

Glycine's Radiolytic Destruction in Ices: First *in situ* Laboratory Measurements for Mars

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Abstract

We report new laboratory studies of the radiation-induced destruction of glycine-containing ices for a range of temperatures and compositions that allow extrapolation to martian conditions. *In situ* infrared spectroscopy was used to study glycine decay rates as a function of temperature (from 15 to 280 K) and initial glycine concentrations in six mixtures whose compositions ranged from dry glycine to H₂O+glycine (300:1). Results are presented in several systems of units, with cautions concerning their use. The half-life of glycine under the surface of Mars is estimated as an extrapolation of this data set to martian conditions, and trends in decay rates are described as are applications to Mars' near-surface chemistry. Key Words: Mars—Ice—Radiolysis—*In situ* measurement—Laboratory investigations. *Astrobiology* 13, 647–655.

1. Introduction

THE FORMATION AND SURVIVAL of molecules in extraterrestrial environments is a subject of considerable interest to the astronomical community. In particular, it would be valuable from the standpoint of astrobiology to know how organic molecules withstand the conditions of high radiation flux and low temperature found in the surface layers of many planets and satellites. Cold planetary near-surface environments could be reservoirs of carbonaceous materials from past impact events or could hold the chemical evidence of past biological systems. It is possible that chemical tracers might decay over the course of time when exposed to radiation, so it is important to determine the radiolytic lifetimes of these molecules when planning missions to other worlds. Understanding the survivability of biological molecules such as amino acids also is valuable to the field of planetary protection, where this information may aid in the determination of the levels of cross contamination of terrestrial and potential extraterrestrial biospheres.

Despite the fact that spacecraft have orbited, landed on, and sampled Mars since the 1960s, large gaps remain in what we know about molecular evolution under martian conditions. For instance, the fate of organic molecules delivered by meteorite impacts and their interactions with the oxidants in martian soil has been the subject of speculation for years (Benner *et al.*, 2000). The relatively high costs of Mars missions have meant that heavy reliance is placed on laboratory measurements and theoretical models to unravel the thermal (Marion *et al.*, 2010), photo- (Stalport *et al.*, 2009), and radiation

(Kminek and Bada, 2006) chemistries of suspected martian molecules. It remains a challenge to design such measurements and models so that they are, to the extent possible, realistic replicas of martian processes and conditions.

Among the external drivers of martian chemistry are photons and particle radiation, both of which readily penetrate the martian CO₂ atmosphere (Molina-Cuberos *et al.*, 2001). The development and application of models of martian chemistry have been hindered by the fact that the radiation environment on Mars has not been measured directly, with the exception of preliminary results from the Mars Science Laboratory Curiosity rover (Hassler *et al.*, 2012). Despite the dearth of radiation data from the surface of the planet, the three processes shown in Fig. 1 appear to be agreed upon as the sources of energy for radiation chemistry on Mars at various depths (Dartnell *et al.*, 2007). Molecules on or within a micrometer to a millimeter of the martian surface are thought to be altered significantly by far-UV solar photons (Stoker and Bullock, 1997; ten Kate *et al.*, 2005). Starting at a few millimeters and extending to a few centimeters, perhaps a meter, solar energetic particles are the dominant source of ionizing radiation (Kminek and Bada, 2006). From at least a meter to several meters, galactic cosmic rays dominate, but their penetration and that of the secondary electrons they produce fall off with depth. By about 4–5 m, the greatest source of radiation in the martian subsurface is from radioactive materials in the soil (Dartnell *et al.*, 2007).

The centimeter-to-meter depths below the martian surface are where conventional radiation chemistry controls molecular change. Theoretical models have been published to aid

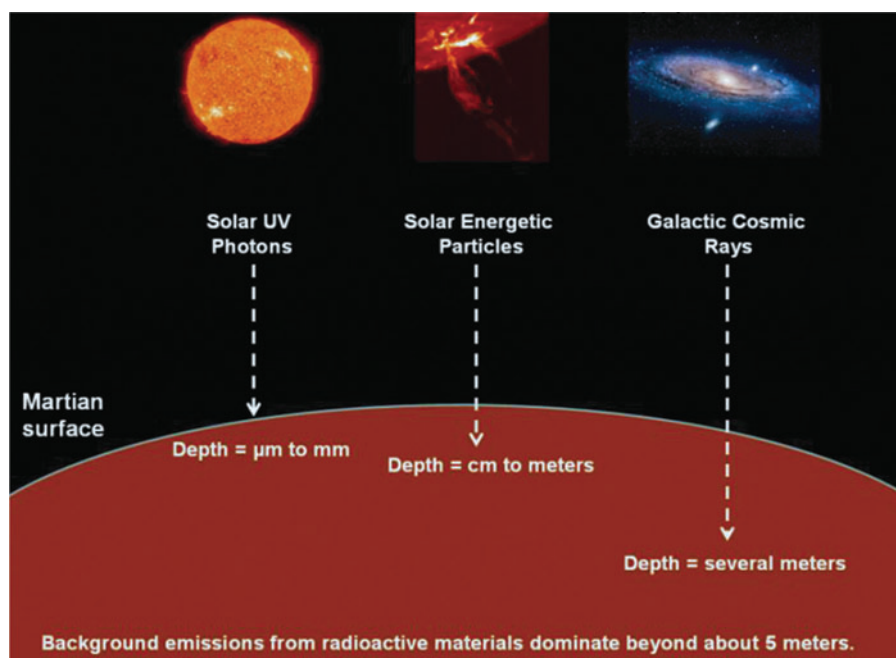


FIG. 1. Sources of radiation in the near-surface environment on Mars.

in understanding molecular destruction at these depths (Dartnell *et al.*, 2007), but in practice these usually are based on room-temperature laboratory data, studies of single-component samples, and *ex situ* methods of chemical analysis. While such models and laboratory investigations contribute to a “baseline understanding” of subsurface radiation chemistry on Mars, experiments on solid-phase mixtures at lower temperatures are also needed, along with *in situ* analytical methods. Somewhat remarkably, few such experiments were performed before the Viking landers’ martian investigations, and few have been published in the 40 years since.

Our research group has used IR spectroscopy and ionizing radiation to study the formation and destruction of organic molecules in ices for many years. In our most recent study, we examined the radiolytic destruction of three amino acids at 15–140 K in the presence and the absence of H₂O-ice (Gerakines *et al.*, 2012). Decay parameters were determined for only two mixtures (dry glycine and H₂O+gly 8.7:1), and lifetimes were estimated for various astronomical environments based on extrapolations of this limited data set. In the present paper, we report new and more extensive *in situ* laboratory measurements of the rate of radiation-induced destruction of the simplest amino acid, glycine (H₂N-CH₂-COOH, abbreviated “gly”), over a larger range of tempera-

tures (15–280 K) and concentrations (0:1 to 300:1). Our aim is to report these new data that allow extrapolation to martian conditions and to illustrate the caution needed in describing radiolytic stabilities.

2. Experimental Methods

The methods used were similar, and in most cases identical, to those of our previous work (Gerakines *et al.*, 2012). Briefly, reagent-grade glycine was sublimed from a Knudsen-type furnace onto a polished aluminum substrate (area $\approx 5\text{ cm}^2$) that was held at the desired temperature within a vacuum chamber ($\sim 10^{-7}$ torr) interfaced to an IR spectrometer. Spectra of ices on the substrate were recorded as 100-scan accumulations at 4 cm^{-1} resolution from 5000 to 700 cm^{-1} with a Thermo Nicolet Nexus 670 spectrometer. A closed-cycle helium cryostat (ARS DE-204) and a resistive heater served to maintain the glycine-containing ice at the desired temperature within the 15–280 K range studied. In most experiments, H₂O vapor was admitted into the vacuum chamber to condense simultaneously with glycine onto the precooled substrate. By careful choices of flow rates, H₂O:glycine ratios from 300:1 down to 0:1 could be attained. A typical H₂O sample growth rate was $5\text{--}10\text{ }\mu\text{m h}^{-1}$. The background pressure inside the vacuum chamber was

TABLE 1. PROPERTIES OF 0.8 MeV PROTONS AND AMINO ACID ICE SAMPLES

Ice's H ₂ O:glycine ratio	Thickness (h) μm	Stopping power (S) $\text{eV cm}^2 \text{g}^{-1} \text{p}^{+1}$	Density (ρ) g cm^{-3}	Average molecular mass (m_{avg}) g molecule^{-1}
0:1	0.9	2.78×10^8	1.61	1.25×10^{-22}
8.7:1	1.8	2.86×10^8	1.06	3.85×10^{-23}
50:1	3.1	2.88×10^8	1.01	3.17×10^{-23}
100:1	3.0	2.89×10^8	1.01	3.08×10^{-23}
200:1	3.0	2.89×10^8	1.00	3.04×10^{-23}
300:1	4.0	2.89×10^8	1.00	3.02×10^{-23}

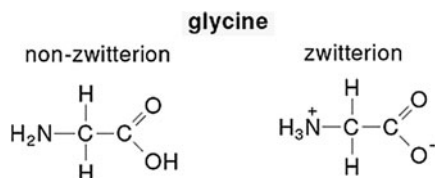


FIG. 2. Molecular structures of glycine.

typically $\sim 10^{-7}$ torr, resulting in a residual H_2O condensation rate of $\sim 0.05 \mu\text{m h}^{-1}$. Ice irradiations were performed with 0.8 MeV protons from a Van de Graaff accelerator, a beam current of $\sim 0.1 \mu\text{A}$, and a fluence of $1 \times 10^{14} \text{ p+cm}^{-2}$ delivered in about 15 min. Direct measurement of the beam current in the underlying aluminum substrate allowed radiation doses to be calculated by using reasonable assumptions for each sample's density (Weast *et al.*, 1984) and stopping power (Ziegler *et al.*, 2010). See Table 1 for the values used in this work. Sample thicknesses were on the order of 1–4 μm , measured by 650 nm laser interferometry. See the work of Gerakines *et al.* (2012) for additional experimental details related to amino acid sublimation, characterization, and irradiation.

It should be emphasized that the design of our experiments allowed amino acid decay data to be collected *in situ* at the temperature of each irradiation. Raising the samples to room temperature for chemical analyses, as is commonly the case, was therefore not necessary. Also, the sublimation and co-condensation method we selected for preparing H_2O +glycine ices avoided the uncertainties in freezing room-temperature aqueous glycine solutions, which will crystallize to give regions of varying glycine concentration in the sample. In short, the value of our results is enhanced both by the way the samples were prepared and by the method of analysis.

3. Results

This paper presents results for solids containing a single amino acid, glycine. Aside from its simplicity, glycine was selected as it is the only amino acid found both in a comet's tail and in meteorites (Degens and Bajor, 1962; Elsila *et al.*,

2009). Also, our earlier work (Gerakines *et al.*, 2012) showed that glycine's rate of radiolytic destruction in H_2O mixtures is somewhat greater than that for more-complex amino acids, so that glycine's behavior could be taken as a limiting case.

Our choice of H_2O -ice was suggested by its possible subsurface presence on Mars and its existence at the martian polar caps. The high-temperature limit of the experiments we report here with H_2O +glycine ices was determined by H_2O -ice's rapid sublimation above about 160 K and its tendency to crystallize above about 155 K, two effects we wished to avoid. The other prominent martian ice, frozen CO_2 , has a more limited range of temperatures for stability in our vacuum system. Nevertheless, future experiments are envisioned with amino acids in solid CO_2 .

Glycine has three common polymorphs (crystalline forms), but the broad IR features of our samples, both with and without H_2O -ice, suggested that all our ices were amorphous. The molecule also has two structures, shown in Fig. 2. Condensation of pure glycine at 140 K gave the zwitterion form of the molecule, while condensation near 15 K gave the non-zwitterion structure. These forms were confirmed in our samples by comparing our measured IR spectra to those found in the literature (*e.g.*, Maté *et al.*, 2011). Warming to about 100 K of any sample of pure glycine deposited near 15 K gave a rapid conversion to the zwitterionic form of the molecule, and samples retained their zwitterionic structure upon recooling. All ices to be irradiated were deposited at 140 K in order to create the zwitterionic structure for all experiments, as it is what is expected at martian temperatures.

Representative IR spectra of irradiated samples are given in Figs. 3 and 4. Figure 3 shows spectra of glycine at 15 K before and after proton irradiation to a final fluence of $8 \times 10^{14} \text{ p+cm}^{-2}$, while Fig. 4 shows the spectra from a similar experiment for a H_2O +glycine (50:1) mixture at 15 K. Doses are given in the figures in units of electronvolts per molecule. See the work of Gerakines *et al.* (2012) and references therein for spectral assignments in the IR.

By integrating the glycine COO^- vibrational feature at 1408 cm^{-1} before and after successive irradiations, data were collected for the decay curves shown in Figs. 5 and 6. In both figures, the left-hand panel displays results for four H_2O :gly mixture ratios (0:1, 8.7:1, 50:1, and 300:1) irradiated at 15 K,

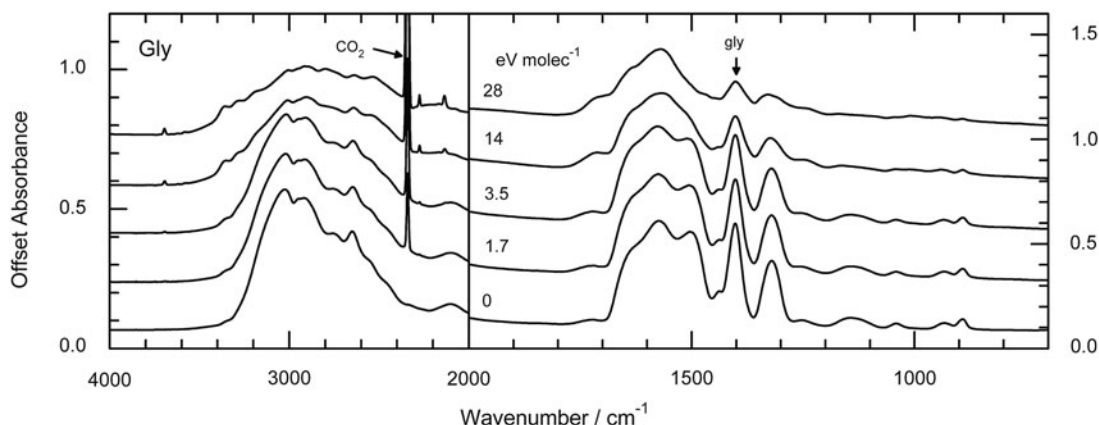


FIG. 3. Infrared spectra of glycine before and after irradiation at 15 K. The 2340 cm^{-1} feature of the radiation product CO_2 and the 1408 cm^{-1} feature of glycine are indicated.

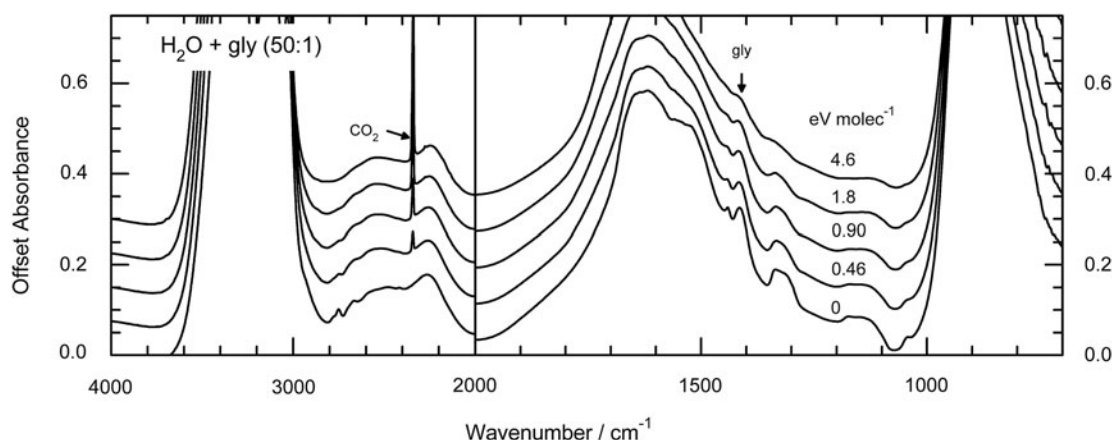


FIG. 4. Infrared spectra of $\text{H}_2\text{O} + \text{gly}$ (50:1) before and after irradiation at 15 K. The 2340 cm^{-1} feature of the radiation product CO_2 and the 1408 cm^{-1} feature of glycine are indicated.

while the right-hand panel shows the data for the 50:1 mixture irradiated at 15, 100, and 140 K. In Fig. 5, the horizontal scale is energy absorbed per gram of sample, found by multiplying proton fluence (F , in p+cm^{-2}) and stopping power (S , in $\text{eV cm}^2 \text{ g}^{-1} \text{ p}^{-1}$) for each ice. The radiation dose is given in megagray units along the upper horizontal scale, where $1 \text{ MGy} = 6.242 \times 10^{21} \text{ eV g}^{-1}$. Doses in the older, but still common, unit megarad (Mrad) can be obtained from the relation $1 \text{ MGy} = 100 \text{ Mrad}$. Although doses in the radiation-chemical literature commonly are given in eV g^{-1} , MGy, and Mrad units, it is not unusual to find doses in the *astrochemical* literature in units of eV molecule^{-1} , calculated from the product $m_{\text{avg}} SF$, where m_{avg} is the average mass (in g) of the ice sample's molecules. Since m_{avg} varies among the ices we studied, the relative positions of decay curves on an eV molecule^{-1} scale will differ from those on an eV g^{-1} (or Mrad or MGy) scale, as can be seen by comparing the curves of Figs. 5 and 6. For example, the decay shown in the $\text{H}_2\text{O} +$

gly (8.7:1) ice sample at 15 K (left-hand panels) is similar to that of the 0:1 sample in Fig. 5; but in Fig. 6 the differing horizontal scales, determined by two very different values of m_{avg} , cause the 8.7:1 sample to display a much steeper slope. However straightforward the relationship between the two horizontal scales may seem, it is important to point out differences due to inconsistencies in the way that relative decay rates are reported in the literature and to stress that the units do matter. They are determined in units equivalent to eV g^{-1} by some (e.g., Kminek and Bada, 2006) and in eV molecule^{-1} by others (e.g., Gerakines *et al.*, 2000), but the two do not necessarily reveal the same relative trends. Tables 2 and 3 give half-life doses measured in our experiments, denoted $D_{1/2}$ and defined as the dose required for the destruction of one-half of each sample's glycine molecules.

The initial density, sample area, and thickness of each ice we investigated could be measured directly by us or approximated from the mixing ratio and the known values for

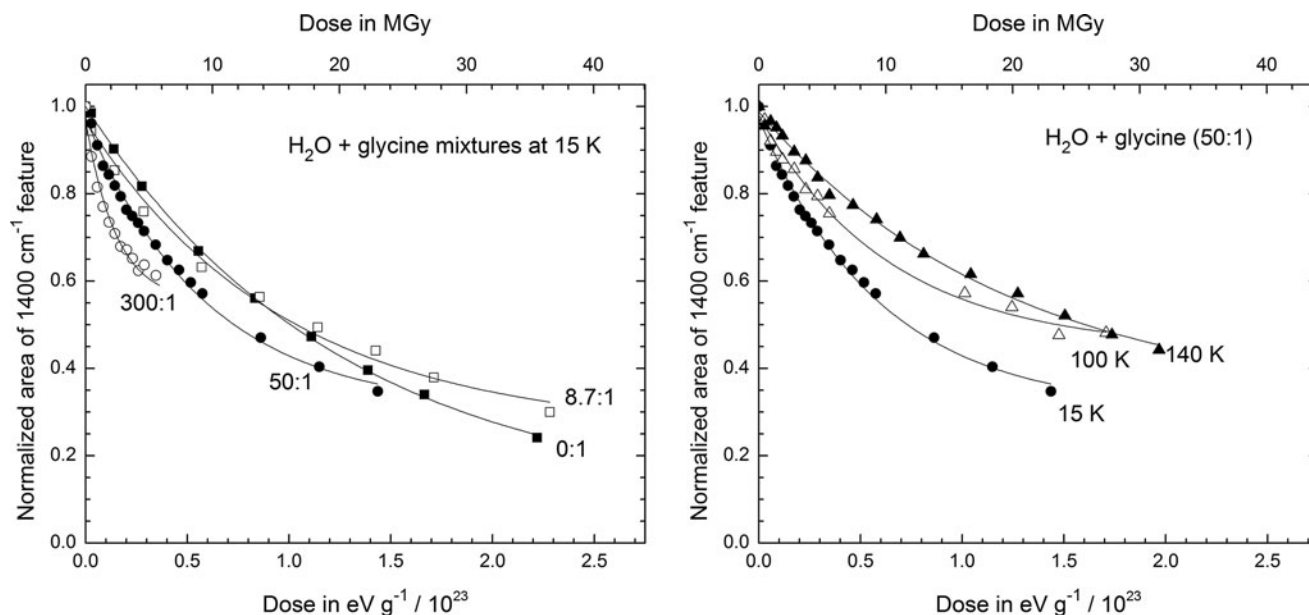


FIG. 5. Area of the glycine feature near 1400 cm^{-1} during irradiation for selected ice compositions and irradiation temperatures, displayed as a function of dose in units of eV g^{-1} (bottom axis) and MGy (top axis).

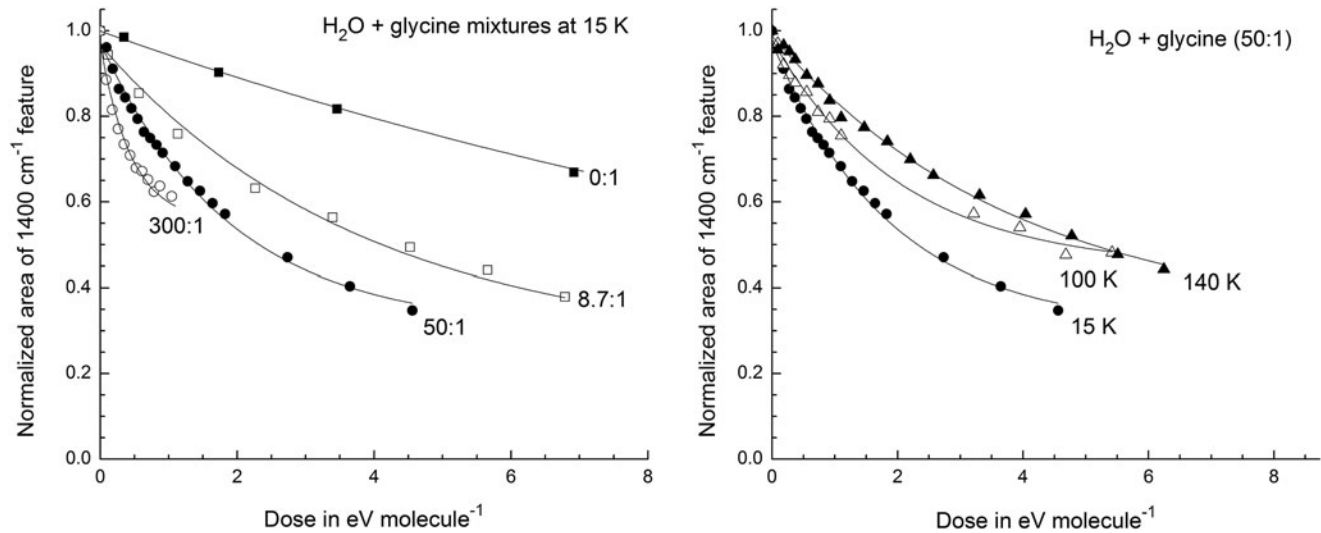


FIG. 6. Area of the glycine feature near 1400 cm^{-1} during irradiation for the same selected ice compositions and irradiation temperatures shown in Fig. 5, but displayed here as a function of dose in units of eV molecule^{-1} . In the left panel, data points for doses higher than $8\text{ eV molecule}^{-1}$ were omitted for clarity.

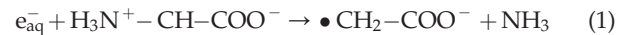
H_2O and glycine. From this, we calculated the number of glycine molecules present after the deposition of each sample, and graphs of the number of glycine molecules versus total energy absorbed, such as shown in Fig. 7, were constructed. The initial slopes of the curves shown lead to glycine's radiation-chemical destruction rate, $G(\text{-gly})$, defined as the number of glycine molecules destroyed per 100 eV absorbed by the sample. This gives another trend, consistent with the manner in which decay rates are typically reported. The $G(\text{-gly})$ values we obtained by this procedure are given in Table 4 for each ice investigated.

4. Discussion

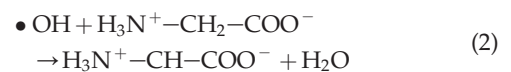
4.1. Radiation chemistry of H_2O +glycine mixtures

Although this paper is not about the reaction products of ices containing glycine, we note that CO_2 formation was readily seen in all our experiments. For example, both Figs. 3 and 4 show an increase in the CO_2 absorbance at 2343 cm^{-1} . The decarboxylation (CO_2 release) of irradiated amino acids has been reported by Gerakines *et al.* (2012) and many others (e.g., Bonifačić *et al.*, 1998). Under our conditions, CO_2 could be formed by molecular decomposition of excited glycine molecules, to make CO_2 and methylamine, or by more-complex pathways as described below.

Our previous paper on amino acid radiolysis (Gerakines *et al.*, 2012) included reactions and mechanisms for amino acid destruction, focusing on direct decomposition of glycine and other molecules. However, here our interest is primarily in glycine at low concentrations in H_2O -rich ices, and a different chemistry applies. The two species directing chemical change in the early stages of radiolysis of H_2O -rich ices are hydrated electrons (e_{aq}^-), which promote reduction, and hydroxyl radicals ($\bullet\text{OH}$), which promote oxidation. The reduction that destroys glycine is



with the intense IR absorptions of H_2O -ice obscuring the IR features of the ammonia produced. Hydroxyl radicals destroy the zwitterionic glycine in our ices through hydrogen-atom abstraction from the glycine's central carbon:



Disproportionation of the resulting zwitterionic radicals gives a molecule of glycine and a protonated imine ($\text{H}_2\text{N}^+ = \text{CHCOO}^-$), the latter reacting with H_2O molecules to produce ammonia and glyoxylic acid:

TABLE 2. HALF-LIFE DOSES IN $\text{eV G}^{-1}/10^{23}$

Ice's H_2O :glycine ratio	Approximate temperature of ice (K)			
	15	100	140	280
0:1	1.01 ± 0.12	1.57 ± 0.38	1.56 ± 0.45	1.21 ± 0.24
8.7:1	1.07 ± 0.22	1.03 ± 0.26	1.13 ± 0.13	
50:1	0.78 ± 0.16	1.46 ± 0.29	1.60 ± 0.25	
100:1	0.56 ± 0.17	0.54 ± 0.11	3.1 ± 5.5	
200:1	0.93 ± 0.73		1.26 ± 0.69	
300:1	—	0.93 ± 0.32	2.3 ± 1.5	

TABLE 3. HALF-LIFE DOSES IN eV MOLECULE^{-1}

Ice's H_2O :glycine ratio	Approximate temperature of ice (K)			
	15	100	140	280
0:1	12.6 ± 1.4	19.5 ± 4.7	19.4 ± 5.6	15.1 ± 3.0
8.7:1	4.24 ± 0.90	4.1 ± 1.0	4.47 ± 0.51	
50:1	2.47 ± 0.52	4.64 ± 0.91	5.06 ± 0.78	
100:1	1.72 ± 0.52	1.65 ± 0.32	10 ± 17	
200:1	2.8 ± 2.2		3.8 ± 2.1	
300:1	—	2.79 ± 0.98	6.9 ± 4.6	

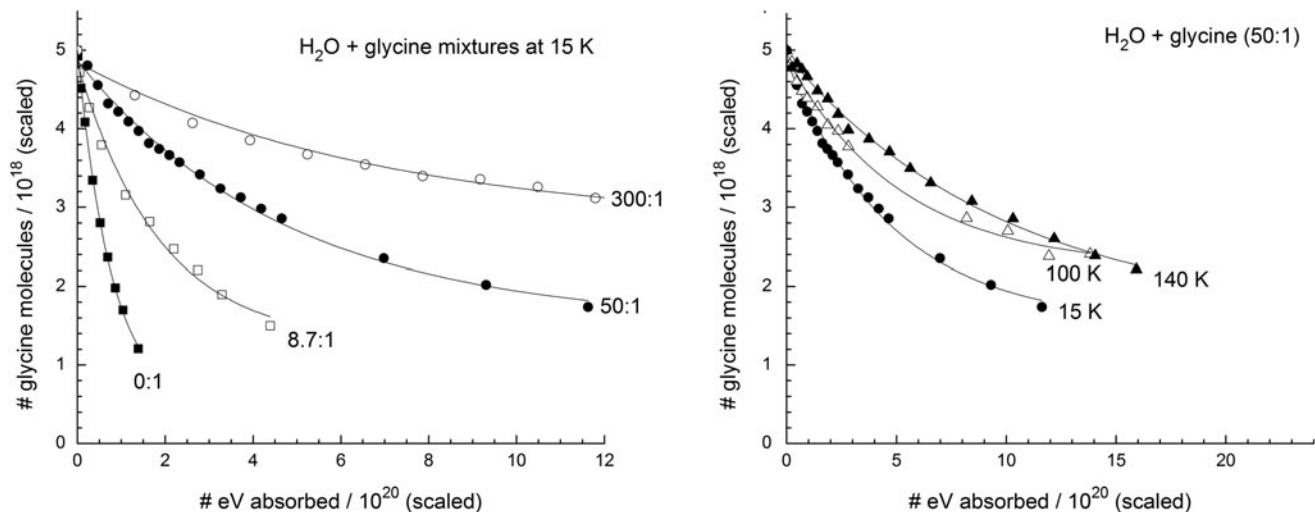
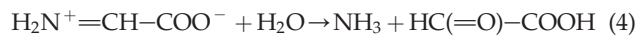
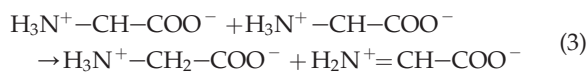


FIG. 7. The decay of glycine as a function of energy absorbed in units of eV for the same selected ice compositions and irradiation temperatures shown in Figs. 5 and 6. Data for each mixture have been scaled to the same initial number of glycine molecules.



It is possible that some of the glyoxylic acid produced is subsequently decomposed into carbon dioxide and formaldehyde (H_2CO), which would account for some of the CO_2 observed. See the work of Neta *et al.* (1970) and Bonifačić *et al.* (1998) for additional information on these reactions.

4.2. Radiation chemistry on Mars

Although the multiple sources of ionizing radiation at Mars, and elsewhere, have been reviewed several times in this journal in recent years, much less attention has been paid to the differences between radiation chemistry, which is our area of study, and the related fields of nuclear chemistry and radiochemistry. Unlike nuclear chemistry, the focus of radiation chemistry is neither transmutations nor radioactive decay; and unlike radiochemistry, the focus is not the reactions of a single radioactive tracer. In radiation chemistry, the emphasis is on the chemical changes induced by an incident ionizing agent. Specifically, as a keV–GeV ion (or electron or photon) passes through an icy solid, it produces a trail of ionizations and excitations as the ion's energy is degraded.

TABLE 4. $G(\text{GLY})$ VALUES

Ice's $\text{H}_2\text{O}:\text{glycine}$ ratio	Approximate temperature of ice (K)			
	15	100	140	280
0:1	5.78 ± 0.54	3.55 ± 0.42	3.32 ± 0.59	4.52 ± 0.49
8.7:1	1.88 ± 0.28	1.80 ± 0.26	1.67 ± 0.10	
50:1	0.716 ± 0.092	0.516 ± 0.058	0.338 ± 0.032	
100:1	0.524 ± 0.118	0.611 ± 0.085	0.163 ± 0.085	
200:1	0.371 ± 0.073		0.139 ± 0.053	
300:1	0.304 ± 0.028	0.088 ± 0.018	0.058 ± 0.015	

Many of these ionizations will produce secondary electrons, which, in turn, travel through the ice creating separate tracks (δ rays) of yet more ionizations and excitations that lead to even more chemical changes. Therefore it must be emphasized that the direct chemical action of the incident radiation is completely overshadowed by the chemistry produced by secondary electrons of energies less than about 20 eV (Pimblott and LaVerne, 2007). This means that, to a robust first approximation, the final products of various ionizing radiations (*e.g.*, e^- , H^+ , He^+ , X-rays, and γ rays) are identical and, however counterintuitive, strongly resemble the products of conventional photochemistry (Spinks and Woods, 1964). For comparisons of IR spectra of ices subjected to radiolytic and photolytic processing, see the work of Hudson and Moore (2001), Hudson *et al.* (2001), and Gerakines *et al.* (2000), and the discussion of Hudson and Moore (2000). For comparisons from other laboratories, see, for example, the work of Baratta *et al.* (2002).

Another point worth emphasizing concerns the particle energies available at Mars and the chemistry they induce. It is perhaps tempting to think that investigations of martian subsurface radiation chemistry must focus on the effects of particles with energies in the gigaelectronvolt and higher ranges, since they dominate the radiation environment at depths of about 1 m (Dartnell *et al.*, 2007), and not on the changes induced by MeV ions such as we have used. However, while GeV ions penetrate more deeply into an ice than MeV ions, the energy delivered by the former per unit distance (*i.e.*, their stopping power) is several orders of magnitude smaller than that of MeV ions of the same type. Combined with the fact that GeV cosmic rays are of much lower abundance than those of smaller energy, the conclusion is that the radiation chemistry on Mars induced by MeV ions dominates that of GeV ions.

The laboratory results we report here apply to depths of a few meters on Mars, where our proton source readily reproduces the radiation doses expected over long times. Specifically, the martian dose rate estimated by Dartnell *et al.* (2007) is on the order of 0.1 Gy year^{-1} , or approximately $2 \times 10^{-8} \text{ eV}$

molecule⁻¹ year⁻¹, at a depth of 1 m. This implies that a molecule under the surface of Mars for about 50 million years will absorb about 1 eV, a dose delivered to an ice by our Van De Graaff accelerator in 20–30 min. In short, our glycine experiments employ martian subsurface radiation doses.

4.3. Laboratory radiation results

Several important points emerge from an examination of our data and calculations. From Figs. 5 and 6, and Tables 2 and 3, it can be seen that glycine's rate of radiation-induced destruction depends on both temperature and the presence of H₂O-ice. This is a trend that is not always obvious in previous work. It suggests that measurements on glycine's radiation stability in the absence of H₂O ice may not be directly applicable to Mars or other environments.

Comparison of the left-hand panels of Figs. 5, 6, and 7 also leads to an important conclusion, namely, that in any discussion of glycine's radiation stability in one chemical environment compared to another the units of dose must be specified carefully. In the present case, Fig. 5 shows little difference in the decay rates of glycine and H₂O + glycine ices on an electronvolt per gram basis, but differences are obvious with Fig. 6's electronvolt per molecule scale. Figure 7, with a scale of total energy absorbed, shows yet a different trend in the initial decay slopes, one which appears to be the opposite of that in Fig. 6. All three figures accurately portray the kinetics, but clearly the choice of *x* axis units can alter the assessment of glycine's radiation stability determined from a half-life dose (*D*_{1/2}). In contrast, the right-hand panels of Figs. 5, 6, and 7 show the decay curves for the same ice composition at three irradiation temperatures. In these cases, the trends are nearly identical, since the conversion from one horizontal scale to the next is primarily dependent on the ice composition.

Another conclusion, taken from Fig. 7 and Table 4, is that *G*(-gly) depends on the initial relative abundance of glycine. This is not surprising since *G*(-M) is a kinetic parameter that refers to the initial destruction of a molecule *M*. Although

commonly termed a radiation-chemical yield, *G*(-M) says nothing about the final equilibrium abundance of a molecule. As a kinetic quantity, it is not surprising that when the concentration of *M* is decreased then its loss rate ($-dM/dt$) falls, as seen in Fig. 8 in moving to the right.

Our experiments also emphasize the importance of distinguishing between the so-called *direct* and *indirect* actions of ionizing radiation. In the former, energy reaches glycine molecules directly, but energy transfers and chemical change in the latter are caused by H₂O molecules and their radiation products (e.g., e⁻, H, OH, H₂O₂). Significantly, Fig. 8 shows that at greater dilutions, where indirect radiolytic action dominates in ices, a limiting value of *G*(-gly) is approached. For a dilution of 300:1, the limits are approximately 0.30, 0.09, and 0.06 at 15, 100, and 140 K, respectively. Other things being equal, glycine will have a greater resistance to radiation-induced change when it is present in an ice in lower abundances. Figure 8 also shows that the value of *G*(-gly) falls as the irradiation temperature increases from 15 to 140 K.

4.4. Glycine on Mars

Macromolecular carbon has been found in martian meteorites (McDonald and Bada, 1995; Glavin *et al.*, 1999), but present conditions do not appear to favor the synthesis of organics on Mars. However, it is expected that amino acids, including glycine, are brought to Mars and other Solar System objects by meteorites and perhaps by comets too. Any surficial martian glycine will be destroyed by photochemistry, but subsurface material can persist until altered by agents such as cosmic radiation, electrical discharges, or geological processes. If surface material is transported downward faster than it can be destroyed photochemically, then its chances of survival will increase due to the shielding provided by soil and ices.

The purpose of this paper has been to report new data that can be used to understand and predict glycine's radiolytic fate on Mars and other worlds. However, the application of laboratory data and theoretical models to martian surface

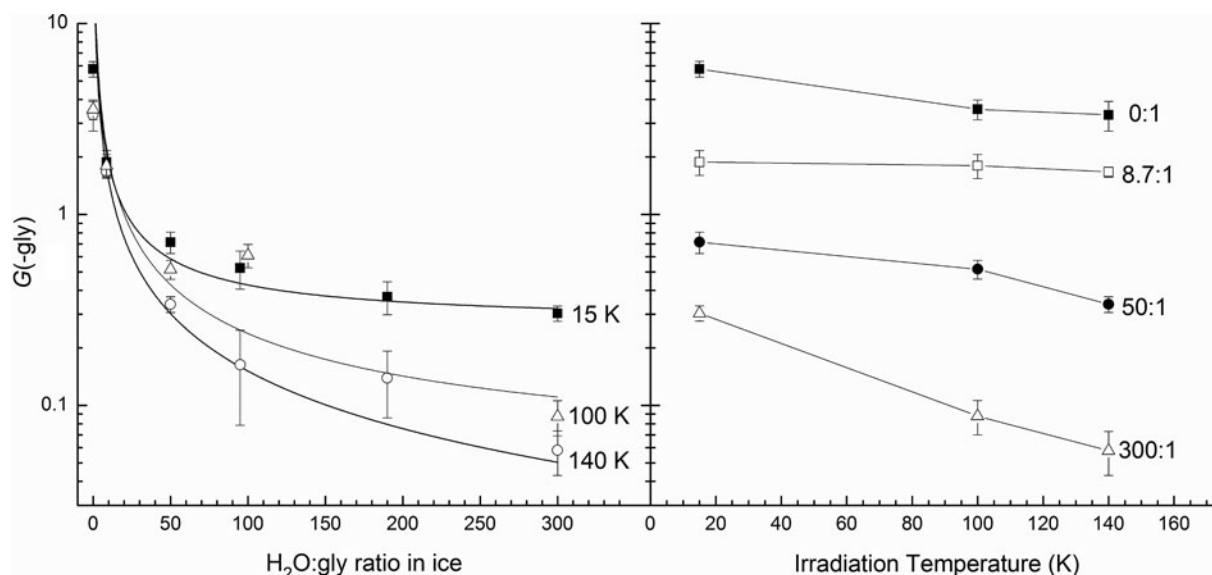


FIG. 8. *G*(-gly) as a function of ice composition (left) and irradiation temperature (right).

chemistry requires assumptions to be made about surface composition. Dartnell *et al.* (2007) published dose-rate estimates for three such models, with their model for pure-H₂O ice being the closest to our experiments. Their calculations show that, over a million years, H₂O-ice at a depth of 1 m receives a dose of 0.31 MGy, falling to about 0.17 MGy at 2 m. Denser surface materials (*e.g.*, rock) further reduce these doses, and the rock-ice model of Dartnell *et al.* (2007) indeed shows doses after 1 million years of 0.08 and 0.01 MGy for depths of 1 and 2 m, respectively. In all cases, background radioactivity is expected to dominate other radiation sources below a depth of about 4.5 m (Dartnell *et al.*, 2007).

Our warmest, most dilute ice in Fig. 8, at 140 K (−133°C), has $G(\text{gly}) \approx 0.05$, and our Tables 2 and 3 give half-life doses for it that can be applied to Mars. Selecting a value of $D_{1/2} = 6.9 \text{ eV molecule}^{-1}$ (Table 3) gives a fractional destruction rate of 0.5 (50%) in $3.5\text{--}4.5 \times 10^8$ years at a depth of 1 m. The limiting G values we report for our most dilute glycine-containing ice can also be extrapolated to high temperature, giving $G = 0.043$ at 210 K on Mars.

In Fig. 9, we show the estimated half-life of glycine as a function of depth in the martian subsurface, where the half-life dose for a H₂O+gly (300:1) mixture at 140 K has been applied, along with absorbed dose data for the rock-ice model by Dartnell *et al.* (2007). Note that doses for the rock-ice model are generally lower than that for pure H₂O-ice but likely represent a more realistic representation of the martian surface. The drill on the Mars Science Laboratory Curiosity rover is capable of sampling the soil to a depth of about 5 cm, where the estimated half-life of glycine in ice is about

$2\text{--}3 \times 10^8$ years, which is a relatively short timescale when compared to the $\sim 3.5\text{--}4$ billion years since the era when Mars was likely the most habitable.

5. Summary and Conclusions

- G values for glycine radiolytic destruction, based on an *in situ* method of analysis, have been presented, and variations with temperature and initial abundance have been explored. As G values for glycine destruction are found to depend on both temperature and the presence of H₂O-ice, published room-temperature data on pure glycine must be used with caution in Mars applications, as must laboratory data not obtained by *in situ* methods.
- Radiation half-life doses in eV g^{-1} and eV molecule^{-1} have been determined for glycine embedded in an amorphous solid (H₂O-ice). Statements about the stability of glycine to radiolysis need to be accompanied by careful attention to dose units.
- Any glycine, or other amino acid, present on Mars is expected to have a low abundance and will likely be partially shielded by its surroundings, whether soil or ice, from the direct effects of ionizing radiation. Our data provide a laboratory basis for estimating glycine destruction times in the case of a H₂O-ice environment.
- The differences in the direct and indirect effects of ionizing radiation are important in any application of laboratory data to problems of solar system chemistry. Moreover, since the existence of extraterrestrial crystalline amino acids is unlikely, due to these molecules' low abundance, attention must be paid in laboratory studies to the amorphous or crystalline nature of amino acid samples.

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Author Disclosure Statement

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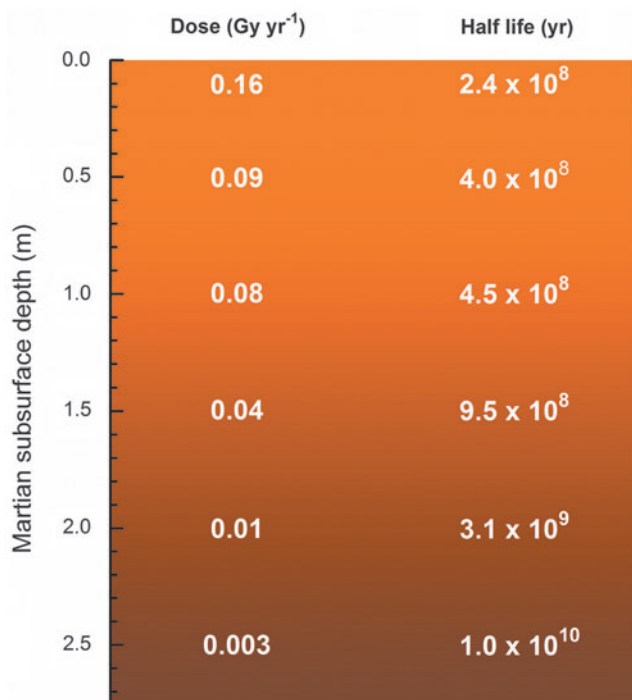


FIG. 9. The estimated half-life for glycine mixed with H₂O-ice under the surface of Mars as a function of depth. Data are based on the half-life dose of glycine in our H₂O+gly (100:1) mixture at 140 K (see Table 2) and the model of dose versus depth for an ice-rock mixture given by Dartnell *et al.* (2007).

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