

# Water Transport and the Evolution of CM Parent Bodies

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# Introduction

- Meteorites have amino acids and hydrated minerals which constrain the peak temperature ranges they have experienced
  - CMs in particular have a narrow range (273-325K)
  - Bulk fluid motion during hydration constrained to small scales (<mm)
- Some asteroids are known to have hydrated minerals on their surfaces
- It is presumed these two facts may be related
- Problem:
  - hydration only occurs (significantly) with liquid water
  - melting water only occurs early on in nebula (1-10 Myrs ANC)
  - in nebula asteroid surface temperature very cold (~150K)
- Can indigenous alteration produce CMs and/or surface hydration?
  - Issue of timescales:  $t_{\text{diffuse}} < t_{\text{heat}} < t_{\text{conduct}} < t_{\text{vap}}$   
(last condition not true, so surface hydration probably not)

Fundamentally a 1D radial thermal diffusivity problem with heat source  $Q$  due to decays and reactions and a heat capacity modified to include latent heat of  $H_2O$ :

$$\frac{\partial T}{\partial t} = \frac{1}{r^2} \frac{\partial}{\partial r} \left( r^2 \kappa \frac{\partial T}{\partial r} \right) - \left( \frac{(\rho c_p)_l}{(\rho c_p)_{ave}} v_l + \frac{(\rho c_p)_g}{(\rho c_p)_{ave}} v_g \right) \frac{\partial T}{\partial x} + \frac{Q}{c_p}$$

where the liquid and gas velocity are found via two-phase Darcy flow:

$$v = \frac{-ks}{\mu} \left( \frac{\partial P}{\partial r} - \rho g \right)$$

Here  $k$  is the permeability (modified by a appropriate relative permeability),  $P$  is the relevant pressure,  $\mu$  is the viscosity,  $s$  is the volume fraction of liquid or gas in the voids (porosity,  $\epsilon$ ) given by  $f_{gas}/\epsilon$  or  $f_{liq}/\epsilon$ ,  $\rho$  is the gas or liquid density and  $g$  is the local acceleration due to gravity.

Flow is mostly driven by capillary pressure gradients:

$$P_{liq} = P_{gas} - P_{cap} = \frac{R_g T M_{gas,j} f_{gas}}{dV_j} - \sigma(T) J(s) \sqrt{\varepsilon/k}$$

where  $\sigma$  is liquid water surface tension,  $J$  is an empirical surface tension correction function,  $M_{gas,j}$  is the mass of water vapor and  $dV_j$  is the total volume of the  $j^{\text{th}}$  cell

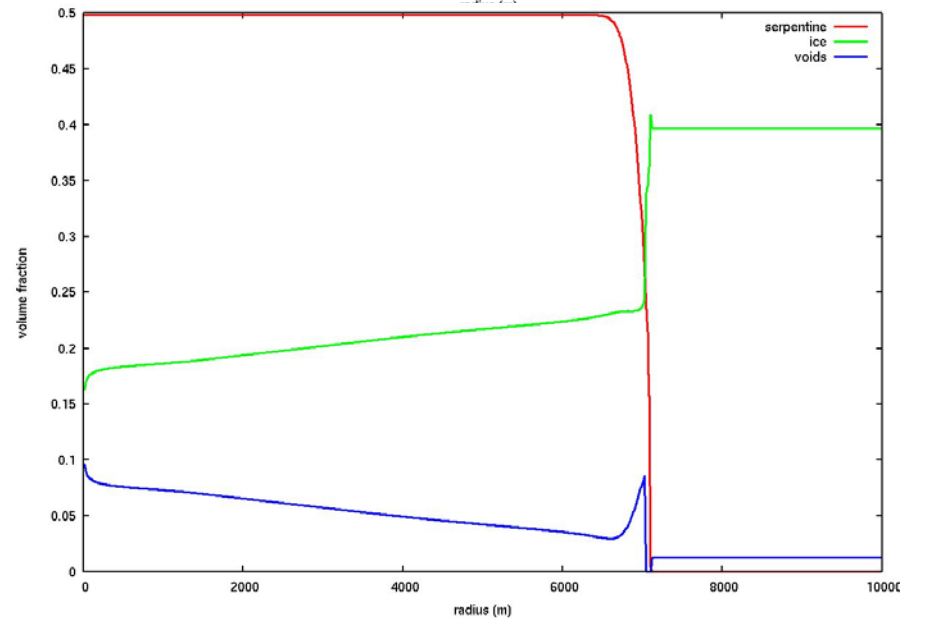
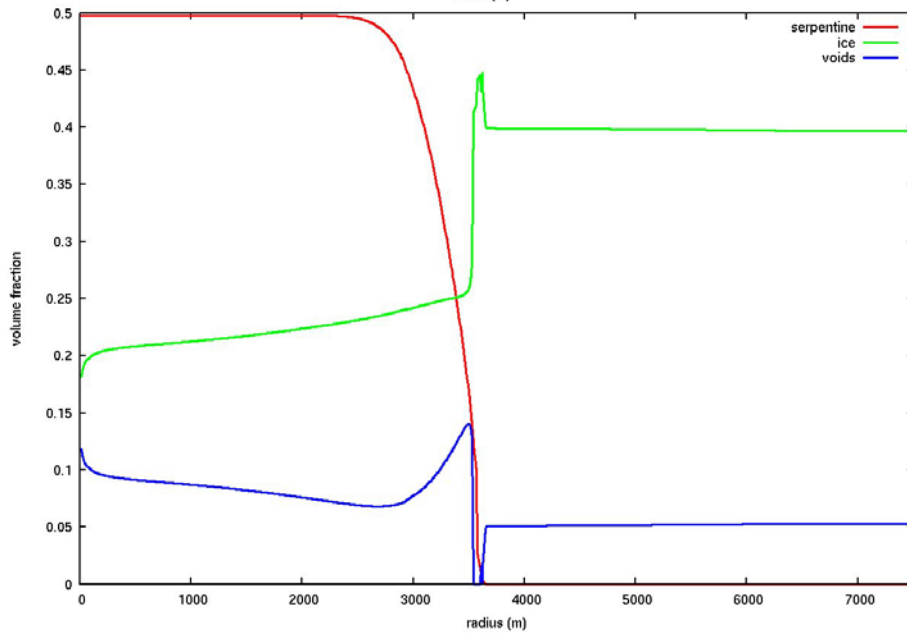
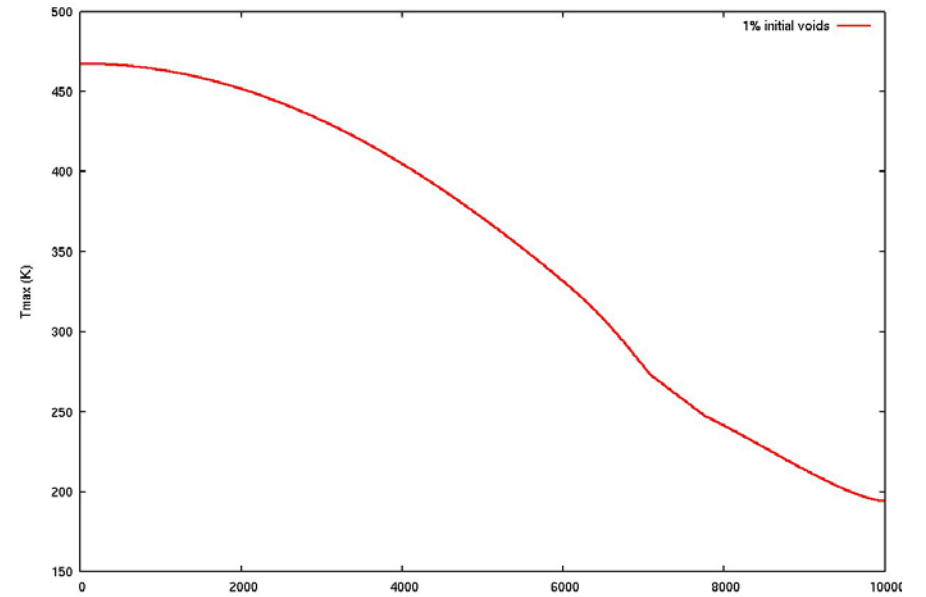
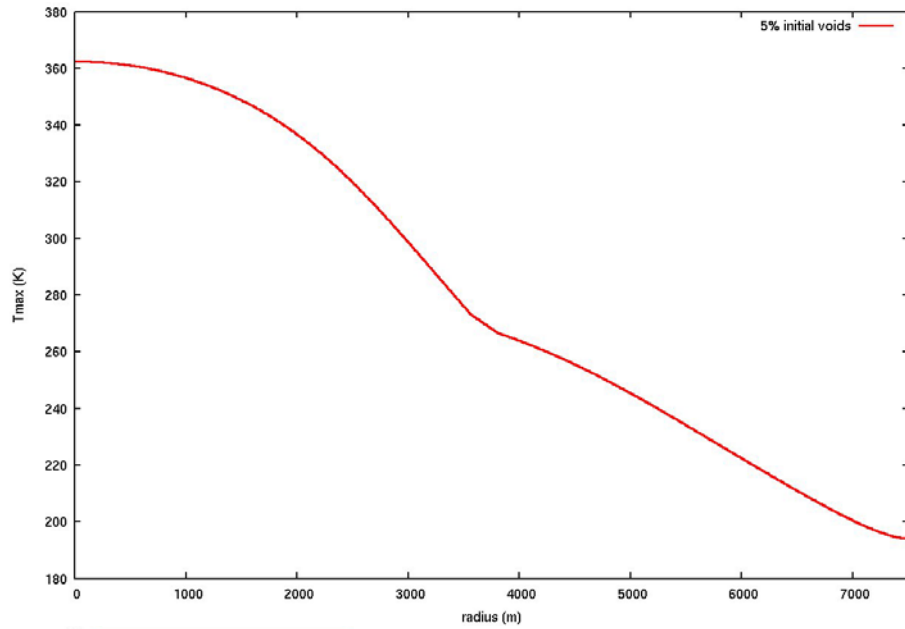
Permeability is given by:

$$k = \frac{\varepsilon r_H^2}{h_{ck}} = \frac{\varepsilon^3 d^2}{36 h_{ck} (1 - \varepsilon)^2}$$

Where  $h_{ck}$  is a geometry factor,  $r_h$  is the mean particle hydraulic radius, and  $d$  is a mean pore diameter

Constraints:  $r_h \sim 0.002\text{-}0.02 \mu\text{m}$  and  $d \sim 20\text{-}200 \text{ nm}$  and  $h_{ck} \sim 5$ ,  
 $\rightarrow k \sim 10^{-19}$  to  $10^{-17} \text{ m}^2$  (relative permeability reduces this)

# Sample of Results



# Takeaways

- Given nebula boundary condition and no pre-accretion hydration the PB:
  - 1) Cannot form less than 0.9 Myr ANC (or no hydration)
  - 2) Cannot be less than 3km in diameter (or no hydration)
  - 3) If forms at 10 Myr ANC, needs to be  $> \sim 180$  km in diameter
  - 4) Thus, minimum diameter is roughly  $20t-13$  km ( $t$  in Myrs ANC)
  - 5) Weak dependence on composition
  - 6) Initial Al (or e.g. Fe) just shifts timescales slightly
  - 7) Weak dependence on permeability
  - 8) Warmer nebula ( $\sim 200$ K) moves rind out and makes thicker
- Given peak temp seen by CMs is  $< 50^\circ\text{C}$  and CMs are  $\sim 50\%$  serpentine:
  - 1) Thin (few kms max) slice of larger ( $> \sim 12$  km diam) PB produces CMs
  - 2) Central  $< 1/6$  volume of smaller ( $< \sim 12$  km diam) PBs produces CMs
  - 3) Pre-accretion hydration just makes '12 km' threshold slightly bigger
  - 4) fills voids above reaction zone with ice (no venting or fracturing)
  - 5) No CMs near surface of PB

# Takeaways (cont.)

- Liquid water
  - 1) Persists for  $\sim 1$  Myr (20 km diam)
  - 2) Capillary flow (using canonical  $k$ ,  $rH$ , pores) peaks at  $\sim$ few m/Myr
  - 3) Transport scale of  $\sim$ m, 1000x too large (not all during reactions)
  - 4) Volume ratio transported  $\sim 1$  ( $\sim 0.1$  for small  $k$ ,  $rH$ ) during reactions
  - 5) More movement results in thicker CM 'rind' if not limiting reagent
  - 6) Sloshing occurs after reactions
  - 7) Low permeability leads to no convection
  - 8) 2<sup>nd</sup> reaction wave  $\sim$ always fills voids above reaction zone with ice
  - 9) Reduction in reaction rate 10x reduces peak T but moves rind inward

No combination gets:

- A CM zone within  $\sim 20\%$  of the surface by radius
  - Larger PBs closer to the surface
- A CM zone more than  $\sim 25\%$  thick by radius
  - Only small PBs get even this much and only in the center

IMPACTS.....