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Prepared by

M. W. CHEN, M. SCHULZ, L. R. LYONS, and D. J. GORNEY
 Space and Environment Technology Center
 Technology Operations

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THE AEROSPACE CORPORATION
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ION RADIAL DIFFUSION IN AN ELECTROSTATIC IMPULSE MODEL FOR STORMTIME RING CURRENT FORMATION

Margaret W. Chen, Michael Schulz, Larry R. Lyons, and David J. Gorney
Space and Environment Technology Center, The Aerospace Corporation

Abstract. Guiding-center simulations of stormtime transport of ring-current and radiation-belt ions having first adiabatic invariants $\mu \gtrsim 15$ MeV/G ($E \gtrsim 165$ keV at $L \sim 3$) are surprisingly well described (typically within a factor of $\lesssim 4$) by the quasilinear theory of radial diffusion. This holds even for the case of an individual model storm characterized by substorm-associated impulses in the convection electric field, provided that the actual spectrum of the electric field is incorporated in the quasilinear theory. Correction of the quasilinear diffusion coefficient D_{LL}^{ql} for drift-resonance broadening (so as to define D_{LL}^{rb}) reduced the typical discrepancy with the diffusion coefficients D_{LL}^{sim} deduced from guiding-center simulations of representative-particle trajectories to a factor ~ 3 . The typical discrepancy was reduced to a factor ~ 1.4 by averaging D_{LL}^{sim} , D_{LL}^{ql} , and D_{LL}^{rb} over an ensemble of model storms characterized by different (but statistically equivalent) sets of substorm-onset times.

Introduction

In a guiding-center simulation designed to investigate the access of particles to form the stormtime ring current [Chen *et al.*, 1992], hereafter referred to as paper 1, we have found that particles having energies $E \lesssim 110$ keV could have gained access to $L \sim 3$ essentially by direct convection [Lyons and Williams, 1980] from the plasma sheet, whereas the inward stormtime transport of particles that attain energies $E \gtrsim 165$ keV at $L \sim 3$ resembles radial diffusion [cf. Cornwall, 1968; Lyons and Schulz, 1989]. We made quantitative comparisons of diffusion coefficients obtained from the simulation with those derived from quasilinear diffusion theory [Fälthammar, 1965] in an effort to show that the transport is indeed appropriately described as radial diffusion. The diffusion coefficients constructed from simulated particle trajectories agreed surprisingly well (typically within a factor of ~ 4) with the predictions of quasilinear radial-diffusion theory, even over a model storm whose main-phase duration (taken as 3 hr) amounted to only about 3–30 drift periods for representative particles of interest.

In the present work we explore two refinements which are expected to produce theoretical diffusion coefficients that agree even better with those obtained from simulated particle trajectories. The first is a resonance-broadening correction [cf. Dungey, 1965] to the quasilinear diffusion theory of Fälthammar [1965] for an individual model storm. The second is an application of quasilinear theory to the mean diffusion coefficients obtained by simulating particle trajectories over an ensemble of 20 statistically similar model storms.

Field Model

We use the magnetic field model that is obtained by adding a uniform southward field ΔB to the geomagnetic dipole field. This superposition leads to the appearance

of a quasi-magnetopause at the boundary between closed and open field lines. The equation of a field line in this model is

$$[1 + 0.5(r/b)^3]^{-1}(r/R_E) \csc^2 \theta = \text{constant} \equiv L \quad (1)$$

where r is the geocentric distance, θ is the magnetic colatitude, R_E is the radius of the Earth, and $b = 1.5L^*R_E = 12.82 R_E$ is the radius of the equatorial neutral line. This value of b which is obtained by mapping the last closed field line (denoted L^*) to a colatitude of 20° on the Earth, corresponds to $|\Delta B| = 14.474$ nT and $L^* = 8.547$. The limit $b \rightarrow \infty$ ($L^* \rightarrow \infty$) would correspond to a purely dipolar B field. In this study, we consider only equatorially mirroring particles. The equatorial field intensity B_0 is given by $B_0 = (\mu_E/r^3) - 14.474$ nT, where $\mu_E = 3.05 \times 10^4$ nT- R_E^3 is the geomagnetic dipole moment. Further details of this field model are given by Schulz [1991, pp. 98–110].

We assume that the total electric field $\mathbf{E} = -\nabla\Phi_E$ is derivable from the scalar potential

$$\Phi_E = -\frac{V_0}{L} + \frac{V_0}{2} \left(\frac{L}{L^*}\right)^2 \sin \phi + \frac{\Delta V(t)}{2} \left(\frac{L}{L^*}\right) \sin \phi, \quad (2)$$

in which the three separate terms correspond to corotation ($V_0 = 90$ kV), the Volland-Stern model of quiescent convection ($V_0 = 50$ kV), and the time-dependent enhancement $\Delta V(t)$ associated with the stormtime convection, respectively. The time-varying term in the potential is assumed to vary as L rather than as L^2 because electric disturbances are expected to be less well shielded than steady-state convection by the inner magnetosphere.

We model the storm-associated enhancement $\Delta V(t)$ in the cross-tail potential drop,

$$\Delta V(t) = \sum_{i=1}^N \Delta V_i \exp[(t_i - t)/\tau] \theta(t - t_i) \quad (3)$$

where $\theta(t)$ is the unit step function ($\equiv 1$ for $t \geq 0$; $\equiv 0$ for $t < 0$), as a superposition of almost randomly occurring impulses that rise sharply and decay exponentially with a "lifetime" $\tau = 20$ min [cf. Cornwall, 1968]. The impulses represent the constituent substorms of a storm. The potential drop ΔV_i associated with any individual impulse is chosen randomly from a Gaussian distribution with a 200-kV mean and a 50-kV standard deviation. We assume that the N start times t_i in (3) are randomly distributed within a 3-hr time interval corresponding to the main phase of a storm, except that we impose a 10-min "dead time" (after each impulse onset) during which no subsequent impulse can start. This constraint imposes a realistic delay between the occurrences of consecutive impulse onsets. Without such a dead-time it would be possible for the next impulse to start immediately after the previous one, and this could lead to the build-up of unrealistically large cross-tail potentials. Further details of this model storm are given in paper 1.

We have constructed 100 such random storms so that on average there are 9 impulses per storm or 3 substorms/hr. We have randomly chosen one model storm for a detailed case study. Figure 1 shows the variation in

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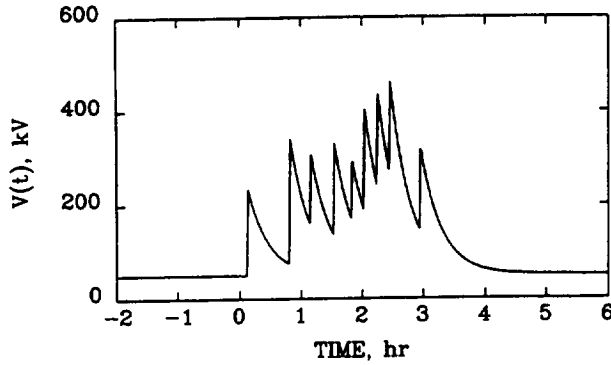


Fig. 1. Cross-tail potential $V(t)$ in our model storm.

cross-tail potential for this prototypical storm. The mean enhancement in the cross-tail potential drop for this particular storm over the time interval $t_1 < t < t_1 + 3$ hr is $\langle \Delta V(t) \rangle = 180$ kV.

Simulated Trajectories

In paper 1 we considered the stormtime transport of singly charged ions with various μ values by numerically tracing their guiding-center trajectories (Burlish-Stoer extrapolation method). We began by tracing the quiescent trajectories obtained by setting $\Delta V(t) \equiv 0$ in (2). The solid curves in Figure 2 illustrate steady-state trajectories of equatorially mirroring ions for a particular μ value (30 MeV/G). This corresponds to an energy of 335 keV at a geocentric radial distance of $r = 3R_E$. For detailed study we chose the quiescent drift shell that intersects the dusk meridian at $R \equiv r/R_E = 3$ for each of 12 μ values ($15 \text{ MeV/G} \lesssim \mu \lesssim 200 \text{ MeV/G}$). We selected this drift shell because it is representative of where particles need to be transported in order to form the stormtime ring current. We computed the quiescent ionic drift period τ_3 , which is 0.73 hr for $\mu = 30 \text{ MeV/G}$. We spaced 12 representative ions equally in time on the drift path

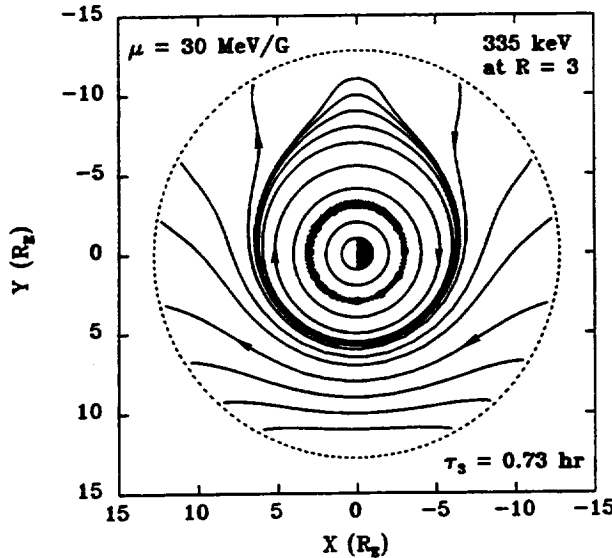


Fig. 2. Solid curves: quiet-time equatorial trajectories of singly charged ions having $\mu = 30 \text{ MeV/G}$. Outer dashed circle: neutral line at $r = b$. Small "filled" circles: the 12 representative ions' "final" positions on the steady-state drift path of interest. Stormtime trajectories computed in our time-reversed simulation span the stippled area.

of interest before starting the simulation. To investigate particle access to the drift shell of interest, we ran the simulation backward in time. Thus, the time-reversed trajectories (which span the stippled area in Figure 2) indicate where the particles must have been prior to the storm in order to have reached the final drift shell ($L \approx 3$) of interest. Indeed, we found a spread among the initial L values of the representative ions. Such particle transport among closed drift paths resembles radial diffusion and was typical for ions with $\mu \gtrsim 15 \text{ MeV/G}$ at $L \sim 3$. We also ran simulations forward in time, so as to follow the dispersal of ions from a common initial drift path. Not surprisingly, we found a qualitatively similar transport, although the particles followed in the time-forward simulations were implicitly different from those followed in the time-reversed simulations.

Diffusion and Quasi-Diffusion Coefficients

Since the quiescent drift shells are not circular (cf. Figure 2), we found it convenient to label drift shells (and thus to express the diffusion coefficients) in terms of the dimensionless third adiabatic invariant defined by Roederer [1970, p. 107] as

$$\frac{1}{L} \equiv \left| \frac{\Phi_B R_E}{2\pi \mu_E} \right| = \left[\frac{1}{2\pi} \oint \frac{d\phi}{L(\phi)} \right], \quad (4)$$

where Φ_B is the magnetic flux enclosed by that drift shell and $L(\phi)$ denotes the field-line label at longitude ϕ on the drift shell. We computed diffusion coefficients for each μ from the distribution of initial L values of the time-reversed trajectories (see Figure 2) by constructing the quantity

$$D_{LL}^{sim} \equiv \left(\frac{L^4}{24T} \right) \left[\sum_{i=1}^{12} (L_i^{-1} - L_f^{-1})^2 - \left[\sum_{i=1}^{12} (L_i^{-1} - L_f^{-1}) \right]^2 \right], \quad (5)$$

where L_i and L_f denote the drift-shell labels of the initial and final trajectories, respectively, of the 12 representative ions and $T (= 3 \text{ hr})$ denotes the duration of the main phase of the model storm. The quantity D_{LL}^{sim} is thus a measure of the variance among the initial third adiabatic invariants of particles situated on the final drift shell of interest. The result for each μ is plotted as an open circle in Figure 3. We also computed diffusion coefficients D_{LL}^{sim} from the time-forward simulations by interchanging the indices i and f in (5), and these results are shown as filled circles in Figure 3. The diffusion coefficients D_{LL}^{sim} obtained from simulations run forward and backward in time are not very different, 80% of them being within 24% of each other, although they pertain implicitly to different sets of particles.

We compare the diffusion coefficients obtained from the simulated trajectories with the resonant-particle formulation [Fälthammar, 1965] of radial-diffusion theory in which the diffusion coefficient is of the form

$$D_{LL}^{ql} = (L^6 R_E^4 / 4\mu_E^2) \hat{E} \left(\frac{\Omega_3}{2\pi} \right), \quad (6)$$

where $\hat{E}(\omega/2\pi)$ is the spectral-density of the (quasi-uniform) equatorial electric field in the inner magnetosphere and $\Omega_3/2\pi$ is the particles' quiescent drift frequency. When we substitute the spectral-density function for our model storm (see paper 1 for derivation) into (6), we obtain the quasilinear diffusion coefficient

$$D_{LL}^{ql} = \frac{\tau^2 L^6 R_E^2}{16T \mu_E^2 (L^*)^2} \sum_{i=1}^N \sum_{j=1}^N \frac{\Delta V_i \Delta V_j \cos[\Omega_3(t_j - t_i)]}{1 + \Omega_3^2 \tau^2} \quad (7)$$

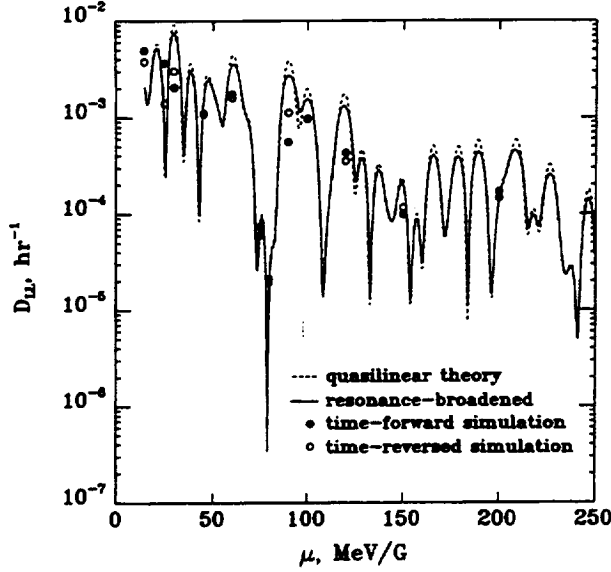


Fig. 3. Diffusion coefficients D_{LL}^{sim} obtained via (5) from time-reversed (open circles) and time-forward (filled circles) simulations, for comparison with the quasilinear diffusion coefficient D_{LL}^q (dashed curve) given by (7) as an implicit function of μ and with D_{LL}^b (solid curve), which represents a correction of the quasilinear diffusion coefficient for resonance-broadening effects.

in which correlations between the impulses lead to cross terms ($j \neq i$). If we neglect the cross terms, we recover essentially the standard diffusion coefficient of Cornwall [1968].

The quasilinear diffusion coefficient at $L \approx 3$, represented by the dashed curve in Figure 3, is not a very smooth function of μ . This is because the impulse onsets (i) associated with any individual storm modeled by (3) occur at specific (although randomly determined) times t_i so that the corresponding spectral density $\tilde{E}(\omega/2\pi)$ is not a very smooth function of frequency. The diffusion coefficients obtained from the simulations can be seen to agree surprisingly well with the quasilinear results despite the strong variability of D_{LL}^q with μ . Typically (i.e., in $> 80\%$ of the available comparisons) we find agreement within a factor of 4. Quasilinear theory even accounts for the μ values (e.g., $\mu = 75$ and 80 MeV/G) for which the diffusion coefficients computed from the simulation are especially small.

However, the diffusion coefficients obtained from the simulated trajectories do not show quite as much variability with μ as quasilinear theory predicts. This is not surprising since quasilinear theory postulates a perfectly sharp resonance at the quiescent drift frequency, whereas the simulated transport leads to an eventual spread among the quiescent drift frequencies of the representative ions for each μ . A rough estimate for the anticipated spread in $\Omega_3/2\pi$ is

$$\Delta\omega/2\pi \approx (D_{LL}T/2\pi^2)^{1/2} |(\partial\Omega_3/\partial L)_\mu| \quad (8)$$

since the mean-square spread in L is $2D_{LL}T$. An estimate for the diffusion coefficient (corrected for resonance-broadening effects) is thus

$$D_{LL} \approx \frac{L^6 R_E^4}{4\mu^2 \Delta\omega} \int_{\Omega_3 - (\Delta\omega/2)}^{\Omega_3 + (\Delta\omega/2)} \tilde{E}\left(\frac{\omega}{2\pi}\right) d\omega. \quad (9)$$

Since the frequency bandwidth given by (8) depends on D_{LL} , we have iterated between (9) and (8) until satisfactory convergence to the desired solution (called D_{LL}^b) is achieved. The result is plotted as a solid curve in Figure 3. We find that the inclusion of nonlinear resonance-broadening effects reduces the discrepancy between quasilinear theory with D_{LL}^{sim} through a smoothing of the sharp relative maxima and minima (with respect to μ). Corrections were typically $\sim 10 - 30\%$ at the relative maxima but were as much as 60% at the relative minima (e.g., near $\mu = 184$ MeV/G). It is not clear whether the remaining discrepancy (typically a factor $\lesssim 3$) between D_{LL}^b and D_{LL}^{sim} is intrinsic to this kind of study or attributable to the neglect of other nonlinear and/or quasilinear effects. A preliminary test suggests that the replacement of L^6 in (9) by its transport-averaged value $\langle L^6 \rangle$ would also be a relatively unimportant nonlinear correction. However, a quasi-linear correction for the variations of $L(\phi)$ and ϕ along a quiescent drift trajectory would introduce harmonic resonances ($\omega = m\Omega_3$) in (9) and diminish the weight of the fundamental drift resonance ($\omega = \Omega_3$). We hope to investigate this correction in a future study.

Diffusion Averaged over an Ensemble of Storms

It could be argued that quasilinear theory is more appropriately applied to an ensemble of model storms than to an individual storm. We tested this hypothesis by randomly choosing 20 different storms from among the 100 storms that we had constructed. We averaged the diffusion coefficients obtained from the simulations, standard quasilinear theory, and the resonance-broadened quasilinear theory over the 20 storms. The results are shown in Figure 4. The ensemble-averaged quasilinear diffusion coefficient \bar{D}_{LL}^q (dashed curve) and its resonance-broadened counterpart \bar{D}_{LL}^b (solid curve) are considerably smoother than D_{LL}^q and D_{LL}^b , respectively, for an individual storm. The ensemble-averaged diffusion coefficients \bar{D}_{LL}^{sim} from the time-reversed and time-forward simulation (open circles and filled circles, respectively) generally agree much

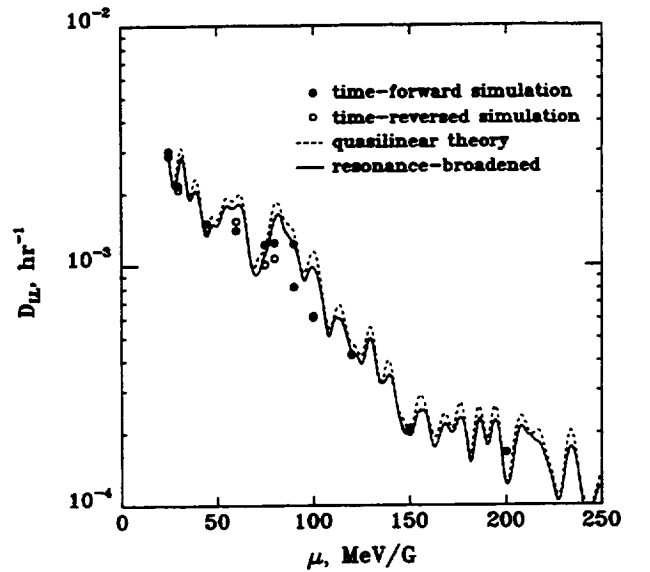


Fig. 4. Ensemble-averaged diffusion coefficients \bar{D}_{LL}^{sim} obtained from the time-reversed (open circles) and time-forward (filled circles) simulations, for comparison with quasilinear theory (\bar{D}_{LL}^q as a function of μ , dashed curve) and with \bar{D}_{LL}^b (solid curve), which represents ensemble-averaged diffusion coefficient corrected for resonance broadening effects.

better with the theoretical diffusion coefficients in Figure 4 than does D_{LL}^{sim} with the theoretical diffusion coefficients for an individual storm in Figure 3. Typically D_{LL}^{sim} and \bar{D}_{LL}^{q1} in Figure 4 agree within a factor of only 1.4. Agreement between \bar{D}_{LL}^{sim} and \bar{D}_{LL}^{rb} is even better (typically within a factor of 1.3).

Discussion and Summary

In this work we have considered two refinements of quasilinear radial diffusion theory for the purpose of interpreting guiding-center simulation results. We found in paper 1 that the quasilinear diffusion theory of Fälthammar [1965] already provides a surprisingly good description of charged-particle transport at first invariants $\mu \gtrsim 15$ MeV/G ($E \gtrsim 165$ keV) around $L = 3$. This holds even for the case of an individual model storm characterized by substorm-associated impulses in the convection electric field provided that the actual spectrum of the electric field is incorporated in the quasilinear theory.

Here we have found that a self-consistent correction of the quasilinear diffusion coefficient D_{LL}^{q1} for resonance-broadening effects accounts even better for the diffusion coefficients D_{LL} obtained from a simulation of particle trajectories (reducing the typical discrepancy from a factor ~ 4 to a factor ~ 3). Resonance broadening tends to reduce the variability of D_{LL} with μ by requiring that the electric spectral density $E(\omega/2\pi)$ be averaged over an appropriate bandwidth $\Delta\omega/2\pi$ surrounding the drift frequency $\Omega_3/2\pi$. The resonance-broadening correction turns out [Schulz, 1975] to be unimportant ($\lesssim 1\%$) when applied to the standard model [Cornwall, 1968] of radial diffusion in which the spectral density is regarded as a smoothly varying function of frequency. However, the spectral-density for any one of our model storms is not smooth at all. Indeed, it typically shows variations of at least 2 orders of magnitude over bandwidths $\Delta\omega/2\pi \sim 1$ hr $^{-1}$, whereas resonance-broadening theory typically requires $E(\omega/2\pi)$ to be averaged over a bandwidth $\Delta\omega/2\pi \sim 0.25$ hr $^{-1}$. The resonance-broadened correction is thus quite significant for our typical model storm.

Even better agreement between D_{LL}^{rb} and D_{LL}^{sim} is found when both are averaged over an ensemble of 20 model storms. Indeed, the typical discrepancy between \bar{D}_{LL}^{rb} and \bar{D}_{LL}^{sim} (and between \bar{D}_{LL}^{q1} and \bar{D}_{LL}^{sim}) is only a factor ~ 1.4 , mainly because \bar{D}_{LL}^{q1} and \bar{D}_{LL}^{rb} are much smoother functions of μ for an ensemble of storms than are D_{LL}^{q1} and D_{LL}^{rb} for an individual storm. The improved agreement achieved by ensemble-averaging is consistent with the usual interpretation of quasilinear theory (with or without the correction of quasilinear broadening) as a statistical description of charged-particle transport. What is surprising is that quasilinear theory provides a reasonably good description of the particle transport even for an individual storm.

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After having made the ensemble-average comparison between \bar{D}_{LL}^{sim} and \bar{D}_{LL}^{q1} , we learned of a recent similar study in which P. Riley and R. A. Wolf (preprint, 1991) have likewise found better agreement between diffusion coefficients obtained from simulated guiding-center trajectories with quasilinear diffusion theory by averaging over an ensemble of model storms than by making the comparison for an individual (in their case, actual) storm. However, they restricted their attention to particle energies $\lesssim 130$ keV at $L = 3$, whereas we have found the transport to be quasi-diffusive only for particles that attain energies $\gtrsim 165$ keV at $L = 3$.

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M. W. Chen, M. Schulz, L. R. Lyons, D. J. Gorney, The Aerospace Corporation, P.O. Box 92957, Los Angeles, CA 90009

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