

Nitrogen Incorporation in CH₄-N₂ Photochemical Aerosol Produced by Far UV Irradiation

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Running Title: High N-Incorporation in CH₄-N₂ Aerosol

Abstract

25 Nitrile incorporation into Titan aerosol accompanying hydrocarbon chemistry is thought to be driven by extreme UV wavelengths ($\lambda < 120$ nm) or magnetospheric electrons in the outer reaches of the atmosphere. Far UV radiation (120 – 200 nm), which is transmitted down to the stratosphere of Titan, is expected to affect hydrocarbon chemistry only and not initiate the formation of nitrogenated species. We have examined the chemical properties of photochemical aerosol produced at far UV wavelengths using a High-Resolution Time-of-Flight Aerosol Mass
30 Spectrometer (HR-ToF-AMS), which allows for elemental analysis of particle-phase products. Our results show that aerosol formed from CH₄/N₂ photochemistry contains a surprising amount of nitrogen, up to 16% by mass, a result of photolysis in the far UV. The proportion of nitrogenated organics to hydrocarbon species is shown to be correlated with that of N₂ in the irradiated gas. The aerosol mass greatly decreases when N₂ is removed, indicating that N₂ plays a
35 major role in aerosol production. Because direct dissociation of N₂ is highly improbable given the immeasurably low cross-section at the wavelengths studied, the chemical activation of N₂ must occur via another pathway. Any chemical activation of N₂ at wavelengths > 120 nm is presently unaccounted for in atmospheric photochemical models. We suggest that reaction with CH radicals produced from CH₄ photolysis may provide a mechanism for incorporating N into
40 the molecular structure of the aerosol. Further work is needed to understand the chemistry involved, as these processes may have significant implications for prebiotic chemistry on the early Earth and similar planets.

Key words: Titan, photochemical aerosol, CH₄-N₂ photolysis, far UV, nitrogen activation

45 Introduction

It has been assumed that the photolysis of CH_4/N_2 in the far ultraviolet (FUV), 120 – 200 nm, would not lead to any nitrogen chemistry due to the difficulty in breaking the N_2 triple bond. Nitriles have been observed on Titan in the thermosphere (Vuitton et al., 2007) and throughout
50 the stratosphere (Coustenis et al., 2007), but the initial formation of the -CN bond is assumed to take place only in the ionosphere, largely from Extreme UV (EUV) photons or the impact of high energy electrons present in Saturn's magnetosphere (Yung et al., 1984). Recent work by Imanaka and Smith (2010) has shown that photoionization initiated by EUV photons ($\lambda < 120$ nm) can lead to nitrogen fixation via the production of the HCCN radical, with a substantial
55 portion of the nitrogenated products partitioning into the solid aerosol phase. These processes would occur mainly above 1000 km in Titan's atmosphere (Sagan et al., 1992).

The photon flux at Ly- α (121.6 nm) and other wavelengths longward of 120 nm is orders of magnitude greater than at the EUV, and the longer wavelength light has a greater penetration depth into the atmosphere (Wilson and Atreya, 2004). However, because the nitrogen chemistry
60 for Titan is thought to be fully accounted for by N_2 dissociation processes in the upper atmosphere (Vuitton et al., 2007; Lavvas et al., 2011), the possibility of chemical activation of N_2 in the FUV has not been carefully examined. Nitrogen activation in the FUV could have exciting implications for prebiotic chemistry on the early Earth and other planetary bodies. Dodonova (1966) found that HCN could be formed via irradiation of CH_4/N_2 mixtures at
65 wavelengths > 125 nm, but successive attempts to replicate these results using ultraviolet line sources were unsuccessful (see Table 1, Raulin et al., 1982). Given the null results, Raulin et al. (1982) summarized that the only contribution of FUV radiation on a CH_4/N_2 atmosphere was the production of hydrocarbons. To simulate the incorporation of nitrogen into solid haze material photochemically, other workers have used HCN and HC_3N in initial gas mixtures (Clarke and
70 Ferris, 1997; Tran et al., 2008).

In previous work by our group the aerosol product of CH_4/N_2 photochemistry at $\lambda > 115$ nm was examined for chemical composition, physical properties, and mass production rates using a variety of analytical techniques, including a unit mass resolution aerosol mass spectrometer (Trainer et al., 2006). The aerosol mass production rate within the photochemical
75 cell was characterized by varying the concentration of CH_4 in N_2 , and the resulting rate of aerosol production was shown to represent a significant source for the delivery of organic

material to the early Earth environment given predicted CH₄ abundances and photon fluxes for the young planet (Trainer et al., 2006). For the chemical analysis the assumption was made that nitrogen chemistry could be ignored and the aerosol product consisted solely of hydrocarbon material. Here we revisit the CH₄/N₂ experiment using a high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) to more closely examine the chemical composition of the aerosol products. The use of the HR-ToF-AMS allows us to provide new insights into the chemical composition of the photochemical aerosol and the role of atmospheric N₂ in aerosol formation.

85 **Materials and methods**

The photochemical flow reactor used to generate the aerosol has been described extensively in previous work (Trainer et al., 2006; DeWitt et al., 2009; Hasenkopf et al., 2010). In the present work, the gas mixture preparation and aerosol generation procedures were followed precisely as described previously for interfacing with the Aerosol Mass Spectrometer (AMS) (Trainer et al., 2004a; 2004b; Trainer et al., 2006). Briefly, a mass flow controller (Mykrolis, FC2900) was used to flow prepared gas mixtures at a rate of 100 standard cm³ min⁻¹ through a 300 cm³ glass reaction cell and into the inlet of the AMS. A water-cooled deuterium lamp (Hamamatsu, L1835) with MgF₂ windows, emitting from 115 - 400 nm, was inserted directly into the reaction cell. The deuterium lamp is a continuum source, but has a peak near 121 nm, analogous to the solar Ly- α line at 121.6 nm which is a primary driver of CH₄ photolysis in planetary atmospheres. The reaction cell was held at close to room pressure (~600 Torr) and room temperature, as measured by an in-line 1000 Torr Baratron (MKS, 626A) and thermocouples attached to the exterior of the reaction cell.

The work described here builds upon previous studies in which the chemical and physical properties of photochemical haze aerosol produced from CH₄/N₂ mixtures were extensively characterized (Trainer et al., 2006). Inherent in the previous work was an assumption that the N₂ molecules were not photolyzed at the experimental wavelengths, and therefore served only as third-bodies for reactions and did not become chemically incorporated in the aerosol molecular structure. To examine the validity of this assumption, the photochemical study was expanded to include high-resolution chemical analysis exploring the effect of the composition of the background gas. Gas mixtures were prepared in which the CH₄ mixing ratio was held constant at 0.1%, and the N₂ concentration was varied in a background gas of argon (Ar). Gas reagents used

were acquired from Airgas, Inc. (Radnor, PA) with minimum purities of 99.99, 99.999, and 99.998% for CH₄, N₂, and Ar, respectively. The CH₄ mixing ratio is lower than that on Titan, but is based on the past experience with the photochemical cell and optimization of aerosol productivity. Our previous work showed that as the CH₄/N₂ ratio is varied in the gas mixture the chemical composition of the aerosol did not change, but higher production rates were observed at 0.1% than at 2%, and the lower mixing ratio has been used to ensure the best signal to noise ratio in our measurements (Trainer et al., 2006).

The following experiments were performed, described by the fractional concentration of N₂ (with Ar bringing the total pressure to 600 torr): 0, 0.01, 0.1, 0.6, 1, 5, 31, 66, and 100% N₂. In addition, one experiment was performed with CH₄ using helium (He) as the background gas but no N₂ present. Each gas mixture was cycled through a minimum of 3 photochemical experiments lasting 1-2 hours each. An experiment consisted of first establishing a flow of the gas mixture through the reaction cell and into the mass spectrometer with no irradiation, providing a procedural background measurement. The light source was then turned on to initiate photochemical aerosol production, and aerosol product was observed in the mass spectrometer within 5-10 minutes. A steady production rate was typically achieved within 20 minutes. The averaged spectra shown in this manuscript represent this steady state.

The distinguishing element of this study is the use of the HR-ToF-AMS for chemical analysis of the photochemical aerosol (DeCarlo et al., 2006). All previous work by this group on CH₄/N₂ photochemical hazes has been conducted using a unit-resolution quadrupole AMS (Q-AMS). While the data produced from the Q-AMS has been extremely valuable in the assessment of the chemical trends of the aerosol as a function of trace gas composition (e.g., CH₄, CO₂, H₂), the high mass resolution of the HR-ToF-AMS allows for the first definitive evaluation of the nitrogen content in the aerosol product. The sample introduction system, particle sizing, and vaporization and ionization systems of the HR-ToF-AMS are the same as that of the Q-AMS used previously (Canagaratna et al., 2007). Following ionization, the generated ions are extracted into a high-resolution, time-of-flight mass spectrometer (H-TOF Platform, from Tofwerk, in Thun, Switzerland). Ions of different m/z values achieve different velocities when they are extracted into the flight path, and therefore are separated by flight time when they impact on the microchannel plate detector. All data presented here were collected with the HR-ToF-AMS ion optics operating in the “W-mode” (named after the shape of the ion path) of the

140 spectrometer, which increases the flight path and improves ion separation and consequently mass resolution, but with reduced sensitivity (DeCarlo et al., 2006). In this mode the resolving power of the instrument ranges from $m/\Delta m = 3000$ to 4300 over the m/z range of interest (10-200), and the mass accuracy has been estimated at approximately ± 7 ppm. Here, only averaged compositional data derived from the mass spectrum (MS) mode are presented.

145 Custom software was used to analyze the HR-ToF-MS data in Igor Pro 6.1 (Wavemetrics, Lake Oswego, OR). The algorithms applied to the data for performing the m/z calibration, Gaussian fit of ion peaks, and extraction of ion intensities for elemental analysis (EA) have been described in detail in the literature (DeCarlo et al., 2006; Aiken et al., 2007; Aiken et al., 2008), and are included in the analysis software. The total ion current of each ion is determined based on the signal peak area (not peak height) in the spectrum. The custom peak-
150 fitting routine fits each ion peak with a modified Gaussian shape that is a function of the ion time-of-flight, applying an array of user-defined possible ions (exact m/z 's) to identify and quantify the ion signal from overlapping ion peaks at each integer m/z (DeCarlo et al., 2006). Two recent studies report good agreement of the EA results of HR-ToF-AMS data with those from other techniques (Chhabra et al., 2010; Kroll et al., 2011).

155 **Results**

The average aerosol mass spectrum for the 0.1% CH₄/N₂ photochemical product is displayed with unit resolution in Figure 1. The spectrum is fundamentally similar to that obtained using the Q-AMS and presented in the previous publication (Trainer et al., 2006). A comparison of the two spectra yields an $R^2 = 0.90$, which is consistent with the level of
160 reproducibility of AMS spectra across instruments (Dzepina et al., 2007). As discussed in the previous work, the pattern observed indicates the aerosol contains unsaturated hydrocarbons with varying chain lengths. Aromatic indicators (fragments at m/z 77 and 91 for phenyl and benzyl ions, respectively) are also observed.

A surprising finding that is revealed in the high-resolution measurement is a substantial
165 nitrogen component in the organic aerosol structure. Figure 2 shows the region of the spectrum around the nominal m/z 27, which was previously assumed to arise solely from a hydrocarbon (i.e. C₂H₃⁺) fragment ion. The high-resolution spectrum retrieved from the HR-ToF and peak-fitted as discussed above (DeCarlo et al., 2006) shows that more than half of the signal at this integer m/z arises from an ion peak centered around m/z 27.011, which is attributed to the HCN⁺

170 ion (exact m/z 27.01035). The remainder of the signal is due to the peak from the hydrocarbon
fragment $C_2H_3^+$ (exact m/z 27.02293). We calculate that a minimum $m/\Delta m$ value of 2148 at full-
width at half-maximum (FWHM) is needed to adequately separate these peaks, and this criterion
is met by the minimum instrument resolving power of 3000.

The spectrum is similarly fitted at each nominal m/z in the range from 20 – 100 where
175 most (88%) of the aerosol signal appears, and the elemental composition of each ion is combined
with the measured ion currents to determine the bulk elemental composition for the average
aerosol spectrum. This compositional analysis allows for a calculation of the relative mass
concentrations of the major elemental species carbon (C), nitrogen (N), and hydrogen (H). Note
that in the AMS the ion signal intensity is proportional to the mass of the aerosol species that
180 created the signal, see Canagaratna et al. (2007) for details. The results for the 0.1% CH_4/N_2
photochemical aerosol are given in Figure 1, and the aerosol is found to have a composition of
75% C, 16% N, and 9% H by mass. Based on the average AMS uncertainty of 20% reported by
Aiken et al. (2007) for the N/C determination, we arrive at an estimate for the contribution of
nitrogen to the aerosol mass of $16 \pm 3\%$. This composition roughly corresponds to an “average”
185 formula of $C_{11}H_{16}N_2$, a metric that is useful for comparisons to other analyses of laboratory-
produced CH_4/N_2 haze analogs (e.g., Imanaka et al., 2004). Most notably, the atomic N/C ratio
of 0.18 for our aerosol is lower than the values of 0.4-0.8 reported for aerosol produced by
irradiation at lower wavelengths, whose higher photon energies facilitate a greater N_2
dissociation rate (Imanaka and Smith, 2010). N/C ratios for Titan-like hazes formed in various
190 types of discharges vary considerably from 0.1 to 0.8 (Imanaka et al., 2004).

Additional understanding of the chemical nature of the aerosol can be obtained by
examining the series of ion peaks identified in the spectrum, rather than just interpreting the bulk
elemental formula. The mass spectrum given in Fig. 1 displays the skeletal formulas (i.e.
without H atoms) derived from the high-resolution analysis for each of the groups of peaks on
195 the “picket fence” pattern in the mass spectrum. Each peak at an integer m/z value represents
that sum of ion signals from molecular fragments that have the skeletal formulas indicated above
that group of peaks. For example, at m/z 55 there are possible contributions from $C_4H_7^+$, $C_3NH_5^+$,
 $C_2N_2H_3^+$, and CN_3H^+ . The peak at m/z 56 contains contributions from ions with an additional
hydrogen atom, namely $C_4H_8^+$, $C_3NH_6^+$, $C_2N_2H_4^+$, and $CN_3H_2^+$. The hydrocarbon fragments $C_1 -$
200 C_7 are represented in the full mass range shown. Nitrogenated organic species are represented by

ions with $N_1 - N_3$ in this mass range, but most (89% of C_xN_y mass) of the ion signal arises from fragments with one N atom. The bars in the mass spectrum are shaded to indicate the contributions of each element to the mass concentration represented by all ions at that integer m/z . It can be seen that with a few exceptions (aromatic-dominated m/z 77, $C_6H_5^+$ and m/z 91, $C_7H_7^+$) almost all of the nominal m/z 's contain a contribution from N. A closer look at the location of the center of each fragmentation packet reveals that the fragment masses skew towards unsaturated organics such as alkenes and nitriles (m/z 40-41, 54-55, 68-69, 82-83), and away from alkanes and amines (m/z 43-44, 57-58, 71-72, 84-85) (McLafferty and Turecek, 1993). This preference for unsaturated organics is consistent with inferences of unsaturated neutral molecules and ions high in Titan's atmosphere based on INMS observations (Vuitton et al., 2007).

The photochemistry experiments discussed above were repeated by varying the N_2 concentration, as detailed in the *Materials and Methods* section. The results from this study were two-fold; both the proportion of N in the aerosol as well as the overall aerosol signal appear to be influenced by the N_2 concentration. Evidence of the latter effect can be observed in Figure 3 in which the aerosol spectra from experiments using 0.1% CH_4 in different Ar/ N_2 mixtures are compared. The spectra are scaled so that the mass loading of the most intense peak, which is at m/z 27 in the top panel, is 100. This scaling highlights the approximately 100-fold decrease in signal when CH_4 is photolyzed in pure Ar as compared to pure N_2 . An experiment performed with 0.1% CH_4 in He (not shown) had similarly low aerosol yield as compared to the Ar results. The low particle yield in Ar is consistent with previous results reported by Bar-Nun and Podolak (1979), in which no condensed phase was observed after hours of irradiation of the same gas mixture by the same FUV wavelength range. Adamkovics and Boering (2003) saw aerosol production from CH_4 -only irradiation with the same FUV source, but the CH_4 pressure, static chamber, and particle detection methods are dissimilar enough from our experiments that it is not possible to compare their aerosol production rates to ours either with or without N_2 . From Figure 3, it is evident that there is a significant leap in aerosol production as even a small amount of N_2 is included in the reaction cell, demonstrated by the differences between the lower two panels (0% N_2 vs. 0.1% N_2). The difference in the spectra indicates a change in the number and abundance of mass peaks, which implies an increase in organic production. These results suggest not only that N_2 can be involved in the molecular structure of the organic aerosol, but that it must

be involved for significant production of aerosol to occur, at least under our experimental conditions.

235 The exact mechanism for the inclusion of nitrogen in the particles is unclear. At the experimental temperature it is extremely unlikely that N_2 would be adsorbed onto the surface of the aerosols or trapped at sufficient mass quantities ($> 10\%$) to account for the observed nitrogen content of the aerosol. Even if trapped, the N_2 molecules would diffuse quickly out of the nm-sized particles in the vacuum system of the HR-ToF-AMS during the 10 ms flight prior to vaporization (Snijder et al., 1995). To investigate whether the nitrile production could be an artifact due to species formation within the ionization region of the AMS, an additional control
240 experiment was performed in which the CH_4/Ar mixture was photolyzed in the reaction cell and N_2 gas was added to the aerosol flow downstream, before the inlet of the HR-ToF-AMS. In this control the HCN^+ peak at m/z 27.01 was absent, and the average mass spectrum did not vary as the proportion of N_2 introduced downstream was varied. These results indicate that the nitrogen-
245 incorporating chemistry is a result of photolysis within the cell.

Select nominal m/z regions – 26, 27, 29, 53, and 77 – are expanded into high-resolution view in Figure 4. Here, the traces and peak identifications of several N_2/Ar combinations are compared. The integer m/z 26, 27, and 29 were chosen since these are the peaks in the fragmentation packet with the highest relative signal, and serve as tracers for the relationship
250 between nitrile (CN^+ , HCN^+) and hydrocarbon ($C_2H_2^+$, $C_2H_3^+$) production. As the proportion of N_2 is decreased, the relative contribution of the nitrile fragments also decreases. At nominal m/z 29, we also observe an example of a peak that cannot be uniquely identified at our resolution. Here, the signal at m/z 29.003 can be due to two possible ions, marked with asterisks: $N^{15}N^+$ (29.00318) and CHO^+ (29.00274). By comparing the aerosol product spectrum to the spectrum
255 of the gas mixture acquired prior to the lamp ignition, we can estimate that approximately 15% of that peak in the 100% N_2 experiment is due to the gas phase nitrogen isotope signal, and the remainder is likely from the contribution of the CHO^+ fragment. This proportion of $N^{15}N^+$ is smaller for the other two experiments shown in Fig 4. The CHO^+ fragment may be evidence of some oxygen contamination in our system from residual H_2O . Fortunately there are few
260 occurrences of oxygen contamination in our results.

The high-resolution view of nominal m/z 53 in Figure 4 shows the multi-peak results typical at higher m/z , where the ion population contains several ions in the 100% N_2 experiments

but only one dominant ion when Ar is the primary background gas. The $C_xH_yN_z^+$ fragments are clearly present at almost all of the m/z values in the 100% N_2 spectrum except when there is an aromatic ion (m/z 77, 91). At these m/z 's the signal from the aromatic ion is sufficiently greater than surrounding ions that it is difficult to resolve the contribution from any possible N-containing ions that may be present. This difficulty is demonstrated in Figure 4. At all N_2 concentrations the peak at m/z 77.039 is identified as the phenyl ion $C_6H_5^+$. However, as shown in the figure, the N-containing ion $C_5H_3N^+$ might be present.

As can be seen in Figure 4, at m/z 27 there is a definitive change in the ratio between the nitrile and hydrocarbon contributions as the N_2 concentration is decreased. This peak has the highest signal intensity in the observed mass spectra, represents a substantial fraction of the total aerosol mass, and does not have any interference from N_2 gas. The ratio of these two fragments at m/z 27 was therefore chosen as a marker for the degree of nitrogen incorporation into the organic aerosol chemical structure. The ratio is shown plotted as a function of the % N_2 in the gas mixture (balance Ar) in Figure 5. Figure 5 also contains a plot of the atomic N/C ratio as a function of N_2 percentage. There is some variability in these observed ratios from experiment to experiment, but the overall trend with N_2 concentration is clear. These data show that a shift in chemical composition occurs between the all-hydrocarbon products generated in Ar and the nitrogenated aerosol formed in N_2 containing mixtures. Above a N_2 concentration of 10%, the contribution of the nitrile fragment to the nominal m/z 27 dominates that from the hydrocarbon fragment.

Photochemical formation of CN bond

The formation of HCN from the irradiation of CH_4 and N_2 mixtures at wavelengths > 125 nm was first reported by Dodonova (1966), with an approximate yield of 10^{-5} HCN molecules per photon (Chang et al., 1979). This result was derived by using a very specific detection technique sensitive only to gas phase HCN, with corroboration from measurement of trapped gases with a mass spectrometer. Another study (Scattergood et al., 1989) using broadband light from a discharge with a lower transmission edge at 115 nm also reported gaseous HCN production with a yield of approximately 10^{14} molecules per J (equivalent to 10^{-4} HCN molecules per Ly- α photon). Neither of these studies proposed a mechanism for HCN formation

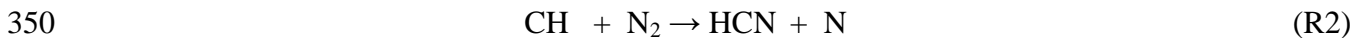
(Scattergood et al., 1989). Using the compositional information retrieved from the HR-ToF-AMS, the total aerosol mass production rates and photon flux reported in our previous work, and an assumption that each nitrogen atom within the aerosol is bonded to a carbon atom (a reasonable assumption given the fragment ions observed), we calculate an approximate yield of 10^{-3} C-N linkages generated per photon within our cell. This value is two orders of magnitude greater than the HCN yield reported by Dodonova (1966), and one greater than Scattergood et al. (1989), which is not surprising given the much broader detection capability of our method. With the HR-ToF-AMS we observe a myriad of particle-phase compounds with -CN groups, which may be produced with a higher yield than HCN gas.

Interestingly, when subsequent attempts in the 1970s and early 80s to repeat the results of Dodonova (1966) by other workers failed to produce nitriles or N-containing organics in the ultraviolet (Ferris and Chen, 1975; Chang et al., 1979; Bossard et al., 1982; Raulin et al., 1982), it was concluded by Bossard et al. (1982) that “UV light is not able to produce nitriles from CH₄-N₂ mixtures” and “UV light is not an efficient source for producing unsaturated carbon chains.” The results presented herein clearly contradict those conclusions. It is difficult to directly compare the results from these previous publications to ours due to variations in experimental conditions (pressure, temperature) and measurement techniques. Table 1, adapted from Raulin et al. (1982), reviews the previous findings alongside the results presented in this work. Notably, the observations summarized in Table 1 show that with the exception of the studies by Dodonova (1966) and Scattergood et al. (1989), the other works were performed using line sources rather than continuum sources such as the one used in our study. While the absorption cross-sections are known in the continuum, there is very little work done on the branching ratio of CH₄ dissociation using continuum radiation because the EUV solar flux is dominated by Ly- α . However, a compilation published by Huebner et al. (1992) shows that across the continuum range the formation rates of CH₄ photolysis products CH₃, CH₂, and CH do vary relative to each other, indicating there is a bias towards or away from particular radical products when a line source is used rather than continuum. Additionally, two of the studies reported in Table 1 were performed using wavelengths above the threshold for CH formation of 137 nm. Imanaka and Smith (2007) explored the formation of gas phase species at EUV and FUV wavelengths and found the production of larger hydrocarbons (C₇-C₈) absent in the FUV, but they were unable to distinguish nitrogen incorporation in the organic species and did not measure the composition of

325 the particle phase. A recent study by Hodyss et al. (2011) has observed –CN formation in water
ice from CH₄/N₂ photolysis at similar wavelengths. Although occurring within different
matrices, this is further evidence that a mechanism to cleave N₂ at λ > 100nm via CH₄ photolysis
needs to be considered. Hodyss et al. (2011) suggest that generation of CH radicals may initiate
the process, and we will explore this possibility in more detail below.

330 Despite early results observing HCN formation, there are no proposed mechanisms for
nitrogen activation at FUV wavelengths put forward in the literature. The N-N triple bond in N₂
is among the strongest of chemical bonds, and direct photolysis of N₂ does not take place unless
the photons have wavelengths shorter than 110 nm, as shown in Figure 6. Although the bond
energy determined thermodynamically is 9.76 eV (compared to 10.2 eV of Ly-α photon), the
335 absorption cross-section of N₂ is negligible in this region of the spectrum (for example, see
Huebner et al., 1992). By far the most extensive chemical modeling of organic chemistry within
a CH₄/N₂ atmosphere addresses the processes on Titan, which includes an array of nitrile
products. In all cases nitrile chemistry is initiated via the dissociation or ionization of N₂ by
energetic electrons within Saturn’s magnetosphere or, more significantly, with solar EUV (Yung
340 et al., 1984; Wilson and Atreya, 2004; Lavvas et al., 2011). Theoretical studies focused on Ly-α
photolysis have only explored hydrocarbon chemistry (Wilson and Atreya, 2000).

It thus appears that there is a mechanism missing from the literature, which is needed to
explain the chemical activation of N₂ within our photochemical cell. One possibility is indirect
N-N cleavage which has been proposed to explain the formation of “prompt” NO in the
345 combustion of CH₄ in air, as discovered by Fenimore (1971). The two-step process begins with
photolysis of CH₄ (e.g., by Ly-α) to produce the CH radical (see, e.g., Yung et al., 1984; Hebrard
et al., 2006; Lodrighito et al., 2009) followed by reaction to form HCN:



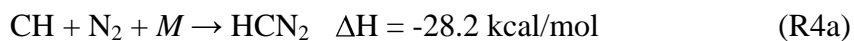
However, reaction R2 is spin-forbidden, as the ground states of CH and N are ²Π and ⁴S,
respectively. Consequently, the reaction has large activation energy. The theoretical rate
coefficient that fits most of the laboratory measurements between 1500 and 4000 K (Miller and
355 Walch, 1997) is $k = 3.68 \times 10^7 \text{ T}^{1.42} \exp(-20723/RT) \text{ cm}^3 \text{ mole}^{-1} \text{ sec}^{-1}$. If we apply this formula at

room temperature as in our reaction cell the rate coefficient is $5 \times 10^{-17} \text{ cm}^3 \text{ s}^{-1}$. This value compares well to those presented in Becker et al. (1996) Table IV, where the authors estimated a rate coefficient for R2 at the lowest temperature (572 K) and lowest pressure (20 Torr) to be $2.9 \times 10^{-20} \text{ cm}^3 \text{ s}^{-1}$. These results led Becker et al. (1996) and other groups to determine that the influence of this bimolecular reaction on the CH + N₂ *k*-values is negligible. In this case reaction of CH with CH₄:



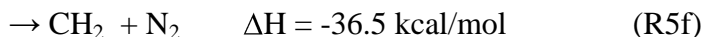
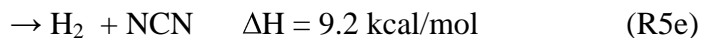
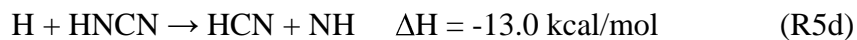
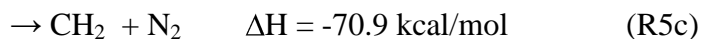
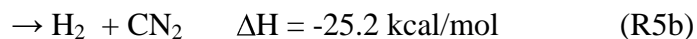
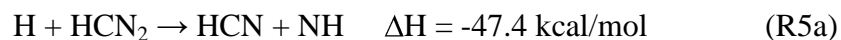
with a rate coefficient around $10^{-10} \text{ cm}^3 \text{ s}^{-1}$ will dominate.

Alternatively, the reaction between CH and N₂ via a termolecular process can lead to the formation of the diazomethyl radical HCN₂ and its isomer the HNCN radical, which is more stable by 34.4 kcal/mol (Tao et al., 1994; Clifford et al., 1998; Moskaleva et al., 2000; Berman et al., 2007):



albeit with unknown branching ratios. The addition reaction is apparently fast, even at low temperatures (Le Picard and Canosa, 1998; Le Picard et al., 1998), suggesting that the reaction has no kinetic barrier. The third order rate constant was found to be $k_0 = (5.7 \pm 0.30) \times 10^{-30} \text{ cm}^6 \text{ s}^{-1}$ at 53 K (Le Picard and Canosa, 1998) with similar values at room temperature (Becker et al., 1996; Brownsword et al., 1996).

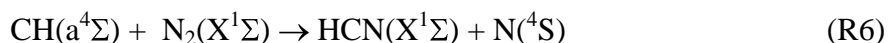
The fate of HCN₂ is most likely reaction with H:



Unfortunately, the branching ratios for the above reactions are completely unknown. Note that pathways (a), (d), and (e) result in the breaking of a N-N bond and the formation of a C-N bond, although (e) is slightly endothermic. Photolysis of HCN₂ may also be possible but with unknown cross-sections or products (Faulhaber et al., 2006). The only atmospheric model that has studied the formation of HCN from CH + N₂ is by Krasnopolsky and Cruikshank (1999). They estimated the HCN yield (Y) relative to total photolysis of CH₄ to be about 1×10⁻⁷ for the atmosphere of Pluto, suggesting this pathway has a small impact in a planetary atmosphere. In the laboratory experiments described herein involving irradiation of a mixture of CH₄ and N₂ by a deuterium lamp, CH radicals are readily produced and the third body concentration is very large. If Y is defined as f₁×f₂×f₃×f₄, where f₁ = branching ratio for producing CH in CH₄ photolysis, f₂ = fraction of CH that reacts with N₂ to form HCN₂ and HNCN, f₃ = fraction of HCN₂ and HNCN that reacts with H, and f₄ = fraction of above reaction that results in the formation of HCN or a CN bond, then the upper limit for Y would be equivalent to f₁, which ranges in the literature from 7 to 51% (Wilson and Atreya, 2000). Yet, given all of the unknown parameters it is difficult to assess whether this reaction scheme is the likely explanation for the formation of -CN bonds and the high nitrogen incorporation observed in the organic aerosol in our experiments.

Likewise, we have explored the possibility of radiative association reactions of CH₃⁺ + N₂ → CH₃N₂⁺ as a mechanism for activation of nitrogen within the photolysis cell (Holtz and Beauchamp, 1971). At Ly-α, CH₃⁺ can be formed from the photoionization of CH₃ radicals. If reaction of CH₃⁺ with N₂ is competitive with reaction with CH₄, this may be a source of activated N₂. Smith and Adams (1978) reported a rate coefficient of 5 × 10⁻²⁹ cm³ s⁻¹ for the formation of CH₃N₂⁺ at conditions relevant to the experimental reaction cell, and this is supported by calculations of Phillips (1990). A rate coefficient for CH₃⁺ + CH₄ is given as 1 × 10⁻⁹ cm³ s⁻¹ in Vuitton et al (2008). Thus in our reaction cell with 0.1% CH₄ in N₂, the reaction with CH₄ will likely dominate. The production of CH₃N₂⁺ may therefore be too low to substantiate the high observed nitrogen content in the laboratory aerosol. Also problematic is that there is not a known pathway to HCN or -CN groups from the CH₃N₂⁺ ion.

The excited state CH(a⁴Σ) is 0.74 eV above the ground state CH(X²Π), and can drive the spin-allowed exoergic reaction:



420 However, only an upper limit of $6 \times 10^{-14} \text{ cm}^3 \text{ s}^{-1}$ has been measured for this reaction (Hou and Bayes, 1993). As the reaction between $\text{CH}(a^4\Sigma)$ and CH_4 is also slow (Hou and Bayes, 1993), it is possible that there is a path for direct formation of HCN by the above reaction in the laboratory and the atmosphere.

425 Surveys of currently known experimental pathways have not yielded a “smoking gun” mechanism to explain the results reported in this work. Clearly further work (experimental and theoretical) is needed to illuminate this problem and provide the missing mechanism. Future studies of these processes at lower pressures would be highly desirable to determine whether the termolecular process proposed here may be the source of the unexpected nitrogen incorporation. Likewise, a systematic exploration of the wavelength dependence of would be enlightening.

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Implications

The formation of HCN or nitrile-rich materials are of great interest for multiple planetary bodies within the Solar System because of the strong spectral features and prebiological potential of these types of molecules (Matthews, 1992). Thus it is important to understand the complete range of formation conditions given the wide array of radiation and compositional constraints of different planetary environments. For the early Earth, there has long been a question as to how Miller-Urey type prebiotic chemistry with N-containing organics may have taken place without a large abundance of NH_3 in the atmosphere (Toupance et al., 1977; Pavlov et al., 2001). FUV light is the presumed preeminent energy source on the early Earth, and these results have significant implications for understanding the chemical nature of a globally available organics source (Trainer et al., 2006; Wolf and Toon, 2010). The results presented here are complementary to recent modeling of enhanced HCN formation on an early Earth with abundant CH_4 , and as suggested by Tian et al. (2011) we have shown that you can get significant $-\text{CN}$ incorporation in the haze particles which may later deposit into oceans for further reaction. Further work regarding the impact of CO_2 and/or H_2 on the nitrogen chemistry needs to be completed (Tian et al., 2005; Haqq-Misra et al., 2008).

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Whether or not FUV production of reactive nitrogen has relevance for Titan is dependent on whether the majority of chemical production of aerosol takes place in the stratosphere, or if the primary energy input that drives organic chemistry and haze formation occurs higher in the thermosphere and ionosphere due to photolysis at EUV wavelengths as some have suggested (Lavvas et al., 2009; Lavvas et al., 2011). It is also dependent on the chemical mechanism responsible for the N₂ activation. Figure 7 demonstrates why the reaction pathways discussed in the previous section (R4 and R5) may not be relevant on Titan. On Titan, where CH₄ is abundant, reactions of CH + CH₄ dominate in haze-forming regions. However, the reactions with N₂ could be relevant for the early Earth. The formation of nitriles from CH via reaction with N₂ can only take place in regions where $k_c[\text{CH}_4] < k_n[\text{N}_2][\text{M}]$. In such regions the pressure is great enough that the N₂ molecules are plentiful relative to CH₄ so that reaction of CH with N₂ can compete. There is a lower concentration of CH₄ on the early Earth, and the haze forms at a lower altitude (ie higher pressure), both of which make the reaction of CH with N₂ more competitive (Wolf and Toon, 2010). Figure 7 shows the relative lifetimes of CH radical for an early Earth CH₄ mixing ratio of 1000 ppmv, and indicates that in the haze formation region this reaction mechanism could contribute to nitrogenated organics forming in the haze on early Earth. However other mechanisms are possible which could have different dependences on altitude. CH₄/N₂ photochemistry also holds interest for the study of Neptune's moon Triton (McDonald et al., 1994).

Interestingly, FUV nitrogen dissociation may have implications for prebiotic chemistry in planets beyond our solar system. Due to their high abundance in the galaxy and long-term stability, low-mass M stars (or M dwarfs) are often considered targets to search for terrestrial planets in the habitable zone (Scalo et al., 2007). Many researchers have evaluated the possibility that a planet like Earth could overcome the strong stellar flares (Segura et al., 2010) and lower insolation (Joshi et al., 1997) in order to maintain habitability in orbit around an M dwarf. An alternate side of this discussion is whether a planet like early Earth would have sufficiently advanced prebiotic chemistry for life to arise in the first place (Buccino et al., 2006). The M dwarf spectrum is shifted longward as compared to our sun and has a significantly decreased ultraviolet flux (Segura et al., 2005). Like the case of Early Earth discussed above, formation of C-N bonds would occur in an environment without strong EUV flux. Buccino et al. (2007) determined that there was not enough UV flux at planets around quiet M dwarf stars to

480 promote the prebiotic synthesis necessary to initiate life but didn't specify chemical reactions of
interest or take into account the wavelengths of light needed for CH₄ dissociation. Stars with
moderate to high flare activity will experience bursts of strong UV flux which may be sufficient
for this type of synthesis (Segura et al., 2010). Although spectra of M dwarfs provided in the
literature don't include radiation shortward of 100 nm, presumably the FUV flux is greater than
485 the EUV radiation otherwise required to break N₂ bonds (Segura et al., 2007; Walkowicz et al.,
2008). As shown in this paper, photolysis of even trace amounts of CH₄ and N₂ in an anoxic
atmosphere can cleave the N₂ bond and produce nitrogen-containing organic molecules without
the presence of EUV.

Ultimately, the unanticipated nitrogen chemistry signals a missing reaction mechanism
for the activation of N₂ via CH₄ photolysis at wavelengths greater than 115 nm. The observed
490 level of nitrogen incorporation, up to 16% by mass with a N/C of approximately 0.18, was not
predicted by current models of CH₄/N₂ chemistry at these wavelengths. Although possible
mechanisms have been proposed here, there is not enough data currently available to assess these
as the source of the nitrile groups in the aerosol. Thus, as indicated by the unexpected nature of
these results, a re-evaluation of the photochemical reaction schemes in this wavelength range and
495 further investigation into the nature of the secondary reactions leading to the chemical activation
of N₂ is clearly needed.

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Table 1. Comparison of results for previous CH₄/N₂ photolysis in Far-UV.

Wavelength Range (nm)	Reference	Reaction pressure (Torr)	Products Reported (phase)
125 – 170	Dodonova, 1966 (Dodonova, 1966)	5-8, 100-200	HCN (gas)
123.6	Chang et al., 1979 (Chang et al., 1979)	<i>not specified</i>	Organics, No N-containing (gas)
147	Bossard, 1979 <i>from Bossard et al., 1982</i> (Bossard et al., 1982)	10	Organics, No N-containing (gas)
184.9	Ferris and Chen, 1975 (Ferris and Chen, 1975)	670	Organics, No N-containing (gas)
112 - 135	Scattergood et al., 1989 (Scattergood et al., 1989)	<i>not specified</i>	HCN, hydrocarbons (gas)
50 – 150	Imanaka and Smith, 2007 (Imanaka and Smith, 2007)	0.05	Organics produced at Ly- α (gas) but with reduced efficiency than lower λ , N chemistry uncertain
115 – 400	This work	600	Nitriles, -CN linkages (solid)

Figure Legends.

720 **FIG. 1.** Elemental analysis from the high-resolution mass spectrum of Titan aerosol analogs produced from photochemistry of 0.1% CH₄ in N₂ background gas. These results show a substantial fraction of nitrogen included in the aerosol material by mass. Atomic ratios are indicated in legend. The mass spectrum shows fragmentation patterns typical of unsaturated hydrocarbons with amine and/or nitrile groups. The figure includes the lists of skeletal formulas
725 (without H) indicated by the high-resolution mass spectra of the integer m/z for each group of peaks on the “picket fence”, showing a variety of molecular fragments.

FIG. 2. The raw Time-of-Flight (ToF) mass spectrum is fit to quantify the ion contributions to the signal. The fitting algorithm determines the peak area to quantify the contribution from each ion and derive the elemental analysis (Aiken et al., 2008). The comparison between the raw
730 data and the total fit provide an indication of the goodness of fit. For m/z 27, the photochemical aerosol arising from a 0.1% CH₄/ 99.9% N₂ atmosphere shows peaks at m/z values of 27.011 and 27.023 which are most likely due to ions HCN⁺ and C₂H₃⁺, respectively.

FIG. 3. Averaged aerosol mass spectra show the effect of N₂ concentration on the aerosol composition and relative mass production rate. The signals for the gases N₂ and Ar have been
735 subtracted from the spectra.

FIG. 4. High resolution ion signal for key fragment ions observed in the aerosol spectrum from 0.1% CH₄ photolysis. The relative contribution to ion peaks as a function of background gas mixture (N₂:Ar) is demonstrated by the solid and dashed traces. The peak at m/z 29.003 (the ion identifications marked with *) may be due to either of two molecular fragments which are
740 indistinguishable at the instrument resolution. Ion signal is shown in arbitrary units, and traces have been normalized so that the height of the largest ion is similar in each plot.

FIG. 5. Plots of the ratio of the contribution of the nitrile (HCN^+) to the hydrocarbon (C_2H_3^+) ion fragments for m/z 27 signal (top panel) and the N/C ratio in the aerosol spectra (bottom panel), as a function of the proportion of N_2 in the background gas during photolysis of 0.1% CH_4 . Argon was used as the balance gas during these experiments and the total pressure in the reaction cell was held constant at 600 Torr. Each point represents an averaged value for one experiment. For the m/z 27 peak height data in the top panel the error bars show the standard deviation to indicate the variability within each experiment. For the N/C data in the bottom panel the error bars represent the 20% error in this ratio cited for the instrument (Aiken et al., 2007). The dashed lines are included as a guide to the eye.

FIG. 6. The absorption cross-sections for CH_4 (Mount et al., 1977; Kameta et al., 2002; Chen and Wu, 2004) and N_2 (Chan et al., 1993) show that direct dissociation of N_2 is not expected in the FUV. The vertical dashed line indicates the lower limit of our lamp emission. Huffman (1969) does report a cross-section measurement at Ly- α of $6 \times 10^{-23} \text{ cm}^2$ (open circle), but provides this only as an upper bound based on the detection limit of the instrumentation. As demonstrated by the graph, even with this upper limit the absorption of N_2 is not competitive as compared to CH_4 and direct dissociation cannot explain the high percentage of N_2 incorporated into the aerosol material.

FIG. 7. The lifetime of the CH radical, produced via CH_4 photolysis, is evaluated for both Titan and the early Earth for two different reactions pathways. Curves labeled $k_c[\text{CH}_4]$ represent the lifetime with respect to the reaction $\text{CH} + \text{CH}_4$, with the second order rate constant $k_c=10^{-10} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Yung et al., 1984). Curves labeled $k_n[\text{N}_2][\text{M}]$ represent the lifetime with respect to the reaction $\text{CH} + \text{N}_2 + \text{M}$ (R3), with the third order rate constant $k_n=5.7 \times 10^{-30} \text{ cm}^6 \text{ molecule}^{-2}$

765 s^{-1} (Le Picard and Canosa, 1998). A concentration of approximately 2% CH_4 is assumed for
Titan, and 1000 ppmv for the Early Earth. The curves cross at the transition where the pressure
is great enough that the number density of N_2 molecules outnumber the CH_4 molecules by
enough to successfully compete for reaction with CH. If this transition occurs above or within
the haze formation region – shown on the plot with the cloud cartoon – then the formation of
770 nitriles can influence the chemical composition of the haze.

Nitrogen Incorporation in CH₄-N₂ Photochemical Aerosol Produced by Far UV Irradiation

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Figures and Figure Legends

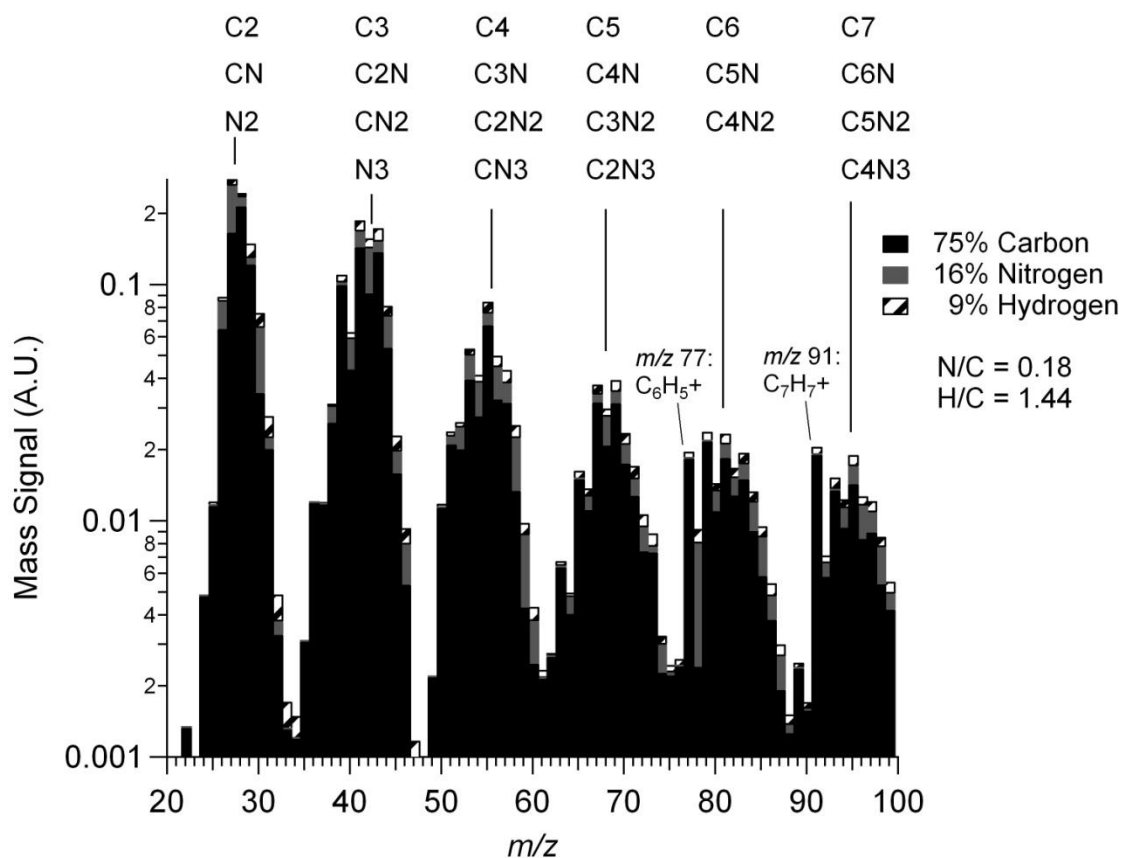


FIG. 1. Elemental analysis from the high-resolution mass spectrum of Titan aerosol analogs produced from photochemistry of 0.1% CH₄ in N₂ background gas. These results show a substantial fraction of nitrogen included in the aerosol material by mass. Atomic ratios are indicated in legend. The mass spectrum shows fragmentation patterns typical of unsaturated hydrocarbons with amine and/or nitrile groups. The figure includes the lists of skeletal formulas (without H) indicated by the high-resolution mass spectra of the integer m/z for each group of peaks on the “picket fence”, showing a variety of molecular fragments.

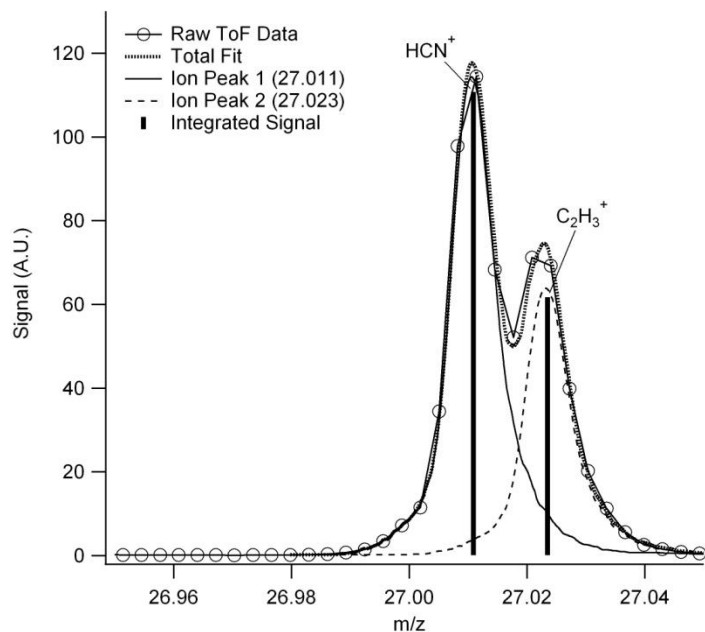


FIG. 2. The raw Time-of-Flight (ToF) mass spectrum is fit to quantify the ion contributions to the signal. The fitting algorithm determines the peak area to quantify the contribution from each ion and derive the elemental analysis (Aiken et al., 2008). The comparison between the raw data and the total fit provide an indication of the goodness of fit. For m/z 27, the photochemical aerosol arising from a 0.1% CH_4 / 99.9% N_2 atmosphere shows peaks at m/z values of 27.011 and 27.023 which are most likely due to ions HCN^+ and C_2H_3^+ , respectively.

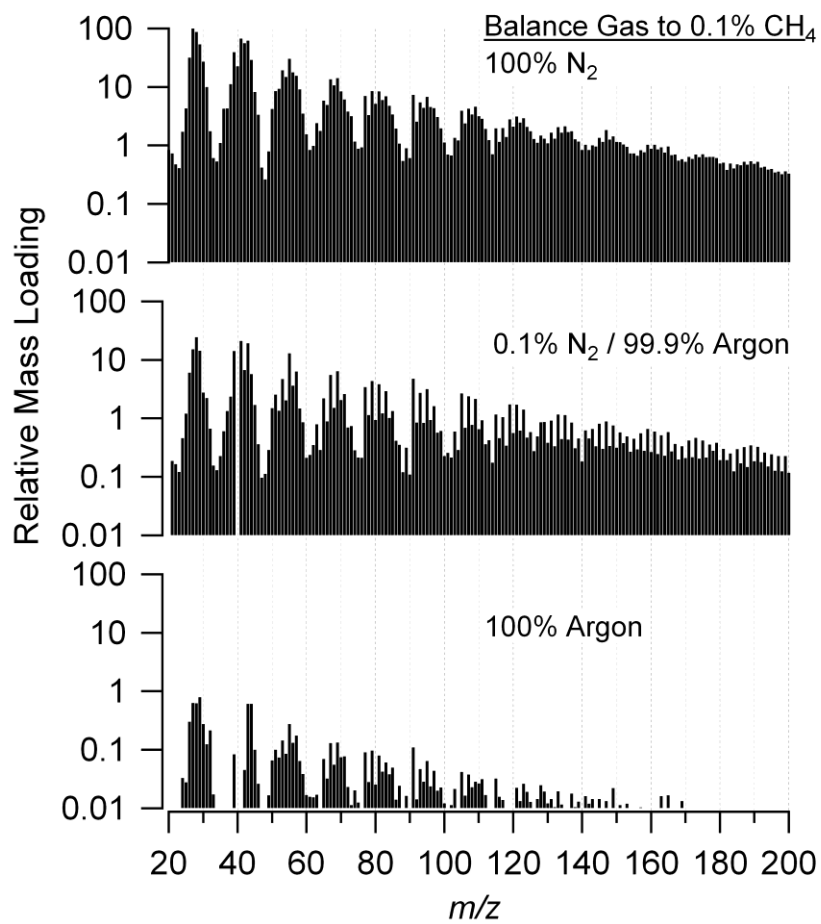


FIG. 3. Averaged aerosol mass spectra show the effect of N₂ concentration on the aerosol composition and relative mass production rate. The signals for the gases N₂ and Ar have been subtracted from the spectra.

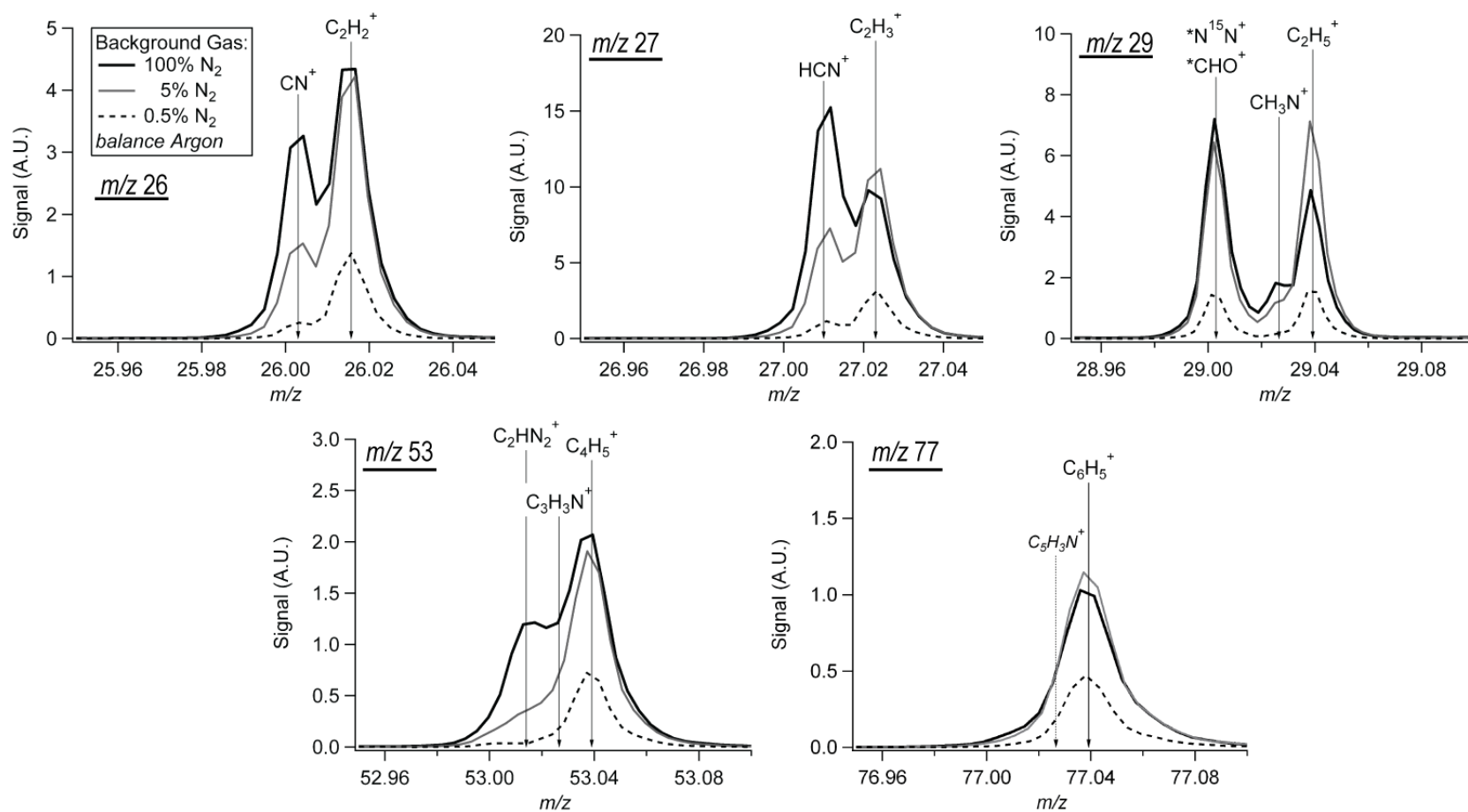


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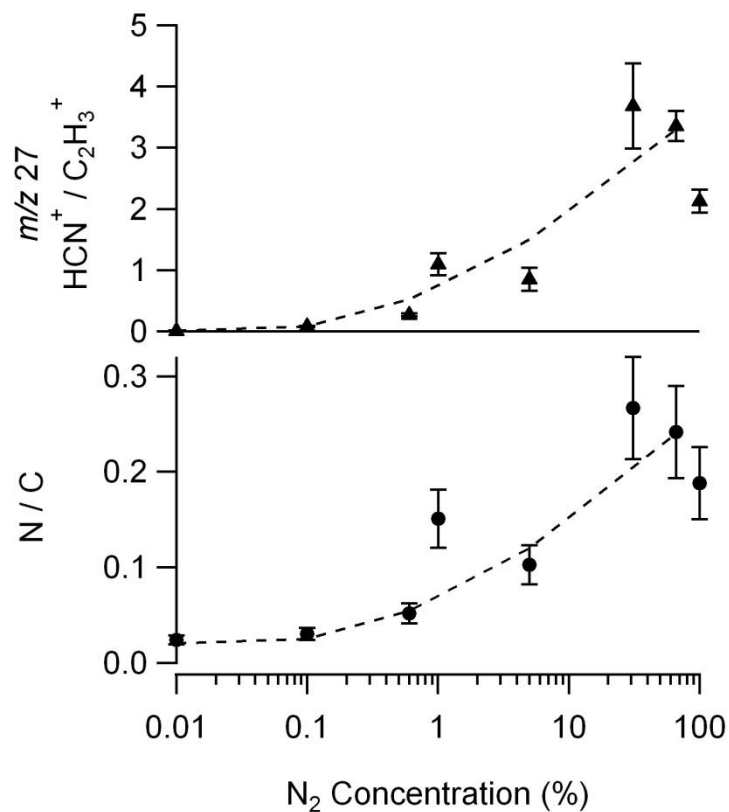


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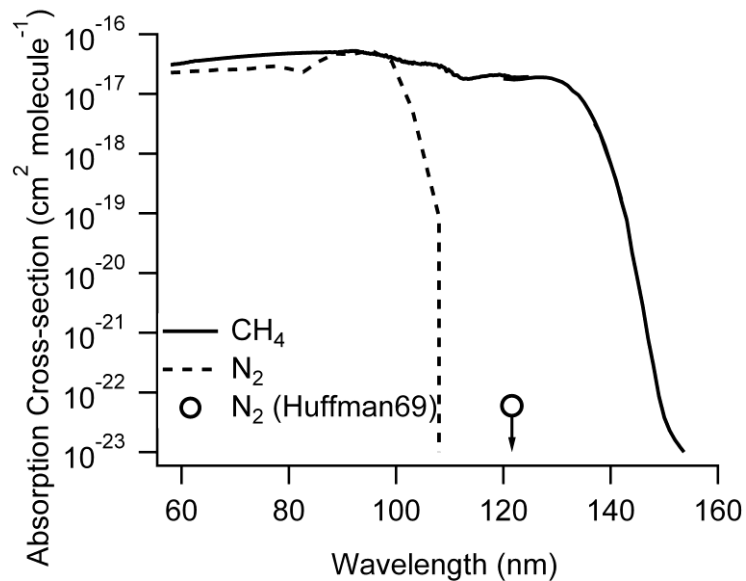


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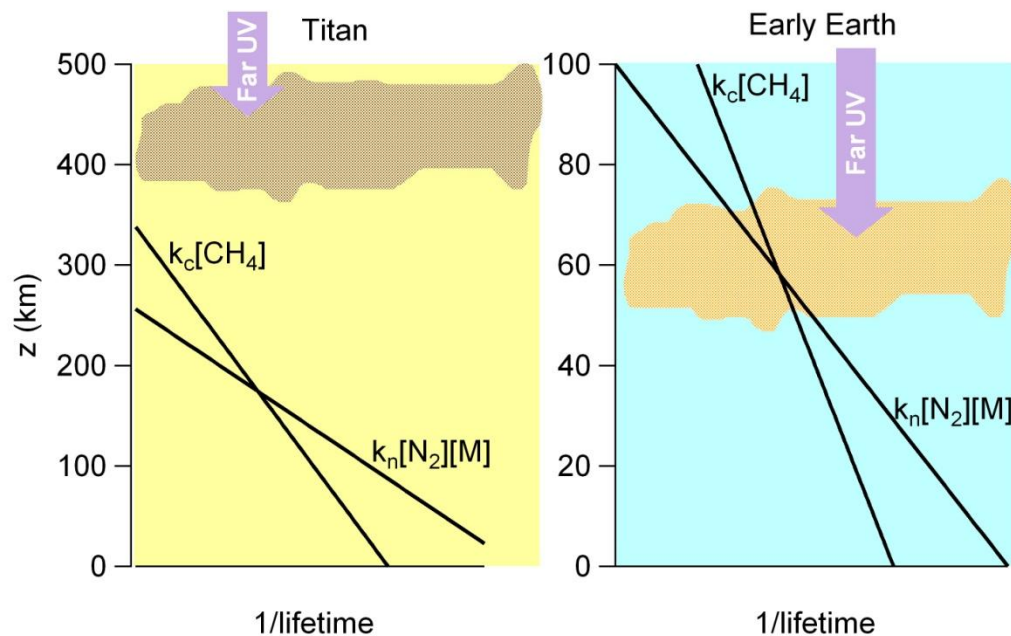


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