



9

Extreme Storms

KEY FINDINGS

1. Human activities have contributed substantially to observed ocean–atmosphere variability in the Atlantic Ocean (*medium confidence*), and these changes have contributed to the observed upward trend in North Atlantic hurricane activity since the 1970s (*medium confidence*).
2. Both theory and numerical modeling simulations generally indicate an increase in tropical cyclone (TC) intensity in a warmer world, and the models generally show an increase in the number of very intense TCs. For Atlantic and eastern North Pacific hurricanes and western North Pacific typhoons, increases are projected in precipitation rates (*high confidence*) and intensity (*medium confidence*). The frequency of the most intense of these storms is projected to increase in the Atlantic and western North Pacific (*low confidence*) and in the eastern North Pacific (*medium confidence*).
3. Tornado activity in the United States has become more variable, particularly over the 2000s, with a decrease in the number of days per year with tornadoes and an increase in the number of tornadoes on these days (*medium confidence*). Confidence in past trends for hail and severe thunderstorm winds, however, is *low*. Climate models consistently project environmental changes that would putatively support an increase in the frequency and intensity of severe thunderstorms (a category that combines tornadoes, hail, and winds), especially over regions that are currently prone to these hazards, but confidence in the details of this projected increase is *low*.
4. There has been a trend toward earlier snowmelt and a decrease in snowstorm frequency on the southern margins of climatologically snowy areas (*medium confidence*). Winter storm tracks have shifted northward since 1950 over the Northern Hemisphere (*medium confidence*). Projections of winter storm frequency and intensity over the United States vary from increasing to decreasing depending on region, but model agreement is poor and confidence is *low*. Potential linkages between the frequency and intensity of severe winter storms in the United States and accelerated warming in the Arctic have been postulated, but they are complex, and, to some extent, contested, and confidence in the connection is currently *low*.
5. The frequency and severity of landfalling “atmospheric rivers” on the U.S. West Coast (narrow streams of moisture that account for 30%–40% of the typical snowpack and annual precipitation in the region and are associated with severe flooding events) will increase as a result of increasing evaporation and resulting higher atmospheric water vapor that occurs with increasing temperature. (*Medium confidence*)

Recommended Citation for Chapter

Kossin, J.P., T. Hall, T. Knutson, K.E. Kunkel, R.J. Trapp, D.E. Waliser, and M.F. Wehner, 2017: Extreme storms. In: *Climate Science Special Report: Fourth National Climate Assessment, Volume I* [Wuebbles, D.J., D.W. Fahey, K.A. Hibbard, D.J. Dokken, B.C. Stewart, and T.K. Maycock (eds.)]. U.S. Global Change Research Program, Washington, DC, USA, pp. 257-276, doi: 10.7930/J07S7KXX.

9.1 Introduction

Extreme storms have numerous impacts on lives and property. Quantifying how broad-scale average climate influences the behavior of extreme storms is particularly challenging, in part because extreme storms are comparatively rare short-lived events and occur within an environment of largely random variability. Additionally, because the physical mechanisms linking climate change and extreme storms can manifest in a variety of ways, even the sign of the changes in the extreme storms can vary in a warming climate. This makes detection and attribution of trends in extreme storm characteristics more difficult than detection and attribution of trends in the larger environment in which the storms evolve (e.g., Ch. 6: Temperature Change). Projecting changes in severe storms is also challenging because of model constraints in how they capture and represent small-scale, highly local physics. Despite the challenges, good progress is being made for a variety of storm types, such as tropical cyclones, severe convective storms (thunderstorms), winter storms, and atmospheric river events.

9.2 Tropical Cyclones (Hurricanes and Typhoons)

Detection and attribution (Ch. 3: Detection and Attribution) of past changes in tropical cyclone (TC) behavior remain a challenge due to the nature of the historical data, which are highly heterogeneous in both time and among the various regions that collect and analyze the data.^{1, 2, 3} While there are ongoing efforts to reanalyze and homogenize the data (e.g., Landsea et al. 2015;⁴ Kossin et al. 2013²), there is still low confidence that any reported long-term (multidecadal to centennial) increases in TC activity are robust, after accounting for past changes in observing capabilities [which is unchanged from the Intergovernmental Panel on Climate Change Fifth Assessment Report (IPCC AR5) assessment statement⁵]. This is not meant to

imply that no such increases have occurred, but rather that the data are not of a high enough quality to determine this with much confidence. Furthermore, it has been argued that within the period of highest data quality (since around 1980), the globally observed changes in the environment would not necessarily support a detectable trend in tropical cyclone intensity.² That is, the trend signal has not yet had time to rise above the background variability of natural processes.

Both theory and numerical modeling simulations (in general) indicate an increase in TC intensity in a warmer world, and the models generally show an increase in the number of very intense TCs.^{6, 7, 8, 9, 10} In some cases, climate models can be used to make attribution statements about TCs without formal detection (see also Ch. 3: Detection and Attribution). For example, there is evidence that, in addition to the effects of El Niño, anthropogenic forcing made the extremely active 2014 Hawaiian hurricane season substantially more likely, although no significant rising trend in TC frequency near Hawai'i was detected.¹¹

Changes in frequency and intensity are not the only measures of TC behavior that may be affected by climate variability and change, and there is evidence that the locations where TCs reach their peak intensity has migrated poleward over the past 30 years in the Northern and Southern Hemispheres, apparently in concert with environmental changes associated with the independently observed expansion of the tropics.¹² The poleward migration in the western North Pacific,¹³ which includes a number of U.S. territories, appears particularly consistent among the various available TC datasets and remains significant over the past 60–70 years after accounting for the known modes of natural variability in the region (Figure 9.1). The migration, which can substantially change patterns of TC hazard exposure and



mortality risk, is also evident in 21st century Coupled Model Intercomparison Project Phase 5 (CMIP5) projections following the RCP8.5 emissions trajectories, suggesting a possible link to human activities. Further analysis comparing observed past TC behavior with climate model historical forcing runs (and with model control runs simulating multidecadal internal climate variability alone) are needed to better understand this process, but it is expected that this will be an area of heightened future research.

In the Atlantic, observed multidecadal variability of the ocean and atmosphere, which TCs are shown to respond to, has been ascribed (Ch. 3: Detection and Attribution) to natural internal variability via meridional overturning ocean circulation changes,¹⁴ natural external variability caused by volcanic eruptions^{15,16} and Saharan dust outbreaks,^{17,18} and anthropogenic ex-

ternal forcing via greenhouse gases and sulfate aerosols.^{19,20,21} Determining the relative contributions of each mechanism to the observed multidecadal variability in the Atlantic, and even whether natural or anthropogenic factors have dominated, is presently a very active area of research and debate, and no consensus has yet been reached.^{22,23,24,25,26,27} Despite the level of disagreement about the relative magnitude of human influences, there is broad agreement that human factors have had an impact on the observed oceanic and atmospheric variability in the North Atlantic, and there is *medium confidence* that this has contributed to the observed increase in hurricane activity since the 1970s. This is essentially unchanged from the IPCC AR5 statement,⁶ although the post-AR5 literature has only served to further support this statement.²⁸ This is expected to remain an active research topic in the foreseeable future.

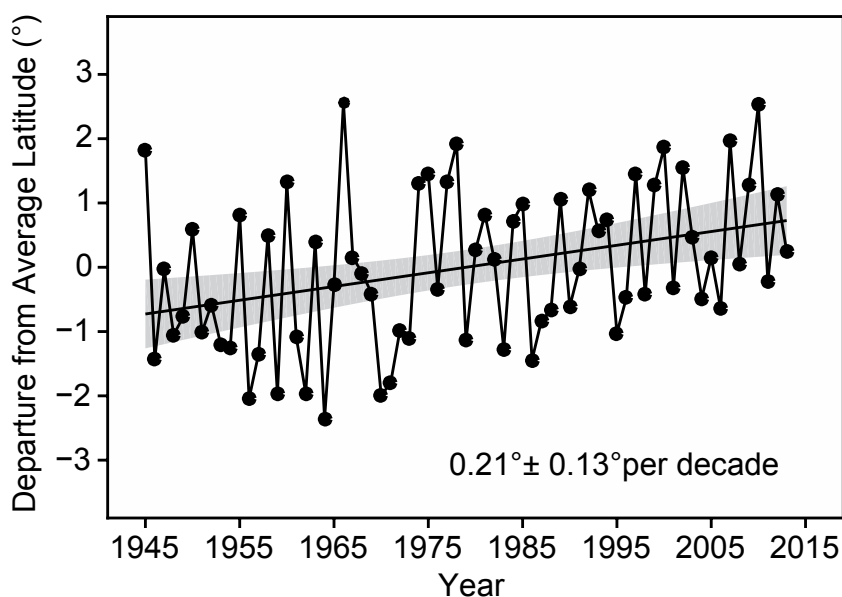


Figure 9.1: Poleward migration, in degrees of latitude, of the location of annual mean tropical cyclone (TC) peak lifetime intensity in the western North Pacific Ocean, after accounting for the known regional modes of interannual (El Niño–Southern Oscillation; ENSO) and interdecadal (Pacific Decadal Oscillation; PDO) variability. The time series shows residuals of the multivariate regression of annually averaged latitude of TC peak lifetime intensity onto the mean Niño-3.4 and PDO indices. Data are taken from the Joint Typhoon Warning Center (JTWC). Shading shows 95% confidence bounds for the trend. Annotated values at lower right show the mean migration rate and its 95% confidence interval in degrees per decade for the period 1945–2013. (Figure source: adapted from Kossin et al. 2016;¹³ © American Meteorological Society. Used with permission.)

The IPCC AR5 consensus TC projections for the late 21st century (IPCC Figure 14.17)⁸ include an increase in global mean TC intensity, precipitation rate, and frequency of very intense (Saffir-Simpson Category 4–5) TCs, and a decrease, or little change, in global TC frequency. Since the IPCC AR5, some studies have provided additional support for this consensus, and some have challenged an aspect of it. For example, a recent study⁹ projects increased mean hurricane intensities in the Atlantic Ocean basin and in most, but not all, other TC-supporting basins (see Table 3 in Knutson et al. 2015⁹). In their study, the global occurrence of Saffir–Simpson Category 4–5 storms was projected to increase significantly, with the most significant basin-scale changes projected for the Northeast Pacific basin, potentially increasing intense hurricane risk to Hawai‘i (Figure 9.2) over the coming century. However, another recent (post-AR5) study proposed that increased thermal stratification of the upper ocean in CMIP5 climate warming scenarios should substantially reduce the warming-induced intensification of TCs estimated in previous studies.²⁹ Follow-up studies, however, estimate that the effect of such increased stratification is relatively small, reducing the projected intensification of TCs by only about 10%–15%.^{30,31}

Another recent study challenged the IPCC AR5 consensus projection of a decrease, or little change, in global tropical cyclone frequency by simulating increased global TC frequency over the 21st century under the higher scenario (RCP8.5).³² However, another modeling study has found that neither direct analysis of CMIP5-class simulations, nor indirect inferences from the simulations (such as those of Emanuel 2013³²), could reproduce the decrease in TC frequency projected in a warmer world by high-resolution TC-permitting climate models,³³ which adds uncertainty to the results of Emanuel.³²

In summary, despite new research that challenges one aspect of the AR5 consensus for late 21st century-projected TC activity, it remains *likely* that global mean tropical cyclone maximum wind speeds and precipitation rates will increase; and it is *more likely than not* that the global frequency of occurrence of TCs will either decrease or remain essentially the same. Confidence in projected global increases of intensity and tropical cyclone precipitation rates is *medium* and *high*, respectively, as there is better model consensus. Confidence is further heightened, particularly for projected increases in precipitation rates, by a robust physical understanding of the processes that lead to these increases. Confidence in projected increases in the frequency of very intense TCs is generally lower (*medium* in the eastern North Pacific and *low* in the western North Pacific and Atlantic) due to comparatively fewer studies available and due to the competing influences of projected reductions in overall storm frequency and increased mean intensity on the frequency of the most intense storms. Both the magnitude and sign of projected changes in individual ocean basins appears to depend on the large-scale pattern of changes to atmospheric circulation and ocean surface temperature (e.g., Knutson et al. 2015⁹). Projections of these regional patterns of change—apparently critical for TC projections—are uncertain, leading to uncertainty in regional TC projections.



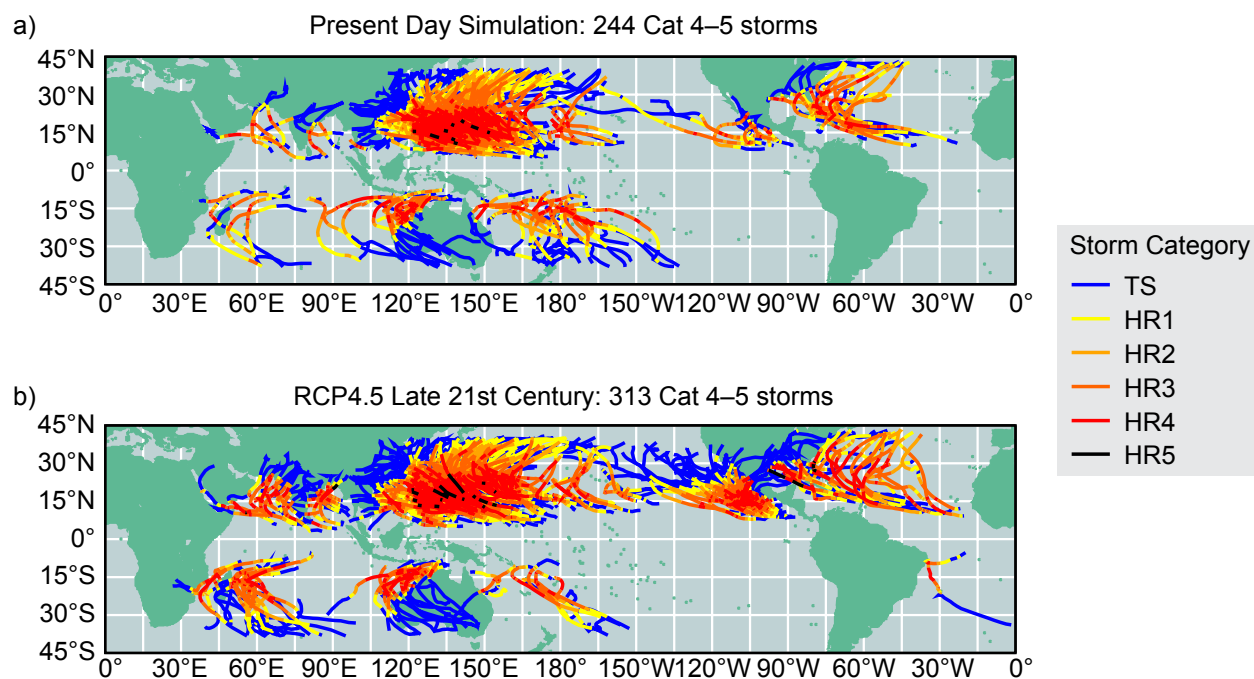


Figure 9.2: Tracks of simulated Saffir–Simpson Category 4–5 tropical cyclones for (a) present-day or (b) late-21st-century conditions, based on dynamical downscaling of climate conditions from the CMIP5 multimodel ensemble (lower scenario; RCP4.5). The tropical cyclones were initially simulated using a 50-km grid global atmospheric model, but each individual tropical cyclone was re-simulated at higher resolution using the GFDL hurricane model to provide more realistic storm intensities and structure. Storm categories or intensities are shown over the lifetime of each simulated storm, according to the Saffir–Simpson scale. The categories are depicted by the track colors, varying from tropical storm (blue) to Category 5 (black; see legend). (Figure source: Knutson et al. 2015;⁹ © American Meteorological Society. Used with permission.)

Box 9.1: U.S. Landfalling Major Hurricane “Drought”

Hurricane Harvey made landfall as a major hurricane (Saffir–Simpson Category 3 or higher) in Texas in 2017, breaking what has sometimes been colloquially referred to as the “hurricane drought.” Prior to Harvey, the last major hurricane to make landfall in the continental United States was Wilma in 2005. The 11-year (2006–2016) absence of U.S. major hurricane landfall events is unprecedented in the historical records dating back to the mid-19th century and has occurred in tandem with average to above-average basin-wide major hurricane counts. Was the 11-year absence of U.S. landfalling major hurricanes due to random luck, or were there systematic changes in climate that drove this?

One recent study indicates that the absence of U.S. landfalling major hurricanes cannot readily be attributed to any sustained changes in the climate patterns that affect hurricanes.³⁴ Based on a statistical analysis of the historical North Atlantic hurricane database, the study found no evidence of a connection between the number of major U.S. landfalls from one year to the next and concluded that the 11-year absence of U.S. landfalling major hurricanes was random. A subsequent recent study did identify a systematic pattern of atmosphere/ocean conditions that vary in such a way that conditions conducive to hurricane intensification in the deep tropics occur in concert with conditions conducive to weakening near the U.S. coast.³⁵ This result suggests a possible relationship between climate and hurricanes; increasing basin-wide hurricane counts are associated with a decreasing fraction of major hurricanes making U.S. landfall, as major hurricanes approaching the U.S. coast are more likely to

Box 9.1 (continued)

weaken during active North Atlantic hurricane periods (such as the present period). It is unclear to what degree this relationship has affected absolute hurricane landfall counts during the recent active hurricane period from the mid-1990s, as the basin-wide number and landfalling fraction are in opposition (that is, there are more major hurricanes but a smaller fraction make landfall as major hurricanes). It is also unclear how this relationship may change as the climate continues to warm. Other studies have identified systematic interdecadal hurricane track variability that may affect landfalling hurricane and major hurricane frequency.^{36, 37, 38}

Another recent study³⁹ shows that the extent of the absence is sensitive to uncertainties in the historical data and even small variations in the definition of a major hurricane, which is somewhat arbitrary. It is also sensitive to the definition of U.S. landfall, which is a geopolitical-border-based constraint and has no physical meaning. In fact, many areas outside of the U.S. border have experienced major hurricane landfalls in the past 11 years. In this sense, the frequency of U.S. landfalling major hurricanes is not a particularly robust metric with which to study questions about hurricane activity and its relationship with climate variability. Furthermore, the 11-year absence of U.S. landfalling major hurricanes is not a particularly relevant metric in terms of coastal hazard exposure and risk. For example, Hurricanes Ike (2008), Irene (2011), Sandy (2012), and most recently Hurricane Matthew (2016) brought severe impacts to the U.S. coast despite not making landfall in the United States while classified as major hurricanes. In the case of Hurricane Sandy, extreme rainfall and storm surge (see also Ch. 12: Sea Level Rise) during landfall caused extensive destruction in and around the New York City area, despite Sandy's designation as a post-tropical cyclone at that time. In the case of Hurricane Matthew, the center came within about 40 miles of the Florida coast while Matthew was a major hurricane, which is close enough to significantly impact the coast but not close enough to break the "drought" as it is defined.

In summary, the absence of U.S. landfalling major hurricanes from Wilma in 2005 to Harvey in 2017 was anomalous. There is some evidence that systematic atmosphere/ocean variability has reduced the fraction of hurricanes making U.S. landfall since the mid-1990s, but this is at least partly countered by increased basin-wide numbers, and the net effect on landfall rates is unclear. Moreover, there is a large random element, and the metric itself suffers from lack of physical basis due to the arbitrary intensity threshold and geopolitically based constraints. Additionally, U.S. coastal risk, particularly from storm surge and freshwater flooding, depends strongly on storm size, propagation speed and direction, and rainfall rates. There is some danger, in the form of evoking complacency, in placing too much emphasis on an absence of a specific subset of hurricanes.

9.3 Severe Convective Storms (Thunderstorms)

Tornado and severe thunderstorm events cause significant loss of life and property: more than one-third of the \$1 billion weather disasters in the United States during the past 25 years were due to such events, and, relative to other extreme weather, the damages from convective weather hazards have undergone the largest increase since 1980.⁴⁰ A particular challenge in quantifying the existence and intensity of these events arises from the data

source: rather than measurements, the occurrence of tornadoes and severe thunderstorms is determined by visual sightings by eye-witnesses (such as "storm spotters" and law enforcement officials) or post-storm damage assessments. The reporting has been susceptible to changes in population density, modifications to reporting procedures and training, the introduction of video and social media, and so on. These have led to systematic, non-meteorological biases in the long-term data record.

Nonetheless, judicious use of the report database has revealed important information about tornado trends. Since the 1970s, the United States has experienced a decrease in the number of days per year on which tornadoes occur, but an increase in the number of tornadoes that form on such days.⁴¹ One important implication is that the frequency of days with large numbers of tornadoes—tornado outbreaks—appears to be increasing (Figure 9.3). The extent of the season over which such tornado activity occurs is increasing as well: although tornadoes in the United States are observed in all months of the year, an earlier calendar-day start to the season of high activity is emerging. In general, there is more interannual variability, or volatility, in tornado occurrence (see also Elsner et al. 2015⁴²).⁴³

Evaluations of hail and (non-tornadic) thunderstorm wind reports have thus far been less revealing. Although there is evidence of an increase in the number of hail days per year, the inherent uncertainty in reported hail size reduces the confidence in such a conclusion.⁴⁴

Thunderstorm wind reports have proven to be even less reliable, because, as compared to tornadoes and hail, there is less tangible visual evidence; thus, although the United States has lately experienced several significant thunderstorm wind events (sometimes referred to as “derechos”), the lack of studies that explore long-term trends in wind events and the uncertainties in the historical data preclude any robust assessment.

It is possible to bypass the use of reports by exploiting the fact that the temperature, humidity, and wind in the larger vicinity—or “environment”—of a developing thunderstorm ultimately control the intensity, morphology, and hazardous tendency of the storm. Thus, the premise is that quantifications of the vertical profiles of temperature, humidity, and wind can be used as a proxy for actual severe thunderstorm occurrence. In particular, a thresholded product of convective available potential energy (CAPE) and vertical wind shear over a surface-to-6 km layer (S06) constitutes one widely used means of

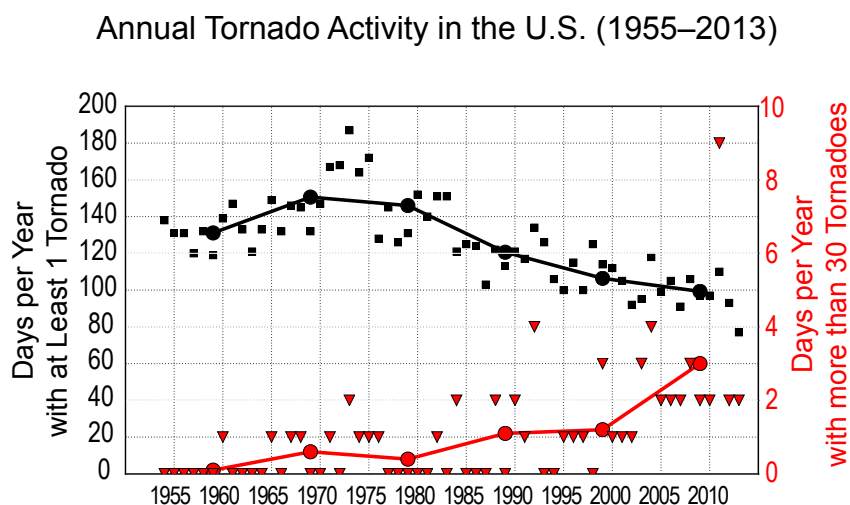


Figure 9.3: Annual tornado activity in the United States over the period 1955–2013. The black squares indicate the number of days per year with at least one tornado rated (E)F1 or greater, and the black circles and line show the decadal mean line of such *tornado days*. The red triangles indicate the number of days per year with more than 30 tornadoes rated (E)F1 or greater, and the red circles and line show the decadal mean of these *tornado outbreaks*. (Figure source: redrawn from Brooks et al. 2014⁴¹).

representing the frequency of severe thunderstorms.⁴⁵ This environmental-proxy approach avoids the biases and other issues with eyewitness storm reports and is readily evaluated using the relatively coarse global datasets and global climate models. It has the disadvantage of assuming that a thunderstorm will necessarily form and then realize its environmental potential.

Upon employing global climate models (GCMs) to evaluate CAPE and S06, a consistent finding among a growing number of proxy-based studies is a projected increase in the frequency of severe thunderstorm environments in the United States over the mid- to late 21st century.^{46, 47, 48, 49, 50, 51} The most robust projected increases in frequency are over the U.S. Midwest and southern Great Plains, during March-April-May (MAM).⁴⁶ Based on the increased frequency of very high CAPE, increases in storm intensity are also projected over this same period (see also Del Genio et al. 2007⁵²).

Key limitations of the environmental proxy approach are being addressed through the applications of high-resolution dynamical downscaling, wherein sufficiently fine model grids are used so that individual thunderstorms are explicitly resolved, rather than implicitly represented (as through environmental proxies). The individually modeled thunderstorms can then be quantified and assessed in terms of severity.^{53, 54, 55} The dynamical-downscaling results have thus far supported the basic findings of the environmental proxy studies, particularly in terms of the seasons and geographical regions projected to experience the largest increases in severe thunderstorm occurrence.⁴⁶

The computational expense of high-resolution dynamical downscaling makes it difficult to generate model ensembles over long time

periods, and thus to assess the uncertainty of the downscaled projections. Because these dynamical downscaling implementations focus on the statistics of storm occurrence rather than on faithful representations of individual events, they have generally been unconcerned with specific extreme convective events in history. So, for example, such downscaling does not address whether the intensity of an event like the Joplin, Missouri, tornado of May 22, 2011, would be amplified under projected future climates. Recently, the “pseudo-global warming” (PGW) methodology (see Schär et al. 1996⁵⁶), which is a variant of dynamical downscaling, has been adapted to address these and related questions. As an example, when the parent “supercell” of select historical tornado events forms under the climate conditions projected during the late 21st century, it does not evolve into a benign, unorganized thunderstorm but instead maintains its supercellular structure.⁵⁷ As measured by updraft strength, the intensity of these supercells under PGW is relatively higher, although not in proportion to the theoretical intensity based on the projected higher levels of CAPE. The adverse effects of enhanced precipitation loading under PGW has been offered as one possible explanation for such shortfalls in projected updraft strength.

9.4 Winter Storms

The frequency of large snowfall years has decreased in the southern United States and Pacific Northwest and increased in the northern United States (see Ch. 7: Precipitation Change). The winters of 2013/2014 and 2014/2015 have contributed to this trend. They were characterized by frequent storms and heavier-than-normal snowfalls in the Midwest and Northeast and drought in the western United States. These were related to blocking (a large-scale pressure pattern with little or no movement) of the wintertime circulation in the Pacific sector of the Northern



Hemisphere (e.g., Marinaro et al. 2015⁵⁸) that put the midwestern and northeastern United States in the primary winter storm track, while at the same time reducing the number of winter storms in California, causing severe drought conditions.⁵⁹ While some observational studies suggest a linkage between blocking affecting the U.S. climate and enhanced arctic warming (arctic amplification), specifically for an increase in highly amplified jet stream patterns in winter over the United States,⁶⁰ other studies show mixed results.^{61, 62, 63} Therefore, a definitive understanding of the effects of arctic amplification on midlatitude winter weather remains elusive. Other explanations have been offered for the weather patterns of recent winters, such as anomalously strong Pacific trade winds,⁶⁴ but these have not been linked to anthropogenic forcing (e.g., Delworth et al. 2015⁶⁵).

Analysis of storm tracks indicates that there has been an increase in winter storm frequency and intensity since 1950, with a slight shift in tracks toward the poles.^{66, 67, 68} Current global climate models (CMIP5) do in fact predict an increase in extratropical cyclone (ETC) frequency over the eastern United States, including the most intense ETCs, under the higher scenario (RCP8.5).⁶⁹ However, there are large model-to-model differences in the realism of ETC simulations and in the projected changes. Moreover, projected ETC changes have large regional variations, including a decreased total frequency in the North Atlantic, further highlighting the complexity of the response to climate change.

9.5 Atmospheric Rivers

The term “atmospheric rivers” (ARs) refers to the relatively narrow streams of moisture transport that often occur within and across midlatitudes⁷⁰ (Figure 9.4), in part because they often transport as much water as in the Amazon River.⁷¹ While ARs occupy less than 10% of the circumference of Earth at any given time, they account for 90% of the poleward moisture transport across midlatitudes (a more complete discussion of precipitation variability is found in Ch. 7: Precipitation Change). In many regions of the world, they account for a substantial fraction of the precipitation,⁷² and thus water supply, often delivered in the form of an extreme weather and precipitation event (Figure 9.4). For example, ARs account for 30%–40% of the typical snowpack in the Sierra Nevada mountains and annual precipitation in the U.S. West Coast states^{73, 74}—an essential summertime source of water for agriculture, consumption, and ecosystem health. However, this vital source of water is also associated with severe flooding—with observational evidence showing a close connection between historically high streamflow events and floods with landfalling AR events—in the west and other sectors of the United States.^{75, 76, 77} More recently, research has also demonstrated that ARs are often found to be critical in ending droughts in the western United States.⁷⁸



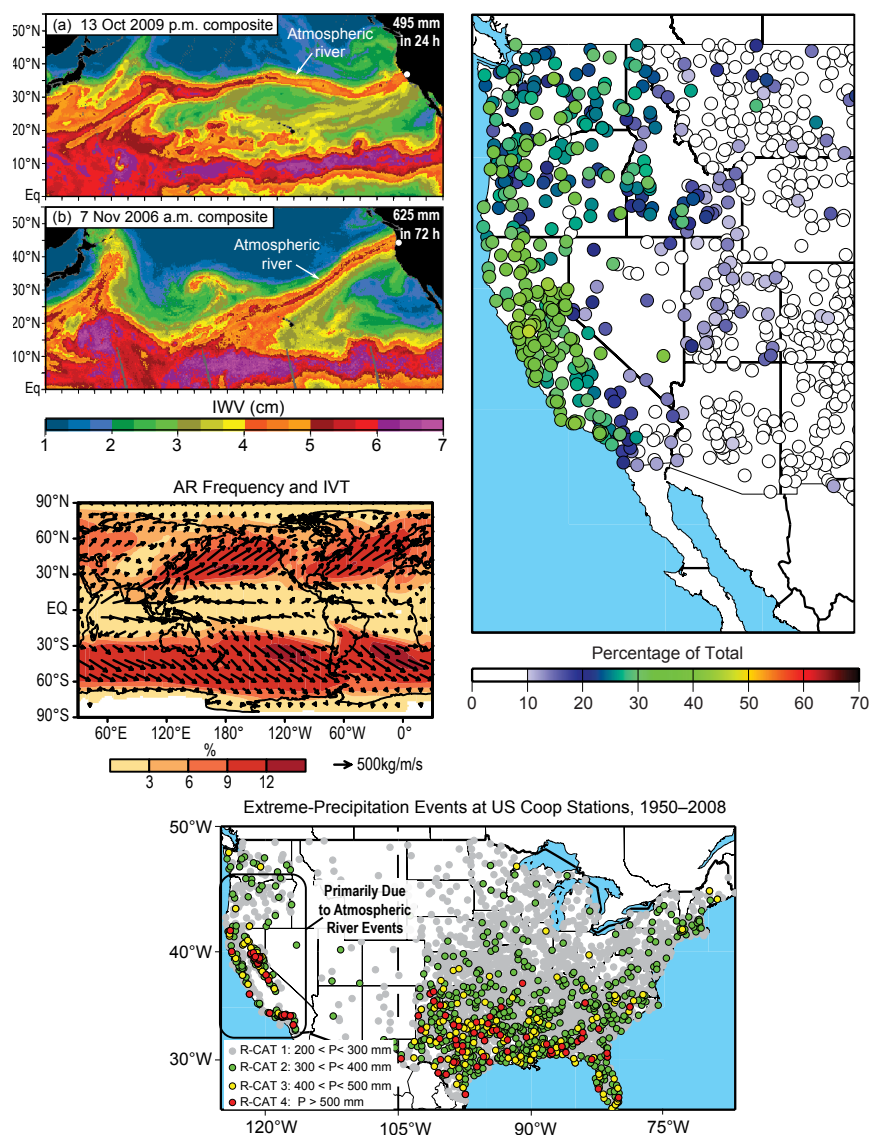


Figure 9.4: (upper left) Atmospheric rivers depicted in Special Sensor Microwave Imager (SSM/I) measurements of SSM/I total column water vapor leading to extreme precipitation events at landfall locations. (middle left) Annual mean frequency of atmospheric river occurrence (for example, 12% means about 1 every 8 days) and their integrated vapor transport (IVT).⁷² (bottom) ARs are the dominant synoptic storms for the U.S. West Coast in terms of extreme precipitation⁹³ and (right) supply a large fraction of the annual precipitation in the U.S. West Coast states.⁷³ [Figure source: (upper and middle left) Ralph et al. 2011,⁹⁴ (upper right) Guan and Waliser 2015,⁷² (lower left) Ralph and Dettinger 2012,⁹³ (lower right) Dettinger et al. 2011;⁷³ left panels, © American Meteorological Society. Used with permission.]

Given the important role that ARs play in the water supply of the western United States and their role in weather and water extremes in the west and occasionally other parts of the United States (e.g., Rutz et al. 2014⁷⁹), it is critical to examine how climate change and the expected intensification of the global water cycle and atmospheric transports (e.g., Held and Soden 2006;⁸⁰ Lavers et al. 2015⁸¹) are projected to impact ARs (e.g., Dettinger and Ingram 2013⁸²).

Under climate change conditions, ARs may be altered in a number of ways, namely their frequency, intensity, duration, and locations. In association with landfalling ARs, any of these would be expected to result in impacts on hazards and water supply given the discussion above. Assessments of ARs in climate change projections for the United States have been undertaken for central California from CMIP3,⁷³ and a number of studies have been

done for the West Coast of North America,^{83, 84, 85, 86, 87} and these studies have uniformly shown that ARs are likely to become more frequent and intense in the future. For example, one recent study reveals a large increase of AR days along the West Coast by the end of the 21st century under the higher scenario (RCP8.5), with fractional increases between 50% and 600%, depending on the seasons and landfall locations.⁸³ Results from these studies (and Lavers et al. 2013⁸⁸ for ARs impacting the United Kingdom) show that these AR changes were predominantly driven by increasing atmospheric specific humidity, with little discernible change in the low-level winds. The higher atmospheric water vapor content in a warmer climate is to be expected because of an increase in saturation water vapor pressure with air temperature (Ch. 2: Physical Drivers of Climate Change). While the thermodynamic effect appears to dominate the climate change impact on ARs, leading to projected increases in ARs, there is evidence for a dynamical effect (that is, location change) related to the projected poleward shift of the subtropical jet that diminished the thermodynamic effect in the southern portion of the West Coast of North America.⁸³

Presently, there is no clear consensus on whether the consistently projected increases in AR frequency and intensity will translate to increased precipitation in California. This is mostly because previous studies did not examine this explicitly and because the model resolution is poor and thus the topography is poorly represented, and the topography is a key aspect of forcing the precipitation out of the systems.⁸⁹ The evidence for considerable increases in the number and intensity of ARs depends (as do all climate variability studies based on dynamical models) on the model fidelity in representing ARs and their interactions with the global climate/circulation. Additional confidence comes from studies that

show qualitatively similar projected increases while also providing evidence that the models represent AR frequency, transports, and spatial distributions relatively well compared to observations.^{84, 85} A caveat associated with drawing conclusions from any given study or differences between two is that they typically use different detection methodologies that are typically tailored to a regional setting (cf. Guan and Waliser 2015⁷²). Additional research is warranted to examine these storms from a global perspective, with additional and more in-depth, process-oriented diagnostics/metrics. Stepping away from the sensitivities associated with defining atmospheric rivers, one study examined the intensification of the integrated vapor transport (IVT), which is easily and unambiguously defined.⁸¹ That study found that for the higher scenario (RCP8.5), multimodel mean IVT and the IVT associated with extremes above 95% percentile increase by 30%–40% in the North Pacific. These results, along with the uniform findings of the studies above examining projected changes in ARs for western North America and the United Kingdom, give *high confidence* that the frequency of AR storms will increase in association with rising global temperatures.



TRACEABLE ACCOUNTS

Key Finding 1

Human activities have contributed substantially to observed ocean–atmosphere variability in the Atlantic Ocean (*medium confidence*), and these changes have contributed to the observed upward trend in North Atlantic hurricane activity since the 1970s (*medium confidence*).

Description of evidence base

The Key Finding and supporting text summarizes extensive evidence documented in the climate science literature and is similar to statements made in previous national (NCA3)⁹⁰ and international⁹¹ assessments. Data limitations are documented in Kossin et al. 2013² and references therein. Contributions of natural and anthropogenic factors in observed multidecadal variability are quantified in Carlsaw et al. 2013;²² Zhang et al. 2013;²⁷ Tung and Zhou 2013;²⁶ Mann et al. 2014;²³ Stevens 2015;²⁵ Sobel et al. 2016;²⁴ Walsh et al. 2015.¹⁰

Major uncertainties

Key remaining uncertainties are due to known and substantial heterogeneities in the historical tropical cyclone data and lack of robust consensus in determining the precise relative contributions of natural and anthropogenic factors in past variability of the tropical environment.

Assessment of confidence based on evidence and agreement, including short description of nature of evidence and level of agreement

Confidence in this finding is rated as *medium*. Although the range of estimates of natural versus anthropogenic contributions in the literature is fairly broad, virtually all studies identify a measurable, and generally substantial, anthropogenic influence. This does constitute a consensus for human contribution to the increases in tropical cyclone activity since 1970.

Summary sentence or paragraph that integrates the above information

The key message and supporting text summarizes extensive evidence documented in the climate science peer-reviewed literature. The uncertainties and points

of consensus that were described in the NCA3 and IPCC assessments have continued.

Key Finding 2

Both theory and numerical modeling simulations generally indicate an increase in tropical cyclone (TC) intensity in a warmer world, and the models generally show an increase in the number of very intense TCs. For Atlantic and eastern North Pacific hurricanes and western North Pacific typhoons, increases are projected in precipitation rates (*high confidence*) and intensity (*medium confidence*). The frequency of the most intense of these storms is projected to increase in the Atlantic and western North Pacific (*low confidence*) and in the eastern North Pacific (*medium confidence*).

Description of evidence base

The Key Finding and supporting text summarizes extensive evidence documented in the climate science literature and is similar to statements made in previous national (NCA3)⁹⁰ and international⁹¹ assessments. Since these assessments, more recent downscaling studies have further supported these assessments (e.g., Knutson et al. 2015⁹), though pointing out that the changes (future increased intensity and tropical cyclone precipitation rates) may not occur in all ocean basins.

Major uncertainties

A key uncertainty remains in the lack of a supporting detectable anthropogenic signal in the historical data to add further confidence to these projections. As such, confidence in the projections is based on agreement among different modeling studies and physical understanding (for example, potential intensity theory for tropical cyclone intensities and the expectation of stronger moisture convergence, and thus higher precipitation rates, in tropical cyclones in a warmer environment containing greater amounts of environmental atmospheric moisture). Additional uncertainty stems from uncertainty in both the projected pattern and magnitude of future sea surface temperatures.⁹



Assessment of confidence based on evidence and agreement, including short description of nature of evidence and level of agreement

Confidence is rated as *high* in tropical cyclone rainfall projections and *medium* in intensity projections since there are a number of publications supporting these overall conclusions, fairly well-established theory, general consistency among different studies, varying methods used in studies, and still a fairly strong consensus among studies. However, a limiting factor for confidence in the results is the lack of a supporting detectable anthropogenic contribution in observed tropical cyclone data.

There is *low* to *medium confidence* for increased occurrence of the most intense tropical cyclones for most ocean basins, as there are relatively few formal studies that focus on these changes, and the change in occurrence of such storms would be enhanced by increased intensities, but reduced by decreased overall frequency of tropical cyclones.

Summary sentence or paragraph that integrates the above information

Models are generally in agreement that tropical cyclones will be more intense and have higher precipitation rates, at least in most ocean basins. Given the agreement between models and support of theory and mechanistic understanding, there is *medium* to *high confidence* in the overall projection, although there is some limitation on confidence levels due to the lack of a supporting detectable anthropogenic contribution to tropical cyclone intensities or precipitation rates.

Key Finding 3

Tornado activity in the United States has become more variable, particularly over the 2000s, with a decrease in the number of days per year with tornadoes and an increase in the number of tornadoes on these days (*medium confidence*). Confidence in past trends for hail and severe thunderstorm winds, however, is *low*. Climate models consistently project environmental changes that would putatively support an increase in the frequency and intensity of severe thunderstorms (a category that combines tornadoes, hail, and winds), especially over

regions that are currently prone to these hazards, but confidence in the details of this projected increase is *low*.

Description of evidence base

Evidence for the first and second statement comes from the U.S. database of tornado reports. There are well known biases in this database, but application of an intensity threshold [greater than or equal to a rating of 1 on the (Enhanced) Fujita scale], and the quantification of tornado activity in terms of tornado days instead of raw numbers of reports are thought to reduce these biases. It is not known at this time whether the variability and trends are necessarily due to climate change.

The third statement is based on projections from a wide range of climate models, including GCMs and RCMs, run over the past 10 years (e.g., see the review by Brooks 2013⁹²). The evidence is derived from an “environmental-proxy” approach, which herein means that severe thunderstorm occurrence is related to the occurrence of two key environmental parameters: CAPE and vertical wind shear. A limitation of this approach is the assumption that the thunderstorm will necessarily form and then realize its environmental potential. This assumption is indeed violated, albeit at levels that vary by region and season.

Major uncertainties

Regarding the first and second statements, there is still some uncertainty in the database, even when the data are filtered. The major uncertainty in the third statement equates to the aforementioned limitation (that is, the thunderstorm will necessarily form and then realize its environmental potential).

Assessment of confidence based on evidence and agreement, including short description of nature of evidence and level of agreement

Medium: That the variability in tornado activity has increased.

Medium: That the severe-thunderstorm environmental conditions will change with a changing climate, but

Low: on the precise (geographical and seasonal) realization of the environmental conditions as actual severe thunderstorms.



Summary sentence or paragraph that integrates the above information

With an established understanding of the data biases, careful analysis provides useful information about past changes in severe thunderstorm and tornado activity. This information suggests that tornado variability has increased in the 2000s, with a concurrent decrease in the number of days per year experiencing tornadoes and an increase in the number of tornadoes on these days. Similarly, the development of novel applications of climate models provides information about possible future severe storm and tornado activity, and although confidence in these projections is low, they do suggest that the projected environments are at least consistent with environments that would putatively support an increase in frequency and intensity of severe thunderstorms.

Key Finding 4

There has been a trend toward earlier snowmelt and a decrease in snowstorm frequency on the southern margins of climatologically snowy areas (*medium confidence*). Winter storm tracks have shifted northward since 1950 over the Northern Hemisphere (*medium confidence*). Projections of winter storm frequency and intensity over the United States vary from increasing to decreasing depending on region, but model agreement is poor and confidence is *low*. Potential linkages between the frequency and intensity of severe winter storms in the United States and accelerated warming in the Arctic have been postulated, but they are complex, and, to some extent, contested, and confidence in the connection is currently *low*.

Description of evidence base

The Key Finding and supporting text summarizes evidence documented in the climate science literature.

Evidence for changes in winter storm track changes are documented in a small number of studies.^{67, 68} Future changes are documented in one study,⁶⁹ but there are large model-to-model differences. The effects of arctic amplification on U.S. winter storms have been studied, but the results are mixed,^{60, 61, 62, 63} leading to considerable uncertainties.

Major uncertainties

Key remaining uncertainties relate to the sensitivity of observed snow changes to the spatial distribution of observing stations and to historical changes in station location and observing practices. There is conflicting evidence about the effects of arctic amplification on CONUS winter weather.

Assessment of confidence based on evidence and agreement, including short description of nature of evidence and level of agreement

There is *high confidence* that warming has resulted in earlier snowmelt and decreased snowfall on the warm margins of areas with consistent snowpack based on a number of observational studies. There is *medium confidence* that Northern Hemisphere storm tracks have shifted north based on a small number of studies. There is *low confidence* in future changes in winter storm frequency and intensity based on conflicting evidence from analysis of climate model simulations.

Summary sentence or paragraph that integrates the above information

Decreases in snowfall on southern and low elevation margins of currently climatologically snowy areas are likely but winter storm frequency and intensity changes are uncertain.

Key Finding 5

The frequency and severity of landfalling “atmospheric rivers” on the U.S. West Coast (narrow streams of moisture that account for 30%–40% of the typical snowpack and annual precipitation in the region and are associated with severe flooding events) will increase as a result of increasing evaporation and resulting higher atmospheric water vapor that occurs with increasing temperature (*medium confidence*).

Description of evidence base

The Key Finding and supporting text summarizes evidence documented in the climate science literature.

Evidence for the expectation of an increase in the frequency and severity of landfalling atmospheric rivers on the U.S. West Coast comes from the CMIP-based

climate change projection studies of Dettinger et al. 2011;⁷³ Warner et al. 2015;⁸⁷ Payne and Magnusdottir 2015;⁸⁵ Gao et al. 2015;⁸³ Radić et al. 2015;⁸⁶ and Hagos et al. 2016.⁸⁴ The close connection between atmospheric rivers and water availability and flooding is based on the present-day observation studies of Guan et al. 2010;⁷⁴ Dettinger et al. 2011;⁷³ Ralph et al. 2006;⁷⁷ Neiman et al. 2011;⁷⁶ Moore et al. 2012;⁷⁵ and Dettinger 2013.⁷⁸

Major uncertainties

A modest uncertainty remains in the lack of a supporting detectable anthropogenic signal in the historical data to add further confidence to these projections. However, the overall increase in atmospheric rivers projected/expected is based to a very large degree on the *very high confidence* that the atmospheric water vapor will increase. Thus, increasing water vapor coupled with little projected change in wind structure/intensity still indicates increases in the frequency/intensity of atmospheric rivers. A modest uncertainty arises in quantifying the expected change at a regional level (for example, northern Oregon vs. southern Oregon) given that there are some changes expected in the position of the jet stream that might influence the degree of increase for different locations along the West Coast. Uncertainty in the projections of the number and intensity of ARs is introduced by uncertainties in the models' ability to represent ARs and their interactions with climate.

Assessment of confidence based on evidence and agreement, including short description of nature of evidence and level of agreement

Confidence in this finding is rated as *medium* based on qualitatively similar projections among different studies.

Summary sentence or paragraph that integrates the above information

Increases in atmospheric river frequency and intensity are expected along the U.S. West Coast, leading to the likelihood of more frequent flooding conditions, with uncertainties remaining in the details of the spatial structure of these along the coast (for example, northern vs. southern California).



REFERENCES

1. Klotzbach, P.J. and C.W. Landsea, 2015: Extremely intense hurricanes: Revisiting Webster et al. (2005) after 10 years. *Journal of Climate*, **28**, 7621-7629. <http://dx.doi.org/10.1175/JCLI-D-15-0188.1>
2. Kossin, J.P., T.L. Olander, and K.R. Knapp, 2013: Trend analysis with a new global record of tropical cyclone intensity. *Journal of Climate*, **26**, 9960-9976. <http://dx.doi.org/10.1175/JCLI-D-13-00262.1>
3. Walsh, K.J.E., J.L. McBride, P.J. Klotzbach, S. Balachandran, S.J. Camargo, G. Holland, T.R. Knutson, J.P. Kossin, T.-c. Lee, A. Sobel, and M. Sugi, 2016: Tropical cyclones and climate change. *Wiley Interdisciplinary Reviews: Climate Change*, **7**, 65-89. <http://dx.doi.org/10.1002/wcc.371>
4. Landsea, C., J. Franklin, and J. Beven, 2015: The revised Atlantic hurricane database (HURDAT2). National Hurricane Center, Miami, FL.
5. Hartmann, D.L., A.M.G. Klein Tank, M. Rusticucci, L.V. Alexander, S. Brönnimann, Y. Charabi, F.J. Dentener, E.J. Dlugokencky, D.R. Easterling, A. Kaplan, B.J. Soden, P.W. Thorne, M. Wild, and P.M. Zhai, 2013: Observations: Atmosphere and surface. *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, and P.M. Midgley, Eds. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 159-254. <http://www.climatechange2013.org/report/full-report/>
6. Bindoff, N.L., P.A. Stott, K.M. AchutaRao, M.R. Allen, N. Gillett, D. Gutzler, K. Hansingo, G. Hegerl, Y. Hu, S. Jain, I.I. Mokhov, J. Overland, J. Perlwitz, R. Sebbari, and X. Zhang, 2013: Detection and attribution of climate change: From global to regional. *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, and P.M. Midgley, Eds. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 867-952. <http://www.climatechange2013.org/report/full-report/>
7. Camargo, S.J., 2013: Global and regional aspects of tropical cyclone activity in the CMIP5 models. *Journal of Climate*, **26**, 9880-9902. <http://dx.doi.org/10.1175/jcli-d-12-00549.1>
8. Christensen, J.H., K. Krishna Kumar, E. Aldrian, S.-I. An, I.F.A. Cavalcanti, M. de Castro, W. Dong, P. Goswami, A. Hall, J.K. Kanyanga, A. Kitoh, J. Kossin, N.-C. Lau, J. Renwick, D.B. Stephenson, S.-P. Xie, and T. Zhou, 2013: Climate phenomena and their relevance for future regional climate change. *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex, and P.M. Midgley, Eds. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 1217-1308. <http://www.climatechange2013.org/report/full-report/>
9. Knutson, T.R., J.J. Sirutis, M. Zhao, R.E. Tuleya, M. Bender, G.A. Vecchi, G. Villarini, and D. Chavas, 2015: Global projections of intense tropical cyclone activity for the late twenty-first century from dynamical downscaling of CMIP5/RCP4.5 scenarios. *Journal of Climate*, **28**, 7203-7224. <http://dx.doi.org/10.1175/JCLI-D-15-0129.1>
10. Walsh, K.J.E., S.J. Camargo, G.A. Vecchi, A.S. Daloz, J. Elsner, K. Emanuel, M. Horn, Y.-K. Lim, M. Roberts, C. Patricola, E. Scoccimarro, A.H. Sobel, S. Strazzo, G. Villarini, M. Wehner, M. Zhao, J.P. Kossin, T. LaRow, K. Oouchi, S. Schubert, H. Wang, J. Bacmeister, P. Chang, F. Chauvin, C. Jablonowski, A. Kumar, H. Murakami, T. Ose, K.A. Reed, R. Saravanan, Y. Yamada, C.M. Zarzycki, P.L. Vidale, J.A. Jonas, and N. Henderson, 2015: Hurricanes and climate: The U.S. CLIVAR Working Group on Hurricanes. *Bulletin of the American Meteorological Society*, **96** (12), 997-1017. <http://dx.doi.org/10.1175/BAMS-D-13-00242.1>
11. Murakami, H., G.A. Vecchi, T.L. Delworth, K. Paffen-dorf, L. Jia, R. Gudgel, and F. Zeng, 2015: Investigating the influence of anthropogenic forcing and natural variability on the 2014 Hawaiian hurricane season [in "Explaining Extreme Events of 2014 from a Climate Perspective"]. *Bulletin of the American Meteorological Society*, **96** (12), S115-S119. <http://dx.doi.org/10.1175/BAMS-D-15-00119.1>
12. Kossin, J.P., K.A. Emanuel, and G.A. Vecchi, 2014: The poleward migration of the location of tropical cyclone maximum intensity. *Nature*, **509**, 349-352. <http://dx.doi.org/10.1038/nature13278>
13. Kossin, J.P., K.A. Emanuel, and S.J. Camargo, 2016: Past and projected changes in western North Pacific tropical cyclone exposure. *Journal of Climate*, **29**, 5725-5739. <http://dx.doi.org/10.1175/JCLI-D-16-0076.1>
14. Delworth, L.T. and E.M. Mann, 2000: Observed and simulated multidecadal variability in the Northern Hemisphere. *Climate Dynamics*, **16**, 661-676. <http://dx.doi.org/10.1007/s003820000075>

15. Evan, A.T., 2012: Atlantic hurricane activity following two major volcanic eruptions. *Journal of Geophysical Research*, **117**, D06101. <http://dx.doi.org/10.1029/2011JD016716>
16. Thompson, D.W.J. and S. Solomon, 2009: Understanding recent stratospheric climate change. *Journal of Climate*, **22**, 1934-1943. <http://dx.doi.org/10.1175/2008JCLI2482.1>
17. Evan, A.T., G.R. Foltz, D. Zhang, and D.J. Vimont, 2011: Influence of African dust on ocean-atmosphere variability in the tropical Atlantic. *Nature Geoscience*, **4**, 762-765. <http://dx.doi.org/10.1038/ngeo1276>
18. Evan, A.T., D.J. Vimont, A.K. Heidinger, J.P. Kossin, and R. Bennartz, 2009: The role of aerosols in the evolution of tropical North Atlantic Ocean temperature anomalies. *Science*, **324**, 778-781. <http://dx.doi.org/10.1126/science.1167404>
19. Booth, B.B.B., N.J. Dunstone, P.R. Halloran, T. Andrews, and N. Bellouin, 2012: Aerosols implicated as a prime driver of twentieth-century North Atlantic climate variability. *Nature*, **484**, 228-232. <http://dx.doi.org/10.1038/nature10946>
20. Dunstone, N.J., D.M. Smith, B.B.B. Booth, L. Hermanson, and R. Eade, 2013: Anthropogenic aerosol forcing of Atlantic tropical storms. *Nature Geoscience*, **6**, 534-539. <http://dx.doi.org/10.1038/ngeo1854>
21. Mann, M.E. and K.A. Emanuel, 2006: Atlantic hurricane trends linked to climate change. *Eos, Transactions, American Geophysical Union*, **87**, 233-244. <http://dx.doi.org/10.1029/2006EO240001>
22. Carslaw, K.S., L.A. Lee, C.L. Reddington, K.J. Pringle, A. Rap, P.M. Forster, G.W. Mann, D.V. Spracklen, M.T. Woodhouse, L.A. Regayre, and J.R. Pierce, 2013: Large contribution of natural aerosols to uncertainty in indirect forcing. *Nature*, **503**, 67-71. <http://dx.doi.org/10.1038/nature12674>
23. Mann, M.E., B.A. Steinman, and S.K. Miller, 2014: On forced temperature changes, internal variability, and the AMO. *Geophysical Research Letters*, **41**, 3211-3219. <http://dx.doi.org/10.1002/2014GL059233>
24. Sobel, A.H., S.J. Camargo, T.M. Hall, C.-Y. Lee, M.K. Tippett, and A.A. Wing, 2016: Human influence on tropical cyclone intensity. *Science*, **353**, 242-246. <http://dx.doi.org/10.1126/science.aaf6574>
25. Stevens, B., 2015: Rethinking the lower bound on aerosol radiative forcing. *Journal of Climate*, **28**, 4794-4819. <http://dx.doi.org/10.1175/JCLI-D-14-00656.1>
26. Tung, K.-K. and J. Zhou, 2013: Using data to attribute episodes of warming and cooling in instrumental records. *Proceedings of the National Academy of Sciences*, **110**, 2058-2063. <http://dx.doi.org/10.1073/pnas.1212471110>
27. Zhang, R., T.L. Delworth, R. Sutton, D.L.R. Hodson, K.W. Dixon, I.M. Held, Y. Kushnir, J. Marshall, Y. Ming, R. Msadek, J. Robson, A.J. Rosati, M. Ting, and G.A. Vecchi, 2013: Have aerosols caused the observed Atlantic multidecadal variability? *Journal of the Atmospheric Sciences*, **70**, 1135-1144. <http://dx.doi.org/10.1175/jas-d-12-0331.1>
28. Kossin, J.P., T.R. Karl, T.R. Knutson, K.A. Emanuel, K.E. Kunkel, and J.J. O'Brien, 2015: Reply to "Comments on 'Monitoring and understanding trends in extreme storms: State of knowledge'". *Bulletin of the American Meteorological Society*, **96** (12), 1177-1179. <http://dx.doi.org/10.1175/BAMS-D-14-00261.1>
29. Huang, P., I.I. Lin, C. Chou, and R.-H. Huang, 2015: Change in ocean subsurface environment to suppress tropical cyclone intensification under global warming. *Nature Communications*, **6**, 7188. <http://dx.doi.org/10.1038/ncomms8188>
30. Emanuel, K., 2015: Effect of upper-ocean evolution on projected trends in tropical cyclone activity. *Journal of Climate*, **28**, 8165-8170. <http://dx.doi.org/10.1175/JCLI-D-15-0401.1>
31. Tuleya, R.E., M. Bender, T.R. Knutson, J.J. Sirutis, B. Thomas, and I. Ginis, 2016: Impact of upper-tropospheric temperature anomalies and vertical wind shear on tropical cyclone evolution using an idealized version of the operational GFDL hurricane model. *Journal of the Atmospheric Sciences*, **73**, 3803-3820. <http://dx.doi.org/10.1175/JAS-D-16-0045.1>
32. Emanuel, K.A., 2013: Downscaling CMIP5 climate models shows increased tropical cyclone activity over the 21st century. *Proceedings of the National Academy of Sciences*, **110**, 12219-12224. <http://dx.doi.org/10.1073/pnas.1301293110>
33. Wehner, M., Prabhat, K.A. Reed, D. Stone, W.D. Collins, and J. Bacmeister, 2015: Resolution dependence of future tropical cyclone projections of CAM5.1 in the U.S. CLIVAR Hurricane Working Group idealized configurations. *Journal of Climate*, **28**, 3905-3925. <http://dx.doi.org/10.1175/JCLI-D-14-00311.1>
34. Hall, T. and K. Hereid, 2015: The frequency and duration of U.S. hurricane droughts. *Geophysical Research Letters*, **42**, 3482-3485. <http://dx.doi.org/10.1002/2015GL063652>
35. Kossin, J.P., 2017: Hurricane intensification along U.S. coast suppressed during active hurricane periods. *Nature*, **541**, 390-393. <http://dx.doi.org/10.1038/nature20783>
36. Colbert, A.J. and B.J. Soden, 2012: Climatological variations in North Atlantic tropical cyclone tracks. *Journal of Climate*, **25**, 657-673. <http://dx.doi.org/10.1175/jcli-d-11-00034.1>

37. Kossin, J.P. and D.J. Vimont, 2007: A more general framework for understanding Atlantic hurricane variability and trends. *Bulletin of the American Meteorological Society*, **88**, 1767-1781. <http://dx.doi.org/10.1175/bams-88-11-1767>
38. Wang, C., H. Liu, S.-K. Lee, and R. Atlas, 2011: Impact of the Atlantic warm pool on United States land-falling hurricanes. *Geophysical Research Letters*, **38**, L19702. <http://dx.doi.org/10.1029/2011gl049265>
39. Hart, R.E., D.R. Chavas, and M.P. Guishard, 2016: The arbitrary definition of the current Atlantic major hurricane landfall drought. *Bulletin of the American Meteorological Society*, **97**, 713-722. <http://dx.doi.org/10.1175/BAMS-D-15-00185.1>
40. Smith, A.B. and R.W. Katz, 2013: U.S. billion-dollar weather and climate disasters: Data sources, trends, accuracy and biases. *Natural Hazards*, **67**, 387-410. <http://dx.doi.org/10.1007/s11069-013-0566-5>
41. Brooks, H.E., G.W. Carbin, and P.T. Marsh, 2014: Increased variability of tornado occurrence in the United States. *Science*, **346**, 349-352. <http://dx.doi.org/10.1126/science.1257460>
42. Elsner, J.B., S.C. Elsner, and T.H. Jagger, 2015: The increasing efficiency of tornado days in the United States. *Climate Dynamics*, **45**, 651-659. <http://dx.doi.org/10.1007/s00382-014-2277-3>
43. Tippett, M.K., 2014: Changing volatility of U.S. annual tornado reports. *Geophysical Research Letters*, **41**, 6956-6961. <http://dx.doi.org/10.1002/2014GL061347>
44. Allen, J.T. and M.K. Tippett, 2015: The Characteristics of United States Hail Reports: 1955-2014. *Electronic Journal of Severe Storms Meteorology*.
45. Brooks, H.E., J.W. Lee, and J.P. Craven, 2003: The spatial distribution of severe thunderstorm and tornado environments from global reanalysis data. *Atmospheric Research*, **67-68**, 73-94. [http://dx.doi.org/10.1016/S0169-8095\(03\)00045-0](http://dx.doi.org/10.1016/S0169-8095(03)00045-0)
46. Diffenbaugh, N.S., M. Scherer, and R.J. Trapp, 2013: Robust increases in severe thunderstorm environments in response to greenhouse forcing. *Proceedings of the National Academy of Sciences*, **110**, 16361-16366. <http://dx.doi.org/10.1073/pnas.1307758110>
47. Gensini, V.A., C. Ramseyer, and T.L. Mote, 2014: Future convective environments using NARCCAP. *International Journal of Climatology*, **34**, 1699-1705. <http://dx.doi.org/10.1002/joc.3769>
48. Seeley, J.T. and D.M. Romps, 2015: The effect of global warming on severe thunderstorms in the United States. *Journal of Climate*, **28**, 2443-2458. <http://dx.doi.org/10.1175/JCLI-D-14-00382.1>
49. Trapp, R.J., N.S. Diffenbaugh, H.E. Brooks, M.E. Baldwin, E.D. Robinson, and J.S. Pal, 2007: Changes in severe thunderstorm environment frequency during the 21st century caused by anthropogenically enhanced global radiative forcing. *Proceedings of the National Academy of Sciences*, **104**, 19719-19723. <http://dx.doi.org/10.1073/pnas.0705494104>
50. Trapp, R.J., N.S. Diffenbaugh, and A. Gluhovsky, 2009: Transient response of severe thunderstorm forcing to elevated greenhouse gas concentrations. *Geophysical Research Letters*, **36**, L01703. <http://dx.doi.org/10.1029/2008GL036203>
51. Van Klooster, S.L. and P.J. Roebber, 2009: Surface-based convective potential in the contiguous United States in a business-as-usual future climate. *Journal of Climate*, **22**, 3317-3330. <http://dx.doi.org/10.1175/2009JCLI2697.1>
52. Del Genio, A.D., M.S. Yao, and J. Jonas, 2007: Will moist convection be stronger in a warmer climate? *Geophysical Research Letters*, **34**, L16703. <http://dx.doi.org/10.1029/2007GL030525>
53. Trapp, R.J., E.D. Robinson, M.E. Baldwin, N.S. Diffenbaugh, and B.R.J. Schwedler, 2011: Regional climate of hazardous convective weather through high-resolution dynamical downscaling. *Climate Dynamics*, **37**, 677-688. <http://dx.doi.org/10.1007/s00382-010-0826-y>
54. Robinson, E.D., R.J. Trapp, and M.E. Baldwin, 2013: The geospatial and temporal distributions of severe thunderstorms from high-resolution dynamical downscaling. *Journal of Applied Meteorology and Climatology*, **52**, 2147-2161. <http://dx.doi.org/10.1175/JAMC-D-12-0131.1>
55. Gensini, V.A. and T.L. Mote, 2014: Estimations of hazardous convective weather in the United States using dynamical downscaling. *Journal of Climate*, **27**, 6581-6589. <http://dx.doi.org/10.1175/JCLI-D-13-00777.1>
56. Schär, C., C. Frei, D. Lüthi, and H.C. Davies, 1996: Surrogate climate-change scenarios for regional climate models. *Geophysical Research Letters*, **23**, 669-672. <http://dx.doi.org/10.1029/96GL00265>
57. Trapp, R.J. and K.A. Hoogewind, 2016: The realization of extreme tornadic storm events under future anthropogenic climate change. *Journal of Climate*, **29**, 5251-5265. <http://dx.doi.org/10.1175/JCLI-D-15-0623.1>
58. Marinaro, A., S. Hilberg, D. Changnon, and J.R. Angel, 2015: The North Pacific-driven severe Midwest winter of 2013/14. *Journal of Applied Meteorology and Climatology*, **54**, 2141-2151. <http://dx.doi.org/10.1175/JAMC-D-15-0084.1>

59. Chang, E.K.M., C. Zheng, P. Lanigan, A.M.W. Yau, and J.D. Neelin, 2015: Significant modulation of variability and projected change in California winter precipitation by extratropical cyclone activity. *Geophysical Research Letters*, **42**, 5983-5991. <http://dx.doi.org/10.1002/2015GL064424>
60. Francis, J. and N. Skific, 2015: Evidence linking rapid Arctic warming to mid-latitude weather patterns. *Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences*, **373**, 20140170. <http://dx.doi.org/10.1098/rsta.2014.0170>
61. Barnes, E.A. and L.M. Polvani, 2015: CMIP5 projections of Arctic amplification, of the North American/North Atlantic circulation, and of their relationship. *Journal of Climate*, **28**, 5254-5271. <http://dx.doi.org/10.1175/JCLI-D-14-00589.1>
62. Perlwitz, J., M. Hoerling, and R. Dole, 2015: Arctic tropospheric warming: Causes and linkages to lower latitudes. *Journal of Climate*, **28**, 2154-2167. <http://dx.doi.org/10.1175/JCLI-D-14-00095.1>
63. Screen, J.A., C. Deser, and L. Sun, 2015: Projected changes in regional climate extremes arising from Arctic sea ice loss. *Environmental Research Letters*, **10**, 084006. <http://dx.doi.org/10.1088/1748-9326/10/8/084006>
64. Yang, X., G.A. Vecchi, T.L. Delworth, K. Paffendorf, L. Jia, R. Gudgel, F. Zeng, and S.D. Underwood, 2015: Extreme North America winter storm season of 2013/14: Roles of radiative forcing and the global warming hiatus [in "Explaining Extreme Events of 2014 from a Climate Perspective"]. *Bulletin of the American Meteorological Society*, **96** (12), S25-S28. <http://dx.doi.org/10.1175/BAMS-D-15-00133.1>
65. Delworth, T.L., F. Zeng, A. Rosati, G.A. Vecchi, and A.T. Wittenberg, 2015: A link between the hiatus in global warming and North American drought. *Journal of Climate*, **28**, 3834-3845. <http://dx.doi.org/10.1175/jcli-d-14-00616.1>
66. Vose, R.S., S. Applequist, M.A. Bourassa, S.C. Pryor, R.J. Barthelmie, B. Blanton, P.D. Bromirski, H.E. Brooks, A.T. DeGaetano, R.M. Dole, D.R. Easterling, R.E. Jensen, T.R. Karl, R.W. Katz, K. Klink, M.C. Kruk, K.E. Kunkel, M.C. MacCracken, T.C. Peterson, K. Shein, B.R. Thomas, J.E. Walsh, X.L. Wang, M.F. Wehner, D.J. Wuebbles, and R.S. Young, 2014: Monitoring and understanding changes in extremes: Extratropical storms, winds, and waves. *Bulletin of the American Meteorological Society*, **95**, 377-386. <http://dx.doi.org/10.1175/BAMS-D-12-00162.1>
67. Wang, X.L., Y. Feng, G.P. Compo, V.R. Swail, F.W. Zwiers, R.J. Allan, and P.D. Sardeshmukh, 2012: Trends and low frequency variability of extra-tropical cyclone activity in the ensemble of twentieth century reanalysis. *Climate Dynamics*, **40**, 2775-2800. <http://dx.doi.org/10.1007/s00382-012-1450-9>
68. Wang, X.L., V.R. Swail, and F.W. Zwiers, 2006: Climatology and changes of extratropical cyclone activity: Comparison of ERA-40 with NCEP-NCAR reanalysis for 1958-2001. *Journal of Climate*, **19**, 3145-3166. <http://dx.doi.org/10.1175/JCLI3781.1>
69. Colle, B.A., Z. Zhang, K.A. Lombardo, E. Chang, P. Liu, and M. Zhang, 2013: Historical evaluation and future prediction of eastern North American and western Atlantic extratropical cyclones in the CMIP5 models during the cool season. *Journal of Climate*, **26**, 6882-6903. <http://dx.doi.org/10.1175/JCLI-D-12-00498.1>
70. Zhu, Y. and R.E. Newell, 1998: A proposed algorithm for moisture fluxes from atmospheric rivers. *Monthly Weather Review*, **126**, 725-735. [http://dx.doi.org/10.1175/1520-0493\(1998\)126<0725:APAFMF>2.0.CO;2](http://dx.doi.org/10.1175/1520-0493(1998)126<0725:APAFMF>2.0.CO;2)
71. Newell, R.E., N.E. Newell, Y. Zhu, and C. Scott, 1992: Tropospheric rivers? – A pilot study. *Geophysical Research Letters*, **19**, 2401-2404. <http://dx.doi.org/10.1029/92GL02916>
72. Guan, B. and D.E. Waliser, 2015: Detection of atmospheric rivers: Evaluation and application of an algorithm for global studies. *Journal of Geophysical Research Atmospheres*, **120**, 12514-12535. <http://dx.doi.org/10.1002/2015JD024257>
73. Dettinger, M.D., F.M. Ralph, T. Das, P.J. Neiman, and D.R. Cayan, 2011: Atmospheric rivers, floods and the water resources of California. *Water*, **3**, 445-478. <http://dx.doi.org/10.3390/w3020445>
74. Guan, B., N.P. Molotch, D.E. Waliser, E.J. Fetzer, and P.J. Neiman, 2010: Extreme snowfall events linked to atmospheric rivers and surface air temperature via satellite measurements. *Geophysical Research Letters*, **37**, L20401. <http://dx.doi.org/10.1029/2010GL044696>
75. Moore, B.J., P.J. Neiman, F.M. Ralph, and F.E. Barthold, 2012: Physical processes associated with heavy flooding rainfall in Nashville, Tennessee, and vicinity during 1–2 May 2010: The role of an atmospheric river and mesoscale convective systems. *Monthly Weather Review*, **140**, 358-378. <http://dx.doi.org/10.1175/MWR-D-11-00126.1>
76. Neiman, P.J., L.J. Schick, F.M. Ralph, M. Hughes, and G.A. Wick, 2011: Flooding in western Washington: The connection to atmospheric rivers. *Journal of Hydrometeorology*, **12**, 1337-1358. <http://dx.doi.org/10.1175/2011JHM1358.1>
77. Ralph, F.M., P.J. Neiman, G.A. Wick, S.I. Gutman, M.D. Dettinger, D.R. Cayan, and A.B. White, 2006: Flooding on California's Russian River: Role of atmospheric rivers. *Geophysical Research Letters*, **33**, L13801. <http://dx.doi.org/10.1029/2006GL026689>
78. Dettinger, M.D., 2013: Atmospheric rivers as drought busters on the U.S. West Coast. *Journal of Hydrometeorology*, **14**, 1721-1732. <http://dx.doi.org/10.1175/JHM-D-13-02.1>

79. Rutz, J.J., W.J. Steenburgh, and F.M. Ralph, 2014: Climatological characteristics of atmospheric rivers and their inland penetration over the western United States. *Monthly Weather Review*, **142**, 905-921. <http://dx.doi.org/10.1175/MWR-D-13-00168.1>
80. Held, I.M. and B.J. Soden, 2006: Robust responses of the hydrological cycle to global warming. *Journal of Climate*, **19**, 5686-5699. <http://dx.doi.org/10.1175/jcli3990.1>
81. Lavers, D.A., F.M. Ralph, D.E. Waliser, A. Gershunov, and M.D. Dettinger, 2015: Climate change intensification of horizontal water vapor transport in CMIP5. *Geophysical Research Letters*, **42**, 5617-5625. <http://dx.doi.org/10.1002/2015GL064672>
82. Dettinger, M.D. and B.L. Ingram, 2013: The coming megafloods. *Scientific American*, **308**, 64-71. <http://dx.doi.org/10.1038/scientificamerican0113-64>
83. Gao, Y., J. Lu, L.R. Leung, Q. Yang, S. Hagos, and Y. Qian, 2015: Dynamical and thermodynamical modulations on future changes of landfalling atmospheric rivers over western North America. *Geophysical Research Letters*, **42**, 7179-7186. <http://dx.doi.org/10.1002/2015GL065435>
84. Hagos, S.M., L.R. Leung, J.-H. Yoon, J. Lu, and Y. Gao, 2016: A projection of changes in landfalling atmospheric river frequency and extreme precipitation over western North America from the Large Ensemble CESM simulations. *Geophysical Research Letters*, **43**, 1357-1363. <http://dx.doi.org/10.1002/2015GL067392>
85. Payne, A.E. and G. Magnusdottir, 2015: An evaluation of atmospheric rivers over the North Pacific in CMIP5 and their response to warming under RCP 8.5. *Journal of Geophysical Research Atmospheres*, **120**, 11,173-11,190. <http://dx.doi.org/10.1002/2015JD023586>
86. Radić, V., A.J. Cannon, B. Menounos, and N. Gi, 2015: Future changes in autumn atmospheric river events in British Columbia, Canada, as projected by CMIP5 global climate models. *Journal of Geophysical Research Atmospheres*, **120**, 9279-9302. <http://dx.doi.org/10.1002/2015JD023279>
87. Warner, M.D., C.F. Mass, and E.P. Salathé Jr., 2015: Changes in winter atmospheric rivers along the North American West Coast in CMIP5 climate models. *Journal of Hydrometeorology*, **16**, 118-128. <http://dx.doi.org/10.1175/JHM-D-14-0080.1>
88. Lavers, D.A., R.P. Allan, G. Villarini, B. Lloyd-Hughes, D.J. Brayshaw, and A.J. Wade, 2013: Future changes in atmospheric rivers and their implications for winter flooding in Britain. *Environmental Research Letters*, **8**, 034010. <http://dx.doi.org/10.1088/1748-9326/8/3/034010>
89. Pierce, D.W., D.R. Cayan, T. Das, E.P. Maurer, N.L. Miller, Y. Bao, M. Kanamitsu, K. Yoshimura, M.A. Snyder, L.C. Sloan, G. Franco, and M. Tyree, 2013: The key role of heavy precipitation events in climate model disagreements of future annual precipitation changes in California. *Journal of Climate*, **26**, 5879-5896. <http://dx.doi.org/10.1175/jcli-d-12-00766.1>
90. Melillo, J.M., T.C. Richmond, and G.W. Yohe, eds., 2014: *Climate Change Impacts in the United States: The Third National Climate Assessment*. U.S. Global Change Research Program: Washington, D.C., 841 pp. <http://dx.doi.org/10.7930/J0Z31WJ2>
91. IPCC, 2013: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press, Cambridge, UK and New York, NY, 1535 pp. <http://www.climatechange2013.org/report/>
92. Brooks, H.E., 2013: Severe thunderstorms and climate change. *Atmospheric Research*, **123**, 129-138. <http://dx.doi.org/10.1016/j.atmosres.2012.04.002>
93. Ralph, F.M. and M.D. Dettinger, 2012: Historical and national perspectives on extreme West Coast precipitation associated with atmospheric rivers during December 2010. *Bulletin of the American Meteorological Society*, **93**, 783-790. <http://dx.doi.org/10.1175/BAMS-D-11-00188.1>
94. Ralph, F.M., P.J. Neiman, G.N. Kiladis, K. Weickmann, and D.W. Reynolds, 2011: A multiscale observational case study of a Pacific atmospheric river exhibiting tropical-extratropical connections and a mesoscale frontal wave. *Monthly Weather Review*, **139**, 1169-1189. <http://dx.doi.org/10.1175/2010mwr3596.1>

