

Monitoring thermal status of ecosystems with MODIS land-surface temperature and vegetation index products

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ABSTRACT

The global land-surface temperature (LST) and normalized difference vegetation index (NDVI) products retrieved from Moderate Resolution Imaging Spectroradiometer (MODIS) data in 2001 were used in this study. The yearly peak values of NDVI data at 5km grids were used to define six NDVI peak zones from -0.2 to 1 in steps of 0.2, and the monthly NDVI values at each grid were sorted in decreasing order, resulting in 12 layers of NDVI images for each of the NDVI peak zones. The mean and standard deviation of daytime LSTs and day-night LST differences at the grids corresponding to the first layer of NDVI images characterize the thermal status of terrestrial ecosystems in the NDVI peak zones. For the ecosystems in the 0.8-1 NDVI peak zone, daytime LSTs distribute from 0-35 °C and day-night LST differences distribute from -2 to 22 °C. The daytime LSTs and day-night LST differences corresponding to the remaining layers of NDVI images show that the growth of vegetation is limited at low and high LSTs. LSTs and NDVI may be used to monitor photosynthetic activity and drought, as shown in their applications to a flood-irrigated grassland in California and an unirrigated grassland in Nevada.

Keywords: MODIS, land-surface temperature, NDVI, ecosystem, photosynthetic activity, drought

1. INTRODUCTION

As a part of the NASA-centered international Earth Observing System (EOS), two MODIS instruments¹ have been launched to provide information for global studies of atmosphere, land, and ocean processes. The first one was launched 18 December 1999 on the morning platform called Terra, and the second one was launched 4 May 2002 on the afternoon platform called Aqua. The strengths of MODIS include its global coverage, high radiometric resolution and dynamic ranges, and accurate calibration in visible, near-infrared and thermal infrared bands. The Terra MODIS data have been used to generate science data products for more than two years and these science data products are available to the public free of charge. A consistent reprocessing has been made for Terra MODIS data since November 1999 to generate science data products (in version 3) in either provisional quality or validated quality. Science data products based on Aqua MODIS data will also be available after the testing and evaluation stage.

NDVI and LST are two of the MODIS land product suite.² The MODIS VI products have been designed for seasonal and interannual monitoring of the Earth's vegetation. The NDVI product maintains continuity with the NDVI data record from the Advanced Very High Resolution Radiometers (AVHRR).^{3, 4} Vegetation indices have been used to conduct land cover classifications,⁵ and derive biophysical properties of the vegetation such as fractional vegetation cover, biomass, leaf area index, fraction of absorbed photosynthetic active radiation and phenology.^{6, 7, 8}

LST is one of the key parameters in the physics of land-surface processes on regional and global scales, combining the results of all surface-atmosphere interactions and energy fluxes between the atmosphere and the ground.^{9, 10} The remotely sensed LST has been used in land cover and land-cover change analysis,^{11, 12} and in the production of the MODIS land cover product. It has also been used in monitoring drought and estimating surface soil moisture,^{13, 14} evaluating water requirements of wheat,¹⁵ and determining frosts in orange groves.¹⁶

The level-2 MODIS LST product (MOD11_L2) is retrieved with the generalized split-window LST algorithm¹⁷ from the calibrated radiance data of bands 31 and 32 data in the MODIS Level-1B 1km resolution (MOD021KM) product. Other inputs to this algorithm include MODIS geolocation product (MOD03), cloudmask product (MOD35), and atmospheric temperature and water vapor profile product (MOD07). The emissivities in bands 31 and 32 are estimated from land cover types¹⁸ in the MODIS land-cover product (MOD12Q) and snow cover product (MOD10) through look-up tables established by TIR BRDF (bidirectional reflectance distribution function) and emissivity modeling.¹⁹ The errors and uncertainties in the classification-based emissivities may be large in semi-arid and arid regions because of the large

temporal and spatial variations in surface emissivities and lack of knowledge on the emissivity variation with viewing angle. To resolve this problem, a physics-based day/night algorithm²⁰ was developed to retrieve surface spectral emissivity and temperature at 5km resolution for the MODIS LST level-3 MOD11B1 product from a pair of daytime and nighttime MODIS data in seven TIR bands, i.e., bands 20, 22, 23, 29, and 31-33. The accuracy of daily MODIS LST product has been validated in nineteen clear-sky cases with in-situ measurement data collected in field campaigns in 2000 and 2001. The MODIS LST accuracy is better than 1K in the range from 263K to 322K over Lake Titicaca in Bolivia, Mono Lake, Bridgeport grassland, and a rice field in Chico, California, Walker Lake and a silt playa in Railroad Valley, Nevada, in the atmospheric column water vapor range from 0.4 to 3.0cm.²¹ The MODIS LST product was also validated with in-situ TIR radiometer measurements in Lake Tahoe, CA/NV, in 27 clear-sky cases in 2000 and 2001 by Simon Hook, JPL, and it was found that the MODIS LST product has a small bias of 0.12K and the standard deviation of its differences from the in-situ values is 0.3K (personal communications). In mid July 2002, the MODIS LST product was also validated at the same accuracy with in-situ measurements in a soybean field near Greenville, Mississippi, in the atmospheric column water vapor range up to more than 4cm.

This paper presents the early results in monitoring thermal status of ecosystems with the MODIS LST and NDVI products.

2. MODIS DATA PRODUCTS USED IN THIS STUDY

Twelve sets of the 16-day MODIS VI product (MOD13A2) were ordered from the NASA EOS Data Gateway (<http://redhook.gsfc.nasa.gov/~imswwww/pub/imswelcome/>) in order to use the global NDVI data sets for year 2001. The file names and the starting dates of the periods of 16 days are listed in Table 1. There are 281 files in each set of the VI product to cover all the land of the Earth between latitudes 60 °S and 80 °N. The MODIS VI product provides NDVI data at the 1km-resolution grids in the integerized sinusoidal (ISIN) projection. The exact size of the 1km grid is 0.928km by 0.928km. The valid range of the NDVI data is from -2000 to 10000, giving NDVI values in the range of -0.2 to 1.0 after divided by a scale factor of 10000. We mosaiced all the files in each set and aggregated the NDVI data at 1km grids into 5km grids by taking average of all the NDVI data in the valid range. The NDVI values at 5km grids in the mosaic images are stored in 8-bit unsigned integer in the valid range of 0 to 240 with the lower limit 0 corresponding to NDVI value -0.2 and the higher limit 240 corresponding to NDVI value of 1.0. A quantization error of 0.0025 associated with the 8-bit unsigned integer is believed to be smaller than the error and uncertainty of the NDVI data. A fill value of 255 is given for grids in ocean and cloudy conditions. After this processing, we have 12 sets of global NDVI images representing monthly NDVI values in this study for the global except the polar regions.

For the whole year of 2001, the daytime and nighttime LSTs were taken from the daily 5km MODIS LST product (MOD11B1) that is also available from the EOS Data Gateway free of charge. The LST data are stored in 16-bit unsigned integer in the valid range of 7500 to 65535, giving LST values in the range of 150 to 1310.7K after multiplying a scale factor of 0.02 (K). A fill value of 0 is given for grids in ocean and cloudy conditions. We mosaiced the LSTs in 317 tiles of the MOD11B1 product into global LST data set each day. Then the daily global LST sets were composited and averaged to obtain the monthly daytime and nighttime LST data sets for 2001. The exact size of the 5km-resolution grids is 4.64km by 4.64km. One set of global LST data at 5km grids are stored in 4320 rows by 8640 columns. In the MODIS LST processing, only the MODIS data in clear-sky conditions at a confidence of 99% defined by the MODIS cloud-mask product (MOD35) are used to retrieve LST values. It is possible that there is no LST value for some grids that have valid NDVI values because a different definition of clear-sky pixels is used in the NDVI processing. A cirrus cloud may have an effect large enough to prevent LST retrieval at the 1K accuracy, but it is not necessarily so for the NDVI estimation. The overpass time of the MODIS nadir observations is around 10:30am local solar time in the daytime and around 10:30pm local solar time at night. The MODIS observations at off-nadir viewing angles are made before or after the overpass time in term of the local solar time for the spots being observed. After the compositing and averaging in the generation of the monthly LST data, the daytime and nighttime LSTs represent the averaged LST values in clear-sky conditions around 10:30am and 10:30pm local solar time, respectively. The MODIS LSTs represent the kinetic temperature determined by the TIR radiations from the surfaces of all components within the pixel or grid, including vegetation, soil, water body, and whatever being observed by the MODIS instrument.

3. NDVI PEAK ZONE APPROACH

Because of the great diversity in the global ecosystems, the vegetation at the grids in a 16-day or monthly NDVI image may be in different stage of its phenology, some in growing phase and some in drying phase. In order to highlight the NDVI change at the global scale, a NDVI peak zone approach is applied in the following steps: (1) we separate the NDVI valid range into six sub-ranges, i.e., -0.2 to 0, 0 to 0.2, 0.2 to 0.4, 0.4 to 0.6, 0.6 to 0.8, and 0.8 to 1.0; (2) to find the maximum NDVI value in the 12 monthly NDVI images for each grid, assign it to one of the six sub-ranges according to its maximum NDVI value, and keep the month as its first index, after processing all grids, we get six images of NDVI peak zones as shown in Fig. 1, each corresponding to one of the six sub-ranges; (3) for all the grids in each NDVI peak zone, the NDVI values in 12 months are sorted in decreasing order and keep the corresponding months in the index images, resulting in 12 layers of NDVI images for each NDVI peak zone; (4) according to the index images, corresponding 12 layers of daytime and nighttime LST images are constructed for each NDVI peak zone. Fig. 1 shows the spatial distributions of the NDVI in the six NDVI peak zones. The images in (a) for the 0.8 - 1.0 zone, (b) for the 0.6 to 0.8 zone, and (c) for the 0.4 - 0.6 zone, were enhanced with the histogram equalization method for better visualization. As shown in Fig. 1 (a), NDVI peak values larger than 0.8 are distributed in tropical forests in South America and Africa and Far East, east part of the North America, northern parts of Europe and Asian, and the east part of Asian. These areas are neighboring by the 0.8 - 1.0 NDVI peak zone shown in Fig. 1 (b). Semi-arid areas such the mid west of US and the interior of Australia are shown in Fig. 1 (c) and (d). Desert areas are shown in Fig. 1 (e) for the 0 - 0.2 NDVI peak zone. The last zone (-0.2 to 0) is distributed mainly in Greenland.

4. CHARACTERISTICS OF THE THERMAL STATUS OF THE ECOSYSTEMS

From the 12 layers of NDVI and corresponding index images, daytime and nighttime LST images, we calculated the histograms of the NDVI, daytime LSTs and day-night LST differences for the five NDVI peak zones (covering NDVI from 0.0 to 1.0). Fig. 2 shows the histograms in layers 1, 4, and 7. The histograms in layers 1, 4, and 7 are shown in columns 1, 2, and 3, respectively. In layer 4 of the NDVI image, the time difference between the month with NDVI peak and the month with the 4th largest NDVI value in 12 months may be 2-3 months, varying with grids. Similarly, the time difference in layer 7 of the NDVI image may be 3-6 months. Comparing the NDVI histograms in layers 1, 4 and 7 (from left to right in the top row) clearly shows the NDVI changes in the five NDVI peak zones. Comparing the histograms of daytime LSTs in the middle row shows that the daytime LSTs change toward low and high sides as the NDVI reduces. This feature in the histograms of daytime LSTs indicates that the growth of vegetation is limited at low LSTs due to cold weather and insufficient photosynthetically active radiation and at high LSTs due to strong evapotranspiration. From the histograms of the day-night LST differences in the bottom row, we can see that the daynight LST difference becomes slightly larger as the NDVI reduces. Because an increase in the day-night LST difference is related to a reduce in the apparent thermal inertia of the Earth surface in general and the apparent thermal inertia reduces as the soil gets dry^{22, 23}, the day-night LST difference may be used to estimate the soil moisture and monitor drought.

Means and standard deviations of the NDVI, daytime LSTs and day-night LST differences were also calculated. They are given in Table 2, and the number of 5km grids with LSTs is given in the last column. The number of 5km grids with LSTs may be different because there may be some grids without LST values due to persistent clouds in one month in some areas. When using the mean and standard deviation values in Table 2, beware that the shapes of the histograms in Fig. 2 are different from a normal distribution.

Analysis indicates that although there are weak correlations between NDVI and LST they are independent variables of the ecosystems at the global scale.

5. POTENTIAL APPLICATIONS

Now we try to apply MODIS LST and NDVI products in applications for monitoring thermal status of ecosystems, photosynthetic activity and drought. We have conducted field measurements with TIR radiometers in a grassland at 1972m above sea level in Bridgeport, California since 2000. The size of the grasslands in the Bridgeport valley is about 5km by 5km. The grasslands are irrigated during April 15 and September 15 in shadow canals with water originated from snow-melt in the Sierra Nevada just north of Mono Lake, approximately 30km from the valley. This irrigated land is used for cattle grazing during the summer months. For a comparison, we also look at an unirrigated grassland at

1840m above sea level near Ione, Nevada. This grassland is about 5km wide and 15km long on a gentle slope by the west edge of the Toiyabe National Forest, and approximately 160km away from Bridgeport Valley.

Fig. 3 shows the NDVI, daytime LST and day-night LST difference values for the central portions of the flood-irrigated grassland in Bridgeport and the unirrigated grassland in Nevada in the spring and summer months of 2001. These values were taken from 1km MODIS LST and NDVI products. From this figure, we can get the following insights: (1) In the Bridgeport site the dips in the daytime LSTs and the day-night LST differences are caused by irrigation. (2) The NDVI increases when the daytime LSTs are in the range up to 35 °C in the spring and summer months. The NDVI in the irrigated grassland increases until the end of June. But the NDVI in the unirrigated grassland increases only in April and May. (3) The cooler daytime LSTs and smaller day-night LST differences in early July (julian day 188) in both sites indicate that there was a major storm event in the area at least from Bridgeport, CA to Ione, NV. This precipitation made a slight increase in the NDVI over the unirrigated grassland. The smaller day-night LST difference corresponds to the increase in the soil moisture condition in the unirrigated grassland. (4) The NDVI decreases as the daytime LST is above 35 °C. This example shows that MODIS LST and NDVI products can be used to monitor thermal status of the ecosystems, photosynthetic activity and drought.

The NDVI peak zone approach was also applied to latitude zones, for example, in the zone between latitudes 30-45 degree N that covers most of the USA, the Mediterranean region, Mid East, China, and Japan. We got similar NDVI and day-night LST difference histograms as shown in Fig. 2, but the daytime LST is distributed in smaller ranges and the standard deviations are also smaller. This NDVI peak zone approach can also be used in regions so that we can get the vegetation and thermal characteristics of the regional ecosystems.

When multi-year MODIS LST and NDVI products are available in the near future, we can get their annual variations in different ecosystems for better global and regional change studies.

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REFERENCES

1. V. Salomonson, W. Barnes, P. Maymon, H. Montgomery, and H. Ostrow, "MODIS: advanced facility instrument for studies of the Earth as a system," *IEEE Trans. Geosci. Remote Sens.*, vol. 27, no. 2, pp. 145-153, 1989.
2. C. O. Justice, E. Vermote, J. R. G. Townshend, R. Defries, D. O. Roy, D. K. Hall, V. V. Salomonson, J. L. Privette, G. Riggs, A. Strahler, W. Lucht, R. B. Myneni, K. Knyazikhin, S.W. Running, P. R. Nemani, Z.Wan, A. R. Huete, W. van Leeuwen, R. E. Wolfe, L. Giglio, J.-P. Muller, and Y. Knyazikhin, M. J. Barnsley, "The Moderate Resolution Imaging Spectroradiometer (MODIS): land remote sensing for global change research," *IEEE Trans. Geosci. Remote Sensing*, vol. 36, pp. 1228-1249, 1998.
3. W. J. D. van Leeuwen, A. R. Huete, and T. W. Laing, "MODIS vegetation index compositing approach: a prototype with AVHRR data," *Remote Sens. Environ.*, vol. 69, pp. 264-280, 1999.
4. T. Miura, A. R. Huete, H. Yoshioka, and B. N. Holben, "An error and sensitivity analysis of atmospheric resistant vegetation indices derived from dark target-based atmospheric correction," *Remote Sens. Environ.*, vol. 78, pp. 284-298, 2001.
5. J. T. Townshend, C. Justice, W. Li, C. Gurney, and J. McManus, "Global land cover classification by remote sensing: present capabilities and future possibilities," *Remote Sens. Environ.*, vol. 35, pp. 243-255, 1991.
6. G. Asrar, M. Fuchs, E. T. Kanemasu, and J. L. Hatfield, "Estimating absorbed photosynthetic active radiation and leaf area index from spectral reflectance in wheat," *Agron. J.*, vol. 76, pp. 300-306, 1984.
7. C. Justice, J. R. G. Townshend, B. N. Holben, and C. J. Tucker, "Analysis of the phenology of global vegetation using meteorological satellite data," *Int. J. Remote Sens.*, vol. 6, pp. 1271-1318, 1985.
8. P. C. Sellers, "Canopy reflectance, photosynthesis and transpiration," *Int. J. Remote Sens.*, vol. 6, pp. 1335-1372, 1985.

9. H. Mannstein, "Surface energy budget, surface temperature and thermal inertia," in *Remote Sensing Applications in Meteorology and Climatology*, ed. R. A. Vaughan and D. Reidel, NATO ASI Ser. C: Math. Phys. Sci. Vol. 201, pp. 391-410, Dordrecht, Netherlands: A Reidel Publishing Co., 1987.
10. P. J. Sellers, F. G. Hall, G. Asrar, D. E. Strebel, and R. E. Murphy, "The first ISLSCP Field Experiment (FIFE)," *Bull. Amer. Meteorol. Soc.*, vol. 69, no. 1, pp. 22-27, 1988.
11. D. Ehrlich and E. F. Lambin, "The surface temperature-vegetation index space for land cover and land-cover change analysis," *Int. J. Remote Sens.*, vol. 17, no. 3, pp. 463-487, 1996.
12. E. F. Lambin and D. Ehrlich, "Land-cover changes in sub-Saharan Africa (1982-1991): Application of a change index based on remotely sensed surface temperature and vegetation indices at a continental scale," *Remote Sens. Environ.*, vol. 61, no. 2, pp. 181-200, 1997.
13. C. M. Feldhake, D. M. Glenn, and D. L. Peterson, "Pasture soil surface temperature response to drought," *Agron. J.*, vol. 88, no. 4, pp. 652-656, 1996.
14. T. R. McVicar and D. L. B. Jupp, "The current and potential operational uses of remote sensing to aid decisions on drought exceptional circumstances in Australia: a review," *Agricultural Systems*, vol. 57, no. 3, pp. 399-468, 1998.
15. R. D. Jackson, R. J. Reginato, and S. B. Idso, "Wheat canopy temperature: a practical tool for evaluating water requirements," *Water Resour. Res.*, vol. 13, pp. 651-656, 1977.
16. V. Caselles and J. A. Sobrino, "Determination of frosts in orange groves from NOAA-9 AVHRR data," *Remote Sens. Environ.*, vol. 29, no. 2, pp. 135-146, 1989.
17. Z. Wan and J. Dozier, "A generalized split-window algorithm for retrieving land-surface temperature from space," *IEEE Trans. Geosci. Remote Sens.*, vol. 34, no. 4, pp. 892-905, 1996.
18. W. C. Snyder, Z. Wan, Y. Zhang, and Y.-Z. Feng, "Classification-based emissivity for land surface temperature measurement from space," *Int. J. Remote Sens.*, vol. 19, no. 14, pp. 2753-2774, 1998.
19. W. Snyder and Z. Wan, "BRDF models to predict spectral reflectance and emissivity in the thermal infrared," *IEEE Trans. Geosci. Remote Sens.*, vol. 36, no. 1, pp. 214-225, 1998.
20. Z. Wan and Z.-L. Li, "A physics-based algorithm for retrieving land-surface emissivity and temperature from EOS/MODIS data," *IEEE Trans. Geosci. Remote Sens.*, vol. 35, no. 4, pp. 980-996, 1997.
21. Z. Wan, Y. Zhang, Q. Zhang, and Z.-L. Li, "Validation of the land-surface temperature products retrieved from Terra Moderate Resolution Imaging Spectroradiometer data," *Remote Sens. Environ.*, in press 2002.
22. J. C. Price, "On the analysis of thermal infrared imagery: the limited utility of apparent thermal inertia," *Remote Sens. Environ.*, vol. 18, pp. 59-73, 1985.
23. A. B. Kahle, "Surface emittance, temperature, and thermal inertia derived from Thermal Infrared Multispectral Scanner (TIMS) data for Death Valley, California," *Geophysics*, vol. 52, no. 7, pp. 858-874, 1986.

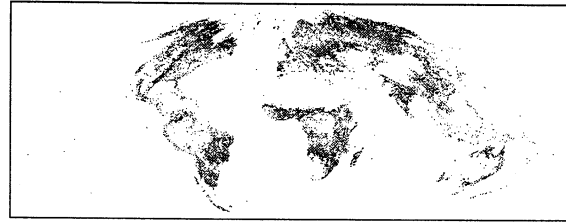
TABLE I. A list of the 16-day NDVI data sets used in this study.

file name	starting date	file name	starting date
MOD13A2.A2001001*.003*.hdf	01/01/2001	MOD13A2.A2001193*.003*.hdf	07/12/2001
MOD13A2.A2001033*.003*.hdf	02/02/2001	MOD13A2.A2001225*.003*.hdf	08/13/2001
MOD13A2.A2001065*.003*.hdf	03/06/2001	MOD13A2.A2001257*.003*.hdf	09/14/2001
MOD13A2.A2001097*.003*.hdf	04/07/2001	MOD13A2.A2001273*.003*.hdf	09/30/2001
MOD13A2.A2001129*.003*.hdf	05/09/2001	MOD13A2.A2001305*.003*.hdf	11/01/2001
MOD13A2.A2001145*.003*.hdf	05/29/2001	MOD13A2.A2001337*.003*.hdf	12/03/2001

Fig. 1 Spatial distributions of the NDVI in the six NDVI peak zones.
Note that the images in (a) - (c) were enhanced with the histogram equalization method.



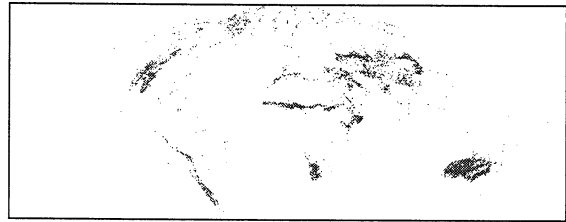
(a) 0.8 - 1.0



(b) 0.6 - 0.8



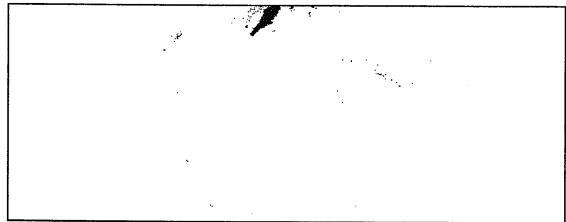
(c) 0.4 - 0.6



(d) 0.2 - 0.4



(e) 0.0 - 0.2



(f) -0.2 - 0.0

Fig. 2, Histograms of the NDVI (a), daytime LST (b), and day-night LST difference (c) in layers 1, 4, and 7 of NDVI images.

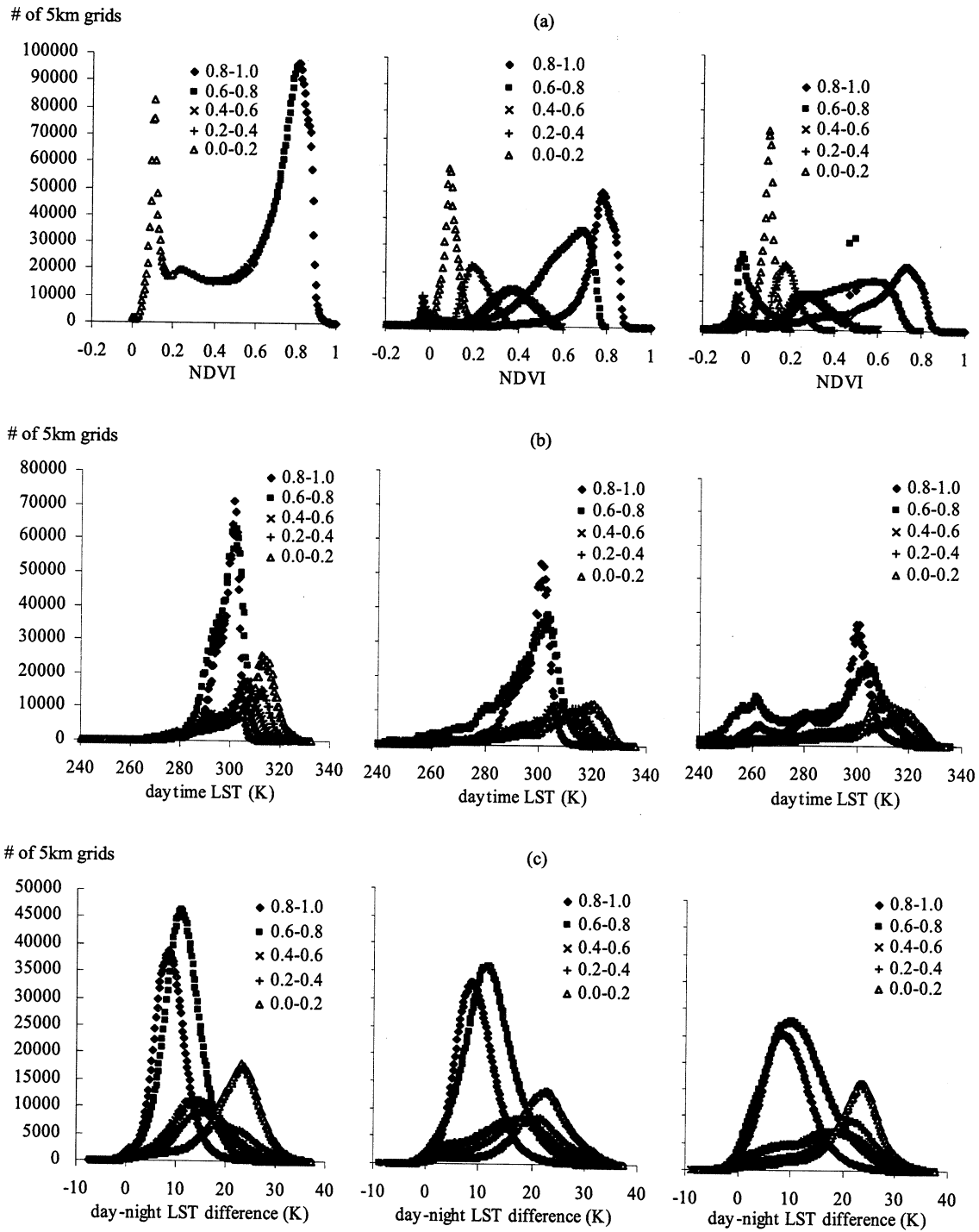


TABLE II. Mean and standard deviation () values of NDVI, daytime LSTs and day-night LSTs in five NDVI peak zones.

Layer	NDVI		daytime LST		day-night LST difference		number of grids with LSTs
	mean	()	mean (°C)	() (°C)	mean (°C)	() (°C)	
in the NDVI peak zone of 0.8 to 1.0							
1	0.8429	0.0284	24.96	5.38	8.724	3.552	1476258
2	0.8136	0.0451	24.79	5.53	8.879	3.599	1481741
3	0.7833	0.0680	24.14	6.52	9.153	3.872	1472496
4	0.7494	0.0951	23.24	7.92	9.370	4.096	1461050
5	0.7069	0.1263	22.11	9.80	9.591	4.402	1447721
6	0.6502	0.1654	20.16	12.16	9.499	4.578	1430770
7	0.5790	0.2053	17.83	14.94	9.334	4.791	1415002
in the NDVI peak zone of 0.6 to 0.8							
1	0.7239	0.0553	25.08	6.47	11.213	3.928	2026743
2	0.6833	0.0753	24.83	7.10	11.561	4.106	2037299
3	0.6256	0.1181	23.18	9.70	11.911	4.652	2023836
4	0.5580	0.1700	20.76	13.06	11.867	5.025	2020403
5	0.4891	0.2064	18.61	15.73	11.817	5.353	2023629
6	0.4196	0.2281	16.10	17.99	11.498	5.529	2024372
7	0.3489	0.2363	13.40	20.56	11.185	5.940	2028772
in the NDVI peak zone of 0.4 to 0.6							
1	0.5075	0.0586	27.42	9.10	14.247	5.552	708130
2	0.4570	0.0765	27.78	9.50	14.807	5.712	711114
3	0.3914	0.1116	26.11	13.05	15.084	6.236	712398
4	0.3249	0.1450	23.39	17.52	14.887	6.937	709320
5	0.2804	0.1535	22.24	18.73	15.106	7.149	707539
6	0.2495	0.1530	21.04	19.42	14.997	7.183	702404
7	0.2268	0.1524	20.01	20.27	14.883	7.436	702083
in the NDVI peak zone of 0.2 to 0.4							
1	0.2955	0.0572	30.22	10.62	16.826	6.294	698779
2	0.2624	0.0620	30.54	11.21	17.287	6.346	698010
3	0.2260	0.0764	29.25	14.45	17.519	6.749	696510
4	0.1958	0.0857	27.66	17.56	17.671	7.212	695923
5	0.1775	0.0874	27.56	18.26	18.076	7.259	694568
6	0.1644	0.0869	27.81	18.55	18.352	7.296	688912
7	0.1530	0.0878	27.37	19.42	18.279	7.378	677752
in the NDVI peak zone of 0.0 to 0.2							
1	0.1149	0.0372	33.64	12.36	20.913	6.008	983021
2	0.1033	0.0376	32.45	12.84	21.294	6.058	963079
3	0.0941	0.0393	32.92	14.75	21.391	6.255	936850
4	0.0868	0.0427	32.92	17.08	20.876	6.546	829195
5	0.0821	0.0464	32.90	18.10	20.900	6.736	758033
6	0.0823	0.0461	32.94	17.26	21.131	6.500	830555
7	0.0817	0.0445	32.18	16.70	21.009	6.342	854213

Fig. 3. Comparisons of NDVI (a), daytime LST and day-night LST difference in an irrigated grassland in Bridgeport (b) and a non-irrigated grassland in NV (c) in 2001.

