## 3.3A AN EVALUATION OF THE ACCURACY OF SOME RADAR WIND PROFILING TECHNIQUES

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## INTRODUCTION

Major advances in Doppler radar measurement in optically clear air have made it feasible to monitor radial velocities in the troposphere and lower stratosphere. For most applications we want to monitor the three dimensional wind vector rather than the radial velocity. Measurement of the wind vector with a single radar can be made assuming a spatially linear, time invariant wind field. The components and derivatives of the wind are estimated by the parameters of a linear regression of the radial velocities on functions of their spatial locations. The accuracy of the wind measurement thus depends on the locations of the radial velocities.

PETERSON and BALSLEY (1979) point out that a tradeoff exists for a given technique between the accuracies of horizontal and vertical component measurements. Because we usually need to measure the three components of wind with different accuracy and as inexpensively as possible, we are led to evaluate the suitability of some of the common retrieval techniques for simultaneous measurement of both the vertical and horizontal wind components. The techniques we will consider are fixed beam, azimuthal scanning (VAD) and elevation scanning (VED).

ERROR ANALYSIS THEORY

The estimation of the parameters of a linear wind field from radial velocities is discussed by KOSCIELNY et al. (1982). The measured radial velocity  $v_r$  can be modelled by a linear regression equation of the form

(1)

where  $P_m$  is a row vector of regressor variables which are functions of range r, azimuth  $\theta$ , and elevation angle  $\theta_e$ ;  $K_m$  is a column vector of m parameters. The measured  $v_r$ , a reflectivity weighted mean of radial velocities within the radar's resolution volume, can contain errors  $\varepsilon$  due to nonuniform reflectivity, turbulence, targets such as hydrometeors that move relative to the wind, and a nonlinear wind. It can be shown that, given n measurements of  $v_r$ , least squares estimates of  $K_m$  are computed by

 $\hat{\mathbf{K}}_{\mathrm{m}} = (\mathbf{P}_{\mathrm{nm}}^{\mathrm{T}} \mathbf{P}_{\mathrm{nm}})^{-1} (\mathbf{P}_{\mathrm{nm}}^{\mathrm{T}} \mathbf{V}_{\mathrm{n}})$ (2)

where T indicates transpose and  $P_{nm}$  is an nxm matrix of the regressor variables corresponding to the n radial velocity measurements in  $V_n$ . Measurement errors in the radial velocities produce uncertainties in the estimate  $K_m$  and the covariances of  $K_m$  about  $K_m$  are given by

$$\mathbf{C}_{\mathrm{mm}} = (\mathbf{P}_{\mathrm{nm}}^{\mathrm{T}} \mathbf{P}_{\mathrm{nm}})^{-1} \sigma_{\varepsilon}^{2}$$
(3)

where  $\sigma_{\epsilon}^{2}$  is the variance of  $\epsilon$ .

If the wind field has variations not modeled by (1), the  $\hat{K}_m$  will be biased and the amount of bias  $B_m$  is given by the product of a known alias matrix  $A_{m,\ell}$ with the vector  $K_{\ell}$  of the  $\ell$  unknown parameters of the wind field not included in  $K_m$ . Thus

$$\mathbf{B}_{\mathrm{m}} = \mathbf{A}_{\mathrm{m}\ell} \mathbf{K}_{\ell} \tag{4}$$

where

$$A_{m\ell} = (P_{nm}^{T} P_{nm})^{-1} (P_{nm}^{T} P_{n\ell}).$$
 (5)

 $P_{n\ell}$  is a matrix of regressor variables for the components not included in (1).

The various techniques referred to in the introduction assume the wind to be uniform (i.e., the first and higher order derivatives are zero) over the data analysis volume. However, because w cannot be uniform for any appreciable depth of the troposphere (i.e., w must be zero at the earth's surface), the horizontal wind can never be uniform at all heights. Thus we must account for errors produced by wind shear. We propose to analyze the errors in these techniques by computing the bias and variance of the least squares estimates  $\hat{K}_m$  with assumptions that  $\sigma_c^{-2}$  is constant and the wind field is actually linear. Thus  $K_3$ contains the three uniform components  $u_0$ ,  $v_0$ ,  $w_0$  and  $K_8$  contains the 8 spatial derivatives  $(u_x, u_z, v_y, v_z, u_y + v_x, w_x, w_y, w_z)$  of the linear wind. In our evaluation and comparison of techniques, we assume that a total of n measurements are available for each and these n measurements are distributed in space to estimate wind at some height h.

#### (a) Fixed Beam

We consider a configuration for the fixed beam technique in which three beams, one vertical and two off-vertical at elevation  $\theta_e$ , are sampled. The offvertical beams usually have perpendicular horizontal projections; for convenience, we will consider them to have azimuths 0° and 90°. The total number of radial velocity measurements for a height h for all three beams is n; for generality, we let the number of vertical measurements be N.

The bias and variance properties of the estimates  $\hat{k}_3^T = (u_0, v_0, w_0)$  are computed in Appendix 1 using (3) and (5), and we find that, for n = 3N,

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{\sigma_{e}^{2}}{n} \cdot 3 \cdot (\sec^{2}\theta_{e} + \tan^{2}\theta_{e})$$
$$VAR(\hat{w}_{o}) = \frac{\sigma_{e}^{2}}{n} \cdot 3 \quad (6)$$

Bias  $\begin{vmatrix} \hat{u}_{o} \\ \hat{v}_{o} \\ \hat{w}_{o} \end{vmatrix} \approx h \begin{bmatrix} u_{x} \cot \theta_{e} + w_{x} \\ v_{x} \cot \theta_{e} + w_{x} \\ v_{y} \cot \theta_{e} + w_{y} \\ 0 \end{bmatrix}$  (7)

The bias equation is approximate because we have used  $h^{\approx} rsin\theta_{e}$  which should be appropriate for r<30 km. From (7) we see that the bias due to spatial derivatives is a linear function of height which is expected because the beam separation is linearly dependent on h. In addition, we see from (6) and (7)

that the variance decreases with n, but that the bias cannot be reduced by data averaging.

# (b) Azimuthal Scanning

In an azimuthal scanning technique, usually called VAD (<u>Velocity</u> <u>Azimuth Display</u>), data along a circle centered on the radar are used to directly estimate the components of the uniform wind field. The results of the bias and variance equation evaluation, in Appendix 2, are that

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} \cdot 2sec^{2}\theta_{e}$$
$$VAR(\hat{w}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} \cdot csc^{2}\theta_{e}$$
(8)

Bias 
$$\begin{bmatrix} \hat{u}_{o} \\ \hat{v}_{o} \\ \hat{w}_{o} \end{bmatrix} \approx h \begin{bmatrix} w_{x} \\ w_{y} \\ 1/2(u_{x} + v_{y}) \cdot \cot^{2}\theta_{e} \end{bmatrix}$$
 (9)

# (c) Elevation Scanning

In the elevation scanning (Velocity Elevation Display) technique, radial velocities at a height h are collected for elevation angles  $\theta_0 \leq \theta_e \leq 180 - \theta_0$ . We assume  $\frac{n}{2}$  data are collected for the two azimuths 0° and 90° so both horizontal components are measured. It is shown in Appendix 3 that

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} \cdot \frac{4(\pi - 2\theta_{o})}{\{\pi - 2\theta_{o} - \sin(\pi - 2\theta_{o})\}}$$
$$VAR(\hat{w}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} \cdot \frac{2(\pi - 2\theta_{o})}{\{\pi - 2\theta_{o} + \sin(\pi - 2\theta_{o})\}}$$
(10)

Bias 
$$\begin{bmatrix} \hat{\mathbf{u}}_{0} \\ \hat{\mathbf{v}}_{0} \\ \hat{\mathbf{w}}_{0} \end{bmatrix} \approx \mathbf{h} \begin{bmatrix} \mathbf{w}_{\mathbf{x}} \\ \mathbf{w}_{\mathbf{y}} \\ (\mathbf{u} + \mathbf{v}_{\mathbf{y}}) \\ \mathbf{x} \\ \mathbf{y} \end{bmatrix} \cdot \begin{bmatrix} \frac{\pi - 2\theta_{0} - \sin(\pi - 2\theta_{0})}{\pi - 2\theta_{0} + \sin(\pi - 2\theta_{0})} \end{bmatrix}$$
 (11)

# ERROR COMPARISON

The results of our analysis of the three techniques, summarized in Table 1 show the variances of the wind estimates all depend on  $\sigma^2/n$ . Since the variance of an average of n independent data is  $\sigma_c^2/n$ , we will divide this quantity by the variance of the estimate of the wind component. Because of its similarity to the usual statistical definition, we term this quantity the efficiency of the estimate.

The variation of the efficiences of the horizontal wind estimates with elevation angle are shown in Figure 1. The VAD technique has the highest efficiency of the techniques for all elevation angles. In addition, the VAD

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Variance and bias equations for horizontal and vertical wind estimates obtained from fixed beam, azimuth scanning or elevation scanning techniques. VAR[ $\hat{v}_{o}$ ]  $\exists$ VAR[ $\hat{v}_{o}$ ] and for Bias ( $\hat{v}_{o}$ ) replace x subscript with y. Table 1.

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	Fixed Beam (one vertical)	Azimuth Scanning (VAD)	Elevation Scanning (VED)
VAR (û <sub>o</sub> )	$\frac{\sigma_{\varepsilon}^{2}}{n} 3(\sec^{2}\theta_{e} + \tan^{2}\theta_{e})$	$\frac{\sigma_{\rm c}^2}{n}$ 2 sec <sup>2</sup> $\theta_{\rm e}$	$\frac{\sigma_{\varepsilon}^{2}}{n} = 2 \left[ \frac{\pi - 2\theta_{o}}{(\pi - 2\theta_{o}) - \sin(\pi - 2\theta_{o})} \right]$
VAR ( $\hat{w_o}$ )	n 8 3	$\frac{\sigma}{n}^2 \csc^{2\theta}_{\theta}$	$\frac{\sigma_{\varepsilon}^{2}}{n} 2 \left[ \frac{\pi - 2\theta_{o}}{(\pi - 2\theta_{o}) + \sin(\pi - 2\theta_{o})} \right]$
Bias $(\hat{u}_{o})$	$h(u_x \cot \theta_e + w_x)$	hw <sub>x</sub>	hw <sub>x</sub>
Bias( $\hat{w}_{o}$ )	O	$h(u_{x} + v_{y}) \stackrel{\cot^{2}\theta}{=} h$	$h(u_{x} + v_{y}) \left[ \frac{\pi - 2\theta_{o} - \sin(\pi - 2\theta_{o})}{\pi - 2\theta_{o} + \sin(\pi - 2\theta_{o})} \right]$

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maintains reasonable efficiency to larger elevation angles (75°) than either fixed beam or VED.

Figure 1 shows efficiencies monotonically increasing as  $\theta_e$  gets smaller. But for measurements at a constant height the range increases when  $\theta_e$  decreases. Because echo power reduces in proportion to the inverse square of range (assuming the echoing layer scatters isotropically and is horizontally homogeneous) the signal-to-noise ratio (SNR) falls as  $\theta_e$  decreases. If measurement errors are solely due to thermal noise and SNR is less than one, ZRNIC' (1979) shows that measurement variance  $\sigma_e^2$  is proportional to (SNR)<sup>-2</sup>. The effect of decreasing SNR as  $\theta_e$  becomes small is to increase  $\sigma_e^2$  as csc  $\theta_e$ ; the efficiency of the horizontal wind measurements thus vanishes as  $\theta_e$  goes to zero. However, the variance  $\sigma_e^2$  includes meteorological effects such as turbulence that, in our experience, places a lower bound on  $\sigma_e^2$  of about 1 ms<sup>2-2</sup>. Because we are mainly concerned in this paper with elevation angles larger than 40° and, consequently, ranges less than about 30 km,  $\sigma_e^2$  can be regarded as a constant.

The efficiencies of the vertical velocity estimates are shown in Figure 2. The VED has the highest efficiency, but for large elevation angles the VAD is comparable. The fixed beam has a constant efficiency of  $\frac{1}{3}$  since the number of vertical estimates is fixed at  $\frac{n}{3}$ .

The biases of the estimates show a linear dependence on the height h. To normalize bias errors, we assume h = 1 km, so the bias for greater heights can be simply computed. The biases depend on the value of unknown spatial derivatives. Following WALDTEUFEL and CORBIN (1979) we use  $u_x = v_y = 10^{-3} \text{s}^{-1}$  and  $w_x = w_y = 10^{-4} \text{s}^{-1}$  as maximum values. The biases thus computed are shown in Figure 3 for the horizontal components and in Figure 4 for the vertical component. The asymmetry of the beam locations about the vertical for the fixed beam technique produces a horizontal wind bias due to  $u_x$  and  $v_y$  which decreases as  $\cot \theta_e$ . The horizontal wind biases for the VAD and VED are the same and are constant with elevation angle. For the fixed beam technique the vertical velocity is not biased by any derivatives. The vertical wind bias in the VED and VAD decreases with increasing elevation angle.

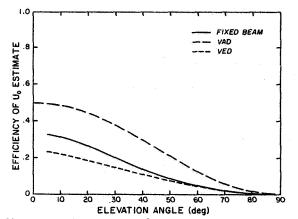
In conclusion, the vertical velocity variance and the bias errors can be decreased by using larger elevation angles. The variance for horizontal components increases with elevation angle but can be controlled to an extent by data averaging. Because bias increases with height, the higher altitudes may require a vertical measurement for vertical velocity.

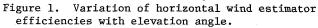
## (a) Example

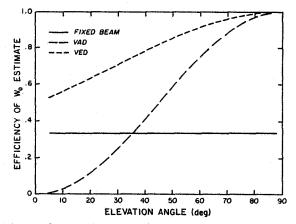
The bias the variance equations can be used to choose an elevation angle for profiling. For example, suppose we wish to profile the winds at 5 km using 360 measurements with  $\sigma_{\rm g} = 1 \, {\rm m} \cdot {\rm s}^{-1}$ . The root mean square errors (bias squared plus variance) for the horizontal and vertical components are shown in Figures 5 and 6 for each of the techniques. We have kept vertical and horizontal errors separate because vertical velocity is much smaller and requires greater accuracy. If we require horizontal and vertical velocity accuracies of  $1 \, {\rm m} \cdot {\rm s}^{-1}$ and 0.1 m·s<sup>-1</sup> respectively, we would use an elevation angle between 83° and 85° for the VAD and 77° and 81° for the VED. Because of the bias error, the fixed beam horizontal wind error is 1.5 m·s<sup>-1</sup> or larger.

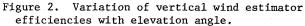
#### POSSIBLE IMPROVEMENTS

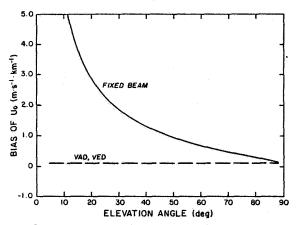
The analysis of the previous section suggests some simple improvements that can be made to increase the accuracy of the measurements.

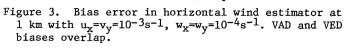


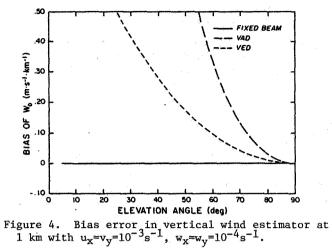












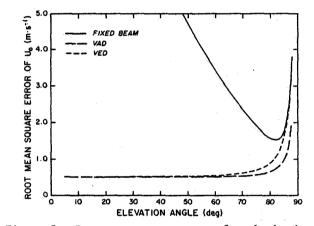


Figure 5. Root mean square error for the horizontal estimator if  $\sigma\text{=}1\ \text{m}\cdot\text{s}^{-1},\ \text{n}\text{=}360,\ \text{and}\ \text{h}\text{=}5\ \text{km}.$ 

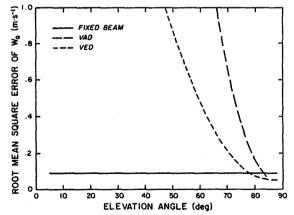


Figure 6. Root mean square error for the vertical estimator is  $\sigma\text{=}1\ \text{m}\text{\cdot}\text{s}^{-1},\ n\text{=}360,\ \text{and}\ h\text{=}5\ \text{km}.$ 

# (a) Fixed Beam with Error Minimization

The wind estimate efficiencies for the fixed beam technique can be improved slightly be collecting a specified number N of vertical data. Minimizing the first diagonal element of the matrix in (Al.1), which is  $VAR(\hat{u}_0) = VAR(\hat{v}_0)$  gives

$$\frac{N}{n} = \frac{\sin\theta}{\sqrt{2} + \sin\theta}$$
(12)

so N would vary from  $\frac{n}{3}$  at  $\theta_e = 45^\circ$  to about 0.4 n for  $\theta_e$  near 90°. The efficiencies are shown in Figure 7 for  $u_0$  and Figure 8 for  $\hat{w}_0$ . It can be seen that for  $\theta_e > 45^\circ$ , the efficiency of estimating  $\hat{u}_0$  is unchanged but is slightly improved for  $w_0$ .

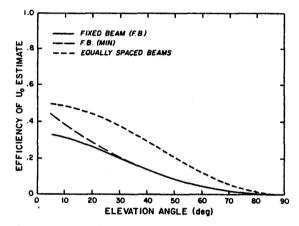


Figure 7. Horizontal wind estimator efficiency for fixed beam, fixed beam with horizontal error minimization, and fixed beam with equally spaced, off-vertical beams.

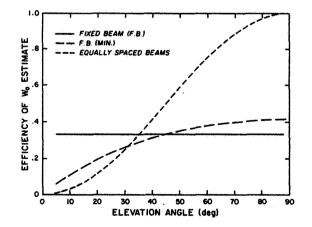


Figure 8. Vertical wind estimator efficiencies for the same techniques as described in Figure 7 caption.

The size of the variance contribution (i.e.,  $\sigma_{\epsilon}^{2} \tan^{2}\theta_{e}/N$ ) from the vertical velocity bias removal appears to indicate that it might be better to ignore the bias error ( $w_{0} \tan \theta_{e}$ ). However, for observation time intervals of several minutes, a mesoscale value of  $w_{0}$  should be used. For large elevation angles, ( $\theta_{e} \gtrsim 75^{\circ}$ ) the bias error could be several meters per second or larger. For tropospheric observation under all conditions, the bias should be removed if a horizontal velocity accuracy of  $1 \text{ m} \cdot \text{s}^{-1}$  is required.

# (b) Three Off-Vertical Beams

For some applications, ground clutter presents a problem for vertical measurements. The fixed beam technique with three off-vertical beams with elevation  $\theta_e$  and azimuths 0°, 120°, 240° is analyzed in Appendix 1b. The analysis shows that the variances (and efficiencies) of the estimates are identical to those for the VAD technique (compare Figures 1, 2 with Figures 7, 8). The biases for the  $u_0$  estimate is very similar to the fixed beam with one vertical but the  $w_0$  estimate is biased as shown in Figure 9.

#### (c) Application of the Continuity Equation to VAD Data

Vertical winds as small as few centimeters per sec are important in forecasting and, as noted earlier, w<sub>o</sub> should be estimated with more accuracy than the horizontal components. Because of ground clutter it may become very difficult to estimate the radial component of air motion when the beam is pointed near the vertical since the radial velocities will have values close to zero. We now show that by assuming a linear wind field and applying the mass gontinuity equation, we can estimate vertical wind, averaged over the circle of measurement, with the required accuracy. When mass continuity is applied, we will call the technique indirect whereas the previously discussed techniques (e.g., VAD) are direct measurements of w.

Two VAD modes in which vertical soundings can be made are fixed  $\theta_e$ variable r, and variable  $\theta_e$  fixed r. With variable r however, the horizontal area for which w<sub>o</sub> is representative varies, so we prefer the second mode. Divergence is estimated by applying Gauss's theorem to the volume V (see Figure 10) enclosed by the area S<sub>1</sub> at constant range from the radar and the area S<sub>2</sub> at constant height (DOVIAK and ZRNIC', 1983). Applying mass continuity and integrating gives an areal averaged w,

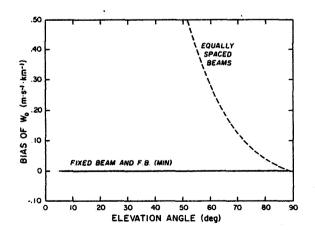


Figure 9. Vertical wind biases for the same techniques as described in Figure 7 caption.

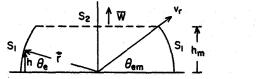


Figure 10. Geometry to estimate vertical velocity averaged over the circular area  $S_2$ .

$$\overline{w} = \frac{-2e^{\Gamma h}}{(1-h^2/r^2)r} \int_{0}^{h} e^{-\Gamma h} C_{0}(h) dh$$
(13)

where  $\Gamma = gM/RT$  is the average lapse rate of air density versus height and

$$C_{o}(h) = \frac{1}{2\pi r} \int_{0}^{2\pi} v_{r} r d\phi$$
 (14)

is the average radial velocity around the circle of measurement.

To fix the number of measurements at n, we assume M values of  $v_r$  are made on each of L circles spaced at intervals  $\Delta h$  from h=0 to h=h<sub>m</sub>. Then

$$\hat{\mathbf{C}}_{o}(\mathbf{h}) = \frac{1}{M} \sum_{m=1}^{M} \hat{\mathbf{v}}_{rm}$$
(15)

If all the radial velocities are independent with the same uncertainty, then

$$\operatorname{VAR}[\overline{\mathbf{w}}] = \frac{4 e^{2\Gamma h} h \operatorname{VAR}[\mathbf{v}_{r}]}{r^{2} (1 - h^{2} / r^{2})^{2} N} \frac{N}{n} \left\{ \frac{1 - e^{-2\Gamma h}}{2\Gamma} \right\}$$
(16)

For a direct measurement with a vertical beam  $VAR[\hat{w}_{o}] = VAR[v_{r}]/N$ . If we require that, for our maximum height  $h_{m}$ ,  $VAR[\overline{w}] = VAR[\hat{w}_{o}]$ , then the range can be found by solving

$$2\Gamma h_{\rm m} = \ell n \left\{ 1 + \frac{\Gamma h_{\rm m} r^2}{2h_{\rm m}^2} \left\{ 1 - \frac{h_{\rm m}^2}{r^2} \right\} \frac{n}{N} \right\}$$
(17)

Because of the accuracy needed for vertical velocity estimates, the number N of vertical data will be much larger than the n-N data for horizontal wind component estimation. Thus in (17) we can assume  $n/N\approx1$ . Using  $h_m = 10$  km,  $\Gamma = 0.113$  km<sup>-1</sup> in (17) gives  $r\approx40$  km, so  $\theta_{em}\approx14^\circ$ . For heights lower than  $h_m$ , VAR[ $\overline{w}$ ] is less than for a direct measurement.

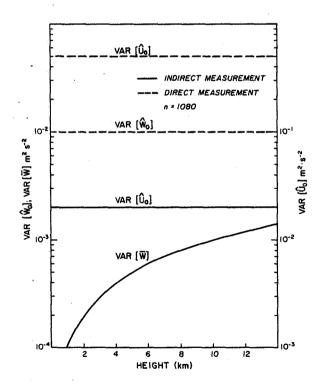
To compare the variances for direct and indirect measurements, assume that we have n/10 measurements at each level for estimating  $u_0$ ,  $v_0$ , and  $w_0$  at each of 10 levels spaced 1 km apart. Solving (16) assuming a specified VAR $[\overline{w}] = 10^{-3} m^2 s^{-2}$  at  $h_m = 10 \ \text{km}$  with VAR $[v_r] = 1 \ m^2 s^{-2}$  gives n = 1080. So there are 108 data at each level and, from (8) VAR $[u_0] \simeq 2 x 10^{-2} m^2 s^{-2}$ . With 108 data at each

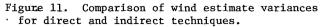
level for the direct VAD measurement techniques, the VAR  $[\hat{w}_0]$  is larger than that obtained using the indirect method even with  $\theta_e = 90^\circ$ . Because  $w_0$  needs to be estimated with better accuracy than  $u_0$ ,  $v_0$ , consider that most measurements are made with a vertical beam. We need at least 4 data spaced 90° apart on a circle at each level in order to remove the bias in  $\hat{u}_0$ ,  $\hat{v}_0$  due to  $u_x$ ,  $v_x$ (Equation 7). Thus from (8) we have VAR $[\hat{u}_0] \ge 0.5 \text{ m}^2 \text{ s}^{-2}$  which should be satisfactory for horizontal wind estimates but which is an order of magnitude larger than obtained by the indirect method. Furthermore, we have at most 104 data available at each level for the vertical beam with the consequence that VAR $[\hat{w}_0] \cong 10^{-2} \text{m}^2 \cdot \text{s}^{-2}$ , an order of magnitude or more (at h<10 km) larger than obtained using the indirect method. A comparison of these variances is shown in Figure 11.

In reality, the variances for the indirect measurement technique may not be this much smaller, since we have neglected the dependence of signal strength (and VAR[ $v_r$ ] with range. However, this analysis has shown that the indirect technique does not require a tradeoff between vertical and horizontal variance. It offers the advantages of low variances, an areal averaged vertical velocity, and requires no assumption about the spatial structure of the wind to measure vertical velocity.

## SUMMARY AND CONCLUSIONS

We have examined the errors in three radar techniques (three fixed beams, VAD, and VED) used to directly measure the three components of the wind. Equations were derived for the bias and variance of the uniform wind components estimates under the assumption of a spatially linear, time invariant wind field





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and a constant radial velocity measurement error. The measurement errors produce variance in the estimates and the linear wind shear biases the estimates. The variance of the estimates can be reduced by averaging more measurements but the biases cannot. Thus, for these direct measurement techniques, the selection of an elevation angle for simultaneous observation requires a compromise based on the required accuracy of the measurement.

We have also examined the errors for an indirect measurement technique based on Gauss's theorem with an equation of continuity constraint. Two advantages this indirect technique offers are that it does not require any assumptions about the spatial structure of the wind to measure vertical velocity and that its error variance can be smaller because it does not require a vertical and horizontal measurement variance compromise.

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## **REFERENCES**

Anton, H. (1981), Elementary Linear Algebra, John Wiley, 375 pp.

Doviak, R. J. and D. S. Zrnic' (1983), <u>Doppler Weather Radar and Weather</u> <u>Observations</u>, Academic Press, New York (in press).

Koscielny, A. J., R. J. Doviak and R. Rabin (1982), Statistical considerations in the estimation of divergence from single Doppler radar and application to pre-storm boundary layer observation, <u>J. Appl. Meteor., 21</u>, 197-210.

Peterson, V. L. and B. B. Balsley (1979), Clear air Doppler radar measurements of the vertical component of wind velocity in the troposphere and stratosphere, <u>Geophys. Res. Lett., 6</u>, 933-946.

Waldteufel, P. and H. Corbin (1979), On the analysis of single Doppler data, J. Appl. Meteor., 18, 532-542.

Zrnic', D. S. (1979), Estimation of spectral moments for weather echoes, <u>IEEE</u> <u>Trans. Geo. Elec., GE-17</u>, 113-128. For the fixed beam technique with one vertical, and two off-vertical beams at elevation  $\theta_e$ , azimuths 0° and 90°, we find, following the notations of KOSCIELNY et al. (1982), that

$$P_{nm} = \begin{bmatrix} \cos\theta_{e} & 0 & \sin\theta_{e} \\ (repeats (n-N/2 \text{ times})^{-1} \theta_{e} \\ 0 & \cos\theta_{e} & \sin\theta_{e} \\ (repeats (n-N)/2 \text{ times})^{-1} \theta_{e} \end{bmatrix} = P_{n3}$$

$$P_{nm}^{T}P_{nm} = \begin{bmatrix} \frac{(n-N)}{2} \cos^{2}\theta_{e} & 0 & \frac{(n-N)}{2} \cos\theta_{e} \cdot \sin\theta_{e} \\ -\frac{(n-N)}{2} \cos^{2}\theta_{e} & \frac{(n-N)}{2} \cos^{2}\theta_{e} \end{bmatrix} \equiv P_{3}^{T}P_{3}$$

Since this is a symmetric matrix, we have not entered the identical terms below the diagonal elements. We invert  $P_3^{TP_3}$  by performing a sequence of row operations to reduce it to the identity matrix. Performing this same sequence on the identity matrix reduces it to  $[P_3^{TP_3}]^{-1}$  (ANTON, 1981). Thus

$$(\mathbf{P}_{3}^{T}\mathbf{P}_{3})^{-1} = \begin{bmatrix} \frac{2\sec^{2}\theta_{e}}{n-N} + \frac{\tan^{2}\theta_{e}}{N} & \frac{\tan^{2}\theta_{e}}{N} & -\frac{\tan^{2}\theta_{e}}{N} \\ & \frac{2\sec^{2}\theta_{e}}{n-N} + \frac{\tan^{2}\theta_{e}}{N} - \frac{\tan^{2}\theta_{e}}{N} \\ & \frac{1}{N} \end{bmatrix}$$
(A1.1)

For an equal number of measurements on each beam, 3N=n and

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{3\sigma_{\varepsilon}^{2}}{n} (\sec^{2}\theta_{e} + \tan^{2}\theta_{e})$$
(A1.2)  
$$VAR(\hat{w}_{o}) = \frac{3\sigma_{\varepsilon}^{2}}{n}$$
(A1.3)

The biasing of the estimates by the derivatives of the linear wind can be computed from the alias matrix of (3). Since the analysis is for constant height, all vertical derivatives cause no bias. Thus, for equal numbers of measurements,

 $A_{34} = [P_{nm}^{T}P_{nm}]^{-1} P_{nm}^{T}P_{nl}$  and performing the multiplication gives

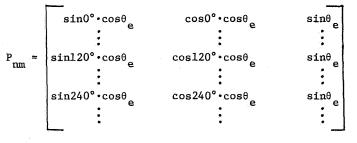
$$A_{34} = \begin{bmatrix} r \cos\theta_e & r \sin\theta_e \\ 0 & r \cos\theta_e & 0 & r \sin\theta_e \\ 0 & 0 & 0 & 0 \end{bmatrix}$$

Using the approximation  $h \approx r = \theta_{ei}$ 

$$A_{34} = \begin{bmatrix} h \cot \theta_{e} & 0 & h & 0 \\ 0 & h \cot \theta_{e} & 0 & h \\ 0 & 0 & 0 & 0 \end{bmatrix}$$
(A1.4)

and  $K_{\ell}^{T} = [u_x, v_y, w_x, w_y]$ .

For the fixed beam technique with three beams at elevation  $\theta_{\mbox{e}}$  and azimuths 0°, 120°, and 240°, we find that



so

and

$$(P_{nm}^{T}P_{nm}) = \begin{pmatrix} \frac{n \cos^{2}\theta_{e}}{2} & 0\\ \frac{n \cos^{2}\theta_{e}}{2} & \frac{n \cos^{2}\theta_{e}}{2} \end{pmatrix}$$

(A1.5)

 $\frac{2 \sec^2 \theta_e}{n} = 0$   $\frac{2 \sec^2 \theta_e}{n}$  $(P_{nm}^{T}P_{nm})^{-1} =$ 0 cscf0 n

0

0

n sin'

The alias matrix can be computed as before. For this case the predictor function for deformation is nonzero so

0					
	0	0	r cos <sup>2</sup> 0e	0	$r \sin\theta_e \cos\theta_e$
_	2	2 2	2 2	•	
P <sub>nl</sub> =				ar $\cos\theta$ sin $\theta$	
	$abr \cos^2 \theta_e$	a <sup>2</sup> r cos <sup>2</sup> 0 e	b <sup>2</sup> r cos <sup>2</sup> <sub>0</sub> e	-ar cosθe sinθ	$-br \sin\theta e \cos\theta e$
			•	÷	: _]

where  $a = \sqrt{3/2}$  and b = 1/2. Computing A as before,

$$A_{35} \approx \begin{bmatrix} \frac{h \cot \theta_{e}}{2} & 0 & 0 & h & 0 \\ 0 & \frac{-h \cot \theta_{e}}{2} & 0 & 0 & \frac{h}{3} \\ 0 & \frac{h \cot^{2} \theta_{e}}{2} & \frac{h \cot^{2} \theta_{e}}{6} & 0 & \frac{h \cot \theta}{3} \end{bmatrix}$$
(A1.6)  
and  $K_{g}^{T} = [u_{y} + v_{x}, u_{x}, v_{y}, w_{x}, w_{y}].$ 

# APPENDIX 2. ANALYSIS OF THE VAD TECHNIQUES.

For the VAD technique, n radial velocity data are collected on a circle at height h. So

$$P_{nm} = \begin{bmatrix} \cos\theta_{e} \cdot \sin\phi_{1} & \cos\theta_{e} \cdot \cos\phi_{1} & \sin\theta_{e} \\ \cos\theta_{e} \cdot \sin\phi_{2} & \cos\theta_{e} \cdot \cos\phi_{2} & \sin\theta_{e} \\ \vdots & \vdots & \vdots \\ \cos\theta_{e} \cdot \sin\phi_{n} & \cos\theta_{e} \cdot \cos\phi_{n} & \sin\theta_{e} \end{bmatrix}$$

and

$$P_{nm}^{T}P_{nm} = \begin{bmatrix} \Sigma \cos^{2}\theta_{e} \cdot \sin^{2}\phi_{i} & \Sigma \cos^{2}\theta_{e} \cdot \sin\phi_{i} & \cos\phi_{i} & \Sigma \cos\theta_{e} \cdot \sin\theta_{e} \cdot \sin\phi_{i} \\ & \Sigma \cos^{2}\theta_{e} \cdot \cos^{2}\phi_{i} & \Sigma \cos\theta_{e} \cdot \sin\theta_{e} \cdot \cos\phi_{i} \\ & & \Sigma \sin^{2}\theta_{e} \end{bmatrix}$$

where all summations are for i=1, 2,..., n. To simplify  $P_{nm}^{T}P_{nm}$ , we approximate the summations by integrals. For example,

$$\Sigma \cos^2 \theta_e \sin \phi_i \approx \cos^2 \theta_e \frac{n}{2\pi} \int_{-\pi}^{\pi} \sin^2 \phi d\phi = \frac{n}{2} \cos^2 \theta_e$$

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A similar evaluation of the remaining summations gives

$$P_{nm}^{T}P_{nm} = \frac{\frac{n \cos^{2}\theta_{e}}{2} \qquad 0 \qquad 0}{\frac{n \cos^{2}\theta_{e}}{2} \qquad 0}$$

$$n \sin^{2}\theta_{e}$$

Thus

.

$$(P_{nm}^{T}P_{nm})^{-1} = \begin{bmatrix} \frac{2 \sec^2 \theta_e}{n} & 0 & 0\\ & \frac{2 \sec^2 \theta_e}{n} & 0\\ & & \frac{2 \sec^2 \theta_e}{n} & 0\\ & & \frac{\csc^2 \theta_e}{n} \end{bmatrix}$$

and

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{h} 2 \sec^{2}\theta_{e}$$
(A2.1)  
$$VAR(\hat{w}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} \csc^{2}\theta_{e}$$
(A2.2)

To find the bias caused by neglecting the parameters of a linear wind field, 'e compute the alias matrix where  $K_5^T = [u_y + v_x, u_x, v_y, w_x, w_y]$ . For analysis t constant height,

	$\cos\theta e \cdot \cos\phi_1 \cdot \sin\phi_1$	$\cos\theta_{e} \cdot \sin\phi_{1}$	$\cos^{\theta} e^{\cos \phi} 1$	$\sin\theta e \sin\phi_1$	$\sin^{\theta} e \cos^{\phi} 1$
$nl = r \cos \theta_{e}$	$\cos\theta e \cos\phi_2 \sin\phi_2$	$\cos\theta_{e} \sin\phi_{2}$	coste cos¢2	$\sin^{\theta} e^{\sin^{\phi}} 2$	$\sin^{\theta} e^{\cos^{\phi} 2}$
		•	•	:	:
	$\cos\theta_e \cdot \cos\phi_n \cdot \sin\phi_n$	$e^{\cos\theta} = \sin \phi$	$\cos\theta e \cos\phi n$	$\sin\theta e \sin\phi n$	$\sin^{\theta} e^{\cdot \cos \phi} n$

erforming the matrix multiplications in (5), approximating summations by integrals, nd using  $h \approx r \sin \theta_e$  gives

$$A_{35} = \begin{bmatrix} 0 & 0 & 0 & h & 0 \\ 0 & 0 & 0 & 0 & h \\ 0 & \frac{h \cot^2 \theta_e}{2} & \frac{h \cot^2 \theta_e}{2} & 0 & 0 \end{bmatrix}$$
(A2.3)

# APPENDIX 3. ANALYSIS OF VED TECHNIQUE.

In the VED technique, the azimuth is fixed and the elevation angle is scanned. The range r is selected so the data are for a constant height h. To measure both horizontal components, another azimuth, preferably at 90° to the first, must be scanned. We assume  $\frac{n}{2}$  data are collected for each scan. For convenience, we introduce the horizontal distance s which is directed along the azimuth  $\phi$ . We make estimates of a horizontal component  $u_0$  (directed along s) and the vertical component  $w_0$ . The predictor function matrix is

$$P_{nm} = \begin{bmatrix} \cos\theta_1 & \sin\theta_1 \\ \cos\theta_2 & \sin\theta_2 \\ \vdots & \vdots \\ \cos\theta_n & \sin\theta_n \end{bmatrix}$$

so

$$P_{nm}^{T}P_{nm} = \begin{bmatrix} \Sigma \cos^{2}\theta_{i} & \Sigma \cos^{2}\theta_{i} \\ & \Sigma \sin^{2}\theta_{i} \end{bmatrix}$$

where the summations are for i=1, 2,...,  $\frac{n}{2}$  and  $\theta_i$  is the elevation angle. Approximating the summations by integrals,

$$\Sigma \cos^2 \theta_i \approx \frac{n}{2(\pi - 2\theta_0)} \int_{\theta_0}^{\pi - \theta_0} \cos^2 \theta d\theta = \frac{n}{4} \left[ \frac{(\pi - 2\theta_0) - \sin 2\theta_0}{\pi - 2\theta_0} \right]$$
  
$$\Sigma \sin \theta_i \cos \theta_i \approx 0$$

$$\Sigma \sin^2 \theta \approx \frac{n}{4} \left[ \frac{(\pi - 2\theta_0) + \sin 2\theta_0}{\pi - 2\theta_0} \right]$$

so

$$\left(P_{nm}^{T}P_{nm}\right)^{-1} = \begin{bmatrix} \frac{4}{n} \begin{bmatrix} \frac{(\pi - 2\theta_{o})}{(\pi - 2\theta_{o}) - \sin 2\theta_{o}} \end{bmatrix} & 0\\ 0 & \frac{4}{n} \begin{bmatrix} \frac{(\pi - 2\theta_{o})}{(\pi - 2\theta_{o}) + \sin 2\theta_{o}} \end{bmatrix} \end{bmatrix}$$

and

$$VAR(\hat{u}_{o}) = VAR(\hat{v}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} 4 \left[ \frac{\pi - 2\theta_{o}}{(\pi - 2\theta_{o}) - \sin\theta 2\theta_{o}} \right]$$

$$VAR(\hat{w}_{o}) = \frac{\sigma_{\varepsilon}^{2}}{n} 2 \left[ \frac{\pi - 2\theta_{o}}{(\pi - 2\theta_{o}) + \sin 2\theta_{o}} \right]$$
(A3.1)

The variance of w is halved since we assume the results from the two scans will be averaged.

The bias by the linear terms is again computed by the alias matrix. The prediction functions corresponding to the excluded parameters are

	$r \cos^{2}\theta_{1}$	$r \sin\theta_1 \cdot \cos\theta_1$
P <sub>nl</sub> =	$r \cos^{2}\theta_{2}$	$r \sin\theta_2 \cos\theta_2$
	r cos <sup>2</sup> θ <sub>n</sub>	$r \sin\theta_n \cos\theta_n$

Performing the matrix multiplications, approximating summation by integrals and using  $h\tilde{\sim}r$  sin0 gives

$$A_{22} = \begin{bmatrix} 0 & h \\ h \begin{bmatrix} (\pi - 2\theta_{o}) & -\sin 2\theta_{o} \\ (\pi - 2\theta_{o}) & +\sin 2\theta_{o} \end{bmatrix} & 0 \end{bmatrix}$$

For the combined analysis of two scans at 0° and 90°,

$$A_{34} = \begin{bmatrix} 0 & 0 & h & 0 \\ 0 & 0 & 0 & h \\ h & \frac{(\pi - 2\theta_0) - \sin 2\theta_0}{(\pi - 2\theta_0) + \sin 2\theta_0} \end{bmatrix} \begin{bmatrix} \frac{(\pi - 2\theta_0) - \sin 2\theta_0}{(\pi - 2\theta_0) + \sin 2\theta_0} \end{bmatrix} 0$$

and the vector of excluded parameters is

-

$$K_4^{T} = (u_x, v_y, w_x, w_y).$$