Space, Atmospheric, and Terrestrial Radiation Environments

J. L. Barth, Senior Member, IEEE, C. S. Dyer, E. G. Stassinopoulos, Member, IEEE

Abstract

The progress on developing models of the radiation environment since the 1960s is reviewed with emphasis on models that can be applied to predicting the performance of microelectronics used in spacecraft and instruments. Space, atmospheric, and ground environments are included. It is shown that models must be adapted continually to account for increased understanding of the dynamics of the radiation environment and the changes in microelectronics technology. The IEEE Nuclear and Space Radiation Effects Conference is a vital forum to report model progress to the radiation effects research community.

I. INTRODUCTION

Topics in radiation environment research cover a broad subject matter because radiation exists throughout the universe originating from many sources and with varying intensities. A survey of the proceedings from the Nuclear Space and Radiation Effects Conference (NSREC), the December Issue of the IEEE Transactions on Nuclear Science (TNS) and TNS Special Issues, shows that radiation environment topics at NSREC are predicated by 1) critical radiation effects and mechanism issues, 2) new atmospheric and space radiation measurements, and 3) availability of resources to transition new findings from basic science research to methods for quantitative predictions of the environment. The survey also reveals that many of the environment papers presented at NSREC were from authors who do not use the proceedings from NSREC as their primary forum for publication and that the NSREC papers were heavily referenced with papers from science journals. This implies that, as new issues in radiation effects were identified, the NSREC community borrowed knowledge from ongoing work in the science research community.

The “environments” session is a relatively new feature at the NSREC. The Space and Terrestrial Environments session was not part of the NSREC program until 1999. From 1991 through 1997, the environment papers were in the Spacecraft and Effects session. In the 1970s and 1980s environment papers were placed in the session of the effects issue that the environment topic addressed.

II. OVERVIEW OF THE RADIATION ENVIRONMENT

First a brief overview of space, atmospheric, and ground radiation environments will be presented. More complete descriptions of the radiation environments can be found in References [1], [2], [3], [4], [5].

A. Space Radiation Environments

The natural space radiation environment can be classified into two populations, 1) the particles trapped by planetary magnetospheres in “belts”, including protons, electrons, and heavier ions and 2) transient particles which include protons and heavy ions of all of the elements of the periodic table. The transient radiation consists of galactic cosmic ray (GCR) particles and particles from solar events, such as, coronal mass ejections and flares. These two types of solar eruptions produce energetic protons, alpha particles, heavy ions, and electrons that are orders of magnitude higher in abundance than the background GCRs.

Table 1 lists the maximum energy of space radiation particles. The table shows that much of the environment is high energy; therefore, shielding is not effective for many radiation source environments. When modeling, particles are treated as isotropic and omnidirectional with the exception of plasma, low altitude trapped protons (<500 km), and cosmic radiation on the ground.

Manuscript received March 21, 2003. This work was supported in part by the NASA NEPP/ERC Project and the Defense Threat Reduction Agency.

Janet L. Barth, E.G. Stassinopoulos are with the NASA Goddard Space Flight Center, Greenbelt, MD 20771.

C.S. Dyer is with QinetiQ, Farnborough, UK.
TABLE 1: MAXIMUM ENERGIES OF PARTICLES

1) Trapped Populations

James Van Allen is credited with discovering the trapped proton and electron regions around the Earth. Fig. 1 is an artist’s drawing, which shows the belt-like structure of these particle regions. The tilt of the Earth’s magnetic pole from the geographic pole and the displacement of the magnetic field from the center cause a dip in the field over the South Atlantic Ocean, resulting in a bulge in the underside of the inner belt. This region (~300 to ~1200 km), not shown in Fig. 1, is called the South Atlantic Anomaly (SAA). In spite of the SAA’s reputation for plaguesing spacecraft, the flux levels there are actually much lower than those at higher altitudes. The E > 30 MeV proton fluxes peak at approximately 2,500 km altitude at the equator. The electrons are trapped into two regions, the inner and outer zones. The E > 2 MeV peak electron fluxes at the equator are at approximately 2,500 km altitude in the inner zone and at 20,000 km altitude in the outer zone. Heavy ions are also trapped in planetary magnetic fields. For most shielded spacecraft systems, the abundances of these ions at energies high enough to penetrate spacecraft materials are too low to be a dominant factor in single event effects rates.

The trapped particle levels and locations are highly dependent on particle energy, altitude, inclination, and activity level of the sun and are highly dynamic. The slot region population, between the inner and outer zones (2 < L' < 2.8), and the outer zone (L >2.8) population can increase above averages by several orders of magnitude due to changes in the magnetosphere induced by solar and magnetic storms. Fig. 2, a plot of measurements of trapped electrons over a 1-year period, shows changes in the extremely dynamic outer zone and the slot region filling periodically with storm electrons. Due to their complex distribution and dependence on long- and short-term solar variability, the trapped particle populations are difficult to model and forecast.

The minimum requirement for the existence of a planetary radiation belt is that the planet’s magnetic dipole must be sufficiently great to arrest the flow of the solar wind before the particles reach the top of the atmosphere where the particles will lose their energy due to collisions. Venus, Mars, and possibly Pluto do not have magnetospheres and, therefore, cannot support particle trapping. The magnetic fields of some of the other planets are similar to the Earth’s, however, they vary in strength. Mercury has a weak magnetic field so it is expected it has a trapped particle population proportionally lower than that of the Earth. Saturn, Uranus, and Neptune have magnetic fields with similar strength to that of the Earth but measurements indicate that the intensities of the trapped radiation environments of Saturn, and Uranus are much lower than the Earth’s and do not pose serious problems to the design of spacecraft systems. On the other hand, Jupiter’s enormous magnetic dipole (428,000 compared to 30,760 nanoTesla for Earth) can support an intense particle environment. Its magnetosphere is the largest object in the solar system. Measurements have shown that the radiation environment is considerably more intense than the Earth’s and is more extensive, therefore, mission planning for spacecraft that will spend even short times in the trapping regions of Jupiter must include careful definitions of the radiation environment. The Probos probe showed that Mars has a radiation environment, however, it is due to the thin atmosphere of Mars, which allows interplanetary GCRs and solar particles to penetrate to the surface. Interaction of these particles with the atmosphere produces neutrons, which penetrate to the planetary surface and then reflect back.

2) Transient Populations

The GCR population is a continuously present, slowly varying population of ions from all elements of the periodic table. The levels of GCRs are modulated by the 11-year solar cycle with the peak GCR populations occurring at solar minimum. Superimposed on the GCR levels are unpredictable, sudden rises in the flux levels due solar energetic particles (SEPs) from solar storms. Galactic and solar particles have free access to spacecraft outside of the magnetosphere. Because the transient particles penetrate the Earth’s magnetosphere, they can reach near-Earth orbiting spacecraft and are particularly hazardous to satellites in polar, highly elliptical, and geostationary (GEO) orbits. Fig. 3 shows measurements from the IMP-8 spacecraft, which illustrates variations in the levels of carbon, oxygen, and nitrogen ions from GCR and solar sources over a 20-year period. The slowly varying, low-level GCR background population is seen to be approximately anti-correlated with the sunspot number. This is because the GCRs originate outside of our solar system and must “fight” against the solar wind to reach us. Fig. 3 also shows the solar particle events seen as spikes superimposed on top of the GCR measurements. These are the sudden increases in particle populations due to coronal mass ejections and/or solar flares.

B. Atmospheric Environments

As cosmic ray and solar particles enter the top of the Earth’s atmosphere, they are attenuated by interaction with nitrogen and oxygen atoms. The result is a “shower” (Fig. 4) of secondary particles and interactions created through the attenuation process.

Dipole shell parameter, L, and rings of constant magnetic field strength are the two coordinates used to map trapped particles. L is most simply described as the value that marks the particle drift shells by their magnetic equatorial distance from the center of the Earth.
performed vacuum chamber experiments to study when the altitude ranges from sea level to mountainous regions.

The variation in the neutron density is 1/500 of the peak flux. Our knowledge of neutron levels comes from balloon, aircraft, and ground based measurements. The energies of neutrons in the atmosphere reach energy levels of 100s of MeV. Taber and Normand [6] give an overview of the neutron environment in the atmosphere.

C. Radiation on Earth

Both natural and man-made radiation are present on the Earth. Table 2 lists these sources with their annual dose equivalents in units of milliSievert. The sources that are most important in producing effects in microelectronics are terrestrial cosmic rays, which induce single event effects. The radiation effects on materials and microelectronics in nuclear facilities present special radiation effects challenges.

Cosmic radiation on the ground is from the products of 6th and 7th generation interactions of GCR and solar particles in the atmosphere. The population has a high vertical directionality. Variations in the primary population intensities cause most of the variations observed in the secondary neutron and proton levels. Levels rise and fall in the same 11-year solar cycle that modulates the GCRs. Sea-level cosmic rays are reduced by 30% during solar active periods. However, during large solar particle events, ground level enhancements (GLEs) of the terrestrial cosmic rays can increase the levels by as much as 500% depending on location and the size of the event [7]. Fig. 5 shows the components of cosmic rays at New York City as a function of energy. Studies have shown that the variation in the neutron flux level is measurable when the altitude ranges from sea level to mountainous regions.

TABLE 2: ANNUAL DOSE FROM GROUND RADIATION IN mSv [8]

<table>
<thead>
<tr>
<th>Source</th>
<th>Dose Equivalent</th>
</tr>
</thead>
<tbody>
<tr>
<td>Dust particles bearing radon daughters</td>
<td>0.005 mSv</td>
</tr>
<tr>
<td>Naturally occurring radioactive elements (primarily potassium-40) other than radon daughters that find their way into our bodies</td>
<td>0.001 mSv</td>
</tr>
</tbody>
</table>

III. Trapped Radiation Belts

Birkeland theorized the existence of radiation trapping in planetary magnetospheres in 1895 when he performed vacuum chamber experiments in 1895 when he performed vacuum chamber experiments to study neutron levels from sea level to mountainous regions.

It was also during the early 1960s that spacecraft electronics were found to be unreliable. Problems from differential charging from the solar wind and from noisy data transmission to the Earth from soft fails were noted. These problems were largely dealt with by building redundancy into systems. As the first scientific satellites were being launched in the late 1950s and early 1960s, the USSR and US also detonated nuclear devices at altitudes above 200 kilometers. The most dramatic of these tests was the US Starfish detonation on July 9, 1962. Ten known satellites were lost because of radiation damage, some immediately after the explosion [11]. The Starfish explosion injected enough fission spectrum electrons with energies up to 7 MeV to
increase the fluxes in the inner Van Allen belt by at least a factor of 100. Effects were observed out to 5 Earth radii. The Starfish electrons that became trapped (modeled by Teague and Stassinopoulos [12]) dominated the inner zone environment (~2.8 Earth radii at the equator) for five years and were detectable for up to eight years in some regions.

The production of enhanced radiation levels by the Starfish explosion and others and the ensuing problem of shortened spacecraft lifetimes emphasized the need for a uniform, quantitative description of the trapped particle environment. Wilmott Hess of the National Aeronautics and Space Administration's (NASA's) Goddard Space Flight Center (GSFC) developed the first empirical models of the trapped radiation belts. Using data from several satellites, he began constructing quantitative radiation models for inner zone protons and electrons. These models were designated as P1, P2, etc. and E1, E2, etc. Starting in 1962 and continuing through the late 1960s, several series of satellites were launched with instruments designed to measure the effects of Starfish, providing a large volume of particle data. In late 1963, James Vette of Aerospace Corporation and later of NASA/GSFC was appointed to lead a trapped radiation environment modeling program jointly funded by NASA and the United States Air Force (USAF). At that time, there were several groups actively involved in trapped particle measurements, including Aerospace Corporation, Air Force Cambridge Research Laboratory (now Air Force Research Laboratory), Johns Hopkins/Applied Physics Laboratory, Bell Telephone Laboratories, GSFC, Lawrence Livermore Laboratory, Lockheed Missile and Space Corporation, the University of California at San Diego, and the University of Iowa. Each organization agreed to make its measurements available to the modeling program.

2) NSREC 1965 – J. I. Vette

At the 1965 NSREC, James Vette of the Aerospace Corporation presented an invited paper [13]. The focus of his paper was on modeling the trapped radiation belts. He referenced 31 papers from various scientific meetings and journals (AGU, Journal of Geophysical Research, etc., see Reference 13) where measurements and modeling results had been presented. Vette reviewed the efforts to model the outer zone particles, stressing the difficulty of developing static maps of the dynamic electrons in that region. He compared the lifetimes of the outer particles (minutes) to those of the inner zone (years). As his paper demonstrates, the modulation of the electrons driven by the 27-day solar rotation period and the fluctuations connected with magnetic activity were known at that time. Rather than presenting maps of the dynamic outer zone, he gave “typical” integral spectra. It was known at the time that the protons in the outer zone are more stable, are more closely confined to the magnetic equator, and have a “soft spectrum”. With respect to damage to spacecraft, he notes that outer zone protons will affect only unshielded devices, but that electron exposure during long missions will result in measurable effects.

Vette presented the AE-1\(^2\) map of the inner zone electron model and the newly developed AP maps for protons with energies greater than 4 MeV. Fig. 6 shows the AP-3 map that he presented for protons for energies greater than 50 MeV. He noted that the natural inner zone electron population was not well known before Starfish and that Starfish electrons dominated the population levels in regions below L of 1.8. He also stated that, before Starfish, protons up to several 100s of MeV dominated the inner zone. Interestingly, as an aside, Vette made reference to a then recent observation by McIlwain that a redistribution of protons at >34 MeV followed a large magnetic storm in the L = 2-3 region [14].

3) Trapped Particle Model Development, 1960s, 1970s, 1980s

Eight trapped proton models, eight trapped electron models, and one Starfish decay model were released during the 27 years that the trapped radiation modeling program was operative. The trapped particle models that are most often used at this time are the AP-8 [15] for protons and the AE-8 [16] for electrons. The AP-8 model, released in 1976, was the culmination of a long-term effort to include all of the previous models under one common approach and to include all of the data after 1970. After 1977, the modeling budget was significantly reduced so a similar effort to consolidate the electron models into the AE-8 model was not completed until 1983. The formal documentation of that model was released in 1991.

The AP-8 and AE-8 models include data from 43 satellites, 55 sets of data from principal investigator instruments, and 1,630 channel-months of data. By the 1970s, scientific interest had shifted from trapped particles to the plasma regime to determine the physical mechanisms of particle energization and transport. As a result, the number of new data sets available for trapped radiation environment modeling was drastically reduced. It was not until the measurement of storm belts by the CRRES mission in 1991 that concerns were renewed about the ability to model the trapped radiation belts to sufficient accuracy for using modern microelectronics in space.

4) The STARFISH Exo-Atmospheric, High-Altitude Nuclear Weapon Test

In 1958 the United States conducted the HARDTACK series of nuclear weapons tests over the Pacific Ocean and the ARGUS series in the South Atlantic Ocean. In 1962 additional tests were conducted\(^1\) The “A” is for Aerospace Corporation where Vette first worked.
in the FISHBOWL series. One test, the STARFISH PRIME, with a yield of 1.4 megatons TNT equivalent, was exploded on July 9, 1962 at a very high altitude (approximately 400 km) over Johnson Island in the Pacific Ocean (about 700 miles from Hawaii). This exo-atmospheric nuclear explosion released about $10^{25}$ fission electrons into the magnetosphere, creating an artificial radiation belt and raising the intensity of the inner zone electron population by several orders of magnitude. This additional radiation increased the radiation damage on spacecraft flying in that region to critical levels. The first failure (due to total ionizing dose) was the TELSTAR satellite, which was launched one day after the STARFISH. It was estimated that the spacecraft experienced a total dose from the explosion that was 100 times larger than was planned for the spacecraft lifetime. Within months after the tests, seven satellites failed, primarily due to solar cell damage.

Initial predictions of the longevity of the STARFISH debris ranged from the overly optimistic of some months to the more realistic of a few years. Studies conducted in the late 1960s [17,18,19,20] attempted to define the rate of decay with varying results. An in-depth evaluation was performed in 1970-1971 [21] using data from the 1963-38C satellite that covered the time span from September 1963 to December 1968. The researchers identified three distinct regions within the inner zone domain that were populated by the artificial electrons and established that their decay lifetime $\tau$ in days could be best expressed by a function of three variables:

$$
\tau_1 = \tau(B, L, E) \quad \text{in Region 1}
$$

$$
\tau_2 = \tau(L, E) \quad \text{in Region 2}
$$

$$
\tau_3 = \tau(E) \quad \text{in Region 3}
$$

where $B$ is the field strength in gauss, $L$ is the magnetic shell parameter, and $E$ is the energy in MeV. Fig. 7 shows the dependence of the decay lifetime on $L$ for $E = 0.28$ MeV. In 1972 a more thorough approach produced a model of the STARFISH flux for September 1964 based on data from several spacecraft (OGO-1, OGO-3, OGO-5, OV3-3, and 1963-38C) [12]. That model distinguished between artificial and natural electrons and provided the artificial flux as a function of equatorial pitch angle, energy, and $L$ value. The decay times for this flux were determined by two separate methods, which were combined to yield average values that are appropriate for the evaluation of the long-term loss process of the artificials. A map of the threshold energy as a function of $L$ for decay cutoff times is shown in Fig. 8.

Numerical values relating to nuclear explosions must include a substantial margin of error. In addition to the difficulties in making measurements of these events at the time of their occurrence, the results are dependent on circumstances that cannot be predicted. Two nuclear weapons of different design may have the same explosive energy yield, but the effects could be markedly different. In the case of the STARFISH, estimates of errors associated with the cutoff-time model are given in Table 3.

**TABLE 3: ACCURACY OF THE STARFISH DECAY MODEL IN MONTHS**

It is interesting to compare the STARFISH effects with those of a Soviet high-altitude test of a low-yield weapon that was performed on October 28, 1962 over Semipalatinsk in Kazakhstan. Fig. 9 shows the integral Van Allen belt electron fluxes before and after the event in the regions form $L$ of 1.8-4.0 for particles with energies of $E > 0.5$ and $E > 1.9$ MeV. This region covers the outer edge of the inner zone, the slot region, and the beginning of the outer zone. The bulk of the fission electrons are concentrated between $L$ of 1.8-2.7, which is to be expected because their injection occurred at a high latitude location. In contrast, the STARFISH debris was concentrated in the inner zone due to the low latitude of the Johnson test site. A schematic of the distribution of the fission electrons from these two tests is shown in Fig. 10 in terms of magnetic shell parameter $L$ and magnetic latitude. Fig. 11 is an attempt to show the average lifetimes of the $E > 2$ MeV electrons from the STARFISH and Soviet experiments. Although it is difficult to draw conclusions from only two tests, the data suggest that longevity is maximum at low $L$ values, decreases rapidly towards the slot region, and is in the range of months and perhaps weeks at larger $L$ values.

5) The Storm Belts - The CRRES Mission

A discussion of developments in our knowledge of the trapped radiation environment cannot be complete without including the contributions that were made by the CRRES mission. Not only did the mission provide data for radiation belt modeling, but it also rekindled interest in radiation belt science. The CRRES satellite carried a full complement of particle detectors, and its geosynchronous transfer orbit of 350 km perigee and over 30,000 km apogee at 18$^\circ$ inclination was ideal for measuring radiation belt particles. Unfortunately, the CRRES launch was too late to measure the large September and October 1989 solar events. However, CRRES was in operation for the second peak of solar cycle 22 and observed the largest magnetic event ever recorded in March of 1991. Analysis of data from the instruments on CRRES before and after the March 1991 magnetic storm showed extremely large redistributions of the trapped particle populations. At the 1991 NSREC, Mullen et al. [22] reported that the CRRES proton instrument had measured a new proton storm belt in the
slot region \((L = 2-3)\). In Fig. 12 this belt for 47 MeV protons is seen as a region outside of the inner zone belt of protons. In the April 1996 Special Issue of the TNS, Gussenhoven et al. [23] summarized all of the CRRES results including a review of this proton belt formation and the production of an electron storm belt during the same storm. Recall from Section III.A that in 1965 McIlwain previously reported that he had observed a redistribution of protons at \(E > 34\) MeV following a large magnetic storm in the \(L = 2-3\) region.

One of the most important conclusions that resulted from the CRRES program was that existing theory could not explain the particle penetrations deep into the magnetosphere observed by McIlwain and Mullen et al. In 1993 Li et al. [24] used a simplified model of the Storm Sudden Commencement (SSC) compression of the magnetosphere to show that electron belts like those measured by CRRES can be created in tens of seconds when the interplanetary shock wave from the storm interacts with the magnetosphere. Later Hudson et al. [25] showed that this shock acceleration theory could also be used to explain the sudden formation of proton storm belts.

The frequency of occurrence of these atypical events is unknown, therefore, applying current models to setting design and operational rad-hard requirements creates uncertainties that are impossible to quantify. Dynamic environment simulations are a method that could be used to address this issue. In the April 1996 Special Issue of the TNS, Boucher et al. [26] reviewed modeling techniques that provide computer generated models of trapped particle transport during storms. Boucher and Bourdarie [27] are developing a 4-D diffusion code to calculate the transport of particles throughout the inner magnetosphere. They have applied the code to several problems, including calculation of the transport of existing and injected particles during storms, determination of the ring current growth using calculations for both protons and electrons, and discovery of the process by which high-energy particles are totally removed from the internal magnetosphere. Fig. 13 gives an example of the output from the code where the diffusion of 100 keV electrons in the radiation belts is plotted after injection of particles at midnight while the field increases (simulating a storm). The differential fluxes are in units of electrons MeV\(^{-1}\) cm\(^{-2}\) s\(^{-1}\). Case studies have been validated using CRRES and STRV-1b measurements [28]. In the future they plan to drive the model with interplanetary magnetic field strength thereby greatly increasing the database for running the code and providing statistics for a range of conditions for a given time period.

6) European Space Agency (ESA) Model Improvements

A paper in the April 1996 Special Issue of the TNS highlighted some of the work by the European Space Agency (ESA) in improving the trapped radiation models. Daly et al. [29] identified errors in the NASA models and documentation, including a source code error in AP-8-MIN and the fact that the AZUR dataset on which the AP-8-MAX is based covered a time span of 3 months, not 6 months. They also noted that the Space Shuttle and the LDEF satellite measured environments in low altitudes (300 to 500 km) were from 60 to 100% higher than those predicted by the AP-8 models. The authors determined that a large source of this error is due to the method used to interpolate between the \(B/B_0\) values in the regions near the atmospheric cutoff. They developed an alternate interpolation method that increased the estimated fluxes by 10 to 40% bringing them closer to the measured levels. They also recommend that an additional \(L\) increment at the low \(L\) values be included in the models to give better resolution at steep gradients. Fig. 14 shows the steep gradients below \(L\) of 1.4. The additional \(L\) increment increased the estimated fluxes by about 40%. When these two refinements to the interpolation scheme were combined, the revised flux levels were much closer to the measured values. Daly’s work renewed interest in improving the radiation belt models for applications to enabling and commercial off the shelf technologies, which can have very low radiation tolerance.

7) Radiation Belt Models with Improved Time Resolution

Analysis of the CRRES instrument and experiment data showed that not only is the environment extremely dynamic but also that electronic parts respond to the short-term changes. The AP-8 and AE-8 radiation belt models, with their 4-6 year averages, were adequate for long mission durations and for long-term degradation effects. However, modern spacecraft and instrument systems have serious problems with short-term effects, such as, interference and data corruption. With a time resolution of 4-6 years, the radiation belt models place a serious restriction on the ability to address design and operation issues of systems.

Several researchers have recognized the need for trapped particle models with finer time resolution. Three empirical models were developed using CRRES data to estimate short-term dynamic changes in the particle population, the CRRESPRO [30], CRRESELE [31], and CRRESRAD [32]. The CRRESPRO gives estimates of the trapped proton levels before and after the March 1991, simulating quiet and active conditions in the magnetosphere. CRRESELE gives estimates for eight conditions of magnetic activity, six ranges of activity as determined by the \(A_{p15}\) magnetic index and for the average and worst case conditions that CRRES measured. The CRRESRAD model was developed from data measured by the CRRES 4-domed dosimeter and includes quiet and active conditions. While the models
are based on data collected over a short 14-month period and during solar maximum conditions only, they give the most comprehensive picture available of the environment resulting from a geomagnetic storm. Reference 23 gives an overview of the models, including a comparison to the AP-8 and AE-8 models and a description of their limitations. Later data from a similar dosimeter flown on the APEX satellite was used to extend the CRRESRAD model to low altitudes and high latitudes. The APEXRAD [33] dose model includes the contributions from outer zone electrons and their variability, which was investigated as a function of magnetic activity using the $A_{\text{ps}}$ index.

Models with improved time resolution for the low altitude portion of the Earth's radiation belts have been developed in recent years. Pfitzer [34] plotted predicted flux values for low inclination orbits as a function of the average atmospheric density. From this index, he interpolated and extrapolated fluxes from the AP-8 and AE-8 models for solar activity conditions. Later Huston and Pfitzer [35] analyzed proton instrument data from the CRRES and TIROS/NOAA satellites with the goal of developing a low altitude (<850 km) trapped proton model with variation over an entire solar cycle as a function of solar activity indices (Fig. 15). At NSREC 1998, Huston presented the first trapped proton model with true solar cycle variation [36]. The proton flux levels were determined by using the solar radio flux proxy for atmospheric heating and included the phase lag between rise and fall of flux levels and solar activity.

Recently, the Huston team joined efforts with Xapsos and others at GSPC to add statistical variations to the solar cycle driven model thereby adding confidence level information to proton levels. This increases the ability to address trapped proton variations for spacecraft design and mission planning. This work was presented at the 2002 NSREC [37]. The authors note that this model is still in development and requires further analysis and validation.

Heynderickx and Lamaire [38] also plotted model fluxes as a function of the average weighted density of the atmosphere (ns) that is encountered by a particle on its drift shell. They found that the relationship between ns and the AP-8 and AE-8 models fluxes is well ordered, especially at low L values. If a practical form for the ns calculation is developed, this method could lead to replacing $B_0$ with $B_0$(ns) when accessing the AP-8 and AE-8 models, thereby reflecting solar activity effects. Using data from the proton instrument on the SAMPEX spacecraft, they also developed a trapped proton model with improved time resolution [39].

B. Trapped Radiation at Other Planets

1) NSREC 1971 – J. W. Haffner

After numerous missions to explore the Earth's radiation belts and interplanetary space, scientists became interested in exploring the outer planets. For the 1971 IEEE TNS, J. W. Haffner contributed a paper, "Natural Radiation Environments for the Grand Tour Missions" [40] that presented an overview of the solar "flare", galactic, and magnetically trapped particle radiation expected during missions to the four outer planets, Jupiter, Saturn, Uranus, and Pluto. He estimated the environment levels from data and models and predicted mission doses. The ability to predict the environments at the outer planets was greatly hampered by the lack of measurements of the interplanetary environment beyond 1 astronomical unit (AU) and by the lack of any measurements of the trapped radiation environments of Jupiter and Saturn.

Haffner gave estimates of the trapped radiation belts of Jupiter and Saturn. It was theorized that Jupiter and Saturn had belts similar to those of the Earth based on decametric (bursts) and decimetric (quasi-steady state) radio frequency radiation emitted by Jupiter and Saturn. Similar rf radiation emission was measured from the Earth's belts. The decimetric radiation is due to synchrotron emission of the electrons trapped in the magnetic field, the decametric radiation is associated with one of Jupiter's moons, lo. Based on assumptions about the limiting particle fluxes, similar relationships in Jupiter's and Earth's magnetic fields, and particle and plasma densities, and by ignoring the effect of the planet itself, Haffner derived relationships between the magnetic field at Jupiter's equator, the particle density relative to the plasma stability limit, and the effective inner radius of the Jovian belts. From that set of parameters, the electron dose rates for mission flybys were estimated.

The same methodology could not be applied to trapped protons because they do not radiate as the electrons do. Haffner pointed out that theory explaining the source and loss mechanisms for protons or electrons in the Earth's belts that could be applied to estimating the proton belts of Jupiter did not exist. Therefore, he had to base the proton estimates for Jupiter and Saturn on the ratios of the protons/electrons in the Earth's Van Allen belts. He used these estimates to calculate doses for the missions.

2) NSREC 1972 – Paul Kase

At the 1972 NSREC Kase followed up Haffner's work with a presentation that focused on concerns about displacement damage on spacecraft electronics due to proton and neutron environments [41]. In addition to natural sources, Kase presented the problem posed by having neutron emitting radio-isotope thermoelectric generators (RTGs) on board spacecraft. Kase also
revisited the problem of modeling the proton belts of Jupiter. By that time, Divine had developed nominal and upper limit models of the proton belts [42]. The three orders of magnitude difference between the two models was an indication of the inaccuracy inherent in the unvalidated theoretical approach. It was expected that the proton spectra were "very hard" near the surface of Jupiter and softer at great distances.

3) Measurements of the Grand Tour Mission

The Pioneer missions to the outer planets carried instruments to measure the radiation environment. The measurements showed that the radiation was orders of magnitude higher than expected. Peak intensities of electrons in the belts, as measured by Pioneer 10, were 10,000 times greater than Earth's maximum. Also, the electron energies were found to be greater than 20 MeV. Protons were several thousand times as intense as Earth's belts. The inner radiation belts of Jupiter, as measured by Pioneer 10, had the highest radiation intensity so far measured, comparable to radiation intensities following an explosion of a nuclear device in the upper atmosphere. Pioneer 11 confirmed these high intensities. In the inner region of the magnetosphere, protons exceeding 35 MeV appear to peak in two shells; the outer shell was detected at 3.5 Jovian radii by Pioneer 10 and confirmed by Pioneer 11, and an inner shell, discovered by Pioneer 11, has a peak at 1.78 radii of Jupiter. Pioneer 11 also found that there is a greater flux of energetic particles at high Jovian latitudes than would have been expected from the measurements made by Pioneer 10. It also discovered that the flux of energetic particles peaks on either side of the dipole magnetic equator [43]. This discovery led to the need to retrofit the Galileo spacecraft with radiation hardened bipolar processors, because the spacecraft design with unhardened processors had been fixed before the arrival of the Pioneer spacecraft at Jupiter [44]. The models of the Jupiter radiation environment were updated using data from the missions to the outer planets [45].

IV. SOLAR PROTON EVENT MODEL DEVELOPMENT

A. Early Estimates of Solar Proton Levels

At NSREC 1965 Vette compared the levels of solar "flare" protons and galactic cosmic rays based on work by Malitson and Webber [46] and noted that the solar protons are low during periods of low solar activity. He also pointed out that they were considered a serious hazard for astronauts on Apollo missions. To predict the solar proton environment, Haffner compiled all of the existing solar proton data at 1 AU for 1956-1968, including data from Webber, McDonald, Lewis, Modisette, and Mosley (see Reference 46). He developed fitting functions to fill in missing data that described the onset, rise time, and decay of solar events as a function of time. He then reconstructed probable values based on sunspot numbers by using predictions of future sunspot numbers developed by Weddell and Haffner [47]. To estimate the diffusion of the solar protons throughout interplanetary space, Haffner assumed a spatial dependence of 1/r^2 where r is measured in AU, which is an approximation still used today. The method was used to predict the expected solar particle environment for each year from 1970 to 1989. Haffner points out the uncertainties in predicting solar particle levels. In fact, as seen in Table 4, the differences between the solar proton levels predicted for 1972 and 1989 at 1 AU and those measured by spacecraft are large.

<table>
<thead>
<tr>
<th>Year</th>
<th>Predicted p/cm²/yr</th>
<th>Actual p/cm²/yr</th>
</tr>
</thead>
<tbody>
<tr>
<td>1972</td>
<td>7.0x10^8</td>
<td>2.4x10^9</td>
</tr>
<tr>
<td>1989</td>
<td>8.0x10^8</td>
<td>4.5x10^9</td>
</tr>
</tbody>
</table>

B. The First Statistical Models

As concern grew over electronics and human exposure to solar protons and as the amount of available on-orbit particle measurements increased, the use of compilations of data for estimating solar proton levels was replaced by modeling efforts. In 1974 King [48] published the first statistical model for solar proton events using Poisson distributions. He concluded from his analysis of proton data from the 20th solar cycle that solar proton events could be classified into "ordinary" and "anomalously large". This was based on the fact that only one anomalously large event occurred in the 20th solar cycle - the August 1972 event. That event alone accounted for 84% of the total proton fluence in the solar cycle at energies E > 30 MeV. Stassinopoulos [49] developed the solar proton model, SOLPRO, based on King's statistical analysis. However, when Feynman et al. [50] added cycle 19 and 21 data to the solar proton event database, they were able to conclude that individual solar proton events actually form a continuum of event severity from the smallest to the largest, blurring the distinction between ordinary and anomalously large events.

C. Engineering Oriented Statistical Models

Many large events similar to the August 1972 event occurred in cycle 22 increasing concern about the validity of the solar proton models. With the goal of improving the ability to address practical aspects of spacecraft reliability, a team led by Xapsos began compiling solar proton data for solar cycles 20, 21 and 22 and using statistical techniques to derive probability distributions of cumulative solar proton fluences. A review of the data sets can be found in Reference 51. At
the 1996 NSREC, Xapsos et al. [52] presented a paper that described the application of extreme value theory to determine probability of encountering a single large event over the course of a mission. They also used compound Poisson process theory to describe the probability of encountering various fluence levels during a mission. The work of the Xapsos team confirmed the Feynman conclusion that a “typical event” cannot be defined.

The Xapsos team then turned their focus to understanding how to define the peaks of solar proton events. To accomplish this, they applied Maximum Entropy Principle (MEP) to select the least biased event probability distribution. The MEP, used for earthquake predictions, is valuable for analyzing incomplete datasets. They validated the results with Lunar Rock Records dating back to ancient times. New peak values for E > 10 MeV were published in the 1998 TNS [53]. The Xapsos team continued their work by establishing worst case solar proton spectra for solar events [54]. When comparing their model with the CREME96 [55] “worst-week” solar proton model, which was based on the October 1989 solar particle event, they found that, statistically, the CREME96 model is closer to a 90% worst case event model [56]. Xapsos et al. have combined the model elements into the Emission of Solar Protons (ESP) model, which is available as a computer code [51,56]. Fig. 16 compares energy spectra for the new ESP model with the King and Feynman models.

V. HEAVY IONS

A. Heavy Ions Prior to NSREC 1975

In the early 1900s, scientists were puzzled by “charge” on electroscope leaves while performing laboratory experiments. The leaves of an electroscope repelled each other without the presence of charge. Scientists assumed that this was the result of the ionization of air by the natural radiation present on Earth (see Table 3). Efforts were made to eliminate the radiation by using radiation pure materials, however, the problem persisted. In 1913 an Austrian scientist, Victor Hess, devised an experiment to put an electroscope in a balloon to get it away from the Earth’s radiation. As Hess and his experiment ascended in the balloon, he observed that the radiation source did not go away rather, as the altitude of the balloon increased, the radiation increased. Hess concluded that the source of this radiation was from outer space. In the summer of 1925, Millikan confirmed with his lake experiments that the radiation source was indeed from outer space. When he presented his lake experiment findings, he called the radiation “cosmic rays” [57]. In 1936 Hess and C. D. Anderson received the Nobel Prize for their discovery of cosmic rays.

Intense interest in understanding cosmic rays continued into the 1950s. Recall from Section IIIA that Van Allen’s interest in getting instruments into space was to study cosmic rays and their origin. The first discussion of GCRs at NSREC was in 1965 when Vette [13] compared the levels of solar “flare” protons and galactic cosmic rays based on work by Malitson and Webber [46]. The primary concern was dose levels on spacecraft components, but it was determined that, in the presence of protons and electrons, the dose from heavy ions is a non-issue.

In the late 1960s, the interest in cosmic rays of solar and galactic origin went beyond basic scientific research and became a safety issue when astronauts on Apollo missions reported visual light flashes. McNulty [58] proposed that Cerenkov radiation generated by individual cosmic ray ions traversing the vitreous of the eye were responsible for the flashes. He and his colleagues proceeded with a series of experiments that exposed human subjects, P. McNulty, V. P. Pease, V. P. Bond, and L. Pinsky, to energetic heavy ions at accelerators. Understanding of the source and mechanism of the light flashes raised the concern for astronaut and safety, which in turn generated interest in measuring and modeling the galactic and heavy ion space environment. In fact, the IMP-8 spacecraft, which has provided the best long-term data set of heavy ions, was planned as a result of these concerns.

Haffner discussed the concern for the GCR contribution to total dose on spacecraft in his 1971 NSREC paper [40]. He presented values for the levels of galactic cosmic rays[1] and estimated that the expected GCR dose for the Grand Tour missions was in the range of a few hundred rads.

B. NSREC 1975 – New Concerns about Heavy Ions

Prior to NSREC 1975, heavy ion populations, whether of galactic or solar origin, were not considered a major concern for the reliability of spacecraft electronics. Regardless of the region of space that missions visited, the contribution of heavy ions to spacecraft charging, ionizing dose, or displacement damage effects were insignificant compared to other sources of radiation, such as, the trapped radiation belts or protons from solar events. That changed at NSREC 1975 when Binder et al. [59] reported, “Anomalies in communication satellite operation have been caused by the unexpected triggering of digital circuits. Although the majority of these events have been attributed to charge buildup from high temperature plasmas, some of the events appear to be caused by another mechanism.” In addition to an analysis of the circuit effects and the basic mechanism of these events, the authors presented cosmic ray spectra of Meyer [60], and they calculated the intensities using abundances from various authors, including Burrell and Wright [61] who investigated

---

1 The source of the data is not referenced.
dose rates of GCRs for astronaut exposure. Previous research on galactic cosmic rays by magnetospheric physicists interested in basic scientific research and by nuclear physicists concerned with astronaut dose became significant for the NSREC community. As explained by McNulty [58], the interaction models used to explain the light flashes observed by astronauts eyes were modified slightly to explain the upsets observed in microelectronic circuits. The Binder paper sparked intense interest in modeling heavy ion environments and interactions. (see other papers in this volume)


A team at Naval Research Laboratory, led by Jim Adams, recognized the need for a comprehensive software package to calculate single event upset (SEU) rates in space that integrated environment predictions with particle interaction models. They embarked on the task of developing a comprehensive computer tool that could be used by researchers and engineers. Due to the extent of the upset problem, the effort had to include compilation of data sets for GCR and solar heavy ion populations, development of GCR and solar heavy ion models, evaluation of solar proton data and models, development of magnetospheric cutoff calculations, analysis of spacecraft shielding effects, and development interaction models for upsets due to heavy ions and protons. Adams et al. produced two Naval Research Laboratory (NRL) Memorandum Reports, which reported on topics related to this task [62, 63]. The first report, published in 1981, contained a comprehensive review of the near-Earth particle environment.

1) The CREME Heavy Ion Environment Models

The GCR environment model was based on data from several researchers collected through 1980. (see Adams [62]). Because of the dissimilar shape of their energy spectra, the hydrogen, helium, and iron ion distributions were treated as separate cases, and the other elements were scaled to one of the three spectra, as appropriate, using the relative abundances of the elements. Four different models of the GCR environment were developed, one of them being the well known “90% worst case environment”.

The solar heavy ion environment was more difficult to model because of the unavailability of a good dataset from spacecraft instrumentation. Adams et al. assumed that the solar particle events with the highest proton fluxes are always heavy ion rich and estimated fluence levels for the higher energy solar heavy ions (>1 MeV) by scaling the abundances to protons. Eight different models of the solar heavy ion environment were developed because, without the benefit of comprehensive space measurements, the authors had to account for all possible solar activity conditions.

2) Extending the Models to a Rate Prediction Tool – CREME86

At NSREC 1982 Adams et al. presented a paper [64] that, with a paper by Petersen et al. [65], essentially laid the groundwork for the CREME86 code. This code was the first end-to-end desktop capability to calculate radiation environments throughout near-Earth regions and to use laboratory test parameters from devices to calculate the rate of a radiation effect. The simple title, “The Natural Radiation Environment Inside Spacecraft,” does not reflect the complexity and groundbreaking nature of the work. The paper reviewed the work on the development of the environment models [62] and presented methods to calculate the transport of particles through the magnetosphere and through spacecraft shielding. The authors also showed the utility of using linear energy transfer (LET) spectra to represent the heavy ion environment in a form that condenses the energy spectra of all ions into a compact expression and that can be applied to calculating energy transfer in microvolumes. Heinrich constructed the first LET spectrum describing the ion environment in space in 1977 for biomedical purposes [66, see also 67]. At 1978 NSREC, Pickel and Blanford showed the applicability of the LET spectrum to the single event upset problem in microelectronics [68]. The resulting CREME [69] code was first released in 1986.

3) Updates to the CREME Code

For ten years the CREME86 code was a standard for calculating heavy ion environments. As knowledge of the radiation environment increased, it became apparent that the CREME86 models could be improved. Using ISEE 3 data, Reames et al. [70] found an inverse correlation between proton intensity and the iron/carbon heavy ion abundance ratio and that the composition of a solar particle event was a result of the location of the event on the sun. This contradicted Adams’ assumption that all solar events are He rich, meaning that the solar heavy ion models were probably over predicting. Dyer et al. [54] measured LET spectra during the March 1991 event with the CREDO instrument on UoSAT-3. When they compared the measurements with LET calculated using the CREME86 solar particle models, they found that, in the LET range important for single event effects analyses, all of the models severely overpredicted the LET levels. Dyer et al. [71] measured LET spectra during several events, including the March 1991 event, with the CREDO instrument on UoSAT-3. When they compared the measurements with the LET spectra calculated using the CREME86 solar particle models, they found that, in the LET range important for single events from ions (greater than 0.5 MeVmg^-2cm^2), all of the models severely overpredicted the fluxes.

In the 1990s NRL recognized the need to improve the environment models in the CREME86 code. The most important update to the code was the solar heavy
ion model. Dietrich from the University of Chicago analyzed the solar heavy ion data from the IMP-8 satellite, providing the most comprehensive set of solar heavy ion space data to date [72]. The dataset is especially important for modeling the fluences at higher energies. A team led by Tylka used the results to model the solar heavy ions based on the October 1989 solar particle event. An analysis of 100 solar heavy ion events in the Dietrich database showed that this event could be used as a representative of a “worst case” environment. Tylka et al. presented the CREME96 model the 1997 NSREC [55]. The CREME96 solar heavy ion estimates are significantly lower than the heavy ion models in CREME86. Fig. 17 compares the LET energy spectra for the CREME86 and CREME96 solar heavy ion models. Recently Dyer et al. [73,74] have presented data on the LET spectra of solar particle events occurring between 1998 and 2001. While for many of the events there are very low enhancements of high LET fluxes, three events do approach or equal the CREME96 worst day model. At low LET, where protons dominate and usually lead to single event effects by nuclear interactions, two events slightly exceed the model. This is consistent with the suggestion by Xapsos et al. (see Section IV.C) that the CREME96 model is a 90% worst case. They are applying the methodology used in the ESP solar proton model to the solar heavy ion measurements to derive a statistically based solar heavy ion model.

The GCR environment model was also updated to include the analysis of the SAMPEX measurements of anomalous cosmic rays (ACRs). With the finding that the ACRs are not singly charged over energies of 20 MeVn⁻¹, four models in CREME86 were replaced with one GCR model in CREME96. The model change significantly reduced the calculated single event effects rates for high threshold devices in low altitude orbits [55].

VI. ATMOSPHERIC AND GROUND ENVIRONMENTS

A. Atmospheric Environments

In the last twenty years it has been discovered that electronics in aircraft systems, which are subjected to increasing levels of cosmic radiation and their secondaries as altitude increases, are also sensitive to single event effects. At the 1984 NSREC Tsao et al. [75] presented methods for calculating LET spectra and SEU rates from cosmic rays and their secondary fragments propagating down to 40,000 feet. Later the calculation of this component of the atmospheric radiation environment was embodied in the Air Propagation (AIRPROP) code [76]. In a companion paper to the Tsao et al. paper, Silberberg et al. [77] presented methods for calculating the SEU rates arising from the secondary neutrons generated by the interactions of cosmic rays with the atmosphere. These papers demonstrated that the ions are important at greater than about 65,000 feet but that SEUs from neutron interactions dominate at lower altitudes. They also predicted that SEU rates would increase during large solar particle events.

An increasing body of data on upsets in avionics systems were accumulated that pointed to neutrons as the being the primary cause of SEUs on aircraft. In an unintentional experiment in 1993, reported by Olsen et al. [78], a commercial computer was temporarily withdrawn from service when bit-errors were found to accumulate in 256 Kbit CMOS SRAMs (D43256 A6U-15LL). Following ground irradiations by neutrons, the observed upset rate of 4.8x10⁸ upsets per bit-day at conventional altitudes (35,000 feet) was found to be explicable in terms of SEUs induced by atmospheric neutrons. In an intentional investigation of single event upsets in avionics, Taber and Normand [79] flew a large quantity of CMOS SRAM devices at conventional altitudes on a Boeing E-3/AWACS aircraft and at high altitudes (65,000 feet) on a NASA ER-2 aircraft. Upset rates in the IMS1601 64Kx1 SRAMs varied between 1.2x10⁻⁷ per bit-day at 30,000 feet and 40° latitude to 5.4x10⁻⁷ at high altitudes and latitudes. Reasonable agreement was obtained with predictions based on neutron fluxes.

Our knowledge of neutron levels comes from balloon, aircraft, and ground based measurements. These studies show that the energies of the neutron flux range from keV to hundreds of MeV. For SEU calculations in aircraft systems, it is usually only the energies greater than 10 MeV that are significant. It is also known that the flux peaks at an altitude of about 60,000 feet, which is the same altitude of the peak of observed SEU rates. Because the shape of the neutron spectrum varies little over altitude, models can be greatly simplified. However it has recently been demonstrated by Dyer and Lei [80] that aircraft structures, passengers and fuel can produce significant thermal neutron fluxes, leading to potentially high rates of SEU in components containing the nuclide boron-10 in borophosphosilicate glass passivation layers.

Two coordinate systems are commonly used to define the neutron distributions, energy-altitude-latitude and energy-atmospheric depth-magnetic rigidity. Taber and Normand [6] have developed an empirical model in the energy-altitude-latitude system based on studies by Mendall and Korff [81], Armstrong [82], and Merker et al. [83]. A model by Wilson-Nealy [80] that aircraft structures, passengers and fuel can produce significant thermal neutron fluxes, leading to potentially high rates of SEU in components containing the nuclide boron-10 in borophosphosilicate glass passivation layers.

Since the discovery of SEUs on at aircraft altitudes, researchers have made significant efforts to monitor the environment. Dyer et al. flew a version of their Cosmic
Ray Environment and Activation Monitor (CREAM) on regular flights on board Concorde G-BOAB between November 1988 and December 1992. At NSREC 1989 Dyer et al. first reported on the results of measurements aboard the Concorde aircraft [84]. Results from 512 flights have been analyzed of which 412 followed high latitude transatlantic routes between London and either New York or Washington DC [71]. Thus some 1,000 hours of observations have been made at altitudes in excess of 50,000 feet and at low cut-off rigidity (< 2 GV**), and these span a significant portion of solar cycle 22. Fig. 18 shows the count rate in CREAM channel 1 (19fC to 46fC, LET 6.1 MeV cm² g⁻¹) plotted as monthly averages for the ranges 54,000-55,000 feet and 1-2 GV. The rates show a clear anticorrelation with the solar cycle and track well with the neutron monitor at Climax Colorado (altitude 3.4 km, cut-off rigidity 2.96 GV). The enhanced period during September and October 1989 comprised a number of energetic solar particle events observed by ground level, high latitude neutron monitors and the Concorde observations are summarized in Table 5 [84,85]. The enhancement factors are the ratios between flight- averaged count rates at greater than 50,000 feet during the solar particle events and the rates during immediately preceding flights on identical routes when only quiet-time cosmic rays were present. The charge-deposition thresholds of each of the channels are also given. Recently Dyer and Lei [80] have attempted to calculate the neutron fluxes in the atmosphere during large solar particle events using Monte-Carlo radiation transport in conjunction with data from ground-level neutron monitors and space borne detectors. Some success has been obtained in fitting the results from Concorde and it has been demonstrated that such events can produce high rates of SEU, up to several hundred times the cosmic ray rates, for high latitude routes even at subsonic altitudes of less than 40,000 feet.

** Magnetic rigidity of a charged particle is defined as the momentum per unit charge and is measured in units of electron volts/number of charge units, i.e., volts or gigavolts (GV).

TABLE 5: ENHANCEMENT FACTORS FOR CREAM ON CONCORDE DURING SOLAR PARTICLE EVENTS

The CREAM detector has also been operated on a Scandinavian Airlines Boeing 767 operating between Copenhagen and Seattle via Greenland, a route for which the cut-off rigidity is predominately less than 2 GV. Approximately 540 hours of data accumulated between May and August 1993 have been analyzed and these are combined with Concorde data from late 1992 to give updated altitude profiles [3]. Using the AIRPROP code [86], they have also shown that cosmic rays and their secondary fragments are not the major contribution to SEUs at aircraft altitudes. Recent work [87] has concentrated on explaining both the altitude dependence and the energy deposition spectra using a microdosimetry code extension to the Integrated Radiation Transport Suite. Fig. 19 shows that atmospheric secondary neutrons are the major contribution but that ions start to become important at the highest altitudes. Fig. 20 shows that at 30,000 feet the charge deposition spectrum is dominated by neutron interactions at the high end while energetic secondary electrons and muons contribute to the low channels. The work of Normand et al. [88] arrived at similar conclusions on the neutron contribution by scaling results of irradiation of silicon detectors obtained at a spallation neutron source.

B. Ground Environments

The first evidence of sea-level soft fails (single event upsets) on 16Kb DRAMs was given by May and Woods [89]. The source of the radiation was traced to alpha particle emission from contaminates in ceramic packing produced in a factory on the Green River in Colorado which was downstream from an old uranium mine. In 1978 Zeigler realized that, if alpha particles from materials could induce soft fails, it is possible that ground cosmic rays could do the same thing. In 1979 Ziegler and Lanford detailed the mechanism by which sea level cosmic rays could cause soft fails [90]. O’Gorman et al. began field testing of soft fails on memories in the early 1980s. In those experiments, they found a distinct increase in soft fail rates with increasing altitude above sea level [91]. An IBM study on the repair rates of memory modules confirmed the altitude effect on soft failure rate, implying that cosmic radiation was responsible for some of the soft fails on devices in ground systems. The results of that study showed that the failure rates in the mountain states (mean altitude 3200 ft) was 5 times higher than in the United States as a whole and that the rates in Denver at 5,280 ft. were 10 times higher [92].

There are many products from the cosmic radiation showers (see Section II.B and Fig. 4) so it was important to determine which were dominant in causing the soft fails. At the 1983 NSREC, Dicello concluded that all types of particle radiation could cause soft fails at the ground level with some probability [93]. At NSREC 1996, Normand correlated measured ground level bit error rate with Weapons Neutron Research neutron beam soft fail rate measurements to show that neutrons are the dominant cause of soft fails on the ground [94]. At the 1997 NSREC, Normand et al. showed that neutrons could also cause destructive failures in microelectronics on the ground [95].

Ziegler and colleagues began an effort to define the cosmic ray terrestrial environment based on the work of many researchers who began studying cosmic radiation
after Hess’ balloon experiments [5]. The focus of the work was on defining neutron spectra as a function of altitude as shown in Fig. 21. In the figure the altitude dependence is expressed as atmospheric depth where 0 g/cm² represents the top of the atmosphere. Ziegler concludes that the energy spectra of neutrons at various locations through the world are known within a factor of 2 below energies of 100 MeV. Two problem areas that he notes are that 1) the accuracy of the levels of neutrons above 100 MeV is no better than a factor of 3 and 2) the flux of neutrons in the energy range of 20-100 MeV is very dependent on local surroundings reducing the accuracy to a factor of 3.

VII. SUMMARY

The state of the knowledge of the space, atmospheric, and ground level radiation environments was reviewed, tying it with the last 40 years of the NSREC and its proceedings. We saw that basic science research often is used to derive the definitions required for understanding radiation effects and for developing models that are useful for designing radiation hardened systems. Often our ability to increase system capability is limited by appropriate models of the radiation environment. Large uncertainty factors in environment definition translate to large design margins. The direct result is reduced system resources due to increased shielding, higher mitigation overhead, and/or the use of less capable components.

The availability of useful models of the radiation environment is dependent on our knowledge of the environment, the availability of appropriate data for modeling, and resources for modeling and validation. The use of radiation sensitive modern microelectronics in systems requires models that can meet the needs of system designers and operators through all phases of the life cycle. Due increasing sensitivity of microelectronics to radiation and increasing complexity of spacecraft systems, it is more difficult to completely avoid the risk of radiation effects on systems. The approach is to reduce risk during mission planning and spacecraft and instrument design and to manage the residual risk during launch and operations.

Fig. 22 shows three mission phases that require environment models to address risk issues in systems. Statistically based specification or “climate” models are used for the design phase, which includes system design and mission planning. The minimum requirements of the models are that they should represent long term variation over the solar cycle with at least 1-month resolution, provide worst case estimates, provide confidence levels, and represent the environment in a broad energy range so they are applicable for surface materials to deeply embedded sensors. The operation phase requires “forecasting” models for protecting the systems investment, for mission planning, and for personnel scheduling. The minimum requirements of these models are spatial coverage from interplanetary to low Earth orbits, information about the level of severity of storms, forecasts of quiet times for maneuvers and on-board operations, and they must be specific to effects on technologies. There is increasing awareness of the need to develop forecasting (space weather) models as part of the space infrastructure [9]. The third type of model for “nowcasting” the environment is used to resolve anomalies so risk can be reassessed for both the operating systems and for other systems that are in development. The minimum requirements of these models are spatial coverage from interplanetary to low earth orbit, time profile of the radiation event, adequate energy range and ion composition/linear energy transfer (LET) range. They should also be specific to the effect on technology. If anomaly resolution is critical to the mission and must be performed in near real-time, monitors that are in close proximity to the system may be required.

There are still radiation environments that need improved models. Funding by NASA’s Space Environments and Effects Program and the Living with a Star Program and ESA/ESTEC has resulted in improvements in “climate” models of the environment in some areas. However, model consolidation, validation, and standardization has not been supported.

The Living with a Star Program is making heavy investments in developing the scientific understanding of the how the sun-earth system is connected. NASA is also working with organizations that need to nowcast and forecast environments for operational systems to transition that knowledge to operational models. It is hoped that this effort will be extended to include the “climate” models that are required for system design.

VIII. ACKNOWLEDGEMENT

The authors wish to thank Dr. Allan Johnston of the Jet Propulsion Laboratory, the General Chairman of NSREC 2003, for urging the NSREC community to revisit some of the outstanding research that has been published over the past forty years. We also acknowledge the support of Chikia Barnes of SGT Inc. and Martha O’Bryan of Raytheon in producing graphs from archival materials.

VIII. REFERENCES


Fig. 1: Artist's drawing of the Earth's Van Allen radiation belts. The drawing does not show the SAA, from BIRA.

Fig. 2: Radiation belt electrons (E > 0.4 MeV) measured by the SAMPEX spacecraft. The measurements are plotted as a function of L and the day of year. The outer zone (L>3), slot region (L=2-3), and the inner zone (L<3) can be clearly seen. The plot shows the dynamic electron population in the outer zone and the numerous episodes of the slot region being filled with storm electrons. S. Kanekal, NASA/GSFC.

Fig. 3: The IMP-8 measurements of carbon, oxygen, and nitrogen ions in interplanetary space (24-hour averaged mean exposure flux). Also plotted is the average sunspot counts for the same time period. Note the sudden rises in the counts due to solar particle events and the anticorrelation of the galactic cosmic ray levels with the sunspot counts.

Fig. 4: Cosmic rays hit the top of the atmosphere and disintegrate into neutrons, pions, muons, etc. Smart and Shea.

Fig. 5: Components in the ground level cosmic ray environment at New York City as a function of energy, Zeigler/IBM.

Fig. 6: The AP-3 proton model map presented by Vette at NSREC 1965.

Fig. 7: Dependence by region of the decay lifetime for electrons in the STARFISH model.

Fig. 8: Decay cutoff times in months for the STARFISH model as a function of threshold energy and L.

Fig. 9: Integral Van Allen belt electrons before and after the Soviet event.

Fig. 10: Schematic of the distribution of fission electrons from the STARFISH and Soviet tests in magnetic coordinates.

Fig. 11: Comparison of the average lifetimes of > 2 MeV electrons for the STARFISH and Soviet tests.

Fig. 12: The second proton belt produced by the March 1991 magnetic storm measured by CRRES for proton energies of 47 MeV, AFR/LVS.

Fig. 13: Results from the Salammbio 4-D diffusion code showing the distribution of 100 keV electrons after an injection (storm) event. The omnidirectional differential fluxes are in units of MeV^-cm^2-s^-1.

Boundaries/ONERA.

Fig. 14: Steep gradients in the AP-8 model at low L values for integral proton fluxes at E > 30 MeV. The fluxes are omnidirectional and integral for E > 0.5 MeV. Daily/ESA-ESTEC.

Fig. 15: Proton distributions on the magnetic equator for protons in the inner zone at E > 80 MeV which were used as the basis for the first trapped proton model with true solar cycle dependence. Hustin and Pfitter.

Fig. 16: Comparison of probability models for solar particles.

Fig. 17: Comparison of CREME86 and CREME96 solar heavy ions models for a geostationary orbit.

Fig. 18: Monthly mean count rates from the CREAM instrument on Concorde from Jan 1989 to Dec 1992 compared with ground level neutron monitor at Climax.

Fig. 19: Average CREAM channel 5 count rates as a function of altitude at 1-2 GV from SAS and Concorde flights. Also shown are the predictions from AIRPPOP and from neutron interactions as calculated using radiation transport and microdosimetry codes (LHIMD/AFRLIVS). Neutrons dominate at 3,000 to 40,000 feet but cosmic ray neutrons dominate at supersonic altitudes.

Fig. 20: The spectrum of charge deposition from CREAM at 30,000-31,000 feet compared with predictions. Neutrons dominate at high values while electrons and muons contribute at the low end.

Fig. 21: "Hess values" of measured energy spectra of neutrons at various altitudes up to outer space. The lowest curve at 1030 g/cm^2 is at sea level and the curve marked 700 g/cm^2 is at an altitude of 2 miles. Zeigler.

Fig. 22: Environment model types that are required in each phase of risk management for missions.
<table>
<thead>
<tr>
<th>Particle Type</th>
<th>Maximum Energy</th>
</tr>
</thead>
<tbody>
<tr>
<td>Trapped Electrons</td>
<td>10s of MeV</td>
</tr>
<tr>
<td>Trapped Protons &amp; Heavy Ions</td>
<td>100s of MeV</td>
</tr>
<tr>
<td>Solar Protons</td>
<td>GeV</td>
</tr>
<tr>
<td>Solar Heavy Ions</td>
<td>GeV</td>
</tr>
<tr>
<td>Galactic Cosmic Rays</td>
<td>TeV</td>
</tr>
<tr>
<td>Source</td>
<td>Type</td>
</tr>
<tr>
<td>-------------------------------</td>
<td>-----------</td>
</tr>
<tr>
<td>Terrestrial</td>
<td>Natural</td>
</tr>
<tr>
<td>Inhaled radionuclides(^1)</td>
<td>Natural</td>
</tr>
<tr>
<td>Internal radionuclides(^2)</td>
<td>Natural</td>
</tr>
<tr>
<td>Cosmic radiation</td>
<td>Natural</td>
</tr>
<tr>
<td>Cosmogenic radionuclides</td>
<td>Natural</td>
</tr>
<tr>
<td>Medical diagnostics</td>
<td>Manmade</td>
</tr>
<tr>
<td>Atmospheric weapons testing</td>
<td>Manmade</td>
</tr>
<tr>
<td>Airline travel</td>
<td>Natural</td>
</tr>
<tr>
<td>Consumer products</td>
<td>Manmade</td>
</tr>
<tr>
<td>Nuclear power</td>
<td>Manmade</td>
</tr>
</tbody>
</table>

\(^1\)Dust particles bearing radon daughters expose lung tissue to alpha radiation.
\(^2\)Naturally occurring radioactive elements (primarily potassium-40) other than radon daughters that find their way into our bodies.
## Table 3: Accuracy of the STARFISH Decay Model in Months

<table>
<thead>
<tr>
<th>Energy (MeV)</th>
<th>0.03</th>
<th>0.05</th>
<th>0.1</th>
<th>0.3</th>
<th>0.5</th>
<th>1.0</th>
<th>3.0</th>
</tr>
</thead>
<tbody>
<tr>
<td>1.3</td>
<td>±4</td>
<td>±4</td>
<td>±4</td>
<td>±5</td>
<td>±8</td>
<td>±10</td>
<td></td>
</tr>
<tr>
<td>1.4</td>
<td>±3</td>
<td>±3</td>
<td>±3</td>
<td>±4</td>
<td>±8</td>
<td>±10</td>
<td></td>
</tr>
<tr>
<td>1.5</td>
<td>±3</td>
<td>±3</td>
<td>±3</td>
<td>±4</td>
<td>±8</td>
<td>±10</td>
<td></td>
</tr>
<tr>
<td>1.6</td>
<td>±3</td>
<td>±3</td>
<td>±3</td>
<td>±4</td>
<td>±6</td>
<td>±7</td>
<td></td>
</tr>
<tr>
<td>1.7</td>
<td>±3</td>
<td>±3</td>
<td>±3</td>
<td>±4</td>
<td>±8</td>
<td>±8</td>
<td></td>
</tr>
<tr>
<td>1.8</td>
<td>±3</td>
<td>±3</td>
<td>±3</td>
<td>±6</td>
<td>±8</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1.9</td>
<td>±8</td>
<td>±7</td>
<td>±5</td>
<td>±7</td>
<td>±8</td>
<td></td>
<td></td>
</tr>
<tr>
<td>2.0</td>
<td>±8</td>
<td>±7</td>
<td>±6</td>
<td>±8</td>
<td>±9</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Year</td>
<td>Predicted $\text{p/cm}^2\text{/yr}$</td>
<td>Actual $\text{p/cm}^2\text{/yr}$</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>------</td>
<td>----------------------------------</td>
<td>---------------------------------</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1972</td>
<td>$7.0 \times 10^9$</td>
<td>$2.4 \times 10^{10}$</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1989</td>
<td>$8.0 \times 10^8$</td>
<td>$4.5 \times 10^{10}$</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Channel Number &amp; Charge Deposition Threshold</td>
<td>29 Sep 1406-1726</td>
<td>19 Oct 1420-1735</td>
<td>20 Oct 0859-1204</td>
<td>22 Oct 1814-2149</td>
<td>24 Oct 1805-2135</td>
<td></td>
<td></td>
</tr>
<tr>
<td>--------------------------------</td>
<td>----------------</td>
<td>----------------</td>
<td>----------------</td>
<td>----------------</td>
<td>----------------</td>
<td></td>
<td></td>
</tr>
<tr>
<td>1 19 fC</td>
<td>3.7 ± 0.02</td>
<td>1.6 ± 0.01</td>
<td>1.4 ± 0.01</td>
<td>1.5 ± 0.01</td>
<td>3.4 ± 0.01</td>
<td></td>
<td></td>
</tr>
<tr>
<td>2 46 fC</td>
<td>4.9 ± 0.1</td>
<td>1.9 ± 0.04</td>
<td>1.6 ± 0.04</td>
<td>1.8 ± 0.04</td>
<td>4.5 ± 0.06</td>
<td></td>
<td></td>
</tr>
<tr>
<td>3 110 fC</td>
<td>5.7 ± 0.1</td>
<td>2.1 ± 0.07</td>
<td>1.8 ± 0.07</td>
<td>1.9 ± 0.07</td>
<td>5.2 ± 0.1</td>
<td></td>
<td></td>
</tr>
<tr>
<td>4 260 fC</td>
<td>5.9 ± 0.2</td>
<td>2.0 ± 0.1</td>
<td>1.8 ± 0.1</td>
<td>2.0 ± 0.1</td>
<td>5.7 ± 0.2</td>
<td></td>
<td></td>
</tr>
<tr>
<td>5 610 fC</td>
<td>5.6 ± 0.6</td>
<td>2.0 ± 0.3</td>
<td>2.0 ± 0.4</td>
<td>2.1 ± 0.3</td>
<td>4.9 ± 0.4</td>
<td></td>
<td></td>
</tr>
<tr>
<td>6 1.50 pC</td>
<td>6.1 ± 1.5</td>
<td>3.0 ± 0.7</td>
<td>1.1 ± 0.8</td>
<td>1.0 ± 0.6</td>
<td>4.3 ± 1.1</td>
<td></td>
<td></td>
</tr>
<tr>
<td>7 3.40 pC</td>
<td>(17.4 ± 17.4)</td>
<td>-</td>
<td>(30.4 ± 30.4)</td>
<td>-</td>
<td>-</td>
<td></td>
<td></td>
</tr>
<tr>
<td>8 8.10 pC</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td></td>
<td></td>
</tr>
<tr>
<td>9 19.3 pC</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td></td>
<td></td>
</tr>
</tbody>
</table>
High Latitude Homs

Inner Belt

Outer Belt

Slot

Fig. 1 – See cap. in Paper

SAMPEX/P1ADC: Electrons E > 0.4 MeV

Fig. 2 – See cap. in Paper

Energy = 25 - 250 MeV/n - IMP-8

Fig. 3 – See cap. in Paper

Primary Heavy Ion Incident on Atmosphere

Low Energy Nucleonic Component

Disintegration Product Neutrons

Low Energy Protons to "Grow"

N, P = High Energy Nucleons

n, p = Disintegration Product Neutrons

= Nuclear Disintegration

Fig. 4 – See cap. in Paper

Energy Feeds Across from Nuclear to Electromagnetic Interactions

Small Energy Feedback from Meson to Nucleonic Component

Fig. 5 – See cap. in Paper

Fig. 6 – See cap. in Paper

Earth Radii
Flux Differential Omindirectional (MeV\(^{-1}\) cm\(^{-2}\) s\(^{-1}\))

**Fig. 13** – See cap. in Paper

Energy: 100 keV

Time: 00:25:00 T=TU

**Fig. 14** – See cap. in Paper

P8 (80 MeV), B/B\(_{min}\)=1.0

**Fig. 15** – See cap. in Paper

**Fig. 16** – See cap. in Paper

**Fig. 17** – See cap. in Paper

**Fig. 18** – See cap. in Paper
Minimize Risk in Design Phase
"Climate" Models

Avoid Risk in Operations
"Forecasting" Models

Reevaluate Risk in Anomaly Resolution
"Nowcasting" Models

Fig. 19 – See cap. in Paper

Fig. 20 – See cap. in Paper

Fig. 21 – See cap. in Paper

Fig. 22 – See cap. in Paper