1	Evaluation of the event driven phenology model coupled with the VegET
2	evapotranspiration model through comparisons with reference datasets in a spatially
3	explicit manner.

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## 12 Abstract

A new model coupling scheme with remote sensing data assimilation was developed for 13 estimation of daily actual evapotranspiration (ET). The scheme represents a mix of the VegET, a 14 15 physically based model to estimate ET from a water balance, and an event driven phenology model (EDPM), where the EDPM is an empirically derived crop specific model capable of 16 producing seasonal trajectories of canopy attributes. In this experiment, the scheme was 17 deployed in a spatially explicit manner within the croplands of the Northern Great Plains. The 18 evaluation was carried out using 2007-2009 land surface forcing data from the North American 19 Land Data Assimilation System (NLDAS) and crop maps derived from remotely sensed data of 20 NASA's Moderate Resolution Imaging Spectroradiometer (MODIS). We compared the canopy 21 parameters produced by the phenology model with normalized difference vegetation index 22

23 (NDVI) data derived from the MODIS nadir bi-directional reflectance distribution function (BRDF) adjusted reflectance (NBAR) product. The expectations of the EDPM performance in 24 prognostic mode were met, producing determination coefficient ( $r^2$ ) of 0.8 ±0.15. Model 25 estimates of NDVI yielded root mean square error (RMSE) of 0.1 ±0.035 for the entire study 26 area. Retrospective correction of canopy dynamics with MODIS NDVI brought the errors down 27 to just below 10% of observed data range. The ET estimates produced by the coupled scheme 28 were compared with ones from the MODIS land product suite. The expected  $r^2=0.7 \pm 0.15$  and 29  $RMSE = 11.2 \pm 4$  mm per 8 days were met and even exceeded by the coupling scheme 30 functioning in both prognostic and retrospective modes. Minor setbacks of the EDPM and 31 VegET performance ( $r^2$  about 0.5 and additional 30 % of RMSR) were found on the peripheries 32 of the study area and attributed to the insufficient EDPM training and to spatially varying 33 34 accuracy of crop maps. Overall the experiment provided sufficient evidence of soundness and robustness of the EDPM and VegET coupling scheme, assuring its potential for spatially explicit 35 applications. 36

#### 37 **1. Introduction.**

38 There is growing consensus in the climate science community that the ability to precisely 39 partition energy and matter fluxes on the land surface is key to improving our understanding of 40 mesoscale atmospheric dynamics, ecosystem, responses to climate change, and interactions with human activities and institutions [Pitman, 2003; Ibanez et al., 2010]. Since Manabe [1969], land 41 42 surface modules (LSM) have become increasingly complex modules within most general circulation models (GCMs). The complexities of LSMs have grown substantially as scientists 43 ask questions about the pace and consequences of climate change that require more precise 44 answers. In pursuit of these answers, researchers have been coupling global and regional climate 45

46 models with a spectrum of modules detailing interactions between the land surface and the lowest level of the atmospheric boundary layer. Modules range from a set of simplified surface 47 energy and water balance procedures to more detailed interactive systems like dynamic soil and 48 vegetation modules, complete with light transfer, photosynthesis, and hydrological schemes. 49 Computational resources often limit the level of detail in LSMs especially in regional studies that 50 51 require finer spatial resolution. Also, there is a trade-off between the number of land surface characteristics that can be tracked and the greater spatial detail often needed for regional to local 52 projects [Stensrud, 2007]. 53

Applications of land surface models in regional studies were often focused on just a few 54 55 variables of interest. In many instances, this narrower focus has led to the use of simplified schemes of land surface processes. Numerous local impact studies are turning to empirical 56 methods based on relationships of modeled land surface characteristics to net radiation, 57 58 precipitation, air temperature and other variables [Nagler et al., 2005; Godfrey et al., 2007; Senay et al., 2007; Abramowitz et al., 2008; Jang et al., 2009; Gao et al., 2010]. However, being 59 developed on microclimatological data, empirical models were often unable to predict well when 60 transferred to a different location, even under similar conditions [Li et al., 2009]. Yet, 61 process-based LSMs to address local questions are often hindered by 62 deployment of 63 computational expense and a lack of appropriate ground level data to calibrate and validate at the level of spatial detail required. Also, several studies have expressed concerns about model 64 65 assumptions, process parameterizations, and a limited range of parameters available for tuning 66 [Sabater et al., 2007; Kiniry et al., 2008; Kang et al., 2009; Stancalie et al., 2010], all of which increase doubts about the likelihood of successful deployment of LSMs in regional to local 67 68 studies. Alternatives solutions are needed to provide robust schemes capable of replacing complex LSMs in finer spatial resolution studies. This paper presents a recent development in
land surface modeling combining both physics-based and empirical approaches to take
advantage of the strengths of each approach while yielding results on an appropriately fine scale.

Our research focuses on how potential futures for rainfed agricultural production in the Northern 72 Great Plains may affect regional hydrometeorology. Actual evapotranspiration  $(ET_a)$  was the key 73 flux of interest. We chose to use a simplified simulator of ET<sub>a</sub> called VegET [Senay 2008]. 74 Similar to Godfrey et al. [2007], Kang et al. [2009] and Yuan et al. [2010], Senay's scheme relies 75 on the Penman-Monteith equation [Monteith, 1964] to calculate reference ET  $(ET_0)$  and handles 76 the influences of soil water status and canopy phenology through the two coefficients: K<sub>s</sub> for 77 soil water status and K<sub>cp</sub>, for canopy phenology. The Penman-Monteith method is a physics-78 based one source model of evapotranspiration in cereal crops with fully developed canopies, 79 used extensively by FAO [Allen et al., 1998]. A key innovation of VegET is the modulation of 80 ET<sub>a</sub> by a canopy phenology coefficient using a climatology of the normalized difference 81 vegetation index (NDVI) as observed from spaceborne sensors [Senav, 2008]. 82

83 The original implementation of VegET, however, could not serve our purpose because we were 84 seeking how ET<sub>a</sub> would change in response to both interannual variability and changes in the 85 crop area. Since they were derived from averages of past observations, a static retrospective climatology for K<sub>cp</sub> would not reflect changes in growing conditions [Godfrey et al., 2007; 86 Wegehenkel, 2009] or in the extent of cultivation [Kovalskyy and Henebry, 2011b]. Therefore, 87 88 we replaced a static phenological parameterization with an interactive vegetation growth module. The use of fully functional crop growing modules with energy balance models in point based 89 studies has been common practice [Maruyama and Kuwagata, 2010; Sancalie et al., 2010]. 90 However, our study case required spatially explicit ET<sub>a</sub> estimates that would entail additional 91

92 parameterization, tuning and running time for the models like ALMANAC [Kiniry et al., 2008], CERES [Mearns et al., 1999], CROPWAT [Sancalie et al., 2010] or MODWht [Kang et al., 93 2009]. Moreover, these specific crop models did not have freely available versions capable of 94 working with raster inputs and producing spatially explicit estimates. Conversely, the vegetation 95 growth modules in global LSMs were developed to deliver spatially explicit results [Dickinson et 96 97 al., 1998; Foley et al., 2000; Bondeau et al., 2007; Campo et al., 2009]. Even the most advanced modules do not provide crop specific canopy behavior; instead, they were designed to mimic 98 seasonal patterns of very broad classes of vegetation functional types [Bonan et al., 2003; 99 100 Lawrence and Chase, 2007].

101 Here we have used the Event Driven Phenology Model (EDPM), which was recently developed as a phenology model that can simulate seasonal dynamics of canopy properties (e.g., in terms of 102 a vegetation index) [Kovalskyy and Henebry, 2011a, 2011b]. The model was shown to capture 103 104 fine temporal details of canopy behavior [Kovalskyy and Henebry, 2011a, 2011b] which has been called "crucial" for ET and other surface fluxes [ Dickinson et al. 1998; Foley et al. 2000; 105 106 Pitman, 2003; Gorfrey et al. 2007; Prihodko et al. 2008; Rosero et al. 2009; Rötzer et al. 2010; Zha et al. 2010]. The EDPM uses virtually the same set of forcings as the Penman-Monteith 107 equation to build seasonal trajectories of canopy properties. Essentially, the model provides a 108 computationally inexpensive replacement for a dynamic vegetation model with a phenology sub-109 module. The model also has an option of a simple, fast 1D data assimilation scheme for satellite 110 111 observations which is a great advantage for spatially explicit simulation studies. The EDPM has been coupled with VegET and evaluated against flux tower observations of ET<sub>a</sub> [Kovalskyy and 112 113 *Henebry*, 2011b], where it performed better or comparable to the results obtained by *Nagler et* al. [2005] and Abramowitz et al. [2008]. 114

115 This paper presents an assessment of the performance of the EDPM on its own and also in 116 conjunction with VegET within a spatially explicit application. Our task was to select appropriate sources of scientifically sound data products that would enable pixelwise 117 comparisons of daily canopy states and ET<sub>a</sub> estimates. We were looking to assess two aspects of 118 the coupled model performance. First and foremost, we focused on temporal and spatial 119 behavior of differences between our estimates and reference data. Analyzing the results from the 120 three years (2007-2009) within the study region delimited by croplands of Northern Great Plains, 121 we tried to capture both inter-annual and intra-annual variability of residuals as well as 122 123 correlation between reference data and estimates produced by our model. Second, we looked at 124 the ability of the EDPM to capture the key dates of the three growing seasons. We contrasted phenological dates reported to National Agricultural Statistics Service (NASS) by farmers with 125 126 the dates produced by the EDPMs phenophase control module. Specifically, the Start of Season (SoS), End of Season (EoS), and Length of Season (LoS) became the main criteria for the 127 evaluation. We also tried to incorporate spatial and temporal variability of phenological metrics 128 129 into the evaluation process. This assessment helped us to identify the strong points of the EDPM and to prioritize directions for model improvement. 130

#### 131 **2. Methods and materials.**

### 132 **2.1 Study area.**

The study area includes Nebraska, Iowa, Minnesota, North and South Dakota entirely and parts of Illinois and Indiana. Together these states have more than half of the nation's maize and soybean crops and comprise the major part of the US corn and soybean belts. There strong gradients of ET across the region. The northern tier has only 600 mm ET annually; whereas at the southern end, the annual ET can reach 1000 mm, but only 400 mm at the western extreme. Maize and soybean are the most prevalent crops across the region. Farmers use different genetic varieties of these crops to match the growing conditions of their farms [*Ransom et al.*, 2004]. The green-up of the area starts in early May on the southeast end but for the northwestern part of the region it can happen as late as mid-June if spring comes late. The length of the growing cycle also varies greatly; it can last almost five months in the south and barely more than three months in the north.



Figure 1. The study area (dark gray) depicted as at least 50% corn/soybean crop cover
during 2007-2009.

## 147 **2.2** Coupling the VegET and the EDPM in a spatially explicit manner.

The idea motivating the development of VegET was the use the time series of remotely sensed vegetation indices to drive the canopy factor that modifies ET<sub>0</sub> as calculated by the Penman-Monteith model. The original design of VegET [*Senay*, 2008] used very simple empirical transformations from the normalized difference vegetation index (NDVI) to phenology driven coefficients based on thresholds and the observed variability in NDVI climatologies derived from long-term AVHRR observations. However, *Kovalskyy and Henebry* [2011b] demonstrated the use of an interactive event driven phenology model [*Kovalskyy and Henebry*, 2011a] to replace the climatologies in the VegET for point-based estimation of daily  $ET_a$ . The coupling scheme was shown to account for contemporaneous fluctuations in the canopy component of evapotranspiration

158 The experiment described here evaluates the performance of the EDPM and the VegET after the coupled models were extended for deployment in a spatially explicit manner. In addition to 159 simulating the temporal dynamics of maize and soybean over the three year period (2007-2009), 160 the EDPM was also providing seasonal canopy trajectories for a third vegetation type: grassland. 161 Using weather forcing from the North American Land Data Assimilation System (NLDAS), the 162 163 EDPM transformed the data in to events (rain, heat stress, frost, insufficient insolation, adequate insolation, and growing degree-days) and further produced daily values of Tower NDVI 164 [Huemmrich et al., 1999], The model simultaneously estimated phenological transition dates in 165 166 the three growing seasons for each vegetation type (Fig. 2). The TNDVI trajectories were mixed linearly based on the proportion of their cover within areal units (0.05 degree pixels) and later 167 transformed into phenology coefficients as described in Kovalskyy and Henebry [2011b]. The 168 169 percentages of cover for each crop and grassland were taken from MODIS based crop maps products [Chen et al., 2007] and aggregated into standard 0.05 degree (~5km) pixels that form 170 the spatial unit of analysis for this investigation. 171





Figure 2. Data processing scheme for the experiment. Rounded boxes are modeling and
data preparation procedures; stacks are image time series (ITS) of data; squared boxes are
maps of results.

Using the workflow shown in Figure 2 the coupled scheme was tested in the prognostic mode 177 (running the forcing only) and the diagnostic mode (involving data assimilation scheme with 178 MODIS NDVI observations). The use of data assimilation techniques is becoming increasingly 179 popular in evapotranspiration studies [Meng et al., 2009; Anderson et al., 2010; Miralles et al., 180 2010; Godfrey and Stensrud, 2010]. Most of these projects relied on remotely sensed data to 181 improve their estimates of ET addressing the spatial variability of land surface. While bringing 182 improvements in performance, these techniques have been criticized as being temporally 183 constrained and scene dependent [Li et al., 2010]. Our study took a more general approach to 184 data assimilation using an unambiguous relationship between Tower NDVI and MODIS NDVI 185

established previously [*Kovalskyy et al.*, 2011]. Relying on this relationship *Kovalskyy and Henebry* [2011a] presented a one-dimensional Kalman filter (1DKF) data assimilation scheme in
which the EDPM used MODIS NBAR NDVI to adjust its estimates of canopy states.

189 2.2.1 Differences from point based deployment of the EDPM.

Several features in the EDPM model were added and modified so that the model could represent 190 spatial variability of canopy development during growing season. First of all, the model received 191 the ability to represent pixels with mixed vegetation cover. The Figure 2 shows the linear mixing 192 procedure used to derive values of vegetation index in a pixel with partial covers of grassland 193 and the two crops. Direct linear mixing of NDVI values based on their share of pixel area has 194 195 been criticized in the literature due to its impact on outcomes [Settle and Campbell, 1998; Roy, 196 2000; Busetto et al., 2008]. Since the NDVI is not a linear function of red and near infrared reflectances, linear mixing should be performed on reflectances first so that "unbiased" NDVI 197 values can be obtained later. However, a relatively small impact from direct linear mixing 198 199 (DLM) of NDVI values may not be entirely prohibitive since the reflectances coming from MODIS products have their own errors of the estimate [Roy et al., 2005]. If the differences 200 201 between the DLM NDVI and the NDVI derived from reflectances can be kept within the margin 202 of error propagated into the unbiased NDVI, then one can successfully perform linear mixing 203 directly using NDVI values. The magnitude of differences with true values depends on the 204 number of endmembers used in linear mixing and varies greatly across space and time. We 205 evaluated a real (empirical) example to demonstrate how the direct linear mixing procedure impacts the resulting values of NDVI. 206

First we took a 1000 by 1000 pixel subset from the MODIS Nadir Bi-directional Reflectance
Distribution Function (BRDF) Adjusted Reflectance (NBAR) MOD43A4 [*Schaaf et al.*, 2002]

209 version product covering the central part of the study area. We screened for snow and clouds and 210 based on averages aggregated the 500m reflectance values from MODIS bands 1 and 2 to produce image time series with 1000 m by 1000 m, 2500 m by 2500 m, 2500 m by 5000 m 211 212 (rectangular shape pixel) and 5000 m by 5000 m pixel sizes. For each time series less than 25 km<sup>2</sup> in size, we calculated NDVI [1] that was later mixed directly into 5000 m pixels. 213 Correspondingly, the four sets of results represented 100, 25, 4, and 2 endmember mixing. Out of 214 5000 m reflectance data we calculated "true" NDVI [1] and expected error [2] propagated from 215 reflectances: 0.004 for band 1 and 0.015 for band 2 [Roy et al., 2005]. 216

217 
$$NDVI = \frac{\rho_N - \rho_R}{(\rho_N + \rho_R)^2}$$
[1]

218 
$$\sigma_{NDVI}^{2} = \left(\frac{-2\rho_{N}}{\left(\rho_{N} + \rho_{R}\right)^{2}}\right)^{2} \sigma_{R}^{2} + \left(\frac{2\rho_{R}}{\left(\rho_{N} + \rho_{R}\right)^{2}}\right)^{2} \sigma_{N}^{2}$$
[2]

where  $\rho_N$  and  $\rho_R$  are the reflectance values of near infrared and red bands respectively, and  $\sigma_N^2$ and  $\sigma_R^2$  are the associated variances. In this setup where all resolutions of NDVI data were nested within 5000 m pixels, we expected to see the difference coming just from linear mixing without other effects such as re-projection or resampling that may otherwise contribute to the difference [*Roy*, 2000]. Figure 3 demonstrates the temporal dynamics of average differences between true NDVI and the four DLM NDVI sets coming from linear mixing with different number of endmembers.



Figure 3. Consequences of linear mixing – an observation based example. Difference NDVI = Mixed
NDVI minus "unbiased" NDVI.

229 It is seen clearly from the figure above that the impacts (differences) from linear mixing increase 230 as the number of endmembers grows. The differences become significant when the number of mixing endmembers reaches 4. In our study we used only 3 endmembers (maize, soybeans and 231 grassland) to be mixed into 0.05 by 0.05 degree pixel representing the trajectory of Tower NDVI 232 values. Considering the magnitude of errors from demonstrated direct linear mixing examples it 233 was safe for us to assume that 3 endmember linear mixing did not make a significant impact on 234 the seasonal trajectories of TNDVI produced by the EDPM. We also have to point out here that 235 the minor impacts from linear mixing appeared to be negligible (20 times smaller) compare to 236 237 the estimate errors of the EDPM reported in *Kovalskyy and Henebry* [2011a]. In this context, the 238 impacts from direct linear mixing could hardly make a difference for comparisons undertaken in this experiment. 239

Next, the transformation of mixed TNDVI into phenology driven coefficient K<sub>cp</sub> had to be 240 generalized (unified). In the prior point based studies we found TNDVI and  $K_{cp}$  relationships to 241 be different for crops and grassland [Kovalskyy and Henebry 2011b]. For the later land cover 242 type, the linear model carried substantial noise that we tried to compensate with modeling of 243 residuals through their relationship with vapor pressure deficit. We did not find the same 244 245 relationship in residuals for crops assuming the bias in grassland was due to differences in equipment calibration. Therefore, we used a single linear model with the slope of 1.22 and offset 246 of 0.01 to transform modeled TNDVI into  $K_{cp}$  in this spatially explicit experiment. This 247 248 relationship was derived on observations of TNDVI and K<sub>cp</sub> on crops and proved its efficacy in Kovalskyy and Henebry [2011b]. 249

Finally, the spatially explicit application of the event driven phenology model required some 250 amendments in the functioning of the phenological phase control module described in Kovalskyy 251 252 and Henebry [2011a]. Trained on specific locations and tested on locations with similar climatic conditions, the EDPM required a supplementary mechanism to match the variability of the 253 growing season dates within a much wider range of conditions than in the initial testing studies. 254 Latitudinal gradients for the emergence of vegetated cover which marks the start of the season 255 were applied to the two controlling variables: thermal time and elapsed days since January 1. 256 257 For the elapsed days we applied a 4 days per degree northward gradient suggested by *Hopkins* [1918]. The thermal time triggers for the onset of greening was also modified with the latitude 258 259 using slopes and intercepts tuned for three vegetation types following the phenological transfer 260 functions described in *Henebry* [2010]. To deal with the southward increase in the duration of the growing seasons, we adjusted the dynamic triggering for transitions between phenological 261 262 phases. The adjustment made the transition probability threshold vary inversely with the latitude.

This helped to postpone transitions between phenological phases for locations to the South oftraining sites, while accelerating the transitions to the North.

#### 265 **2.3 Data sources and preparations for the experiment**

The experiments conducted within this investigation had to use various data sources to reach 266 their goals: (1) running the EDPM plus VegET coupling scheme required weather forcing data; 267 (2) percent crop cover data were necessary for the EDPM to produce seasonal canopy trajectories 268 of pixels with mixed vegetation cover; (3) NDVI observations were needed to verify the 269 EDPM's prognosis of seasonal canopy trajectories and later to produce retrospective outcomes; 270 (4) observations of actual ET were needed to evaluate the quality of estimates produced by the 271 EDPM plus VegET coupling scheme; and (5) crop progress reports were crucial for assessment 272 273 of the EDPM module responsible for estimating dates of phenological transitions.

The meteorological forcings for the EDPM and the VegET were supplied by the North American 274 275 Land Data Assimilation System (NLDAS) in native GRIB1 format (1 hour temporal and 0.125 276 degree spatial resolutions). The choice of NLDAS [Mitchell et al., 2004] was based in part on the 277 fact that these forcings were validated on the southern Great Plains adjacent to our study region 278 [Luo et al., 2003]. The original time series of weather data were aggregated into daily image 279 time series and resampled into 0.05 degree grid using nearest neighbor procedure to match the 280 MODIS Climate Modeling Grid (CMG) projection. The last transformation preserved most of the original data and allowed for the fusion of MODIS NDVI data with the EDPM produced 281 seasonal canopy trajectories and later calculation of the ET<sub>a</sub> at 0.05 degree resolution. The list of 282 283 forcing variables included: 2 meter air temperature [K]; 2 meter specific humidity [kg/kg]; 284 surface pressure [Pa]; U wind component [m/s]; V wind component [m/s]; downward shortwave radiation  $[W/m^2]$ ; downward longwave radiation  $[W/m^2]$ ; total precipitation  $[kg/m^2]$ . The forcing 285

dataset and LSM simulations of NASA's Mosaic model from NLDAS Phase 2 were obtained
from the NASA Goddard Earth Sciences Data and Information Services Center (GES DISC) at: *http://disc.sci.gsfc.nasa.gov/hydrology/data-holdings.*

289 Land cover data came in the form of MODIS based crop maps for 2007 through 2009. The 0.5 290 km resolution maps were provided directly by members of the product development team [Chang et al, 2007]. The procedures for deriving the percent covers of maize and soybeans were 291 based on decision tree techniques applied on level 2 MODIS reflectance in seven bands covering 292 visible and infrared portions of the electromagnetic spectrum. Additional metrics capturing the 293 temporal development of land surface properties relevant to the vegetation were also fused into 294 295 the procedure. The 2007 paper reported drawbacks in using the universal sampling approach in the development of their decision tree model which were related mostly to the differences in 296 cropping timing and radiometric properties of underlying soils in different areas of CONUS. As 297 298 an alternative, they proposed a single state based modeling of percent crop type cover which significantly improved the performance of the procedure in independent tests. 299

300 Here we were able to use the latest versions of crop cover maps derived from state based 301 decision tree models of Nebraska, Iowa, Minnesota, North and South Dakota, where the most 302 variability was captured in the least amount of training. Some adjacent areas of other states were 303 combined in the final areal compositing procedure. Within delineated study area where each 304 pixel had at least 50% crop cover, the proportion of grassland was assumed to be the remainder 305 of a pixel cover. This assumption was based on the NLDAS land cover scheme that considered grassland as the second most abundant land cover within our region [Luo et al., 2003]. We also 306 masked out all non-grassland or non-cropland land covers from our study area based on the 307

308 MODIS land cover product (MCD12C1, IGBP classification type available at 309 ftp://e4ftl01.cr.usgs.gov/MOTA/MCD12C1.005).

Verifying the estimated canopy states and actual evapotranspiration on a spatially explicit basis 310 posed some difficulties only because there are so few observational datasets available for such 311 analyses. One such comparison was the TNDVI time series generated by the EDPM in 312 prognostic mode with the MODIS NDVI time series. The reference NDVI image time series 313 NBAR data (MCD43C4 314 were produced from MODIS version) available at ftp://e4ftl01.cr.usgs.gov/MOTA/MCD43C4.005. First, bands 1 and 2 of NBAR data in CMG 315 projection were extracted and screened for insufficient quality records using QA bits. Then, the 316 317 NDVI [1] was computed out of screened red and near infrared reflectance data and organized into image time series. 318

Potentially, daily ET<sub>a</sub> estimates from the EDPM plus VegET scheme had several sources of 319 320 reference data since at the time of our study two ET monitoring products were on their way to 321 public release [Mu et al., 2007; Anderson et al., 2010]. However, only the MODIS 322 evapotranspiration product (MOD16) data had become publically accessible [Mu et al., 2009]. 323 Therefore the MOD16 product had become our first choice for reference when assessing the 324 quality of results from the coupled EDPM+VegET scheme. This product presents estimates of 8 325 day sums of actual and reference ET modeled from weather forcings and remotely sensed 326 properties of the land surface [Mu et al., 2007]. Standard HDF files were obtained from 327 ftp.ntsg.umt.edu/pub/MODIS/Mirror/MOD16/MOD16A2.105\_MERRAGMAO/. The original actual ET layers of MOD16A2 version of the product with 1 km resolution were spatially 328 aggregated and then resampled into 0.05 degree grid again to match the MODIS CMG projection 329 330 adopted as the basis for this experiment.

331 The MOD16 product has been closely approaching the *in situ* measured ET<sub>a</sub> with each improvement to its procedures [Mu et al., 2009, 2011], yet it is still a product with varying 332 degree of spatial and temporal uncertainty. Therefore, we retained an alternative set of ET<sub>a</sub> 333 estimates with which to compare our results. We selected the outcomes of NASA's Mosaic LSM 334 [Koster and Suarez, 1994, 1996] from NLDAS as an alternative reference point for comparison 335 based on the validation studies of Mosaic LSM [Koster and Suarez, 2003; Koster et al., 2004]. 336 To match the formatting of the first reference product (MOD16), the daily ET<sub>a</sub> estimates from the 337 coupled EDPM + VegET scheme and from Mosaic LSM were each temporally aggregated into 8 338 339 day ET<sub>a</sub> totals.

340 The accuracy in estimating phenological dates has always been a subject of point based validation studies [Menzel et al., 2006; Schwartz et al., 2006; Richardson et al., 2009; Zhang et 341 al., 2009; White et al., 2009; Dufour and Morin, 2010]. In the search of reference data we 342 343 examined the National Agricultural Statistics Service (NASS) weekly Crop Progress reports on the percentages of crops achieving crop specific phenophases. From this source we could only 344 obtain information about growing season progress on the state level since the county level 345 reports were inconsistent. Therefore, we reorganized the pixel based EDPM reports into the daily 346 state level growing season progression time series to see the parallels between reported and 347 348 observed dates. These dates were compared with two available years (2008, 2009) of state level crop progress reports obtained for the five states (Nebraska, Iowa, Minnesota, North and South 349 350 Dakota) from the NASS archives: 351 http://www.nass.usda.gov/Data\_and\_Statistics/Quick\_Stats\_1.0/index.asp

Considering the spatial mismatch, temporal precision differences, and the differences in biogeophysical meaning between reported events and dates estimated by the EDPM, we have 354 chosen to rely mostly on the midpoints of distributions in phenological metrics for our comparison. Therefore in the analysis we used midpoint dates (when 50% of crops went through 355 start of season [SoS] or end of season [EoS] ) and their inter-quartile range (IRQ) as a measure of 356 data variability. Based on SoS and EoS dates we also calculated the lengths of seasons (LOS) 357 together with their inter-quartile ranges. The LoS values from the NASS reports were calculated 358 by subtracting the 50% EoS date from 50% EoS date and the IQR 75% EoS date minus 25% EoS 359 date and 25% EoS date minus 75% EoS date. The IQR in the LoS data generated during our 360 experiment were collected directly from the EDPM reported pixel phenology dates. 361

# 362 2.4 Road map for analysis.

Resulting test runs of the EDPM and the coupled scheme with VegET produced several sets of 363 results for the evaluation. First, the image time series of TNDVI estimated by the EDPM in 364 prognostic mode were compared with MODIS NBAR NDVI data. Despite the discrepancy in 365 temporal resolution (8 day for MODIS products and daily for our estimates), the comparison 366 367 could give a good idea of how close our predictions were to the observations. In preparation for 368 such comparison, the EDPM outcomes went through the transformation into MODIS NDVI 369 using the relationship developed in *Kovalskyy and Henebry* [2011a] and confirmed in *Kovalskyy* 370 et al. [2011]. Avoiding the comparison of data beyond the growing season where the EDPM 371 cannot produce TNDVI, we allocated only the results and reference data representing the period from early March (97<sup>th</sup> day of the year) to late October (305<sup>th</sup> day of the year). In addition to that, 372 373 only the dates matching the beginnings of 8 day compositing periods of MODIS products (not the averages over compositing period) were selected for comparison. 374

In diagnostic mode the EDPM used the former reference--MODIS NBAR NDVI data—to correct its outcomes via the built-in data assimilation scheme [*Kovalskyy and Henebry*, 2011a]. Therefore, to assess the model performance in diagnostic mode, we had to rely on error propagation to infer the accuracy of the assimilation-enhanced EDPM estimates of TNDVI.

Prognostic and diagnostic versions of the EDPM outcomes were used to parameterize VegET to produce corresponding  $ET_a$  outcomes. Aggregated into 8 day totals to match the format of first reference data, the  $ET_a$  estimates from both prognostic and diagnostic runs of the scheme were compared with the temporally matching image time series of actual evapotranspiration from MOD16 product validated by *Mu et al.* [2009] and Mosaic LSM validated by *Luo et al.* [2003]. Only the time series of  $ET_a$  from early March to late October were used in the comparisons.

In our assessment we relied generally on the two most common measures of performance: 385 coefficient of determination  $(r^2)$  and root mean square error (RMSE). The first measure showed 386 387 the ability of produced estimates to follow the observed developments of the modeled variable. RMSE showed the overall level of departure of modeled TNDVI and ET<sub>a</sub> from what we assumed 388 to be the reality (reference datasets). Based on the results received in Kovalskyy and Henebry 389 390 [2011a] and Kovalskyy and Henebry [2011b], the expected performance levels for the canopy state estimates (viz., NDVI) simulated by the EDPM were  $r^2=0.8 \pm 0.1$  and RMSE = 0.1 \pm 0.025. 391 For ET<sub>a</sub>, the expected performance levels were  $r^2=0.7 \pm 0.15$  and RMSE = 1.4 \pm 0.5 mm per day, 392 393 but transformed into 8 day values by simple multiplication yields RMSE =  $11.2 \pm 4$  mm per 8 394 days. Additionally, the results were examined for the presence of biases in the residuals. Analyzing differences with reference data, we aimed to assess both temporal and spatial aspects 395 396 of their distributions to receive clear contrasts between sets of our modeling results and reference 397 data.

In its collection, the NASS archive offered emergence and maturity dates for maize as well as emergence and leaf drop dates for soybeans. We assumed these phenological turn points to be closely related to the SoS and EoS dates produced by the EDPM. Comparing phenological data
 we plotted our estimates against references expecting to see connections between plant
 physiological events and their manifestation in the temporal dynamics of optical properties of the
 vegetated surface.

404 **3. Results.** 

### 405 3.1 Contrasting the EDPM derived NDVI against MODIS product.

The maps representing performance measures for each year were produced to show how the ability of the EDPM to represent the canopy conditions varies in space. We also included the maps of average seasonal propagated errors into Figure 4 from results received after the data assimilation (retrospective mode) to contrast those with RMSE obtained during uncorrected (prognostic) estimation.



412 Figure 4. Comparison of the EDPM produced vegetation index against MODIS NDVI

413 within the study area. (a) Coefficient of determination  $(r^2)$ ; (b) Root mean square error; (c)

414 Seasonally averaged propagated daily NDVI error after assimilation of MODIS NDVI

415 **observations.** 

416 The figure above clearly demonstrates that the EDPM was well fit for to the task of following the dynamics of observed MODIS NDVI. Maps in the left column are dominated by dark color 417 representing  $r^2$  of 0.8 and higher. The  $r^2$  values had a tendency to decrease toward the borders of 418 419 the study area and whereas 2007 was the year with the worst performance, 2008 the best. The same conclusion was supported by the RMSE maps in Figure 4. The overall level of error 420 reached 0.18 for 2007, but dropped to just above 0.11 for 2008. The right column of Figure 4 421 shows the uniform distribution of average seasonal propagated errors throughout the study area 422 after EDPM predictions were updated with MODIS NDVI observations. The general level of 423 propagated errors was very close for all three years and constituted slightly less than 0.1. 424



Figure 5. Spatial distributions of residuals (NDVI<sub>EDPM</sub> – NDVI<sub>MODIS</sub>): (*a*) seasonal means of 426 427 residuals; (b) standard deviations of residuals.

425

428 Figure 5 above reveals that the EDPM was mostly underestimating the value of NDVI. Again the picture changed for different years and the character of bias reversed towards the peripheral areas 429 430 of the study region. The year of 2007 came out as the most biased having the mean of residuals -431 0.2 to -0.3 spread along the western Iowa and Minnesota borders. For 2008 and 2009, most of the seasonally averaged differences between observed and modeled NDVI varied between -0.2 432 and 0.1. The variability of the residuals grew from the center towards the borders for each year. 433

However, similar absolute values of RMSE (Fig. 4b) and mean residuals (Fig.5a) point that the
bias was rather uniform in time for most of the study area.

A closer look into intra-annual dynamics of residuals (Fig.6) reveals similarities in developments
seen in both the mean difference with observations and the standard deviation of residuals within
the three growing seasons.



Figure 6. Temporal dynamics of residuals (NDVI<sub>EDPM</sub> - NDVI<sub>MODIS</sub>) during the 2007-2009
growing seasons. Light grey squares represent season of 2007; darker grey diamonds are
2008; and black triangles are 2009.

The trajectories in Figure 6 represent temporal dynamics of residuals averaged over the entire study region (18.7k pixels). It is seen clearly that the biases from different years went through similar seasonal patterns. The graphs show that the EDPM in prognostic mode was starting up seasons with minor underestimation and kept it at this level till the growing season started for

maize and soybeans. The mean of residuals was dropping at every phenological transition point 447 which constitutes a source of performance problems in the EDPM [Kovalskyy and Henebry, 448 2011a, 2011b]. After the change of the phenological phase, the differences with observations 449 came back to the initial level. This pattern means that corrections of the model outcomes during 450 phase change were needed to decrease the bias and make the bias more stable. Overall, the 451 analysis of the EDPM performance suggests that although the errors from EDPM were higher, 452 they were still within the expected range based on prior performance. However, the results from 453 the EDPM can be found reasonably accurate for prognosis or retrospective temporal gap filling 454 455 in observations, considering the fact that the NDVI from EDPM carries the uncertainty from 456 transformation to MODIS NDVI (standard error of the slope of 0.11[Kovalskyy and Henebry, 2011a]), and the uncertainties of the crop maps proliferated through the mixing process. 457

#### 458 3.2 Contrasting $ET_a$ estimates from the EDPM+VegET scheme against MOD16 product.

Before comparing the results from the EDPM+VegET with references, it is important to note that 459 460 the gaps between the two reference datasets were substantial. Plot (a) in Figure 7 clearly shows 461 that compared with MOD16 product, Mosaic ET<sub>a</sub> first overestimated and then brought bias close 462 to 0 in the middle of the growing season, but later it returned to overestimation. The two versions 463 of the EDPM+VegET estimates representing ET<sub>a</sub> derived with and without assimilation via 1DKF scheme also had their differences shown in plot (b) of the figure 7. Following the 464 previously noted pattern of underestimation of canopy properties by the EDPM working in 465 466 prognostic mode, the prognosis of ET<sub>a</sub> values was lower than ET<sub>a</sub> produced in diagnostic mode (with 1DKF). The variability of residuals in Figure 7b exhibited similar temporal behavior to the 467 one found in the bottom plot of Figure 6. 468









# 473 <sub>EDPM with 1DKF +VegET</sub> - ET<sub>a MOD16</sub>. Light grey squares represent season of 2007; darker grey 474 diamonds are2008; and black triangles are 2009.

Retaining the main features from plots a and b, the remaining graphics of Figure 7 show the 475 temporal dynamics of differences between two reference datasets and the two sets of 8 day  $ET_{a}$ 476 estimates from the EDPM+VegET scheme. In prognostic mode the EDPM+VegET results were 477 starting growing seasons with underestimation of 15 mm per 8 days compare to ET<sub>a</sub> produced by 478 Mosaic. In the midseason the difference came close to zero, but later a smaller (~10 mm per 8 479 days) underestimation prevailed again (Fig. 7c). Meanwhile compared to the ET<sub>a</sub> from MOD16, 480 the prognosis from the EDPM+VegET showed close to 0 difference for most of the season with 481 482 slight overestimation in early June (up to 7 mm per 8 days) and underestimation of the same magnitude in late August (Fig. 7d). The variability of residuals for prognostic ET<sub>a</sub> estimates 483 remained high and had a clear temporal pattern driven by phenology. 484

The estimates of ET<sub>a</sub> obtained with the EDPM+VegET working in diagnostic mode (with 1DKF) 485 486 exhibited similar behavior of residuals when compared to reference datasets. Differences with 487 Mosaic were negative at the beginnings of growing seasons (Fig. 7e), but in the mid-season the 488 curves drifted toward slight (up 5 mm per 8 days) overestimation which later changed back to the 489 underestimation of 15 mm per 8 days again (Fig. 7e). Compared with MOD16 the 490 EDPM+VegET diagnostic estimates produced residuals that signal slight overestimation early in 491 the growing season. Later, however, the residuals came close to 0 and remained there till the end 492 of growing season indicating good match (Fig. 7f). The variability of residuals for retrospective/ diagnostic ET<sub>a</sub> estimates from EDPM+VegET dropped quite dramatically in both comparisons 493 494 (Fig. 7e,f) showing the relative efficacy of data assimilation for this method of  $ET_a$  estimation.

495 Overall, the EDPM+VegET scheme showed closer temporal resemblance with MOD16 product
496 and therefore further we present figures representing the spatial particularities of the coupled
497 model performance compared to the MODIS product. (Analogous figures showing the
498 comparison with Mosaic can be found in Appendix A.)



499



Figure 8. Comparison of MOD16 ET<sub>a</sub> with the ET<sub>a</sub> produced by EDPM+VegET working
in (A) prognostic mode and (B) diagnostic mode involving 1DKF assimilation. (*i*)

503 Coefficient of determination  $(r^2)$ ; *(ii)* Root mean square error (mm per 8 days).

Both parts of figure 8 show that EDPM+VegET scheme was able to follow the dynamics of  $ET_a$ in the reference dataset and produced high values of determination coefficient exceeding the expectations set in previous section. Average coefficient of determination was above 0.8 level for the scheme working in both prognostic and diagnostic modes. In 2008, however, the average value of r<sup>2</sup> dropped to the expected 0.7 level for both versions of derived  $ET_a$  (Fig. 8A and B). The distribution of r<sup>2</sup> values within the study area was more even in the results from the coupled

scheme working in diagnostic mode involving 1DKF assimilation with MODIS NDVI data (Fig. 510 8A). In both modes the EDPM+VegET scheme showed lower  $r^2$  in the western peripheral 511 regions where the accuracy of crop cover maps was lower. Correspondingly, the RMSE values in 512 513 those regions were higher especially in the results of the scheme working in prognostic mode. In the results from diagnostic mode RMSE had more uniform distribution and constituted around 6 514 mm per 8 days on average which is half of what was expected. The average RMSE for 515 EDPM+VegET outcomes derived in prognostic mode was about 8 mm per 8 days. Transformed 516 into corresponding units, this performance would be comparable to Nagler et al. [2005] or 517 Abramowitz et al. [2008], if the ET<sub>a</sub> data from MOD16 product approximated the reality with the 518 accuracy of flux tower instruments [Mu et al., 2009]. A point based flux tower validation study 519 has shown that the scheme can approximate daily ET<sub>a</sub> in crops with similar accuracy [Kovalskyy 520 521 and Henebry, 2011b].





Figure 9. Spatial distributions of residuals (A) ET<sub>a EDPM+VegET</sub> – ET<sub>a MOD16</sub> (B) ET<sub>a EDPM with</sub>
1<sub>DKF+VegET</sub> – ET<sub>a MOD16</sub>. (*i*) annual mean of residuals (mm per 8 days); (*ii*) standard
deviation of residuals (mm per 8 days).

The contrast between the two sets of  $ET_a$  estimates from the EDPM+VegET scheme can be easily depicted from the Figure 9 (A and B). In the left column (*i*) of panel A of Figure 9, the prognoses of  $ET_a$  had mostly negative bias changing to overestimation in the peripheral areas of the study region (both east and west). The magnitudes of the mean residuals deviated not too far from 0 giving a peak of up to 12 mm per 8 days in 2007 in the central part of the study region. Left column (*ii*) of Figure 9A shows uneven distribution of variability in residuals revealing clusters of instability in performance coming from EDPM+VegET scheme working in prognostic mode. Panel B of the Figure 9 shows that performance of the EDPM+VegET scheme was more stable during the work in diagnostic mode. The bias in the left column of the Figure 9A was mostly positive fluctuating no more than 9 mm per 8 days. There was less contrast between years and also less difference between various parts of the study region. Smaller and more homogenously distributed standard deviations of residuals (Fig. 9B column *ii*) also indicated a greater stability in performance compare to prognostic mode (Fig. 9A column *ii*).

Contrasted with the ET<sub>a</sub> estimates from Mosaic (Appendix A) the results from the 540 EDPM+VegET scheme were less correlated and had greater spatial variability in RMSE and 541 542 residuals. Figures in Appendix A clearly demonstrate the problem in the central part of the study area (especially during 2007 growing season) that came from numerous differences in 543 approaches to the ET<sub>a</sub> modeling and the associated assumptions made about the parameter 544 545 datasets e.g. land cover types, soil types, LAI, etc [Koster and Suarez, 1996; Mitchell et al., 2004]. Nevertheless, the expected performance of  $r^2=0.7 \pm 0.15$  and RMSE = 11.2 \pm 4 mm per 8 546 days were achieved by the coupled models working only in retrospective/diagnostic mode using 547 MODIS observations for correction of simulated TNDVI trajectories. 548

#### 549 3.3 Comparison of growing season parameters.

The need to evaluate the performance of the phenological control module in the EDPM was well motivated by the patterns in residuals seen in Figures 6 and 7. Therefore, we highlight the contrasts between the EDPM estimated and in situ start of season [SoS] and end of season [EoS] dates reported to NASS.



#### 554

Phenological Dates reported by NASS (DOY)

# Figure 10. Contrasting start and end dates of the growing season for the two crops and two years.

Figure 10 shows fairly good agreement between observed and estimated parameters of the two 557 growing seasons. It also reveals the persisting delays in SoS for maize crops within all five 558 states. Nevertheless, the 2 weeks delays in SoS prognoses were comparable with errors 559 encountered in retrospective analyses by Fisher et al. [2006] Zhang et al. [2009] and Kovalskyy 560 et al. [2011]. Meanwhile, the estimates of both SoS and EoS for soybeans were even more 561 precise and consistent. Figure 10, however, does not show the variability of the start and end of 562 season dates where dramatic differences arise between NASS reports and the EDPM. To 563 564 condense the graphical information, we brought the variability measure, the interquartile range

565 (IQR) into Figure 11, which also shows the scatterplots in length of season (LoS). Similar
566 patterns occur in the variability of SoS and EoS (data not shown).



567

# Figure 11. Contrasts between the EDPM estimates and NASS reports in the length of season and its variability for the two crops and two years.

The most apparent feature of Figure 11 is the error bars showing the inter-quartile range of the length of the season. The contrast between NASS and the EDPM dates went to the edge of the anticipated differences due to disparities between compared datasets. We expected the variability in LoS to be driven by gradients in some climatic factors such as rain, temperatures, duration of daylight etc. What we found in NASS reports was that the states with more variability in seasonal precipitation (viz., Nebraska and the Dakotas) had more variability in 576 phenological timing. The EDPM did not have the precipitation in the list of phenological controls [Kovalskyy and Henebry, 2011a] and, therefore, the vast difference between observed 577 and estimated IQRs in LoS came as a result of limitations in number of factors considered as 578 drivers of phenological timing. Moreover, the EDPM could not take into account the progress of 579 agricultural work in spring as well as other anthropogenic factors affecting the development of 580 crops. Nevertheless, the central tendencies were captured quite well for soybeans. The SoS 581 delays in maize became the reason for underestimation of LoS for this crop. Yet, with all the 582 shortcomings, the EDPM estimates of phenological dates for all crops and years managed to stay 583 584 within the range of state reports from NASS.

#### 585 **4. Discussion.**

586 Planned as a validation study this experiment took the form of a comparison between products while still providing insight on the performance of the EDPM +VegET scheme. In this context 587 588 the discrepancies between estimates found in this study have to be considered just as relative 589 indicators of better or worse performance. Lacking the actual spatially explicit observations, we 590 managed to obtain the reference points for the future application studies where the results will 591 receive interpretation. It is clear now that the outcomes of this experiment helped reaching the 592 goal of this investigation, yet they raised a number of other issues that need to be clarified. In 593 each of the three sets of comparisons we presented spatial and temporal dynamics of error 594 measures but we did not talk in details about the structure of uncertainties or about the reasons 595 behind the observed patterns. Many of these issues are interconnected, and therefore we kept them for this section where the linkages can be explained. Every issue here is discussed in terms 596 of its impact on the abilities of the EDPM alone and the EDPM plus VegET scheme to meet 597 598 nominal performance expectations. We also present ideas about how these impacts can be

mediated at this point and draw perspectives on possible corrections of the problems in futureversions of the event driven phenology model.

Comparison between the MODIS NDVI and the vegetation index produced by the EDPM had 601 both temporal and spatial issues in performance. High  $r^2$  was definitely a plus to the EDPM, but 602 603 the RMSE and bias of 2007 in prognostic mode pushed the performance to the edge of what was expected of the model. Introduction of noise from the NDVI-TNDVI relationship could not be 604 the reason for this error jump since such noise should have been present constantly and not just 605 during late season drought on just about one-fifth of the study area. Apparently, the reaction of 606 the EDPM to this development was too strong (residuals dropped to -0.25), most likely due to 607 608 inability to account for irrigation. An appropriate solution for the 2007 error spike problem would be extra training of the EDPM on irrigated flux tower sites during the drought years. 609 During other years, the bias appeared to be quite consistent throughout the area and could be 610 611 arithmetically removed from the results. Possibly, the bias can be corrected by obtaining better estimates of background vegetation-free TNDVI values for growing season initiation as 612 suggested by Zhang at al., [2003]. This correction would, most probably, draw the overall 613 RMSE close to 0.1 level. This performance mark was also achieved through the data 614 assimilation. 615

Patterns in temporal dynamics of residuals constitute a problem that cannot be corrected with a simple transformation. It requires collecting new data for parameterization of phenophase control module in the EDPM. Inclusion of precipitation as a control variable for phase transitions should help to address the issue of temporal variability in PTPs within states in addition to increasing the overall accuracy of the phenophase control procedures. With the current level of accuracy, we should refrain from interpreting the results based on uncorrected (prognostic) daily NDVI 622 data in places where the variability of residuals goes beyond the level of two seasonal standard deviations. This warning, however, would be less applicable for time averaged (weekly or 623 monthly) or composited prognoses. Meanwhile, the NDVI outcomes received from the EDPM's 624 data assimilation scheme carried significantly smaller traces of phase control errors. Therefore 625 further analysis can be conducted on the retrospective 1DKF corrected daily VI records and 626 627 interpretations would be valid throughout the study region. The issues with temporal stability in performance also came out in the ET<sub>a</sub> estimates produced by the EDPM+VegET scheme. 628 Exceeding the expectation in  $r^2$  and RMSE in comparison with MOD16 product, the results from 629 prognostic mode exhibited a small bump and a dip of similar magnitude in temporal dynamics of 630 the residuals. These fluctuations appeared exactly in the times of phenological transitions from 631 green-up to reproductive phase and then from reproductive phase to senescence respectively. 632 633 Present in results from all three years, the features indicated a systematic problem in phenological control module of the EDPM that, if removed, could further increase the 634 performance of the coupling scheme. In retrospective mode the results still had the issue of the 635 636 early season overestimation. This indicates that while decreasing the level of variability in residuals the assimilation of MODIS NDVI could not completely suppress all the setbacks for 637 the EDPM+VegET scheme. Improvements in the functioning and parameterization of 638 phenological phase control module requires further training on long term flux tower records that 639 will be undertaken in the future. However, all observed magnitudes of the deviations in temporal 640 pattern would not pose a significant obstacle for the use of these results in further analyses. 641

From the comparison of the EDPM+VegET scheme outcomes with  $ET_a$  estimates from Mosaic LSM, we received the diverse spatial dynamics in  $r^2$  and RMSE complemented with clear seasonal patterns in temporal dynamics of residuals. These discrepancies persisted even after the 645 assimilation of MODIS data into the EDPM and VegET results. In fact, the pattern became even more pronounced since the variability in residuals dropped. It is most likely that a better 646 sensitivity of the EDPM to ongoing weather conditions contributed to the temporal dynamics of 647 differences between two ET<sub>a</sub> estimates as the energy balance scheme in NASA's Mosaic LSM 648 [Koster and Suarez, 1996] uses static climatological trajectories of leaf area index as a 649 phenology driven factor of canopy resistance. However, other patterns could not be explained 650 entirely by the lack of sensitivity to contemporaneous vegetation development in the Mosaic 651 model. It is also possible that numerous discrepancies came out as consequences of different 652 653 assumptions about land cover on the 0.125-degree NLDAS grid and/or the ET flux partitioning between canopy and underlying soil. 654

A unique feature of this study was the comparison of growing season metrics estimated by the 655 EDPM with ones reported to NASS. In our analysis, we were missing proper geographic and 656 657 temporal precision in the NASS reports for each of the five states. Nevertheless, we tried to preserve the temporal and spatial variability of growing season dates by organizing our SoS and 658 EoS estimates to match the structure of reference data. We also kept in mind the fact that the 659 transition points in NDVI dynamics and the actual phenological event for crops had different 660 physical meanings. A good matching was achieved between reported and estimated state 661 662 averaged SoS and EoS. Their variability, however, became problematic for the EDPM giving the ground to include more controlling variables into the automatic estimation of phenophase 663 664 transition dates.

Despite all the issues listed in this section the overall impression from the comparisons is quite positive for the VegET+EDPM coupling scheme. The scheme managed to keep the departures from references within nominal boundaries. The results matched and even exceeded most of the 668 expected measures of model performance obtained on point based validations [Kovalskyv and Henebry, 2011a, 2011b]. The biggest problem for the TNDVI trajectories estimated by the 669 EDPM was the model's overreaction to late season drought in 2007 that accentuated the usually 670 small underestimation. Meanwhile, the ET<sub>a</sub> estimates followed closely the reference records 671 from MOD16 products. Even in the worst cases, the error measures in ET<sub>a</sub> were also comparable 672 with those of Senay [2008], Mu et al. [2007], and Abramowitz et al. [2008]. Remarkably, this 673 level of performance was achieved during the spatially explicit deployment of the coupled 674 models. Plus, the results from the scheme were complemented with estimates of phenological 675 676 metrics for grassland and crops that matched well the central tendencies of NASS reports. Combined with the ability of the scheme to produce daily estimates of vegetation index and 677 actual evapotranspiration the performance characteristics of the VegET+EDPM coupling scheme 678 679 justified its use in a real life application study.

680 The lessons learned from this experiment will help to analyze and interpret the results of the greater investigation of recent shifts in the phenology and ET regime in the Northern Great 681 Plains. After the undertaken comparisons we can confidently say that consistency of received 682 errors still allows for the trend analysis especially after correcting with MODIS observations. 683 The delays of season starts in maize will be accounted for in the assessment of inter-annual 684 variability of growing season parameters. Also, we intend to scale the variability in phenological 685 dates from the EDPM to match the variability in NASS reports through inclusion of precipitation 686 687 in the phenophase control mechanism. Special attention will be paid to the peripherals of the 688 study region as those are most likely to carry land cover mapping errors. Finally, we will use appropriate testing methods and critical values when relating the shifts in ET<sub>a</sub> regime to crop 689 690 cover change insuring a more conservative interpretation of their correlation.

692 The purpose of the experiment described in this paper was to provide the rationale for the use of the EDPM+VegET coupling scheme in a spatially explicit application. Such rationale was 693 694 attained via assessing the performance of the scheme through comparison of modeled variables 695 with reference data. First, we compared the image time series of vegetation index produced by the phenology model with MODIS NDVI derived from MCD43C4 product. The expectations of 696 model performance in producing seasonal NDVI trajectories were met vielding r<sup>2</sup> of 0.8  $\pm 0.15$ 697 and RMSE of 0.1 ±0.035 for the entire study area. Retrospective correction of canopy dynamics 698 with MODIS NDVI brought the variability in errors closer to the 0.1 level. Estimation of 699 700 growing season metrics by the EDPM matched the NASS reports with reasonable accuracy – up to 2 weeks of difference in key dates. The estimates of actual evapotranspiration produced by the 701 coupled scheme were compared with ET<sub>a</sub> from NASA's Mosaic model from NLDAS and with 702 MOD16 data from MODIS land product suite. In both comparisons, the expected  $r^2$ =0.7 ±0.15 703 and RMSE =  $1.4 \pm 0.5$  mm per day were met by the coupling scheme working in retrospective 704 mode using MODIS observations for correcting seasonal trajectories of canopy development. 705

Minor issues of model performance were encountered during this experiment as well. The 706 707 EDPM produced trajectories of vegetation index biased towards underestimation but the bias was 708 relatively uniform in space and time and therefore removable. Actual ET estimates from the 709 VegET+EDPM were closer to MOD16 product while producing greater differences with Mosaic 710 LSM that also had persisting spatial and temporal patterns in them. While spatial patterns in 711 differences could be attributed to distinct assumptions about land cover in Mosaic LSM [Mitchel 712 et al., 2004], the seasonal profiles of differences between our estimates and reference data 713 exhibited clear patterns driven by phenology. The impacts of these issues on performance of the

EDPM and the VegET models, however, were relatively small and therefore they could not pose an obstacle for the analysis and interpretation of the outcomes. In general, this study provided sufficient assurance that the interpretations of future results derived the the planned spatially explicit application study will be valid and sound, provided that the detected issues are properly addressed in the analysis.

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- 726

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Figure A1. Comparison of ET<sub>a</sub> from Mosaic LSM with the ET<sub>a</sub> produced by EDPM plus
VegET coupling scheme deployed in (A) prognostic mode and (B) diagnostic mode
involving 1DKF assimilation. (*i*) Coefficient of determination (r<sup>2</sup>); (*ii*) Root mean square
error (mm per 8 days).

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911 Figure A2. Spatial distributions of residuals (A) ET<sub>a EDPM+VegET</sub> – ET<sub>a Mosaic</sub> (B)ET<sub>a EDPM with</sub>
912 <sub>1DKF+VegET</sub> – ET<sub>a Mosaic</sub>. (*i*) seasonal means of residuals (mm per 8 days); (*ii*) standard
913 deviations of residuals (mm per 8 days).