provided us with her accurate radiation routine.

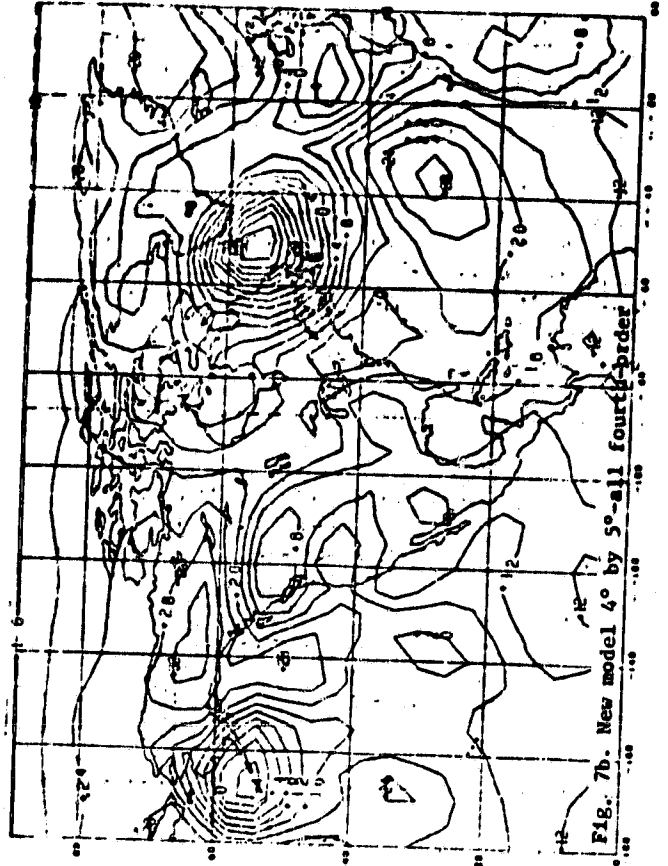
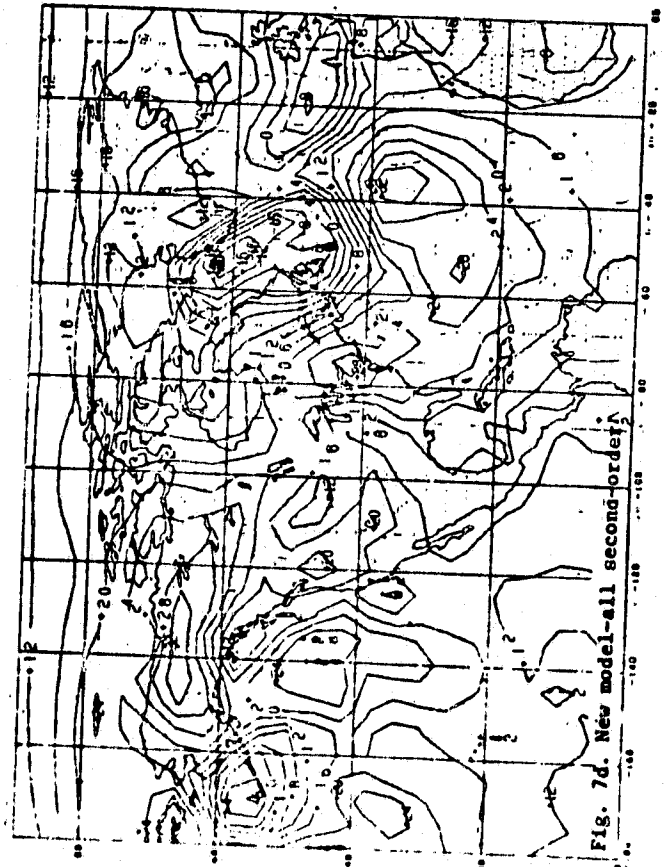
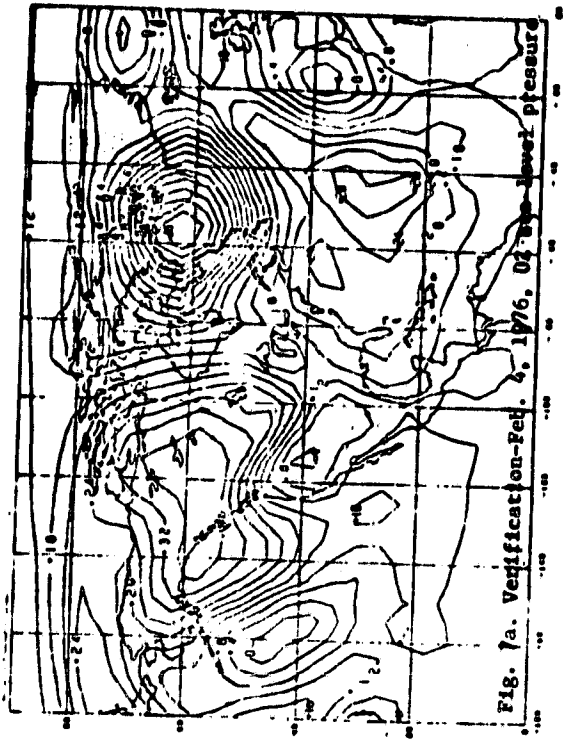
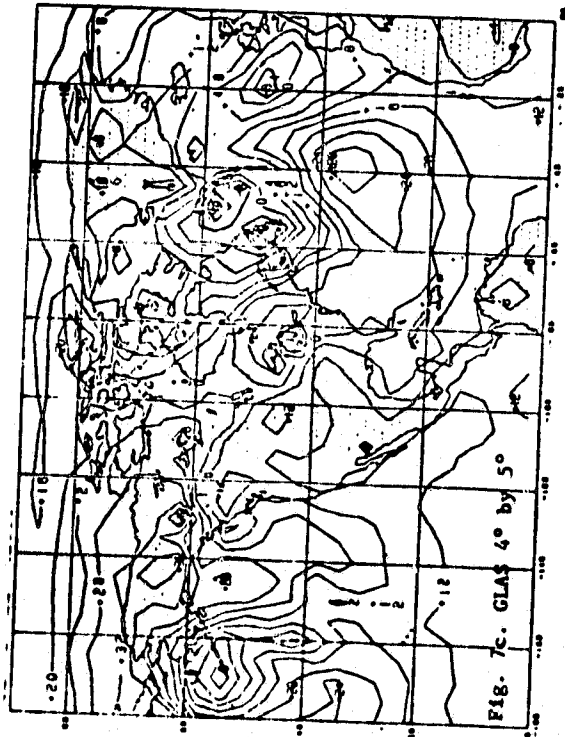
Fig. 5 Sea level pressure maps corresponding to 3-day forecasts with February 19, 1976, OZ initial conditions.

Fig. 6 Same as Fig. 5 but 500 mb geopotential height.

Fig. 7 Same as Fig. 5 but for February 1, 1976, OZ initial conditions.







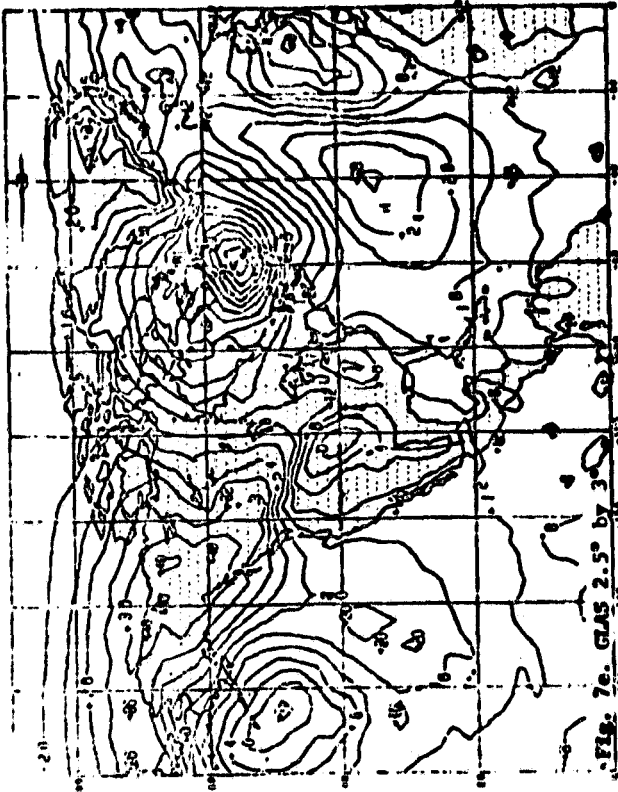


Fig. 7e. GLAS 2.5° by 3°

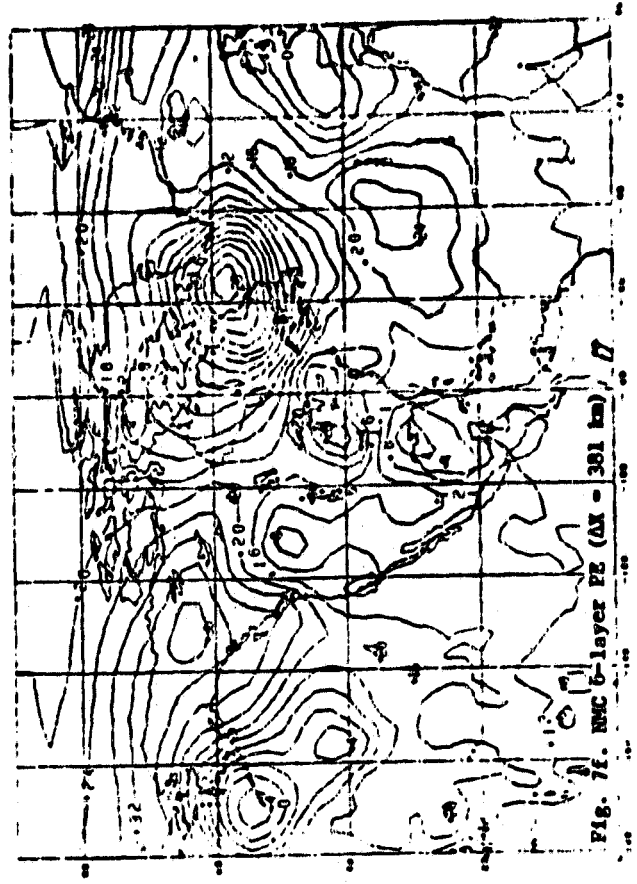


Fig. 7f. NMC 6-layer PE (ΔX = 361 km)

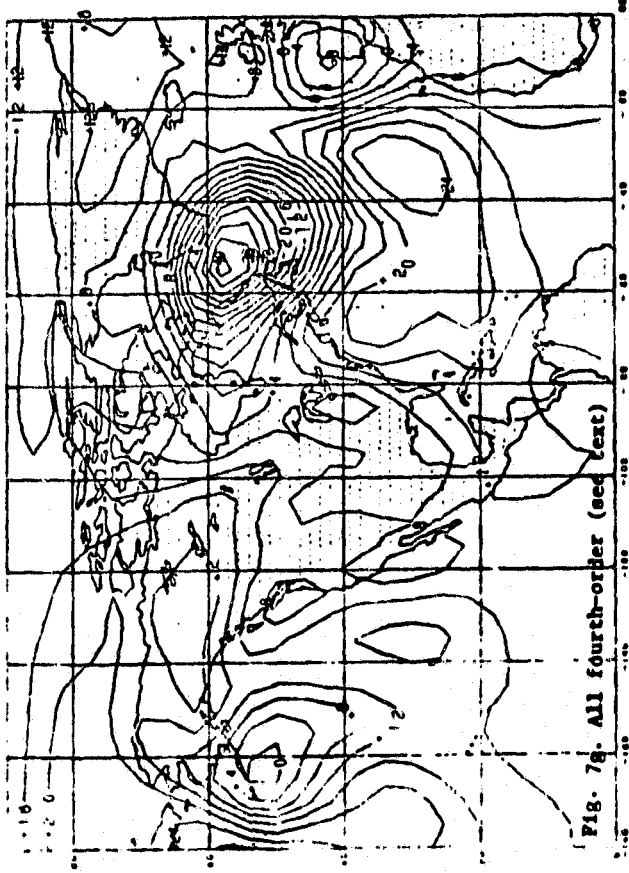


Fig. 7g. All fourth-order (see text)

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1SN 0031      EVENT(ITAU)=MOD(ITAU,INTFX(XTAU)),LT,(DT+P  
1SN 0032      CALL SUPND  
1SN 0033      CALL DATE(ADATE)  
1SN 0034      CALL TIME(ATIME)  
1SN 0035      CALL CLOCKS(ITM)  
1SN 0036      NDS=1  
1SN 0037      ITIME=ITM  
1SN 0038      LTIME=ITM  
1SN 0039      CALL INPUT  
1SN 0040      CALLAVE  
1SN 0041      DT=24*60  
1SN 0042      NDC=0 FOR PREDICTOR STEP  
1SN 0043      NDC=0 FOR CORRECTOR STEP  
1SN 0044      TAU=TAU+DT  
1SN 0045      DIMED=1/3600  
1SN 0046      IDT=INT(XIDTHR)  
1SN 0047      XST=INT(1/2*IDT)  
1SN 0048      NSTEP=3+AVG(AUTAU)/DTHR  
1SN 0049      WRITE(3,26)NSTEP  
1SN 0050      FORMAT(1/30*,FIRST NSTEP=,15//)  
1SN 0051      26  
1SN 0052      NSTEP=0  
1SN 0053      I1=ISEG(0),I1M1F(2)=MAISUN  
1SN 0054      PRTSNK=MLF(1)  
1SN 0055      ITAVENSTP=IDTHR  
1SN 0056      TAU=FLOAT(ITAU)/XINT  
1SN 0057      IDAY=TAU/12  
1SN 0058      IDAY=FLOAT(ITAU-IDAY*12)/XINT  
1SN 0059      CALL DAILY0  
1SN 0060      IF(I1AL ISR,GI,0) CALL ALTER  
1SN 0061      WRITE(3,20)PUTZ)  
1SN 0062      TAU=TAU  
1SN 0063      TAPNUM=LABEL(20)  
1SN 0064      WRITE(3,901)TAPNUM,TAUST,TAU,TAUT,XLABB  
1SN 0065      #CODE=BLANK  
1SN 0066      WRITE(3,24)  
1SN 0067      IF(I1START,LE*) GO TO 45  
1SN 0068      C**** MAIN LOOP  
1SN 0069      45  
1SN 0070      #CODE=BLANK  
1SN 0071      IF(I1ALTER,GT,0) CALL ALTER2  
1SN 0072      45 CONTINUE  
1SN 0073      NSTEPT=NSTEP+NSTEPC  
1SN 0074      CONG = .ALX  
1SN 0075      IF(.NOT.EVENT:TAUT) GO TO 50  
1SN 0076      CONG = .TRUE  
1SN 0077      50  
1SN 0078      REPLACE DT WITH MAISUN, TIME-STEP SIZE SINCE  
1SN 0079      C ANY RESTART NEEDS TO BEGIN WITH A MAISUN( STEP  
1SN 0080      TEMP = DT  
1SN 0081      DT = DTSAVE  
1SN 0082      CALL CLOCKS(ITM)  
1SN 0083      #CODE=INTAPE  
1SN 0084      CALL TWRITE  
1SN 0085      C INITIALIZE PRECIPITATION ARRAY.  
1SN 0086      ITAU = TAU + 0.01  
1SN 0087      IF (MOD(ITAU,12) *NE, 3) GO TO 47  
1SN 0088      DO 46 I = 1, IM  
1SN 0089      SDO(I, J, 3) = 0.  
1SN 0090      SDO(I, J, 4) = 0.  
1SN 0091      46 CONTINUE  
1SN 0092      47 CONTINUE  
1SN 0093      C  
1SN 0094      C REPLACE DT WITH CURRENT STEP-SIZE  
1SN 0095      C DT = TEMP  
1SN 0096      C  
1SN 0097      50 CALL CLOCKS(ITM)  
1SN 0098      TDM=FLOAT(ITM-NTH)/100.  
1SN 0099      STP=FLOAT(LNT-NTH)/100.  
1SN 0100      LNT=INTM  
1SN 0101      WRITE(3,902)NSTEP, IDAY, TDFDAY, JDATE, JMONTH, STM, TOTM, TAU  
1SN 0102      I * XLABEL(20), #CODE  
1SN 0103      CALL SBTCH(4,ISS)  
1SN 0104      IF (ISS=EQ) WRITE (15,906)TAU, IDAY, TDFDAY, TAUST, TAUE  
1SN 0105      IF (ISS=EQ) CALL SSCFF  
1SN 0106      CALL SBTCH(16,KSS)  
1SN 0107      IFLKSS=EQ)WRITE(3,27)NSTEP  
1SN 0108      FORMAT(1/30*,LAST NSTEP=,15//)  
1SN 0109      IFLKSS=EQ)GOTO 87  
1SN 0110      IF(TAU=91.5*DT,TAUE)GCTC 88  
1SN 0111      IF(NSTEP=NE)GOTO 86  
1SN 0112      DO 69 J=1, JNP  
1SN 0113      60 CALL COMP3(U,V,T,S,H,P,U,UPDL,VPDL,TPUL,SPDL,PPDL,QPDL,J)
```

```

ISN 0117 86 ITAU=ITAU + IOTHR
ISN 0118 NPC=0
ISN 0119 TAU=FLOAT(ITAU)/X*NT
C * *
ISN 0120 IDAY=ITAU/124
ISN 0121 YFDAY=FLCAT(ITAU-IDAY*124)/XINT
ISN 0122 NSTEP=NSTEP+1
ISN 0123 IF(INSEQ.GT.1)GO TO 12
ISN 0124 MATSUN=MAISNX
ISN 0125 MATSNX=MLF(12)
ISN 0126 GO TO 11
ISN 0127 12 NTH=MOD(NSTEP-1,NSEQ)+1
ISN 0128 NEXT=MOD(NSTEP,NSEQ)+1
ISN 0129 MATSUN=MLF(NTH)
ISN 0130 MATSNX=MLF(NEXT)
ISN 0131 CONTINUE
ISN 0132 DT=DT*SAVE
ISN 0133 IF(NATSUN.EQ.0)DT=DT*52
ISN 0134 CALL COMPOU.V.T.SM.P.Q.UPOL.VPOL.TPOL.SMPOL.SPOL.OPOL.VT.VT.T.
ISN 0135 $ SMT.P.Q.UT.VPOL.VT.VT.T.PPOL.SHTPOL.PIPOL.OTPOL)
ISN 0136 IF((NATSUN.EQ.1)OR.(NSTEP.EQ.1))GO TO 84
ISN 0137 IF(INDS.EQ.0)GO TO 14
ISN 0138 IF(NDS.EQ.0)GO TO 14
ISN 0139 NSM=MOD(NSTEP-NSM1,(SMTH)
ISN 0140 IF(NATSUN.EQ.1)GO TO 14
ISN 0141 IF(NSM.GT.1)GO TO 14
ISN 0142 CALL SMSHAP(TT.PT.OT.VT.VT.PPOL.PPOL.PPOL.PPOL)
ISN 0143 14 CONTINUE
ISN 0144 N*NS=1
ISN 0145 IF(NATSUN)OR FIRST LEAP FRAG (STEP)
ISN 0146 C *** NPC=1
ISN 0147 IF((NATSUN.EQ.1)OR.(NSTEP.EQ.1))GO TO 84
ISN 0148 NSTEP=NSTEP+1
ISN 0149 IF(INSEQ.GT.1)GO TO 15
ISN 0150 MATSUN=MAISNX
ISN 0151 MATSNX=MLF(12)
ISN 0152 GO TO 18
ISN 0153 15 NTH=MOD(NSTEP-1,NSEQ)+1
ISN 0154 NEXT=MOD(NSTEP,NSEQ)+1
ISN 0155 MATSUN=MLF(NTH)
ISN 0156 MATSNX=MLF(NEXT)
ISN 0157 CONTINUE
ISN 0158 ITAU=ITAU + IOTHR
ISN 0159 TAU=FLOAT(ITAU)/X*NT
ISN 0160 IDAY=ITAU/124
ISN 0161 TOFDAY=FLOAT(ITAU-IDAY*124)/XINT
ISN 0162 CALL COMPOU.V.T.SM.P.Q.UPOL.VPOL.TPOL.SMPOL.SPOL.OPOL.VT.VT.T.
ISN 0163 $ OTPOL.V.T.SM.P.Q.UT.VPOL.VT.VT.T.PPOL.SHTPOL.PIPOL.OTPOL)
ISN 0164 IF((SMTH.EQ.0)GO TO 17
ISN 0165 IF(NSM.GT.1)GO TO 17
ISN 0166 IF(NATSUN)OR.(NSTEP.EQ.1)GO TO 17
ISN 0167 CALL SMSHAP(TT.PT.OT.VT.VT.PPOL.PPOL.PPOL.PPOL)
ISN 0168 17 CONTINUE
ISN 0169 IF(NATSUN.EQ.1)INDS=0
ISN 0170 IF(EVENT(24)) CALL DAILY
ISN 0171 GOTO 40
ISN 0172 ***** END OF MAIN LOOP
ISN 0173 87 WRITE(3,905)
ISN 0174 88 CONTINUE
ISN 0175 WRITE(3,904) IDAY,TOFDAY
ISN 0176 STOP
ISN 0177 901 FORMAT(0,TAPE,144,TAUS,2F10.3,F3.3,5X,20A4,/,
ISN 0178 WCODE=,2A4)
ISN 0179 902 FORMAT(5X,STEP,14,5X,SIMUL TIME IS DAY,14,7X,HOUR,F7.3,15,1X,2A,/,
ISN 0180 12F,2,6X,TAU=,F9.3,TAPE=,F4, WCODE=,2A4)
ISN 0181 904 FORMAT(1,PROGRAM HAS TERMINATED AT DAY,14, HOUR,1F,3)
ISN 0182 905 FORMAT(1,SENSE SWITCH 6 ON,1)
ISN 0183 906 FORMAT(1,SSA ON,TAU=F8.2,2X,DAY,14,2X,HOUR,F6.2)
ISN 0184 * RUN START,F8.2,2X,BILL END,F6.2)
ISN 0185 END

```

NAME TAG	TYPE	ADD.	1	2	3	NAME	TAG	TYPE	ADD.	1	2	3	NAME	TAG	TYPE	ADD.	
C	SFA	CE	R44	000000						R44	000388		J	SFA	CE	R44	000388
P	SFA	CE	R44	000210						R44	000388		N	SFA	CE	R44	000388
U	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
DT	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
IM	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
NR	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
PU	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
TT	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
DXM	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
ITM	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
J24	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
J22	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388
MLF	SFA	CE	R44	000570						R44	000388		N	SFA	CE	R44	000388

NAME OF COMMON BLOCK * 1MR#E OF BLOCKCOMMON 0001C HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 CSWRT L#4 000000 N#4 NOC3 L#4 000004 SHSMO L#4 000008 FAST L#4 00000C
 TMP R#4 SHAFTA L#4 000014 ADPCLO L#4 000018

NAME OF COMMON BLOCK * 80T#9E OF BLOCKCOMMON 0385AB HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 S00T R#4 000000 OMEGA I#2 000004 MNMXTS I#4 000008
 COMC L#4 0385AA

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE
 ISDOT 00CF00

NAME OF COMMON BLOCK * -NTR#E OF BLOCKCOMMON 083130 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JSP L#4 000000 JNP L#4 000004 JIN L#4 000008 NLAY L#4 00000C
 PTOP R#4 000010 ESTART I#4 000014 JSP1 L#4 000018 NLAJ L#4 00002C
 FIM R#4 000014 NLAYMI L#4 000018 NLAJMI L#4 000022 JI L#4 000026
 JM L#4 000020 JTEST I#4 000024 ITES I#4 000028 (ROI L#4 00003C
 MROT L#4 000024 INCS L#4 000028 JSD L#4 000032 JNB L#4 000036
 JAVS L#4 000028 DLON R#4 000032 IDI L#4 000036 JTAU L#4 000040
 DLAT R#4 000032 XINI R#4 000036 JDMNTH I#4 000040 JDEAR L#4 000044
 ITAU L#4 000036 NCYCLE I#4 000040 TAUE R#4 000044 NMOGAN L#4 000048
 TOFDAY R#4 000044 NCYCLE I#4 000048 TAUE R#4 000052 NMOGAN L#4 000056
 NSTEP L#4 000048 TAUI R#4 000052 PI R#4 000056 RAS L#4 000060
 STAU R#4 000052 PSL R#4 000056 PSF R#4 000060 RMCH L#4 000064
 TMULT R#4 000056 NLAJ L#4 000060 COSD R#4 000064 RMAX L#4 000068
 KAPA R#4 000064 NLAJ L#4 000068 COSD R#4 000072 RMAX L#4 000076
 NFB R#4 000072 NLAJ L#4 000076 COSD R#4 000080 RMAX L#4 000084
 NLD R#4 000080 NLAJ L#4 000084 COSD R#4 000088 RMAX L#4 000092
 SIND R#4 000088 NLAJ L#4 000092 COSD R#4 000096 RMAX L#4 000100
 DUMMYC R#4 000096 NLAJ L#4 000100 COSD R#4 000104 RMAX L#4 000108
 SIG R#4 000104 NLAJ L#4 000108 COSD R#4 000112 RMAX L#4 000116
 JPS L#4 000108 NLAJ L#4 000112 COSD R#4 000116 RMAX L#4 000120
 KPS R#4 000112 NLAJ L#4 000116 COSD R#4 000120 RMAX L#4 000124
 KBS R#4 000116 NLAJ L#4 000120 COSD R#4 000124 RMAX L#4 000128
 DXP R#4 000120 NLAJ L#4 000124 COSD R#4 000128 RMAX L#4 000132
 F R#4 000124 NLAJ L#4 000128 COSD R#4 000132 RMAX L#4 000136
 PHIS R#4 0C0D50 ALBEDC I#2 N#2

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE
 0 000E70
 SHS 002300 U 0024FC
 C 000000?

NAME OF COMMON BLOCK * #6R#E OF BLOCKCOMMON 01D1C0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 WORXX R#4

NAME OF COMMON BLOCK * =0#E OF BLOCKCOMMON 081400 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 PU R#4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE
 QT 000120
 GW 001680 TY 0017A0

NAME OF COMMON BLOCK * #56#E OF BLOCKCOMMON 000008 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 ISMTH I#4 000000 NPC I#4 000004

NAME OF COMMON BLOCK * * T#E OF BLOCKCOMMON 0035C0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 TOPOG R#4

NAME OF COMMON BLOCK * #DS_7E OF BLOCKCOMMON 000004 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 ISWICH I#4 000000

NAME OF COMMON BLOCK * * S1#E OF BLOCKCOMMON 0005E0 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME

VARIABLE OFFSETE
 V 0019B0
 SM 003030

VARIABLE OFFSETE
 VT 000C50
 SMT 0022E0

VARIABLE OFFSETE
 TS 001890
 P 002F10

VARIABLE OFFSETE
 GT 0008A0
 PT 0021C0

VARIABLE OFFSETE
 UY 00012C
 TY 0017A0

VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 JMIN I N.R. JMAX I N.R. JSUM I V.R. SMTM R N.R.

NAME OF COMMON BLOCK # *M3|,E OF BLOCKCOMMON 300C38 HEXADECIMAL BYTES

VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 NSEQ I 000005 MLF I 000094 MATSNX I 00002C NSM1 I 000030

LABEL ADDR
40 000A44
50 000B7A
60 000C9A
70 000D9E
80 001032

LABEL ADDE
45 000A52
55 000B7A
65 000C9A
75 000D9E
85 001032

LABEL ADDR
90 000B7A
95 000D9E
100 001032

LABEL ADDR
107 000B82
117 000E4A
126 000FAE
135 00111C

OPTIONS IN EFFECT NAME= MAIN,OPT=0,L INECNT=5%,SIZE=100K.

OPTIONS IN EFFECT SOURCE=EBCDIC,NOLIST,NODECK,LCAD,MAP,NCEDIT,LD,NOXREF

STATISTICS SOURCE STATEMENTS = 191 PROGRAM SIZE= 4460

STATISTICS NO DIAGNOSTICS GENERATED

***** END OF COMPILATION *****

LEVEL 19.6-APR 71

OS/360 FORTRAN H AT GISS

852K BYTES OF CORE NOT USED

DATE 12/12/79-07803.03

```
COMPILER OPTIONS - NAME= MAIN,OPT=0,L INECNT=5%,SIZE=100K,
SOURCE=EBCDIC,NOLIST,NODECK,LCAD,MAP,NCEDIT,LD,NOXREF
SUBROUTINE DAILY.ECDDIC,NOLIST,NODECK,LOAD,MAP,NCEDIT,LD,NOXREF
FINE NINE LAYER GLOBAL CDE BLOCK FOR VM MODEL, FEBRUARY 1976
INTEGER*2 ALBERO
REAL*8 KAPA,LAI
COMMON/FOURTH/CON1,CUR2,CON3,CON4,CON5,TP8AR,ALPHA,BETA,I,MD2PI,
$ ADLDP,JMS2,JMS1,CPTH,PSKAPA,RAD,CP,INC,NSTART,CPD2,
$ INO2,RAD1M,JEND(2),JPI,JP2,MAISUN,JNE XI(2),MDPKI(2,2),
$ PSIGN(4),POLES1(48),W(72,9,5),PK(72,9,3),DIFF(9),
$ EQUIVALENCE : OPOL(1,1),PHISPL(2),SUM(2,3)
$ (VPOL(1,1),POLES(1,1)),(VPOL(1,1),POLES(1,1)),(VPOL(1,1),
$ POLES(1,1)),(SHPOL(1,1),POLES(28,1)),(UTPOL(1,1),POLES(73,1)),
$ (VTPOL(1,1),POLES(82,1)),(TTPOL(1,1),POLES(91,1)),(SHPOL(1,1),
$ POLES(100,1)),(PPOL(1,1),POLES(14,1)),(PTPOL(1,1),POLES(147,1))
$ DIMENSION PPOL(2),PTPOL(2),UPOL(36,1),VPOL(36,1),TPOL(36,1),
$ SHPOL(36,1),OPOL(9,4,2),Q(72,10,4,1)
$ DIMENSION PTPOL(2),UTPOL(36,1),VPOL(36,1),TPPOL(36,1),
$ SHPOL(36,1),QTPPOL(9,4,2),QT(72,10,4,1)
COMMON /CNTRL/
$ JSP,JNP,IM,NLAY,PTOP,ISTART,JSPPI,JNPMI,FIM,NLAYMI,NLAYPI,
$ J1,JM,KM,TAUT,IRCT,MROT,JTEST,ITEST,
$ NR,JAYS(12),INCS(1),JSB,JNB,DLAT,DLON,
$ DT,TAU,ITAU,XINT,(DAY,JDAY,TOPDAY,JDAT,MONTH(2),JYEAR,NSTEP,
$ NCYCLE,NCOMP3,MOGAT,TAUP,TAUI,TAUE,TAUO,DTMULT,
$ PI,GRAY,KGAS,KAPA,PSL,ED,FMU,NFL,SPF,PRCH,RESDIST,SIND,COSD,
$ RHMAX,CDX,DUMMYC(18),IALTER,DUMMYA(99),
$ XLABEL(20),SIG(20),DSIG(20),SIGE(21),DELGO(19),
$ J1PS(11),JMP3(11),JUS(11),JMS(11),KSES(11),KNBS(11),
$ *LAT(46),DXU(46),DXPI(46),DYU(46),CYPI(46),DAYP(46),F(46),SINL(46),
$ *COSL(46),DUMMY(72),PHIS(288,46),ALBEDO(72,46)
$ *DIMENSION U(72,40,1),V(72,40,1),T(72,40,1),SHI(72,46,1),
$ *P(288,1),TS(288,1),SHS(288,1),GHI(288,1),GWI(288,1),C(300)
$ EQUIVALENCE : U(1,1),PHIS(73,1),TS(1,1),PHIS(72,1,1),
$ (V(1,1),PHIS(793,1)),(SHS(1,1),PHIS(1,1)),(PHIS(1,1)),
$ (SH(1,1,1),PHIS(223,1)),(GT(1,1),PU(72,1,1)),(GW(1,1),
$ PU(144,1)),(C(1),JSP)
COMMON /WORK2/PU(288,46)
INTEGER AMONTH,DAYS,SPM,DAYS,SPY
DIMENSION AMONTH(2,12),DAYS,SPM(12)
DATA AMONTH,'JANU','ARY','FEBR','MAY','MARC','M','APR','
$ 'L','MAY','JUN','JULY',
$ 'ST','SEPT','EMBR','OCTO','BER','NOVE','MBER','DECE',
$ 'MBER',
$ DATA DAYS,SPM/31,28,31,30,31,30,31,31,30,31,30,31,30,31/
DATA IDAYO/302/,DAYS,SPY/365/
DATA SOLS,APHEL,ECRN/173,183,173,183,173,183,173,183/
IFIRST=0
C**** CALCULATE GLOBAL MEAN SURFACE PRESSURE
C ** *
$ SAREA=0
$ SMAX=0
$ INCE=1
$ FINE=INC
DO 20 J=2, JM
SPRESS=0
DO 10 I=INC,IM,INC
$ SPRESS=SPRESS+PI(I,J)
$ SAREA=SAREA+DXYP(I,J)*FIM
$ SMAX=SMAX+SPRESS*DAYP(J)
DO 20 J=2, JM
$ SMAX=PI * SIN(.5*DLAT) ** 2
DAP = 4.0
```


NSTART C I#4 No.R. PHISPL C R#4 No.R. P#1POL No.R. M#KAPA No.R.
 RSDIST S C R#4 No.R. SEASON SFA C R#4 No.R. 000174 No.R. SMPOL No.R. SINLON C R#4 No.R.
 SPRESS SF C R#4 No.R. TOFDAY C R#4 No.R. 000178 No.R. XLABEL No.R. 0005A0 No.R.

***** COMMON INFORMATION *****

NAME OF COMMON BLOCK * F:U#Z OF BLOCKCOMMON 00524 HEXADECIMAL BYTES

VAR.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.
CON1	CON1	R#4	No.R.				CON1	R#4	No.R.				CON1	R#4	No.R.			
CON2	CON2	R#4	No.R.				CON2	R#4	No.R.				CON2	R#4	No.R.			
CON3	CON3	R#4	No.R.				CON3	R#4	No.R.				CON3	R#4	No.R.			
CON4	CON4	R#4	No.R.				CON4	R#4	No.R.				CON4	R#4	No.R.			
CON5	CON5	R#4	No.R.				CON5	R#4	No.R.				CON5	R#4	No.R.			
CPT1	CPT1	R#4	No.R.				CPT1	R#4	No.R.				CPT1	R#4	No.R.			
CP	CP	R#4	No.R.				CP	R#4	No.R.				CP	R#4	No.R.			
IMD2P1	IMD2P1	I#4	No.R.				IMD2P1	I#4	No.R.				IMD2P1	I#4	No.R.			
IMD2P2	IMD2P2	I#4	No.R.				IMD2P2	I#4	No.R.				IMD2P2	I#4	No.R.			
JPOL	JPOL	I#4	No.R.				JPOL	I#4	No.R.				JPOL	I#4	No.R.			
INDEX	INDEX	I#4	No.R.				INDEX	I#4	No.R.				INDEX	I#4	No.R.			
PK	PK	R#4	No.R.				PK	R#4	No.R.				PK	R#4	No.R.			
PHISPL	PHISPL	R#4	No.R.				PHISPL	R#4	No.R.				PHISPL	R#4	No.R.			

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 OPOL 00051A
 TPOL 00045C
 TTPOL 00057C

VARIABLE OFFSETS
 UPOL 000414
 UTPOL 000534
 PPOL 000654

VARIABLE OFFSETS
 VPOL 00043B
 VTPOL 00055B
 VPPOL 00065C

NAME OF COMMON BLOCK * NTR#E OF BLOCKCOMMON 0B3C0 HEXADECIMAL BYTES

VAR.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.
JSP	JSP	I#4	No.R.				JSP	I#4	No.R.				JSP	I#4	No.R.			
PTOP	PTOP	R#4	No.R.				PTOP	R#4	No.R.				PTOP	R#4	No.R.			
FIM	FIM	R#4	No.R.				FIM	R#4	No.R.				FIM	R#4	No.R.			
JN	JN	I#4	No.R.				JN	I#4	No.R.				JN	I#4	No.R.			
MROT	MROT	I#4	No.R.				MROT	I#4	No.R.				MROT	I#4	No.R.			
JAYS	JAYS	I#4	No.R.				JAYS	I#4	No.R.				JAYS	I#4	No.R.			
DLAT	DLAT	I#4	No.R.				DLAT	I#4	No.R.				DLAT	I#4	No.R.			
ITAU	ITAU	I#4	No.R.				ITAU	I#4	No.R.				ITAU	I#4	No.R.			
TOFDAY	TOFDAY	I#4	No.R.				TOFDAY	I#4	No.R.				TOFDAY	I#4	No.R.			
NSTEP	NSTEP	I#4	No.R.				NSTEP	I#4	No.R.				NSTEP	I#4	No.R.			
TAUP	TAUP	I#4	No.R.				TAUP	I#4	No.R.				TAUP	I#4	No.R.			
DTMULT	DTMULT	R#4	No.R.				DTMULT	R#4	No.R.				DTMULT	R#4	No.R.			
KAPA	KAPA	R#4	No.R.				KAPA	R#4	No.R.				KAPA	R#4	No.R.			
MFLW	MFLW	R#4	No.R.				MFLW	R#4	No.R.				MFLW	R#4	No.R.			
SIND	SIND	R#4	No.R.				SIND	R#4	No.R.				SIND	R#4	No.R.			
DUNNYC	DUNNYC	R#4	No.R.				DUNNYC	R#4	No.R.				DUNNYC	R#4	No.R.			
SIG	SIG	R#4	No.R.				SIG	R#4	No.R.				SIG	R#4	No.R.			
JIPS	JIPS	I#4	No.R.				JIPS	I#4	No.R.				JIPS	I#4	No.R.			
KSBS	KSBS	I#4	No.R.				KSBS	I#4	No.R.				KSBS	I#4	No.R.			
DXP	DXP	R#4	No.R.				DXP	R#4	No.R.				DXP	R#4	No.R.			
F	F	R#4	No.R.				F	R#4	No.R.				F	R#4	No.R.			
PHIS	PHIS	R#4	No.R.				PHIS	R#4	No.R.				PHIS	R#4	No.R.			

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 U 000E70
 T 0024F0

VARIABLE OFFSETS
 V 0019B0
 SM 003030

VARIABLE OFFSETS
 SMS 002303
 C 000000

NAME OF COMMON BLOCK * #03#E OF BLOCKCOMMON 0B160 HEXADECIMAL BYTES

VAR.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.	NAME	TYPE	REL.	ADDR.	PE	REL.
PU	PU	R#4	No.R.				PU	R#4	No.R.				PU	R#4	No.R.			

VARIABLE OFFSETS
 Y 0019B0
 SM 003030

VARIABLE OFFSETS
 SMS 002303
 C 000000

VARIABLE OFFSETS
 VAR. NAME TYPE REL. ADDR. PE REL.

VARIABLE OFFSETS

LABEL ADDR
30 000354
70 0006C4

LABEL ADDR
20 000314
60 00049A

LABEL ADDR
10 0002F4
50 000430

LABEL ADDR
5 000298
40 0003FE

DATE 12/12/79-0783.04

868K BYTES OF CORE NOT USED

OPTIONS IN EFFECT NAME= MAIN,OPT=02,LINECMT=SS,SIZE=1000K.

OPTIONS IN EFFECT SOURCE,EBCDIC,NOLIST,NODECK,LOAD,MAP,NOEDIT,NOXREF

STATISTICS SOURCE STATEMENTS = 71 PROGRAM SIZE= 1660

STATISTICS NO DIAGNOSTICS GENERATED

***** END OF COMPILATION *****

LEVEL 19.6-APR 71

OS/360 FORTRAN M AT GLSS

```
COMPILER OPTIONS - NAME= MAIN,OPT=02,LINECMT=SS,SIZE=1000K,  
SUBROUTINE INPUT  
LOGICAL FLAGS  
COMMON/FLD/FLAGS(5)  
COMMON/FOURTH/CON1,CON2,CON3,CON4,CON5,CON6,CON7,CON8,ALPHA,BETA,IMD2P1,  
ADLDP,JMS2,JMS1,CPTH,PSKAPA,RRAD,CP,INC,NSTART,CPD2,  
IND2,RADIN,JEND(2),JPI,JP2,MATSUN,JNE XI(2),MODPK(2,2),  
JPOL(2,2),JPMOD(2,2),SINLON(72),COSLOPK(2,2),  
PSIGN(4),POLES(148),W(72,9,5),PK(72,9,3),DIFF(9),  
CONVPL(9),SDPOL(9,2),PHISPL(2),SUM(2,5)  
EQUIVALENCE (OPOL(1,1),POLES(1,1)),(OTPOL(1,1),POLES(73)),  
(UPOL(1,1),POLES(1,1)),(VPOL(1,1),POLES(10)),(TPOL(1,1),  
POLES(19)),(SHPOL(1,1),POLES(28)),(UTPOL(1,1),POLES(73)),  
(VTPOL(1,1),POLES(62)),(WTPOL(1,1),POLES(91)),(SWTPOL(1,1),  
POLES(100)),(PPOL(1,1),POLES(149)),(PPDL(1,1),POLES(147))  
DIMENS CON PPOL(2),PTPOL(2),UPOL(36,1),VPOL(36,1),TPOL(36,1),  
SHPOL(36,1),OPOL(9,4,2),O(72,10,4,1)  
DIMENS LON PTPOL(2),UTPOL(36,1),VPOL(36,1),TPPOL(36,1),  
SHPOL(36,1),OTPOL(9,4,2),OT(72,10,4,1)  
INTEGER*2 ALBEDO  
REAL*4 KAPA,LAY  
REAL*8 SINLON,COSLON,SUM  
REAL*8 ALONG,TWOPI  
COMMON/INSMH/FSD6,F17,F98  
COMMON/WORK2/PU(2880,46)  
COMMON/SMITH/OSM(72,9,4,5),PSM(72,5),OSMPOL(9,4,2),PSMPOL(2)  
COMMON/SDOT48/SDOT(72,46,8),OMEGA(72,46,5),MMXSTS,RADT(72,46,1,0)  
+ ,COMG  
LOGICAL COMG  
INTEGER*2 RADT,OMEGA  
INTEGER*2 ISDOT(72,46,8)  
EQUIVALENCE (SDOT(1,1,5),ISDOT(1,1,1))  
COMMON /CNTRL/  
JSP,JNP,IM,NLAY,PTOP,ISTART,JSPFI,JNPMI,FIM,NLAYMI,NLAYPI,  
JI,JK,KM,TAUT,(ROT,MROT),JTEST,ITEST,  
NR,JAYS(12),INCS(11),JSB,JNB,OLAT,DLON,  
DT,TAU,ITAU,XINT,TDAY,JDAY,TOFDAY,JDATE,JMONTH(2),JYEAR,NSTEP,  
NCYCLE,NCOMP3,NHOGAN,TAUP,TAUI,TAUE,TAUO,DTMULT,  
PI,GRAY,KGAS,KAPA,PSL,ED,FMU,NFLB,PSF,DRCH,RSDIST,SIND,COSO,  
RHMAL,CDX,DUMATYC(18),ALTER,DUMVAL(99)  
XLABEL(20),SIG(20),OSIG(20),SAGE(21),D(160,19)  
JIPS(11),JMPS(11),JUS(11),JMU(11),KSES(11),KNBR(11)  
*LAT(46),DX(46),DYP(46),DRUT(46),LXP(46),LXVP(46),F(46),SINL(46),  
*COSL(46),DUMNY(72),PHIS(2880,46),ALBEDO(72,46)  
DIMENSION U(72,40,1),V(72,40,1),T(72,40,1),SH(72,40,1),  
*P(2880,1),TS(2880,1),SMS(2880,1),GT(2880,1),GM(2880,1),C(330)  
DIMENSION UT(72,40,1),VT(72,40,1),TT(72,40,1),SHT(72,40,1),  
*PT(2880,1)  
COMMON/S/JMIN,JMAX,JSUM,SMTH(37,17)  
COMMON/MLF/NSQ,MLF(10),MATSX,NSMI,NCMI  
EQUIVALENCE  
{(1,1,1),U(1,1,1),PHIS(73,1)},(TS(1,1),PHIS(72,1)),  
{(1,1,1),PHIS(79,1)},(SMS(1,1),PHIS(1,44,1)),  
{(1,1,1),PHIS(15,1)},(PI(1,1),PHIS(216,1)),  
{(SM(1,1),PHIS(233,1)),(OT(1,1,1),JUT(1,1,1),PU(73,1)),  
{(GM(1,1),PU(72,1)),(VT(1,1,1),PU(793,1)),  
{(W(1,1),PU(144,1)),(TT(1,1,1),PU(15,1)),  
{(PT(1,1),PU(216,1)),(SHT(1,1),PU(233,1)),(C(1,1),JSP)  
DIMENSION CI(300)  
DIMENSION INCP(46),ICU(46)  
REAL*8 RECORD,ANDEND,DPI  
DIMENSION RECORD(10)
```


BAND 850
BAND 860
BAND 870
BAND 880
BAND 890
BAND 900
BAND 910
BAND 920
BAND 930
BAND 940
BAND 950
BAND 960
BAND 970
BAND 980
BAND 990
BAND 500
BAND 510
BAND 520
BAND 530
BAND 540
BAND 550
BAND 560
BAND 570
BAND 580
BAND 590
BAND 600
BAND 610
BAND 620
BAND 630
BAND 640
BAND 650
BAND 660
BAND 670
BAND 680
BAND 690
BAND 700
BAND 710
BAND 720

```
WRITE (6,901) IVER  
READ (5,902) XLABEL  
WRITE (3,903) XLABEL  
COPY INPUTZ NAMELIST CNTC CORE TAPE AND TITLE PAGE  
C****  
20  
WRITE (11,904) RECORD  
WRITE (13,905) RECORD  
IF (RECORD(1).NE.ANDEND) GO TO 20  
REWIND 11  
READ (11,INPUTZ)  
IF (ISTART.EQ.7) WRITE (3,938) ISTART  
IF (ISTART.EQ.7.AND.PAWZE.NE.0.) WRITE (15,538) ISTART  
C****  
READ IN VARIABLE ALBEC  
READ (140):(ALBEC(I),J),J=1,JNP),I=1,IM)  
C****  
READ IN SURFACE GEOPOTENTIAL  
READ (17):(PHIS(I),J),J=1,JNP),I=1,IM)  
READ (46):(TPOG(I),J),J=1,JNP),I=1,IM)  
WRITE (6,1900)(J,PHIS(I),J),I=1,IM),J=1,JNP)  
FORMAT(1X,46(J),I=1,IM),PHIS(I),J=1,JNP)  
1900  
WRITE (6,1941)(J,CALBEC(I),J),I=1,IM),J=1,JNP)  
FORMAT(1X,46(J),I=1,IM),CALBEC(I),J=1,JNP)  
1941  
MMXTS = TAUT + XINT * 01  
IF (PAWZE.EQ.1) PAUSE , MOUNT TAPES  
CALL SSM5CH (5,MSSWS)  
IF (KSSW5.EQ.1) ISTART=8  
KTR=8  
C ** * ON INITIAL START ENDFILE MODEL OUTPUT TAPE  
C  
IF (ISTART.EQ.5.AND.ISTART.LE.7) ENDFILE PTR  
GO TO (30,120,120,40,110,110,40,50,90), ISTART  
C****  
30 STOP  
40 DLON=2.*DPI/IM  
FIM=IM  
KMM=4*NLAY+5  
KMA=KX-4  
IMA=IM*4  
JSPPI=JSP+1  
JNPMI=JNP-1  
JNPPI = JNP + 1  
JMPID2 = JNP + 1 / 2  
JSBPI = JSB + 1  
JNBMI = JNB + 1  
NLAYMI = NLAY + 1  
NLAYPI = NLAY + 1  
C *--- IF (DSIG(I).NE.0.) GO TO 60  
50 DSIG(I) = 0./NLAY  
60 STGE(I) = 0./NLAY  
70 STGE(I+1) = STGE(I) + DSIG(I)  
SIG(I) = 5.*STGE(I) + STGE(I+1)  
80 DSIG(I) = SIG(I+1) - SIG(I)  
IF (ISTART.EQ.4) GO TO 120  
IF (ISTART.EQ.7) GO TO 110  
C**** MACHINE CHECK RESTART ISTART = E OR 9  
90 WRITE (15,906) ISTART  
READ (15,907) TAUP  
TAUP=TAUP-AMOD(TAUP,TAUT)  
100 WRITE (13,908) TAUP  
GO TO 120  
C**** INITIAL CONDITIONS FROM UNIT 12. ISTART= 5, 6, OR 7  
110  
TAUT=1  
TAUT=TAUT  
IF (TAUP.LI.0.) TAUP=TAUT  
TAPNUM=ISK  
IF (PAWZE.EQ.0.) GO TO 120  
WRITE (15,909) ISTART  
1909  
READ (15,902) TAPNUM  
WRITE (13,910) TAPNUM  
READ TAPE ON UNIT KTR  
C  
C 120 READ(KTR,ERR=800,END=130)TAUX,C1 ((PI(I),J),J=1,JNP),I=1,IM),  
S ((OT(I),J),J=1,JNP),I=1,IM), (SAS(I),J),J=1,JNP),I=1,IM),  
5 ((OT(I),J),J=1,JNP),I=1,IM), ((GW(I),J),J=1,JNP),I=1,IM),  
CALL PDIAPLETRI  
IF (ISTART.LE.9) GO TO 120  
IF (TAUX.LE.TAUP*.01) GO TO 120  
IF (TAUX.LE.TAUP*.01) GO TO 120  
121 CONTINUE  
WRITE (6,912) TAUP
```

ISN 0102
ISN 0103
ISN 0104
ISN 0105
ISN 0106
ISN 0107
ISN 0108
ISN 0109
ISN 0110
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ISN 0114
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ISN 0189


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15N 0190
15N 0191
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15N 0193
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15N 0196
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15N 0209
15N 0210
15N 0211
15N 0212
15N 0214
15N 0215
15N 0216
15N 0217
15N 0218
15N 0219
15N 0220
15N 0221
15N 0222
15N 0223
15N 0224
15N 0225
15N 0226
15N 0227
15N 0228
15N 0229
15N 0230
15N 0231
15N 0232
15N 0233
15N 0234
15N 0235
15N 0236

```

```

CALL EXIT
C 130 IF(I,START,GE,SIG)GTO 830
BACKSPACE KTR
140 TAPLTAUX
TAPLTAUX
DO 145 J=1,JM
BACKSPACE KTR
145 BACKSPACE KTR
140 TAPLTAUX
IF(I,START,EG,DR,ISTART,EG,7) GC TO 160
C**** COPY C ARRAY FROM CI ARRAY
DO 150 K=1,300
C(K)=C(CI,K)
150 IF(I,START,EG,2,DR,ISTART,EG,5) GC TO 160
REWIND 11
READ (11,INPUTZ)
IF(I,TR,EG,8) C(16)=C(115)
IF(I,TR,EG,8) ALABEL(20)=C(1220)
IF(I,TR,EG,12) ALABEL(20)=TAPRUM
C**** CALCULATE DISTANCE PROJECTION ARRAYS
150 FJ=0.5*(JSP/JNP)
DO 200 J=1,JNP
LAT(J)=DLAT*(J-FJEO)
DO 210 I=1,JNP
SIN(I,J)=SIN(LAT(I,J))
COS(I,J)=COS(LAT(I,J))
DO 220 J=1,JM
DXP(J)=RAD*DLON*COSL(J)
DO 230 J=2,JM
DYU(J)=RAD * DLAT
DYU(J)=RAD*(LAT(J)-LAT(J-1))
CONTINUE
JMPID2 = (JM + 1)/2
JNP1 = JNP + 1
DO 240 J = 2,JMPID2
DXP(J) = 12. * DXP(J) * DYU(J)
DXP(JNP1-J) = DXYP(J)
DXYP(1) = 0.
DXYP(JNP1) = 3.
WRITE(3,2200)(DXYP(J),J=1,JNP)
FORMAT(1X,*,DXYP=/(1X,SEI6.8))
2200 RETURN
C * * * AUG 23 1977 ADDITIONS HERE TILL RETURN
ENTRY INPUT*

```

```

84M05730
84M05740
84M05750
84M05760
84M05770
84M05780
84M05790
84M05800
84M05810
84M05820
84M05830
84M05840
84M05850
84M05860
84M05870
84M05880
84M05890
84M05900
84M05910
84M05920
84M05930
84M05940
84M05950
84M05960
84M05970
84M05980
84M05990
84M06000
84M06010
84M06020
84M06030
84M06040
84M06050
84M06060
84M06070
84M06080
84M06090
84M06100
84M06110
84M06120
84M06130
84M06140
84M06150
84M06160
84M06170
84M06180
84M06190
84M06200
84M06210
84M06220
84M06230
84M06240
84M06250
84M06260
84M06270
84M06280
84M06290
84M06300
84M06310
84M06320
84M06330
84M06340
84M06350
84M06360
84M06370
84M06380
84M06390
84M06400
84M06410
84M06420
84M06430
84M06440
84M06450
84M06460
84M06470
84M06480
84M06490
84M06500
84M06510
84M06520
84M06530
84M06540
84M06550
84M06560
84M06570
84M06580
84M06590
84M06600

```

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15N 0237
15N 0238
15N 0239
15N 0240
15N 0241
15N 0242
15N 0243
15N 0244
15N 0245
15N 0246
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15N 0258
15N 0259
15N 0260
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15N 0263
15N 0264
15N 0265
15N 0266
15N 0267
15N 0268
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15N 0271
15N 0272
15N 0273
15N 0274
15N 0275
15N 0276

```

```

INC = 1
BETA = 1. - 2. * ALPHA
FSD6 = 15. / 6.
F17 = 17. / 18.
F98 = .325 / 981.
R = RGAS
RADIM = 3 * RAD * IM
CON1 = 4. * COTAN(5*DLAT) / RADIM
CON2 = -COTAN(DLAT) / RADIM
CON3 = 1. / (RADIM * DLAT)
INC(1,1) = 2
JPOL(1,1) = 3
JPOL(1,2) = JM - 1
JPOL(2,2) = 2
JMOD(1,1) = 3
JMOD(2,1) = 3
JMOD(1,2) = MOD(JM-1,5) + 1
JMOD(2,2) = MOD(JM-2,5) + 1
MODPK(1,1) = 2
MODPK(1,2) = 3
MODPK(2,1) = MOD(JM-1,3) + 1
MODPK(2,2) = MOD(JM-2,3) + 1
JNP=JMT
JEND(1) = 1
JEND(2) = JNP
JNEXT(1) = 2
JNEXT(2) = JM
ADLDP = 12. * RAD * DLAT * DLON
IMD2=IM/2
IMD2PI = IMD2 + 1
JMS1 = JM - 1
JMS2 = JM - 2
JMS1=6
JMS2=6
JMIN = 2 + JSMT
JMAX = JM + JSMT
JSUM = JMAX + JMIN
TROJN = 2. * JM
JLTM = JMIN - 1

```

* * *

ISN 0277
ISN 0278
ISN 0279
ISN 0280
ISN 0281
ISN 0282
ISN 0283
ISN 0284
ISN 0285
ISN 0286
ISN 0287
ISN 0288
ISN 0289
ISN 0290
ISN 0291
ISN 0292
ISN 0293
ISN 0294
ISN 0295
ISN 0296
ISN 0297
ISN 0298
ISN 0299
ISN 0300
ISN 0301
ISN 0302
ISN 0303
ISN 0304
ISN 0305
ISN 0306
ISN 0307

DO 301 J = 1, JLM
JDH = J + 1
IF (J.GE.JMIN) JDM = JMAX + 1 - JMIN
IMAX = INT(FACCCOS(LJDM))
IMAXM1 = IMAX - 1
IMAXP3 = IMAX + 4
DO 310 I = 1, IMD2
SMTH(I, J) = 1.
IF (I.GE.IMAX) AND (I.LE.(IMAXP3) SMTH(I, J) = 1 - FLOAT(I - IMAXM1) / 6.
IF (I.GT.IMAXP3) SMTH(I, J) = 0.
CONTINUE
CONTINUE
PIKAPA = KAPA + 1
DO 320 L = 1, NLAY
DIFF(L) = PIKAPA * DSIG(L)
CONTINUE
PSKAPA = EXPBYK(1000.)
THBAR = 280. / PSKAPA
CP = RGAS / KAPA
CPD2 = 5 * CP
CPH = CP * THBAR
WRITE(6, 9778) PSKAPA, THBAR, RGAS, KAPA, CP, CPTH
PHISPL(1) = PHIS(1, 1)
PHISPL(2) = PHIS(1, 1)
IMD2 = IM / 2
IMD1 = 2 * PI
TWOP1 = 6.283185307175567
DO 1000 I = 1, IMD2
ALONG = FLOAT(I-1) * TWOP1 / FLOAT(IM)
C * * * * ALONG = FLOAT(I-1) * TWOP1 / FLCAT(IM) - PI
C * * * *
SINLON(I) = -DSIN(ALONG)
COSLON(I) = -DCOS(ALONG)
SINLON(IMD2+I) = -SINLON(I)
COSLON(IMD2+I) = -COSLON(I)
INDEX(I) = I + IMD2
INDEX(IMD2+I) = I
CONTINUE
OHM2 = 2 * TWOP1 / SDAY
DO 1010 J = 1, JM
F(J) = OHM2 * SINL(J)
F(JNP) = OHM2
PSIGN(1) = -1.
PSIGN(2) = -1.
PSIGN(3) = 1.
PSIGN(4) = 1.
C * * * * INITIAL VELOCITIES GIVEN ON A STAGGERED GRID.
C * * * * INTERPOLATION TO UNSTAGGERED GRID IS USED.
C * * * * MAKE USE OF EQUIVALENCING AND OVERLAPPING THE U ARRAY.
C * * * *
DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
DO 1018 I = 1, IM
IF (I.START.EQ.7) GO TO 1017
Q(I, L, NM2, JEND(M)) = QT(I, L, NM2, JEND(M))
Q(I, L, N, JEND(M)) = QT(I, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
P(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q(I, L, NP2, J) = QT(I, L, NP2, J)
IF (I.START.EQ.7) GO TO 1026
Q(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J+1) +
\$ QT(I, L, N, J-1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
P(I, J) = PT(I, J)
CONTINUE
C * * * * SCALE UT
C * * * *

ISN 0308
ISN 0309
ISN 0310
ISN 0311
ISN 0312
ISN 0313
ISN 0314
ISN 0315
ISN 0316
ISN 0317
ISN 0318
ISN 0319
ISN 0320
ISN 0321
ISN 0322
ISN 0323

DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
DO 1018 I = 1, IM
IF (I.START.EQ.7) GO TO 1017
Q(I, L, NM2, JEND(M)) = QT(I, L, NM2, JEND(M))
Q(I, L, N, JEND(M)) = QT(I, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
P(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q(I, L, NP2, J) = QT(I, L, NP2, J)
IF (I.START.EQ.7) GO TO 1026
Q(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J+1) +
\$ QT(I, L, N, J-1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
P(I, J) = PT(I, J)
CONTINUE
C * * * * SCALE UT
C * * * *

ISN 0324
ISN 0325
ISN 0326
ISN 0327
ISN 0328
ISN 0329
ISN 0330
ISN 0331
ISN 0332
ISN 0333
ISN 0334
ISN 0335
ISN 0336
ISN 0337
ISN 0338
ISN 0339
ISN 0340
ISN 0341
ISN 0342
ISN 0343
ISN 0344
ISN 0345
ISN 0347
ISN 0348
ISN 0349
ISN 0350
ISN 0351
ISN 0352
ISN 0353
ISN 0354

DO 1015 M = 1, 2
DO 1018 N = 3, 4
NM2 = N - 2
DO 1018 L = 1, NLAY
DO 1018 I = 1, IM
IF (I.START.EQ.7) GO TO 1017
Q(I, L, NM2, JEND(M)) = QT(I, L, NM2, JEND(M))
Q(I, L, N, JEND(M)) = QT(I, L, N, JEND(M))
CONTINUE
DO 1015 I = 1, IM
P(I, JEND(M)) = PT(I, JEND(M))
CONTINUE
DO 1020 J = 2, JM
JPI = J + 1
DO 1025 N = 1, 2
NP2 = N + 2
DO 1025 L = 1, NLAY
ISI = IM
DO 1025 I = 1, IM
Q(I, L, NP2, J) = QT(I, L, NP2, J)
IF (I.START.EQ.7) GO TO 1026
Q(I, L, N, J) = 0.25 * (QT(I, L, N, J) + QT(I, L, N, J+1) +
\$ QT(I, L, N, J-1) + QT(I, L, N, J))
ISI = I
GO TO 1025
Q(I, L, N, J) = QT(I, L, N, J)
CONTINUE
DO 1030 I = 1, IM
P(I, J) = PT(I, J)
CONTINUE
C * * * * SCALE UT
C * * * *

BAND 6610
BAND 6620
BAND 6630
BAND 6640
BAND 6650
BAND 6660
BAND 6670
BAND 6680
BAND 6690
BAND 6700
BAND 6710
BAND 6720
BAND 6730
BAND 6740
BAND 6750
BAND 6760
BAND 6770
BAND 6780
BAND 6790
BAND 6800
BAND 6810
BAND 6820
BAND 6830
BAND 6840
BAND 6850
BAND 6860
BAND 6870
BAND 6880
BAND 6890
BAND 6900
BAND 6910
BAND 6920
BAND 6930
BAND 6940
BAND 6950
BAND 6960
BAND 6970
BAND 6980
BAND 6990
BAND 7000
BAND 7010
BAND 7020
BAND 7030
BAND 7040
BAND 7050
BAND 7060
BAND 7070
BAND 7080
BAND 7090
BAND 7100
BAND 7110
BAND 7120
BAND 7130
BAND 7140
BAND 7150
BAND 7160
BAND 7170
BAND 7180
BAND 7190
BAND 7200
BAND 7210
BAND 7220
BAND 7230
BAND 7240
BAND 7250
BAND 7260
BAND 7270
BAND 7280
BAND 7290
BAND 7300
BAND 7310
BAND 7320
BAND 7330
BAND 7340
BAND 7350
BAND 7360
BAND 7370
BAND 7380
BAND 7390
BAND 7400
BAND 7410
BAND 7420
BAND 7430
BAND 7440
BAND 7450
BAND 7460
BAND 7470
BAND 7480

ISN 0355
ISN 0356
ISN 0357
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ISN 0359

ISN 0360
ISN 0361
ISN 0362
ISN 0363
ISN 0364
ISN 0365
ISN 0366
ISN 0367
ISN 0368

ISN 0369
ISN 0370
ISN 0371
ISN 0372
ISN 0373
ISN 0374

ISN 0375

ISN 0376
ISN 0377
ISN 0378
ISN 0379
ISN 0380
ISN 0381
ISN 0382
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ISN 0391
ISN 0392
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ISN 0395
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ISN 0397
ISN 0398

ISN 0399
ISN 0400

ISN 0401
ISN 0402
ISN 0403
ISN 0404

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ISN 0409

ISN 0410
ISN 0411
ISN 0412

ISN 0413
ISN 0414
ISN 0415
ISN 0416
ISN 0417

ISN 0418

ISN 0419

ISN 0420

ISN 0421

ISN 0422

```
DO 1020 N = 1.4  
DO 1020 L = 1.4  
DO 1020 I = 1.4  
OT(I,L,N,J) = DXP(J) * PI(I,J) * Q(I,L,N,J)  
CONTINUE  
C *** * COMPUTATION OF STEREOGRAPHIC POLAR VELOCITIES: UPOL(L,M).  
C *** * VPOL(L,M), WHERE M=1 CORRESPONDS TO SOUTH POLE  
C *** *  
DO 1040 M = 1.2  
COEF1 = (-1) ** M  
DO 1040 L = 1.4  
IF(I,START,NE,7) GO TO 1052  
S1 = 0.  
S2 = 0.  
DO 1050 I = 1.4  
S1 = S1 - U(I,L,JNEXT(M)) * SIN(LN(I)) - COEF1 * V(I,L,JNEXT(M))  
S2 = S2 + COEF1 * U(I,L,JNEXT(M)) * COS(LN(I)) + V(I,L,JNEXT(M)) * SIN(LN(I))  
CONTINUE  
VPOL(L,M) = S1 / FLOAT(I,M)  
GO TO 1053  
1052 UPOL(L,M) = -U(I,L,JEND(M)) * SIN(LN(I)) -  
1 VPOL(L,M) = -U(I,L,JEND(M)) * COEF1 * V(I,L,JEND(M)) * COS(LN(I)) +  
1 VPOL(L,M) = -U(I,L,JEND(M)) * COEF1 * V(I,L,JEND(M)) * SIN(LN(I))  
CONTINUE  
TPOL(L,M) = T(I,L,JEND(M))  
SPOL(L,M) = SH(I,L,JEND(M))  
PPOL(L,M) = PI(I,L,JEND(M))  
WRITE(6,1968)M,L,PPOL(M),(OPOL(L,N),M),N(1.4)  
1968 FORMAT(1,968)M,L,PPOL(M),(OPOL(L,N),M),N(1.4)  
TPOL(M) = PPOL(M)  
DO 1045 N = 1.4  
C *** * THE OT(I,L,JEND(M)) = OPOL(L,N,M) + TPOL(L,M)  
C *** * WITH THE OT(I,L,JEND(M)) COMPUTATION BELOW AGREES  
C *** *  
IF(I,START,NE,7) GO TO 1054  
U(I,L,JEND(M)) = -UPOL(L,M) * SIN(LN(I)) + COEF1 * VPOL(L,M) * COS(LN(I))  
V(I,L,JEND(M)) = -COEF1 * UPOL(L,M) * COS(LN(I)) - VPOL(L,M) * SIN(LN(I))  
CONTINUE  
1054 CONTINUE  
DO 1040 N = 1.4  
OT(I,L,N,JEND(M)) = Q(I,L,N,JEND(M)) + PPOL(M)  
CONTINUE  
C *** * ERROR ENCOUNTERED READING TAUC RECORD ON UNIT KTR  
800 WRITE(7,911)  
CALL EXIT  
C *** * ERROR ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR  
810 WRITE(7,916) J  
CALL EXIT  
C *** * END-OF-FILE ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR  
820 WRITE(7,917) J  
CALL EXIT  
C *** * LATER PICKUP TAPE NEEDED  
630 WRITE(7,913) TAUP  
STOP  
CALL EXIT  
901 FORMAT(7,40X,'GISS N LAYER WEATHER MODEL//100X *VERSION* (4//)  
902 FORMAT(20A4)  
903 FORMAT(25X,20A4)  
904 FORMAT(15A8)  
905 FORMAT(15A8)  
906 FORMAT(15X,10A8)  
* E.C.G. ENTER THE PESTART VALUE OF TAU AS AN 8 CHARACTER NUMBER  
907 FORMAT(426.2 AS 04026*20*)  
908 FORMAT(7F8.2)  
909 FORMAT(7MACHINE CHECK RESTART, TAUP=.F9.2//  
910 *NUMBER, )  
911 FORMAT(7//) THE NUMBER OF THE NEW TAPE IS (4//)  
912 FORMAT(7ERROR ENCOUNTERED READING TAUC RECORD ON UNIT KTR. )  
913 FORMAT(7EARLIER PICKUP TAPE NEEDED. TAUP = .F10.2)  
914 FORMAT(7LATER PICKUP TAPE NEEDED. TAUP = .F10.2)  
* IROTE (7,22)  
915 FORMAT(7C,12, REGION SPLIT GRID LIMITS JIP,1115/31X,20A4,3X,  
* 1115/31X,JIU,1115/31X,JMU,1115/31X,INC,1115/31X,KSB,.  
916 FORMAT(7ERROR ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR. )  
917 BEING READ WAS (4)  
*KTR, (7END-OF-FILE ENCOUNTERED READING LATITUDE RECORD ON UNIT KTR.  
918 FORMAT(7BEING READ WAS (12, ) THE DECK IS SET UP AS AN INITIAL START,  
* /, IF IT SHOULD BE A RESTART DC JOB NO. AND INSERT AN ISTART=3  
*ARD BEFORE SEND AND RELOAD. THANK YOU!  
END
```


IMD2 144 000050 RADIN R44 000054 JEND 144 000058
 JPD2 144 N4R 000069 MATSUN R44 000069 JPI 144 000069
 JPOL 144 000084 JPMOD 144 000094 JNEXT 144 000094
 INDEX 144 000224 PSIGN R44 000404 SINLN R44 000404
 PK 144 N4R 000404 DIFR R44 000404 POLES R44 000414
 PHISPL R44 000574 SUM R44 000574 CONVPL R44 000574
 SDPOL R44 N4R 000574

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 OPOL 000414
 TPOL 00045C
 TTPOL 00057C
 SHTPOL 0005A0

NAME OF COMMON BLOCK * 1156, ZE OF BLOCK COMMON 00060C HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 F5D6 R44 000000 F17 R44 000004 F98 R44 000008

NAME OF COMMON BLOCK * 0085E OF BLOCK COMMON 081600 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 PU R44 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 OT 000120
 GW 001680
 UT 00120
 TT 0017A0

NAME OF COMMON BLOCK * 00018 OF BLOCK COMMON 00018 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 QSH R44 000000 PSN R44 000000 GMPOL R44 000000

NAME OF COMMON BLOCK * 00019 OF BLOCK COMMON 0385AE HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 SDOT R44 000000 OMEGA I42 000000 MNMXTS I44 0286E0
 CONG L44 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 ISDOT 00CF00

NAME OF COMMON BLOCK * 00019 OF BLOCK COMMON 083E30 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JSD R44 000000 JAP R44 000004
 PTD R44 000010 ISTART I44 000014
 FJM R44 000020 NLAYMI I44 000024
 MROT I44 000030 JTEST I44 000034
 JAVE I44 000040 IPCS I44 000044
 DLAT I44 000050 DLON R44 000058
 TOTAU R44 N4R 000068
 INSTP R44 N4R 000068
 S44 R44 000078
 DTMUT R44 000108
 KAPA R44 000118
 NPL9 R44 000128
 SIND R44 N4R 000138
 DUMNYC R44 N4R 000148
 S4S R44 000150
 JIBS R44 N4R 000158
 KSB5 R44 N4R 000168
 D4E R44 000178
 PHIS R44 000050 ALBEDO I42 082350

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETS
 Q 000E70
 SHS 002300
 C 000000

NAME OF COMMON BLOCK * 00050 OF BLOCK COMMON 000500 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JMIN I44 000000 JMAX I44 000004 JSUN I44 000008
 TS 000F10
 P 002F10

NAME OF COMMON BLOCK * 00050 OF BLOCK COMMON 000500 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 JMIN I44 000000 JMAX I44 000004 JSUN I44 000008
 SM 000300
 SH 000300

VARIABLE OFFSETS
 VPOL 000238
 WPOL 000550
 PTPOL 00065C

VARIABLE OFFSETS
 VI 000C40
 SMI 0022E0

VARIABLE OFFSETS

VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
 NLAY I44 00000C
 JNPHI I44 00001C
 JI I44 00002C
 IRGI I44 00003C
 NR I44 00004C
 JMD I44 000050
 FAU R44 000050
 JDAY I44 N4R
 JYER I44 N4R
 NHOGAN I44 0000F4
 TAUO R44 000104
 RGAS R44 000114
 FPU R44 000124
 RSDIST R44 N4R
 CDX R44 00144
 LABEL R44 000320
 DSIGG R44 000350
 JMSU I44 N4R
 DXU R44 N4R
 DXPV R44 000950
 DUMNY R44 N4R

VARIABLE OFFSETS
 V 001980
 SH 000300

VARIABLE OFFSETS
 TS 000F10
 P 002F10

VARIABLE OFFSETS
 V 001980
 SH 000300

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 000038 HEXADECIMAL BYTES
NS1 184 000000 VAR. NAME TYPE REL. ADDR. HE
NS2 184 000034 MFLF 184 000004 VAR. NAME TYPE REL. ADDR. HE
NS3 184 000034 MATSNX 184
VAR. NAME TYPE REL. ADDR. HE
NS4 184 000034 NSM1 184 000030

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 00001C HEXADECIMAL BYTES
CMRT 184 000000 VAR. NAME TYPE REL. ADDR. HE
TMP 184 000010 MUC3 184 000004 VAR. NAME TYPE REL. ADDR. HE
SMAFTA 184 000014 SHKOD 184 000008
ADPCLD 184 000018

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 000004 HEXADECIMAL BYTES
ISWCH 184 000000 VAR. NAME TYPE REL. ADDR. HE
VAR. NAME TYPE REL. ADDR. HE

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 000008 HEXADECIMAL BYTES
ITC 184 000000 VAR. NAME TYPE REL. ADDR. HE
G5W 184 000004 VAR. NAME TYPE REL. ADDR. HE
VAR. NAME TYPE REL. ADDR. HE

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 000008 HEXADECIMAL BYTES
ISWTH 184 000000 VAR. NAME TYPE REL. ADDR. HE
NPC 184
VAR. NAME TYPE REL. ADDR. HE
VAR. NAME TYPE REL. ADDR. HE

NAME OF COMMON BLOCK * * * * *
VAR. NAME TYPE REL. ADDR. HE 000000 HEXADECIMAL BYTES
ISWTH 184 000000 VAR. NAME TYPE REL. ADDR. HE
VAR. NAME TYPE REL. ADDR. HE
VAR. NAME TYPE REL. ADDR. HE

```

LEVEL 19.6-APR 71

```

```

*OPTIONS IN EFFECT*
*OPTIONS IN EFFECT*
*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****

```

```

NAME= MAIN,OPT=02,ALINECNT=55,SIZE=1000K.
SOURCE:EBDC(C),NOLIST,NODECK,LOAD,MAP,NOEDIT,NOXREF
SOURCE STATEMENTS = 421 PROGRAM SIZE= 12544

```

```

*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****

```

```

772K BYTES OF CORE NOT USED
DATE 12/12/79-0783.20

```

```

05/360 FORTRAN H AT GISS

```

```

COMPILER OPTIONS -
SUBROUTINE TWRITE
INTEGER*2 ALBEDO
REAL*8 KAPA,LAT
COMMON/CONTR/
* JSP,JNP,M,NLAY,PIOP,I,START,JSCPI,JNPM1,FIM,NLAYM,NLAY*2,
* NR,JAYS,I2),INCS(1),JTB,JNB,DIAT,DION,
* DT,TAU,ITAU,XINT,IDAY,JDAY,TOEGAY,JDATE,JMONTH(2),JYEAR,NSTEP,
* NCYCLE,NCOMP,KAPA,PSI,ED,FMU,NFLB,PSF,PRCH,RSDIST,SIND,COSD,
* PI,GRAY,RCAS,KAPA,PSI,ALTER,DUMMYA(99),
* RHMAX,CDX,DUMMYC(18),SIG(20),SIGE(21),D,SIGD(19),
* XLABEL(1),JMPS(1),JBUS(1),JMUS(1),KMS(1),KNBS(1),
* *LATT(4),DUMY(4),DXP(46),DYU(46),DYP(46),DYP(46),F(46),SINL(46),
* *COSL(46),DUMMY(72),PHIS(280),ALBEDO(12,46)
* DIMENSION U(172,40,1),V(172,40,1),T(172,40,1),SM(72,40,1),
* *P(280,1),TS(280,1),S(280,1),G(1280,1),GM(280,1),C(300)
* COMMON/WORK/PUI(280,46)
+ COMG
LOGICAL COMG
INTEGER*2 RADT,OMEGA
INTEGER*2 ISDOT(72,46,8)
EQUIVALENCE (SDUT(1,1,5), (SDOT(1,1,1)))
DIMENSION SDFIL(172,12),RADFIL(72,16),OMF(172,9)
EQUIVALENCE
( U(1,1,1), PHIS(73,1) ), ( TS(1,1,1), PHS(73,1) ),
( V(1,1,1), PHIS(73,1) ), ( S(1,1,1), PHS(1,1,1) ),
( P(1,1,1), PHIS(23,1) ), ( P(1,1,1), PHIS(261,1) ),
( G(1,1,1), PUI(23,1) ), ( G(1,1,1), FU(1,41,1) ), ( C(1,1,1), JSP )
COMMON/OMG/TS(1),PHIS(73),PHIS(1,1),PHIS(1,1)
EQUIVALENCE(C(100),MACHIN)
MACHIN = 105
IMPI = 14 + 1
IM2P = 2 * IM
IM3P = 5 * IM
IM3P = IM3 + 1
IM4P = IM4 + 4
IM5P = IM5 + 4
TAUR = INT(TAU + .5)
BACKSPACE 8

```

```

*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****

```

```

772K BYTES OF CORE NOT USED
DATE 12/12/79-0783.20

```

```

05/360 FORTRAN H AT GISS

```

```

COMPILER OPTIONS -
SUBROUTINE TWRITE
INTEGER*2 ALBEDO
REAL*8 KAPA,LAT
COMMON/CONTR/
* JSP,JNP,M,NLAY,PIOP,I,START,JSCPI,JNPM1,FIM,NLAYM,NLAY*2,
* NR,JAYS,I2),INCS(1),JTB,JNB,DIAT,DION,
* DT,TAU,ITAU,XINT,IDAY,JDAY,TOEGAY,JDATE,JMONTH(2),JYEAR,NSTEP,
* NCYCLE,NCOMP,KAPA,PSI,ED,FMU,NFLB,PSF,PRCH,RSDIST,SIND,COSD,
* PI,GRAY,RCAS,KAPA,PSI,ALTER,DUMMYA(99),
* RHMAX,CDX,DUMMYC(18),SIG(20),SIGE(21),D,SIGD(19),
* XLABEL(1),JMPS(1),JBUS(1),JMUS(1),KMS(1),KNBS(1),
* *LATT(4),DUMY(4),DXP(46),DYU(46),DYP(46),DYP(46),F(46),SINL(46),
* *COSL(46),DUMMY(72),PHIS(280),ALBEDO(12,46)
* DIMENSION U(172,40,1),V(172,40,1),T(172,40,1),SM(72,40,1),
* *P(280,1),TS(280,1),S(280,1),G(1280,1),GM(280,1),C(300)
* COMMON/WORK/PUI(280,46)
+ COMG
LOGICAL COMG
INTEGER*2 RADT,OMEGA
INTEGER*2 ISDOT(72,46,8)
EQUIVALENCE (SDUT(1,1,5), (SDOT(1,1,1)))
DIMENSION SDFIL(172,12),RADFIL(72,16),OMF(172,9)
EQUIVALENCE
( U(1,1,1), PHIS(73,1) ), ( TS(1,1,1), PHS(73,1) ),
( V(1,1,1), PHIS(73,1) ), ( S(1,1,1), PHS(1,1,1) ),
( P(1,1,1), PHIS(23,1) ), ( P(1,1,1), PHIS(261,1) ),
( G(1,1,1), PUI(23,1) ), ( G(1,1,1), FU(1,41,1) ), ( C(1,1,1), JSP )
COMMON/OMG/TS(1),PHIS(73),PHIS(1,1),PHIS(1,1)
EQUIVALENCE(C(100),MACHIN)
MACHIN = 105
IMPI = 14 + 1
IM2P = 2 * IM
IM3P = 5 * IM
IM3P = IM3 + 1
IM4P = IM4 + 4
IM5P = IM5 + 4
TAUR = INT(TAU + .5)
BACKSPACE 8

```

```

772K BYTES OF CORE NOT USED
DATE 12/12/79-0783.20

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```

05/360 FORTRAN H AT GISS

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```

*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****

```

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772K BYTES OF CORE NOT USED
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05/360 FORTRAN H AT GISS

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*STATISTICS* NO DIAGNOSTICS GENERATED
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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

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05/360 FORTRAN H AT GISS

```

```

*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****

```

```

772K BYTES OF CORE NOT USED
DATE 12/12/79-0783.20

```

```

05/360 FORTRAN H AT GISS

```

```

70 WRITE (8,END=80,ERR=90)TAUR,C, (2)(1,J),J=1,JNP,I=1,IM,
$ ((TS(I,J),J=1,JNP),I=1,IM), (PHS(I,J),J=1,JNP),I=1,IM),
$ ((G(I,J),J=1,JNP),I=1,IM), (GM(I,J),J=1,JNP),I=1,IM)

```


JM	104	000000	NR	000030	TAUT	104	000030	104	000030	IRDT	104	000030
MROT	104	NR	NR	NR	ITEST	104	NR	104	NR	NR	104	NR
JAYS	104	NR	NR	NR	JSB	104	NR	104	NR	JNB	104	NR
DLAT	104	NR	NR	NR	DT	104	NR	104	NR	JAU	104	NR
ITAU	104	NR	NR	NR	DAY	104	NR	104	NR	JDAY	104	NR
TOFDAY	104	NR	NR	NR	MONTH	104	NR	104	NR	JTEAR	104	NR
NSTEP	104	NR	NR	NR	COMP3	104	NR	104	NR	MHGAN	104	NR
TAUP	104	NR	NR	NR	TAU	104	NR	104	NR	TAUO	104	NR
DTMULT	104	NR	NR	NR	GRAV	104	NR	104	NR	RGAS	104	NR
KAPA	104	NR	NR	NR	ED	104	NR	104	NR	FRU	104	NR
NFLN	104	NR	NR	NR	MRCB	104	NR	104	NR	ASD1ST	104	NR
SIND	104	NR	NR	NR	RHMYA	104	NR	104	NR	CDX	104	NR
DUMMYC	104	NR	NR	NR	DUMNYA	104	NR	104	NR	XLABEL	104	NR
SIG	104	NR	NR	NR	SIGE	104	NR	104	NR	DSIGD	104	NR
JIPS	104	NR	NR	NR	JIUS	104	NR	104	NR	JMUS	104	NR
KSBS	104	NR	NR	NR	LAT	104	NR	104	NR	DXU	104	NR
DXP	104	NR	NR	NR	DYP	104	NR	104	NR	DXYP	104	NR
F	104	NR	NR	NR	CSL	104	NR	104	NR	DUMNY	104	NR
PHS	104	000050	NR	NR	ALBDDO	104	NR	104	NR			

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 001890
 U 0024F0
 T 00018C
 MACHIN

NAME OF COMMON BLOCK * 000000 OF BLOCK COMMON 001400 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 PU 104 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 00168C
 GT 000B8C

NAME OF COMMON BLOCK * 00019E OF BLOCK COMMON 0385AE HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
 SOOT 104 000200
 COMG 104 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 00CF00
 ISDCT

NAME OF COMMON BLOCK * 00000B OF BLOCK COMMON 00000B HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE

VARIABLE OFFSETE
 SMS 0023D0
 C 900003

VARIABLE OFFSETE

VARIABLE OFFSETE

VARIABLE OFFSETE


```

C *** IF(LEAP.AND.NEXT STEP IS LEAP)
C *** IF(MATSUN+MATSX.EQ.0)GO TC 45
C *** IF(NEXT CALL IS TO MATS PRED.)
C *** IF((NPC.EQ.1).AND.(MATSX.EQ.1))GO TO 58
C *** DO 28 I = INC.IM,INC
C *** P(I.JS2) = PSM:I.JS2MOD)
C *** CONTINUE
C *** DO 30 N = 1,4
C *** DO 30 L = 1,NLAY
C *** DO 30 I = INC.IM,INC
C *** Q(I.L.N.JS2) = QSM:I.L.N.JS2MOD)
C *** CONTINUE
C *** DO 61 I = INC.IM,INC
C *** P(I.JS2) = PT(I.JS2)
C *** GO TC 67
C *** DO 62 N = 1,4
C *** DO 62 L = 1,NLAY
C *** DO 62 I = INC.IM,INC
C *** Q(I.L.N.JS2) = QT(I.L.N.JS2) * (DXYP(JS2) * PT(I.JS2))
C *** CONTINUE
C *** GO TC 64
C *** DO 45 DO 48 I = INC.IM,INC
C *** P(I.JS2) = BETA * P(I.JS2) + ALPHA * (PSM(I.JS2MOD) + PT(I.JS2))
C *** CONTINUE
C *** DO 48 DO 50 N = 1,4
C *** DO 50 L = 1,NLAY
C *** DO 50 I = INC.IM,INC
C *** Q(I.L.N.JS2) = BETA * DXYP(JS2) * P(I.JS2) * Q(I.L.N.JS2)
C *** + ALPHA*(QSM(I.L.N.JS2MOD)+QT(I.L.N.JS2)*DXYP(JS2)*PT(I.JS2))
C *** CONTINUE
C *** DO 63 CONTINUE TERM CORRECTION DUE TO LEAPFROG TIME SCHEME
C *** DO 63 * * * SOURCE
C *** IF(MOD(NSTEP-NCM1,NCOMP3).NE.0)GC TO 67
C *** DO 57 N = 1,4
C *** DO 57 L = 1,NLAY
C *** DO 57 I = INC.IM,INC
C *** Q(I.L.N.JS2) = Q(I.L.N.JS2) - QT(I.L.N.JS2) *
C *** IDXP(JS2)*PT(I.JS2)
C *** CONTINUE
C *** DO 57 I = INC.IM,INC
C *** GO TC TO 67
C *** IF(MOD(NSTEP-NCM1,NCOMP3).NE.0)GC TO 67
C *** IF(MATSUN.PRED. STEP)
C *** IF((MATSUN.EQ.1).AND.(NPC.EQ.0))GO TO 67
C *** CALL COMP3:UT.VT.T.SHT.PT.UT.FOL.VT.POL.HTPOL.SHTPOL.PT.POL.
C *** $ OTPOL.JS2)
C *** IF(MATSUN.DR NOT COMP3 STEP)
C *** IF((MATSX.NE.0).OR.(MOD(NSTEP-NCM1,NCOMP3).NE.0))GO TO 67
C *** DO 65 N = 1,4
C *** DO 65 L = 1,NLAY
C *** DO 65 I = INC.IM,INC
C *** Q(I.L.N.JS2) = Q(I.L.N.JS2) + DXYP(JS2) * PT(I.JS2) *
C *** $ CONTINUE
C *** CONTINUE
C *** IF (J .LT. JM)
C *** JS2 = JS2 + 1
C *** JS2MOD = MOD(JS2-1,5) + 1
C *** GO TO 29
C *** IF (JS2 .LE. JM)
C *** * * * POLES
C *** DO 70 M = 1
C *** IF (J .EQ. JM)
C *** COEF1 = (-1.) ** M
C *** IF(LEAP.AND.NEXT STEP IS LEAP)
C *** IF(MATSUN+MATSX.EQ.0)GO TO 85
C *** IF(MATS.PRED.OR(MATS CORR. AND NEXT STEP IS LEAP))
C *** IF((MATSUN.EQ.1).AND.(NPC.MATSX.EQ.0))GC TO 78
C *** PPOL(M) = PTPOL(M)
C *** DO 75 N = 1,4
C *** DO 75 L = 1,NLAY
C *** OPOL(L,N,M) = OTPOL(L,N,M)
C *** CONTINUE
C *** GO TC 54
C *** DO 78 PPOL(M) = PSMOL(M)
C *** DO 80 N = 1,4
C *** DO 80 L = 1,NLAY
C *** OPOL(L,N,M) = OSMOL(L,N,M)
C *** CONTINUE
C *** GO TC 54

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ISN 0120
ISN 0122
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84M 0870
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84M 0950
84M 0960
84M 0970
84M 0980
84M 0990
84M 1000
84M 1010
84M 1020
84M 1030
84M 1040
84M 1050
84M 1060
84M 1070
84M 1080
84M 1090
84M 1100
84M 1110
84M 1120
84M 1130
84M 1140
84M 1150
84M 1160
84M 1170
84M 1180
84M 1190
84M 1200
84M 1210
84M 1220
84M 1230
84M 1240
84M 1250
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84M 1270
84M 1280
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84M 1360
84M 1370
84M 1380
84M 1390
84M 1400
84M 1410
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84M 1460
84M 1470
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84M 1490
84M 1500
84M 1510
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84M 1570
84M 1580
84M 1590
84M 1600
84M 1610
84M 1620
84M 1630
84M 1640
84M 1650
84M 1660
84M 1670
84M 1680
84M 1690
84M 1700
84M 1710
84M 1720
84M 1730
84M 1740

```

ISN 0200      PPOL(M) = BETA * PPOL(M) + ALPHA * (PSNPOL(M) + PTPOL(M))
ISN 0201      DO 87 N = 1,4
ISN 0202      DD 87 L = 1,NLAY
ISN 0203      OPOL(L,N,M) = BETAPPOL(M) * OPOL(L,N,M) + AL * PA
ISN 0204      $ * (OSMPOL(L,N,M) + QTPOL(L,N,M))
C * * * SOURCE TERM CORRECTION DUE TO LEAPFROG TIME SCHEME
C * * *
C *** 12/12/77
ISN 0205      93 CONTINUE
ISN 0206      IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 94
ISN 0207      DD 92 N = 1,4
ISN 0208      DD 92 L = 1,NLAY
ISN 0209      OPOL(L,N,M) = OPOL(L,N,M) - QTPOL(L,N,M)
ISN 0210      92 * * * UNSCALE
C * * *
ISN 0211      94 CONTINUE
ISN 0212      DD 96 N = 1,4
ISN 0213      DD 96 L = 1,NLAY
ISN 0214      QPOL(L,N,M) = QTPOL(L,N,M) / PTFOL(M)
ISN 0215      IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 98
ISN 0216      96 IF (MOD(INSTEP - NCM1, NCOMP3), NE, 0) GC TO 98
ISN 0217      IF (MATSUNO PRED, STEP)
ISN 0218      CALL COMP3(UT, VT, TT, SHT, PT, DT, UTPOL, VTPOL, TTPOL, SHTPOL, PTPOL)
ISN 0219      $ QTPOL(J, S2)
ISN 0220      IF (MATSUNO, NE, 0) GO TO 58
ISN 0221      DD 195 N = 1,4
ISN 0222      DD 195 L = 1,NLAY
ISN 0223      QPOL(L,N,M) = QPOL(L,N,M) + PTPOL(M) * QTPOL(L,N,M)
ISN 0224      10E * * * COPY QTPOL INTO QT
C * * *
ISN 0225      98 DO 100 I = INC, IM, INC
ISN 0226      PT(I, JEND(M)) = PTPOL(M)
ISN 0227      100 DO 110 N = 3,4
ISN 0228      DO 110 L = 1,NLAY
ISN 0229      DO 110 I = INC, IM, INC
ISN 0230      QT(I, L, N, JEND(M)) = QTPOL(L, N, M)
ISN 0231      110 DO 120 L = 1,NLAY
ISN 0232      DO 120 I = INC, IM, INC
ISN 0233      UT(I, L, JEND(M)) = UTPOL(L, N, M) * SIN(PI * I) + COS(PI * I) * COS(LON(I))
ISN 0234      VT(I, L, JEND(M)) = -COS(PI * I) * UTPOL(L, N, M) * SIN(LON(I))
ISN 0235      120 JS2 = JS1
ISN 0236      JS1 = J
ISN 0237      10 FORMAT( '0 PRESSURE DIAGNOSTIC. (J,PT,PE=.2I4.2E15.5)
ISN 0238      302 FORMAT(5X, 2E14.4)
ISN 0239      RETURN
ISN 0240      END
ISN 0241
ISN 0242
ISN 0243
ISN 0244
ISN 0245
ISN 0246
ISN 0247

```

NAME TAG TYPE ADD	COMP	SIZE OF PROGRAM	HEXADECIMAL BYTES
000000			
000100			
000200			
000300			
000400			
000500			
000600			
000700			
000800			
000900			
001000			
001100			
001200			
001300			
001400			
001500			
001600			
001700			
001800			
001900			
002000			
002100			
002200			
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004200			
004300			
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006800			
006900			
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007200			
007300			
007400			
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007700			
007800			
007900			
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008100			
008200			
008300			
008400			
008500			
008600			
008700			
008800			
008900			
009000			
009100			
009200			
009300			
009400			
009500			
009600			
009700			
009800			
009900			

NAME OF COMMON BLOCK * SROT#9E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
SROT R#4 000003 OMEGA I#2 N#R. MNMXTS I#4 N#R. RADT I#2 N#R.
CORG L#4 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 00CF00 VARIABLE OFFSETE

NAME OF COMMON BLOCK * #56#7E OF BLOCKCOMMON 000006 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
ISMTH I#4 NPC I#4 000004

NAME OF COMMON BLOCK * #M3#E OF BLOCKCOMMON 000038 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
NSE0 I#4 000034 MLP I#4 N#R. MATSNX I#4 00002C NSM1 I#4 N#R.

NAME OF COMMON BLOCK * # T-7E OF BLOCKCOMMON 0035C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
TOPOG R#4 N#R.

NAME OF COMMON BLOCK * #W6R#E OF BLOCKCOMMON 0101C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
FD R#4 N#R.

NAME OF COMMON BLOCK * #0#5#E OF BLOCKCOMMON 081400 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E VAR. NAME TYPE REL. ADDR.#E
PU R#4 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE VARIABLE OFFSETE

LABEL	ADDR	ADDR
15	000B1E	000B1E
16	000D4A	000E2A
17	0011CA	0017D4
18	00144B	001750
19	0018FA	0019EE
20	001C12	001C6E
21	001F2E	001F72
22	0020D4	002142
23	002278	0022F0
24	0024B2	00254C
25	00278C	0027CC

LABEL	ADDR	ADDR
26	000B66	000BDE
27	000E2A	000E84
28	0017D4	00155E
29	001750	0018A4
30	0019EE	001A3E
31	001C6E	001D82
32	001F72	002098
33	002142	0022A4
34	0022F0	0023A2
35	00254C	00263E
36	0027CC	

```

*OPTIONS IN EFFECT* NAME= MAIN,OPT=02,L INECLT=55,SIZE=100K.
*OPTIONS IN EFFECT* SOURCE=EBCOIC,NOLIST,NODECK,LCAD,MAP,NCECREDIT,LD,NOXREF
*STATISTICS* SOURCE STATEMENTS = 246 PROGRAM SIZE= 11120
*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILE *****
LEVEL 19.6-APR 71
  
```

```

COMPILER OPTIONS - NAME= MAIN,OPT=02,L INECLT=55,SIZE=100K.
SUBROUTINE CO.SPI:UAV,T,SH,P,C,UPCL,VPCL,JPCL,SM,PPOL,PPOL,QPOL,
UT,V7,TT,SH,PT,QT,UTPOL,VTPOL,TPOL,STPOL,STPOL,PTPOL,PTPOL,PTPOL,
COMMON/FGURTH/CONI,CON2,CON3,CON4,CONST,PBAR,ALPHA,BETA,IMDZPI,
ADLDP,JMS2,JMS1,CPTH,PSKAPA,R,RAD,CP,INC,NSTART,CPD2,
IND2,RADIM,JEND(2),JPI,JP2,MATSUN,JNEXT(2),MODPK(2,2),
POLIGN(4),POLLS(148),W(72,9,5),PK(72,9,3),DIFF(72),
CONVPL(9),SDPOL(9,2),PHISPL(2),SUM(2,5),
DIMENSION PPOL(2),PIPOL(2),UPOL(36,1),VPOL(36,1),TPOL(36,1),
SHPOL(36,1),QPOL(9,4,2),0(72,10,4,1),
DIMENSION PIPOL(2),VTPOL(36,1),VTPOL(36,1),TPOL(36,1),
SHTPOL(36,1),QPOL(9,4,2),GT(72,10,4,1)
COMMON/SMITH/OSM(72,9,4,5),PSMI(72,5),OSMFOL(9,4,2),PSMPOL(2)
INTEGER*2 ALBEDO
REAL*4 KAPA,LAT
COMMON /CNTRL/,
  
```

```

* JSP, JNP, IM, NLAY, PTOPI, ISTART, JSPPI, JNPM, FIM, NLAYMI, NLAYPI,
* JI, JM, KM, TAUT, IROT, MROT, JTEST,
* NR, JAYS(12), INCS(11), JDB, JNB, DLAT, DLON,
* DT, JAU, ITAU, XINT, IDAY, JDAY, TOFCAY, JDATE, JMONTH(2), JYEAR, NCTEP,
* NCVCL, NCOMP3, MDGAN, I'AUP, TAU, I'FAUE, I'ALO, DTMULT,
* PI, GRAY, RGAS, KAPA, PSL, ED, FNU, NPL, W, P, PS, MRCH, RSDIST, SIND, COSD,
* RHMAX, CDX, DUMMYC(18), I'ALTER, DU, NYA(99),
* XLABEL(20), SIG(20), OSIG(20), I'JUS(11), JMS(11), PK(5(11), D, SIG(19),
* JIPS(11), JMS(11), JUS(11), JMS(11), PK(5(11), D, SIG(19),
* LAT(46), DXU(46), DXP(46), DYU(46), DYP(46), CXP(46), F(46), SINL(46),
* COSL(46), DUMMY(72), PHIS(2680,46), ALBEDO(72,46)
* DIMENSION U(72,40,1), V(72,40,1), W(72,40,1), SM(72,40,1),
* P(2880,1), TS(2880,1), SHS(2880,1), GT(2680,1), G(2880,1), C(300)
EQUIVALENCE (TS(1,1), PHIS(72,1)), (G(1,1), SMI(1,1)), (C(1,1), JSP)
( GT(1,1), PU(72,1)
COMMON/COMP/2/PI,TPOL
COMMON/WORK/ CONV(72,5,46)
COMMON/WORK2/PU(2880,46)
DIMENSION UT(72,40,1), VT(72,40,1), TT(72,40,1), SH(72,40,1),
*PT(2880,1), SD(72,9,1), PIT(640,1), PV(72,2)
EQUIVALENCE (SD(1,1,1), CONV(1,1,1)), (CONV(1,9,1),
* PIT(1,1))
COMMON /DEBUG/ISNTCH
C
COMMON/SDOT48/SDOT(72,46,8), OMEGA(72,46,5), MNXTS, RADT(72,46,10)
+ ,COMG
LOGICAL COMG
C
INTEGER*2 RADT, OMEGA
INTEGER*2 ISDOT(72,46,8)
C
EQUIVALENCE (SDOT(1,1,5), ISDOT(1,1,1))
C
COMMON/SMO00/ISMTH,NPC
QSAT(1M,PR) = .622 * EXP(21.65604 - 5417.983 / TM) / PR
DATA J272/
JMS2 = JM - 2
FICO = 0 * DT
F2CO = -.5 * DT
  
```

DATE 12/12/79-0783.48

780K BYTES OF CORE NOT USED

LABEL	ADDR
20	000C6B
202	000F5C
29	00156A
58	0018A4
48	001AC2
64	001D82
75	002098
87	0022A4
96	0023A2
110	00263E

LABEL	ADDR
BAMI 2320	
BAMI 2330	
BAMI 2340	
BAMI 2350	
BAMI 2360	
BAMI 2370	
BAMI 2380	
BAMI 2390	
BAMI 2400	
BAMI 2410	
BAMI 2420	
BAMI 2430	
BAMI 2440	
BAMI 2450	
BAMI 2460	
BAMI 2470	
BAMI 2480	
BAMI 2490	
BAMI 2500	
BAMI 2510	
BAMI 2520	
BAMI 2530	
BAMI 2540	
BAMI 2550	
BAMI 2560	
BAMI 2570	
BAMI 2580	
BAMI 2590	
BAMI 2600	
BAMI 2610	
BAMI 2620	
BAMI 2630	
BAMI 2640	
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BAMI 2670	
BAMI 2680	
BAMI 2690	
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BAMI 2740	
BAMI 2750	
BAMI 2760	
BAMI 2770	
BAMI 2780	
BAMI 2790	
BAMI 2800	
BAMI 2810	
BAMI 2820	
BAMI 2830	
BAMI 2840	
BAMI 2850	
BAMI 2860	
BAMI 2870	

ORIGINAL PAGE IS
OF POOR QUALITY

```

ISN 0029      L = J + 2
ISN 0030      JP2 = J + 1
ISN 0031      JS1 = J + 1
ISN 0032      JS1 = J - 1
ISN 0033      JS2 = J - 2
ISN 0034      JP2MOD = MOD(JP2-1,S) + 1
ISN 0035      JPMOD = MOD(JP1-1,S) + 1
ISN 0036      JMOD = MOD(J-1,S) + 1
ISN 0037      JS1MOD = MOD(JS1-1,S) + 1
ISN 0038      JS2MOD = MOD(JS2-1,S) + 1
ISN 0039      IF (J .GT. 2)
ISN 0040      JS1 = 1
ISN 0041      JS2 = 1
ISN 0042      JS1MOD = 1
ISN 0043      JS2MOD = 1
ISN 0044
C * * * BEGIN OF LAYER LOOP
C * * * CALC OF PV
C
2150 IF (J .EQ. JM)
DO 2154 I = INC,IM,INC
PV(I,1) = EXP(J) * P(I,J) + DAP(JP1) * P(I,JP1)
S * V(I,L,JPI)
PV(I,2) = EXP(J) * P(I,J) + EXP(JP2) * P(I,JP2)
S * V(I,L,JP2)
CONTINUE
C * * * PU CALCULATION : OMIT FOR J=1, SOUTH POLE)
C
2154 IF (J .EQ. 1)
GO TC 2150
C * * * PU CALCULATION : OMIT FOR J=1, SOUTH POLE)
C
2158 IF (J .EQ. 1)
GO TC 2225
I = IM - 1
PUI1 = DYU(J) * P(1,I,J) + U(1,I,L,J)
PUI2 = DYU(J) * P(I,J) + U(I,L,J)
DO 2160 IPI = INC,IM,INC
PUIPI = DYU(J) * P(IPI,J) + U(IPI,L,J)
PUI(I,1) = PUI1 + PUI
PUI(I,2) = PUI1 + PUIPI
PUI = PUIPI
I = IPI
CONTINUE
C
2160
C * * * HORIZONTAL ADVECTION OF MOMENTUM, TEMPERATURE, MOISTURE
C * * * COMPUTE FLUXES FIRST IN THE LONGITUDINAL (I) DIRECTION
C
I = IM - 1
DO 2223 IPI = INC,IM,INC
FLUX1 = FICO * PUI(I,1)
FLUX2 = F2CC * PUI(I,2)
DO 2223 N = 1,4
QFLUX1 = FLUX1 * (Q(1,I,L,N,J) + Q(I,L,N,J))
QFLUX2 = FLUX2 * (Q(1,I,L,N,J) + Q(I,L,N,J))
IF (N.EQ.4) GO TO 110
IF (Q(1,I,L,N,J).GT.0.) QFLUX1=0.
IF (QFLUX2.GT.0.) QFLUX2=0.
IF (Q(I,I,L,N,J).LE.0.) AND. (QFLUX1.LT.0.) QFLUX1=0.
IF (Q(I,I,L,N,J).LE.0.) AND. (QFLUX2.LT.0.) QFLUX2=0.
CONTINUE
Q(I,I,L,N,J) = Q(I,I,L,N,J) + QFLUX1
Q(I,I,L,N,J) = Q(I,I,L,N,J) + QFLUX2
CONTINUE
I = IPI
CONTINUE
IF (J .EQ. JM)
GO TO 2237
C * * * NOW DO FLUX CALCULATION IN THE LATITUDINAL (J) DIRECTION
C
2225 DO 2230 I = INC,IM,INC
FLUX1 = FICO * PV(I,1)
FLUX2 = F2CC * PV(I,2)
DO 2230 N = 1,4
QFLUX1 = FLUX1 * (Q(I,L,N,J) + Q(I,L,N,JPI))
QFLUX2 = FLUX2 * (Q(I,L,N,J) + Q(I,L,N,JPI))
IF (N.EQ.4) GO TO 115
IF (Q(I,L,N,J).GT.0.) QFLUX1=0.
IF (QFLUX2.GT.0.) QFLUX2=0.
IF (Q(I,L,N,JPI).LE.0.) AND. (QFLUX1.LT.0.) QFLUX1=0.
IF (Q(I,L,N,JPI).LE.0.) AND. (QFLUX2.LT.0.) QFLUX2=0.
CONTINUE
Q(I,L,N,JPI) = Q(I,L,N,JPI) + QFLUX1
Q(I,L,N,JPI) = Q(I,L,N,JPI) + QFLUX2
CONTINUE
ISN 0045      ISN 0051
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ISN 0117
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ISN 0123

2230 QT(I,L,N,J) = QT(I,L,N,J) - QFLUX1 - QFLUX2
CONTINUE
C * * * FOLLOWING IS CORRECTION FOR CASE J=2
C * * * P(I,0) = P(INDEX(I),2)
C * * * V(I,L,2) = -V(INDEX(I),L,2) ETC.
C * * * PV(I,2) FOR J=0 IS EQUAL TO: -CONV(I,L,2) * SEE ABOVE
C * * *
IF(J.GT.1)
DO 2240 I = INC,IM,INC
CONV(I,L,2)=0.
CONTINUE
C * * * CONTINUITY EQUATION
C * * * DUE TO 1 LINE OF LATITUDE AT A TIME THE DEFINITION OF PV,
C * * * AND 4TH ORDER FORM OF CONV, WE COMPUTE IT AS FOLLOWS:
C * * *
2250 I22 = IM - 1
IS1 = IM
DO 2400 I = INC,IM,INC
CONV(I,L,JP2) = -PV(I,I,J2)
CONV(I,L,JP1) = CONV(I,L,JP1) + E * PV(I,I,1)
CONV(I,L,J) = CONV(I,L,J) - 8 * PV(I,I,1) + PV(I,I,2)
\$ 8 * PV(I,I,1) - PV(I,I,1) + (PV(I,2) - PV(I,2,2))
I22 = IS1
IS1 = I
CONTINUE
C * * * FOLLOWING IS CORRECTION FOR J=JM.
C * * * U(I,L,JM2) = -U(INDEX(I),L,JM), ETC.
C * * *
2237 IF (J.EQ.1) JM1
I22 = IM - 2
I = IM - 1
I1 = IM
DC 2250 IP2 = INC,IM,INC
PVLJM = DXPLJM * V(I,L,JM)
PV2JM = 0.
CONV(I,L,JM) = CONV(I,L,JM) - 8 * PVLJM
\$ + PV2JM + DYU(I,I) * (P(I,1,JM) * U(I,1,L,JM)
\$ - P(I,1,JM) * U(I,1,L,JM)) + (P(I,2,J) * U(I,2,L,JM)
\$ - P(I,2,JM) * U(I,2,L,JM))
CONTINUE
C * * * CALCULATION OF THE REMAINDER OF QT(I,L,N,JM)
C * * *
FLUX1 = FICO * PVLJM
FLUX2 = F2CO * PVLJM
DO 2255 N = 1,4
DELUX1 = FLUX1 * (I(L,N,JM) + S(I,L,N,NP1))
DELUX2 = FLUX2 * (I(L,N,JM) + S(I,L,N,NP1))
IF(N.EQ.1) GO TO 126
IF(I(L,N,JM).GT.0) GO TO 126
IF(DELUX1.GT.0) DELUX1=0.
IF(DELUX2.GT.0) DELUX2=0.
IF(I(L,N,NP1).GT.0) AND (QFLUX1.LT.0) DELUX1=0.
IF(I(L,N,NP1).GT.0) AND (QFLUX2.LT.0) DELUX2=0.
1: QFLUX2(LI,0) = DELUX2 *
1: QFLUX2(LI,0) = QFLUX2 *
QT(I,L,N,JM) = QT(I,L,N,J) - OFLUX1 - OFLUX2
CONTINUE
C * * *
IS2 = IS1
IS1 = I
I = IPI
IPI = IP2
CONTINUE
C * * * NOTE: MULTIPLY BY DSIG(L) FOR LAT J ONLY.
C * * *
2405 DO 2410 I = INC,IM,INC
CONV(I,L,J) = CONV(I,L,J) * DSIG(L)
CONTINUE
2410 I = L + 1
IF (L.EQ. NLAY) GO TO 2150
C * * * END OF LAYER LOOP
C * * * COMPUTATION OF SIGMA DOT AND NEW SURFACE PRESSURE
C * * * SKIP THIS COMPUTATION IF (J.EG.1) (SOUTH POLE)
C * * * P(I,I,J) IS EQUIVALENT TO CONV(I,9,J)
C * * *
IF (J.EQ.1)
DO 2440 I = INC,IM,INC
\$ P(I,I,J) = CONV(I,9,J)
DO 2420 L = 1,NLAYM1
P(I,I,J) = P(I,I,J) + CONV(I,L,J)
CONTINUE
2420 SD(I,L,J) = CONV(I,I,J) - DSIG(L) * P(I,I,J)
\$ (I,COMG) OMEGA(I,J,1) = SD(I,I,J)/DXPLJM * E6

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ISN 0182 IF (NLAY * EQ, 2) GO, TC 2440
ISN 0184 DO 2430 L = 2, NLAYM1
ISN 0185 SD(I, L, J) = SD(I, L-1, J) + CONV(I, L, J) - (SIG(L) * PIT(I, J))
ISN 0186 IF (COMG) OMEGA(I, J, L) = SD(I, L, J) / DXYP(J) * #10.ES
ISN 0187 CONTINUE
ISN 0188 CONTINUE
ISN 0189
ISN 0190 GO 2450 I = INC, IM, INC
ISN 0191 DELPT = DT * PIT(I, J) / DXYP(J)
ISN 0192 IF I = 1
ISN 0193 IF I * EQ, III = 1
ISN 0194 PT(I, J) = PT(I, J) + DT * PIT(I, J) / DXYP(J)
ISN 0195 CONTINUE
ISN 0196
ISN 0197 IF (NOT, COMG) GO TO 2456
ISN 0198 DO 2455 L = 1, NLAY
ISN 0199 DC 2455 I = I, IM
ISN 0200 OMEGA(I, J, L) = OMEGA(I, J, L) + SIG(L) * PIT(I, J) / DXYP(J)
ISN 0201 CONTINUE
ISN 0202
C * * * VERTICAL ADVECTION OF MOMENTUM AND TEMPERATURE
C * * *
DO 2490 L = 1, NLAYM1
LPI = L + 1
C
C FIRST DO MOMENTUM
DO 2470 N = 1, 2
DO 2470 I = INC, IM, INC
SFLUX = F2CO * SD(I, L, J) * (I, L, N, J) + Q(I, L, P1, N, J)
QT(I, L, P1, N, J) = QT(I, L, P1, N, J) - SFLUX / (SIG(L, P1))
QT(I, L, N, J) = QT(I, L, N, J) + SFLUX / (DSIG(L))
CONTINUE
C * * * VERTICAL ADVECTION OF TEMPERATURE
C * * * TRYING GISS CODE FOR PK CALC. INSTEAD OF 4TH ORDER CODE
DO 2480 I = INC, IM, INC
PL1 = SIG(L) * P(I, J) + PTOP
PL2 = SIG(L, P1) * P(I, J) + PTOP
PK1 = EXPBK(P, 1)
PK2 = EXPBK(PL2)
TFLUX = F2CO * SD(I, L, J) * (I, L, J) / P1 + T(I, L, P1, J) / PK2
TT(I, L, P1, J) = TT(I, L, P1, J) - PK1 * TFLUX / DSIG(L)
TT(I, L, J) = TT(I, L, J) + PK1 * TFLUX / DSIG(L)
CONTINUE
RETURN
IF (J, LT, JM)
C * * * CALCULATIONS AT THE POLES
C * * *
C * * * WHEN (J=1) OR (J=JM) WE DO THE CALCULATIONS AT THE POLES.
C * * * MOMENTUM, TEMPERATURE, SPECIFIC HUMIDITY AT SOUTH POLE.
C * * * (M = 1) AND NORTH POLE (M = 2)
C * * * JPOL(K, M) = (2, 3, JM, JM-1)
C * * * COM1 = DXPI2 / (3 PI * A * A * SIN(OLAT / 2) ** 2)
C * * * COM2 = -DXPI3 / (12 PI * A * A * SIN(OLAT) ** 2)
C * * * (K = 1) CORRESPONDS TO THE LATITUDE LINE NEAREST TO THE POLE.
C * * *
2600 M = 1
IF (J, EQ, JM) M = 2
PITPOL(M) = 0.
COEF1 = (-1.) ** M
COEF2 = -COEF1
DO 2500 L = 1, NLAY
DO 2505 N = 1, 5
SUM(K, N) = 0.
C * * * HORIZ. ADVECS. OF MOMENTUM, TEMPERATURE, SPECIFIC HUMIDITY
C * * * AT THE NORTH AND SOUTH POLES.
C * * * 4/4/78 - PIV TO BE CHANGED TO PIV FOR: -PIEQLONGPIE
CONTINUE
DO 2510 I = INC, IM, INC
DO 2510 K = 1, 2
IF (M, EQ, 2) JKP = JNP
JK = JPOL(K, M)
PIV = P(I, JK) * V(I, L, JK)
SUM(K, 1) = SUM(K, 1) - PIV * (U(I, L, JK) * SIN(LON(I)
+ COEF1 * V(I, L, JK) * COS(LON(I))
+ SUM(K, 2) = SUM(K, 2) - PIV * (COEF2 * U(I, L, JK) * COS(LON(I)
+ V(I, L, JK) * SIN(LON(I))

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C N

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ISN 0243 SUM(K,3) = SUM(K,3) + PIV * T(I,L,JK)
ISN 0244 SUM(K,4) = SUM(K,4) + PIV * SH(L,JK)
          DC 2515 N = 3.0
          CONTINUE
          CONTINUE
ISN 0245 SUM(K,5) = SUM(K,5) + PIV
ISN 0246 DC 2520 N = 1.0
ISN 0247 ODPOL(L,N,M) = ODPOL(L,N,M) + COEF1 * DT * (CON1 * SUM(L,N)
ISN 0248 + CON2 * SUM(L,N))
ISN 0249 CONTINUE
          CONTINUE
          * * * CONTINUITY EQUATION AT THE POLES.
          CONVP(L) = COEF1 * (CON1 * SUM(L,5) + CON2 * SUM(2,5)) * DSIG(L)
          PITPOL(M) = PITPOL(M) + CONVP(L)
          CONTINUE
          * * * COMPUTATION OF SIGMA DCT AND NEW SURFACE PRESSURE AT POLES.
          * * *
          SDPOL(L,M) = CONVP(L) - DSIG(L) * PITPOL(M)
          IF (CON1) OMEGA(I,JKP,1) = SDPOL(L,M) * 1.E2
          IF (NLAY .EQ. 2)
            DC 2530 L = 2, NLAYMI
          SDPOL(L,M) = SDPOL(L,1,M) + CONVP(L) - DSIG(L) * PITPOL(M)
          IF (CON1) OMEGA(I,JKP,1) = SDPOL(L,M) * 1.EE
          CONTINUE
          * * *
          PTPOL(M) = PTPOL(M) + DT * PITPOL(M)
          IF (NOT .CON1) GO TO 2537
          DC 2539 L = 1, NLAY
          OMEGA(I,JKP,L) = OMEGA(I,JKP,L) + SIG(L) * PITPOL(M)
          2537 CONTINUE
          * * * VERTICAL ADVECTION OF MOMENTUM AND TEMPERATURE AT POLES
          * * *
          DC 2540 L = 1, NLAYMI
          LPI = L + 1
          DO 2550 N=1,2
            SFLUX = F2CO * SDPOL(L,M) * (OPOL(L,N,M) + QPOL(LPI,N,M))
            ODPOL(L,N,M) = ODPOL(L,N,M) - SFLUX / DSIG(LPI)
            ODPOL(L,N,M) = ODPOL(L,N,M) + SFLUX / DSIG(L)
            CONTINUE
          * * * VERTICAL ADVECTION OF TEMPERATURE AT THE POLES.
          * * *
          P1 = SIG(L) * PPOL(M) + PTCO
          P2 = EXPBK(P1) * PPOL(M) + PTOP
          PK2 = EXPBK(P2)
          TFLUX = F2CO * SDPOL(L,M) * (TPOL(L,M) / PK1 + TPOL(LPI,M) / >K2)
          TTPOL(LPI,M) = TTPOL(LPI,M) - PK2 * TFLUX / DSIG(LPI)
          TTPOL(L,M) = TTPOL(L,M) + PK1 * TFLUX / DSIG(L)
          CONTINUE
          RETURN
          END
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NAME TAG TYPE ADD. / COMPI / SIZE OF PROGRAM 002602 HEXADECIMAL BYTES

NAME TAG TYPE ADD.	COMPI	SIZE OF PROGRAM 002602 HEXADECIMAL BYTES
ISN 0243	C	000000
ISN 0244	CE	001104
ISN 0245	C	001104
ISN 0246	XR	000000
ISN 0247	XR	000000
ISN 0248	XR	000000
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ISN 0267	XR	000000
ISN 0268	XR	000000
ISN 0269	XR	000000
ISN 0270	XR	000000
ISN 0271	XR	000000
ISN 0272	XR	000000
ISN 0273	XR	000000
ISN 0274	XR	000000
ISN 0275	XR	000000
ISN 0276	XR	000000
ISN 0277	XR	000000
ISN 0278	XR	000000
ISN 0279	XR	000000
ISN 0280	XR	000000
ISN 0281	XR	000000
ISN 0282	XR	000000
ISN 0283	XR	000000
ISN 0284	XR	000000
ISN 0285	XR	000000

ORIGINAL PAGE IS
OF POOR QUALITY

```

NAME OF COMMON BLOCK * F:U:R:Z:E OF BLOCKCOMMON 005E24 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CON1 R24 R24 000000 CON2 R24 R24 000004 ALPHA R24 R24 000028
CON2 R24 R24 000004 THBAE R24 R24 000028 BETA R24 R24 000028
CON3 R24 R24 000004 ADLDP R24 R24 000028 JMS1 R24 R24 000028
CPH1 R24 R24 000004 PSKAPA R24 R24 000028 R R24 R24 000028
CP R24 R24 000004 IMC R24 R24 000028 JMS2 R24 R24 000028
IMD2 R24 R24 000066 RADIM R24 R24 000028 JEND R24 R24 000028
JP2 R24 R24 000066 MATSUN R24 R24 000028 JEXT R24 R24 000028
JPOL R24 R24 000066 JPMOD R24 R24 000028 SMLN R24 R24 000028
INDEX R24 R24 000024 PSIGN R24 R24 000028 POLES R24 R24 000028
PHISPL R24 R24 000024 DIFF R24 R24 000028 SUM R24 R24 000028
SUM R24 R24 000024

```

**** COMMON INFORMATION ****

NAME OF COMMON BLOCK * F:U:R:Z:E OF BLOCKCOMMON 005E24 HEXADECIMAL BYTES

```

VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CON1 R24 R24 000000 CON2 R24 R24 000004 ALPHA R24 R24 000028
CON2 R24 R24 000004 THBAE R24 R24 000028 BETA R24 R24 000028
CON3 R24 R24 000004 ADLDP R24 R24 000028 JMS1 R24 R24 000028
CPH1 R24 R24 000004 PSKAPA R24 R24 000028 R R24 R24 000028
CP R24 R24 000004 IMC R24 R24 000028 JMS2 R24 R24 000028
IMD2 R24 R24 000066 RADIM R24 R24 000028 JEND R24 R24 000028
JP2 R24 R24 000066 MATSUN R24 R24 000028 JEXT R24 R24 000028
JPOL R24 R24 000066 JPMOD R24 R24 000028 SMLN R24 R24 000028
INDEX R24 R24 000024 PSIGN R24 R24 000028 POLES R24 R24 000028
PHISPL R24 R24 000024 DIFF R24 R24 000028 SUM R24 R24 000028
SUM R24 R24 000024

```

NAME OF COMMON BLOCK * F:U:R:Z:E OF BLOCKCOMMON 00D148 HEXADECIMAL BYTES

```

VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CON1 R24 R24 000000 CON2 R24 R24 000004 ALPHA R24 R24 000028
CON2 R24 R24 000004 THBAE R24 R24 000028 BETA R24 R24 000028
CON3 R24 R24 000004 ADLDP R24 R24 000028 JMS1 R24 R24 000028
CPH1 R24 R24 000004 PSKAPA R24 R24 000028 R R24 R24 000028
CP R24 R24 000004 IMC R24 R24 000028 JMS2 R24 R24 000028
IMD2 R24 R24 000066 RADIM R24 R24 000028 JEND R24 R24 000028
JP2 R24 R24 000066 MATSUN R24 R24 000028 JEXT R24 R24 000028
JPOL R24 R24 000066 JPMOD R24 R24 000028 SMLN R24 R24 000028
INDEX R24 R24 000024 PSIGN R24 R24 000028 POLES R24 R24 000028
PHISPL R24 R24 000024 DIFF R24 R24 000028 SUM R24 R24 000028
SUM R24 R24 000024

```

NAME OF COMMON BLOCK * F:U:R:Z:E OF BLOCKCOMMON 083C30 HEXADECIMAL BYTES

```

VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CON1 R24 R24 000000 CON2 R24 R24 000004 ALPHA R24 R24 000028
CON2 R24 R24 000004 THBAE R24 R24 000028 BETA R24 R24 000028
CON3 R24 R24 000004 ADLDP R24 R24 000028 JMS1 R24 R24 000028
CPH1 R24 R24 000004 PSKAPA R24 R24 000028 R R24 R24 000028
CP R24 R24 000004 IMC R24 R24 000028 JMS2 R24 R24 000028
IMD2 R24 R24 000066 RADIM R24 R24 000028 JEND R24 R24 000028
JP2 R24 R24 000066 MATSUN R24 R24 000028 JEXT R24 R24 000028
JPOL R24 R24 000066 JPMOD R24 R24 000028 SMLN R24 R24 000028
INDEX R24 R24 000024 PSIGN R24 R24 000028 POLES R24 R24 000028
PHISPL R24 R24 000024 DIFF R24 R24 000028 SUM R24 R24 000028
SUM R24 R24 000024

```

EQUILANCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETS
IS

VARIABLE OFFSETS
C 000000

VARIABLE OFFSETS
C 000000

NAME OF COMMON BLOCK * C:W799E OF BLOCKCOMMON 000C08 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
PITPOL R#4 000000

NAME OF COMMON BLOCK * *W6R9E OF BLOCKCOMMON 01D1C0 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
CONV R#4 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 000900
SD 000000 PIT

NAME OF COMMON BLOCK * *D599E OF BLOCKCOMMON 081C00 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
PU R#4 000000

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 001600
GT 000B40

NAME OF COMMON BLOCK * *D5_7E OF BLOCKCOMMON 000C04 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
ISBTR I#4 NoR.

NAME OF COMMON BLOCK * *S0T19E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME
SOOT R#4 000003 OMEGA I#2 019E00 MIMXTS I#4 NoR.
COMG L#4 0389A4

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE 00CF00
ISSOT

NAME OF COMMON BLOCK * *S619E OF BLOCKCOMMON 000C08 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME VAR. NAME TYPE REL. ADDR. ME


```

(SN 0029 JSI = J - 1
SN 0030 JS2 = J - 2
SN 0031 JP2MOD = MOD(JP2-1.5) + 1
SN 0032 JPIMOD = MOD(JPI-1.5) + 1
SN 0033 JMOD = MOD(J-1.5) + 1
SN 0034 JS1MOD = MOD(JS1-1.5) + 1
SN 0035 JS2MOD = MOD(JS2-1.5) + 1
C * *
SN 0036 JPKP2 = MOD(JP2-1.3) + 1
SN 0037 JPKP1 = MOD(JPI-1.3) + 1
SN 0038 JPK = MOD(J-1.3) + 1
SN 0039 DTKAPA = DT * KAPA
SN 0040 IF(JGT1)
SN 0041 JP2 = 1
SN 0042 JP2MOD = 1
SN 0043 JPKP2 = 1
SN 0044
SN 0045
C * * * CORIOLIS FORCE FOR 4TH ORDER DOES NOT FOLLOW GISS
C * * * METHOD OF COMPUTATION
C * * * ADLDP = (12 * RAD * DLAT * DLON = 12. * DYU(J) * DLON)
C
3001 DO 3000 L = 1,NLAY
DO 3000 I = INC,IM,INC
FX = DT * (F(J) * DXYP(J) + ADLDP * SIML(J) * U(I,L,J))
UT(I,L,J) = UT(I,L,J) + FX * P(I,J) * V(I,L,J)
VT(I,L,J) = VT(I,L,J) - FX * P(I,J) * U(I,L,J)
CONTINUE
3000
C * * * M A I N L O O P
C
C IF(JGTJ) GO TO 3032
C
3005 DO 3030 LX = 1,NLAY
LPI = L + 1
IF(JEQ1,OR,JEQJNP) GO TO 3141
COMPUTATION OF THE ENERGY TERM FOR THE THERMODYNAMIC
EQUATION FOR J=2,3,....JM-LVJM
SET UP CYCLIC I-INDICES
IS2 = IM - 3
IS1 = IM - 2
I = IM - 1
IPI = IM
SET UP J-INDICES
NDXJP2 = JP2
IF(JEQJ) NDXJP2 = JM
NDXJS2 = JS2
IF(JEQ2) NDXJS2 = 2
PERFORM CALCULATIONS
DO 3140 IP2 = INC,IM,INC
SKT = DTKAPA*SIG(L)*(I,L,J)/(PTCP+SIG(L)*P(I,J))
PIV = DXP(J)*V(I,L,J)
SET UP PROPER I-INDEX AND PRESSURE TERMS
NDXI = I
IF(JEQJ) NDXI = INDEX(I)
GI = P(NDXI,NDXJP2)
NDXI = I
IF(JEQ2) NDXI = INDEX(I)
G2 = P(NDXI,NDXJS2)
TT(I,L,J) = TT(I,L,J) + SKT*P(I,J)*PIT(I,J) + DYU(J) *
I U(I,L,J)*G2*(P(I,J)-PI(I,J)) - P(IP2,J) +
I P(IS2,J) + PIV*(G*(P(I,JPI)-P(I,JSI)) - GI + G2)
RESET CYCLIC INDICES
IS2 = IS1
IS1 = I
I = IPI
IPI = IP2
C
3140 CONTINUE
C

```

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ISN 0086 3141 CONTINUE
ISN 0087 IF(J.EQ.JM) GO TO 3030
ISN 0089 C * * COMPUTE PK AT LATITUDE JP2 (USING GISS METHOD)
IF(L.LT.NLAY) GO TC 3055
3007 DO 3010 I = 1,IM
PLI = SIGE(I) * P(I,JP2) + PTOP
PKI = PLI * EXPBK(PLI)
LL = I
DO 3010 LLPI = 2,NLAYPI
PL2 = SIGE(LLPI) * P(I,JP2) + PTEP
PK2 = LL * PL2 * EXPBK(PL2)
PK1(LL,JKPK2) = (PK2 - PK1) / (DIFF(LL) * P(I,JP2))
PK1 = PK2
LL = LLPI
CONTINUE
3010 C * * COMPUTATION OF PHI (GEOCENTRAL) AT JF2 FOR L=NLAY,
C * * PSKAPA=1000*K, SIGE(I)=0,
C * * HERE PHI IS NORMALIZED, I.E., EQUALS (STANDARD PHI)-PHIBAR.
IF (L.LT. NLAY) GO TC 3055
PHI(LL,JP2MOD) = PHIS(I,JP2) + CPTH*(PK(I,L,JKPK2) - PSKAPA)
LLM1 = I
DO 3050 LL = 1,NLAY
LLPI = LL + I
IF(LL.EQ.NLAY) LLPI = LL
DUM1 = DSIG(LL)*(PK(I,LLPI,JKPK2) - PK(I,LL,JKPK2))
DUM2 = SIGE(LL)*(PK(I,LLPI,JKPK2) - PK(I,LL,JKPK2))
WL = SIG(LL) * P(I,JP2)
DUM3 = WL * R * DSIG(LL) / (WL + PTOP)
LLM1 = LL
C PHI(I,L,JP2MOD) = PHI(I,L,JP2MOD) + DUM3 - (DUM1 - DUM2)
3050 I = DUM1 + DUM2 / PK(I,LL,JKPK2)
CONTINUE
GO TC 3065
C * * COMPUTATION OF PHI AT JF2
3055 DO 3060 I = INC,IM,INC
PHI(I,L,JP2MOD) = PHI(I,L,JP1,JP2MOD) + CFC2 * (PK(I,L,JP1,JP2) -
$ PK(I,L,JP2))
$ PK(I,L,JP2) / PK(I,L,JP2)
$ PK(I,L,JP1,JP2) - (THBAR)
CONTINUE
3060 C * * FOR PHI(I,L,0) USE INDEX(I)
C * * COMPUTE W(I,L,JP2MOD) FOR PRESSURE GRADIENT TERM AND SKT.
C * * PIV FOR ENERGY TERM IN THERMODYNAMICS EQ.
3065 IF(JP2.EQ.1.OR.JP2.EQ.JNP) GO TO 3030
DO 3070 I = INC,IM,INC
THBAR = THBAR * PK(I,L,JKPK2)
PRESME = I(L,JP2) - THBAR
PRESUR = PTOP + SIGE(I) * R * (I,JP2)
W(I,L,JP2MOD) = SIGE(I) * R
CONTINUE / PRESUR
GO TO 3030
3070 C * * IF (J=1) RETURN AND COMPUTE PHI(I,L,2), W(I,L,2),.....
C * * FOR SOUTH POLE CALCULATIONS
CONTINUE
GO TC 3032
3074 IF(J.EQ.JM)
JP2 = JP2 + 1
JP2MOD = JP2MOD + 1
JKPK2 = JKPK2 + 1
C PRESSURE GRADIENT (V EQUATION) FOR J=2
3032 DO 3031 LX = 1,NLAY
L = NLAYPI - LX
IF(J.EQ.1) GO TO 3111
3080 IF (J.GT. 2)
DO 3083 I = INC,IM,INC

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BAMI 7550
BAMI 7560
BAMI 7570
BAMI 7580
BAMI 7590
BAMI 7600
BAMI 7610
BAMI 7620
BAMI 7630
BAMI 7640
BAMI 7650
BAMI 7660
BAMI 7670
BAMI 7680
BAMI 7690
BAMI 7700
BAMI 7710
BAMI 7720
BAMI 7730
BAMI 7740
BAMI 7750
BAMI 7760
BAMI 7770
BAMI 7780
BAMI 7790
BAMI 7800
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BAMI 7890
BAMI 7900
BAMI 7910
BAMI 7920
BAMI 7930
BAMI 7940
BAMI 7950
BAMI 7960
BAMI 7970
BAMI 7980
BAMI 7990
BAMI 8000
BAMI 8010
BAMI 8020
BAMI 8030
BAMI 8040
BAMI 8050
BAMI 8060
BAMI 8070
BAMI 8080
BAMI 8090
BAMI 8100
BAMI 8110
BAMI 8120
BAMI 8130
BAMI 8140
BAMI 8150
BAMI 8160
BAMI 8170
BAMI 8180
BAMI 8190
BAMI 8200
BAMI 8210
BAMI 8220
BAMI 8230
BAMI 8240
BAMI 8250
BAMI 8260
BAMI 8270
BAMI 8280
BAMI 8290
BAMI 8300
BAMI 8310
BAMI 8320
BAMI 8330
BAMI 8340
BAMI 8350
BAMI 8360
BAMI 8370
BAMI 8380
BAMI 8390
BAMI 8400
BAMI 8410
BAMI 8420

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7/31/78 USING PHILLIPS GEOPOTENTIAL AT POLES **

** PRESSURE GRADIENT AND ENERGY CONVERSION TERMS FOR
 ** MOMENTUM TEMPERATURE EQUATIONS
 ** JPMOD(K,M) = (2.3*PDD(JM-1.5)+1)*MOD(JM-1)-1.5+1
 ** JMODPK(K,M) = (2.3*PDD(JM-1.5)+1)*MOD(JM-2.5)+1

PRPOL = SIG(L) * PPOL(M) + PTOP
 SKPOL = SIG(L) * KAPA + TPOL(L,P) / PRPOL
 WPOL = W(I,L,JP2MOD)

COMPUTE SUMMATION FOR TEMPERATURE EQUATION
 DO 3211 I=1,IM
 SUM(I,3) = SUM(I,3) + COEF1*WPOL(L,M)*COS(LON(I))
 1 * VPOL(L,M)*SIN(LON(I))*B*EP(I,JPOL(I,3))-S(I,JPOL(2,4))

MODIFY TEMPERATURE AT POLES
 TPOL(L,M) = TPOL(L,M) + CT*SKTFOL*PPOL(M)*
 1 (PTPOL(M) + CONS*SUM(I,3))

DO 3210 I=1,IM
 DO 3220 K=1,2
 JK = JPOL(K,M)
 JKMOD = JPMOD(K,M)
 MOD3J = JMODPK(K,M)
 TERM = PHI(I,L,JKMOD) + WPOL*PHI(I,JK)
 SUM(K,1) = SUM(K,1) + COS(LON(I)) * TERM
 SUM(K,2) = SUM(K,2) + SIN(LON(I)) * TERM
 CONTINUE
 CONTINUE
 DO 3222 K=1,2
 SUM(K,2) = COEF1 * SUM(K,2) / SUM(K,1) CONTINUE

DO 3225 N=1,2
 CHANGED SIGN BELOW TO -
 OTPOL(L,N,M) = OTPOL(L,N,M) - PPOL(M) * (-CONS*SUM(I,N)) +
 COM4 * SUM(I,2,N)

OMEGA(1,JKP,L) = OMEGA(1,JKP,L) + SIG(L)*IPOL(M)*CONS*SUM(I,3)
 CONTINUE
 CONTINUE
 RETURN
 END

ISN 0199
 ISN 0200
 ISN 0201

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 ISN 0203

ISN 0204

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 ISN 0218

ISN 0219
 ISN 0220

ISN 0221
 ISN 0222
 ISN 0223
 ISN 0224

NAME TAG TYPE ADD. COMP 2 / SIZE OF PROGRAM 005398 HEXADECIMAL BYTES

NAME	TAG	TYPE	ADD.	COMP	2	SIZE OF PROGRAM	HEXADECIMAL BYTES
C	C	SF	F	1	00000		
K	F	F	C	1	0000E		
P	L	SF	C	1	00000		
J	Q	XR	C	1	00000		
U	V	F	C	1	00000		
DD	DT	F	C	1	00000		
FX	F1	SF	C	1	00000		
GW	G2	SF	C	1	00000		
IM	G3	SF	C	1	00000		
KN	J1	SF	C	1	00000		
PI	LX	SF	C	1	00000		
QT	PT	SF	C	1	00000		
UD	RS	SF	C	1	00000		
CDU	TX	CF	C	1	00000		
DYU	W1	SF	C	1	00000		
I21	IS2	SF	C	1	00000		
JK1	IS2	SF	C	1	00000		
JPI	JPK	S	C	1	00000		
JSI	JPK	S	C	1	00000		
NPC	LPI	SF	C	1	00000		
PK1	PL1	SF	C	1	00000		
PSE	PL2	SFA	C	1	00000		
RAD	QSM	F	C	1	00000		
SKI	SHT	F	C	1	00000		
COM3	TAU	F	C	1	00000		
COM3	COM2	F	C	1	00000		
COM3	COM3	F	C	1	00000		
COSL	COM4	F	C	1	00000		
DLAT	CPD2	F	C	1	00000		
DUM2	DXVP	SF	C	1	00000		
IDAY	DUM3	SF	C	1	00000		
ITAU	IUM2	SF	C	1	00000		
JMDS	JAVS	F	C	1	00000		
JMDS	JMDS	F	C	1	00000		
KAPA	JMDS	F	C	1	00000		
LLP1	JMDS	F	C	1	00000		
MLW	JMDS	F	C	1	00000		
PTSD	JMDS	F	C	1	00000		
SDDT	JMDS	F	C	1	00000		
TAUE	JMDS	F	C	1	00000		
TAUT	JMDS	F	C	1	00000		

UPOL	R#4	000030	000000C	N#R#	000170	XINT	C	R#4	N#R#
ADLDP	R#4	000024	N#R#	000174	CON#2	CON#2	C	R#4	000178
DSIG0	R#4	N#R#	N#R#	0002E4	INDEX	INDEX	CE	1#2	00CP00
TSMTX	C	N#R#	N#R#	0002E4	JMSI	JMSI	SF	1#4	00017C
JNEXT	C	N#R#	N#R#	0001B0	JTEST	JTEST	SF	1#4	0001B6
JMWOD	C	000074	N#R#	N#R#	JTEST	JTEST	SF	1#2	01CE00
MOPK	C	000074	N#R#	N#R#	MSIC	MSIC	SF	1#4	01CE00
POL#S	C	N#R#	00018C	N#R#	PSIG	PSIG	SF	1#4	000030
OTPOL	C	000000	N#R#	000000	RHMAX	RHMAX	XR	R#4	000030
SHPOL	C	000000	N#R#	000000	TTPOL	TTPOL	SF	R#4	000000
CSLON	XR	000000	N#R#	N#R#	ALBED0	ALBED0	XR	R#4	000000
DUNNYC	C	0001C	N#R#	N#R#	DTMULT	DTMULT	C	R#4	N#R#
ISTART	C	N#R#	00019C	N#R#	FALTER	FALTER	C	1#4	N#R#
JMOWTS	C	N#R#	00019C	N#R#	JMONTH	JMONTH	C	1#4	00019A
MUXI#S	C	N#R#	00019C	N#R#	JSMOD	JSMOD	SF	1#4	N#R#
NSTART	C	00018	N#R#	00019C	NLAPL	NLAPL	SF	1#4	0601A6
PSKAPA	C	N#R#	N#R#	N#R#	OTPOL	OTPOL	F	1#4	00003C
SKAPA	C	00003	N#R#	000000	OTPOL	OTPOL	SF	R#4	00012C
SKAPD	C	000000	N#R#	0001B0	SKTPOL	SKTPOL	SF	R#4	N#R#
PRIME	SF	000184	00018E	N#R#	XLABEL	XLABEL	C	R#4	N#R#

***** COMMON INFORMATION *****

NAME OF COMMON BLOCK * #F#R#ZE OF BLOCKCOMMON

VAR.	NAME	TYPE	REL.	ADDR.#	PE	VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME	
	CON#	R#4	000010	N#R#		CON#	ALPHA	R#4	000008	N#R#		CON#	BETA	R#4	00000C	N#R#		
	INDZP1	R#4	000030	N#R#		JMS2	R	R#4	000038	N#R#		JMS1	RAD	R#4	N#R#			
	CPH	R#4	000030	N#R#		INSTART	R	R#4	000058	N#R#		CPD2	CPD2	R#4	00004C	N#R#		
	CP	R#4	000054	N#R#		JEND	JEND	R#4	000058	N#R#		JPI	JPI	R#4	000060	N#R#		
	IND2	R#4	000064	N#R#		JNEXT	JNEXT	R#4	0000A4	N#R#		MODPK	MODPK	R#4	000074	N#R#		
	JPOL	R#4	000084	N#R#		STNOLN	STNOLN	R#4	0000A4	N#R#		COSLON	COSLON	R#4	3002C4	N#R#		
	INDEX	R#4	003904	N#R#		CONVPL	CONVPL	R#4	N#R#		N#R#		SOPGL	SOPGL	R#4	300664	N#R#	
	PHISPL	R#4	N#R#	00C57C														

NAME OF COMMON BLOCK * #M.#ZE OF BLOCKCOMMON

VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME
	QSM	R#4				QSMPL	QSMPL	R#4				PSMPOL	PSMPOL	R#4			

NAME OF COMMON BLOCK * #NTR#E OF BLOCKCOMMON

VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME
	JSP	R#4	000000	N#R#		IM	IM	R#4	000008	N#R#		NLAY	NLAY	R#4	00000C	N#R#	
	PTOP	R#4	000010	N#R#		JSPPI	JSPPI	R#4	000028	N#R#		JNPHI	JNPHI	R#4	N#R#		
	FIM	R#4	000030	N#R#		NLAYP1	NLAYP1	R#4	000028	N#R#		J	J	R#4	N#R#		
	JMOT	R#4	N#R#	N#R#		ITEST	ITEST	R#4	N#R#			IRD7	IRD7	R#4	N#R#		
	JAYS	R#4	N#R#	N#R#		JNB	JNB	R#4	N#R#			JNB	JNB	R#4	N#R#		
	DLAT	R#4	N#R#	N#R#		DT	DT	R#4	0000BC	N#R#		TAU	TAU	R#4	N#R#		
	ITAU	R#4	N#R#	N#R#		IOAY	IOAY	R#4	N#R#			JDAY	JDAY	R#4	N#R#		
	TOFDAY	R#4	N#R#	N#R#		JMONTH	JMONTH	R#4	N#R#			JWEAR	JWEAR	R#4	N#R#		
	NSTEP	R#4	N#R#	N#R#		NCOMP3	NCOMP3	R#4	N#R#			MUGGAN	MUGGAN	R#4	N#R#		
	TAUP	R#4	N#R#	N#R#		TAUE	TAUE	R#4	N#R#			TAUO	TAUO	R#4	N#R#		
	OTMULT	R#4	N#R#	N#R#		GRAV	GRAV	R#4	N#R#			RGAS	RGAS	R#4	N#R#		
	KAPA	R#4	000118	N#R#		ED	ED	R#4	N#R#			FRU	FRU	R#4	N#R#		
	NFLW	R#4	N#R#	N#R#		HRCH	HRCH	R#4	N#R#			RSDIST	RSDIST	R#4	N#R#		
	SIND	R#4	N#R#	N#R#		RHMAX	RHMAX	R#4	N#R#			CDX	CDX	R#4	N#R#		
	DUNNYC	R#4	N#R#	N#R#		DUMNYA	DUMNYA	R#4	N#R#			XLABEL	XLABEL	R#4	N#R#		
	SIG	R#4	000370	N#R#		SICE	SICE	R#4	000010	N#R#		DSIG0	DSIG0	R#4	N#R#		
	JIPS	R#4	N#R#	N#R#		JUS	JUS	R#4	N#R#			JMSU	JMSU	R#4	N#R#		
	KSBS	R#4	N#R#	N#R#		LAT	LAT	R#4	N#R#			DXU	DXU	R#4	N#R#		
	DXP	R#4	000728	N#R#		DYP	DYP	R#4	N#R#			DAYP	DAYP	R#4	N#R#		
	F	R#4	000A0B	N#R#		COSL	COSL	R#4	N#R#			DUNNY	DUNNY	R#4	000950	N#R#	
	PHIS	R#4	000050	N#R#													

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSET# 001890
TS
VARIABLE OFFSET# 002300

NAME OF COMMON BLOCK * #WR#E OF BLOCKCOMMON

VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME	VAR.	NAME	TYPE	REL.	ADDR.#	ME
	CONV	R#4	000000	N#R#													

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSET# 000900
PIT
VARIABLE OFFSET#

NAME OF COMMON BLOCK * 005#E OF BLOCKCOMMON 081600 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
PU R#A 000000
EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE
GT 000B40 CB 00168C VARIABLE OFFSETE

NAME OF COMMON BLOCK * 005_7E OF BLOCKCOMMON 000604 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
ISWTCB I#A N#R.

NAME OF COMMON BLOCK * S#OT#9E OF BLOCKCOMMON 0385A8 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
SDOT R#A 000000 ONEGA I#2 019E00 MINMXTS I#4
COMG L#A 0389A4
EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
VARIABLE OFFSETE
ISDOT 00CF00 VARIABLE OFFSETE

NAME OF COMMON BLOCK * S#6#E OF BLOCKCOMMON 000C08 HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE
ISMTH I#A N#R. NPC I#4

NAME OF COMMON BLOCK * C#M79#E OF BLOCKCOMMON 000C0E HEXADECIMAL BYTES
VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE VAR. NAME TYPE REL. ADDR. HE

LABEL ADDR
3001 0036E8
3141 0036F8
3055 004130
3030 004370
3063 004670
3135 004A10
3211 004AF0
3225 004FE0

LABEL ADDR
3000 003AE0
3007 003E50
3060 004250
3074 0043AE
3085 0046F0
3150 004B70
3220 004F0A
3031 005050

LABEL ADDR
3005 003AF0
3610 003F60
3655 00427A
3632 0043E2
3690 004856
3111 004B98
3210 004F2E

LABEL ADDR
3140 003E10
3050 004130
3070 00435A
3080 004552
3100 0049F4
3200 004C0A
3222 004F6C

```
*OPTIONS IN EFFECT*
*OPTIONS IN EFFECT*
*STATISTICS*
*STATISTICS* NO DIAGNOSTICS GENERATED
***** END OF COMPILATION *****
LEVEL 19.6-APR 71
```

NAME= MAIN,OPT=02,LINECNT=55,SIZE=100K.

SOURCE=EBCDIC,NOLIST,NODECK,LOAD,MAP,NOEDIT,ID,NOMREF

SOURCE STATEMENTS = 223 PROGRAM SIZE= 21400

808K BYTES OF CORE NOT USED
DATE 12/12/79-0783

```
COMPILER OPTIONS -
SUBROUTINE AVRX(U,V,1,SH,P,Q,J)
INTEGR*2 ALBEDO
REAL*4 KAPA,LAI
COMMON/FOURTH/CONI,COSIN,COSON,SUM
$ ADLDP,JMS2,JMS1,COS2,COS3,COS4,COS5,TSBAR,ALPH,BETA,I,MO2PI,
$ INDE,RADIN,JEND(2),JP1,JP2,MAJSUN,INEAT(2),MORS,CPD2,
$ JPOL(2,2),JPMOD(2,2),SINL,OMI(2),COSL,OMI(2),INDEX(2,2),
$ PSIGN(4),POLES(8),W(72,9,5),PMI(72,9,3),DIFF(5),
$ CONVP(9),SDPOL(9,2),PHISPL(2),SUM(2,5)
* JSP,JNP,IM,NLAY,PTOP,ISTART,JSPPL,JNPM,I,FIN,NLAY,MLAY,PI,
* J1,JM,KM,T,U,TROT,PROT,JTEST,ITEST,
* DT,JAU,ITAU,XINT(1),J5B,JNB,DIAT,OLON,
* N,CYCLE,NCOMP3,NUOGAN,TAUP,TAUI,TAUE,TAU LO,DTMA,
* PI,GRAY,RCGAS,KAP,PL,EDAFMU,NFLW,PSF,MRCH,RODST,SIND,COSD,
* XLABEL(20),SIG(20),SIZE(21),D(10),KES(11),
* ELAT(46),DXU(46),DXV(46),DXW(46),DXP(46),ALBEDO(72,46),
* COSL(46),DUMNY(72,46),PMIS(2880,46),GT(2880,1),GM(2880,1),C(300)
$ DIMENSION U(72,46),V(72,46),W(72,46),X(72,46),Y(72,46),Z(72,46),
$ (GT(1),PU(72,1),PHIS(72,1)),GM(1),FU(144,1),PMIS(144,1),
$ SMPOL(36,1),OPOL(9,4,2),OPOL(36,1),VFDL(36,1),TPOL(36,1),
$ (PUOLS(1),PMOLS(1)),VPOLS(1),POLES(28),
$ (PPOL(1),PMPS(15))
COMMON/INSM/SYS TO,PI7:59
COMMON/DBUG/IS,TG,PI7:59
COMMON/SM00/ISMT,ICM
COMMON/S/JMIN,JMAX,JSUM,SMTH(37,17)
COMPLEX TRAN(37),JAX,JSUM,SMTH(37,17)
EQUIVALENCE(TRAN(1),DATA(1))
COMMON/WORK2/PVI(2880,46)
COMMON/S/MOTH/DAT(150) CAT:150,SLP(72,46)
SLXP(PI,110) = EXP(PI / (710 + F98 * PMI) / 287.)
TIO (PP,TB,T9) = LOG(PI + PROPI) / (F5DC * PP + PTOP)
$ + T9 * LOG((PI + PP + PTOP) / (F5DC * PP + PTOP))
MISS = 1
IF (MISS.EQ.0)
INC = 1
DRAT=DTN:(J/DSP:J) GO TO 40
IF(DRAT.GT.1) GO TO 40
ALP = 125 * DRAT - 1.0
NNE=DRAT
FNR=NEN
ALP=NALP/FNR
DO 30 N=1,4
DO 30 L=1,NLAY
I=IN
I=IN
DO 10 IPINC=INC,IM,INC
```

```
ISN 0002
ISN 0003
ISN 0004
ISN 0005
ISN 0006
ISN 0007
ISN 0008
ISN 0009
ISN 0010
ISN 0011
ISN 0012
ISN 0013
ISN 0014
ISN 0015
ISN 0016
ISN 0017
ISN 0018
ISN 0019
ISN 0020
ISN 0021
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ISN 0023
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ISN 0025
ISN 0026
ISN 0027
ISN 0028
ISN 0029
ISN 0030
ISN 0031
ISN 0032
ISN 0033
ISN 0034
ISN 0035
ISN 0036
```

```
BAMI 9760
BAMI 9780
BAMI 9790
BAMI 9800
BAMI 9810
BAMI 9820
BAMI 9830
BAMI 9840
BAMI 9850
BAMI 9860
BAMI 9870
BAMI 9880
BAMI 9890
BAMI 9900
BAMI 9910
BAMI 9920
BAMI 9930
BAMI 9940
BAMI 9950
BAMI 9960
BAMI 9970
BAMI 9980
BAMI 9990
BAM2 0000
BAM2 0010
BAM2 0020
BAM2 0030
BAM2 0040
BAM2 0050
BAM2 0060
BAM2 0070
BAM2 0080
BAM2 0090
BAM2 0100
BAM2 0110
BAM2 0120
BAM2 0130
BAM2 0140
BAM2 0150
BAM2 0160
BAM2 0170
BAM2 0180
BAM2 0190
BAM2 0200
BAM2 0210
BAM2 0220
BAM2 0230
BAM2 0240
BAM2 0250
BAM2 0260
BAM2 0270
BAM2 0280
BAM2 0290
BAM2 0300
BAM2 0310
BAM2 0320
BAM2 0330
BAM2 0340
```

```

ISN 0037      DUMNY(I)=Q(I,L,N,J)*ALPHS(Q(IPINC,L,N,J)+Q((IMINC,L,N,J)-
$      Q(I,L,N,J)-Q(I,L,N,J))
IMINC=I
10 I=IPINC
DO 20 A=INC,IM,INC
20 Q(I,A,J)=DUMNY(I)
30 CONTINUE
IMINC=IM-INC
I=IM
DO 35 IPINC=INC,IM,INC
DUMNY(I)=P(I,J)+ALPHS(P(IPINC,J)+P(IMINC,J)-P(I,J)-P(I,J))
IMINC=I
35 I=IPINC
P(I,40)=INC,IM,INC
40 CONTINUE
GO TO 60
C * * * FOURIER SMOOTHING NEAR POLES
C 12/12 VERSION 3 SMSHAP ON P AND T
IF(I,J,GT,JMIN)ANDS(J,LT,JMAX)*OR(J,EQ,J)*OR(J,EG,JMP) RETURN
JSUB=J
IF(J,GE,JMAX)
N=225NN/2
DO 50 I=1,NLAY
DO 50 I=1,IM
DATA(I)=Q(I,L,N,J)
650 DO 670 I=1,IMD2PI
TRAN(I)=SMTH(I,JSUB)*TRAN(I)
670 DO 680 I=1,IMD2PI / FLOAT(IM)
CONTINUE
CALL FOURT2(CATA,IM,1,-1.0)
CONTINUE
CALL FOURT2(TRAN,IM,1,+1,-1)
CONTINUE
CALL SMSHAP
CONTINUE
680 IF(J,EQ,JM)
CONTINUE
690 DO 100 I=1,IM
SLP(I,J)=P(I,J)
C * * * TRANSFORM TO SEA LEVEL PRESSURE
CONTINUE
C
DO 110 I=1,IM
TSURF=T(SLP(I,J),T(L,NLAY,I),T(L,NLAY,J))
SLP(I,J)={SLP(I,J)+PTOP}*SLEXP(PHS(I,J),TSURF)
CONTINUE
110 IF(IM,GT,0)GO TO 781
DATA(I)=SLP(I,J)
760 CONTINUE
CALL FOURT2(DATA,IM,1,-1.0)
DO 770 I=1,IMD2PI
TRAN(I)=SMTH(I,JSUB)*TRAN(I)
770 CONTINUE
CALL FOURT2(TRAN,IM,1,1,-1)
DO 790 I=1,IM
SLP(I,J)=DATA(I)/FLOAT(IM)
780 CONTINUE
781 CONTINUE
C * * * SMOOTHING ALONG LONGITUDE (SEE SMSHAP)
C
NSM=8
IF(J,LT,8)*OR(J,GT,JM-6) NSM=4
IF(J,LT,6)*OR(J,GT,JM-4) NSM=2
IF(J,LT,4)*OR(J,GT,JM-2) NSM=1
SGN=I-1
CX=4
DO 130 I=1,IM
DATA(I)=SLP(I,J)
130 DO 140 N=1,NSM
I=IM
DO 145 IP=I,IM
CATA(IP)=DATA(IP)-DATA(I)
I=IP
CONTINUE
145 IS=IM
DO 150 I=1,IM
DATA(I)=CATA(I)-CATA(IS)
IS=I
CONTINUE
150 DO 160 I=1,IM
CONTINUE
160

```

ISN 0038
ISN 0039
ISN 0040
ISN 0041
ISN 0042
ISN 0043
ISN 0044
ISN 0045
ISN 0046
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ISN 0118

BAM20350
BAM20350
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BAM20380
BAM20390
BAM20400
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BAM20690
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BAM21000
BAM21010
BAM21020
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BAM21070
BAM21080
BAM21090
BAM21100
BAM21110
BAM21120
BAM21130
BAM21140
BAM21150
BAM21160
BAM21170
BAM21180
BAM21190
BAM21200
BAM21210
BAM21220

JPOL 104 No.R
 INDEX 104 No.R
 PK 104 No.R
 PHISPL 104 No.R
 CUSLJN 104 No.R
 SDPOL 104 No.R

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000414
 VPOL 000430

VARIABLE OFFSETE 000438
 VPOL

SINLUN 104 No.R
 CONPOL 104 No.R
 POLES 104 No.R
 200414
 No.R

VARIABLE OFFSETE 00045C
 TPOL

NAME OF COMMON BLOCK * 083C30 HEXADECIMAL BYTES
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JSC 104 000000 000004
 P10 104 000010 000004
 P14 104 000020 000024
 P18 104 000030 000024
 JTEST 104 000040 000044
 JINC 104 000050 000044
 JLN 104 000060 000044
 JDATE 104 000070 000044
 JCYCLE 104 000080 000044
 JTAUT 104 000090 000044
 JSL 104 000100 000044
 JPSF 104 000110 000044
 JCOED 104 000120 000044
 JALTER 104 000130 000044
 JSIG 104 000140 000044
 JAPS 104 000150 000044
 JNBS 104 000160 000044
 JYU 104 000170 000044
 JLN 104 000180 000044
 JALBDD 104 000190 000044
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 IM 104 000008
 JSPPI 104 000008
 RLAYPI 104 000008
 JTAUT 104 000008
 JTEST 104 000008
 JDB 104 000008
 JDAY 104 000008
 JMONTH 104 000008
 JNCMP 104 000008
 JTAUT 104 000008
 JGRAY 104 000008
 JED 104 000008
 JMRCH 104 000008
 JSHAX 104 000008
 JUMNYA 104 000008
 JSIG 104 000008
 JIUS 104 000008
 JLAT 104 000008
 JYU 104 000008
 JCSL 104 000008

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000000
 VPOL

NAME OF COMMON BLOCK * 115670 OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 F5DE 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 F17 104 000004

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 ISATCP 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 NPC 104 000004

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 ISMTH 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JMAX 104 000004

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JMIN 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 JSUM 104 000004

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 PU 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 CATA 104 000004

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000840
 VPOL

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 DATA 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 SLP 104 000004

EQUIVALENCED VARIABLES WITHIN THIS COMMON GROUP
 VARIABLE OFFSETE 000C30
 VPOL

NAME OF COMMON BLOCK * 0527E OF BLOCKCOMMON
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 DUMNY 104 000000 000004
 VAR. NAME TYPE REL. ADDR. ME ADDR. ME
 DUMNY 104 000004


```

ISN 0033      GO TO 140
ISN 0034      DO 110 I = 1,NLAY
ISN 0035      DO 110 I = 1,IM
ISN 0036      POT(I,J,L) = T(I,L,J)
ISN 0037      CONTINUE
ISN 0038      DO 100 I = 1,IM
ISN 0039      SLP(I,J) = P(I,J)
ISN 0040      CONTINUE

C * * TRANSFORM TO SEA LEVEL PRESSURE AND POTENTIAL TEMPERATURE * *
C * *
C
C 140 DO 210 I = 1,IM
C      TSURF = T10(S(I,J),POT(I,J,8),POT(I,J,5))
C      DO 220 L = 1,NLAY
C      PKAPA = EXPB*(SLP(I,J) * SIGIL) + PTOP)
C      POT(I,J,L) = POT(I,J,L) / PKAPA
C      CONTINUE
C 220 PHISX = PHIS(I,J)
C      SLP(I,J) = (SLP(I,J) + PTOP) * SLEXP(PHISX,TSURF)
C      IF (J.EQ.1 OR J.EQ.JNP) GO TO 230
C      CONTINUE
C      GO TO 200

C 230 DO 240 I = 2,IM
C      DO 250 L = 1,NLAY
C      POT(I,J,L) = POT(I,J,L)
C      CONTINUE
C 250 SLP(I,J) = SLP(I,J)
C      CONTINUE
C 260 NF = -1
C * * SMOOTHING ALONG LONGITUDE * *
C * *
C * * *
C * * * 16TH ORDER AT ALL J VALUES. * *
C
C 1000 NF = NF + 1
C      J = 2,JM
C      JNM = J - 1
C      IF (J.LT.8 OR J.GT.JM-6) NSM = 4
C      IF (J.LT.6 OR J.GT.JM-4) NSM = 2
C      IF (J.LT.4 OR J.GT.JM-2) NSM = 1
C      SGN = (-1)**NSM
C      CX = NF * GT * NSM
C      IF (NF * GT * 0)
C      DO 1020 I = 1,IM
C      DATA(I) = SLP(I,J)
C      CONTINUE
C      GO TO 1024

C 1020 CONTINUE
C      GO TO 1028

C 1024 DO 1026 I = 1,IM
C      DATA(I) = POT(I,J,NF)
C      CONTINUE
C 1026 DO 1030 N = 1,NSM
C      I = IM
C      DO 1035 IP = 1,IM
C      CATA(IP) = DATA(IP) - DATA(I)
C      I = IP
C      CONTINUE
C 1035 IS = IM
C      DO 1040 I = 1,IM
C      DATA(I) = CATA(I) - CATA(IS)
C      IS = I
C      CONTINUE
C      GO TO 1054

C 1040 CONTINUE
C      GO TO 1054

C 1050 IF (NF * GT * 0)
C      DO 1050 I = 1,IM
C      SLP(I,J) = SLP(I,J) -
C      DATA(I) / CX
C      GO TO 1010

C 1054 DO 1056 I = 1,IM
C      POT(I,J,NF) = POT(I,J,NF) -
C      DATA(I) / CX
C      CONTINUE
C 1056 CONTINUE
C 1010 CONTINUE

C * * SMOOTHING ALONG LATITUDE * *
C * * * ALSO USE PERIODICITY ALONG COMPLETE MERIDIAN (J=1,2*JM) * *
C * *
C
C      CY=4**NSM
C      IMD2 = IM / 2
C      JMT2 = JM / 2
C      DO 1110 I = 1,IMD2
C      IF (NF * GT * 0)
C      DO 1120 J = 1,JM
C      DATA(J) = SLP(I,J)
C      DATA(J+JM) = SLP(I+IMD2,JM-J+2)
C      CONTINUE
C      GO TO 1124

C 1120 CONTINUE
C      GO TO 1128

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C * * * SHOULD PUT (-1) FOR WIND SMOOTHING IN 1126
C
1124 DO 1126 J = 1, JM
DATA(J) = POT(I, J, NF)
DATA(J+JM) = POT(I+IMD2, JM-J+2, NF)
CONTINUE
1126
1128 DO 1130 N=1, NSM
J = JMT2
DO 1135 JP = 1, JMT2
CATA(J) = DATA(JP) - DATA(J)
J = JP
CONTINUE
1135 JS = JMT2
DO 1140 J = 1, JMT2
DATA(J) = CATA(J) - CATA(JS)
JS = J
CONTINUE
1140
1130 IF (NF .GT. 0)
CONTINUE
GO TC 1154
1140
1150 DO 1150 J = 1, JM
POT(I, J, NF) = POT(I, J, NF) - DATA(J) / CY
POT(I+IMD2, JM-J+2, NF) = SLP(I+IMD2, JM-J+2) - DATA(J+JM) / CY
CONTINUE
GO TC 1110
1150
1154 DO 1170 J = 1, JM
POT(I, J, NF) = POT(I, J, NF) - DATA(J) / CY
POT(I+IMD2, JM-J+2, NF) = POT(I+IMD2, JM-J+2, NF) - DATA(J+JM) / CY
CONTINUE
GO TC 1110
1170
1110 IF (NF .GT. 0)
S1 = 0.
S2 = 0.
DO 1160 I = 1, IM
S1 = S1 + SLP(I, JNP)
S2 = S2 + SLP(I, JNP)
CONTINUE
1160 S1 = S1 / IM
S2 = S2 / IM
DO 1165 I = 1, IM
SLP(I, J) = S1
SLP(I, JNP) = S2
CONTINUE
1165
1164 S1 = 0.
S2 = 0.
DO 1173 I = 1, IM
S1 = S1 + POT(I, J, NF)
S2 = S2 + POT(I, JNP, NF)
CONTINUE
1173 S1 = S1 / IM
S2 = S2 / IM
DO 1175 I = 1, IM
POT(I, J, NF) = S1
POT(I, JNP, NF) = S2
CONTINUE
1175
1180 IF (NF .LT. NLAY)
CONTINUE
GO TC 1000
C
C * * * SECOND MAIN J-LOOP
C * * * TRANSFORM BACK TO SURFACE PRESSURE
C
MAXIT = 50
DO 2000 J = 1, JNP
DO 1320 I = 1, IM
PHSX = PHIS(I, J)
PSFC = SLP(I, J) - PTOP
DO 1330 NITX = 1, MAXIT
PI = PSFC
T8 = POT(I, J, 8) * EXPBK(P5FC * SIG(8) + PTOP)
T9 = POT(I, J, 9) * EXPBK(P5FC * SIG(9) + PTOP)
TSURF = T10(P5FC, T8, T9)
PSFC = SLP(I, J) / SLEXP(PHISX, TSURF) - PTOP
IF (ABS(PSFC-PI)/1000. .LE. 1.E-5) GO TC 1340
PRINT 1335, MAXIT
FORMAT(IX, ' MORE THAN ', IS, ' ITERATIONS.', I)
PRINT 7990, J
PRINT 7991, I
FORMAT(IX, ' J = ', I5)
7990 FORMAT(IX, ' I = ', I5)
7991 FORMAT(IX, ' I = ', I5)
1340 SLP(I, J) = PSFC
IF (J.EQ.1 .OR. J.EQ.JNP)
GO TC 1350
CONTINUE
1350 DO 1360 I = 2, IM
CONTINUE
GO TC 1310
1360

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1360 SLP(I,J) = SLP(I,J) CONTINUE  
1370 CONTINUE  
C * * * TRANSFORM BACK TO TEMPERATURE  
C  
DO 1400 L = 1,NLAY  
DO 1420 I = 1,IM  
PKAPA = EXPBKISLP(I,J) * SIGIL) + PTC(I)  
POT(I,J,L) = POT(I,J,L) * PKAPA  
IF (J.EQ.1 .OR. J.EQ.JNP) GO TO 1430  
CONTINUE  
GO TO 1400  
1420 DC 1440 I = 2,IM  
1430 POT(I,J,L) = POT(I,J,L) CONTINUE  
1440 CONTINUE  
1400  
C * * * RETURN SMOOTHED VALUES  
C  
DO 1520 L = 1,NLAY  
DO 1520 I = 1,IM  
T(I,L,J) = POT(I,J,L)  
1520 P(I,J) = SLP(I,J) CONTINUE  
1530 CONTINUE  
2000  
DO 1550 L = 1,NLAY  
TPOL(L,1) = POT(1,1,L)  
TPOL(L,J2) = POT(1,JNP,L) CONTINUE  
1550 PPOL(1) = SLP(1,1)  
PPOL(2) = SLP(1,JNP)  
C * * *  
12714 RUN HERE TILL RETURN  
IF(SATSNX.EQ.0)RETURN  
DO 500 J=2,JM  
DO 500 N=1,4  
DO 500 K=1,NLAY  
DO 500 I=1,IM  
GT(I,K,N,J)=P(I,J)*DXYP(J)*Q(I,K,N,J)  
500 CONTINUE  
DO 501 J=2,JM  
DC 501 I=1,IM  
PT(I,J)=G(I,J)  
501 CONTINUE  
DO 502 L=1,NLAY  
DO 502 N=1,4  
DO 502 M=1,2  
QPOL(L,N,M)=PPOL(M)*QPOL(L,N,M)  
502 CONTINUE  
PTPOL(1)=PPOL(1)  
PTPOL(2)=PPOL(2)  
599 FORMAT(1X,' NEW SMOOTHER ENTERED. TAU=',F12.3)  
RETURN  
END
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