

## 12.4 USING C-BAND DUAL-POLARIZATION RADAR SIGNATURES TO IMPROVE CONVECTIVE WIND FORECASTING AT CAPE CANAVERAL AIR FORCE STATION AND NASA KENNEDY SPACE CENTER

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### 1. INTRODUCTION AND BACKGROUND

The United States Air Force's 45<sup>th</sup> Weather Squadron (45WS) is the organization responsible for monitoring atmospheric conditions at Cape Canaveral Air Force Station and NASA Kennedy Space Center (CCAFS/KSC) and issuing warnings for hazardous weather conditions when the need arises. One such warning is issued for convective wind events, for which lead times of 30 and 60 minutes are desired for events with peak wind gusts of 35 knots or greater (i.e., Threshold-1) and 50 knots or greater (i.e., Threshold-2), respectively (Roeder et al. 2014).

45WS forecasters use a variety of instrumentation to increase protection for personnel, facilities, space launches, and space mission payloads at CCAFS/KSC, including a C-band dual-polarization radar (45WS-WSR) (Roeder et al. 2009) and the Cape Weather Information Network Display System (Cape WINDS) – a network of 29 weather observation towers strategically placed throughout and around the CCAFS/KSC complex (Fig. 1). Even with this wealth of technology available, forecasting convective wind events (e.g., downbursts) is a difficult process. It can be challenging to identify which storms will produce threshold-level downbursts (i.e., peak wind gust of 35 knots or greater) from those that will not, and the lead times offered for threshold-level events are often shorter than desired by the 45WS.

The purpose of this study is to identify C-band dual-polarization radar signatures that can be used by 45WS forecasters in real-time to:

- 1) Provide increased lead times for downburst events at CCAFS/KSC;
- 2) Decrease false alarm ratios (FARs) for 45WS convective wind warnings;
- 3) Differentiate between storm cells that will produce below-threshold wind gusts, Threshold-1 gusts, and Threshold-2 gusts.

Past studies have noted various radar signatures that, when placed in an environmental context, can provide insight into the physical processes occurring within downburst-producing thunderstorms. Some examples include the peak value of radar reflectivity factor ( $Z_h$ ) within a storm cell (Loconto 2006), a column-like region

of positive differential reflectivity ( $Z_{dr}$ ) values above the environmental 0 °C level (Tuttle et al. 1989), and a region of near-0 dB  $Z_{dr}$  values extending below the 0 °C level within the descending precipitation core of the storm (Wakimoto and Bringi 1988, Scharfenberg 2003). Additional studies have examined the effects that environmental conditions have on precipitation processes within downburst-producing thunderstorms. Srivastava (1987) showed that precipitation ice can be very important to the formation and intensification of wet downbursts, especially in warm humid environments. Meischner et al. (1991) showed that the melting of falling precipitation ice over a shallow layer beneath the 0 °C level can significantly enhance downburst intensity, and White (2015) indicated how  $Z_{dr}$  can be used to examine the melting of precipitation ice.

These and additional radar signatures were explored in this study. Given that the typical lifecycle of ordinary convective cells in the southeastern United States is around 20 – 40 minutes (Smith et al. 2004), attention was given to longer-lived thunderstorm systems (e.g., multicellular convection) in an effort to increase lead times for 45WS convective wind warnings. It was hypothesized that the aforementioned radar signatures can be observed within the multiple updraft-downdraft cycles that occur within multicellular (multicell) convection, which can be used by 45WS forecasters to identify storm systems that may eventually produce a threshold-level downburst at the CCAFS/KSC complex.

Preliminary results of this study indicate that certain dual-polarization radar signatures may offer increased lead times for threshold-level downbursts. Specifically, it was found that the peak height of the 1 dB contour within a  $Z_{dr}$  column, the peak height of the 30 dBZ  $Z_h$  contour co-located with  $Z_{dr}$  values around 0 dB, the peak  $Z_h$  value, and the gradient in  $Z_{dr}$  below the 0 °C level within a descending reflectivity core (DRC) may offer average lead times of 40 – 46 minutes, especially in multicellular systems.

### 2. DATA AND METHODOLOGY

#### 2.1 Data

All radar data used in this study were from the C-band 45WS-WSR, and were provided by the 45WS. Environmental conditions were analyzed using atmospheric soundings from the KXMR site at the KSC

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### 3. RESULTS AND DISCUSSION

#### 3.1 Overview

Four radar signatures common amongst threshold-level downburst-producing storm systems at CCAFS/KSC have been identified in this study:

- 1) Peak height of the 1 dB  $Z_{dr}$  contour within a  $Z_{dr}$  column;
- 2) Peak height of co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$ ;
- 3) Peak  $Z_h$  value within the storm system;
- 4) Gradient in  $Z_{dr}$  below the 0 °C level within a DRC, including an increase in  $Z_{dr}$  to a value of at least 3 dB.

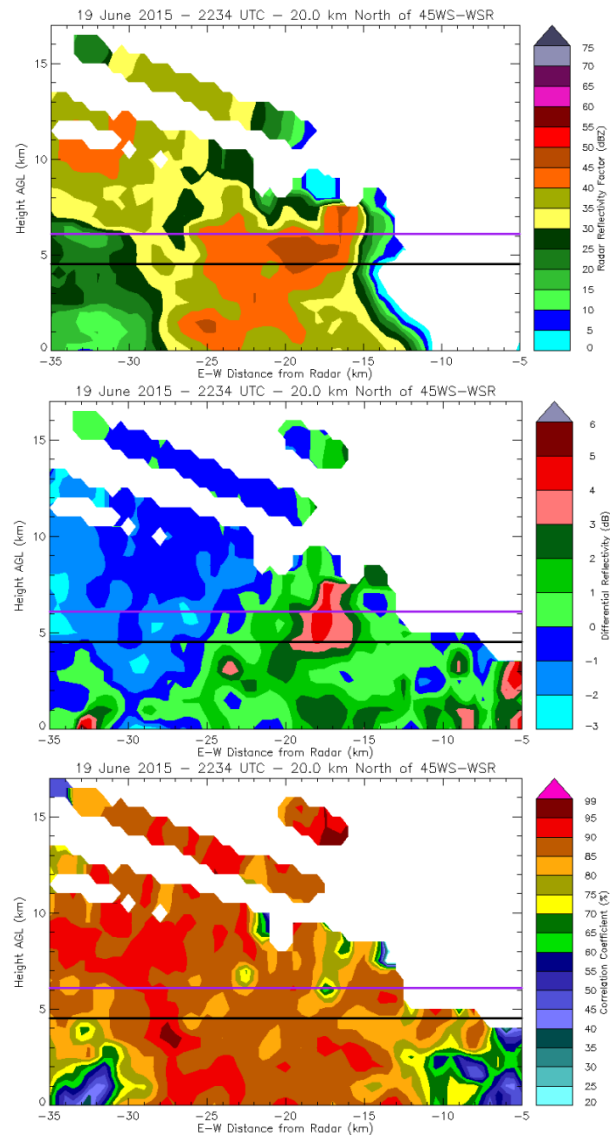
These signatures imply the importance of precipitation ice in threshold-level downbursts at CCAFS/KSC, as discussed below. The lead times offered by these signatures were relatively large in multicell events, due to the multiple updraft-downdraft cycles observed.

#### 3.2 Signature #1 – Peak $Z_{dr}$ Column Height

The first radar signature identified was the peak heights of  $Z_{dr}$  columns within threshold-level downburst-producing thunderstorms. An example of this signature is presented in Fig. 2, where the 1 dB  $Z_{dr}$  column can be seen approximately 16 to 20 km west of the radar extending to roughly 7.5 km above ground level (AGL), or about 3 km above the 0 °C level. The signature seen in Fig. 2 was present before a 35-knot downburst wind gust was recorded on Cape WINDS Tower 1007 about 48.5 minutes later at 2322:30 UTC. Since positive  $Z_{dr}$  values are typical of oblate-shaped liquid hydrometeors and  $Z_{dr}$  values around 0 dB are typical of spherical or ice hydrometeors (Herzogh and Jameson 1992),  $Z_{dr}$  columns indicate regions where liquid hydrometeors are lofted above the environmental 0 °C level by the storm's updraft (Tuttle et al. 1989). These lofted liquid hydrometeors will either freeze and subsequently melt as they descend back below the 0 °C level and/or evaporate which, along with mass loading, will contribute to negative buoyancy within the storm and lead to formation and intensification of the downburst (Srivastava 1987).

In this study, it was observed that the 1 dB contour within a  $Z_{dr}$  column extended at least 1 km above the environmental 0 °C level in 12 of the 14 (85.71%) threshold-level events analyzed. This suggests that the lofting of liquid hydrometeors above the 0 °C level and the subsequent freeze-melting, evaporation, and mass loading processes are important to the formation of threshold-level downbursts at CCAFS/KSC.

The lead times offered by this signature ranged from 11.50 minutes to 78.50 minutes, with a mean lead time of 40.67 minutes and a median lead time of 42.50 minutes. Thus, the mean lead time offered by the 1 dB  $Z_{dr}$  contour extending at least 1 km above the 0 °C level is more than 10 minutes longer than desired by the 45WS for wind gusts that meet their warning Threshold-1. Furthermore, the maximum observed lead time for this signature is nearly 20 minutes longer than desired



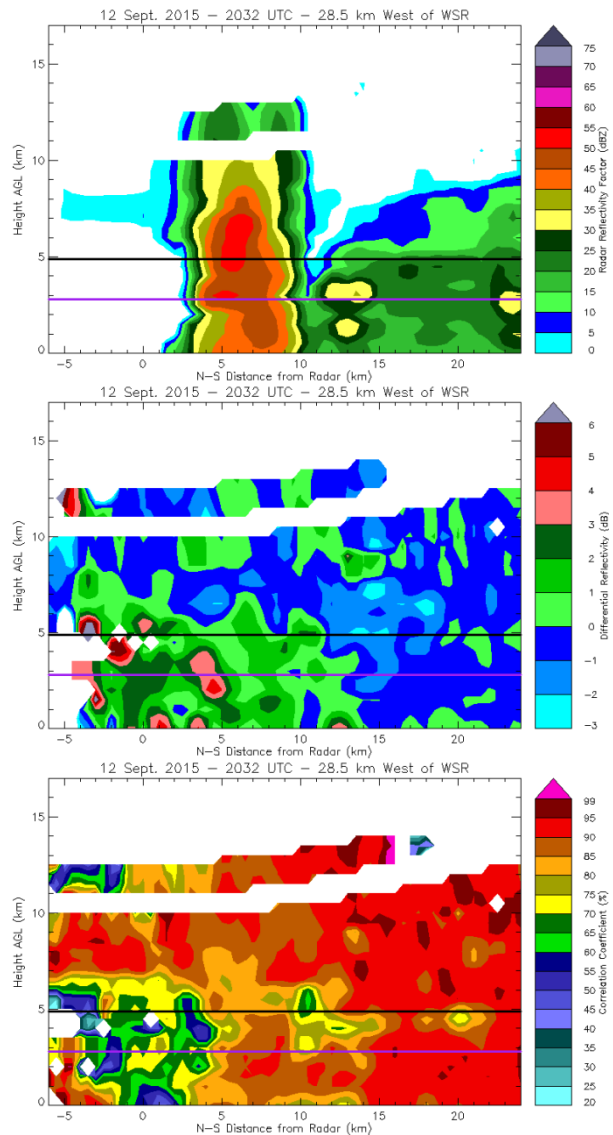
**Figure 2.** Vertical cross sections taken in the east-west direction (i.e., x-z plane shown) 20.0 km north of the 45WS-WSR during the volume scan that ended around 2234 UTC on 19 June 2015. The top image shows radar reflectivity factor in dBZ, the center image shows differential reflectivity in dB, and the bottom image shows correlation coefficient in percentage form. In each image, the black horizontal line marks the 0 °C level and the purple horizontal line marks the height of minimum equivalent potential temperature calculated from the 0000 UTC KXMR sounding on 20 June 2015.

by the 45WS for events that meet their warning Threshold-2. Therefore, this signature may offer the lead times desired by the 45WS for threshold-level downbursts. However, a peak 1 dB  $Z_{dr}$  column height of at least 1 km above the 0 °C level was also observed in 4 of the 4 (100%) null events analyzed in this study. Therefore, while the 1 dB  $Z_{dr}$  column height may offer increased lead times for threshold-level downbursts, it may not serve as well for distinguishing storm cells that

will produce threshold-level events from those that will not. This signature could possibly lead to a high FAR for 45WS convective wind warnings if used as the sole downburst prediction parameter, but may be very useful if used in combination with other signatures.

### 3.3 Signature #2 – Height of 30 dBZ $Z_h$ and 0 dB $Z_{dr}$

The second radar signature identified in this study was the peak height of approximately co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$ . An example of this signature can be seen in Fig. 3. In Fig. 3, the region of



**Figure 3.** Vertical cross sections taken in the north-south direction (i.e., y-z plane shown) 28.5 km west of the 45WS-WSR during the volume scan that ended around 2032 UTC on 12 September 2015. The variables and horizontal lines shown are the same as in Fig. 1, with the environmental conditions calculated using the 0000 UTC KXMR sounding on 13 September 2015.

co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$  can be seen from 3 to 10 km north of the 45WS-WSR extending to 10 km AGL, or about 5 km above the 0 °C level. The signature seen in Fig. 3 was present before a 42-knot downburst wind gust was recorded on Cape WINDS Tower 0421 about 50.5 minutes later at 2122:30 UTC. This signature implies the presence of a significant amount of precipitation ice aloft in the storm cell. Past studies have shown that graupel and small hailstones typically have a  $Z_h \geq 29 - 33$  dBZ (Deierling et al. 2008), so a  $Z_h$  value of 30 dBZ was used in this study as the lower boundary for precipitation ice.  $Z_{dr}$  values within regions of precipitation ice are generally around 0 dB due to the spherical shape of the ice hydrometeors and/or the lower dielectric within the ice hydrometeors as a result of their lower bulk densities (Herzogh and Jameson 1992).

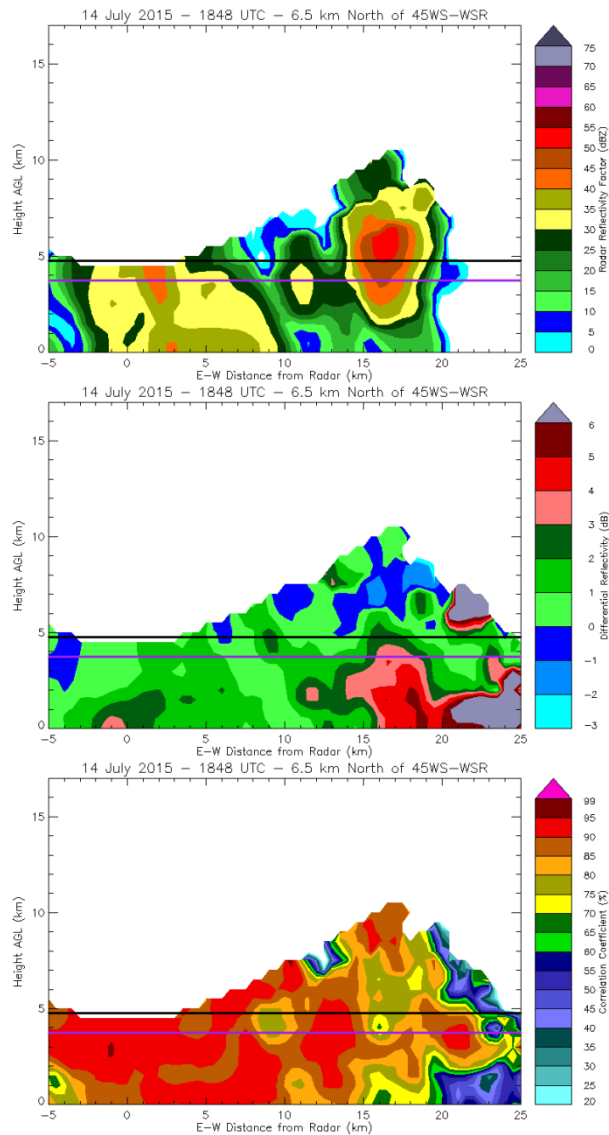
Past studies have shown that the melting of ice hydrometeors may contribute more greatly to negative buoyancy within a downburst than evaporation of liquid hydrometeors, despite the latent heat of vaporization being roughly 8.5 times larger than latent heat of fusion, especially in regions where high relative humidity values enhance melting and suppress evaporation (Srivastava 1987). Therefore, the presence of a large quantity of precipitation ice within a storm cell may indicate the potential for negative buoyancy to be enhanced within the downburst via latent heat of melting and downward forcing from mass loading (Srivastava 1987).

In this study, the co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$  extended at least 3 km above the 0 °C level in 13 of the 14 (92.86%) threshold-level events analyzed. This suggests that ice precipitation is important to the formation of threshold-level downbursts at CCAFS/KSC, as indicated by several past studies.

The lead times offered by this signature ranged from 3.50 to 78.50 minutes, with a mean lead time of 40.88 minutes and a median lead time of 35.50 minutes. As with Signature #1, the mean lead time offered by co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$  was more than 10 minutes longer than desired by the 45WS for Threshold-1 events and the maximum lead time offered by this signature was nearly 20 minutes longer than desired for Threshold-2 events. However, co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$  extended at least 3 km above the 0 °C level in 4 of the 4 (100%) null events analyzed. Therefore, as with Signature #1, Signature #2 may be very useful for offering the increased lead times desired by the 45WS for threshold-level events, but may cause FAR to be relatively high if used alone.

### 3.4 Signature #3 – Peak $Z_h$ Value

The third radar signature identified in this study was the peak value of  $Z_h$  within the threshold-level downburst-producing storm system. An example of this signature can be seen in Fig. 4, where the peak  $Z_h$  is around 50 – 55 dBZ. The signature seen in Fig. 4 was present before a 35-knot downburst wind gust was recorded on Cape WINDS Tower 0001 about 24.5 minutes later at 1912:30 UTC.



**Figure 4.** Vertical cross sections taken in the east-west direction (i.e., x-z plane shown) 6.5 km north of the 45WS-WSR during the volume scan that ended around 1848 UTC on 14 July 2015. The variables and horizontal lines shown are the same as in Figs. 1 – 2, with the environmental conditions calculated using the 1500 UTC KXMR sounding on 14 July 2015.

This signature implies the presence of large-sized hydrometeors and/or a large concentration of hydrometeors within the storm cell. A region of large  $Z_h$  that develops above the 0 °C level implies the presence of a large quantity of precipitation ice in the cell, while a region of large  $Z_h$  that develops below the 0 °C level implies the presence of a large quantity of liquid hydrometeors within the cell (Loconto 2006). The presence of a large quantity of hydrometeors (liquid and/or ice) can enhance the negative buoyancy within a downdraft through mass loading (Srivastava 1987). Additionally, melting from the large amount of ice hydrometeors and evaporation from the large amount of

liquid hydrometeors can enhance negative buoyancy through latent heat absorbed by these processes (Srivastava 1987).

For this study, only the peak  $Z_h$  value was identified; the location of peak  $Z_h$  relative to the 0 °C level will be examined in future work. A peak  $Z_h$  of at least 50 dBZ was present in 13 of the 14 (92.86%) threshold-level events analyzed, which suggests the importance of a large quantity of hydrometeors in producing threshold-level downbursts at CCAFS/KSC.

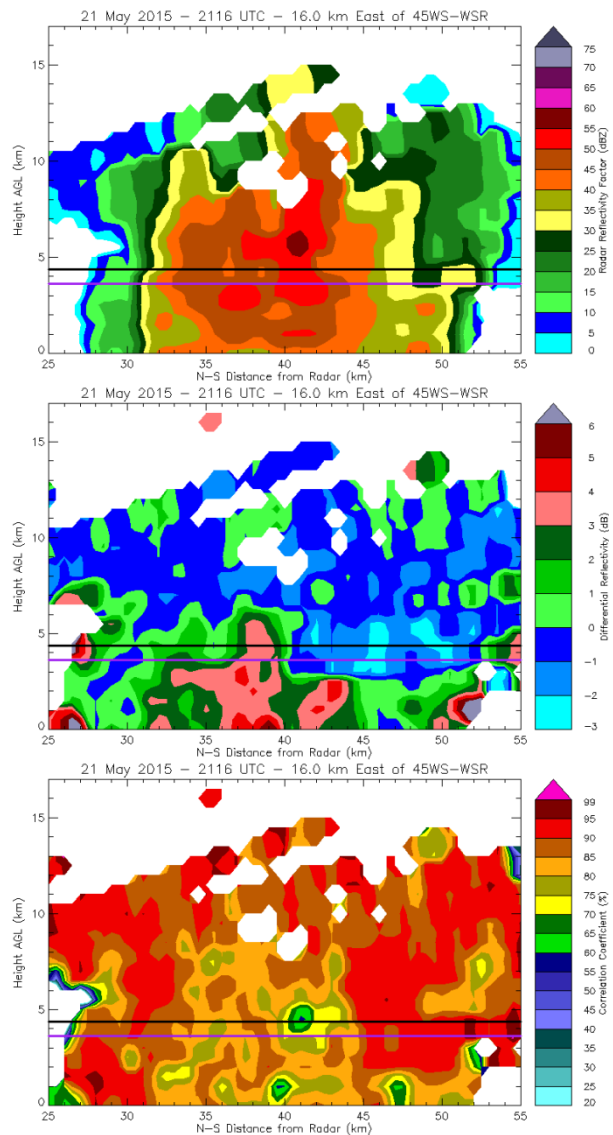
The lead times offered by this signature ranged from 11.50 to 78.50 minutes, with a mean lead time of 45.88 minutes and a median lead time of 48.50 minutes. The mean lead time offered by the peak  $Z_h$  value within a storm system, as with the previous two signatures, is more than 10 minutes longer than desired by the 45WS for Threshold-1 events and the maximum lead time is nearly 20 minutes longer than desired for Threshold-2 events. However, a peak  $Z_h$  value of at least 50 dBZ was observed in 3 of the 4 (75%) null events analyzed. This indicates that peak  $Z_h$  of at least 50 dBZ in a storm system may offer longer lead times for threshold-level downbursts, but may lead to a high FAR if used as the sole predictor of threshold-level events.

### 3.5 Signature #4 – $Z_{dr}$ Gradient within the DRC

The final radar signature identified in this study was the gradient in  $Z_{dr}$  within the DRC of a downburst-producing storm cell. An example of this signature is provided in Fig. 5, where  $Z_{dr}$  values can be seen increasing with decreasing height from about -1 dB to values greater than 3 dB within the DRC located about 41 – 44 km north of the 45WS-WSR around 2 – 3 km AGL. The signature seen in Fig. 5 was present before a 51-knot downburst wind gust was recorded on Cape WINDS Tower 0019 16.5 minutes later at 2132:30 UTC. Note that Tower 0019 is not shown in Fig. 1, but is located approximately halfway between Towers 0022 and 0015.

This signature indicates the melting of small precipitation ice hydrometeors (e.g., graupel) over a relatively shallow layer beneath the 0 °C level. Since  $Z_{dr}$  is a measure of the reflectivity-weighted mean diameter of a given drop size distribution (DSD) (Jameson 1983), and smaller ice hydrometeors typically melt completely before larger ice hydrometeors do, an increase in  $Z_{dr}$  from near-0 dB around the 0 °C level to positive values with decreasing height below the 0 °C level implies the melting of smaller ice hydrometeors (Meischner et al. 1991). Given that the latent heat absorbed by melting ice hydrometeors may contribute more to negative buoyancy within a downburst than the latent heat absorbed by evaporating liquid hydrometeors in regions where relative humidity values are high (Srivastava 1987), regions where  $Z_{dr}$  increases sharply over a shallow layer below the 0 °C level may suggest the melting of a large quantity of smaller ice hydrometeors, which can lead to significant downburst intensification (Meischner et al. 1991). Past studies have also noted that  $Z_{dr}$  values increase from near-0 dB around the 0 °C level to at least 3 dB below the 0 °C level within

downburst-producing storms in warm humid climate regions (White 2015).



**Figure 5.** Vertical cross sections taken in the north-south direction (i.e., y-z plane shown) 16.0 km east of the 45WS-WSR during the volume scan that ended around 2116 UTC on 21 May 2015. The variables and horizontal lines shown are the same as in Figs. 1 – 3, with the environmental conditions calculated using the 0000 UTC KXMR sounding on 22 May 2015.

In this study, it was observed that  $Z_{dr}$  values increased from around 0 dB near the 0 °C level to at least 3 dB within 2.5 km below the 0 °C level in a DRC in 13 of the 14 (92.86%) threshold-level events analyzed. This indicates the importance of the melting of smaller ice hydrometeors to the intensification of threshold-level downbursts at CCAFS/KSC.

The lead times offered by this signature ranged from 1.50 minutes to 78.50 minutes, with a mean lead time of 40.42 minutes and a median lead time of 41.50

minutes. As with the previous signatures, the mean lead time offered by this signature is more than 10 minutes longer than desired by the 45WS for Threshold-1 events and the maximum lead time was nearly 20 minutes longer than desired by the 45WS for Threshold-2 events. However, an increase in  $Z_{dr}$  to at least 3 dB within 2.5 km below the 0 °C level was observed in the DRC of 4 of the 4 (100%) null events analyzed. Thus, the gradient in  $Z_{dr}$  below the 0 °C level, including an increase in  $Z_{dr}$  to at least 3 dB, may provide the lead times desired by the 45WS for threshold-level downburst events, but may also lead to a high FAR if used as the sole downburst prediction parameter.

#### 4. SUMMARY AND FUTURE WORK

The purpose of this study was to identify C-band dual-polarization radar signatures that can be used in real-time by 45WS forecasters to increase lead times and decrease FAR for their convective wind warnings and help distinguish which storms will produce a Threshold-1 peak wind gust (35+ knots) from those that will also produce a Threshold-2 peak wind gust (50+ knots). So far, four common radar signatures have been identified amongst threshold-level downburst-producing storm systems at CCAFS/KSC:

- 1) The 1 dB  $Z_{dr}$  contour within a  $Z_{dr}$  column extending at least 1 km above the 0 °C level;
- 2) Co-located values of 30 dBZ  $Z_h$  and near-0 dB  $Z_{dr}$  extending at least 3 km above the 0 °C level;
- 3) Peak  $Z_h$  value of at least 50 dBZ within the storm system;
- 4)  $Z_{dr}$  increasing to at least 3 dB within 2.5 km below the 0 °C level in the DRC.

The mean lead times offered by each of these signatures ranged from 40 – 46 minutes, with median lead times between 35 – 49 minutes. The longer lead times were mainly due to the large number of multicell events within the storm sample, each of which contained several updraft-downdraft cycles and allowed these signatures to be studied well in advance of the threshold-level gust recorded on the Cape WINDS.

These signatures are all related to precipitation ice, which suggests that the melting and mass loading of precipitation ice is very important in the formation of threshold-level downbursts at CCAFS/KSC.

This work is part of an ongoing research project. While the radar signatures identified in this study may offer the lead times desired by the 45WS, future work in this project will include identifying signatures that are more exclusive to threshold-level events and less common in null events in an effort to decrease FAR for 45WS convective wind warnings. Signatures will also be examined that may help distinguish Threshold-1 events from Threshold-2 events. Given the relatively small sample size of 14 threshold-level events and 4 null events used in this study, more cases will be added to each of these categories in the near future to further examine the occurrence of these four radar signatures and any other new signatures that are identified.

Other future work will include examining the location of peak  $Z_h$  relative to the 0 °C level in threshold-level events. Through not discussed herein, several events in this study occurred within the presence of sea breeze fronts, gust fronts, and/or storm mergers. The presence of these features may also be useful in forecasting threshold-level downbursts at CCAFS/KSC. Furthermore, environmental data will be examined in more detail to better understand how certain conditions (e.g., relative humidity profile, temperature lapse rate,  $\theta_e$  profile) can be used to meet the three primary goals of this study. The results of this research project will eventually be integrated into a new algorithm that can be used by 45WS forecasters to aid in the forecasting of downbursts at CCAFS/KSC.

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