Introduction to radiative transfer theory and models (Optical Domain)

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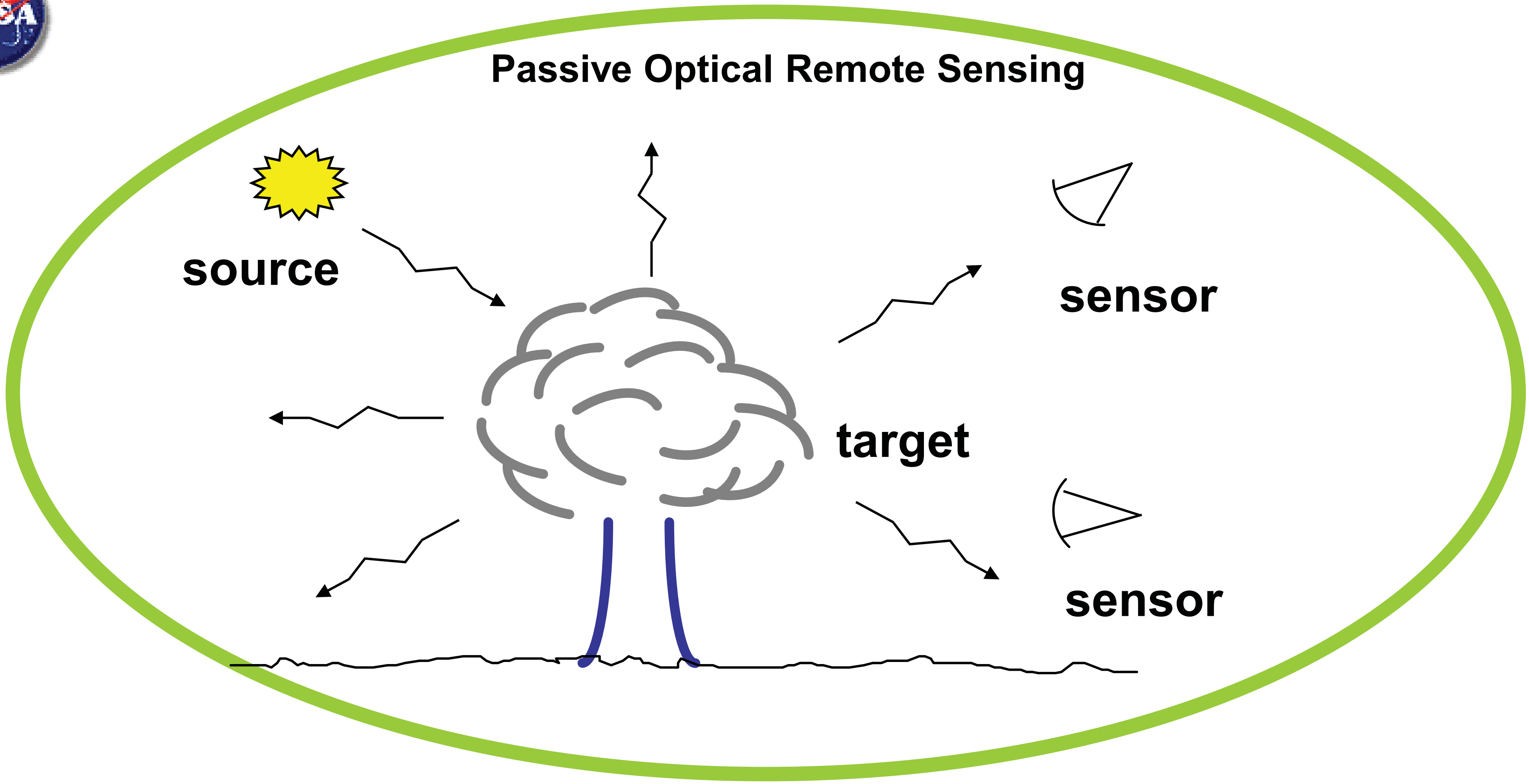
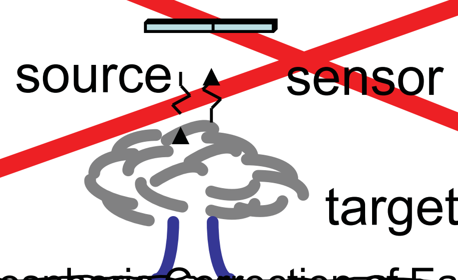
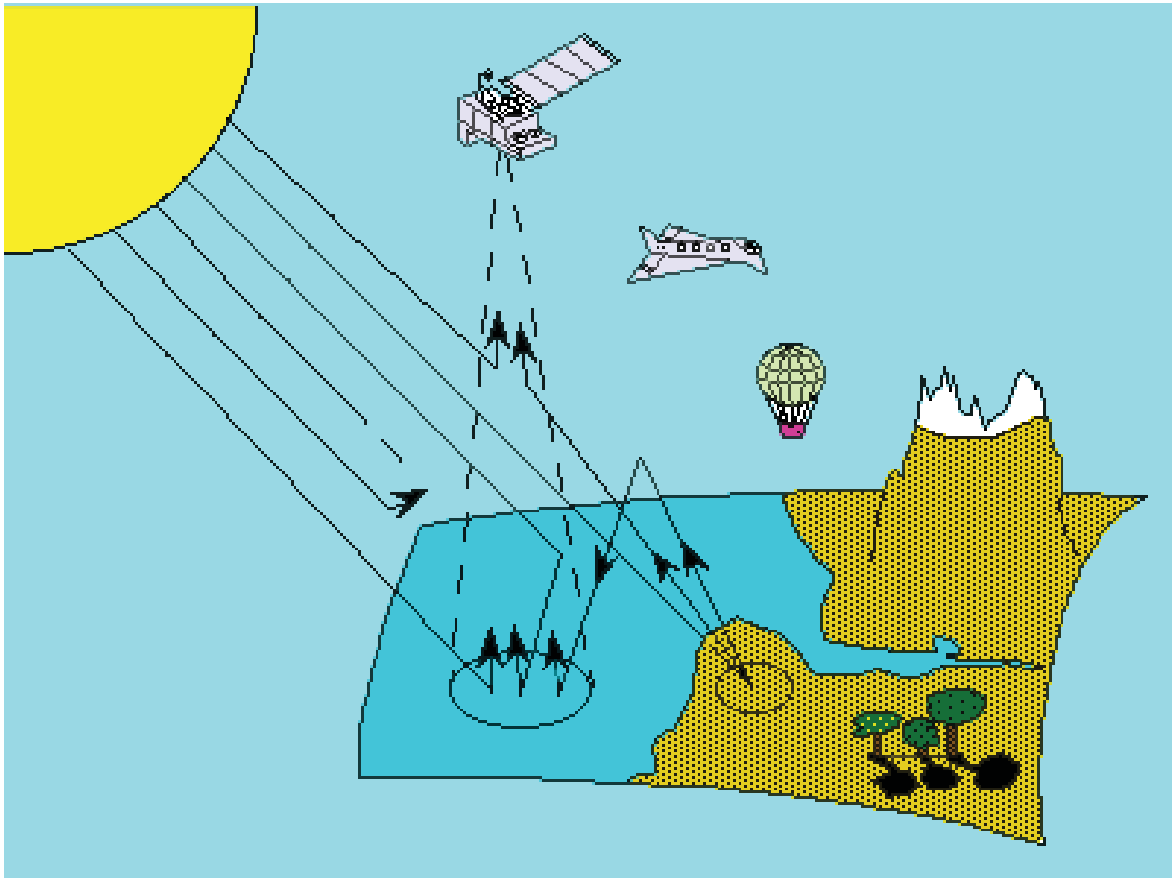


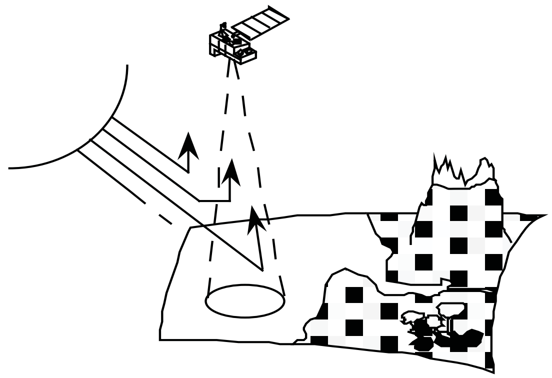
~~RADAR: radiation generated next to sensor~~



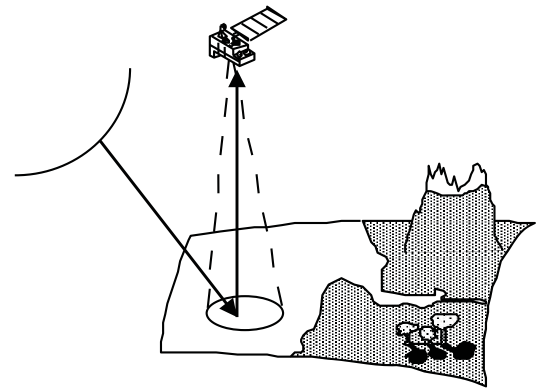
~~Atmospheric Correction of Earth~~
Observation Data for
Environmental Monitoring: Theory

Solar Energy Paths

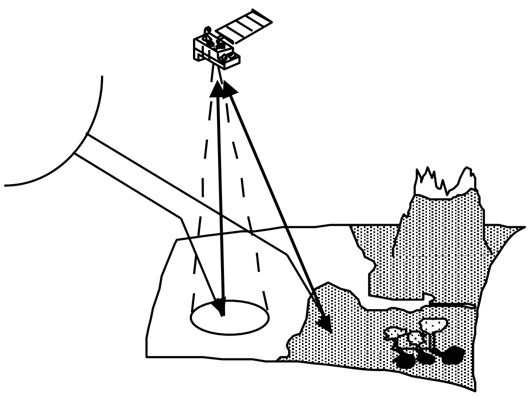




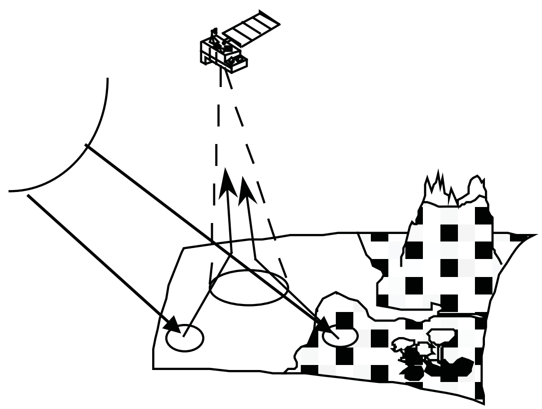
atmospheric contribution



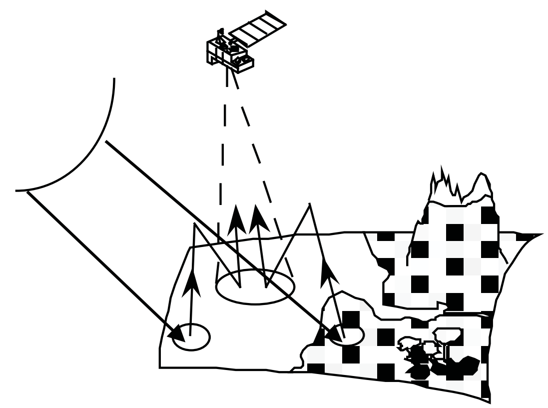
direct + direct



diffuse + direct



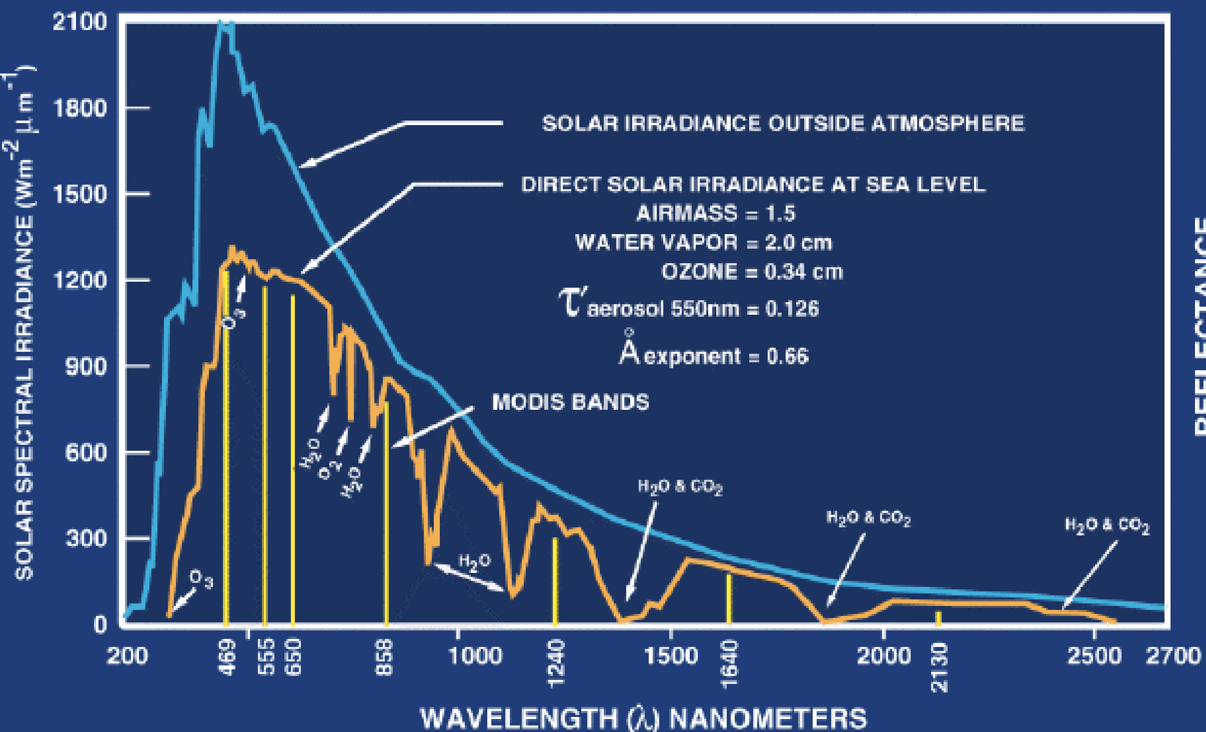
direct + diffuse

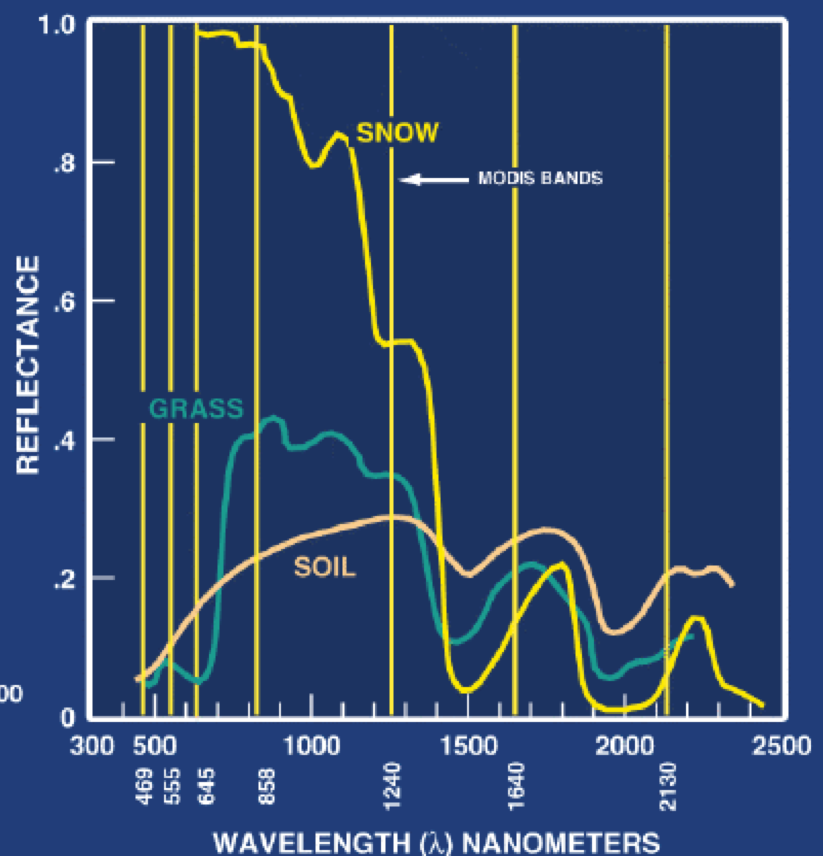


multiple scattering

Atmospheric Correction of Earth
Observation Data for
Environmental Monitoring: Theory

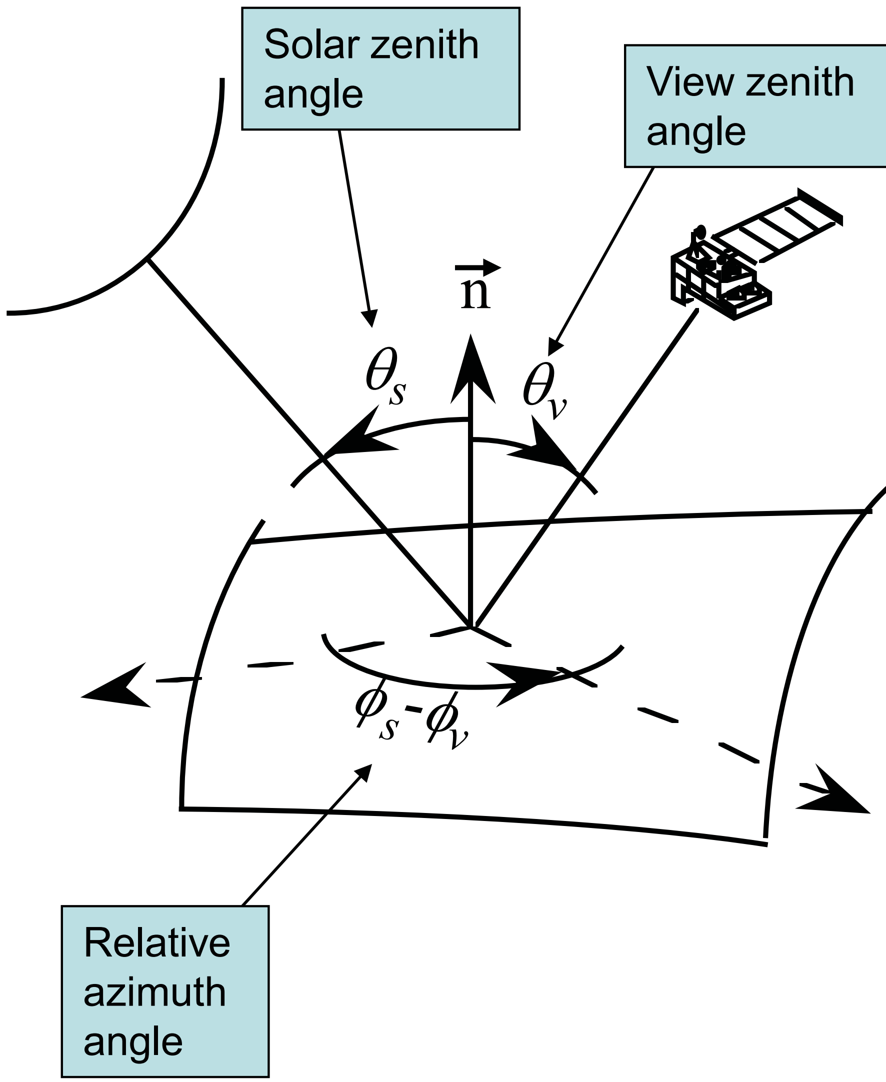
Solar (reflective) spectral domain





Atmospheric Correction of Earth Observation Data for Environmental Monitoring: Theory

Observation Geometry



Solution of the Radiative Transfer in the reflective domain for non absorbing atmosphere and lambertian ground

Atmospheric reflectance

Ground reflectance (= albedo for lambertian)

$$\rho_{app}(\theta_s,\theta_v,\phi) = \rho_{atm}(\theta_s,\theta_v,\phi) + T_{atm}(\theta_s)T_{atm}(\theta_v)\frac{\rho_{ground}}{1-\rho_{ground}S_{atm}}$$

Atmospheric Transmissions

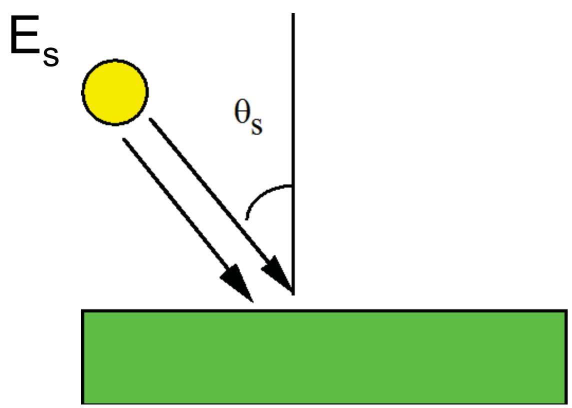
Apparent reflectance at satellite level

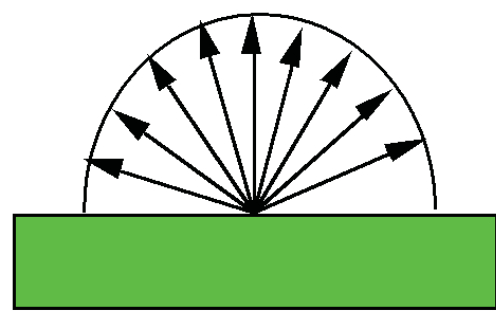
Atmosphere spherical albedo

Perfect Lambertian Reflector

Radiance of the Perfect Lambertian Reflector

$$\int_0^{\pi}\int_0^{2\pi} RPLF(\theta_S,\theta,\phi)\cos(\theta)\sin(\theta)\,d\theta\,d\phi = E_s\cos(\theta_s)$$



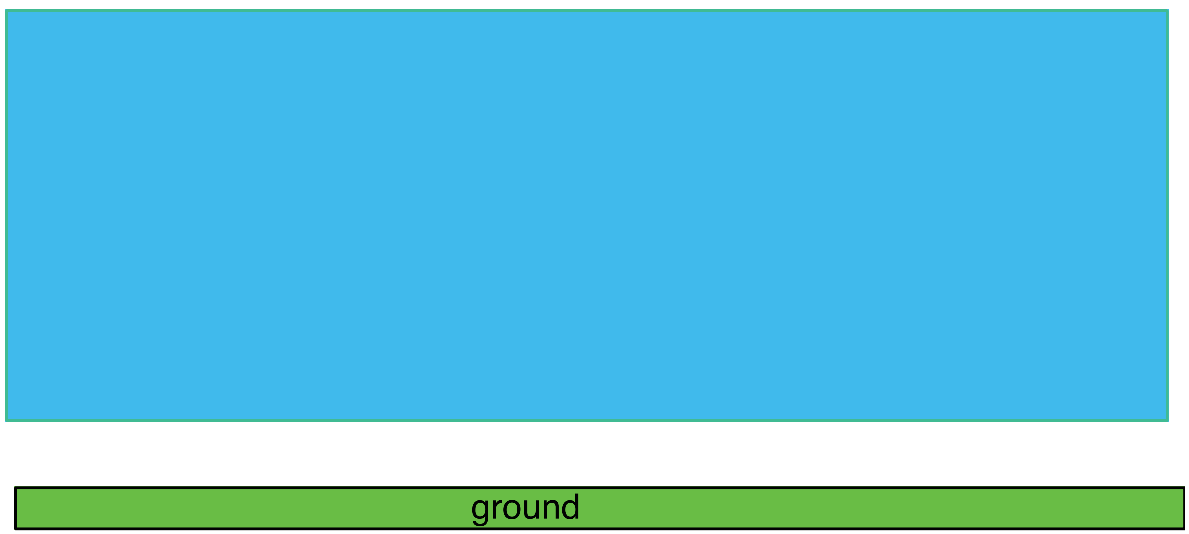


Isotropic radiation

$$\rho_{Perfect\quad Lambertian\quad reflector}(\theta_S,\theta_V,\phi) = 1$$

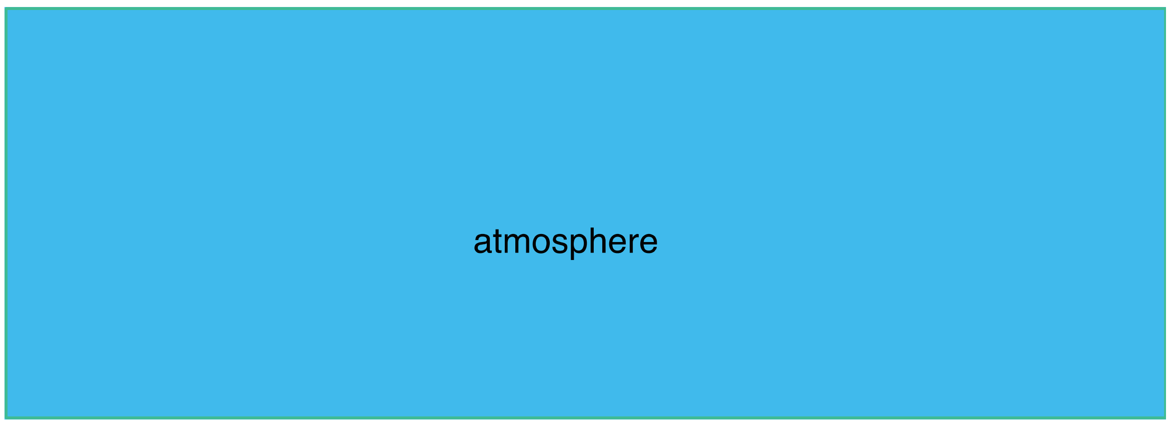
$$\rho_{Lambertian\quad reflector}(\theta_S,\theta_V,\phi) = \rho$$

Simple Radiative Transfer Equation



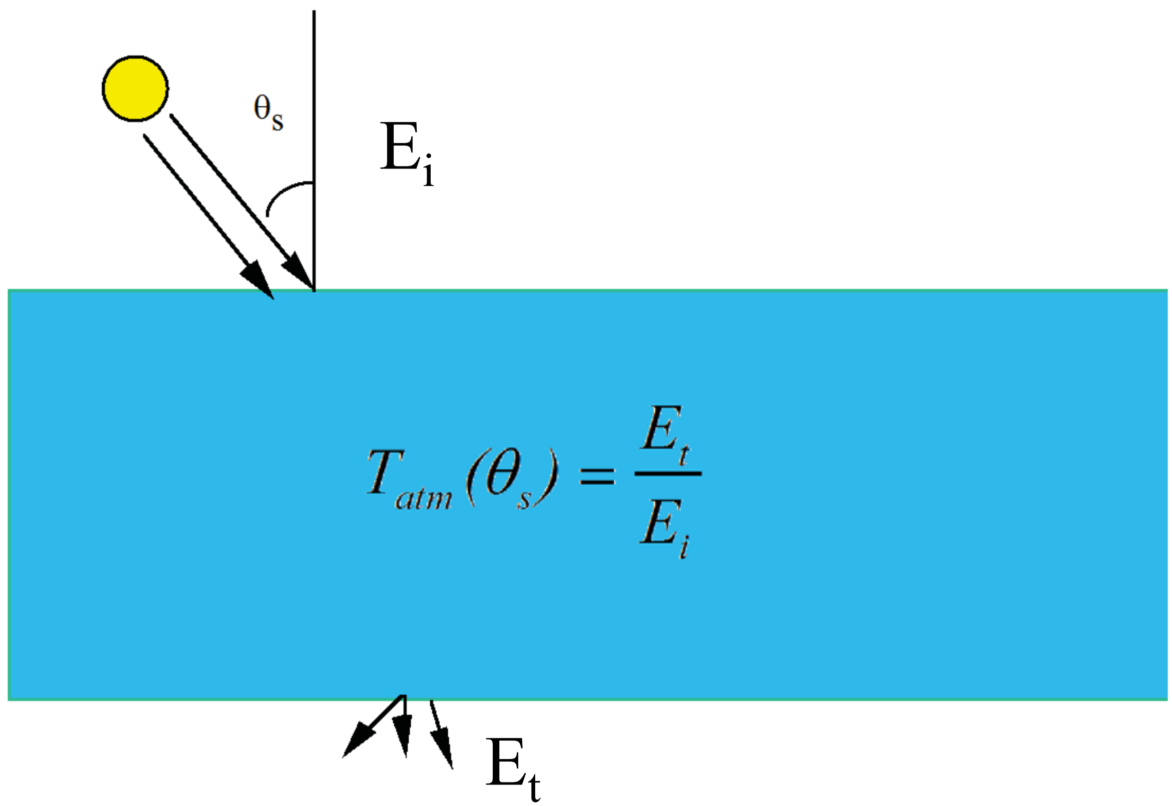
SRTE (cont.)

$$\rho_{atm}(\theta_s, \theta_v, \phi)$$



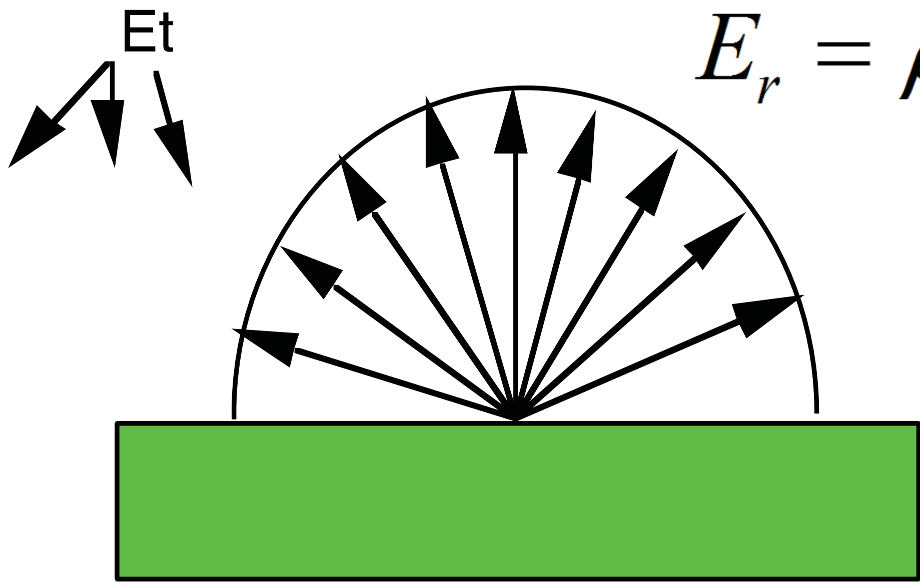
Absorbing ground

SRTE (cont.)



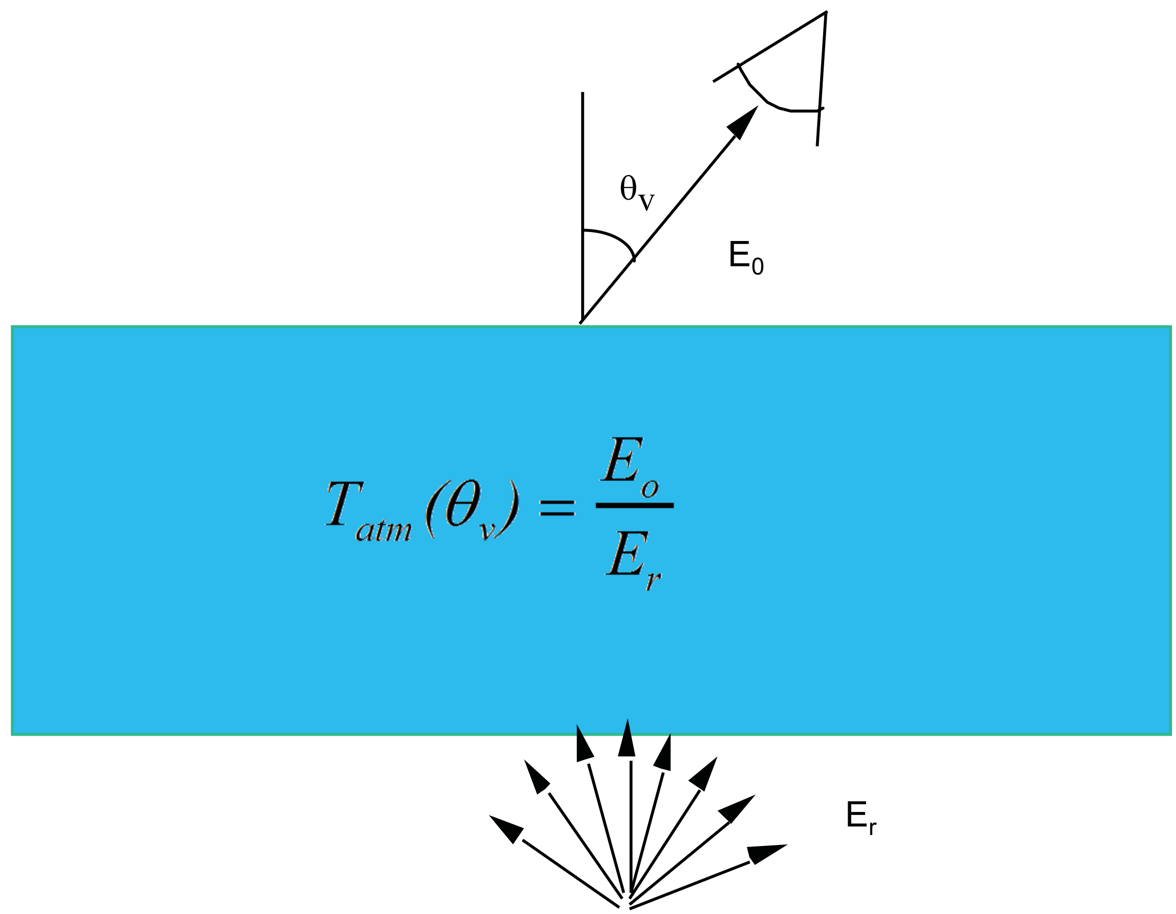
$$T_{atm}(\theta_s) = \frac{E_t}{E_i}$$

SRTE (cont.)

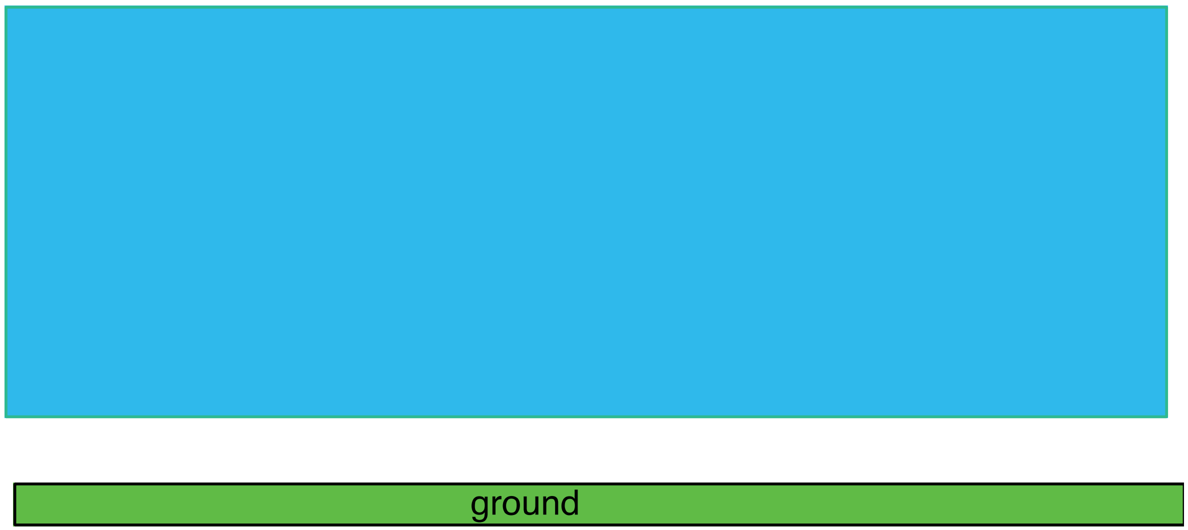


$$E_r = \rho_{ground} E_t$$

SRTE (cont.)



SRTE 1 interaction (cont.)

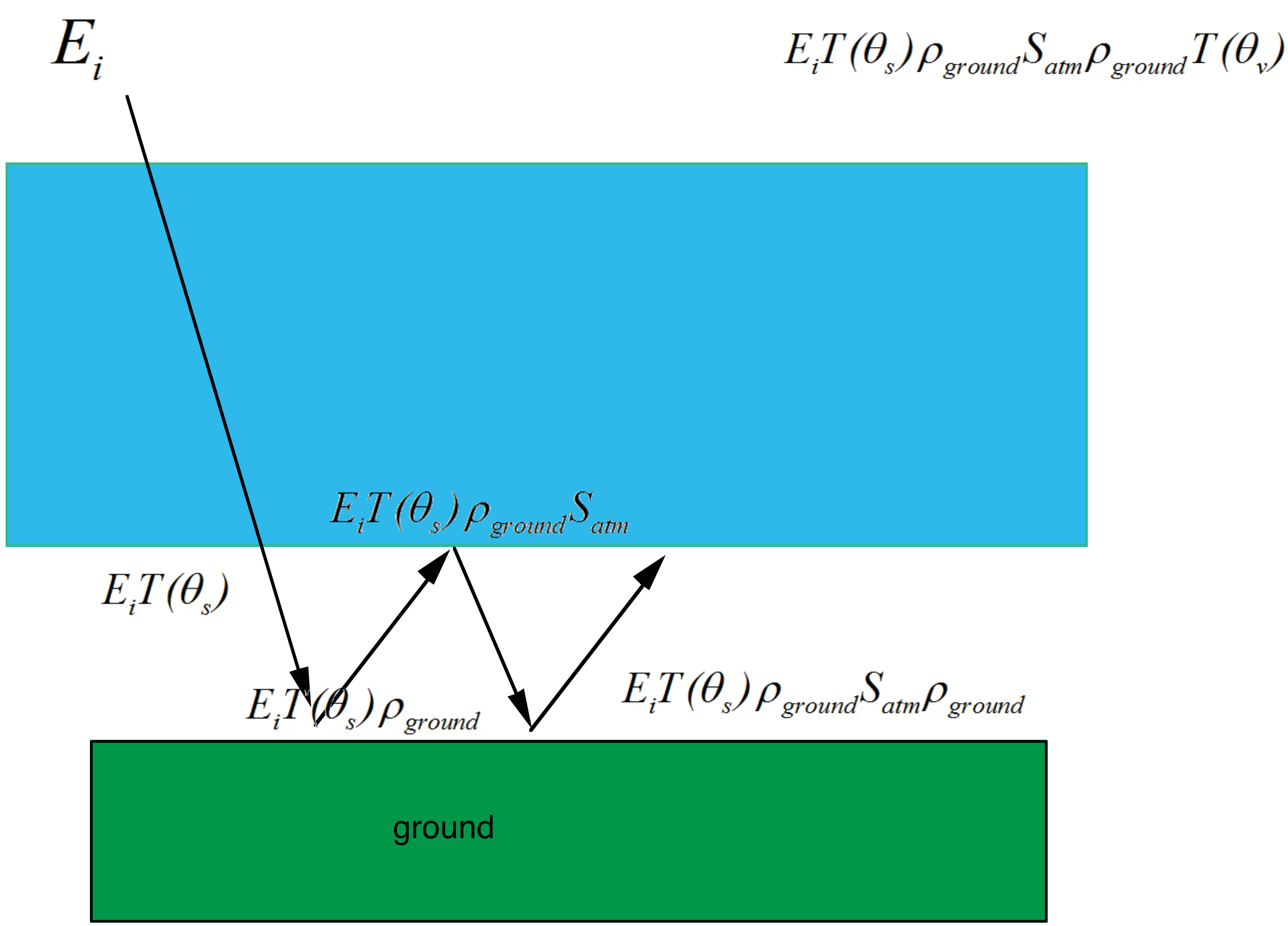


$$\rho_{app} = \rho_{atm} + \frac{E_o}{E_i}$$

$$\frac{E_o}{E_i} = \frac{T(\theta_v)E_r}{E_i} = \frac{T(\theta_v)\rho_{ground}E_t}{E_i} = T(\theta_v)\rho_{ground}T(\theta_s)$$

$$\rho_{app} = \rho_{atm} + T(\theta_v)\rho_{ground}T(\theta_s)$$

SRTE 2 interactions



SRTE Multiple Interactions

$$\rho_{app} = \rho_{atm} + T(\theta_s)T(\theta_v)\rho_{ground}\left[1+\rho_{ground}S_{atm} + \left(\rho_{ground}S_{atm}\right)^2 + \left(\rho_{ground}S_{atm}\right)^3 \ldots\right]$$

$$1 + r + r^2 + r^3 + \ldots r^{n-1} = \frac{1-r^n}{1-r}$$

$\rho_{ground}S_{atm} < 1$ so when n->∞ then $(\rho_{ground}S_{atm})^n$ ->0

Therefore $$\left[1+\rho_{ground}S + \left(\rho_{ground}S\right)^2 + \left(\rho_{ground}S\right)^3 \ldots\right] = \frac{1}{1-\rho_{ground}S}$$

$$\rho_{app} = \rho_{atm} + T_{atm}(\theta_s)T_{atm}(\theta_v)\frac{\rho_{ground}}{1-\rho_{ground}S_{atm}}$$

STRE for non absorbing atmosphere and lambertian ground

Atmospheric reflectance

Ground reflectance (= albedo for lambertian)

$$\rho_{app}(\theta_s, \theta_v, \phi) = \rho_{atm}(\theta_s, \theta_v, \phi) + T_{atm}(\theta_s) T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground} S_{atm}}$$

Atmospheric Transmissions

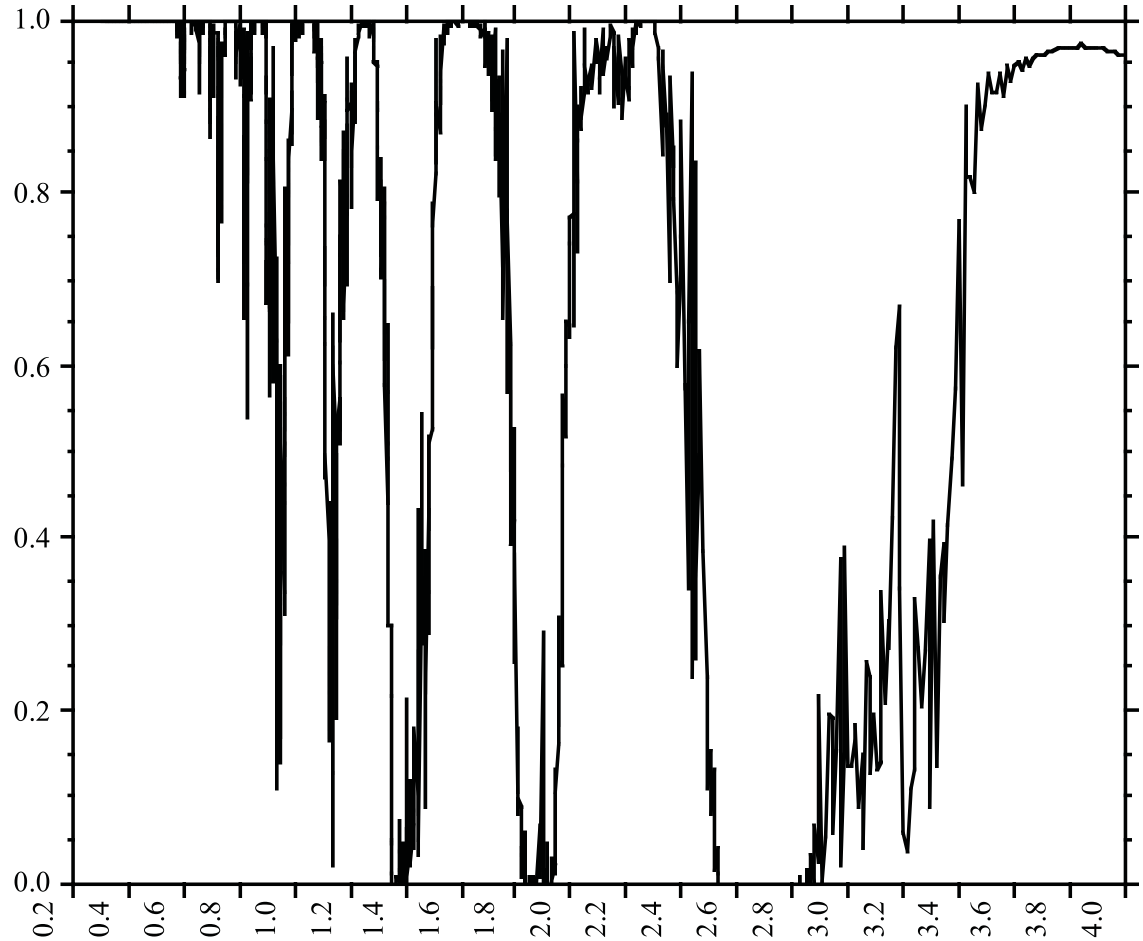
Apparent reflectance at satellite level

Atmosphere spherical albedo

The composition of the atmosphere

Permanent Constituents		Variable constituents	
Constituent	% by volume	Constituent	% by volume
Nitrogen (N_2)	78.084	Water Vapor (H_2O)	0.04
Oxygen (O_2)	20.948	Ozone (O_3)	12×10^{-4}
Argon (Ar)	0.934	Sulfur dioxide (SO_2)[b]	0.001×10^{-4}
Carbon dioxide (CO_2)	0.033	Nitrogen dioxide (NO_2)	0.001×10^{-4}
Neon (Ne)	18.18×10^{-4}	Ammonia (NH_3)	0.001×10^{-4}
Helium (He)	5.24×10^{-4}	Nitric oxide (NO)	0.0005×10^{-4}
Krypton (Kr)	1.14×10^{-4}	Hydrogen sulfide (H_2S)	0.00005×10^{-4}
Xenon (Xe)	0.089×10^{-4}	Nitric acid vapor	trace
Hydrogen (H_2)	0.5×10^{-4}		
Methane (CH_4)	1.5×10^{-4}		
Nitrous Oxide (N_2O)	0.27×10^{-4}		
Carbon Monoxide (CO)	0.19×10^{-4}		

Gaseous Absorption (H_2O)



Modified SRTE to account for absorption

In case of a pure molecular atmosphere (no aerosol) we can write:

$$\rho_{app}(\theta_s,\theta_v,\phi) = Tg^{othergases}(m,U_{gaz})\left[\rho_{atm}(\theta_s,\theta_v,\phi) + Tg^{H_2O}(m,U_{H_2O})T_{atm}(\theta_s)T_{atm}(\theta_v)\frac{\rho_{ground}}{1-S_{atm}\rho_{ground}}\right]$$

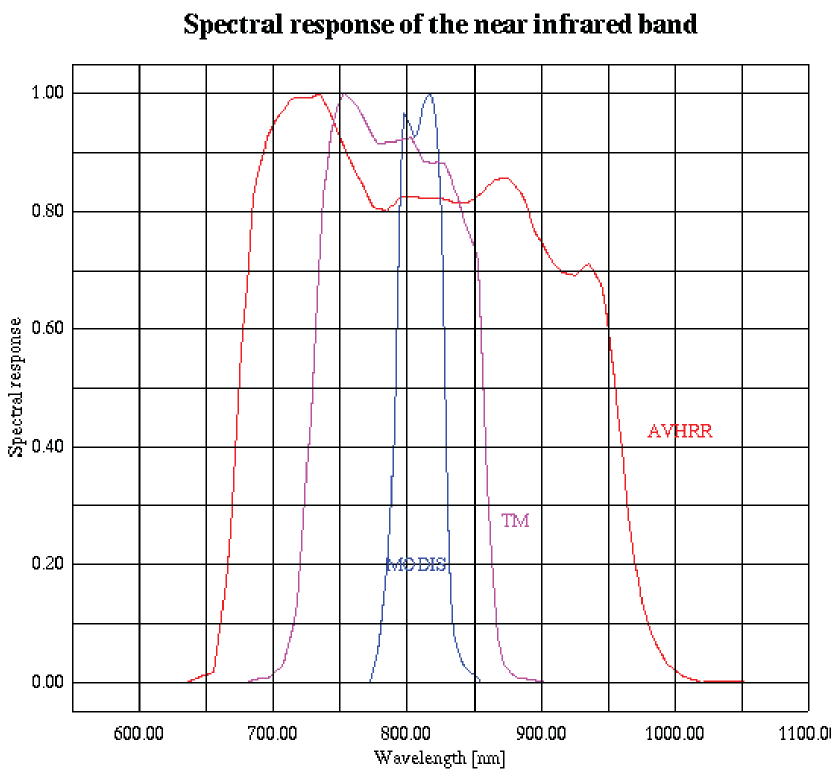
m is the air mass = $1/\cos(\theta_s)+1/\cos(\theta_v)$

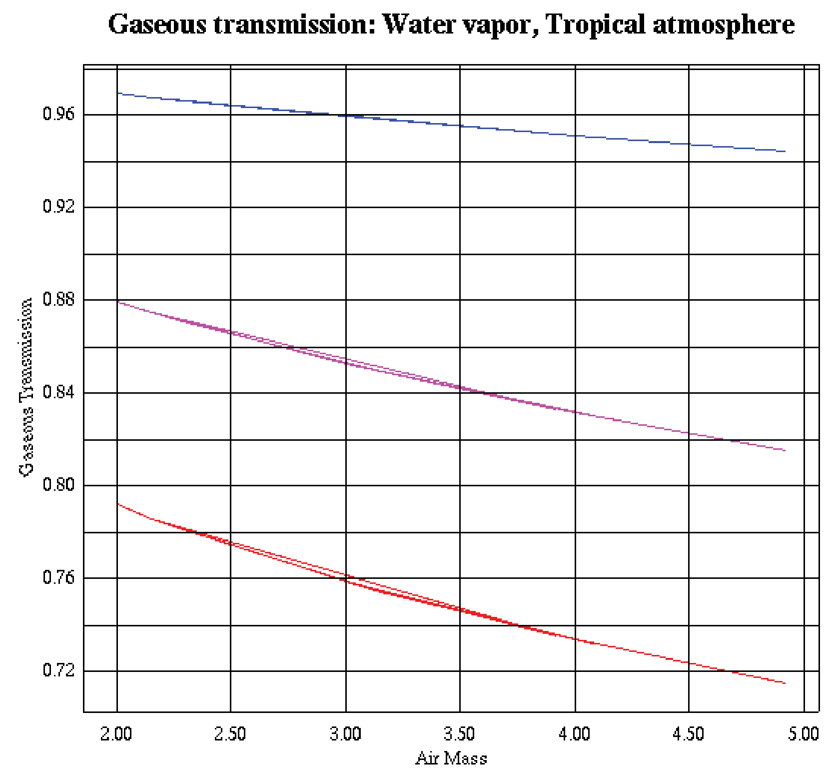
U_{gaz} is the gaz concentration

Final SRTE approximation

$$\rho_{app}(\theta_s,\theta_v,\phi) \sim Tg^{othergases}(m,U_{gaz})\left[\begin{array}{l}\rho_R(\theta_s,\theta_v,\phi) + Tg^{H_2O}(m,U_{H_2O}/2)\rho_A(\theta_s,\theta_v,\phi) \\ +Tg^{H_2O}(m,U_{H_2O})T_A(\theta_s)T_A(\theta_v)T_R(\theta_s)T_R(\theta_v)\dfrac{\rho_{ground}}{1-S_{R+A}\rho_{ground}}\end{array}\right]$$

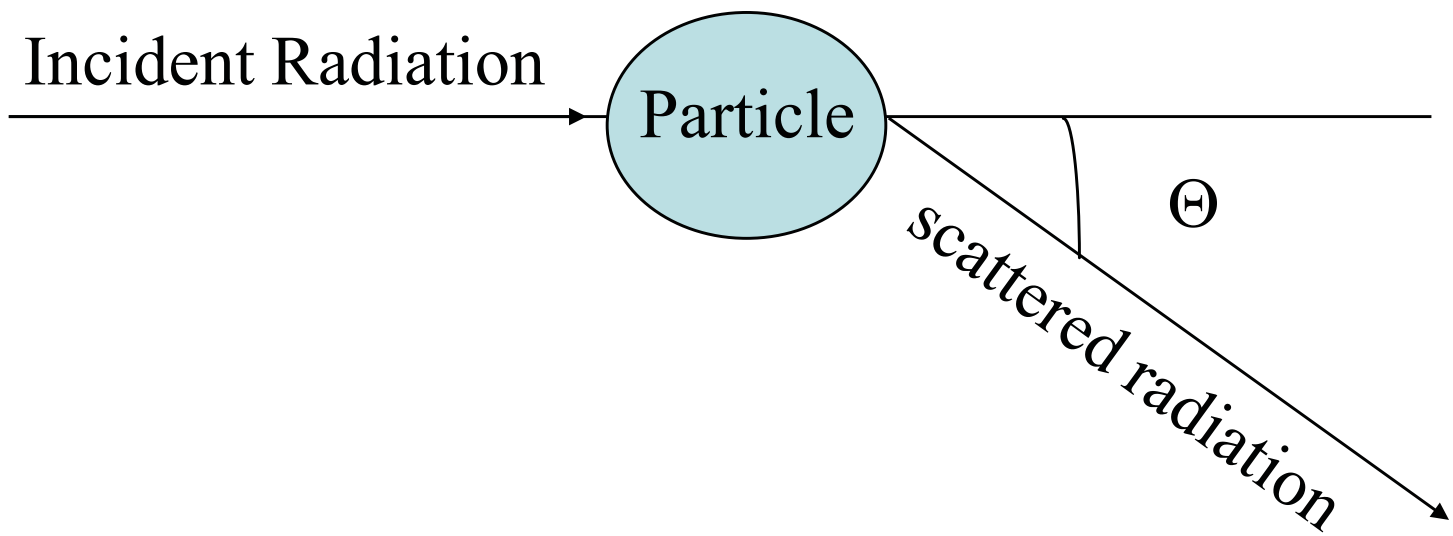
Water vapor effect for different sensors in the near infrared





Scattering angle ,

- The scattering angle, Θ, is the relative angle between the incident and the scattered radiation



Phase function

- The phase function, $P(\Theta)$, describe the distribution of scattered radiation for one or an set of particles. It is normalized such as:

$$\int_0^{2\pi}\int_0^{\pi} P(\Theta)d\omega = 4\pi$$

since

$$\int_0^{2\pi}\int_0^{\pi} P(\Theta)\sin(\theta)d\theta d\phi = 2\pi\int_0^{\pi} P(\theta)\sin(\theta)\,d\theta$$

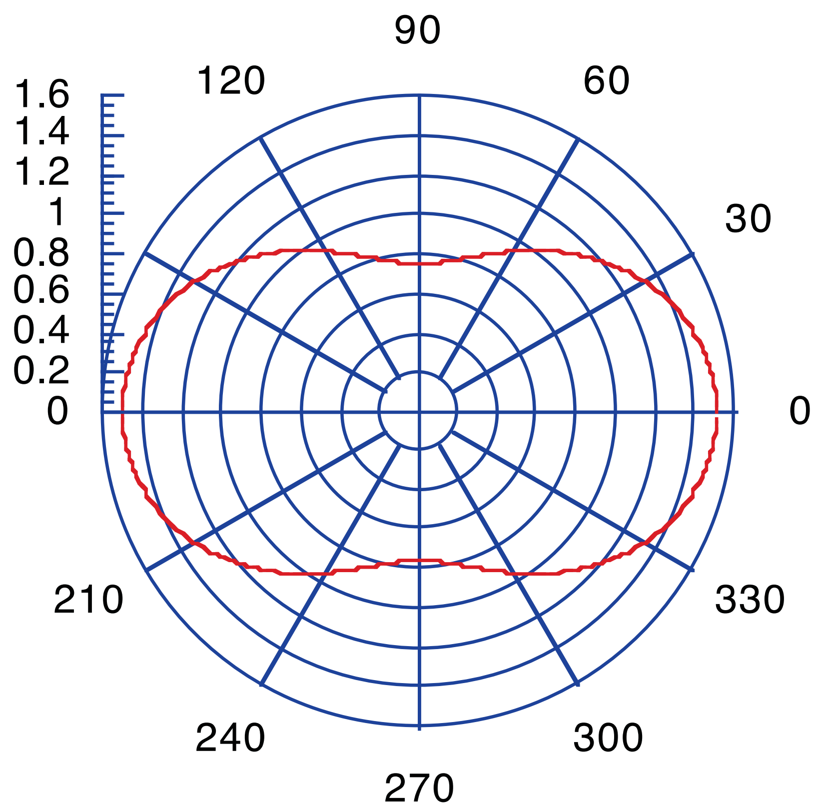
we have

$$\int_0^{\pi} P(\theta)\,\sin(\theta)d\theta = 2$$

Rayleigh/molecular scattering 1/4

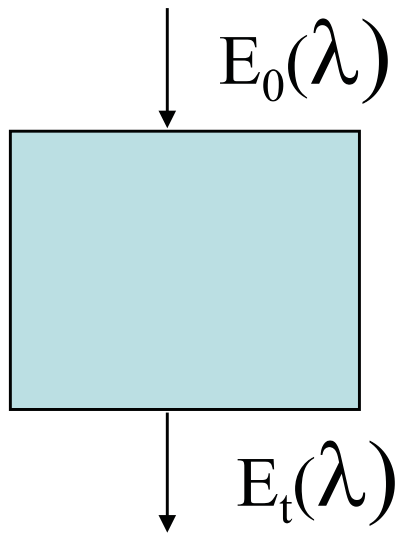
- Rayleigh or molecular scattering refers to scattering by atmospheric gases, in that case:

$$P(\Theta) = \frac{3}{4}\left(1 + \cos^2(\Theta)\right)$$



Rayleigh/molecular scattering 2/4

- The concentration in scatterer is better described by the efficiency they scatter at a certain wavelength or the proportion of direct transmission which is related to the spectral optical thickness $\tau(\lambda)$



$$E_t(\lambda)/\,E_0(\lambda)=e^{-\tau(\lambda)}$$

- For Rayleigh $\tau(\lambda)$ is proportional to λ^{-4} and for standart pressure is ~ 0.235 at 0.45 μm

Rayleigh/molecular scattering 3/4

- The rayleigh reflectance, ρ_R, could be crudely approximated by:

$$\rho_R^{\lambda}(\theta_s, \theta_v, \phi) \sim \frac{\tau_R^{\lambda} P(\Theta)}{4\cos(\theta_s)\cos(\theta_v)}$$

Rayleigh/molecular scattering 4/4

- Compute the reflectance of the sky (assumed clear no aerosol) at solar noon at 45degree latitude at vernal equinox looking straight up at 0.45μm, 0.55μm, 0.65μm

$$\rho_R^{0.45\,\mu m}(45°,0°,\phi)=\frac{0.235P(45°)}{4\cos(45°)\cos(0°)}$$

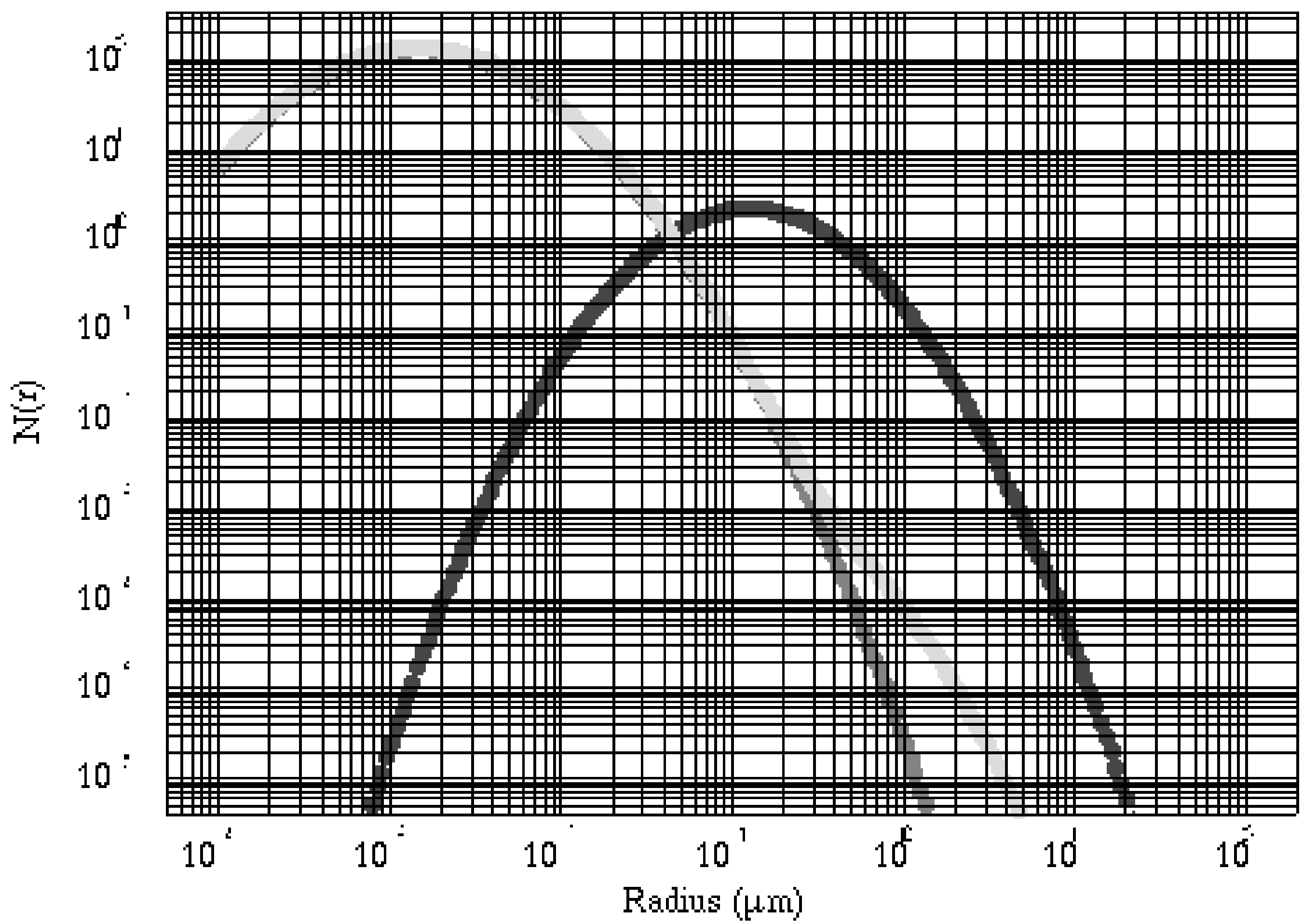
$$=\frac{0.235x0.75(1+\cos^2(45))}{4\cos(45°)\cos(0°)}=\frac{0.265}{2.8}\sim 0.1$$

$$\rho_R^{0.55\,\mu m}=0.1x(0.45/0.55)^4\sim 0.1x0.4=0.04$$

$$\rho_R^{0.65\,\mu m}=0.04x(0.55/0.65)^4\sim 0.04x0.5=0.02$$

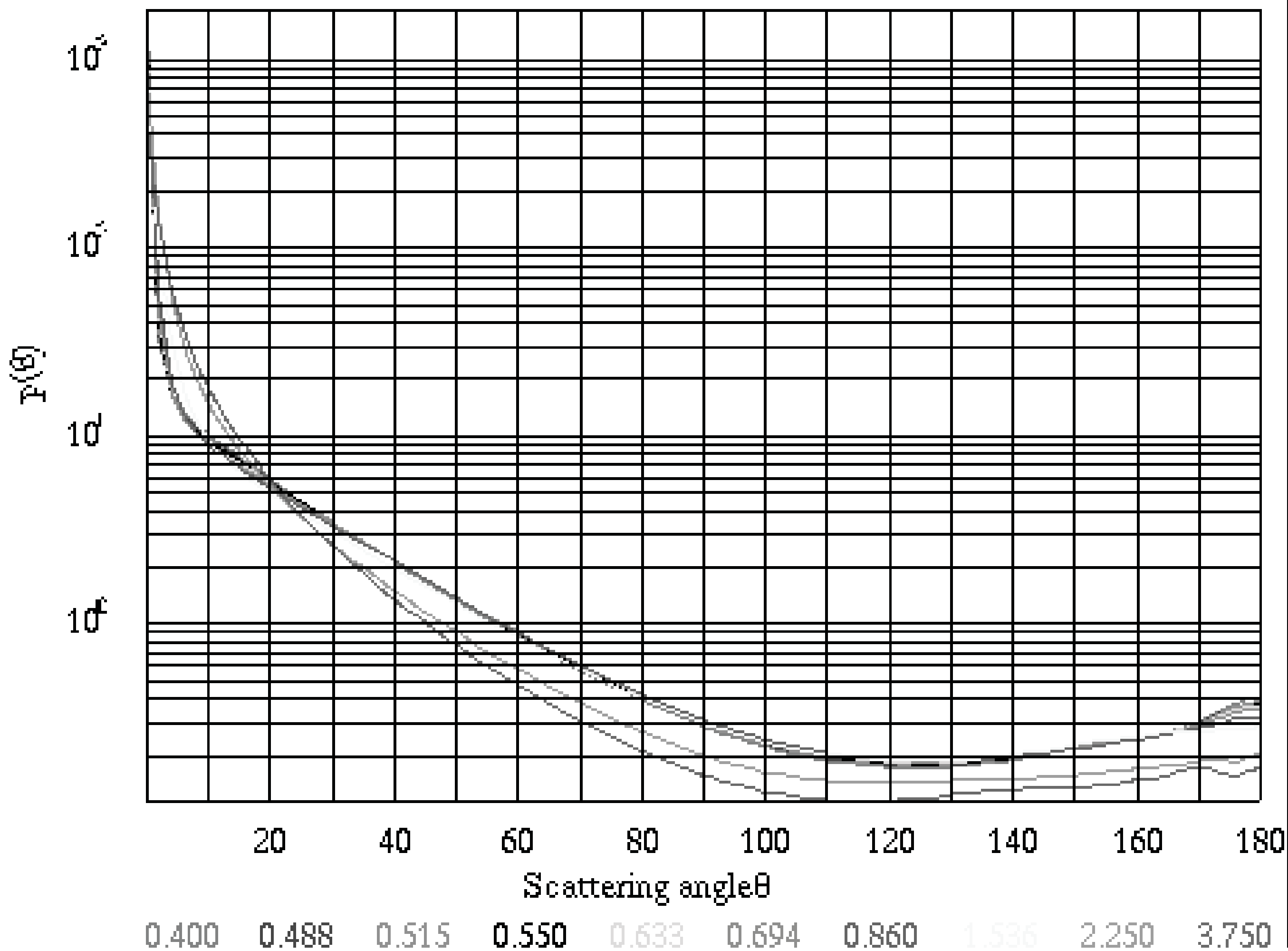
Aerosol scattering 1/5

- aerosol scattering refers to scattering by particles in suspension in the atmosphere (not molecules). The MIE scattering theory could be applied to compute the aerosol phase function and spectral optical depth, based on size distribution, real and imaginary index.



Aerosol scattering 2/5

Continental aerosol phase function



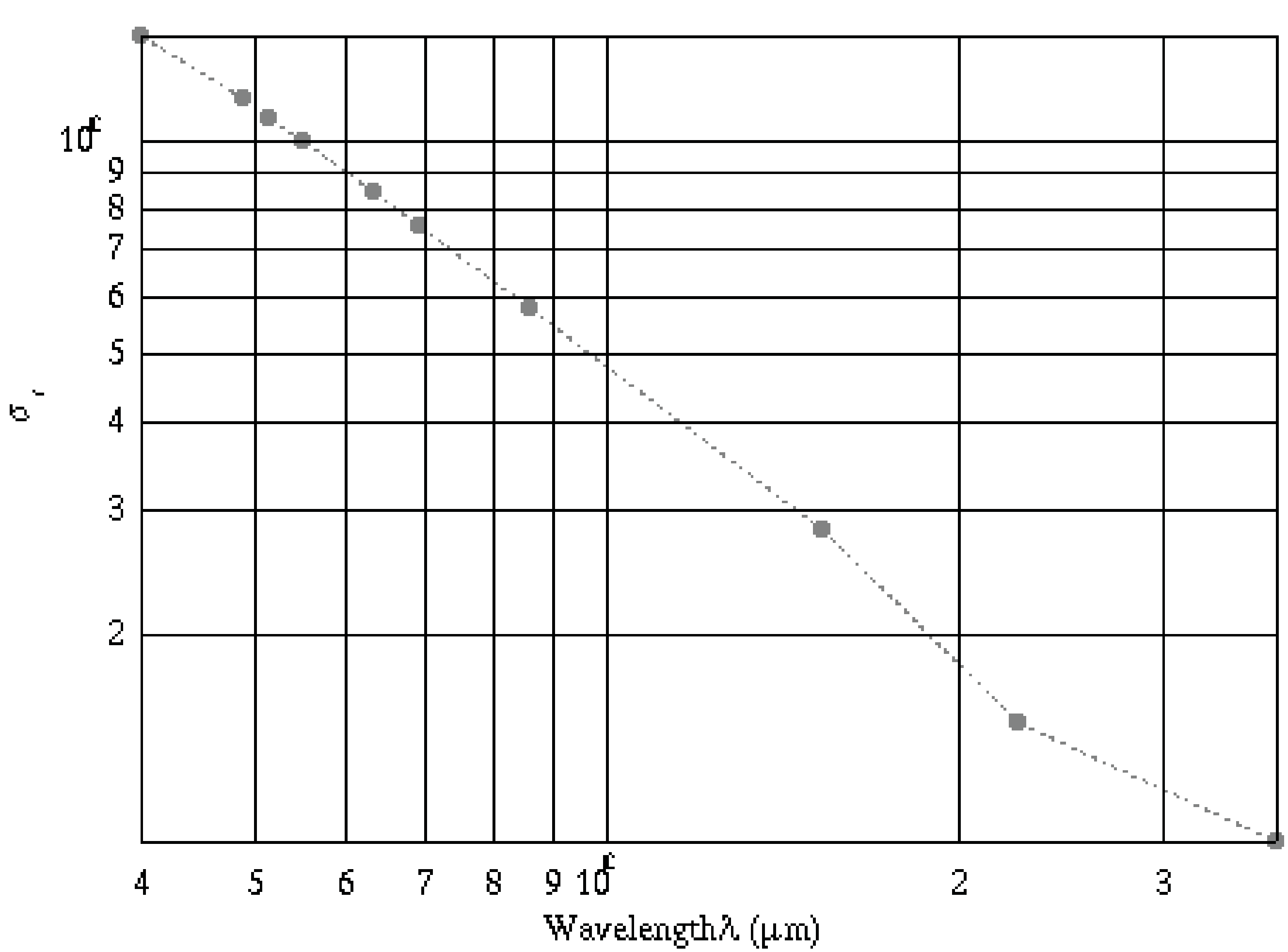
Aerosol scattering 3/5

single scattering albedo (0.2-1.0) to account for absorbing particles

$$\rho_A^\lambda(\theta_s, \theta_v, \phi) \sim \frac{\omega_A^\lambda \tau_A^\lambda P(\Theta)}{4\cos(\theta_s)\cos(\theta_v)}$$

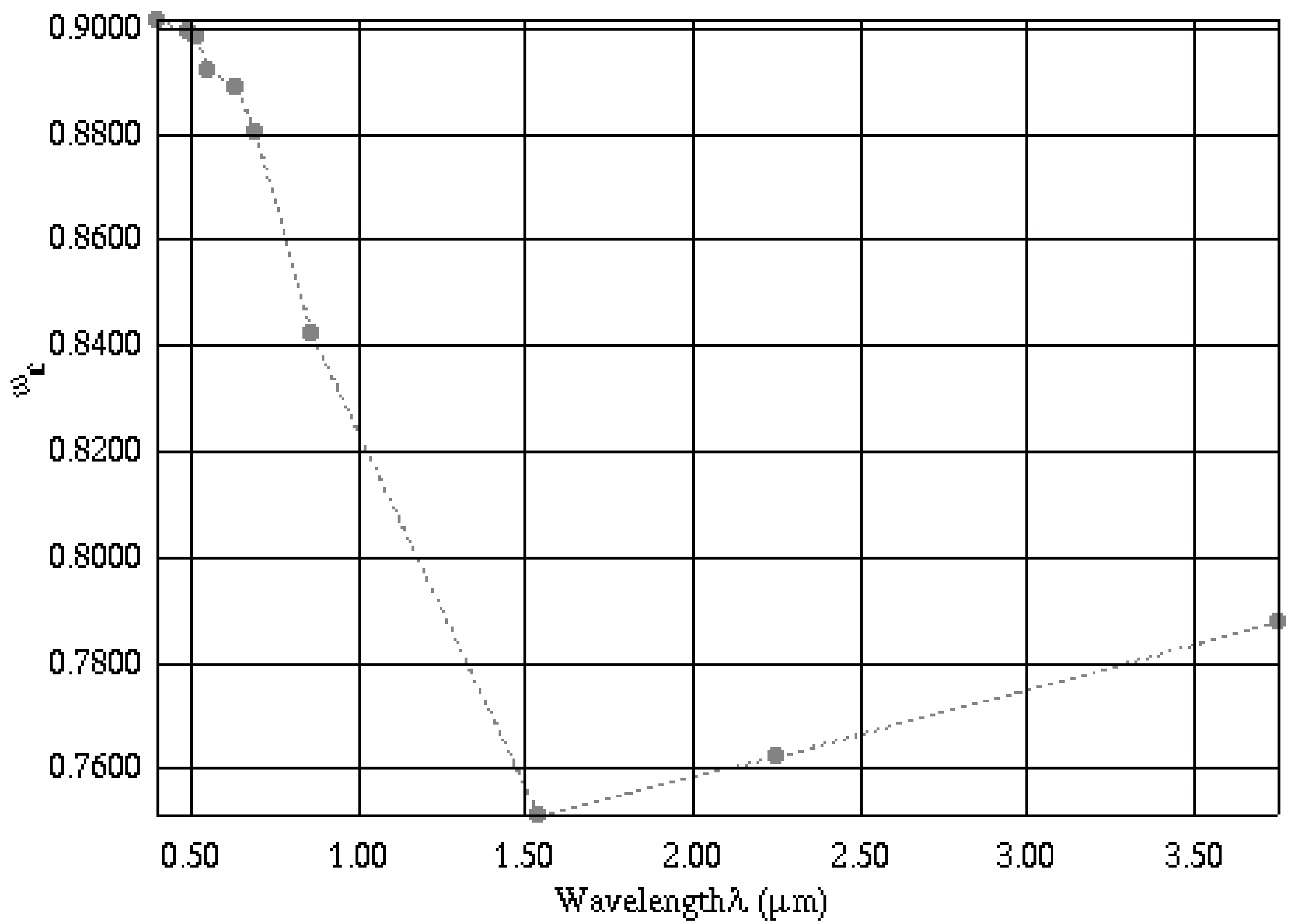
Aerosol scattering 4/5

Continental aerosol optical thickness spectral variation

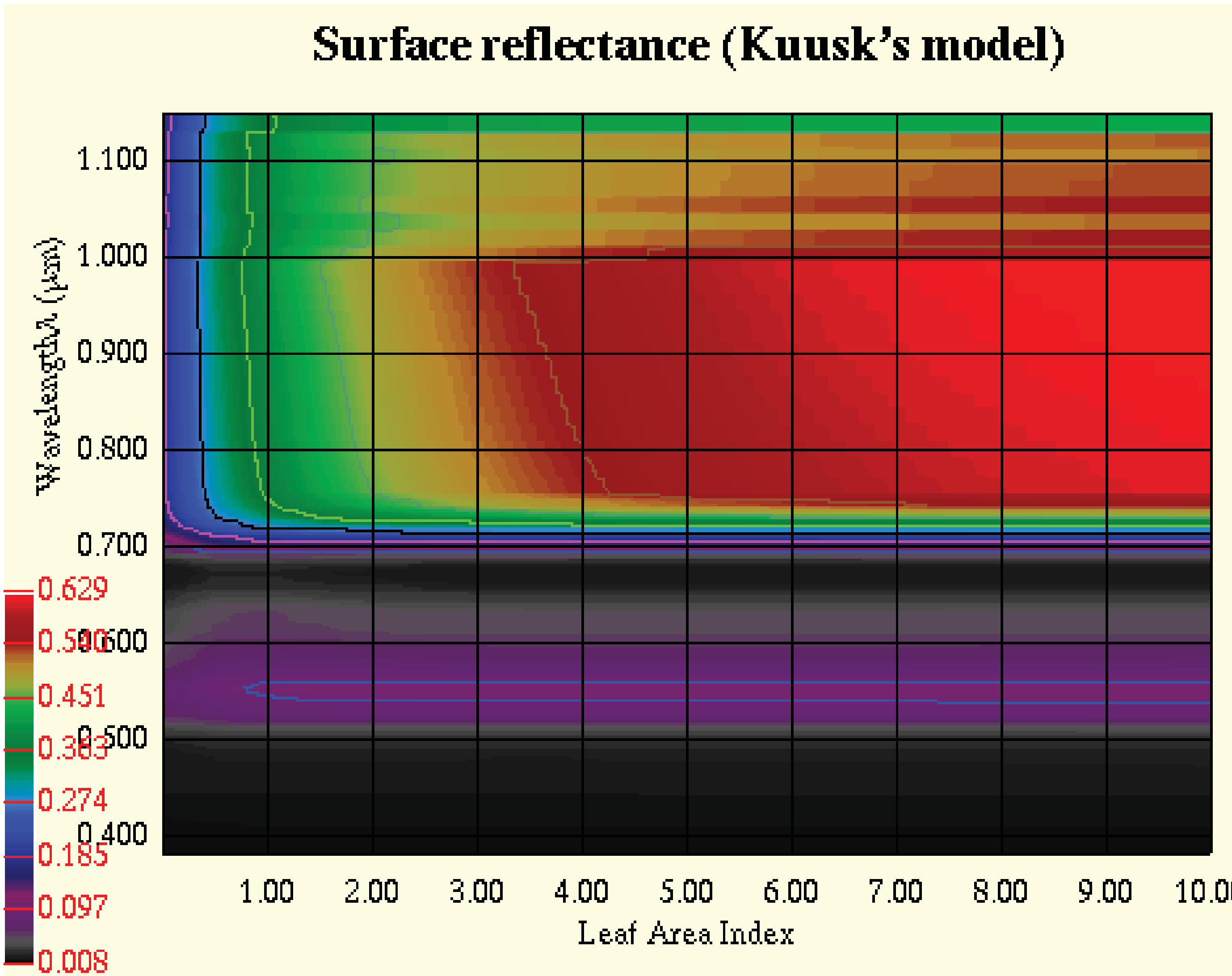


Aerosol scattering 5/5

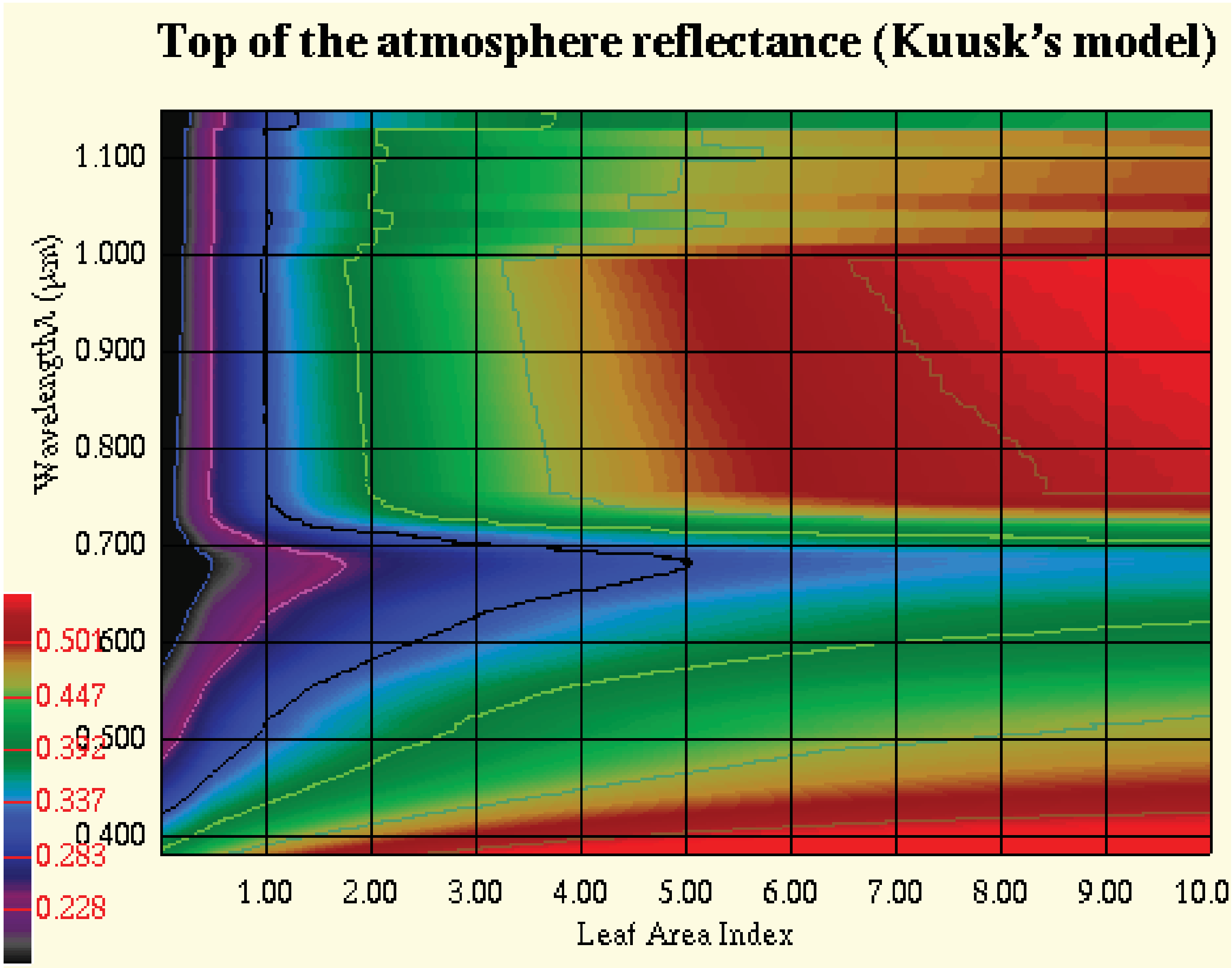
Continental aerosol single scattering albedo spectral variation



Atmospheric effect: Vegetation 1/3

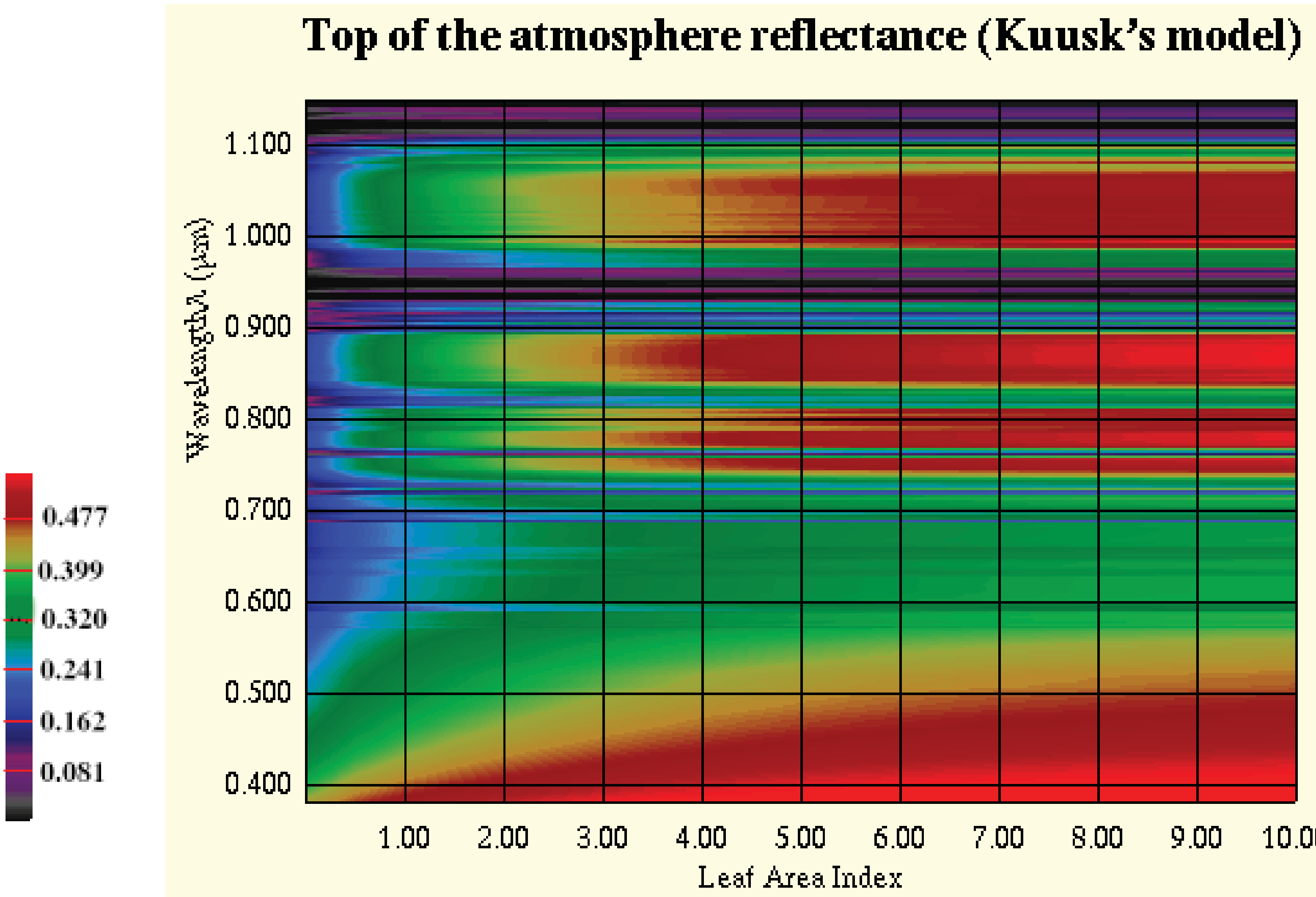


Atmospheric effect: Vegetation 2/3



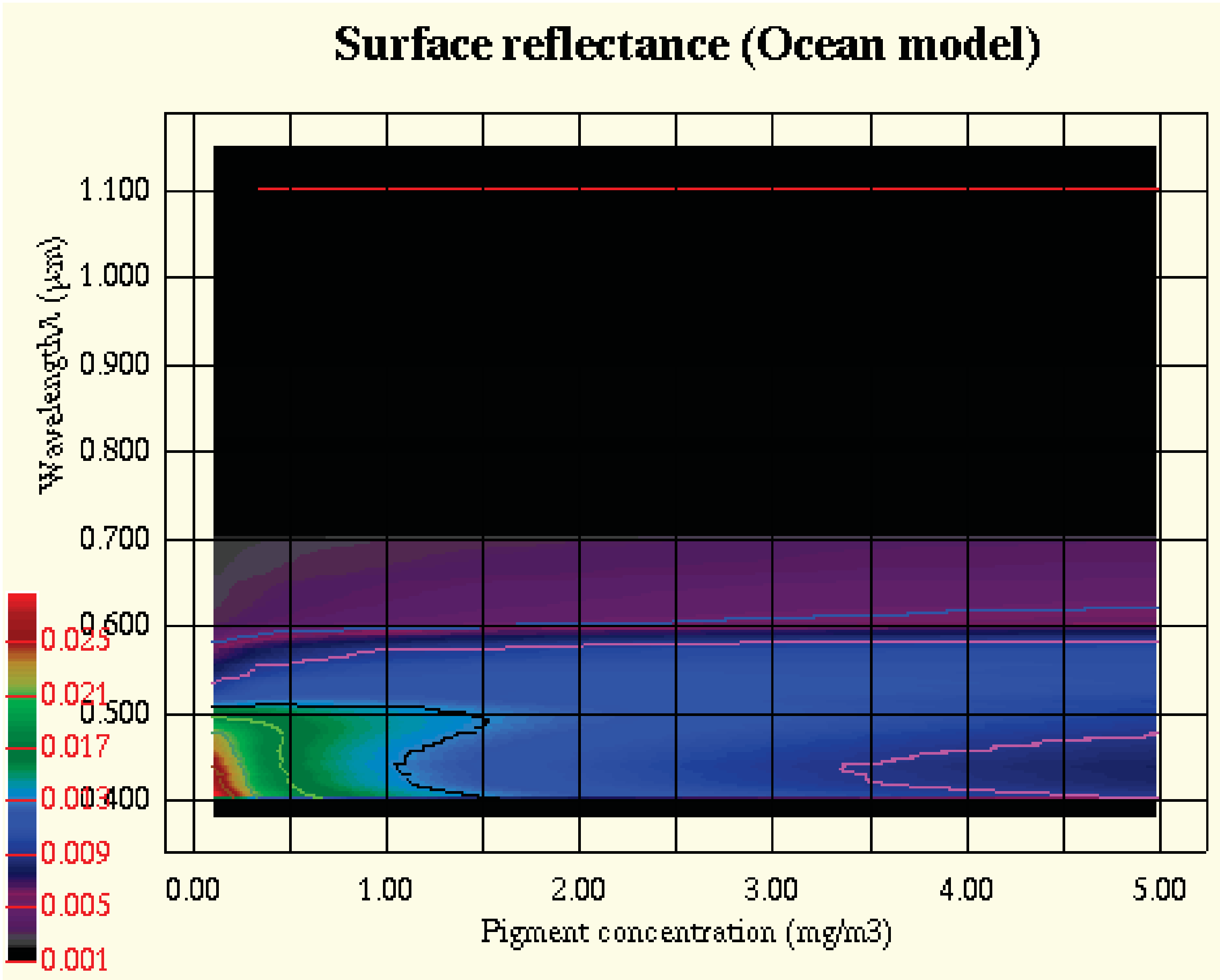
No absorption, Continental aerosol

Atmospheric effect: Vegetation 3/3

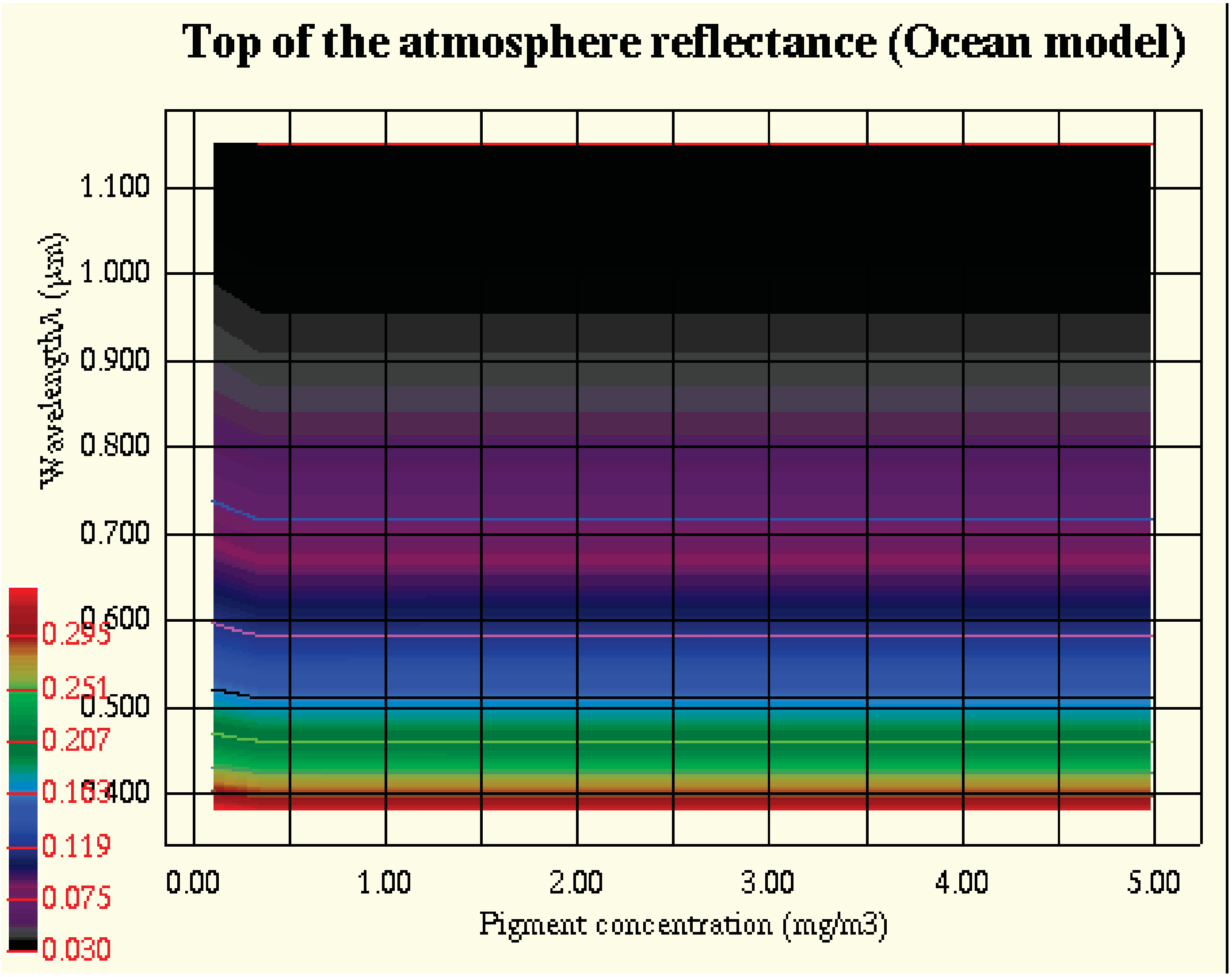


Absorption tropical atmosphere, Continental aerosol

Atmospheric effect: Ocean 1/2

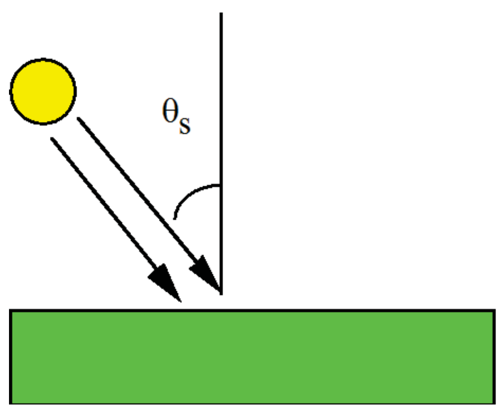


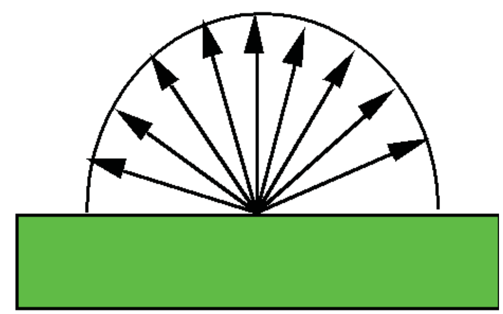
Atmospheric effect: Ocean 2/2



Perfect Lambertian Reflector

$$\int_0^{\pi}\int_0^{2\pi} RPLF(\theta_S, \theta, \phi)\cos(\theta)\sin(\theta)d\theta d\phi = E_s \cos(\theta_s)$$

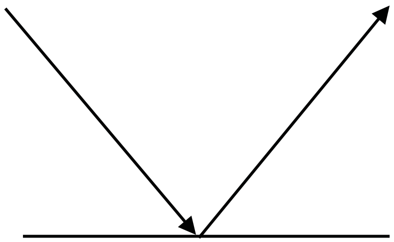




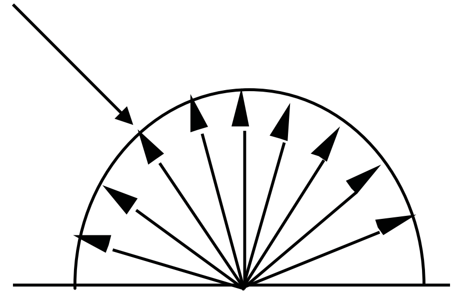
Isotropic radiation

$$RPLF(\theta_S, \theta_V, \phi) = \frac{E_S \cos(\theta_S)}{\pi}$$

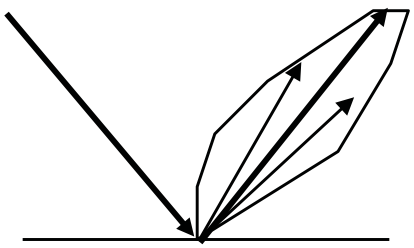
Different Types of Reflectors



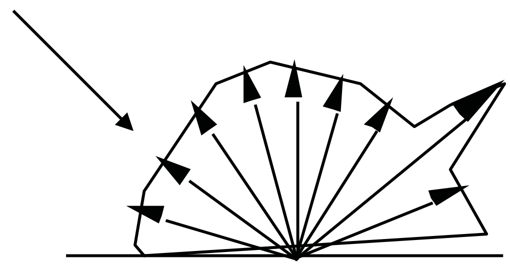
Specular reflector (mirror)



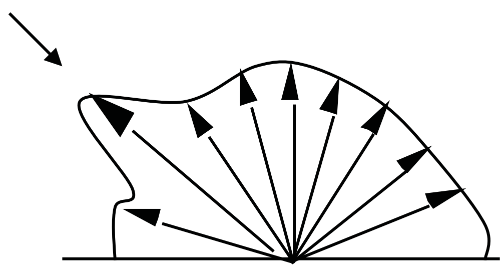
diffuse reflector (lambertian)



Nearly Specular reflector (water)

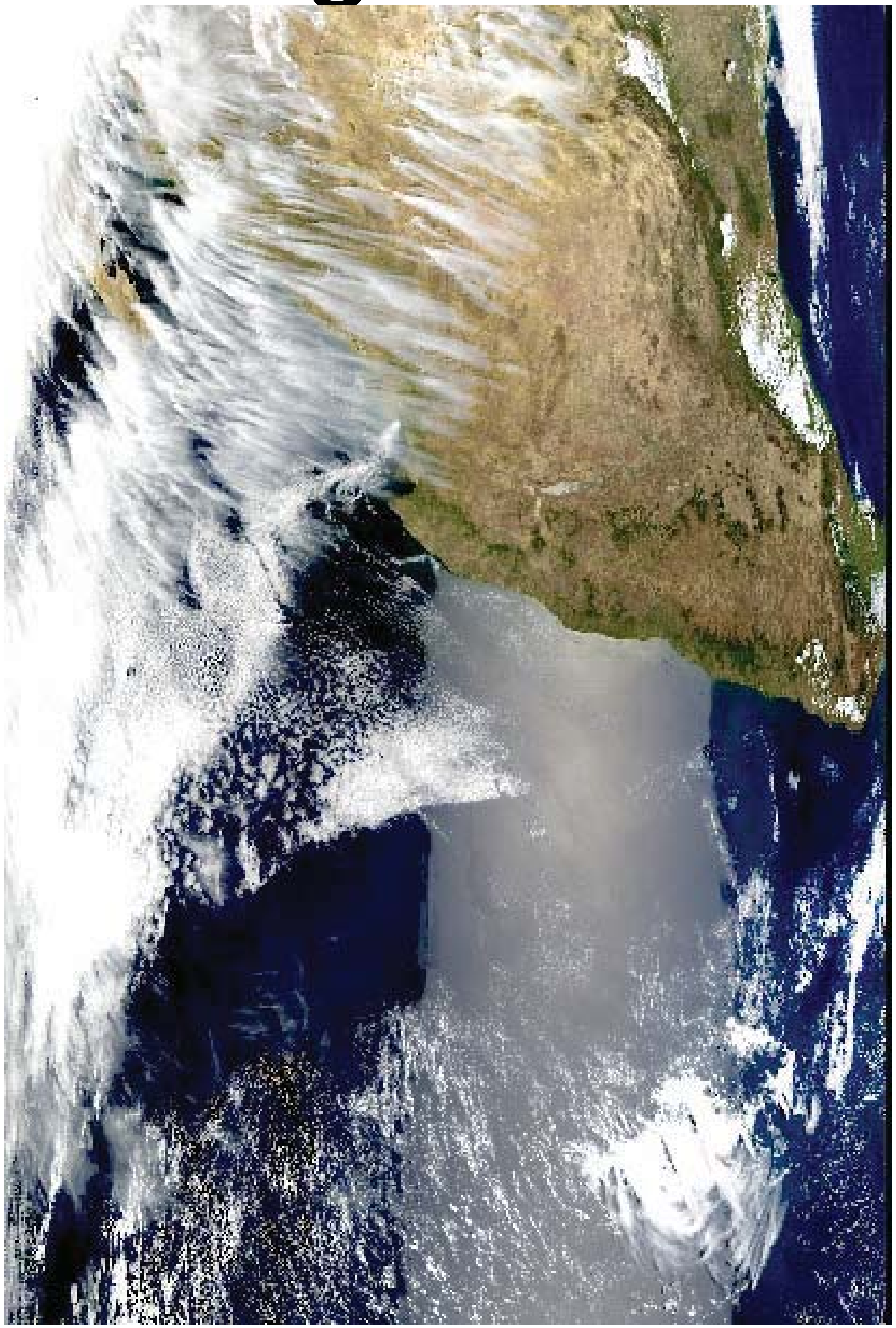


nearly diffuse reflector



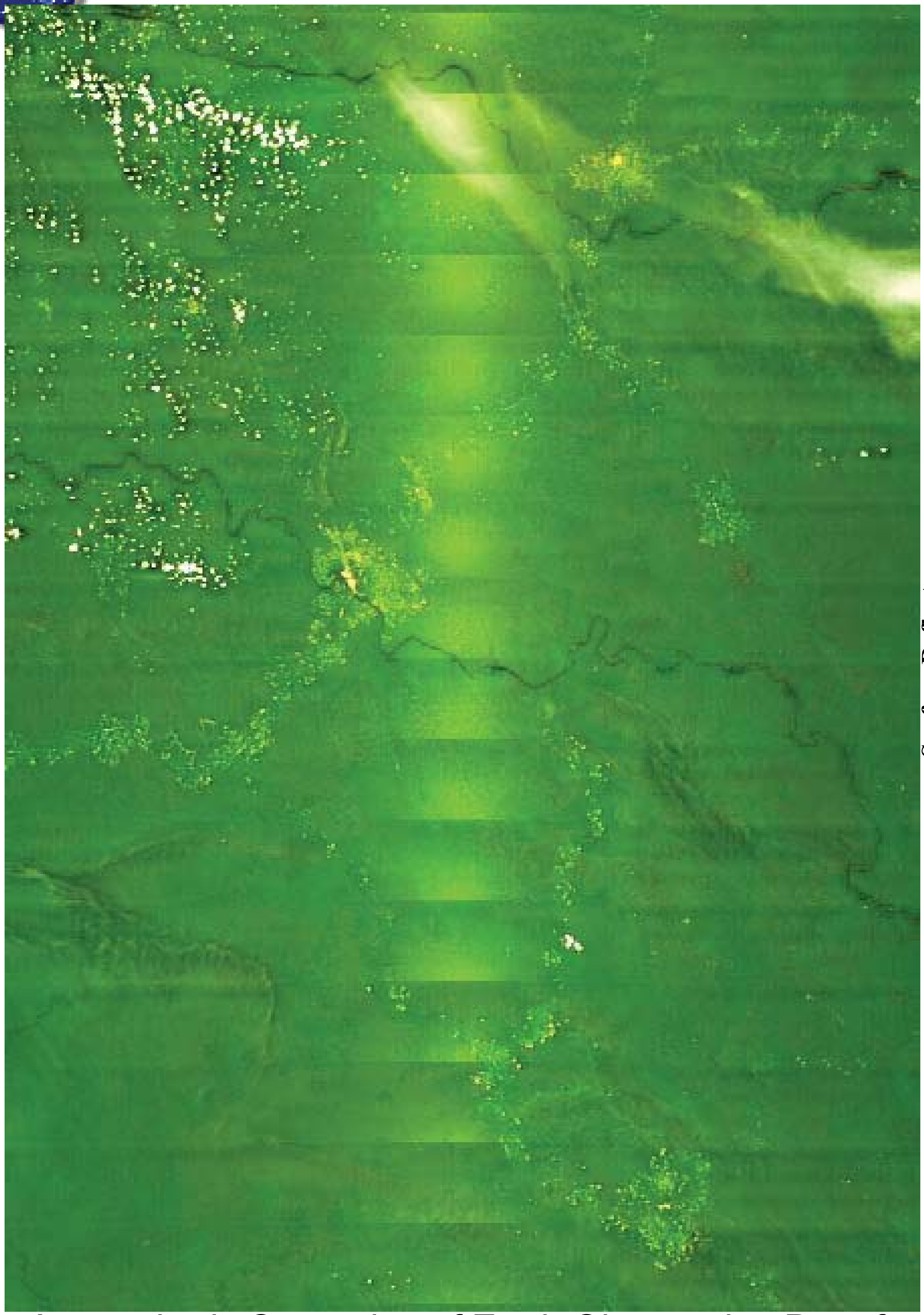
Hot spot reflection

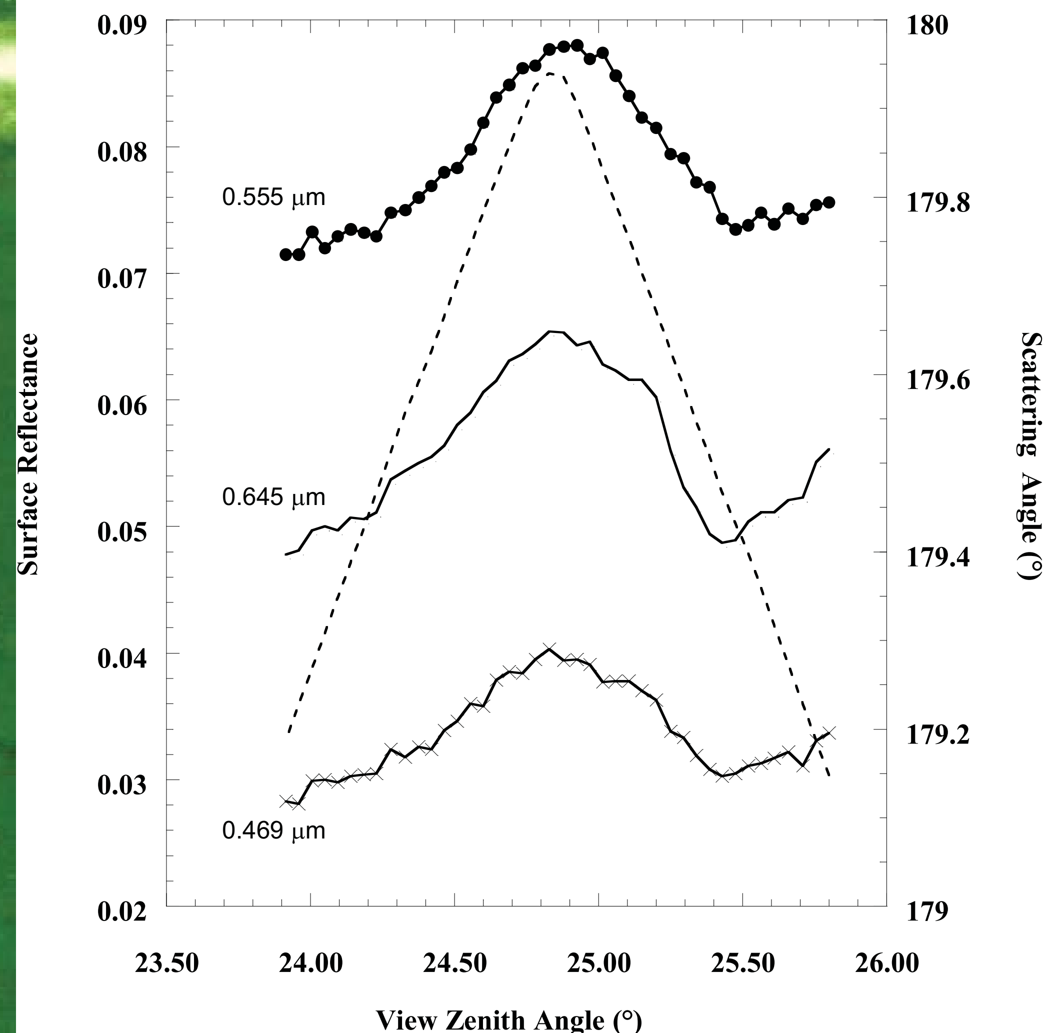
Sun glint as seen by MODIS



Gray level temperature image

MODIS data illustrating the hot-spot over dense vegetation





BRDF atmosphere coupling correction

Lambertian infinite target approximation

$$\rho_{app}\left(\theta_s,\theta_v,\phi\right)=\rho_{atm}\left(\theta_s,\theta_v,\phi\right)+T_{atm}(\theta_s)T_{atm}(\theta_v)\frac{\rho_{ground}}{1-\rho_{ground}S_{atm}}$$

BDRF atmosphere coupling approximation

$$\rho_{app}(\theta_s,\theta_v,\phi)=\rho_{atm}(\theta_s,\theta_v,\phi)+e^{-\tau/\mu_s}e^{-\tau/\mu_v}\rho_s(\theta_s,\theta_v,\phi)$$

$$+e^{-\tau/\mu_v}t_d(\theta_s)\overline{\rho}_s+e^{-\tau/\mu_s}t_d(\theta_v)\overline{\rho}'_s+t_d(\theta_v)t_d(\theta_s)\overline{\overline{\rho}}_s$$

$$+\frac{T_{atm}(\theta_s)T_{R+A}(\theta_v)S_{atm}\left(\overline{\overline{\rho}}_s\right)^2}{1-S_{atm}\overline{\overline{\rho}}_s}$$

$$\overline{\rho}_s(\mu_s,\mu_v,\phi)=\frac{\int_0^{2\pi}\int_0^1\mu L_{atm}^{\downarrow}(\mu_s,\mu,\phi')\rho_s(\mu,\mu_v,\phi'-\phi)\,d\mu\, d\phi'}{\int_0^{2\pi}\int_0^1\mu L_{atm}^{\downarrow}(\mu_s,\mu,\phi')\,d\mu\, d\phi'}$$

$$\overline{\overline{\rho}}_s(\mu_s,\mu_v,\phi)=\overline{\overline{\rho}'_s(\mu_s,\mu_v,\phi)}\qquad\qquad \overline{\rho}'_s(\mu_s,\mu_v,\phi)=\overline{\rho}_s(\mu_s,\mu_v,\phi)$$

Adjacency effect correction

Lambertian infinite target approximation

$$\rho_{app}(\theta_s, \theta_v, \phi) = \rho_{atm}(\theta_s, \theta_v, \phi) + T_{atm}(\theta_s) T_{atm}(\theta_v) \frac{\rho_{ground}}{1 - \rho_{ground} S_{atm}}$$

adjacency effect approximation

$$\rho_{app} = \rho_{atm} + \frac{T_{atm}(\theta_s)}{1 - S_{atm}\rho_e}\left(e^{-\tau/\mu_v}\rho_s + t_{atm}^d(\theta_v)\rho_e\right)$$

$$\rho_e = \frac{1}{2\pi}\int_0^{2\pi}\int_0^{\infty} \rho(r,\psi)\frac{dF(r)}{dr}dr\,d\psi$$

Adjacency effect correction (practical implementation)

$$\rho_i = \frac{\rho_{app} - \rho_{atm}}{T_{atm}(\theta_s) T_{atm}(\theta_v)}$$

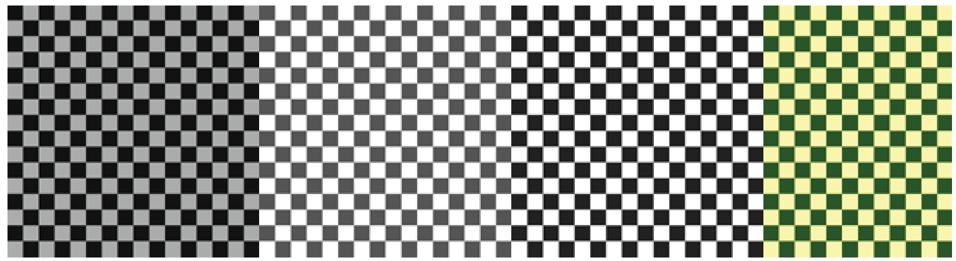
$$\rho_s^{\inf} = \frac{\rho_i}{1 + S_{atm}\rho_i}$$

$$\rho_i = \frac{\rho_s e^{-\tau/\mu_v} + t_{atm}^d(\theta_v)\rho_e}{T_{atm}(\theta_v)\left(1 - S_{atm}\rho_e\right)}$$

$$\rho_s = \frac{\rho_i T_{atm}(\theta_v)\left(1 - S_{atm}\rho_e\right) - t_{atm}^d(\theta_v)\rho_e}{e^{-\tau/\mu_v}}$$

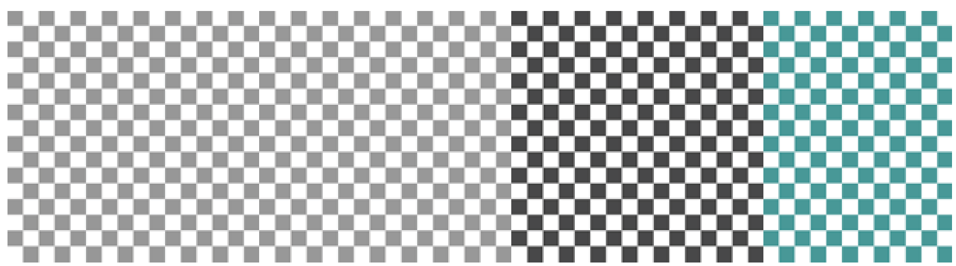
$$\rho_e = \sum_{j=-n}^{n}\sum_{i=-n}^{n} \frac{dF(r(i,j))}{dr}\rho_s^{\inf}(i,j)$$

Adjacency effect correction (testing)



(a) (b) (c) (d)

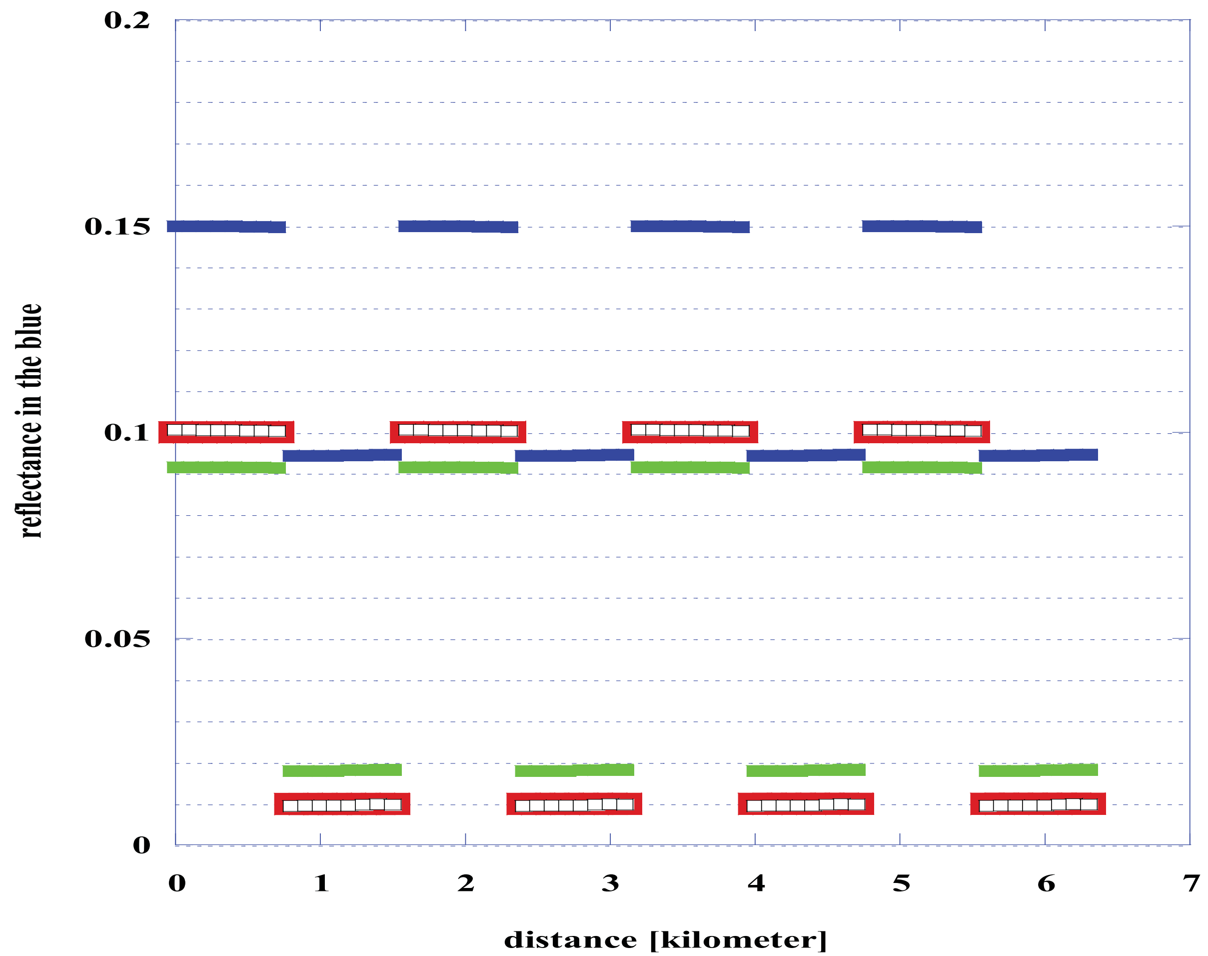
Synthetic data set for surface reflectance in the blue (a), green (b) and red (c), and in RGB (d) corresponding to bright soil (yellow squares) and dense vegetation (green squares).



(a) (b) (c) (d)

Typical atmospheric effect on the synthetic surface reflectance shown above

Adjacency effect correction (testing)



Reflectance's observed over a horizontal transect on the checkerboard. The red bars are the "true" surface reflectance, the blue bars correspond to the top of the atmosphere signal including adjacency effect. The green bars correspond to the corrected data using the infinite target assumption. The open square correspond to the data corrected for the adjacency effect using the operational method developed.

Adjacency effect correction (validation)