Radiation belt electron dynamics: modeling atmospheric losses
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The first year of work on this project has been completed. This report provides a summary of the progress made and the plan for the coming year. Also included with this report is a preprint of an article that was accepted for publication in *Journal of Geophysical Research* and describes in detail most of the results from the first year of effort.

The goal for the first year was to develop a radiation belt electron model for fitting to data from the SAMPEX and Polar satellites that would provide an empirical description of the electron losses into the upper atmosphere. This was largely accomplished according to the original plan (with one exception being that, for reasons described below, the inclusion of the loss cone electrons in the model was deferred). The main concerns at the start were to accurately represent the balance between pitch angle diffusion and eastward drift that determines the dominant features of the low altitude data, and then to accurately convert the model into simulated data based on the characteristics of the particular electron detectors. Considerable effort was devoted to achieving these ends. Once the model was providing accurate results it was applied to data sets selected from appropriate periods in 1997, 1998, and 1999. For each interval of ~30 to 60 days, the model parameters were calculated daily, thus providing good short and long term temporal resolution, and for a range of radial locations from $L = 2.7$ to $3.9$.

The results of the program just described provided some new insights into radiation belt dynamics. First, it become clear early on, that in order to obtain accurate fits to the data it would be necessary to include a local time dependence to the pitch angle diffusion rates in the model. This led to some additional computational complexity and necessitated a fairly considerable effort to achieve, although there was no fundamental difficulty in its implementation. The results showed that electron losses by diffusion into the atmosphere are generally stronger at daytime than at nighttime by a considerable and variable factor. Secondly, the temporal variations in the loss rates are significantly greater than those observable from the decay of the stably-trapped electrons, and are positively correlated with geomagnetic activity. Finally, for electrons inside the plasmasphere, there is not a considerable variation in loss rate with radial location. These results are generally consistent with the view that radiation belt electron losses inside the plasmasphere are due to scattering by whistler mode plasma waves ("plasmaspheric hiss") that are excited largely on the dayside of the magnetosphere during times of recovery from geomagnetic storms. While plasma wave data are generally sufficient to confirm this picture only on a statistical basis, the results described above show that the low altitude electron observations can provide a proxy for the wave data and for determining the scattering rates into the atmosphere on a real-time basis. All of these results are described in detail in the attached preprint.
The initial goal was to include loss cone electrons in the empirical modeling described above. However, this proved to be impractical because of substantial additional complexity that would be required. It was possible to show that their neglect did not significantly impact the quasi-trapped and stably-trapped electron fluxes and therefore the empirical loss rates determined from the model. Therefore, the calculations described above were completed with the simpler model and the inclusion of the loss cone electrons was delayed for the next stage of analysis which is briefly described below. Some progress has already been made in this stage and the rest will be completed in the second year.

The inclusion of the loss cone electrons, those that will encounter the dense atmosphere prior to completing their bounce motion, will allow for the simulation of all electron populations observed by a low-altitude satellite (such as SAMPEX). The bounce dependence of this population requires the inclusion into the model of an additional coordinate (this distance along the magnetic field), hence the additional complexity. It will also be necessary to account for electron backscattering from the atmosphere. This will be done using results from a Monte Carlo transport code (EGS4) that have already been provided by a colleague at no cost to this project (M. D. Looper, private communication). They will provide the atmospheric boundary condition for the model solution. In addition, a theoretical model for the pitch angle diffusion rate based on plasma wave theory will be incorporated, as will a local time dependence based on plasma wave data. These improvements will place the work on a stronger theoretical basis while reducing the required number of adjustable parameters in the model (possibly to zero if accurate plasma wave data are available). Therefore it will be possible, for selected event intervals, to provide a complete picture of the connection between plasma wave scattering and electron losses, all the way from the stably trapped population, through the drifting quasi-trapped population, and into the atmospheric loss cone.
Atmospheric losses of radiation belt electrons

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Abstract. A numerical model of the low altitude energetic electron radiation belt, including the effects of pitch angle diffusion into the atmosphere and azimuthal drift, predicts lifetimes and longitude dependent loss rates as a function of electron energy and diffusion coefficient. It is constrained by high altitude (\(\sim 20,000 \text{ km}\)) satellite measurements of the energy spectra and pitch angle distributions and then fit to low altitude (\(\sim 600 \text{ km}\)) data that are sensitive to the longitude dependence of the electron losses. The fits provide estimates of the parameterized diffusion coefficient. The results show that the simple drift-diffusion model can account for the main features of the low altitude radiation belt inside the plasmasphere during periods of steady decay. The rate of pitch angle diffusion is usually stronger on the day side than on the night side, frequently by a factor \(\sim 10\). The average derived lifetimes for loss into the atmosphere of \(\sim 10\) days are comparable to the observed trapped electron decay rates. Considerable variability in the loss rates is positively correlated with geomagnetic activity. The results are generally consistent with electron scattering by plasmaspheric hiss as the primary mechanism for pitch angle diffusion.

1. Introduction

The substantial role of pitch angle diffusion as a mechanism for radiation belt electron loss into the atmosphere has been qualitatively demonstrated by low altitude satellite data [Imhof, 1968; Sheldon, 1991]. However, limitations in spatial coverage and the problem of determining which electrons subsequently will be lost make it difficult to directly measure the loss rate. Observations of trapped electron decay rates [Williams et al., 1968; Roberts, 1989; Pesnell et al., 2001] provide an indirect estimate but can be influenced by other processes such as radial diffusion. Theoretical calculations of pitch angle diffusion based on scattering by plasma waves and Coulomb collisions [Lyons et al., 1972; Abel and Thorne, 1998] are in reasonable agreement with the observed decay rates, but a direct determination of the electron losses probably requires a combination of the theoretical and experimental approaches. Fortunately, the non-dipolar component of the geomagnetic field (or, equivalently, the offset of the dipole from the center of the Earth) causes losses to occur non-uniformly in longitude and the resulting redistribution of electrons by azimuthal drift provides the necessary observational constraints for theoretical modeling of the atmospheric losses. The validity of this approach to modeling the observed low altitude electron distribution was demonstrated by Imhof [1968], based on an approximate analytic solution to the diffusion equation. A similar concept with a simplified construction was described by Sheldon [1991]. Detailed simulations by Abel and Thorne [1999] accounted for both diffusion and drift but were not compared directly to electron data. These works suggested that a simple balance between azimuthal drift and pitch angle diffusion can represent the low altitude electron distribution. Our goal in this work is to further explore the validity of this idea using an extensive satellite electron data set. Time periods of steady electron decay and locations inside the plasmasphere but near the peak intensity of the outer radiation belt are chosen for study. We start by describing a parameterized theoretical model, then apply constraints from high altitude electron data taken on the Polar satellite to simulate observations from the low altitude SAMPEX satellite, and finally obtain estimates of the diffusion coefficients and associated loss rates by determining the model parameters that best fit the data.

2. Drift-diffusion model

A simplified description of radiation belt electron dynamics above the upper atmosphere is

\[
\frac{\partial f}{\partial t} + \omega_d \frac{\partial f}{\partial \phi} = \frac{1}{x T'(y) \partial x} \left( x T(y) D_\alpha \frac{\partial f}{\partial x} \right)
\]

(1)

where the left hand side is an advective time derivative including the effects of azimuthal drift and the right hand side is the rate of pitch angle diffusion. The equation models the electron phase space density \(f(x, \phi, t)\) that is averaged over the phases of the gyration and bounce motions, but it retains the dependence on the drift phase \(\phi\) and azimuth \(\phi\). It also depends on time \(t\) and \(x = \cos \alpha_0\), where \(\alpha_0\) is the equatorial pitch angle. The \(x\) diffusion coefficient due to pitch angle scattering is \(D_\alpha\). For the diffusion term [Schulz and Lanzerotti, 1974, p. 56], \(T\) is proportional to the electron bounce period and we approximate a dipole magnetic field with [Davidson, 1976] \(T(y) = 1.380173 - 0.639693 y^{0.737}\), where \(y = \sin \alpha_0\). The electron energy determines the drift frequency \(\omega_d\), which we also approximate as being independent of \(\phi\) as in a dipole field.

Electron losses are included in the model by the low altitude boundary condition, for which we assume \(f = 0\) at the edge of the bounce loss cone (BLC). (High or low altitude is equivalent to high or low \(\alpha_0\) and to low or high \(x\)). We define the BLC as the range of equatorial pitch angles where an electron’s adiabatic mirroring point reaches below an altitude of 100 km in either hemisphere. The non-dipolar nature of the Earth’s magnetic field at low altitudes causes the BLC angle for a given drift shell to vary with \(\phi\). For the high altitude boundary condition we require \(df/dx = 0\) at \(x = 0\) (\(\alpha_0 = 90^\circ\)).

The nature of the solutions to the model equation (1) is that the details of any initial condition are rapidly lost, after which a characteristic dependence of \(f\) on \(x\) and \(\phi\) is obtained and decays away exponentially in time at a rate determined by the diffusion coefficient.
The characteristic shape during the exponential decay phase is also the lowest order eigenmode of the combined drift-diffusion operator in (1), with the corresponding eigenvalue representing the decay rate. Higher order eigenmodes may be included in the initial condition but have faster decay rates. The lowest order eigenmode is a function of \( x \) and \( \phi \) that depends only on the dimensionless ratio \( D_{\text{ex}}/\omega_0 \). Some examples are shown in Figure 1. They were obtained by numerically solving (1) by a finite-differencing technique for a time sufficiently long to reach the exponential decay phase and then normalizing each solution by its value at the high altitude boundary \( \alpha_0 = 90^\circ \). The low altitude boundary is the bounce loss cone for a given \( L \) shell value, calculated from the IGRF magnetic field model [Barton, 1997]. The azimuthal angle \( \phi \) is the dipole east longitude which increases in the same direction as the electron drift and is measured relative to the Earth’s dipole axis, starting from the intersection of the dipole equator with the geographic prime meridian (thus it is nearly equivalent to geographic east longitude). The solutions are shown for only a small range of \( x \) near the BLC. At lower \( x \) values (higher mirror point altitudes) the \( \phi \) dependence is insignificant (the solutions in this region are described in the next section).

The variation with longitude of the BLC angle that forms the lower boundary of the model is determined primarily by the South Atlantic anomaly (SAA). It is the region of weak magnetic field that, for these fixed \( L \) values, is centered near \( \phi = 0^\circ \) and causes the significant peak in the BLC angle at that location.

The choice of a functional dependence for \( D_{\text{ex}} \) on \( x \) and \( \phi \) offers a great variety of possible solutions. Guided by observations as detailed in the following sections, we have chosen the parameterized form

\[
D_{\text{ex}} = D_{\text{day}/\text{night}} \left( E^{-\eta} \right)^{1/(\epsilon + x^\sigma)}
\]

where the four free parameters are \( D_{\text{day}}, D_{\text{night}}, \mu, \) and \( \sigma \). The electron kinetic energy in MeV is \( E \). Independent values of \( D_{\text{day}} \) and \( D_{\text{night}} \) are allowed for daytime and nighttime respectively, as determined by the magnetic local time (MLT). The day-night, or MLT, dependence is translated into a corresponding dipole longitude (\( \phi \)) dependence based on a given universal time (UT). A fixed value of \( \epsilon = 10^{-4} \) is included only for numerical stability reasons. It prevents \( D_{\text{ex}} \) from becoming too large at low \( x \) values without significantly impacting the final solution. In Figure 1 the values of \( D_{\text{ex}}/\omega_0 \) refer to \( x \approx 1 \) (the region shown in the figure) and to the day side value of \( D_{\text{ex}}/\omega_0 \). Note that the energy dependence \( E^{-\eta} \) is included in this ratio but it is not significant for the individual model solutions. Its significance will become apparent when the solutions are combined for modeling observations over a range of electron energies.

Referring to Figure 1, the three solutions on the left hand side (labelled a, b, and c) have equal day and night values of \( D_{\text{ex}} \) (day/nights = 1) but the ratio \( D_{\text{ex}}/\omega_0 \) decreases by a factor of 10 in each case. In the first case (a), \( D_{\text{ex}}/\omega_0 \) has the high value of \( 10^{-5} \) and diffusion is dominant over drift. The solution reflects the \( \phi \) dependence of the BLC angle, although there is a small influence of the drift as shown by the small eastward shift of the higher altitude contours. As the ratio \( D_{\text{ex}}/\omega_0 \) decreases (b and c), drift becomes increasingly important relative to diffusion. This is shown by the contours becoming independent of \( \phi \) in the stable trapping region, that is the region of \( \alpha_0 > 8.5^\circ \), the highest BLC angle for \( L = 3.5 \), and by the relative lack of electrons in the drift loss cone (DLC) to the east of the SAA. The DLC is the range of \( \alpha_0 \) values where electrons are not in the BLC but cannot complete a full drift without entering the BLC at another longitude. The DLC is relatively empty when \( D_{\text{ex}}/\omega_0 \) is small because the diffusion is too slow to fill it before drift takes the electrons toward the SAA where they are lost. At the low value of the ratio \( D_{\text{ex}}/\omega_0 = 10^{-5} \) (c) the \( f \) value decreases by several orders of magnitude with decreasing \( \alpha_0 \) in the DLC. These three cases show that if diffusion is dominant then losses into the atmosphere are distributed evenly in longitude, but that as drift become more significant the losses, though still caused by diffusion, occur over a decreasing range of longitudes to the west of the SAA.

The three solutions on the right hand side of Figure 1 (labelled d, e, and f) all have a day side value of \( D_{\text{ex}}/\omega_0 = 10^{-5} \) and a night side value that is 10 times smaller (day/night = 10). They differ only in the UT hours of the simulation which are (d) 6, (e) 12, and (f) 18. Now diffusion is dominant on the day side while drift is dominant on the night side, as can be seen from the MLT (magnetic local time) labels in each case. In the first two (d and e) the DLC region to the east of the SAA (0 to 90° dipole longitude) is on the dayside and is filled in by diffusion. Then the eastward drift fills in the rest of the drift loss cone so that these solutions are similar to the case (a) in which diffusion was dominant throughout. In the third case (f) the DLC region to the east of the SAA is on the nightside where diffusion is weak and is therefore relatively empty, while to the east of 180° dipole longitude the DLC is on the dayside and is filled in by diffusion. There is no possibility, at any UT hour, for the eastern half (>180°) of the DLC to be relatively empty compared to the western half (<180°), because the eastern half is always filled in either by diffusion from lower \( x \) or by drift from the west.

The examples of Figure 1 also illustrate the variation in the BLC angles with \( L \) and the variation in the model solution with the \( \sigma \) parameter. The first three cases (a, b, and c) are for \( L = 3.5 \) and \( \sigma = 40 \). The next three cases (d, e, and f) are for \( L = 3.1 \) and \( \sigma = 20 \).

3. Model lifetimes

The decay lifetimes of the model solutions, or equivalently the eigenvalues associated with the lowest order eigenmodes, depend primarily on \( D_{\text{ex}} \) and are nearly independent of \( \omega_0 \). This is because the BLC angles for a given \( L \) vary over only a small range of \( x \) and the lifetimes are therefore well approximated by those that would be obtained with an average value of the BLC angle, independent of \( \phi \). In the case that \( D_{\text{ex}} \) is also independent of \( \phi \) we can drop the \( \phi \) dependence of the solution altogether and solve a simpler equation in \( x \) only. These solutions are illustrated in Figure 2. It shows the normalized pitch angle distributions for various values of \( \alpha \), where \( D_{\text{ex}} \) is the same \( x \) dependence as before. The BLC angle is that for \( L = 3.5 \) in a dipole magnetic field.

The increasing slope of the pitch angle distributions near the BLC angle with increasing \( \alpha \) suggests an inverse dependence of the model lifetimes on \( \alpha \). This is illustrated in Figure 3 which shows the normalized e-folding lifetimes \( \tau D \) versus \( \alpha \), where \( \tau \) is the lifetime, versus \( \alpha \) for various dipole \( L \) values. (If the approximation \( T = 1 \) is made in equation (1) then the 1-dimensional eigenmodes and eigenvalues can be obtained analytically [Roberts, 1969].) The lifetimes \( \tau \) are similar to the numerical results obtained here.

In the case where \( D_{\text{ex}} \) has a day/night asymmetry the 1-dimensional solutions (Figure 2) and the normalized lifetimes (Figure 3) are still good approximations for \( x \) values other than those near the BLC because they are independent of \( D \). The lifetimes \( \tau \) are then simply obtained from the normalized lifetimes after division by the mean of the day and night \( D \) values.

4. High altitude data

At high altitudes, where equatorial pitch angles \( \alpha_0 \) are well away from the loss cones, phase mixing by azimuthal drift causes the electron intensity to be independent of drift phase. Therefore measurements made at a single longitude can be used to normalize the model for simulating low altitude data. Electron energy spectra from \( L = 3.5 \) measured at high altitude on the Polar satellite are shown in Figure 4. They were taken during a 40-day period in 1998 following an electron injection event \(~10 \) days earlier. These data were obtained by HISST (the High Sensitivity Telescope) [Blake et al., 1995] and analyzed using techniques described previously [Selesnick and Blake, 2000]. During this period the altitude of the elliptical Polar
orbit at \( L = 3.5 \) was \( \sim 20,000 \) km. We also use similar data from other \( L \) shells.

We have also chosen two other time periods, from mid-1996 and mid-1997, with similar characteristics of electron decay. The measured electron intensity at \( L = 3.5 \) from all three periods, for selected energies, is shown versus time in Figure 5. It demonstrates the generally steady decay at the lower energies and relatively constant levels at the higher energies.

The periods of generally decreasing intensity illustrated in Figure 5 were chosen so that the pitch angle distributions (PADs) are likely to have reached the exponential decay phase discussed earlier. Examples of the PADs measured by HIST near the end of the 1998 time interval are shown in Figure 6 for several electron energies. The shapes of these distributions are generally independent of energy and measurements from earlier times show that they are also generally independent of time during the decaying period.

The Polar data are used to normalize the model at high altitudes. The Polar altitude at the time of the measurement was such that equatorial pitch angles between \( \alpha_0 \approx 50^\circ \) and \( 130^\circ \) were not observable. Therefore, the normalization is done at \( \alpha_0 = 40^\circ \). This is easily sufficient for the distributions to be independent of drift phase. Measurements from other time periods show that the distributions at the angles that are unobserved in this case are essentially flat [e.g., \cite{Selesnick and Blake, 1998}], consistent with the high-\( L \) distributions (Figure 2).

The measured PADs are not well resolved near the loss cone angles, \( \alpha_0 \approx 20^\circ \) and \( \alpha_0 \approx 160^\circ \), so we cannot expect an accurate comparison with the model distributions in those ranges. Low altitude data are necessary for such comparisons.

5. Low altitude data

Once the model is normalized by the high altitude data it can be used to simulate a particular set of low altitude electron data. We use data taken by PET (the Proton/Electron Telescope) on the SAMPEX satellite [\cite{Cook et al., 1993}]. Its high-inclination, low-altitude (\( \sim 600 \) km) orbit crosses each radiation belt \( L \) shell approximately 60 times per day covering all longitudes and providing a suitable data set for comparison with the model results, as illustrated in Figure 7 for \( L = 3.5 \).

The equatorial pitch angles of electrons mirroring at SAMPEX (upper part of Figure 7) are a little above the BLC angles at the same longitude in each hemisphere. However, PET does not measure only the mirroring electrons (or the pitch angle distribution) but a weighted average of electrons mirroring at and below SAMPEX that depends on the satellite pointing direction relative to local magnetic field. The variations in the pointing direction due to the four separate crossings of \( L = 3.5 \) per orbit cover a wide range of local pitch angles (lower part of Figure 7). During the 1998 period of days 150 to 170 the local times of the SAMPEX orbit were within approximately 1 hour of the noon-midnight plane. The satellite was rotating at 1 revolution per orbit, hence the regular pattern of pointing directions. During the 1996 and 1997 periods it was rotating at 1 revolution per minute, providing a more random distribution of pointing directions at a given \( L \) shell for those periods.

We consider data taken by three PET electron rate counters, labelled P1, ELO, and EHI, that measure a low, medium, and high range of electron energies respectively, together covering the range from \( \sim 0.5 \) to 14 MeV. Samples of the data are shown in the next section. The flux (counting rate) from each counter is related to the electron intensity \( f(E, \alpha_0) \) at each SAMPEX location by response functions that are described in the Appendix. These take into account the energy and angular response of each counter for a given pointing direction and are combined with the model results to simulate the counting rate at a given position on the orbit. The model is parametrized by the ratio \( D_e/\alpha_0 \), so the energy \( E \) is determined for a given value of the diffusion coefficient \( D_e \) from the drift frequency. Therefore a range of model solutions, corresponding to the appropriate range of energies, enters into the simulation of each rate counter.

6. Low altitude simulations

To provide strong constraints on the model parameters the low altitude data must sample the full range of longitudes and UT hours, so that a full day of data are required. A sample set of data and the corresponding model simulation from 1998 day 159, \( L = 3.5 \) (the data corresponding to Figure 7) is shown in Figure 8. All of the PET flux measurements from the three rate counters (interpolated to \( L = 3.5 \)) are shown, as a function of dipole longitude as in the previous figures, by the filled data points. The open points show the simulation that is the best fit to the data. The variations with longitude can be understood with reference to Figure 7. High fluxes from the southern hemisphere in the vicinity of the SAA (\( \sim 0^\circ \) and \( 360^\circ \) dipole east longitude) are stably trapped electrons that are primarily at equatorial pitch angles above the highest BLC angle. Low fluxes in the same longitude range but from the northern hemisphere are electrons in the BLC. They are generally \( \lesssim 1\% \) of the stably trapped southern hemisphere fluxes because, with backscattering from the atmosphere, electrons in the BLC can remain trapped for at most a few bounces. The model does not include any electrons in the BLC so there are no simulations corresponding to the data in that region. Low fluxes to the east of the SAA (dipole east longitude \( \geq 60^\circ \)), from both north and south, gradually increase with longitude as the DLC is filled in by diffusion. On this date, the day side diffusion coefficient was significantly (~30 times) greater than the night side. Therefore, during the first half of the day the western part of the DLC region (\( \lesssim 180^\circ \)) showed substantially higher fluxes than during the second half of the day (compare Figures 1d and 1f). The day-night asymmetry accounts for most of the large range of observed and model fluxes at a given longitude in that region, although substantial (factor \( \sim 10 \)) variations can also be caused by differences in either the SAMPEX pointing direction or the equatorial pitch angle of locally mirroring electrons.

To further emphasize the strong day-night asymmetry apparent in the data from 1998 day 159, the same case is shown in Figure 9 but without the day-night asymmetry in the model simulation. The model parameters are the same as in the previous figure except that the night side diffusion coefficient is increased to the same value as on the day side. Now there is a clear mismatch between the lower data points observed in the DLC region during the second half of the day and the much higher simulated values, that are similar to those observed and simulated during the first half of the day. The model with the strong day-night asymmetry was clearly a much better fit to the data.

The example of 1998 day 159 is also useful in illustrating the sensitivity of the model parameters to a range of diffusion coefficients. The factor \( \sim 30 \) variation from day to night produced a substantially larger flux variation in the western longitudes of the DLC region. The model diffusion coefficient is therefore well constrained by the data. The same would be true if there were no local time dependence and the diffusion coefficient were either low throughout or high throughout (as shown by the simulation in Figure 9).

The best fitting model for 1998 day 159 produced a reasonably accurate simulation for most of the data points, considering that there are only 4 model parameters and ~50 data points. However, there are some systematic deviations between the simulations and the data. This is seen more clearly in Figure 10 which shows the best fit to the data from 1998 day 150 at \( L = 2.7 \). Despite the lower \( L \) shell the general nature of the data and simulation are similar to those of Figure 8. As in that case, and in other similar cases, a subset of the data are substantially below the model simulations. These points occurred at similar times (\( \sim 2300 \) UT) and longitudes (\( \sim 180^\circ \) to \( 270^\circ \)) on each day, suggestive of systematic errors in the model. For example, assuming a greater local time extent of the low diffusion region on the night side, or a more gradual transition from the day to night diffusion coefficients, may improve the fits.

The similar effect on the simulated data to that of the day-night asymmetry can be obtained by assuming that the diffusion coefficient has no local time asymmetry but varies with UT, being higher in the first 12 UT hours of the day than in the second 12 UT hours.
(this is evident from the similarity of Figures 1a with 1d and 1c with 1f). In some cases, usually those showing significant systematic errors with the day-night asymmetry model such as on day 150 (Figure 10), the fit can be slightly improved by using this alternate model. In other cases the reverse is true. The day-night asymmetry seems preferable to a UT periodicity in the diffusion coefficient primarily because of our prior knowledge of the diffusion mechanism, as discussed further below.

A final example is shown in Figure 11 from 1998 day 151 and L = 2.7, the day after the case of Figure 10. The stably trapped electrons (near 0 and 360° dipole longitude) are almost unchanged from the previous day, but the data and simulations in the DLC region (~60° to 300°) are substantially lower. In the simulations, this is a consequence of a significant decrease in the diffusion coefficients, both at day and night, as shown in the model parameters. This must be a true decrease, rather than modeling error, because it occurred in the data taken throughout the day. While the change in $D_{\text{as}}$ occurred near the boundary of days 150 and 151, it is likely that such variations can also occur at other times during a day, causing some mismatches between the data and simulations.

### 7. Parameter estimates

To find the model parameters that best fit each 1-day subset of the SAMPEX data we minimize the function $\chi^2 = \sum (y_i - f_i)^2 / \sigma_i^2$ where $y_i$ = ln $F$, $f_i$ is the observed flux, $\sigma_i$ is the simulated value of $y_i$, and $\sigma_i$ (unrelated to $\sigma$ in (2)) is the standard uncertainty in $y_i$. For $\sigma$, we include uncertainties due to Poisson counting statistics, which are significant for the low fluxes, and an additional 10% modeling uncertainty to prevent the fit from being constrained by only the high flux points. This can include the small uncertainties (~5%) in the instrumental response functions (Appendix) because they lead to systematic relative errors in the simulation of the same type as any modeling errors. However, the assumed uncertainties are meant only to provide reasonable weighting factors to the data and clearly do not account for the actual mismatch from the simulation (e.g. Figure 10). Uncertainties in the final model parameters are estimated by the standard least-squares techniques, but to account for the systematic deviations of the model from the data, at least in an approximate way, the parameter variances are multiplied by the reduced $\chi^2$, that is by $\chi^2 / \nu$ where $\nu$ is the number of degrees of freedom (number of data points minus number of fit parameters). The fit parameters were allowed to vary in the ranges of $5 \times 10^{-10}$ to $1 \times 10^{-6}$ s$^{-1}$ for the larger of $D_{\text{day}}$ and $D_{\text{night}}$, 0.01 to 100 for the day/night ratio, -5 to 5 for $\mu$, and 10 to 20 for $\sigma$.

Examples of model fits were shown in the previous section. Results from the full 1998 time period and $L = 3.5$ are shown in Figure 12. The day side diffusion coefficients for electron energies of $E = 1, 2, 4$ MeV and for $x = 1$ are calculated from the estimated parameters by $D_{\text{as}} = D_{\text{day}} E^x$. The 2 MeV values are generally the best constrained by the particular response functions of the three PET rate counters. The 1 and 4 MeV uncertainties are therefore somewhat higher than for the 2 MeV case but, for clarity, only the uncertainties at 2 MeV are shown in the figure. The day/night ratio is simply $D_{\text{day}} / D_{\text{night}}$ and neither it nor $\sigma$ (the parameter for the $x$ dependence) are functions of energy. The uncertainties shown for $\sigma$ are probably larger than necessary because that parameter is constrained primarily by the stably trapped electrons near the SAA that are less influenced by the systematic modeling errors in the DLC region. The variations of electron flux with longitude in the DLC region are the primary constraints on both the day side diffusion coefficient and the day/night ratio. As described above, there are variations in the estimated parameters from day to day, but there are also trends that extend over periods of several days. The relative quality of fits can be judged from the values of $\chi^2$ also shown in the figure, but the general trend of decreasing $\chi^2$ is a result of the gradually decreasing trapped electron flux relative to the fixed background levels and does not indicate improving fit quality.

Estimated parameters from 1998, $L = 3.1$ are shown in Figure 13. The trends are similar to those at $L = 3.5$ (Figure 12), with some systematic differences. The corresponding results from 1996, $L = 3.9$ are shown in Figure 14. Rather than including all of the available results, a summary is given in Table 1. For each of the three time periods the means and standard deviations of the estimated $D_{\text{as}}$ (at $E = 2$ MeV and $x = 1$), $\mu$ and $\sigma$ values are listed. The standard deviations of the daily estimates are included, rather then the uncertainties in the means (which would be much smaller), in order to approximate the degree of variability in each case. The lower $L$ values are not available for the 1996 and 1997 periods because the Polar satellite orbit did not reach low L at those times.

The estimated model parameters can be used to derive 1/fold lifetimes, as in Figure 3, based on the mean of the day and night diffusion coefficient from each fit. These are shown for $L = 3.5$ and each of the time intervals in Figure 15. The geomagnetic index $D_s$ is also shown in the figure for comparison with the variations in lifetime. To show correlations more clearly, the $E = 2$ MeV lifetimes are shown versus $D_s$ for the four selected L shells and all three time periods combined (where available) in Figure 16. This figure also provides a summary of all of the lifetime results from the three time periods.

### 8. Discussion

We have shown that a combination of azimuthal drift and pitch angle diffusion into the upper atmosphere describes the essential features of low altitude radiation belt energetic electron data during times of steady decay. The loss rate is determined by the rate of diffusion, but drift must also be included to account for the observed azimuthal distribution of electrons in the drift loss cone and that distribution constrains the value of the diffusion coefficient.

The model was also constrained by data at high altitudes that provided the upper boundary condition for the model simulations. The high altitude data are not essential because the electron intensity there could be an additional free parameter, but their inclusion provides a much stronger constraint on the equatorial pitch angle ($\theta_0$) dependence of the diffusion coefficient than would the low altitude data alone, while the agreement of the model with both data sets increases our confidence in the final results.

The low altitude data provided most of the critical information for determining the diffusive loss rates because they were obtained within and just above the electron drift loss cone (DLC), that is the region where the balance between drift and diffusion determines the longitudinal electron distribution. It is also only at low altitude that the satellite orbital period is short enough to provide good longitudinal coverage on a daily time scale. It was critical to the accuracy of the final results to include accurate and detailed information on the angular and energy dependencies of the electron detectors as well as the satellite orbital and pointing data.

The solutions of the drift-diffusion model used to fit the data are applicable to steady electron decay. Transient effects of varying magnetospheric conditions were not included. However, the resulting parameter estimates in fact did show significant variability in the diffusion coefficient on a daily time scale. For a diffusion coefficient of the form $D_{\text{as}} = D_c E^x (x = \cos \theta_0)$, a change in the coefficient $D_c$ causes a change in the diffusion rate and the longitudinal electron distribution in the DLC, but not in the shape of the high-altitude pitch angle distribution. Therefore, the electron distribution can readjust to the change in $D$ on a time scale of only one drift period, or ~10 to 30 minutes for an electron energy range of 5 to 0.5 MeV. Daily variations in the fit parameters relating to the diffusion rate, that is $D_{\text{day}}$, $D_{\text{night}}$, and $\mu$ in the notation of our model, can therefore be accommodated by steady decay solutions. The same is true for the daily variation at a fixed longitude caused by any day-night asymmetry in $D_{\text{as}}$.

Variations in $\sigma$, which provides the model $x$ dependence, require longer time scales for the pitch angle distribution to adjust to new steady decay state. Time scales on the order of the electron lifetimes would be required for significant changes. Therefore, changes in the estimated $\sigma$ values should be considered reliable only on time scales of several days or more. However, since we are concerned primarily with the diffusive loss rates obtained from
the low altitude electron distributions, this is not a significant short-
coming of our analysis.

The model included a day-night asymmetry in the diffusion co-
efficient. This was required to fit 1-day subsets of the data, al-
though it would also be possible to use a daily periodicity in the
time dependence of the diffusion coefficient rather than the local
time dependence. The data required significantly faster diffusion
on the day side than on the night side for most of the time. Oc-
casionally, the night side diffusion was faster. On the average (Ta-
ble 1), the day/night ratio was ~2 to 3 during the 1996 and 1997
periods but was ~4 to 20 during the 1998 period (being higher at
lower L) due to higher day side diffusion. Daily variations in the
day/night ratio were comparable to the ratio itself but large changes
tended to be isolated rather than randomly distributed. The higher
diffusion rates in daytime compared to nighttime appear to be con-
sistent with observations [Russell et al., 1969; Andr é et al., 2002] of
a similar local time distribution in the plasmaspheric hiss. It is
the ELF electromagnetic emission thought to be the primary cause
of electron pitch angle scattering within the part of the outer zone
radiation belt that is inside the plasmasphere (L ~ 2 to 4) [Abel and
Thorne, 1998].

The estimated values of the diffusion coefficient correspond to
electron diffusive loss lifetimes (averaged over day and night)
that vary over a wide range. Given the uncertainties in the estimates,
the lifetimes are well correlated with the $D_\alpha$ geomagnetic index
(Figure 16), the shorter lifetimes generally corresponding to higher
geomagnetic activity. The average activity level was lowest during
the 1996 period and highest during the 1998 period, leading to gen-
erally weaker and stronger diffusion rates respectively, but the cor-
relation was also evident with the specific small magnetic storms
that occurred during the 1997 and 1998 periods (Figure 15). The
correlation is consistent with the interpretation that scattering by
plasmaspheric hiss is the principal diffusion mechanism, because
the hiss intensity increases during the recovery phase of geomag-
netic storms [Smith et al., 1974].

Observed trapped electron decay rates of ~10 to 20 days (Fig-
ure 5), are in reasonable agreement with average estimated life-
times. However, the variability in the estimated lifetimes (Fig-
ures 15 and 16) is greater than that of the observed decay rates. The
difference is significant for only the occasional isolated shells. It is
usuallly due primarily to the lower L shells, because in those in-
stances there should be significant reductions of the trapped
electron intensity during a single day. Such reductions are not ob-
erved in either the high altitude Polar data or in the trapped com-
ponent of the low altitude SAMPEX data. The clearest example is
1998 day 158 at L ~ 2.7 (Figure 10), for which the estimated
lifetime is 1.5 days, which is significantly shorter than ~6 days for
the stably trapped intensity. Then the stably trapped intensity
should decrease by a factor of ~6 during that day, but in fact it is
not greatly decreased a day later (Figure 11). A possible explana-
tion is that the trapped electrons are replenished during times of
rapid losses into the atmosphere, by either radial diffusion or local
acceleration.

The estimated values of $\sigma$ showed that the diffusion coeffi-
cient increases rapidly with decreasing $x$, or increasing mirror point
altitude, in the low altitude region. This is consistent with the mea-
sured pitch angle distributions that are flat at intermediate pitch an-
gles and decrease rapidly near the loss cones (compare Figures 2 and
6), the flat section being due to relatively fast diffusion at high alti-
itudes.

The $\sigma$ values generally increase with $L$ (Table 1). This may
coincide with a greater relative contribution to electron scattering
from lightning generated whistlers at lower $L$ compared with plas-
maspheric hiss at higher $L$ [Abel and Thorne, 1998; Blake et al.,
2001], on the other spatial variation in the diffusion mechanism.
However, it would be possible to use a daily periodicity in the
time dependence of the trapped electron fluxes measured on the
low and high altitude satellites, its value is sensitive to any errors
in this comparison. The same is not true of the other model param-
eters that are constrained primarily by the low altitude longitudinal
electron distribution.

The energy dependence of the diffusion coefficient was modeled
only in a crude way because of the limited constraints available
from the broad energy response of the three PET electron counters.
Also, the accuracy of the energy dependence can be compromised
by systematic modeling errors and by the assumptions that $\sigma$ and
the day-night ratio are not energy dependent. However, the esti-
mated $\mu$ values (Figure 16) do show some potentially signifi-
cant trends. The mean values are generally positive ($\mu > 0$) increas-
ing with energy) at the lower $L$ values and negative ($\mu < 0$, decreas-
ing with energy) at the higher L values, with significant daily variations.
The diffusion coefficient is best constrained by the data near $E = 2\text{ MeV}$.
It is possible that the results described above could be improved
by more detailed observations, particularly in the energy depen-
dence of the low altitude data. However, the main source of uncer-
tainty currently is in the model. At the expense of added complexity
there are several areas of possible improvement: time dependent
solutions could be used, rather than the steady decay approxima-
tion; the local time dependence of the diffusion coefficient could
be more realistic, possibly guided by plasma wave data; finally, the
model could include other physical processes, such as atmospheric
scattering with diffusion into the bounce loss cone, radial diffusion,
and local acceleration.

Appendix A: PET response

The rate $r$ measured by any of the PET counters, P1, ELO, or
EHI [Cook et al., 1993], is related to the local electron intensity $j$ and the counter response function $R$ by

$$r = \int_0^L \int_0^{2\pi} \int_0\pi \int_0^{\pi/2} \int_0^{\pi/2} R(E, \theta, \phi) d\phi d\theta dB dL.$$  

where $\theta$ and $\phi$ are the spherical polar and azimuthal angles, re-
spectively, relative to the telescope axis, and $\alpha$ is the local pitch angle that is related to the angle $\theta_B$ between the telescope axis and the local magnetic field direction by

$$\cos \alpha = \sin \theta_B \cos \theta - \cos \theta_B \cos \phi.$$  

The response functions were determined experimentally prior to the
SAMPEX launch. For energies from 0.3 to 3 MeV, this was done
using a $^{197}\text{Au}$ radioactive source and a magnet $\beta$-spectrometer at
the California Institute of Technology. For energies from 1.5 to
27 MeV, the Department of Energy linear electron accelerator op-
erated by EG&G in Santa Barbara was used. The calibration results
are shown in Figure A1. The energy response curves (left side of
figure) can be characterized by the approximate energy ranges of
$>0.6\text{ MeV}$ for P1, 1.5 to 6 MeV for ELO, and 2.5 to 14 MeV for
EHI, but the actual energy range of a given measurement depends
on the spectrum. The angular response (right side of figure) also
varies somewhat with energy and between rate counters. The com-
plete calibrated response functions from Figure A1 were used to
build the model intensity $j$ by numerical evaluation of (A1) to sim-
ulate the PET data.

There are some additional corrections that must be made to the
PET rates. Instrumental deadtime, which is significant at high
counting rates, is measured by PET onboard the satellite and we
make a correction for it in the rate data. The single detector P1 data
have a significant background due to penetrating cosmic rays, while
the multi-detector ELO and EHI data have small backgrounds. We
have added empirically determined background values of 80, 0.1,
and 2 counts per $6\text{ s}$ interval to the simulated P1, ELO, and EHI
data respectively. The ELO counter requires coincident measure-
ments in each of the P1 and P2 detectors. At high rates the ELO
data include a significant fraction of chance coincidences, due to
separate low-energy electrons arriving in the P1 and P2 detectors
within the resolving time of the counter, $\tau \approx 2.25\text{ ms}$. The rate
of chance coincidences, which we add to the simulated ELO rate, is
[Knoll, 1989, p. 636] $2\pi r_1 r_2$ where $r_1$ and $r_2$ are the P1 and
P2 rates determined from their respective response functions (the
P2 response function, not shown in Figure A1, was also determined
by calibrations). It was not necessary to correct for chance coinci-
dences in EHI. We have also not made any correction for pileup of
low energy electrons in the P1 measurement because that effect is
significant only for relatively soft spectra.
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References


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Figure 1. Color coded solutions of the model drift-diffusion equation with selected values of the model parameters. They are normalized to $f = 1$ at $\alpha_0 = 90^\circ$. The lower boundary where $f = 0$ is the bounce loss cone angle. Dashed contours are shown every decade in $f$. Cases a, b, and c have equal day and night diffusion coefficients but decreasing rates of diffusion relative to azimuthal drift. Cases d, e, and f all have the same diffusion coefficient, with the daytime value 10 times greater than the nighttime one, but are for different universal times (UT). The $L$ shells and $\sigma$ values also differ ($D_{\nu,t} \sim x^{-0}$).
Figure 2. Model pitch angle distributions for selected $\sigma$ values ($D_{\text{sf}} \sim x^{-\sigma}$) assuming a dipole magnetic field at $L = 3.5$.

Figure 3. Normalized $e$-folding lifetime $\tau D$ versus $\sigma$, where $\tau$ is the lifetime and $D_{\text{sf}} = D x^{-\sigma}$, for selected $L$ values.
Figure 4. Electron energy spectra measured at high altitude on the Polar satellite, averaged over 5-day intervals centered on the labelled day numbers.
Figure 5. Electron intensity versus time for three selected time intervals, measured at α₀ = 40° and at the indicated energies, from the Polar satellite (units as in Figure 4). The Dₙ geomagnetic index is also shown.
Figure 6. Electron equatorial pitch angle distributions measured simultaneously at the labelled energies from the Polar satellite, with normalization factors chosen for clarity of display. The smooth curves, with dashed uncertainty ranges, are the pitch angle distributions after deconvolution from the instrumental angular response and the satellite rotation sector width [Selesnick and Blake, 2000]. Data points are centered on each sector.
Figure 7. (above) Equatorial pitch angles as a function of longitude for electrons mirroring at the low altitude SAMPEX satellite. Each orbital crossing of $L = 3.5$ in the northern or southern hemisphere during a one day period is indicated, with the loss cones angles for each hemisphere. (below) Local pitch angles of the SAMPEX (and PET instrument) pointing directions at the same locations.
Figure 8. SAMPEX/PET electron data and model simulations for 1998 day 159. The data points are 6 s rate accumulations, interpolated to $L = 3.5$, from the 3 rate counters, P1, ELO, and EHI, that respond to electrons from the approximate indicated energy ranges. The calibrated energy and angular response functions (see Appendix) were used in the simulations. Dashed lines connect each data point to its corresponding simulated value. The model parameters were chosen to give the best fit to the data and included a high day/night diffusion ratio.
1998 day 159, $L = 3.5$, high $D = 6.3 \times 10^{-8} \text{ s}^{-1}$, day/night = 1.0, $\mu = -0.90$, $\sigma = 52.0$

**SAMPEX data:** ▲ North ▼ South

**Simulation:** ▲ North ▼ South

**Figure 9.** Similar to Figure 8 except that the nighttime diffusion coefficient was set equal to the daytime value from the best fit.
Figure 10. Similar to Figure 8 but for 1998 day 150 and \( L = 2.7 \).
Figure 11. Similar to Figure 10 but for 1998 day 151.
Figure 12. Model parameter estimates from 1998, $L = 3.5$. Day and night diffusion coefficients are evaluated at $x = 1$ and at energies of 1, 2, and 4 MeV, as indicated. Day/night is the ratio of diffusion coefficients. For clarity, error bars are shown only in the 2 MeV case. Horizontal bars show the time range (1 day) of each fit.
$L = 3.1$, as in Figure 12.

**Figure 13.** Model parameter estimates from 1998, $L = 3.1$, as in Figure 12.
Figure 14. Model parameter estimates from 1996, $L = 3.9$, as in Figure 12.
Figure 15. $e$-folding lifetimes (data points), derived from the model parameter estimates for three time periods at $L = 3.5$, and the $D_{st}$ geomagnetic index.
Figure 16. Derived e-folding lifetimes at $E = 2$ MeV versus the $D_{st}$ geomagnetic index, for four selected $L$ shells. The lifetimes are for 1-day periods and points are plotted at the mean $D_{st}$ value for the day. The error bars on the lifetimes are derived from the data fits and those on $D_{st}$ indicate the minimum and maximum value during the day.
Figure A1. (left) Calibrated PET energy response curves integrated over solid angle, or geometry factors, for each of the three electron rate counters. (right) Polar plots of the corresponding angular response $R \sin \theta$, where $\theta$ is the angle from the telescope axis, at selected energies. Solid (dashed) curves are used where the response is increasing (decreasing) with energy.
<table>
<thead>
<tr>
<th>L shell</th>
<th>$D_{st}$ ($\times 10^{-9}$ s$^{-1}$)</th>
<th>$\mu$</th>
<th>$\sigma$</th>
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<tbody>
<tr>
<td></td>
<td>Day</td>
<td>Night</td>
<td></td>
</tr>
<tr>
<td>1996 days 143–176</td>
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<td></td>
<td></td>
</tr>
<tr>
<td>3.5</td>
<td>4.3 ± 3.3</td>
<td>1.9 ± 2.7</td>
<td>0.4 ± 0.5</td>
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<tr>
<td>3.9</td>
<td>7.5 ± 7.1</td>
<td>2.1 ± 3.1</td>
<td>-0.9 ± 1.0</td>
</tr>
<tr>
<td>1997 days 180–219</td>
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<td></td>
</tr>
<tr>
<td>3.1</td>
<td>16 ± 14</td>
<td>6.4 ± 5.3</td>
<td>1.7 ± 0.9</td>
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<tr>
<td>3.5</td>
<td>9.1 ± 10</td>
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<tr>
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<td>7.4 ± 8.3</td>
<td>3.0 ± 4.2</td>
<td>-0.6 ± 0.6</td>
</tr>
<tr>
<td>1998 days 150–177</td>
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<td></td>
<td></td>
</tr>
<tr>
<td>2.7</td>
<td>58 ± 120</td>
<td>3.6 ± 3.1</td>
<td>0.3 ± 0.8</td>
</tr>
<tr>
<td>3.1</td>
<td>24 ± 32</td>
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<tr>
<td>3.5</td>
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</tr>
<tr>
<td>3.9</td>
<td>15 ± 16</td>
<td>3.3 ± 5.6</td>
<td>-1.5 ± 0.8</td>
</tr>
</tbody>
</table>

$^a$ Mean and standard deviation of daily values in each period.

$^b$ Evaluated at $E = 2$ MeV and $x = 1$. 

Table 1. Summary of Estimated Model Parameters$^a$