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AURORAL THERMOSPHERE TEMPERATURES FROM  
OBSERVATIONS OF 6300 Å EMISSIONS

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Science and Engineering Directorate

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TABLE OF CONTENTS

	Page
I. INTRODUCTION .....	1
II. INSTRUMENTATION .....	1
III. OBSERVATIONS .....	2
IV. DATA REDUCTION .....	3
V. RESULTS AND DISCUSSION .....	6
A. MSIS .....	7
B. SHADOW HEIGHT .....	7
C. VOLUME EMISSION RATE PROFILES PREVIOUSLY MEASURED .....	8
VI. FUTURE WORK .....	10
VII. CONCLUSION .....	10
REFERENCES .....	12
APPENDICES	
A. THEORETICAL GAUSSIAN PROFILES .....	27
B. SOLAR GEOPHYSICAL DATA .....	30
C. MSIS EXAMPLE OUTPUT .....	31
D. PROGRAM AND EXAMPLE OUTPUT .....	32
E. SHADOW HEIGHT .....	49
F. EXAMPLE OF RAW DATA .....	59
G. VERIFICATION OF EQUATION (A-1), APPENDIX A .....	60

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## LIST OF ILLUSTRATIONS

Figure	Title	Page
1.	Fabry-Perot Interferometer Control Schematic . . . . .	14
2.	Fabry-Perot Interferometer Piston Diagram . . . . .	15
3.	Fabry-Perot Interferometer Controller and Data Acquisition and Display System Schematic . . . . .	16
4.	Optimization of Number of Fourier Coefficients . . . . .	17
5.	Comparison of Theoretical Gaussian Profiles to Smoothed Data . . . . .	18
6.	Comparisons of Observed Temperatures with MSIS Temperatures, Plotted Versus Universal Time. (a) March 7, 1984, (b) March 10, (c) March 11, (d) March 12, (e) March 13, (f) March 14, (g) March 15 . . . . .	19
7.	Shadow Height as a Function of Local Time in Different Look Directions for Days Encompassing the Period of the Observations . . . . .	26
E-1.	Local Shadow Height Geometry . . . . .	57
E-2.	Global Geometric Parameters Used in Shadow Height Program . . . . .	58

## I. INTRODUCTION

This report summarizes the results of a set of observations of the atomic oxygen OI 15,867 K 6300 Å  $^3P-^1D$  thermospheric emission to determine the temperature of the emitting species using a Fabry-Perot interferometer at the University of Alaska Geophysical Institute (64.86° latitude, -147.85° longitude) during March 1984. Spectral profiles obtained from the interferometer are used to determine the Doppler temperature by means of the technique reported by Hays and Roble [1] and Roble [2].

This general concept of using Doppler widths of airglow lines to find temperatures has been used by many investigators such as Wark [3], Nilson and Shepherd [4], Turgeon et al. [5], Zwick and Shepherd [6], Hernandez [7], Hernandez and Roble [8-11], and Smith et al. [12]. Analytical descriptions of Fabry-Perot spectrometers have been presented by Born and Wolfe [13] and Hernandez [14]. A review of temperatures and winds measured by Fabry-Perot spectrometry was done by Hernandez [7].

## II. INSTRUMENTATION

The Fabry-Perot interferometer was located on the top floor of the Geophysical Institute, Fairbanks, Alaska, allowing observations through a variable geometry periscope system and a Plexiglass dome. The interferometer is described by Roble [2] and Sivjee et al. [15] and is shown in Figures 1 and 3. Spectral profiles were obtained by pressure scanning of the etalon in steps. A stepper motor controlled the pistor (shown in Figure 2), and position data were provided by an A/D encoder. During a pressure scan, the refractive index in the etalon changed. This caused the light in the center of the interferometer pattern to change intensity, creating an intensity/pressure fringe. This light from the interferometer was reflected through a 1/8-inch aperture, then filtered, and finally detected by a photomultiplier tube. Simultaneously, a photometer recorded the intensity of the 6300 Å lines in the same direction as the interferometer observation. Also, the etalon temperature was monitored. Pressure scanning, mirror positioning, and integration times were controlled by computer. Times, pressures, observation directions,

temperatures, and counts were automatically stored in the format shown in Appendix F.

An He-Ne ( $6328 \text{ \AA}$ ) frequency-stabilized Tropel laser was used for calibration. To prevent photomultiplier tube saturation, the beam was incident on a ground glass sphere above the interferometer. The glowing sphere simulated the sky, so pressure scanning produced fringes without changing the experimental setup from the sky observation configuration. If, during calibration, the fringes were found to be asymmetric, the aperture or, in extreme cases, the etalon plates were adjusted. Neither of these adjustments were required between the observations included in this report, although such adjustments were required between earlier trial runs.

### III. OBSERVATIONS

Observations were made at an inclination of  $30^\circ$  from the horizontal in the north, south, east, and west magnetic directions. Observations were also made at local zenith. All observations were made between sunset and sunrise from March 6 to March 16, 1984. At the beginning of this period the moon was only 9% full; but by March 16, it was 97% full.

Most nights were clear, and slight to moderate auroral activity was observed in the zenith and to the north. On clear nights, diffuse aurorae were usually seen to the north. Data were also taken on cloudy and foggy nights, even though the observed light was scattered from all sky sources prior to detection. If there is mass average velocity with the emitting atoms, the temperature measurement is still correct. Winds could cause an error if they were unusually strong under these cloudy conditions.

In addition, the raw data were monitored during all runs. For example, an X-Y plotter was used to record the intensity of the interferometer image as a function of pressure to monitor the fringes and to provide a quick look at the raw data. After each night of observations, the records stored in the computer were transferred to disks. The parameters that were monitored and the formatting of this data are discussed in the following section.

#### IV. DATA REDUCTION

The computer program for data analysis included: reading of data from disk, temperature and pressure compensations, normalization, Fourier transform smoothing, deconvolution, and least squares fitting to theoretical profiles. A listing of this program is included in Appendix D along with a sample output.

##### A. Data Format, Reading Data

Each pressure scan of 64 steps included one free spectral range or fringe, which covered only 34 steps. The other 30 steps corresponded to other fringes and were, therefore, ignored. These fringes occurred as variations in photomultiplier counts as functions of pressure. Data at each step included etalon pressure and temperature, photometer reading, and photomultiplier reading. Times were recorded at the beginning of each pressure scan. All these data were stored in files such as the one shown in Appendix F. The first step in the data analysis program was to read the data corresponding to one fringe.

##### B. Data Modification

First, the reading of the etalon pressure was adjusted to compensate for variations of etalon temperature. Then the photomultiplier counts (i.e., intensity) were adjusted to account for temporal intensity variations of the sky by dividing by the photometer reading, which was recorded in kilorayleighs. Finally, the background signal was subtracted.

##### C. Fourier Cosine Transform Smoothing

The Fabry-Perot interferometer produces a series of concentric rings. The center of this image is monitored by a photomultiplier tube and as the etalon pressure is changed, the measured intensity changes. A series of 34 different pressure readings were taken in each scan. The resulting observed profile is a convolution of the actual sky profile and the

instrument function, with noise superimposed. To extract this noise a Fourier technique was used.

A Fourier cosine transform is fit to the data from one fringe. When the first six coefficients are used to reconstruct the curve, the result is a smoothed version of the data as shown in the equations below. If not enough coefficients are taken, the data are not accurately represented. On the other hand, if too many coefficients are taken, the resulting curve fits statistical noise as well as the source variation.

For a theoretical emission profile which is perfectly smooth, the coefficients decrease as the wavenumber (i.e.,  $m$  in equation (1)) increases, as shown by Hays and Roble [1]. For our data, the coefficients decrease until  $m$  is 6, whereupon it begins to increase, as shown in Figure 4. Therefore, to represent the actual data with a curve that is as accurate as possible with minimum noise, the optimum number of coefficients is 6. Each of the three curves shown in Figure 4 are from different pressure scans, all taken successively; therefore, each curve corresponds to a different fringe.

The smoothed profiles of the observed data are given by:

$$Y = \sum_{m=1}^6 Y_m \cos (mx), \quad (1)$$

where the  $Y_m$ 's are the Fourier coefficients from fits to the observed data. These coefficients are found using:

$$Y_m = 1/\pi \sqrt{Y_{cm}^2 + Y_{sm}^2},$$

where

$$Y_{cm} = \sum_{i=1}^{34} \frac{C}{m} [2 \sin (c2 P_1) (\sin c2 \frac{DPS}{2})^2 + \cos (c2 P_1) \sin (c2 DPS)]$$



where

$C$  = counts at  $i$ th step

$P_i$  = pressure at  $i$ th step

$$c2 = \frac{2\pi \times m}{SRS}$$

SRS = free spectral range, pressure units

$$DPS = P_i - P_{i-1};$$

and similarly,

$$Y_{sm} = \sum_{i=1}^M \frac{C}{m} [2 \cos (c2 P_i) (\sin c2 \frac{DPS}{2})^2 - \sin (c2 P_i) \sin (c2 DPS)] .$$

#### D. Deconvolution

The smoothed profile is assumed to be a convolution of the theoretical Doppler profile of the sky emission ( $D$ ) and the instrument profile ( $L$ ) (obtained by observing through the instrument a diffuse illumination provided by an He-Ne laser); i.e.,

$$Y_m = D_m * L_m .$$

To deconvolve the laser profile from the observed data, the coefficients from the observed data were divided by the laser coefficients using the convolution theorem (e.g., see Reference 16); i.e.,

$$D_m = Y_m / L_m ,$$

where the laser coefficients are found by running a slightly modified version of the program (in Appendix D) independently for the laser fringe coefficients only.

#### E. Least Squares Fit

After fitting the data to a smoothed profile using the Fourier coefficients and deconvolving the instrument function, the profile is then compared to a series of theoretical Gaussian emission profiles. These theoretical profiles are discussed in Appendix A. They have a known Doppler width for any given temperature. These theoretical profiles are convolved with a Gaussian instrument function obtained by calibrating with an He-Ne laser (as in equation (2)). The theoretical curve that gives the least square error with the smoothed data was used to obtain the temperature.

To ensure that the deconvolved profiles are Gaussian, they are plotted with theoretical Gaussians as shown in Figure 5. Non-Gaussian results were not included in temperature plots. Also, the half widths of all deconvolved profiles have been compared to those in Table A-1, Appendix A to ensure that no major errors occurred in the deconvolving calculation.

#### V. RESULTS AND DISCUSSION

After converting the fringes to temperatures, the temperatures were plotted as a function of time for each night of observation, as shown in Figures 6a-g. Each direction of observation: zenith, north, south, east, and west (all magnetic) were included. The times given are in Universal Time (UT), where local time was 9 hours behind UT.

It is apparent from the results in Figures 6a-g that the data points are scattered, but are typically within a band that is  $\pm 100$  K about the mean. The data for March 7 (Figure 6a) are highly scattered, but on March 10, the temperature is seen to be about  $800 \pm 100$  K. The next three nights give temperatures of 750, 750, and 800 K. All the above data were taken on clear nights. March 14 was cloudy and the apparent temperature was found to be 750 K. March 15 was foggy and the apparent temperature was found to

be 700 K. On March 16, only four zenith points were taken and three of these were at about 680 K.

Contributions to the spread in the data may be from both the instrument and actual variations in OI emission intensity. It is difficult to determine the fractional contributions of these components, although it has been considered by Hays and Roble [1].

#### A. MSIS

Also plotted in Figures 6a-g are the temperature profiles derived from the MSIS (Mass Spectrometer Incoherent Scatter) model of the neutral atmosphere [17-18]. Parameters utilized in this model include local time, altitude, location, 3-month average  $F_{10.7}$  flux, daily  $F_{10.7}$  flux, and daily  $A_p$  value (see Appendix B). An example of the MSIS output is given in Appendix C.

The purpose of running MSIS was to deduce the predicted model thermospheric temperature for the geomagnetic conditions at the observing location. Comparisons of the MSIS temperatures with our results show that ours are in a reasonable range.

TABLE 1. PLOT SUMMARY

Date (UT)	MSIS Height (km)	Temp. (K)	Sky
March 7		high dispersed	clear
March 10	180	800±100	clear
March 11	170	750±100	clear
March 12	170	750±100	clear
March 13	170	800±100	clear
March 14	170	750±100	cloud
March 15	150	700±100	fog
March 16	minimal data	680	clear

#### B. Shadow Height

In order to determine whether the observed volumes were sunlit, the Earth's "shadow height" was calculated. Heights for zenith, north, south, east, and west (magnetic) observations were calculated. It was determined

that the observed volumes in some of the observations (taken shortly after sundown and shortly before sunrise) were sunlit. Plotted in Figures 6a-g are heights of the Earth's shadow. This shadow height is the distance from the point where the line of observation intersects the surface between sunlight and shadow to the ground directly below. (See Appendix E for a detailed explanation of how shadow height is determined.)

Each curve corresponds to a given direction of observation (e.g., east). Where the curve intersects an MSIS altitude curve, the corresponding time indicates when the shadow height was at the altitude of the MSIS curve. For example, in Figure 6a, the vertical curve shown corresponds to observation toward the east. This curve intersects the MSIS 170 km curve at 1400 UT, indicating the shadow height at that time. As shown, the shadow height passed through lower altitudes at later times. The overall curve indicates sunrise, because observation points to the left of the curve are in darkness, while observation points to the right, i.e., later in time, are sunlit. In the following figures, different directions of observation are included. N represents north, while W and E represent west and east, respectively. Also shown in the figures are the curves for sunset. Of all directions observed, points to the west were the last to remain sunlit on any given night, and points to the east were the first to become sunlit.

For sunrise observations, the lowest shadow heights were to the east. It can be seen that on March 11 and 13, observations to the east, an hour after the emission region passed into sunlight, indicate pre-dawn temperature enhancement. After sunset, the lowest shadow heights were to the west, but sunlit dusk observations were made only on March 14, which was a cloudy night.

Shadow height calculations were based on Chamberlain's work [19]. Figure 7 shows the resulting shadow heights as a function of time for all directions observed. Programs in Appendix E were used to obtain these results.

### C. Volume Emission Rate Profiles Previously Measured

Altitude profiles of the  $6300 \text{ \AA}$  volume emission rate have been reported by various investigators using data from the Atmosphere Explorer

program. Selected results are noted here for comparison with the results presented in the previous section. For example, Abreau et al. [20] found altitude profiles of the volume emission rate using a photometer onboard the AE-E satellite. From November 1980 to February 1981, averaged data gave a peak rate at altitudes of 250-280 km.

Hays et al. [21] reported altitude profiles of the 6300 Å volume emission rate for various solar zenith angles during evening twilight. Peak emissions were in the 220-250 km altitude region. Their theoretical models using O<sub>2</sub> photodissociation and photoelectron impact indicated that below 200 km, the dominant mechanism of excitation was photodissociation.

Mid-latitude orbit data for summer and winter were examined by Torr et al. [22]. Theoretical and measured peak emission altitudes were approximately 200 km in winter and 200 km in summer.

Hernandez and Roble [10] obtained temperatures from the 6300 Å emission as a function of time. Their temperatures were found to be greater than MSIS model predictions for an altitude of 250 km during summer and equinox nights. Their data for March 1976 (the same month as our data) were scattered to about the same degree as our data, and their data were centered on 800 K in agreement with our results. Further results were later reported by Hernandez and Roble [11]. They found that the 6300 Å line peak emission rate occurred below the F<sub>2</sub> peak at 250 km, at about 100 R, decreasing throughout the night.

Recently, Sipler et al. [23] also measured the neutral F-region temperature using a Fabry-Perot interferometer. During geomagnetically quiet nights from 1975 to 1979, equinox solar minimum average temperatures were found to be about 750 K.

Emission rate profiles for 6300 dayglow were reported by Killeen et al. [24]. They reported that a typical profile had a peak of 180 cm<sup>-3</sup> s<sup>-1</sup> at about 210 km altitude.

The above results indicate a peak volume emission rate at altitudes of 200-300 km. As discussed earlier, these results are in rough agreement with the present observations, given the spread in our temperatures and the combined uncertainties of MSIS temperatures, as well as our lack of statistical data. Assuming the glow intensity is originating from 200 km, our data suggest the MSIS model is predicting slightly higher temperatures than observed.

## VI. FUTURE WORK

Many improvements could be made to the apparatus. With CCD arrays now readily available, it is possible to image the Fabry-Perot fringes. Imaging of fringes has been done by Sivjee et al. [15] and Rees and Greenaway [25]. The latter investigators developed a Doppler imaging system (DIS) which used a Fabry-Perot interferometer and a 120° field of view all-sky camera. The multiring image contains spectral and spatial information. By imaging the fringe from the interferometer (of this investigation), all the information could be digitized and analyzed by computer. By using the entire image of the fringe rather than just the central fringe, a much stronger signal is obtained. This system would provide a convenient means of monitoring winds using the Doppler shift. Yet another advantage of this system would be that fringes would be made much more quickly than with pressure scanning. Further, integration times could be easily varied to an optimum duration. This system could be further enhanced by using a pyramid with mirror surfaces rather than a scanning mirror. This would allow monitoring of all four directions simultaneously. With both of the above modifications, all moving parts would be eliminated, greatly simplifying the entire system and reducing its size.

An extension of the pyramid system would be to use an all-sky lens so that all directions could be monitored simultaneously. This lens could be coupled to an interference filter, an image intensifier, and finally to a CCD array via a fiber optic plug. In this arrangement, the apparatus could be considerably reduced in size compared with the current system. However, as size decreases, integration time increases. An image intensifier could be used to reduce integration time.

Cooling of the array would be imperative to reduce noise accumulation. A Peltier electric cooler, for example, would be convenient for this task.

## VII. CONCLUSION

The temperatures of the 6300 line on the clear nights of March 10, 11, 12, and 13 were 800, 750, 750, and 800 K, respectively, with a spread of

$\pm 100$  K. Despite the large spread in the data, the temperatures and the spread in the temperatures were consistent from day to day, and consistent with previous observations.

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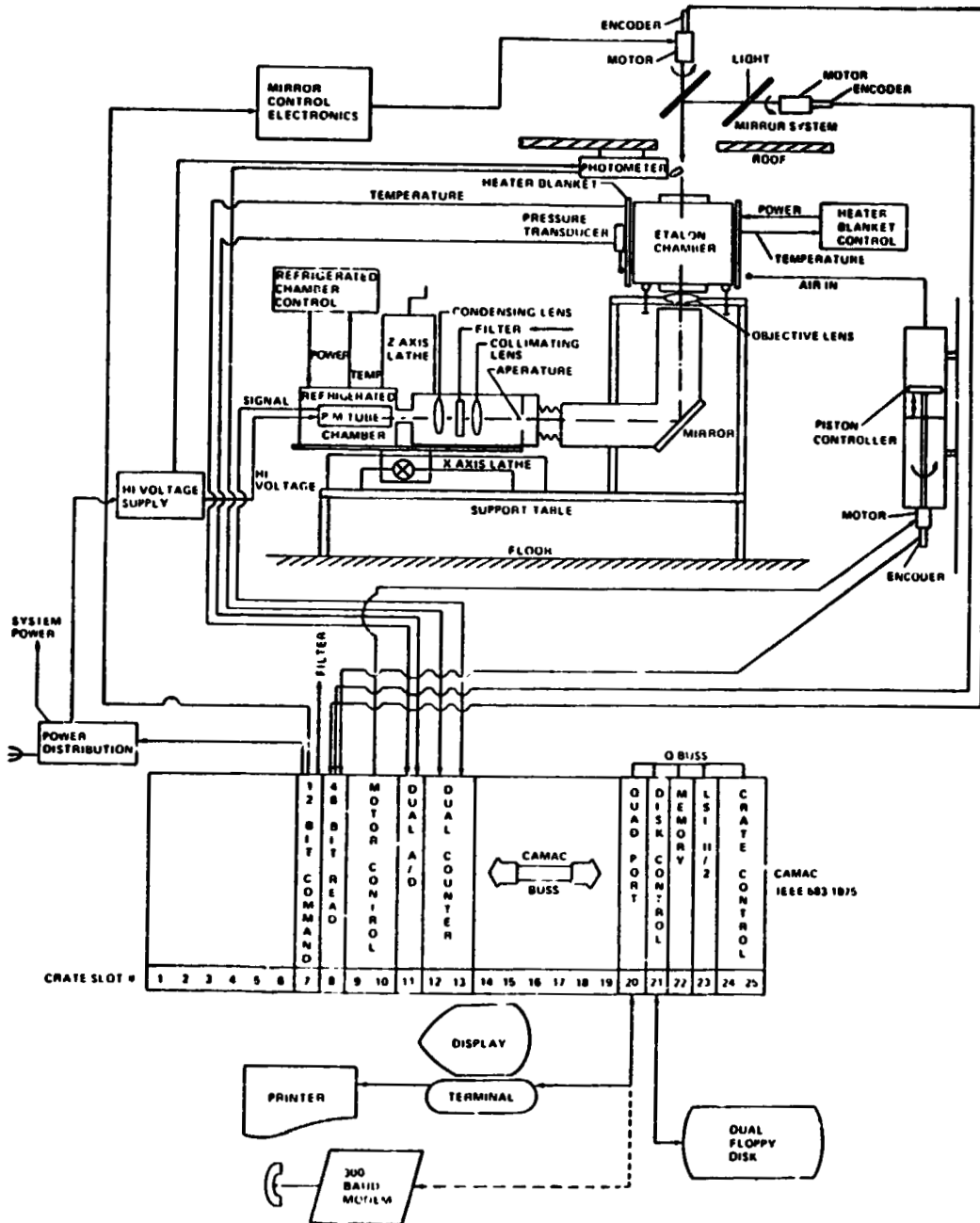


Figure 1. Fabry-Perot Interferometer control schematic.

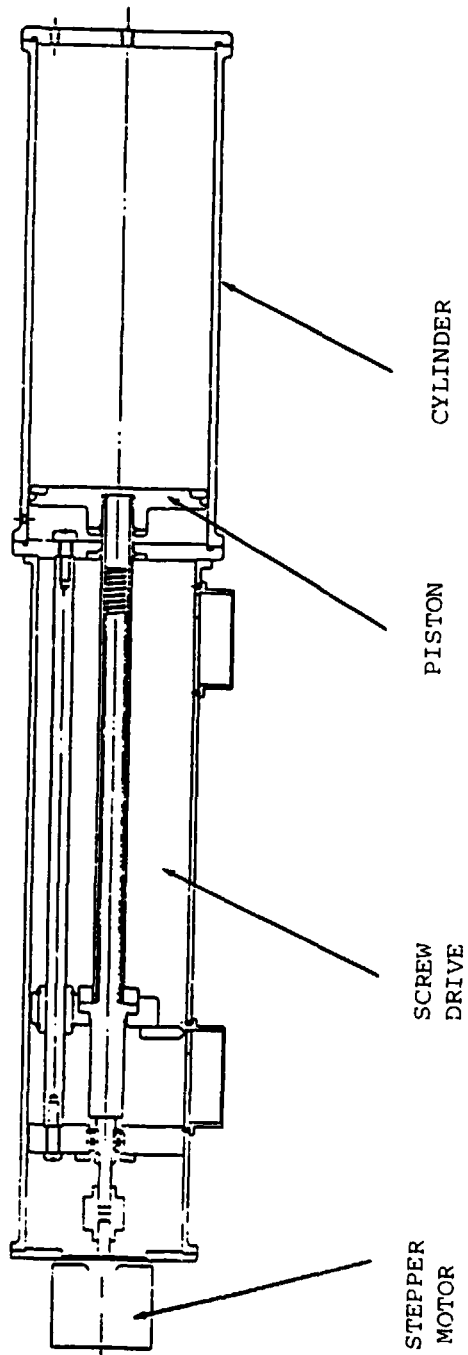


Figure 2. Fabry-Perot Interferometer piston diagram.

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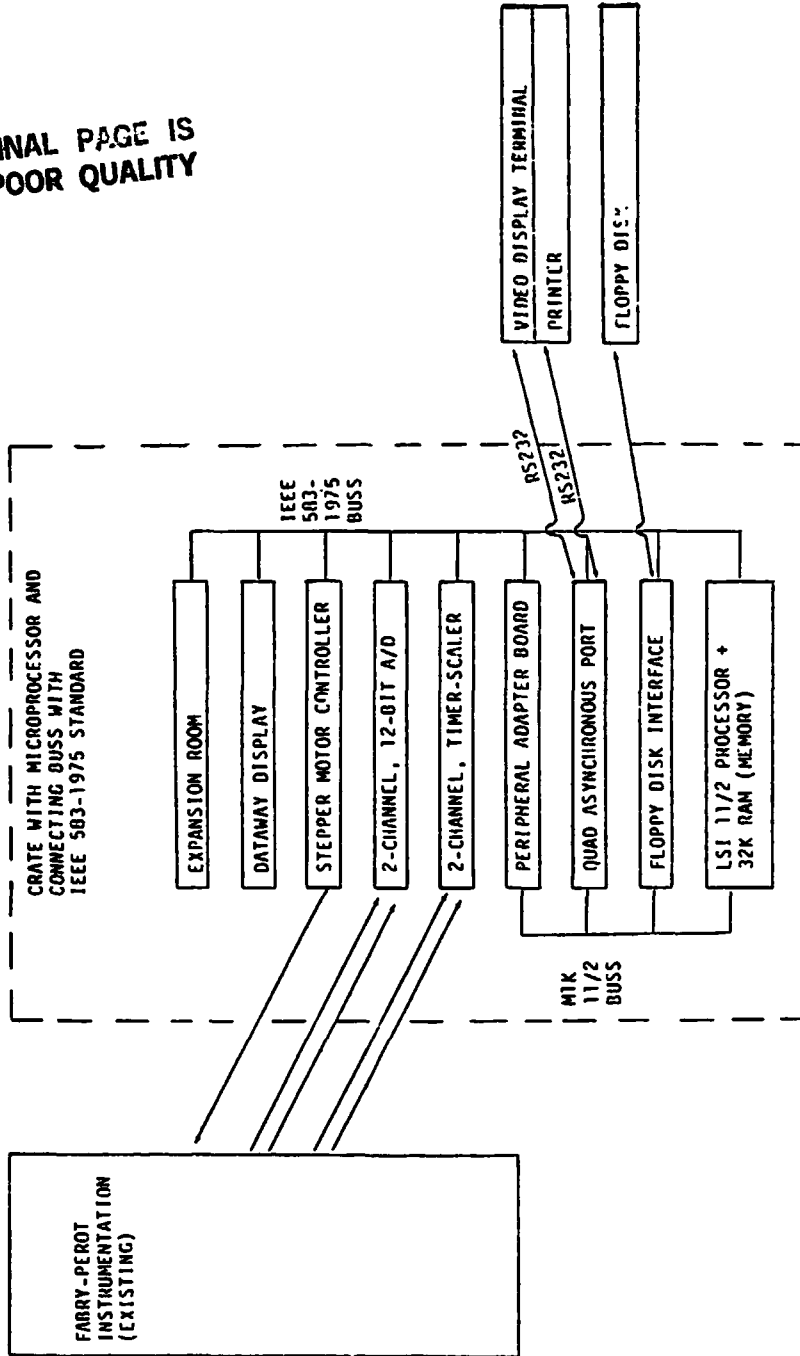


Figure 3. Fabry-Perot Interferometer controller and data acquisition and display system schematic.

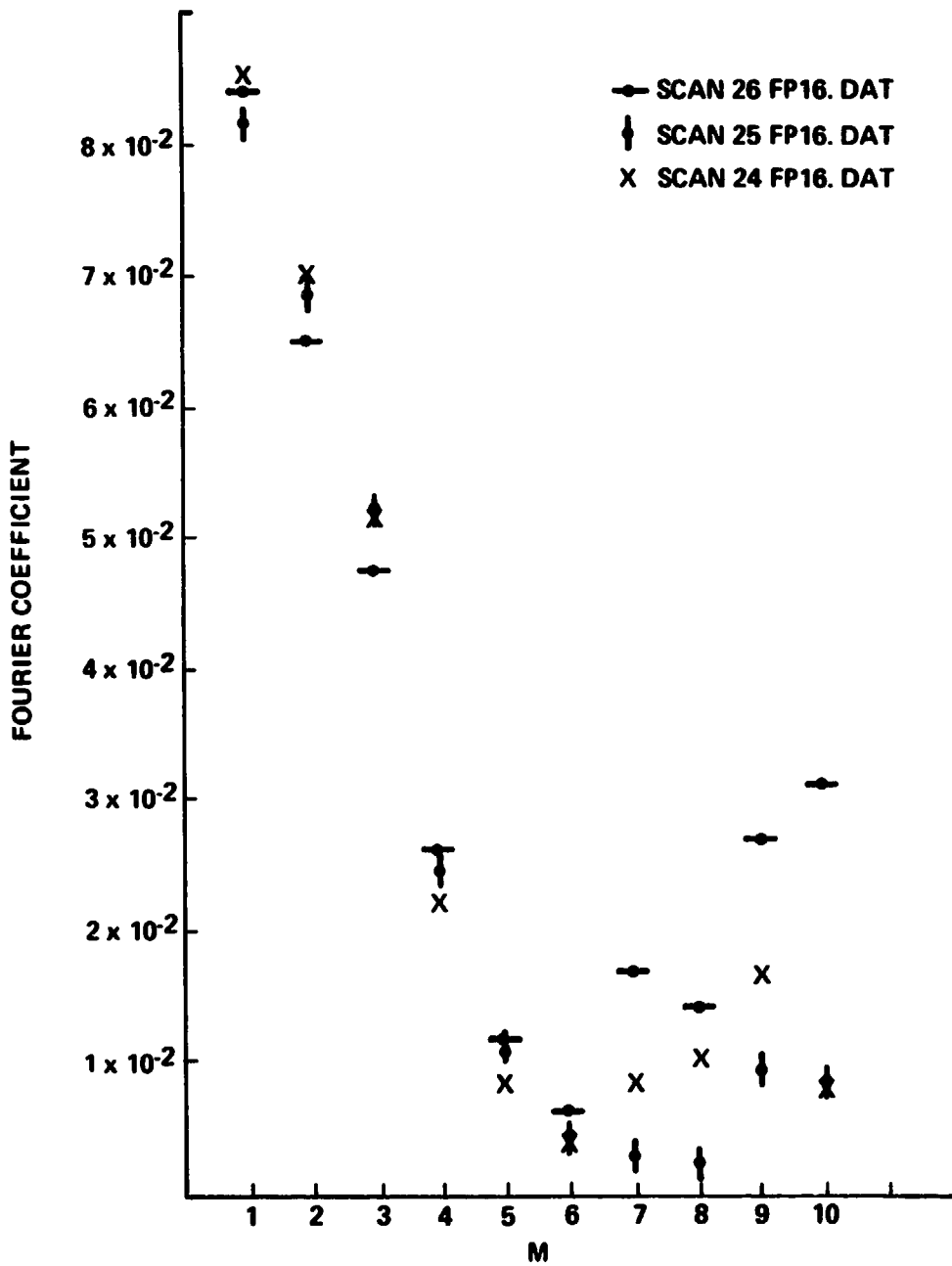


Figure 4. Optimization of number of Fourier coefficients.

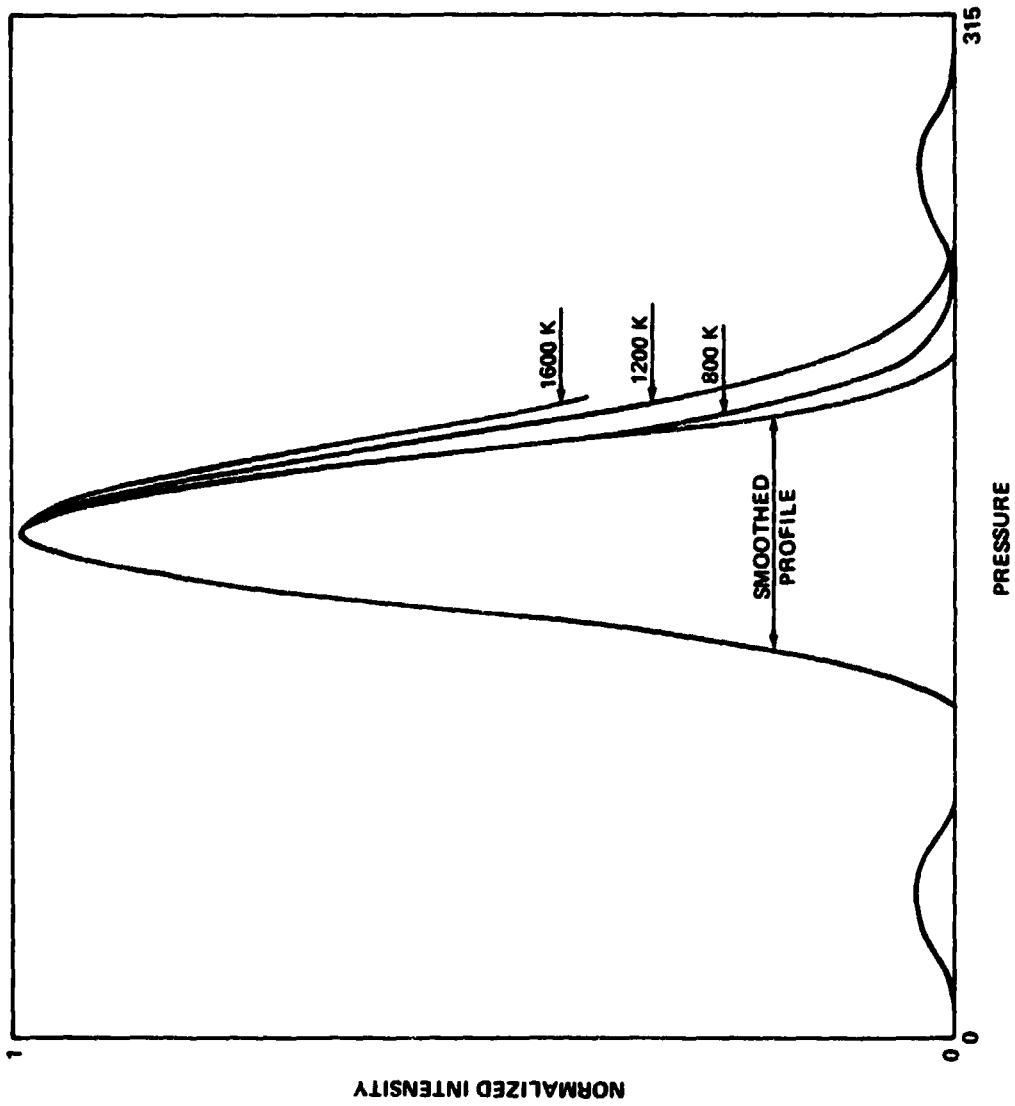
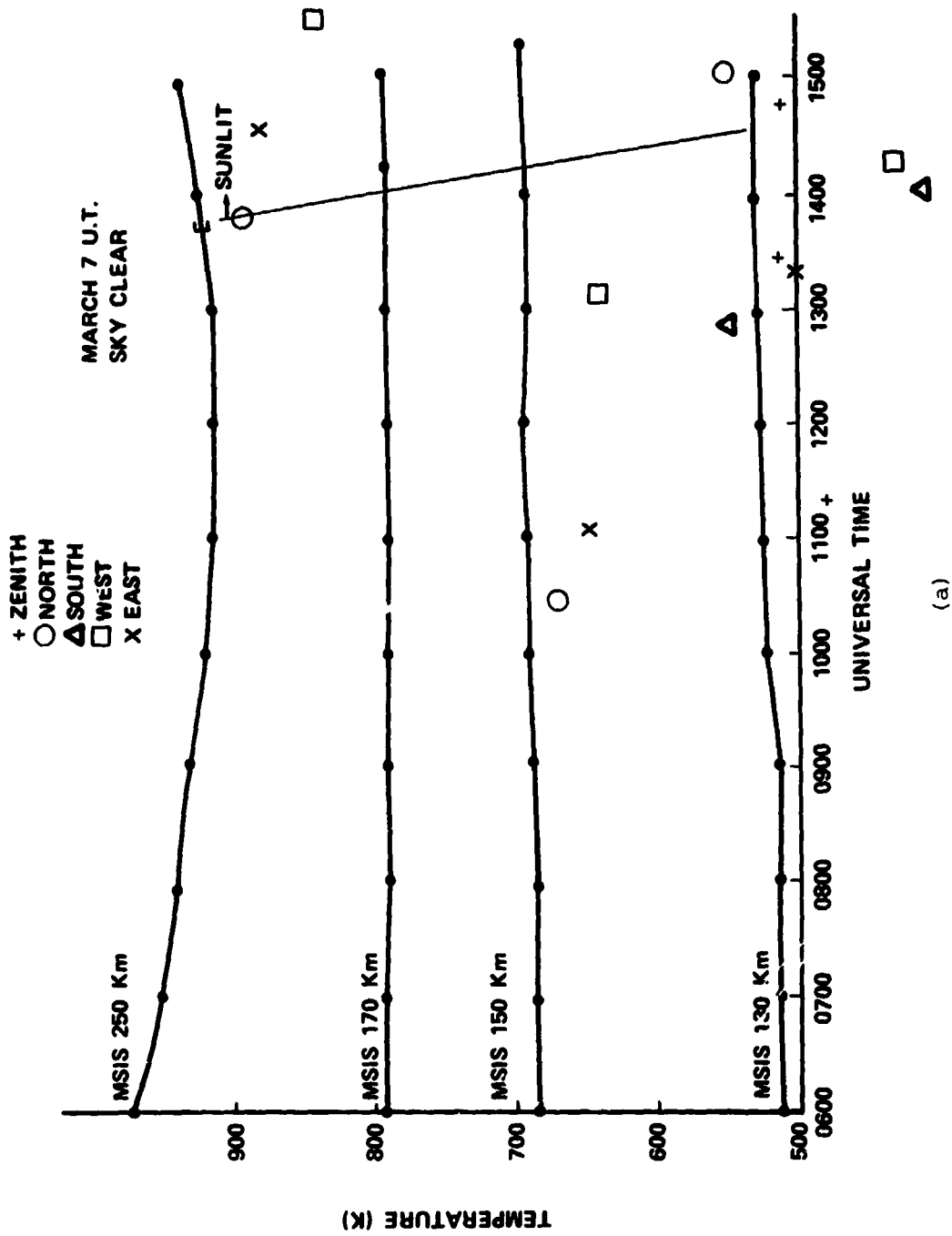


Figure 5. Comparison of theoretical Gaussian profiles to smoothed data.



(a)

Figure 6. Comparisons of observed temperatures with MSIS temperatures, plotted versus Universal Time.  
 (a) March 7, 1984, (b) March 10, (c) March 11, (d) March 12, (e) March 13, (f) March 14,  
 (g) March 15.

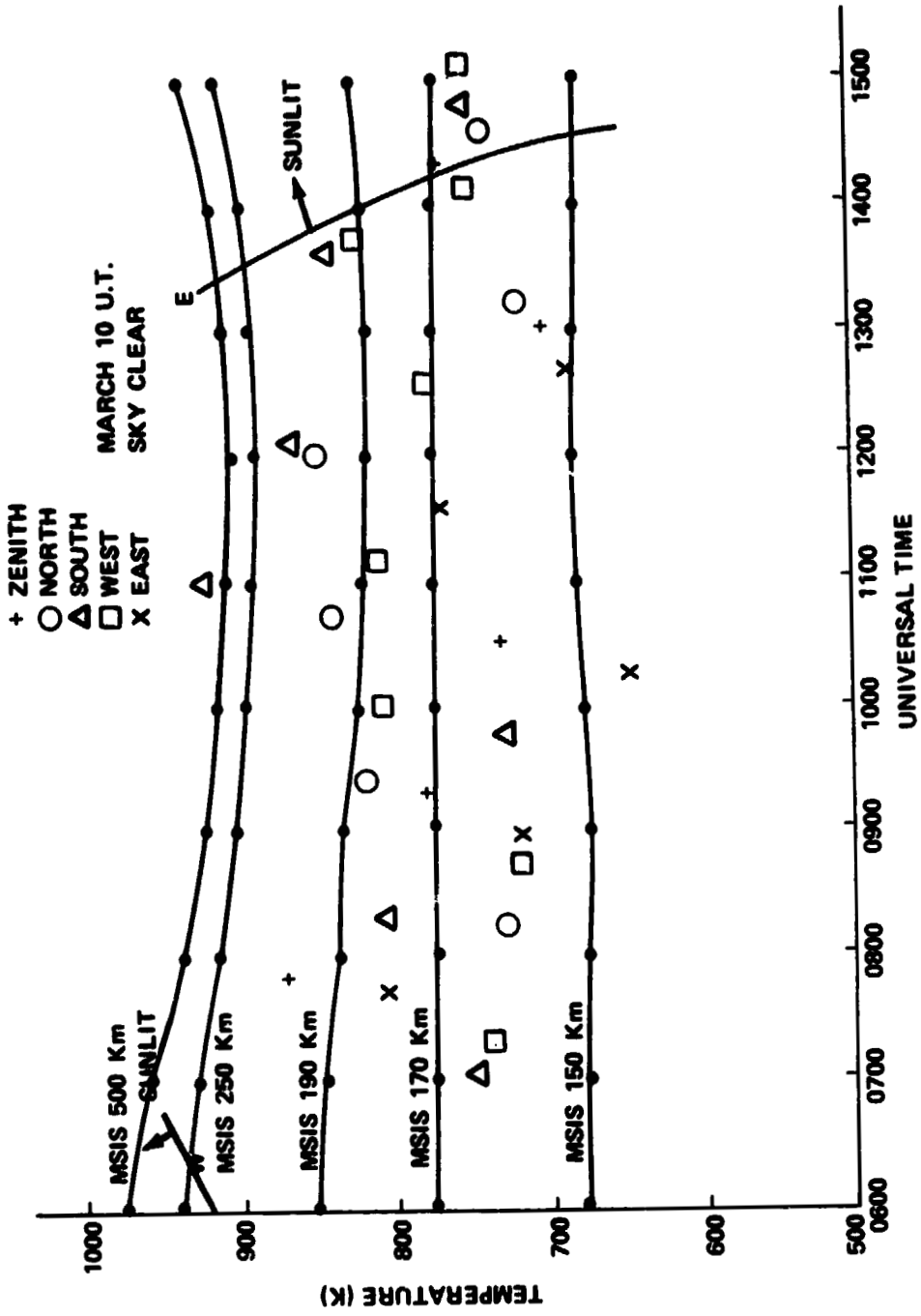


Figure 6b. March 10



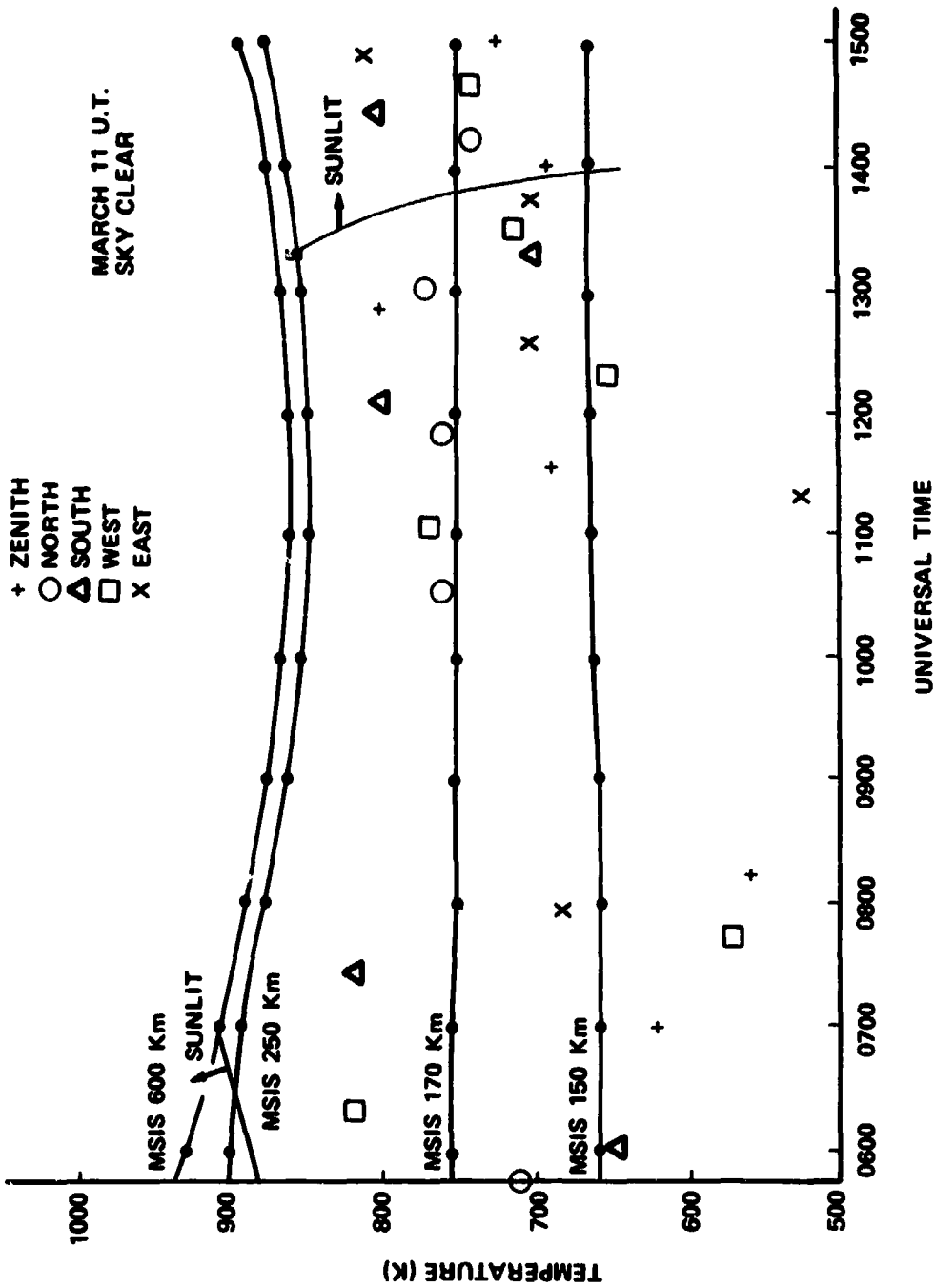


Figure 6c. March 11.

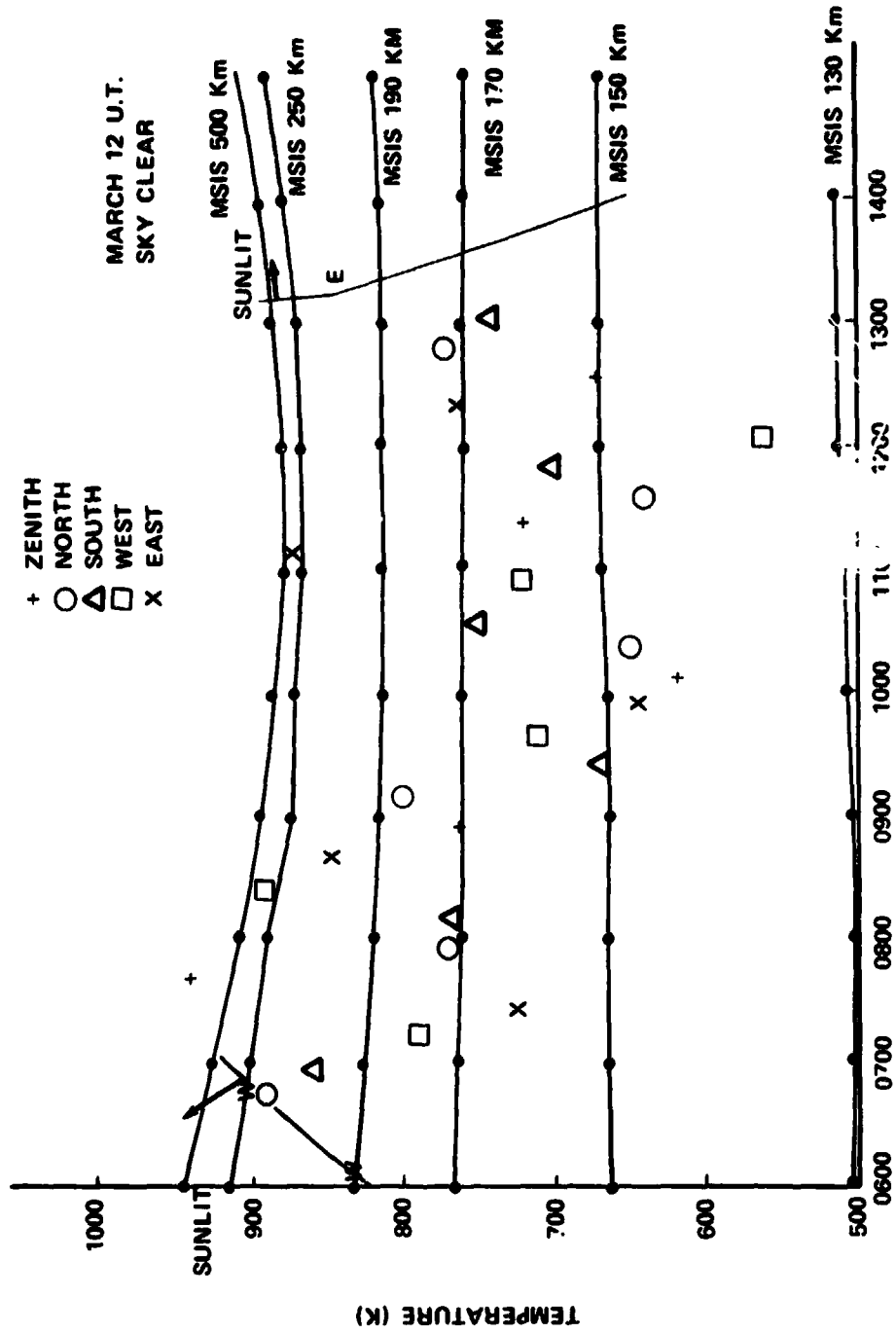


Figure 6d. March 12.

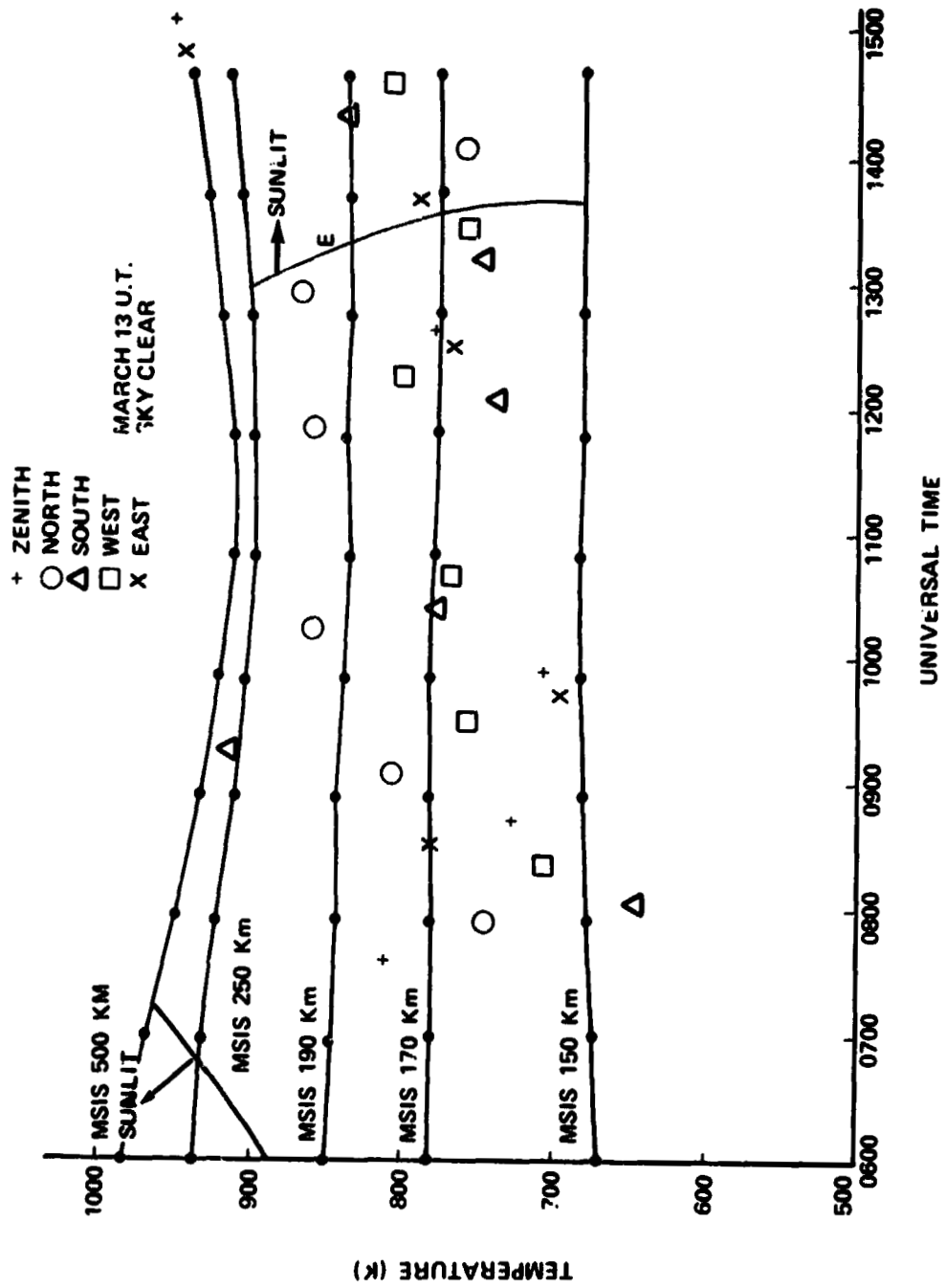


Figure 6e. March 13.

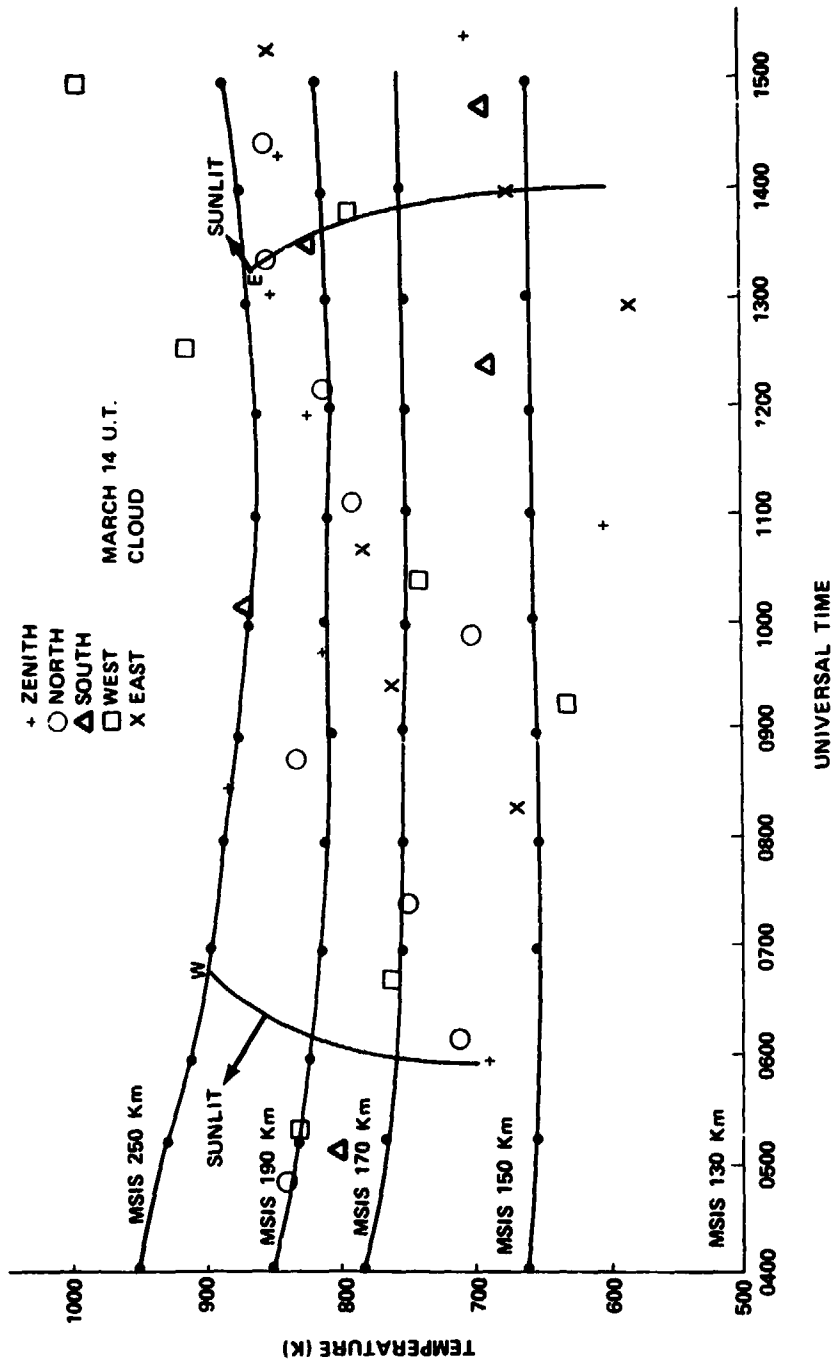


Figure 6f. March 14.

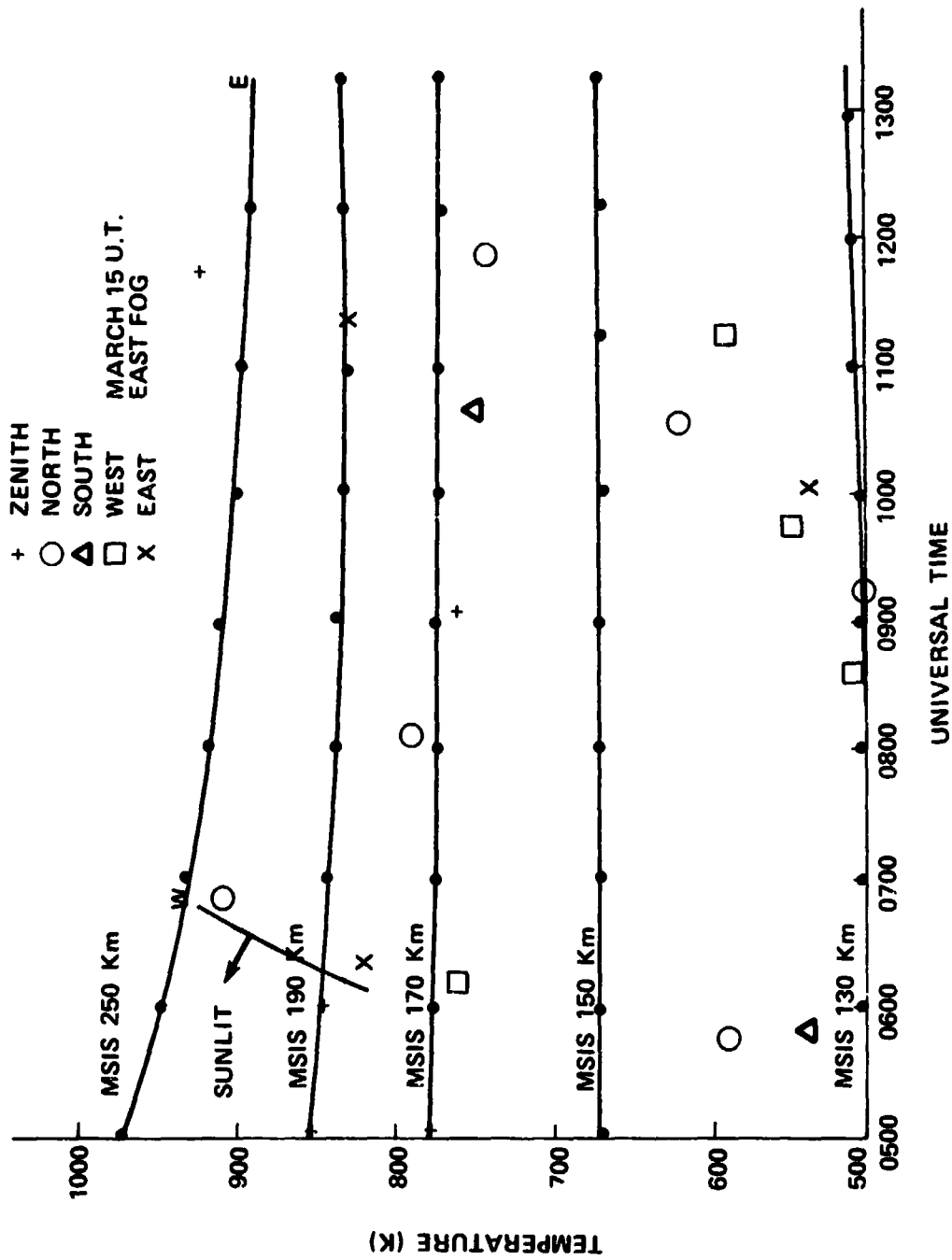


Figure 6g. March 15.

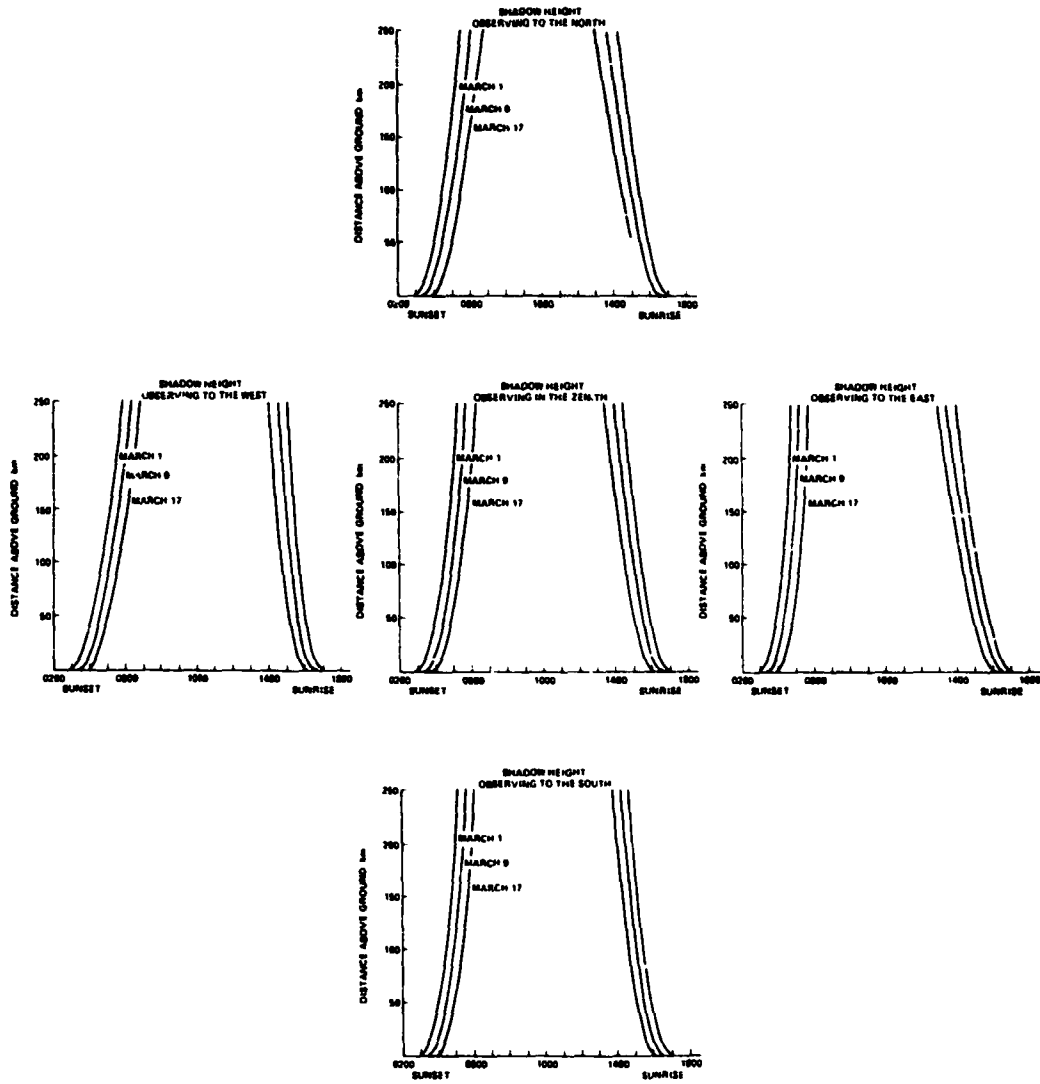


Figure 7. Shadow height as a function of Local Time in different look directions for days encompassing the period of the observations.

## APPENDIX A. THEORETICAL GAUSSIAN PROFILES

A Doppler shift of wavelength results from the motion of a radiating particle towards or away from an observer. In gas or plasmas, the random thermal motions of all particles lead to a Maxwellian velocity distribution. This results in a Gaussian distribution in observed frequency due to the Doppler shifting. Hence, the spectral profile has a Doppler broadened component that is a function of temperature. From Griem [26], the Doppler shape is given by

$$I(\Delta\lambda) = \left\{ \frac{M_c^2}{2\pi k t \lambda_0^2} \right\}^{\frac{1}{2}} \exp \left\{ \frac{-M_c^2}{2k y \lambda_0^2} (\Delta\lambda)^2 \right\}, \quad (\text{A-1})$$

where  $\lambda_0 = 6300 \text{ \AA} = 6.3 \times 10^{-7} \text{ m}$ ,  $c = 2.99792 \times 10^8 \text{ m/s}$ ,  $M = \text{mass of emitting species (0)} (= 2.6776 \times 10^{-26} \text{ kg})$ ,  $k = \text{Boltzman's constant} (= 1.38 \times 10^{-23} \text{ J/K})$ , and  $T = \text{temperature in K}$ .

Normalized intensity as a function of  $T$  and  $\Delta\lambda$  is found from:

$$\text{Normalized Intensity} = \exp \left\{ -2.19682568 \times 10^{26} \frac{(\Delta\lambda)^2}{T} \right\}.$$

The above equation is verified as shown in Appendix G, by comparing it with Wark's [3] version of this equation.

### Theoretical Gaussian Profiles

To convert the form of this equation to one in terms of etalon pressure instead of wavelength,  $\lambda$ , the conversion

$$\Delta\lambda = \frac{1.98 \times 10^{-11} \text{ mFSR}}{74.09 \text{ pressure units FSR}} \times \Delta \text{ pressure units} \quad (\text{A-2})$$

is substituted into equation (A-1) to obtain:

$$\text{Normalized Intensity} = \exp \left\{ -15.689416 \frac{(\Delta \text{ pressure units})^2}{T} \right\},$$

where FSR is Free Spectral Range and T is temperature in K.

The factor  $(1.98 \times 10^{-11} \text{ m FSR})$  is a fixed instrument parameter. The FSR observed with the calibration was 75.53 pressure units, and this value is used throughout the program and in the calculations except in the above step. Here, a slightly smaller FSR was used to compensate for the fact that the 6328 Å laser was used for calibration, while the observations were of 6300 Å. Since the FSR is proportional to the square of the incident wavelength, the 6328 He-Ne laser which gave a FSR of 75.53 would, at 6300 Å, have given an FSR of:

$$75.53 \left( \frac{6300}{6328} \right)^2 = 74.86$$

So the FSR of 75.53 was  $75.53 - 74.86 = 0.67$  pressure units too high. To compensate for this, the FSR used in the above step (equation (A-2)) is approximately 75 pressure units.

To input  $\Delta$  pressure units in radians  $\times 100$  (where  $2\pi$  rad of transformed data corresponds to 1 FSR of data), the conversion for each set of data

$$\frac{75.53 \text{ pressure units}}{2\pi \text{ rad} \times 100} = \frac{12.02097}{100} \quad (\text{A-3})$$

is substituted into equation (A-2) to obtain:

$$\text{Normalized Intensity} = \exp \left\{ 0.2267 \frac{(\Delta \text{ rad})^2}{T} \right\},$$

where the numerical factor 0.2267 is labeled "C0" in the program. The reason for scaling the pressure units by the factor 100 is to make a unit change in input (equation (A-3)) small enough to allow many ( $628 \sim 100 \times 2 \times \pi$ ) steps per fringe. This facilitated programming of the graphics.

To check the results of the program, the HWHMs of the laser and theoretical Gaussians can be easily convolved. The laser width of 3.5 pressure units was obtained from a pressure and temperature compensated laser calibration scan. To find the HWHM of the theoretical Gaussians start with equation (A-2).



$$\text{Normalized Intensity} = \exp \left\{ -15.689 \frac{(\text{pressure units})^2}{T} \right\} .$$

Let normalized intensity = 0.5 (i.e., HWHM), then solve for pressure units to obtain:

$$\text{HWHM in pressure units} = \sqrt{\frac{T \ln 0.5}{-15.689}} .$$

Convolving this with the laser calibration scan width provides a quick look determination of the temperature from the observed width. Examples are shown in Table A-1.

TABLE A-1. QUICK-LOOK DETERMINATION OF TEMPERATURE FROM OBSERVED WIDTH

Temperature (K)	Theoretical Gaussian HWHM (pressure units=psi)	Convolution of Theoretical Gaussian and Laser HWHM <sup>2</sup> + 3.5 <sup>2</sup> (pressure units=psi)
1000	6.646	7.51
1100	6.97	7.80
1200	7.28	8.078
1300	7.57	8.3477
1400	7.86	8.608
1500	8.14	8.8611
1600	8.407	9.107
1700	8.666	9.346
1800	8.917	9.5799
1900	9.162	9.8077
2000	9.400	10.030

APPENDIX B. SOLAR GEOPHYSICAL DATA

The following data were used in the MSIS program and were obtained from NOAA [27]:

March 1984	6	7	8	9	10	11	12	13	14	15	16
Daily Average Indices Ap	31	26	29	13	17	9	11	19	6	9	16

Daily Solar Flux at 2800 MHz (10.7 cm)	109.5	105.0	103.8	102.4	98.8	98.6	102.3*	114.7	121.1	134.4	124
--	-------	-------	-------	-------	------	------	--------	-------	-------	-------	-----

\*Adjusted for burst in progress at time of measurement.

Mean  $F_{10.7}$  for December 1983, January 1984, February 1984, and March 1984 was 90.5, 112.4, 137.2, and 120.8, respectively.

APPENDIX C. MSIS EXAMPLE OUTPUT

MARCH 7 UT

84067

INPUT UNIVERSAL TIME IN SEC

14440 (0400 UT)

INPUT MIN, MAX ALTITUDES AND ALT. STEP IN KM

130,290 20

INPUT GEODETIC LAT., EAST LONG. IN DEG

64.9,212.2

INPUT LOCAL APPARENT SOLAR TIME IN HRS

18.2

INPUT 3-MO. AVE. OF  $F_{10.7}$  FLUX

112

INPUT DAILY  $F_{10.7}$  FLUX FROM PREVIOUS DAY

109.5

INPUT DAILY AP VALUE

26

DATE	UT(SEC)	LATITUDE	LONGITUDE	LOCAL TIME	$\langle F_{10.7} \rangle$	$F_{10.7}$	AP
84067	14440	64.9	212.2	18.2	112.0	109.5	26

ALT	[HE]	[O]	[N2]	[O2]	[AR]	[H]	T
-----	------	-----	------	------	------	-----	---

130.	1.7502E+07	3.2964E+10	1.1277E+11	2.0486E+10	4.8818E+08	1.4128E+06	524.
150.	1.2866E+07	1.3711E+10	3.0166E+10	4.7443E.09	8.6045E+07	5.0339E+05	692.
170.	1.0393E+07	7.1768E+09	1.1209E+10	1.5684E+09	2.2538E+07	2.5966E+05	806.
190.	8.8380E+06	4.2385E+09	4.8880E+09	6.1621E+08	7.1885E+06	1.7916E+05	883.
210.	7.7383E+06	2.6843E+09	2.3331E+09	2.6699E+08	2.5662E+06	1.4700E+05	936.
230.	6.8957E+06	1.7755E+09	1.1773E+09	1.2287E+08	9.8243E+05	1.3188E+05	972.
250.	6.2138E+06	1.2081E+09	6.1598E+08	5.8823E+07	3.9375E+05	1.2358E+05	996.
270.	5.6406E+06	8.3788E+08	3.3030E+08	2.8927E+07	1.6285E+05	1.1829E+05	1013.
290.	5.1460E+06	5.8885E+08	1.8022E+08	1.4498E+07	6.8877E+04	1.1441E+05	1025.

APPENDIX D. PROGRAM AND EXAMPLE OUTPUT

.TYPE G.G

```

C      PROG TO GENERATE THEORETICAL GAUSSIAN PROFILES
C      AND PLOT OBSERVED PROFILES AND FIND TEMPERATURE
C      Following is the starting point (SP) value used to
C      establish the place in the data to start reading.
C      SP=1668.
C      Following is a table that appears in the report, and is not essential
C      to this program. It is included here for reference only.
C      GENERATE A TABLE TO FIND APPROX TEMP FROM HWHM OF RAW DATA
CC     TE=700.
CC     WRITE(5,*)'*****'
C      X IS TEMP, XX IS HWHM PSI, XXX IS HWHM CONVOLUTED WITH 3.5 PSI LASER
CC     WRITE(5,*)'TEMP,      HWHM THEORETICAL          HWHM CONVOLUTED '
C3     X=((TE*(ALOG(.5)))/-15.689)**.5
C      XX=X*315./27.67
C      XXX=((X*X)+(3.5*3.5))**.5
C      WRITE(5,*)TE,X,XX,XXX
C      TE=TE+100.
C      IF (TE .LE. 1700) GO TO 3
C      *****
C      The program starts here. YSM, YSSM, and YSCM are the fourier coefficients
DIMENSION YSM(15),YSCM(15),YSSM(15)
REAL F(10)
C      The following step is required for graphing only and is not essential
CALL INITT(30)
C      RETREIVE OBSERVED DATA
C      RETREIVE OBSERVED TEMP(TL), PRESSURE(PL), COUNTS(CL),
C      KILORAYLEIGHS(KL), NP IS NEW PRESSURE, NC IS NEW # OF COUNTS
REAL TL(34)
REAL PL(34)
REAL CL(34)
REAL KL(34)
REAL NP(34)
REAL NC(34)
OPEN(UNIT=3,NAME='FW2:FP16.DAT',READONLY,TYPE='OLD')
C      READ: TEMP PRES COUNTS KRAYS
C      READ BLANKS UNTIL START OF DESIRED DATA*****
ZW=1
WRITE(5,*)'ZW  ZS'
187  READ(3,*,ERR=187)
C      WRITE(5,*)ZW,ZS
ZW=ZW+1
C      MAX ZW VALUE (SP) BELOW DETERMINES STARTING PT TO READ DATA
C      For convenience, the following step was moved to near besinins of prog.
C      SP=1185.
C      IF (ZW .LE. SP) GO TO 187
ZY=1
ZY=ZY+1
C      C      *****

```

```

C      Blanks have now been read to beginning of data. Start reading data now.
      ZZ=1
189   READ(3,150)TL(ZZ),PL(ZZ),CL(ZZ),KL(ZZ)
C     PRESSURE AND INTENSITY COMPENSATION *****
      NP(ZZ)=PL(ZZ)*(298./((273.+TL(ZZ)))
      NC(ZZ)=CL(ZZ)/KL(ZZ)
C     *****
      ZZ=ZZ+1
C     The following step tells computer to read 34 lines of data (the fringe).
      IF (ZZ .LE. 34) GO TO 189
150   FORMAT (9X,F6.3,F7.2,15X,F6.0,4X,F5.3)
C     CALCULATE BACKGROUND INTENSITY ZB=AVERAGE OF 1ST AND LAST PTS
C     OF NC AFTER COMPENSATIONS, BEFORE SUBTRACTION OF BG AND NORMALIZATION
      ZB=(NC(1)+NC(34))/2
      WRITE(5,*)'TTTTTTTTTTTT'
      WRITE(5,*)'BACKGROUND INTENSITY IS'
      WRITE(5,*)ZB
C     Following step not required, so ignore.
C     WRITE(5,*)'NP(34)-NP(1)='
C     Following is specification of free spectral range in pressure units
      SRS=75.53
      WRITE(5,*)SRS
C     NORMALIZE NUMBER OF COUNTS AND SUBTRACT BACKGROUND
C     NORMALIZE BY DIVIDING BY 11TH ELEMENT OF NC (18TH FOR 34ELEMENTS ASBELOW)
      WRITE(5,*) 'NC BEFORE NORMALIZING AND AFTER SUBTRACTION OF BG'
      ZA=1
149   WRITE(5,*) NC(ZA)-ZB
      ZA=ZA+1
      IF (ZA .LE. 34) GO TO 149
C     NORMALIZE TO NN=NC(11)-ZB=MIDDLE NO. OF COUNTS - Z BACKGROUND
      WX=NC(18)-ZB
      WRITE(5,*)'WX IS'
      WRITE(5,*)WX
C     The next 4 steps are used as checks only and may be ignored.
C     NN=NC(18)-ZB
C     NN=WX
CC    WRITE(5,*)'WX IS'
CC    WRITE(5,*)WX
C     The next 4 lines are the normalization described 15 lines above
      ZW=1
152   NC(ZW)=(NC(ZW)-ZB)/WX
      ZW=ZW+1
      IF (ZW .LE. 34) GO TO 152
      WRITE(5,*)'PRESSURE, COUNTS, KILORAYLEIGHS, IN RAW DATA FORM'
      WRITE(5,*)'LAST 2 COLUMNS ARE NEW PRESSURE AND NEW COUNTS'
      WRITE(5,*)'      PL              CL              KL              NP NC'
      ZA=1
151   WRITE(5,*)PL(ZA),CL(ZA),KL(ZA),NP(ZA),NC(ZA)
      ZA=ZA+1
      IF (ZA .LE. 34) GO TO 151
      WRITE(5,*)'THIS TEXT CAUSES LAST LINE TO PRINT'
C     100+++++

```

```

C      Next two lines are free spectral range in pressure units
      WRITE(5,*)'SRS'
      WRITE(5,*)SRS
C      Next line applies only to graphing on terminal
C      CALL FINITT(0,760)
C      NOW FIND COEFFS OF OBSERVED DATA ++++++
      PI=3.1415926535
      WRITE(5,*)'
C      The following is used if the Delta Pressure Step is desired output
C      WRITE(5,*)'      DPS      NP(J)'
C      MAX M IS NO. OF FOURIER COEFFS
      DO 106 M=1,10
C      FSR IN PRESSURE UNITS
C      SRS=75.53
      SUNC=0.
      SUNS=0.
      C2=6.28318*M/SRS
      DO 107 J=1,34
      IF (J .NE. 1) GO TO 104
C      Following line calculates the difference in pressure between 1st 2 steps
      DPS=NP(J+1)-NP(J)
C      DPS=(SRS-(NP(20)-NP(1)))/2
      GO TO 105
104    CONTINUE
C      Following line calculates the difference in pressure between steps
      DPS=NP(J)-NP(J-1)
C      IF (J .NE. 20) GO TO 105
C      DPS=(SRS-(NP(20)-NP(1)))/2
105    CONTINUE
      IF (M .NE. 1) GO TO 116
C      The following step is used if Delta Pressure Step is desired output.
C      WRITE(5,*)DPS,NP(J)
116    CONTINUE
C      The next two steps calculate the Fourier cos coeffs, and sin coeffs
      SUNC=SUNC+(NC(J)/M)*(2.*SIN(C2*NP(J))*((SIN(C2*DPS/2.))**2
      *)+COS(C2*NP(J))*SIN(C2*DPS))
      SUNS=SUNS+(-NC(J)/M)*(2.*COS(C2*NP(J))*((SIN(C2*DPS/2.))**2
      *)-SIN(C2*NP(J))*SIN(C2*DPS))
107    CONTINUE
C      The next line calculates the Fourier coeffs of the fringe
      YSM(M)=SQRT(SUNS*SUNS+SUNC*SUNC)/PI
      YSCM(M)=SUNC/PI
      YSSM(M)=SUNS/PI
106    CONTINUE
C      Following are the Fourier coeffs
      WRITE(5,*)'YSM'
      WRITE(5,*)YSM(1)
      WRITE(5,*)YSM(2)
      WRITE(5,*)YSM(3)
      WRITE(5,*)YSM(4)
      WRITE(5,*)YSM(5)
      WRITE(5,*)YSM(6)

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```
WRITE(5,*)YSM(7)
WRITE(5,*)YSM(8)
WRITE(5,*)YSM(9)
WRITE(5,*)YSM(10)
C WRITE(5,*)YSM(11)
C WRITE(5,*)YSM(12)
C WRITE(5,*)YSM(13)
C WRITE(5,*)YSM(14)
C WRITE(5,*)YSM(15)
C BELOW IRRELEVANT EXCEPT FOR LASER COEFFS*****
WRITE(5,*)'NOW PLOT OBSERVED DATA USING FOURIER COEFFS'
CALL INITT(30)
CALL DWINDO(0.,629.,-1.,1.)
CALL TWINDO(50,900,50,800)
CALL MOVEA(0.,0.)
X=.01
C FIND FIRST VALUE FW AND CENTRAL CW TO N PLOT ONLY
C FW=YSM(1)*COS(.01/100)+YSM(2)*COS(2*.01/100)
C *+YSM(3)*COS(3*.01/100)+YSM(4)*COS(4*.01/100)
C *+YSM(5)*COS(5*.01/100)+YSM(6)*COS(6*.01/100)
C *+YSM(7)*COS(7*.01/100)+YSM(8)*COS(8*.01/100)
C CW=YSM(1)*COS(PI)+YSM(2)*COS(2*PI)+YSM(3)*COS(3*PI)
C *+YSM(4)*COS(4*PI)+YSM(5)*COS(5*PI)+YSM(6)*COS(6*PI)
C *+YSM(7)*COS(7*PI)+YSM(8)*COS(8*PI)+YSM(9)*COS(9*PI)
900 Y=((YSM(1)*COS(X/100)+YSM(2)*COS(2*X/100)+YSM(3)*COS(3*X/100)
*+YSM(4)*COS(4*X/100)+YSM(5)*COS(5*X/100)+YSM(6)*COS(6*X/100)
*+YSM(7)*COS(7*X/100)+YSM(8)*COS(8*X/100)+YSM(9)*COS(9*X/100)
*+YSM(10)*COS(10*X/100)+YSM(11)*COS(11*X/100)
*+YSM(12)*COS(12*X/100)+YSM(13)*COS(13*X/100)
*+YSM(14)*COS(14*X/100)+YSM(15)*COS(15*X/100))
C910 CALL DRAWA(X,Y)
X=X+1
C IF (X .LE. 629) GO TO 900
C NOW PLOT LASER PROFILE WITH FOURIER COEFFS
X=.01
REAL L(10)
C LASER COEFFS FOR FSR OF 75.7332 PSI
L(1)=.2254809
L(2)=.1613598
L(3)=.1071364
L(4)=.094267033
L(5)=.074299119
L(6)=.055807292
L(7)=.029683163
L(8)=.02819389
L(9)=.018680153
L(10)=.018559434
C 200+++++
C LASER FIRST VALUE (LF), LASER CENTRAL VALUE (LC) TO N PLOT ONLY
C FF=.7633933
C CC=-0.1496794
C920 Y=((L(1)*COS(X/100)+L(2)*COS(2*X/100)+L(3)*COS(3*X/100)
```

```

920  Y=L(1)*COS(X/100)+L(2)*COS(2*X/100)+L(3)*COS(3*X/100)
      *L(4)*COS(4*X/100)+L(5)*COS(5*X/100)
      *L(6)*COS(6*X/100)+L(7)*COS(7*X/100)
      *L(8)*COS(8*X/100)+L(9)*COS(9*X/100)
      *L(10)*COS(10*X/100)
CCC  CALL DRAWA(X,Y)
      X=X+1
CCC  IF (X .LE. 629) GO TO 920
C     *****
C     10 Fourier coeffs were found but any number up to 10 may be used
C     Now select the number of Fourier coeffs to be used:
C     YSM(3)=0.
C     YSM(4)=0.
C     YSM(5)=0.
C     YSM(6)=0.
C     YSM(7)=0.
C     YSM(8)=0.
C     YSM(9)=0.
C     YSM(10)=0.
C     PLOT CONVOLUTED PROFILE
      CALL MOVEA(0.,0.)
      DO 940 M=1,10
C     DECONVOLUTE LASER FROM OBSERVED DATA (/20 NORMALIZES ONLY)
940  YSM(M)=(YSM(M)/L(M))/20
      WRITE(5,*)'DECONVOLUTE COEFFS ARE'
      WRITE(5,*)YSM(1)
      WRITE(5,*)YSM(2)
      WRITE(5,*)YSM(3)
      WRITE(5,*)YSM(4)
      WRITE(5,*)YSM(5)
      WRITE(5,*)YSM(6)
      WRITE(5,*)YSM(7)
      WRITE(5,*)YSM(8)
      WRITE(5,*)YSM(9)
      WRITE(5,*)YSM(10)
      WRITE(5,*)'THIS TEXT CAUSES LAST LINE TO PRINT
C     TO AVERAGE COEFFS WITH ANOTHER PROFILE jjjjjjjjjjjjjjjjjjjjjjjjj
C     C1=8.738/100
C     C2=4.583/100
C     C3=2.346/100
C     YSM(1)=(YSM(1)+C1)/2
C     YSM(2)=(YSM(2)+C2)/2
C     YSM(3)=(YSM(3)+C3)/2
      WRITE(5,*)'AVERAGE COEFFS ARE'
      WRITE(5,*)YSM(1)
      WRITE(5,*)YSM(2)
      WRITE(5,*)YSM(3)
C     jjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjjj
C     *****TO PLOT LASER PROFILE INSTEAD***
C     YSM(1)=.2433401
C     YSM(2)=.1676189
C     YSM(3)=.1367357

```



```

C *****
C X=.01
C FIRST PLOT IS NON NORMALIZED VERSION
950 Y=YSM(1)*COS(X/100)+YSM(2)*COS(2*X/100)+YSM(3)*COS(3*X/100)
    *+YSM(4)*COS(4*X/100)
    *+YSM(5)*COS(5*X/100)
    *+YSM(6)*COS(6*X/100)+YSM(7)*COS(7*X/100)+YSM(8)*COS(8*X/100)
CC CALL DRAWA(X,Y)
C WRITE(S,*)'X,Y'
C WRITE(S,*)X,Y
C FIND FIRST AND CENTRAL VALUES TO NORMALIZE IN STEP 980
  IF (X .NE. .01) GO TO 960
  FV=Y
960 IF (X .NE. 315.01) GO TO 970
  WRITE(S,*)'X IS 315 X AND Y ARE'
  C WRITE(S,*)X,Y
  C WRITE(S,*)'*****'
  CV=Y
  C WRITE(S,*)'CENTRAL VALUE IS CV='
  C WRITE(S,*)CV
  C WRITE(S,*)'*****'
970 X=X+1
  IF (X .LE. 629) GO TO 950
  WRITE(S,*)'FIRST VALUE, FV, CENTRAL VALUE, CV'
  WRITE(S,*)FV,CV
  C WRITE(S,*)'CV IS'
  C WRITE(S,*)CV
  WRITE(S,*)'THIS TEXT CAUSES LAST LINE TO PRINT'
  CO=(15.689416*((SRS/(2*PI))**2))/10000
  WRITE(S,*)'CO'
  WRITE(S,*)CO
  WRITE(S,*)'XXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXXX'
  CALL MOVEA(0,0)
  C CALL DRAWA(629,0)
  C CALL MOVEA(0,2)
  C CALL DRAWA(629,2)
  C CALL MOVEA(0,4)
  C CALL DRAWA(629,4)
  C CALL MOVEA(0,5)
  C CALL DRAWA(629,5)
  C CALL MOVEA(0,6)
  C CALL DRAWA(629,6)
  C CALL MOVEA(0,8)
  C 300+++++
  C CALL DRAWA(629,8)
  C CALL MOVEA(0,1)
  C CALL DRAWA(629,1)
  C NOW PLOT NORMALIZED CONVOLUTED DATA
  CALL DWINDO(-315,315,0,1)
  X=.01-315.
980 Y=((YSM(1)*COS((X/100))+YSM(2)*COS(2*X/100)
    *+YSM(3)*COS(3*X/100))

```

```

      *+YSM(4)*COS(4*X/100)
      *+YSM(5)*COS(5*X/100)
      *+YSM(6)*COS(6*X/100)
      *+YSM(7)*COS((7*X/100))+YSM(8)*COS(8*X/100)
      *+YSM(9)*COS(9*X/100)+YSM(10)*COS(10*X/100)
      *)-CV)/(FV-CV)
      CALL DRAWA(X,Y)
C      WRITE(5,*)X,Y
      X=X+1
      IF (X .LE. 315) GO TO 980
C      PLOT THREE REFERENCE CURVES T=800,1200,1600
      CALL MOVEA(0.,1.)
C      Followins is temperature of first reference curve
      T=800
984  X=0.
      CALL MOVEA(0.,1.)
C      Below is the theoretical saussian frinse profile
985  Y=EXP(-CO*(X*X)/T)
      CALL DRAWA(X,Y)
      X=X+10
C      In the followins line 315 is used because pi radians of fringe plotted
      IF (X .LE. 315) GO TO 985
      T=T+400
      IF (T .LE. 1600) GO TO 984
C      FIND SQUARE ERROR FOR T=
C      Start searching for correct temperature at the temperature below
      T=400
      WRITE(5,*)'TEMP, ERROR'
986  ER=0.
      X=0.
987  ER=ER+((((YSM(1)*COS((X/100))+YSM(2)*COS(2*X/100)
      *+YSM(3)*COS((3*X/100))+YSM(4)*COS(4*X/100)
      *+YSM(5)*COS((5*X/100))+YSM(6)*COS(6*X/100)
      *+YSM(7)*COS((7*X/100))+YSM(8)*COS(8*X/100)
      *+YSM(9)*COS(9*X/100)+YSM(10)*COS(10*X/100)
      *)-CV)/(FV-CV))-
      *(EXP(-CO*(X*X)/T))**2)
      X=X+10
      IF (X .LE. 315.) GO TO 987
C      WRITE(5,*)'TEMP, ERROR'
      WRITE(5,*)T,ER
C      WRITE(5,*)'THIS TEXT CAUSES LAST LINE TO PRINT'
      T=T+10
      IF (T .LE. 1700) GO TO 986
      CLOSE(UNIT=3,DISPOSE='SAVE')
C      Next line applies to graphins on terminal
      CALL FINITT(0,760)
      END

```

.TYPE G.COM

```

FORTRAN/LIST:FW2:G.LST/SHOW:3 G.G
LINK G/LINKLIBRARY:FSPLIB
R G

```

60

.  
FORTRAN/LIST:FW2:G.LST/SHOW:3 G.G  
.MAIN.

. LINK G/LINKLIBRARY:FSPLIB  
?LINK-W-Undefined globals:  
\*VIRSZ

. R G

ZW ZS

+++++

BACKGROUND INTENSITY IS

988.7231

75.53000

NC BEFORE NORMALIZING AND AFTER SUBTRACTION OF BG

41.25928

136.7249

11.27686

120.3678

72.98279

119.8583

125.0291

413.9391

433.7892

816.1156

1174.443

1839.478

2851.599

4722.567

6891.408

8991.732

10968.37

11536.11

11110.94

9288.200

6731.890

4509.587

2715.975

1797.112

918.6843

663.2294

319.3577

169.7096

303.6644

149.3887

141.9388

114.2964

104.1340

-41.25934

WX IS

11536.11

PRESSURE, COUNTS, KILOBAYLEIGHS, IN RAW DATA FORM  
LAST 2 COLUMNS ARE NEW PRESSURE AND NEW COUNTS

PL	CL	KL	NP	NC
957.1900	584.0000	0.5670000	960.7752	3.5765325E-03
954.9500	628.0000	0.5580000	958.5333	1.1851901E-02
952.7000	555.0000	0.5550000	956.2716	9.7752654E-04
950.3500	610.0000	0.5500000	953.9160	1.0434001E-02
948.1900	585.0000	0.5510000	951.7607	6.3264635E-03
946.2000	633.0000	0.5710000	949.7921	1.0389834E-02
943.8500	656.0000	0.5890000	947.4012	1.0838059E-02
941.6300	843.0000	0.6010000	945.1664	3.5882030E-02
939.2400	872.0000	0.6130000	942.7611	3.7602723E-02
937.3100	1119.000	0.6200000	940.8429	7.0744425E-02
934.8000	1339.000	0.6190000	938.3013	0.1018058
932.5200	1745.000	0.6170000	936.0159	0.1594539
930.6400	2381.000	0.6200000	934.1635	0.2471890
928.0900	3541.000	0.6200000	931.5568	0.4093725
925.7500	4799.000	0.6090000	929.2112	0.5973770
923.8500	6128.000	0.6140000	927.3290	0.7794423
921.5500	7246.000	0.6060000	925.0110	0.9507859
919.3300	7565.000	0.6040000	922.7765	1.000000
917.4700	7163.000	0.5920000	920.9374	0.9631442
914.9400	6012.000	0.5850000	918.3515	0.8051413
912.7600	4532.000	0.5870000	916.1726	0.5835493
910.6800	3255.000	0.5920000	914.0910	0.3909105
908.3000	2208.000	0.5960000	911.6835	0.2354324
906.3900	1652.000	0.5930000	909.7880	0.1557814
904.0000	1133.000	0.5940000	907.3646	7.9635531E-02
901.9500	973.0000	0.5890000	905.3192	5.7491589E-02
899.3400	777.0000	0.5940000	902.6629	2.7683303E-02
897.0800	680.0000	0.5870000	900.4006	1.4711161E-02
895.0400	747.0000	0.5780000	898.3924	2.6322946E-02
892.6400	651.0000	0.5720000	895.9683	1.2949656E-02
890.3300	649.0000	0.5740000	893.6648	1.2303872E-02
888.2900	621.0000	0.5630000	891.6171	9.9077048E-03
885.7800	612.0000	0.5600000	889.0557	9.0267882E-03
883.5200	523.0000	0.5520000	886.7994	-3.5765378E-03

THIS TEXT CAUSES LAST LINE TO PRINT

SRS

75.53000

YSM

0.3696032

0.2238235

0.1115526

4.6207257E-02

1.6699495E-02

5.2267537E-03

1.8393506E-03

1.2564564E-03

ORIGINAL PAGE IS  
OF POOR QUALITY

3.5860834E-03  
3.0901073E-03

NDW@PLOT OBSERVED DATA USING FOURIER COEFFS

7 @

DECONVOLUTED COEFFS ARE

8.1958868E-02  
6.9355413E-02  
5.2061040E-02  
2.4508704E-02  
1.1238017E-02  
4.6828589E-03  
0.0000000  
0.0000000  
0.0000000  
0.0000000

THIS TEXT CAUSES LAST LINE TO PRINT  
FIRST VALUE, FV, CENTRAL VALUE, CV

0.2438049 -4.6711199E-02

THIS TEXT CAUSES LAST LINE TO PRINT  
CD

0.2267180

ZXZBZIRRRZNRZKXNMZBRZUWXNZXZYZZBZAZBZEBZEMZIKZMLMWPQxST!x#TgUWzXZ(\;f\_)#@BC DF"  
'GaHJbKlCnDqOfErsfUVswhYZiC JjLjZAKBCIEFGImJ"mZJKMNOQR TUvXYZ\1J]f1&@KABDJEFiHIhJLgM  
OfFeQSDtCuwBxay'\[\ ]\*\_)'@A:CCD'('IzEGyHxJwKlvNuOPtRsSTrUr\*PPLQvSxTzU:W X"ayd[\s\j  
fn\_at@uByC)D#aFeG#e+GjHoJtNyL N#eOKPaRwS\*TXeVlWsy(Z&bCj]st(\\_d,A1BuC E(hFrG(I)eJp  
NzM)z,M\*dNoOzQ+eRfTCU,fVaX:Y-hZs\ J.Jtv=@/bAmByI0dEeF(HifIrJ)l2hMsN)P3hQ3h-QrS)T4  
sUpWzX5cYl[Cu\^]6f\_n.@vA)C7dDKEqGwH:I8kflKNoOrPvRxS(T)V^W X8 .X9`Z[\8 f\*\_)/@BxCu  
DrFoGk IfJaK7!MwNaDjQdR6)SvUnVfW5~YuZl[Cj4zfp\_4p/\_f0A3:BrChE2)FrHhI1)JaLfm0(NoPdQ/  
xRmTaU.vVjX- YsZh\,:Jatf1@+zAoCdD+d1D\*yeGdH)zIoNeL( (MaOHp'QuS1TcU&CWrXjYb[Zz\st  
l\_e2@\*BwCqDkFeG# HyJ#y2JtKoLjNeOaP" )RySuTqVnWjXsZd[aj' f!\_z3AxBvCtErFr6AsBCtDuFG  
vHwJKxM!::6MyNzQQ(R:S)UV^W Y" \Z[a]bfc\_c7AdBeCEfFgHIhJLlMnJpQkRTU1VXYmZ\J]f1m8@A"m8AB  
DEGHIKLlMOFGKSTJUWX1YCh\]s\_f9@BeCDdFGcHbJKaLN\0' PR^ST)U:WX(Z[z ]]z9]fy\_y:AxBCwEFG  
vI.JuKMNOlQRSUsVWYZ\Jrfr;@ABDEFHr:BC9'.[[[BK/H6pV6p/V3w0C0oQ-cf\*^1L'mYZ12G#^T^"3Bj  
0' ]:4JuXs5ESr6@Ncr7IVr8DQ\_r9LZr:GUr:Ur;BFJ9'.[[[Br/H71V5K0C2:Q0f+~n1L\*(YCr2G&tTz  
c3B#^0cJ" a4-JdX' )5ExS!x5Su6@tNsEr7IVr8DQ\_r9LZr:GUr;BFJ9'.[[[Bu/H7wv6f0C4hU2^f/t1L-1  
Y+a2G

TEMP.	ERROR
400.0000	0.2097060
410.0000	0.1945095
420.0000	0.1801474
430.0000	0.1665846
440.0000	0.1537879
450.0000	0.1417260
460.0000	0.1303690
470.0000	0.1196891
480.0000	0.1096597
490.0000	0.1002555
500.0000	9.1452442E-02
510.0000	8.3228029E-02
520.0000	7.5560570E-02
530.0000	6.8429396E-02
540.0000	6.1815027E-02
550.0000	5.5698719E-02
560.0000	5.0062709E-02
570.0000	4.4890013E-02
580.0000	4.0164385E-02
590.0000	3.5870302E-02
600.0000	3.1992738E-02
610.0000	2.8518166E-02
620.0000	2.5432343E-02
630.0000	2.2722503E-02
640.0000	2.0376189E-02
650.0000	1.8381432E-02
660.0000	1.6726822E-02
670.0000	1.5401303E-02
680.0000	1.4394365E-02
690.0000	1.3695851E-02
700.0000	1.3296020E-02
710.0000	1.3185506E-02
720.0000	1.3355302E-02
730.0000	1.3796733E-02
740.0000	1.4501466E-02
750.0000	1.5461468E-02
760.0000	1.6668990E-02
770.0000	1.8116590E-02
780.0000	1.9797074E-02
790.0000	2.1703515E-02
800.0000	2.3829229E-02
810.0000	2.6167762E-02
820.0000	2.8712889E-02
830.0000	3.1458605E-02
840.0000	3.4399092E-02
850.0000	3.7528761E-02
860.0000	4.0842198E-02
870.0000	4.4334162E-02
†C880.0000	4.7999568E-02
†C	

TYPE 0.0

```
C   PROG TO READ AND MODIFY LASER PROFILES
C   TEMP(LASER) PRESSURE(LASER) COUNTS(LASER)
    DIMENSION YSM(35),YSCM(35),YSSM(35)
    REAL TL(130)
    REAL PL(130)
    REAL CL(130)
    REAL NP(130)
    REAL NC(68)
    OPEN(UNIT=3,NAME='FW2:FP11.DAT',READONLY,TYPE='OLD')
    WRITE(5,*)'  TL          NP          CL          NC'
    DO 100 I=1,68
      READ(3,50) TL(I),PL(I),CL(I)
      IF (I .LE. 4 ) GO TO 100
C   START TEMPERATURE COMPENSATION
C   NEW PRESSURE (NP)=PRESSURE OF LASER(PL)*TEMP FACTOR
      NP(I)=PL(I)*(298./.(273.+TL(I)))
C   NORMALIZE THE COUNTS
      NC(I)=CL(I)/32350
50  WRITE(5,*)TL(I),NP(I),CL(I),NC(I)
100  FORMAT (9X,F6.3,F7.2,15X,F6.0)
C   CONTINUE
C   NOW SMOOTH THE LASER SCAN USING FOURIER/ROBLE METHOD
      FI=3.1415926535
      WRITE(5,*)'NP AT 1,2,3,4,5'
      WRITE(5,*)NP(1)
      WRITE(5,*)NP(2)
      WRITE(5,*)NP(3)
      WRITE(5,*)NP(4)
      WRITE(5,*)NP(5)
      WRITE(5,*)'NP AT 68,69'
      WRITE(5,*)NP(68)
      WRITE(5,*)NP(69)
      NP(1)=NP(5)
      NP(20)=NP(68)
      NP(2)=NP(17)
      NP(3)=NP(19)
      NP(4)=NP(21)
      NP(5)=NP(23)
      NP(6)=NP(25)
      NP(7)=NP(27)
      NP(8)=NP(29)
      NP(9)=NP(31)
      NP(10)=NP(33)
      NP(11)=NP(35)
      NP(12)=NP(37)
```

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NP( 13 )=NP( 39 )
NP( 14 )=NP( 41 )
NP( 15 )=NP( 43 )
NP( 16 )=NP( 45 )
NP( 17 )=NP( 47 )
NP( 18 )=NP( 49 )
NP( 19 )=NP( 51 )
W=1
WRITE( 5,*)'NEW PRESSURE VALUES NP'
103 WRITE( 5,*)W,NP( W )
W=W+1.
IF ( W .LE. 20 ) GO TO 103
DO 106 M=1,21
C FSR IN PRESSURE UNITS
C SRS=53.2464
C SRS=43.3838
SRS=75.5332
SUNC=0.
SUNS=0.
C2=6.28318*M/SRS
DO 107 I=5,68
IF ( I .NE. 5 ) GO TO 104
DPS=NP( I+1 )-NP( I )
GO TO 105
104 CONTINUE
DPS=NP( I )-NP( I-1 )
105 CONTINUE
IF ( M .NE. 1 ) GO TO 116
CC WRITE( 5,*)DPS
CC WRITE( 5,*)NP( I )
116 CONTINUE
SUNC=SUNC+( NC( I )/M )*( 2.*SIN( C2*NP( I ) )*( ( SIN( C2*DPS/2. ) )**2
* )+COS( C2*NP( I ) )*SIN( C2*DPS ) )
SUNS=SUNS+( -NC( I )/M )*( 2.*COS( C2*NP( I ) )*( ( SIN( C2*DPS/2. ) )**2
* )-SIN( C2*NP( I ) )*SIN( C2*DPS ) )
107 CONTINUE
YSM( M )=SQRT( SUNS*SUNS+SUNC*SUNC )/PI
YSCM( M )=SUNC/PI
YSSM( M )=SUNS/PI
106 CONTINUE
WRITE( 5,*)'YSM'
WRITE( 5,*)YSM( 1 )
WRITE( 5,*)YSM( 2 )
WRITE( 5,*)YSM( 3 )
WRITE( 5,*)YSM( 4 )
WRITE( 5,*)YSM( 5 )
WRITE( 5,*)YSM( 6 )
WRITE( 5,*)YSM( 7 )
WRITE( 5,*)YSM( 8 )
WRITE( 5,*)YSM( 9 )

```



```

WRITE(5,*)YSM(10)
WRITE(5,*)YSM(11)
WRITE(5,*)YSM(12)
WRITE(5,*)YSM(13)
WRITE(5,*)YSM(14)
WRITE(5,*)YSM(15)
WRITE(5,*)YSM(16)
WRITE(5,*)YSM(17)
WRITE(5,*)YSM(18)
WRITE(5,*)YSM(19)
WRITE(5,*)YSM(20)
WRITE(5,*)YSM(21)
CALL INITT(30)
CALL DWINDO(0.,629.,-1.,1.)
CALL TWINDO(50,900,50,800)
CALL MOVEA(0.,0.)
X=.01
900  Y=YSM(1)*COS(X/100)+YSM(2)*COS(2*X/100)+YSM(3)*COS(3*X/100)+YSM(4)
    **COS(4*X/100)+YSM(5)*COS(5*X/100)+YSM(6)*COS(6*X/100)+YSM(7)*COS
    *(7*X/100)
    *+YSM(8)*COS(8*X/100)
    *+YSM(9)*COS(9*X/100)+YSM(10)*COS
    *(10*X/100)
CC   *+YSM(11)*COS(11*X/100)+YSM(12)*COS(12*X/100)
C    *+YSM(13)*COS(13*X/100)+YSM(14)*COS(14*X/100)+YSM(15)*COS(15*X/100)
C    *+YSM(16)*COS(16*X/100)+YSM(17)*COS(17*X/100)+YSM(18)*COS(18*X/100)
C    *+YSM(19)*COS(19*X/100)+YSM(20)*COS(20*X/100)+YSM(21)*COS(21*X/100)
C    IF (X .GE. 2) GO TO 910
    WRITE(5,*)'X,Y'
    WRITE(5,*)X,Y
910  CALL DRAWA(X,Y)
    X=X+1
    IF (X .LE. 629) GO TO 900
    CLOSE(UNIT=3,DISPOSE='SAVE')
    CALL FINITT(0,760)
    END

```

.TYPE 0.COM

```

FORTRAN/LIST:FW2:0.LST/SHOW:3 0.0
LINK 0/LINKLIBRARY:FSPLIB
R 0

```

.@0

. FORTRAN/LIST:FW2:0.LST/SHOW:3 0.0  
.MAIN.

. LINK 0/LINKLIBRARY:FSPLIB  
?LINK-W-Undefined globals:  
\$VIRSZ

. R 0  
TL NP CL NC  
?Err 5  
in routine ".MAIN." line 10

?Err 5  
in routine ".MAIN." line 10

23.77200	939.0898	3509.000	0.1084699
23.88800	940.5798	3409.000	0.1053787
23.99800	941.9172	3434.000	0.1061515
24.10500	943.1827	3317.000	0.1025348
24.21600	944.3544	3374.000	0.1042968
24.31800	945.4337	3304.000	0.1021329
24.44800	946.3330	3333.000	0.1030294
24.64100	946.7205	3189.000	9.8578051E-02
24.72700	947.6982	3211.000	9.9258117E-02
24.77300	948.9128	3386.000	0.1046677
24.81400	950.1931	3363.000	0.1039567
24.84900	951.4221	3287.000	0.1016074
24.88700	952.6312	3385.000	0.1046368
24.97300	953.2164	3454.000	0.1067697
24.99300	954.5724	3323.000	0.1027202
25.02300	955.7662	3642.000	0.1125811
25.05100	956.8363	3622.000	0.1119629
25.06800	957.8914	3722.000	0.1150541
25.07500	959.3185	3837.000	0.1186090
25.08100	960.7289	4064.000	0.1256260
25.10700	961.6247	4104.000	0.1268024
25.11100	963.1113	4476.000	0.1383617
25.11600	964.4546	4941.000	0.1527357
25.12700	965.7284	5274.000	0.1630294
25.12900	967.0714	5498.000	0.1699536
25.13400	968.3645	6284.000	0.1942504
25.14800	969.1286	6991.000	0.2161051
25.14300	970.1844	8488.000	0.2623802
25.13200	971.5097	11030.00	0.3409583
25.12700	972.7354	16357.00	0.5056260
25.12600	974.1281	24998.00	0.7727357
25.12400	975.5840	32280.00	0.9978362
25.12700	976.8936	32312.00	0.9988254

25.14000	977.7606	28974.00	0.8956414
25.14000	979.2099	22393.00	0.6922102
25.14600	980.5896	16566.00	0.5120866
25.15100	981.8824	12226.00	0.3779289
25.15600	983.2053	9171.000	0.2834930
25.15900	984.4948	7474.000	0.2310355
25.16800	985.0446	6617.000	0.2045441
25.16500	986.5237	5754.000	0.1778671
25.16200	987.8430	5349.000	0.1653478
25.15800	988.9757	4901.000	0.1514992
25.15800	990.1851	4684.000	0.1447913
25.16000	991.4576	4362.000	0.1348377
25.16000	992.7970	4224.000	0.1305719
25.15200	993.5732	4035.000	0.1247295
25.15100	994.8059	3970.000	0.1227202
25.14900	996.0220	3882.000	0.1200000
25.14500	997.3847	3711.000	0.1147141
25.14300	998.7208	3633.000	0.1123029
25.14100	999.9669	3773.000	0.1166306
25.14200	1000.813	3606.000	0.1114683
25.13800	1002.316	3611.000	0.1116229
25.13900	1003.622	3613.000	0.1116847
25.14200	1004.941	3541.000	0.1094590
25.14500	1006.290	3732.000	0.1153632
25.15300	1007.503	3567.000	0.1102628
25.15700	1008.749	3707.000	0.1145904
25.15800	1009.565	3607.000	0.1114992
25.15800	1011.004	3609.000	0.1115611
25.16200	1012.160	3605.000	0.1114374
25.15600	1013.529	3551.000	0.1097682
25.14300	1014.623	3434.000	0.1061515
NP AT 1,2,3,4,5			
0.0000000			
0.0000000			
0.0000000			
0.0000000			
939.0898			
NP AT 68.69			
1014.623			
0.0000000			
NEW PRESSURE VALUES NP			
1.000000	939.0898		
2.000000	952.6312		
3.000000	954.5724		
4.000000	956.8363		
5.000000	959.3185		
6.000000	961.6247		
7.000000	964.4546		

8.000000	967.0714
9.000000	969.1286
10.00000	971.5097
11.00000	974.1281
12.00000	976.8936
13.00000	979.2099
14.00000	981.8824
15.00000	984.4948
16.00000	986.5237
17.00000	988.9757
18.00000	991.4576
19.00000	993.5732
20.00000	1014.623

YSM

0.2254809  
0.1613598  
0.1071364  
9.4267033E-02  
7.4299119E-02  
5.5807292E-02  
2.9683163E-02  
2.8193893E-02  
1.8680153E-02  
1.8559434E-02  
4.4040936E-03  
9.1322120E-03  
4.7645955E-03  
7.8142360E-03  
4.0516402E-03  
2.6422290E-03  
5.1584304E-04  
2.9768483E-03  
3.2191928E-03  
3.7323066E-03  
1.9601721E-03

X,Y

9.9999998E-03	0.8134672
X,Y	
1.010000	0.8127617
X,Y	
2.010000	0.8106764
X,Y	
3.010000	0.8072207
X,Y	
4.010000	0.8024113
X,Y	
5.010000	0.7962701
X,Y	
6.010000	0.7888262
X,Y	
7.010000	0.7801136
X,Y	
8.010000	0.7701731
X,Y	
9.010000	0.7590501
X,Y	
10.01000	0.7467955
X,Y	
11.01000	0.7334650
X,Y	
7 12.01000	0.7191188

## APPENDIX E. SHADOW HEIGHT

### Shadow Height in Two Dimensions

To explain the concept of shadow height, this section shows how shadow height is calculated for the simplified case of two dimensions.

The first step is to find:

$M$  = distance from observer to shadow along line of observations from  $\alpha$  and  $\theta$ , where

$\alpha$  = angle from local vertical to line of observation, where positive values are toward the Sun, and

$\theta$  = change in latitude between observer and sunset as shown.

$M$  is found in terms of  $\alpha$  and  $\theta$ . Then, shadow height will be found in terms of  $M$  and  $\alpha$ .

Let  $\beta \equiv \alpha - \theta$ ,  $h \equiv M \cos \beta$

$$M = \frac{h}{\cos \beta} = \frac{h}{\cos (\alpha - \theta)} = \frac{r - r \cos \theta}{\cos (\alpha - \theta)}$$

which applies for  $\alpha \neq 90^\circ$ .

Shadow height (SH) is determined from  $M$  and  $\alpha$  (see Figure E-1), using the cosine law, where  $\psi \equiv 90 - |\alpha|$

$$(SH + r)^2 = r^2 + M^2 - 2rM \cos (90 + \psi)$$

$$SH + r = [r^2 + M^2 - 2rM \cos (90 + \psi)]^{\frac{1}{2}}$$

$$SH = [r^2 + M^2 - 2rM \cos (90 + \psi)]^{\frac{1}{2}} - r$$

$$SH = [r^2 + M^2 - 2rM \cos |\alpha|]^{\frac{1}{2}} - r$$

$$\begin{aligned}
\text{because } \cos (90 + \psi) &= \cos (90 + (90 - |\alpha|)) \\
&= \cos (180 - |\alpha|) \\
&= -\cos |\alpha|
\end{aligned}$$

#### Shadow Height Program Rotations Used

Two shadow height programs are included. One of these programs (L.SH) is a simple version, but it requires the solar depression angle T1, and the azimuth (or bearing, i.e., E from N) angle from the Sun of the observation be input (in degrees). These two angles may not be known, but they are not required in the more complex version of the program (M.SH). Instead, the time and place are input and these two angles are calculated in M.SH. The rest of M.SH is the same as program L.SH.

The components of the vector from the Earth to the Sun in "inertial space" are found in M.SH. This vector is called VI. By performing a series of rotations, this vector is transformed into the frame of the observer, so the resulting vector VL gives the direction to the Sun from the observer in terms of local zenith and azimuth.

To demonstrate this, an example is shown below. In Figure E-2 the components of VI are X, Y, and Z. With an angle of  $66^\circ$  for  $\phi$ , these components are 0, 0.9, and 0.4, respectively. The next rotation produces the "Geographic Vector," VG, which is fixed with respect to Earth and points toward the Sun. The amount of rotation, -GSTR, depends on the time.

The next rotation is the longitude rotation. If the observer is at the  $0^\circ$  longitude, in England, as he is in this example, the rotation is 0.

The next rotation is the latitude rotation. Here, the coordinate system is rotated  $90^\circ$  about Y; then the coordinate system is rotated through an angle that equals latitude, so the Z axis is in the local vertical. The X axis points south and the Y axis points east. As in the other rotations, the new components for the unit vector pointing to the Sun in this rotated coordinate frame are computed. These new vector components are VL(1), VL(2), and VL(3).

Finally, the observer has the components of a unit vector pointing toward the Sun in terms of local zenith, south, and east, so he can easily

calculate the zenith and bearing angles to the Sun. The zenith angle; i.e., the angle from the zenith to the Sun is

$$\phi = \tan^{-1} \frac{\sqrt{y^2 + x^2}}{z}$$

which can be visualized by translating the XY plane component along Z to the maximum Z value.

To find the bearing, or azimuth angle,

$$\theta = \tan^{-1} \frac{y}{x}$$

is computed, where  $\theta$  is measured to the E from S. To convert this to a "bearing" angle which is defined as angle E of N,

$$\text{Azimuth} = \text{Bearing} = 180 - \theta$$

TABLE E-1. SHADOW HEIGHT PROGRAM VECTOR COMPONENTS

	X	Y	Z	
VI	0	0.9	0.4	inertial vector
VG	0.9	0	0.4	geographic vector
VJ	0.9	0	0.4	after longitude rotation
VL	0.05	0	0.95	after latitude rotation

## Shadow Height Programs

```
TYPE M.SH
C   m.sh SHADOW HEIGHT PROG. BASED ON LOCAL TIME
    DIMENSION VI(3),BV(3),VG(3),VL(3),VJ(3),PV(3),ZZ(3,3)
    DIMENSION C(3,3),CT(3,3)
    CALL INITT(30)
    CALL DWINDO(7200,67000.,0.,250.)
    CALL TWINDO(50,900,50,800)
    CALL MOVEA(0.,0.)
    CALL MOVEA(3600.,0.)
    CALL DRAWA(3600.,10.)
    CALL MOVEA(7200.,0.)
    CALL DRAWA(7200.,0.)
    RAD=57.295779
C   LATITUDE AND LONGITUDE OF OBSERVER IS
    ALA=64.86001/RAD
    ALO=-147.84711/RAD
C   ALA=49./RAD
C   ALO=-122./RAD
    IDAY=61
C   IDAY=46
    IYR=1984
    3   SECS=0.
        CALL MOVEA(0.,0.)
C   4   WRITE(5,*)'ENTER YEAR (INTEGER), DAY (INTEGER), SECONDS (REAL)'
    4   MS=1
C   *****
C   ?????????????????
C   5   READ(5,*)   IYR, IDAY, SECS
        IF (IYR.LT.1901.OR.IYR.GT.2099) STOP
        CALL SUN(IYR, IDAY, SECS, GST, SLONG, SRASN, SDEC)
C   WRITE(5,200)
C   WRITE(5,300) IYR, IDAY, SECS, GST, SLONG, SDEC
C   GO TO 5
C00  FORMAT(2I4,F10.2)
200  FORMAT(8X,'IYR',6X,'IDAY',6X,'SECS',7X,'GST',5X,'SLONG',
        *5X,'SDEC')
300  FORMAT(1X,2I10,4F10.3)
C   *****
C   NOW FIND INERTIAL COORDS (VI) FROM SDEC AND SLONG
    P=((3.141592/2)-(SDEC/RAD))
    SLONG=SLONG/RAD
    VI1=SIN(P)*COS(SLONG)
    VI2=(SIN(P)*SIN(SLONG))
    VI3=COS(P)
    VI(1)=VI1
    VI(2)=VI2
    VI(3)=VI3
C   WRITE(5,*)'INERTIAL VECTOR VI X Y Z COMPS ARE'
C   WRITE(5,*)VI
C   *****
C   NOW TRANSFORM THE INERTIAL COORDINATES (VI) TO GEOGRAPHIC (VG)
    GSTR=GST/RAD
    CALL ROTXYZ(-GSTR,ZZ,3)
```



```

      CALL TRF(VI,ZZ,VG)
C     WRITE(5,*)' GEOGRAPHIC VECTOR VG X Y Z COMPS TO SUN FROM EARTH'
C     WRITE(5,*)VG
C     *****:*****
C     NOW TRANSFORM THE GEOGRAPHIC COORDS (VG) TO LOCAL AZIMUTH AND ZENITH
C NEXT ROTATE LONG ABOUT Z AXIS
      XI=90./RAD
      CALL ROTXYZ(-AL0,ZZ,3)
      CALL TRF(VG,ZZ,VJ)
C     WRITE(5,*)' VJ IS'
C     WRITE(5,*)VJ
C NEXT ROTATE LAT, AND 90 DEGREES ABOUT Y
      E=-AL0+XI
      CALL ROTXYZ(E,ZZ,2)
      CALL TRF(VJ,ZZ,VL)
C     WRITE(5,*)'ZENITH FROM ACOS Z TO SUN IS '
      X=VL(1)
      Y=VL(2)
      Z=VL(3)
      SQ=SQRT((X**2)+(Y**2))
      AZ=(ATAN2(SQ,Z))*RAD
C     WRITE(5,*)AZ
      TH=(ATAN2(Y,X))*RAD
      B=180-TH
C     WRITE(5,*)'AZIMUTH TH IS'
C     WRITE(5,*)TH
C     WRITE(5,*)'BEARING TO SUN IS'
C     WRITE(5,*)B
C     *** NOW FIND SHADOW HEIGHT*****
C     FIND SOLAR DEPRESSION ANGLE
      TI=AZ-90.
C     TI=45.
C     WRITE(5,*)'SOLAR DEPRESSION ANGLE IS'
C     WRITE(5,*)TI
C     INPUT GEOGRAPHIC BEARING ANGLE OF OSERVATION
C     WRITE(5,*)'INPUT GEOGRAPHIC BEARING ANGLE OF OBS (DEG)'
C     READ(5,*)GB
      GB=.01
C     WRITE(5,*)'=====
C00  WRITE(5,*)'GEOGRAPHIC BEARING ANGLE OF OBS IS'
C     WRITE(5,*)GB
900  MS2=1
C     INPUT ENITH OF OSERVATION (DEGREES)
C     WRITE(5,*)'INPUT ZENITH OF OBS'
C     READ(5,*)ZI
      ZI=60.
C     WRITE(5,*)'ZENITH OF OBSERVATION IS '
C     WRITE(5,*)ZI
C     CALCULATE BEARING ANGLE OF OBSERVATION FROM SUN
910  SI=GB-B
C     SI=45.
C     WRITE(5,*)'BEARING ANGLE OF OBSERVATION FROM SUN IS'
C     WRITE(5,*)SI

```

```

C      NOW RUN SHADOW HEIGHT PROG L.SH
C      Z=ZENITH ANGLE, S=AZIMUTH, T=SOLAR DEPRESSION ANGLE
C      RADIUS OF EARTH IS KM IS:
C      RA=6378.
C      THE FOLLOWING 3 VALUES ARE INPUTS TO PROG (ABOVE)
C      ZI IS ZENITH INPUT (DEGREES) OF OBSERVATION
C      SI IS AZIMUTH INPUT (DEGREES) BEARING ANGLE FROM SUN OF OBSERVATION
C      TI IS SOLAR DEPRESSION ANGLE INPUT (DEGREES)
C      NOW CONVERT THESE TO RADIAN:
C      Z=ZI/RAD
C      S=SI/RAD
C      T=TI/RAD
C      C1=1/((SIN(T))**2)
C      C2=((COS(Z))**2)*(COS(T))**2)/((SIN(Z))**2)*(SIN(T))**4)
C      C3=(2*COS(Z)*COS(T)*COS(S))/(SIN(Z)*(SIN(T))**3)
C      C4=(COS(T)*COS(S))/SIN(T)
C      C5=COS(Z)/((SIN(T))**2)*SIN(Z)
C      C6=C4+C5+((C1+C2+C3)**.5)
C      AP=(SIN(Z)*C6)-(COS(Z))
C      P=RA/AP
C      AG=((RA**2)+(P**2)+(2*RA*P*(COS(Z))))**.5)-RA
C      WRITE(5,*)'HEIGHT ABOVE GROUND IS'
C      WRITE(5,*)AG
C      GB=GB+90.
C      WRITE(5,*)'=====
C      IF (GB .EQ. 90.01) GO TO 950
C      IF (GB .EQ. 270.01) GO TO 955
C      IF (GB .EQ. 360.01) GO TO 960
C      IF (GB .EQ. 450.01) GO TO 970
C      GO TO 1000
950    S2=AG
C      GO TO 1010
955    S1=AG
C      GO TO 1010
960    W1=AG
C      GO TO 1010
970    S4=AG
C      GO TO 1020
1000   S3=AG
1010   IF(GB .LT. 280) GO TO 900
C      NOW ZENITH OBSERVATION
C      ZI=0.01
C      GO TO 910
C      WRITE IN ORDER OF N E S W ZENITH
C020   WRITE(5,*)S2,S3,S1,W1,S4
1020   LL=1.
C      IF (SECS .GT. 81200.) GO TO 2000
C      IF (SECS .LT. 10800.) GO TO 2000
C      CALL DRAWA (SECS,S4)
2000   SECS=SECS+1200.
C      IF(SECS .LT. 86400.) GO TO 4
C      IDAY=IDAY+8
C      WRITE(5,*)'DAY'

```

```
C WRITE(5,*)IDAY
GO TO 3
CALL FINITT(0,760)
END
```

```
.TYPE M.COM
FORTRAN/LIST:FW2:M.LST/SHOW:3 M.SH
LINK M,ALL11/LIBRARY:FSPLIB
R M
```

```
.TYPE SUN.CAL
SUBROUTINE SUN(IYR, IDAY, SECS, GST, SLONG, SRASN, SDEC)
DATA RAD /57.29578/
DOUBLE PRECISION DJ, FDAY
IF( IYR.LT.1901.OR.IYR.GT.2099) RETURN
FDAY=SECS/86400.
DJ=365*( IYR-1900 )+( IYR-1901 )/4+IDAY+FDAY-0.5D0
T=DJ/36525.
VL=DMOD( 279.696678+0.9856473354*DJ, 360.D0)
GST=DMOD( 279.690983+0.9856473354*DJ+360.*FDAY+180., 360.D0)
G=DMOD( 358.475845+0.985600267*DJ, 360.D0)/RAD
SLONG=VL+( 1.91946-0.004789*T)*SIN( G)+0.020094*SIN( 2.*G)
OBLIQ=( 23.45229-0.0130125*T)/RAD
SLP=( SLONG-0.005686 )/RAD
SIND=SIN( OBLIQ)*SIN( SLP)
COSD=SQRT( 1.-SIND**2)
SDEC=RAD*ATAN( SIND/COSD)
COTAN=COS( OBLIQ )/SIN( OBLIQ)
SRASN=180.-RAD*ATAN2( COTAN*SIND/COSD, -COS( SLP)/COSD)
RETURN
END
```

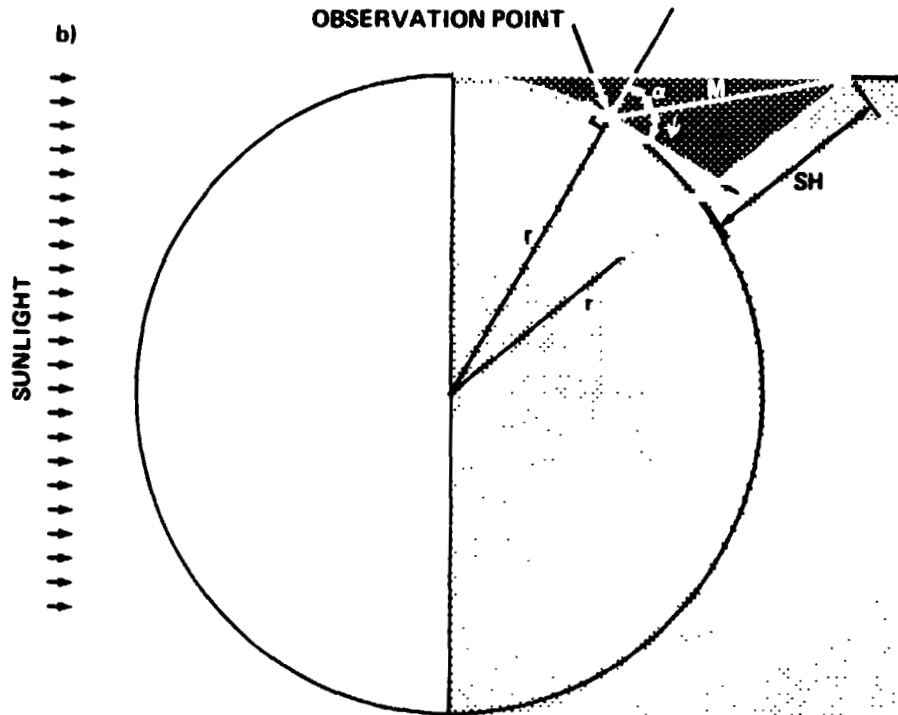
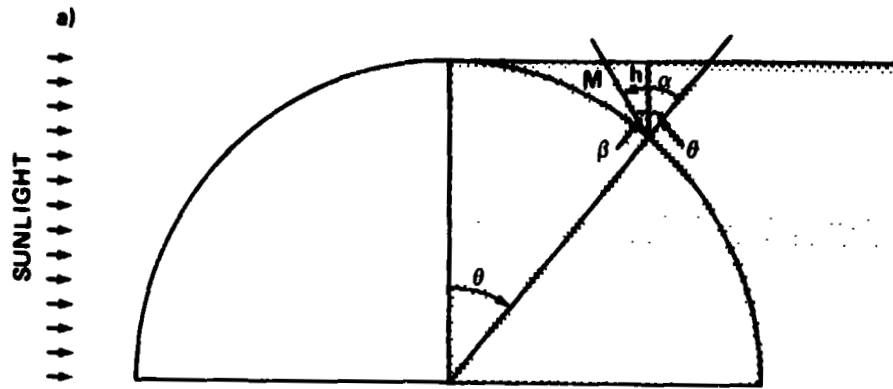
```

SUBROUTINE TRF(X,A,XT)
DIMENSION X(3),A(3,3),XT(3)
DO 1 I=1,3
XT(I)=0.0
DO 1 J=1,3
XT(I)=XT(I)+A(I,J)*X(J)
1 CONTINUE
RETURN
END

SUBROUTINE ROTXYZ(A,B,IROT)
DIMENSION B(3,3)
A1=COS(A)
A2=SIN(A)
GO TO (1,2,3),IROT
C ROTATION ABOUT THE X-AXIS
1 B(1,1)=1.0
B(1,2)=0.0
B(1,3)=0.0
B(2,1)=0.0
B(2,2)=A1
B(2,3)=-A2
B(3,1)=0.0
B(3,2)=A2
B(3,3)=A1
RETURN
C ROTATION ABOUT THE Y-AXIS
2 B(1,1)=A1
B(1,2)=0.0
B(1,3)=-A2
B(2,1)=0.0
B(2,2)=1. ^
B(2,3)=0.0
B(3,1)=A2
B(3,2)=0.0
B(3,3)=A1
RETURN
C ROTATION ABOUT THE Z-AXIS
3 B(1,1)=A1
B(1,2)=-A2
B(1,3)=0.0
B(2,1)=A2
B(2,2)=A1
B(2,3)=0.0
B(3,1)=0.0
B(3,2)=0.0
B(3,3)=1.0
RETURN
END

```

ORIGINAL PAGE IS  
OF POOR QUALITY



SHADOW HEIGHT GEOMETRY

Figure E-1. Local shadow height geometry.

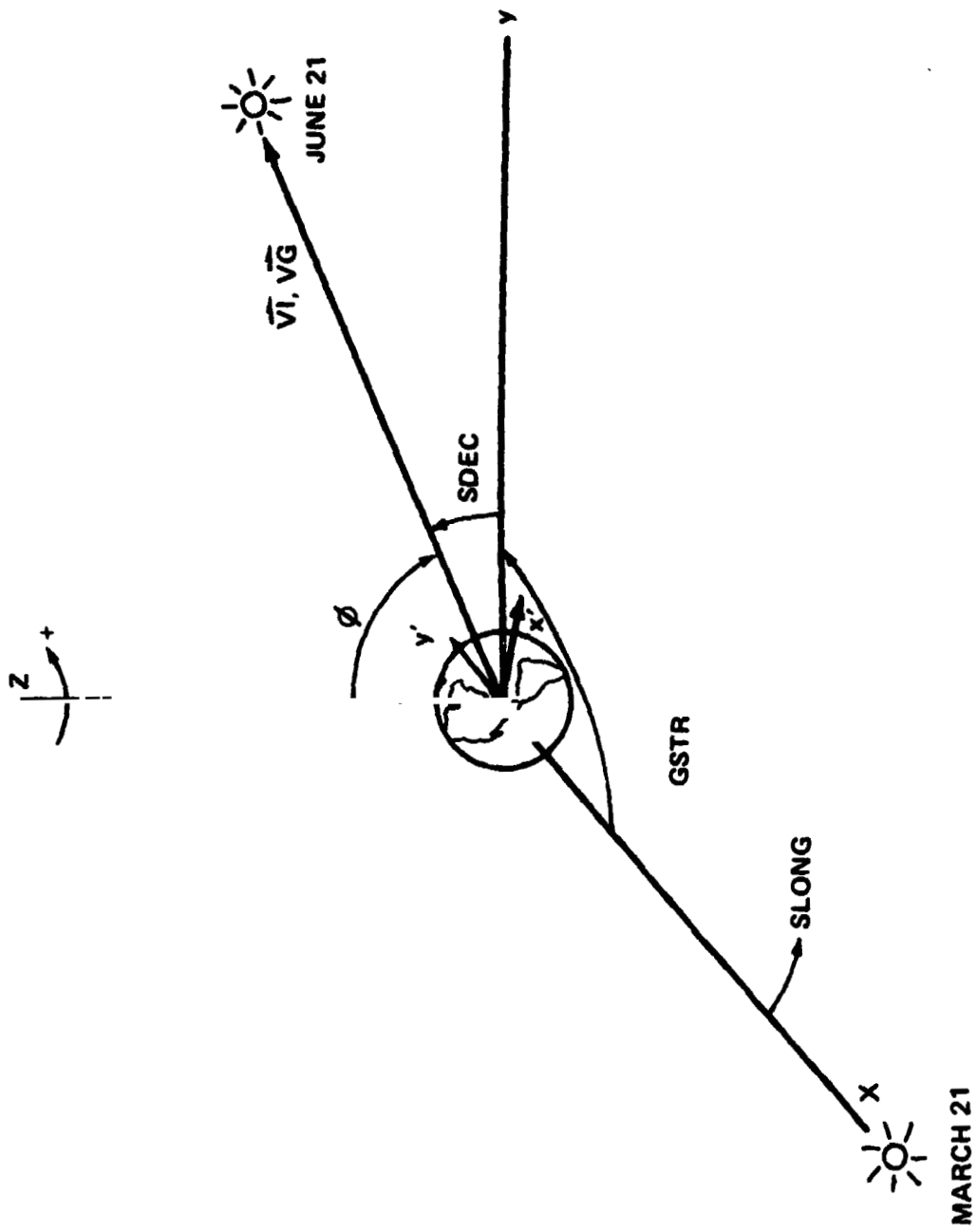


Figure E-2. Global geometric parameters used in shadow height program.

APPENDIX F. EXAMPLE OF RAW DATA

Scan Number	T, °C	Pressure	Encoder Position	Counts	Photometer	Mirror Position
3	24	246	897.78	3876	0	0
3	35	24	248	899	89	3868
3	38	24	245	902	33	4058
3	37	24	246	904	17	4150
3	38	24	251	908	85	4240
3	39	24	240	908	46	4330
2	40	24	247	911	79	4420
3	41	24	254	914	02	4510
3	42	24	251	918	51	4600
3	43	24	250	918	92	4690
3	44	24	250	920	89	4778
3	45	24	240	923	32	4864
2	46	24	229	925	82	4952
3	47	24	242	928	33	5040
3	48	24	250	930	42	5124
3	49	24	243	932	89	5210
3	50	24	246	935	17	5292
3	51	24	241	937	18	5378
3	52	24	242	939	55	5458
3	53	24	235	942	09	5542
3	54	24	240	944	49	5624
3	55	24	239	946	89	5704
3	56	24	236	948	87	5784
3	57	24	228	951	23	5862
3	58	24	227	953	86	5942
3	59	24	235	956	00	6024
3	60	24	191	958	33	6104
3	61	24	198	961	14	6184
3	62	24	225	963	38	6264
3	63	24	226	965	85	6342
3	64	24	228	967	87	6418

SCAN	AT TIME	94956	PERISCOPE AZ	AND EL	0 00	0 00	FOR/BACK	1
4	65	24	174	967	70	6420	-6	449
4	66	24	162	964	65	6344	-6	438
4	67	24	141	961	94	6266	-121	441
4	68	24	089	959	66	6188	0	448
4	69	24	070	957	59	6110	-6	458
4	70	24	077	955	10	6032	0	469
4	71	24	072	952	75	5952	0	414
4	72	24	063	950	31	5872	-6	498
4	73	24	069	947	89	5792	-3	430
4	74	24	059	945	43	5712	-3	441
4	75	24	057	943	18	5630	-3	498
4	76	24	057	941	01	5548	0	478
4	77	24	070	938	32	5464	-3	503
4	78	24	063	936	10	5380	-6	582
4	79	24	068	933	47	5298	-6	664
4	80	24	063	931	01	5210	-134	872
4	81	24	057	928	77	5126	3	1103
4	82	24	051	926	83	5040	-3	1428
4	83	24	067	924	20	4954	-134	1881
4	84	24	065	921	94	4870	3	2234
4	85	24	059	919	70	4784	-3	2502
4	86	24	089	917	21	4698	-134	2414
4	87	24	065	914	99	4612	3	2246
4	88	24	060	913	03	4528	-6	1771
4	89	24	079	910	42	4438	-6	1358
4	90	24	077	908	08	4346	0	1019
4	91	24	072	905	73	4256	-9	848

APPENDIX G. VERIFICATION OF EQUATION (A-1), APPENDIX A

We can find the temperature T in terms of width by substituting  $\frac{1}{2} I_{\max}$  for I in equation (A-1)

$$\text{where } I_{\max} = \left\{ \frac{M_c^2}{2\pi k T \lambda_o^2} \right\}^{\frac{1}{2}}$$

Substituting into equation (A-1) gives:

$$\frac{1}{2} \left\{ \frac{M_c^2}{2\pi k T \lambda_o^2} \right\}^{\frac{1}{2}} = \left\{ \frac{M_c^2}{2\pi k T \lambda_o^2} \right\}^{\frac{1}{2}} \exp \left\{ \frac{-M_c^2 (\Delta\lambda)^2}{2kT\lambda_o^2} \right\}$$

Solving for T, we obtain:

$$\ln\left(\frac{1}{2}\right) = \frac{-2.196825683 \times 10^{26}}{T} (\Delta\lambda)^2$$

$$T = 3.169349519 \times 10^{26} (\Delta\lambda)^2$$

where  $\Delta\lambda$  is the HWHM in meters.

To input  $\overset{\circ}{\text{A}}$  rather than m, we have:

$$T = 3.169349519 \times 10^6 (\Delta\lambda)^2$$

where  $\Delta\lambda$  is the HWHM specified in  $\overset{\circ}{\text{A}}$ .

To put the formula in terms of FWHM, multiply the numerical constant by  $(\frac{1}{2})^2 = \frac{1}{4}$



$$T = 7.92337 \times 10^5 (\Delta\lambda)^2$$

where  $\Delta\lambda$  is the FWHM in  $\text{\AA}$ .

This is consistent with the result of Wark [3] who obtained  $T = 7.89 \times 10^5 (\Delta\lambda)^2$ . To find a characteristic value for  $\Delta\lambda$ , we solve for HWHM ( $\text{\AA}$ ) from above:

$$\begin{aligned} (\Delta\lambda)^2 &= \frac{T}{3.16 \times 10^6} \\ \text{HWHM } (\text{\AA}) &= \Delta\lambda = \frac{T}{3.16 \times 10^6} \end{aligned}$$

$$\begin{aligned} \text{For } T = 1000 \text{ K, HWHM } (\text{\AA}) &= 0.01776 \\ &= 17.76 \text{ m\AA} \end{aligned}$$

APPROVAL

AURORAL THERMOSPHERE TEMPERATURES FROM OBSERVATIONS OF 6300 Å EMISSIONS

By John C. Bird, Gary R. Swenson, and Richard H. Comfort

The information in this report has been reviewed for technical content. Review of any information concerning Department of Defense or nuclear energy activities or programs has been made by the MSFC Security Classification Officer. This report, in its entirety, has been determined to be unclassified.



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