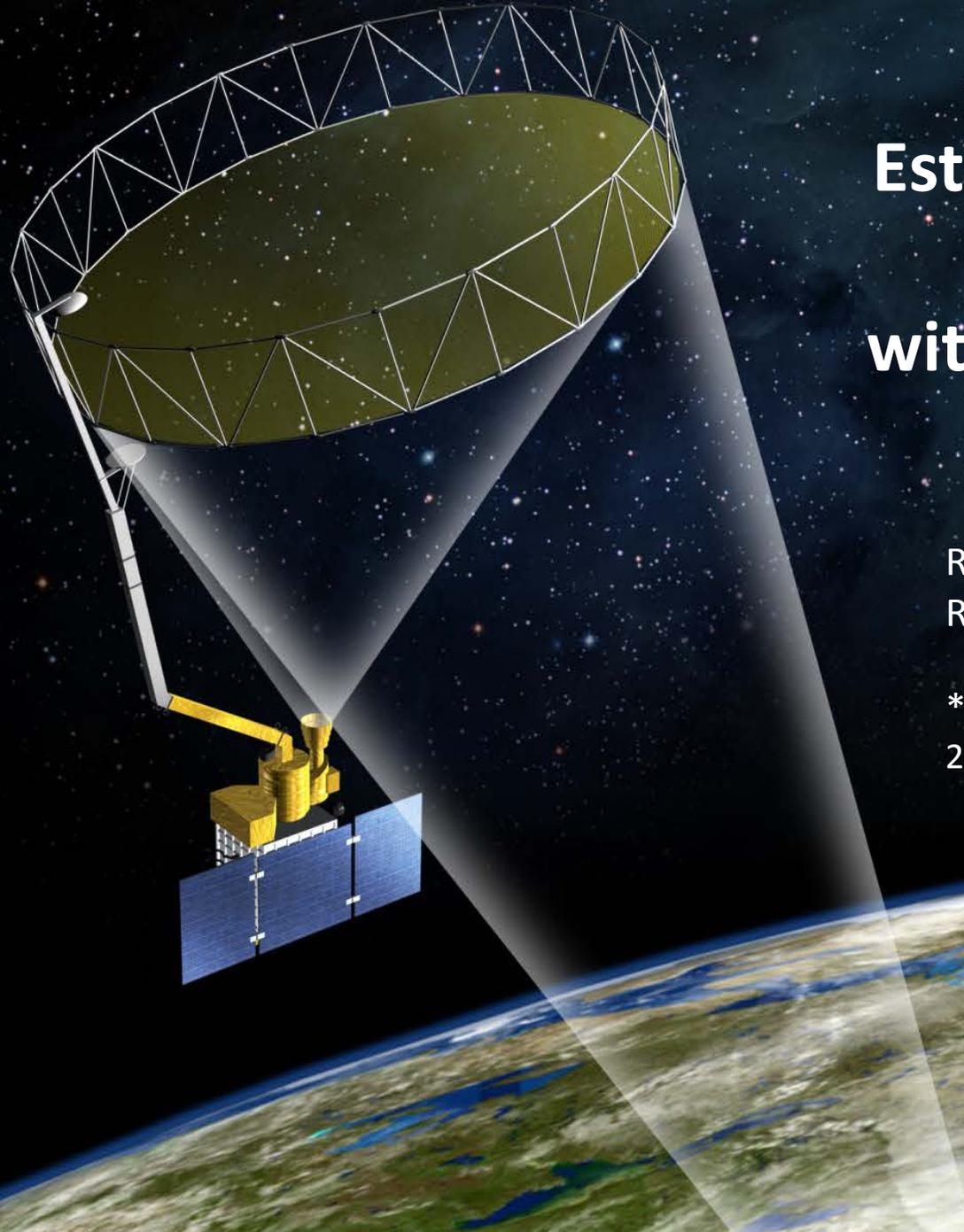


Soil Moisture  
Active Passive  
Mission  
**SMAP**



# Estimating Basin-Scale Water Budgets with SMAP Level 2 Soil Moisture Data

Randal Koster\*, Wade Crow, Rolf Reichle, and Sarith Mahanama

\*GMAO, NASA/GSFC, Greenbelt, MD 20771; [randal.d.koster@nasa.gov](mailto:randal.d.koster@nasa.gov)

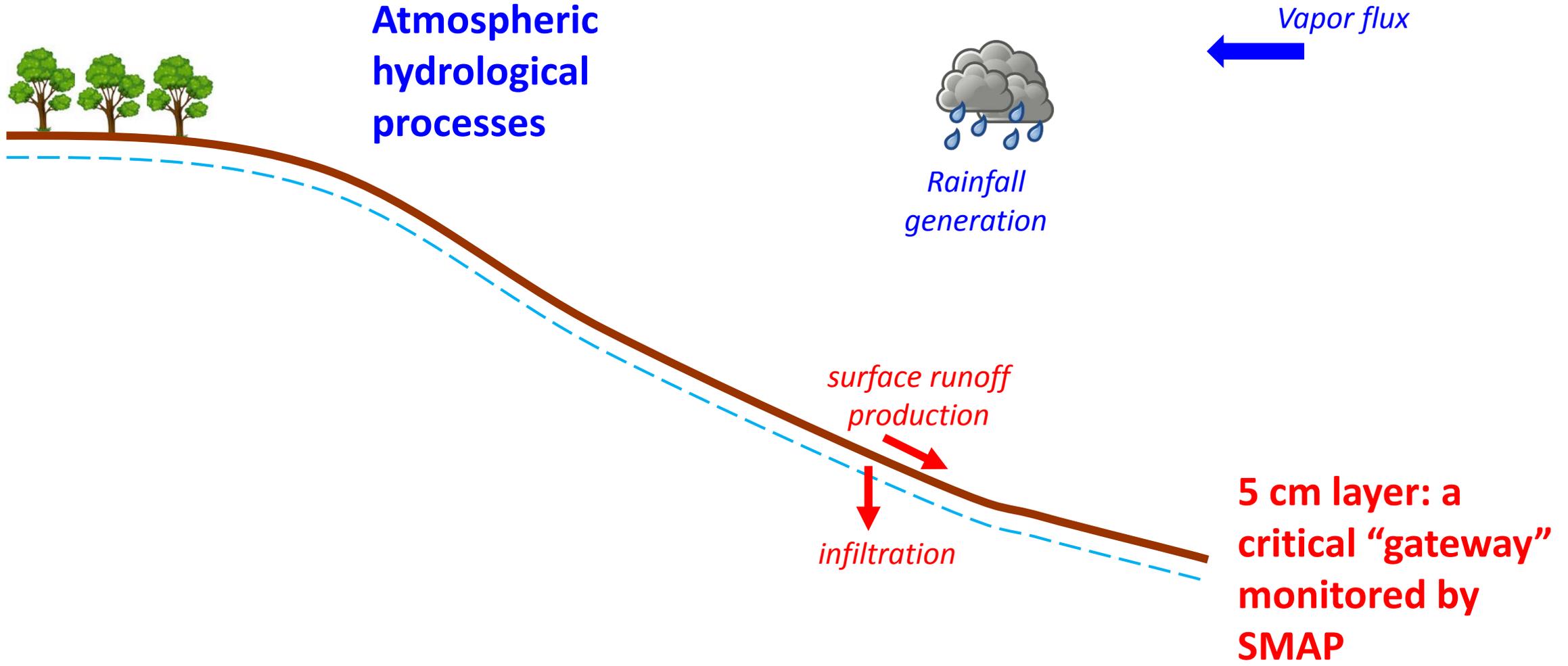
April 18, 2018  
JPL, Pasadena

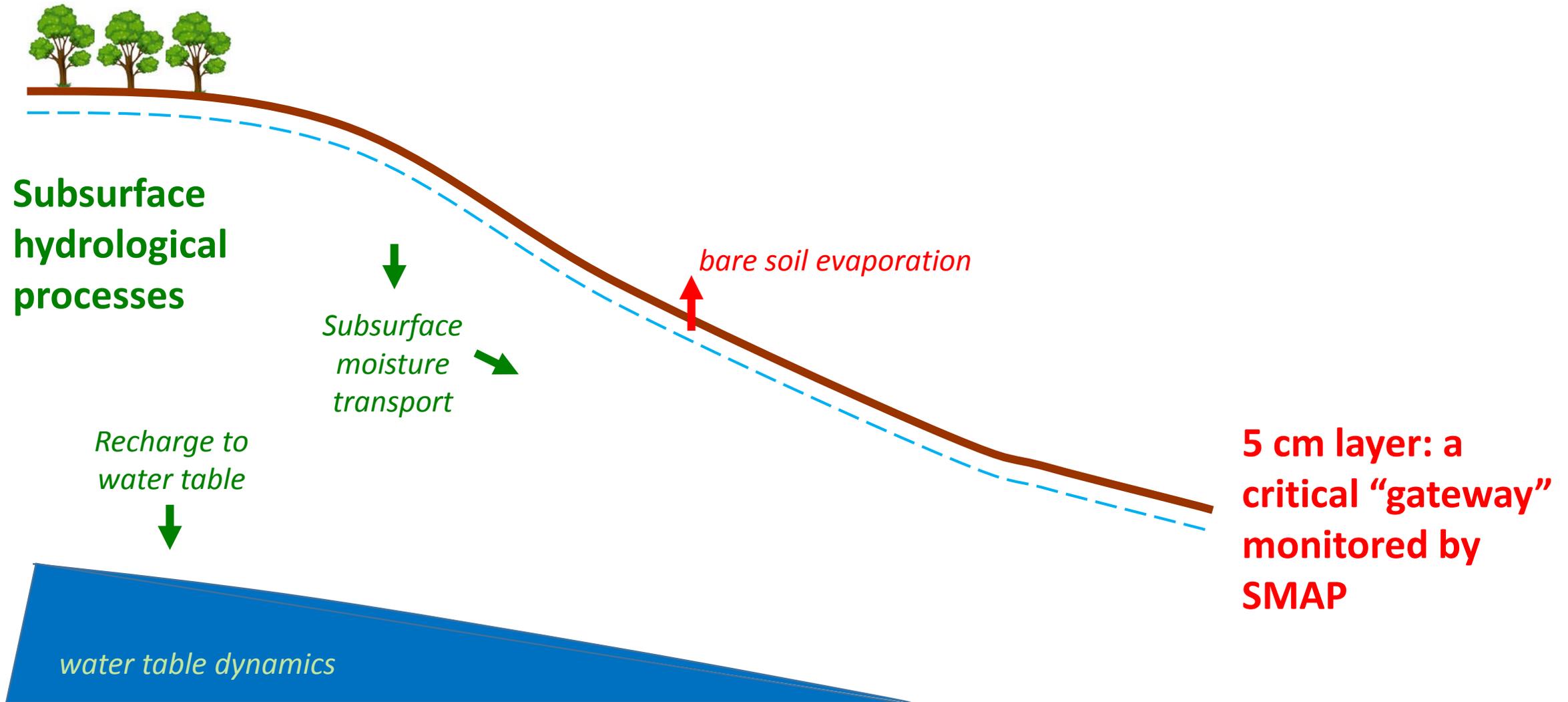


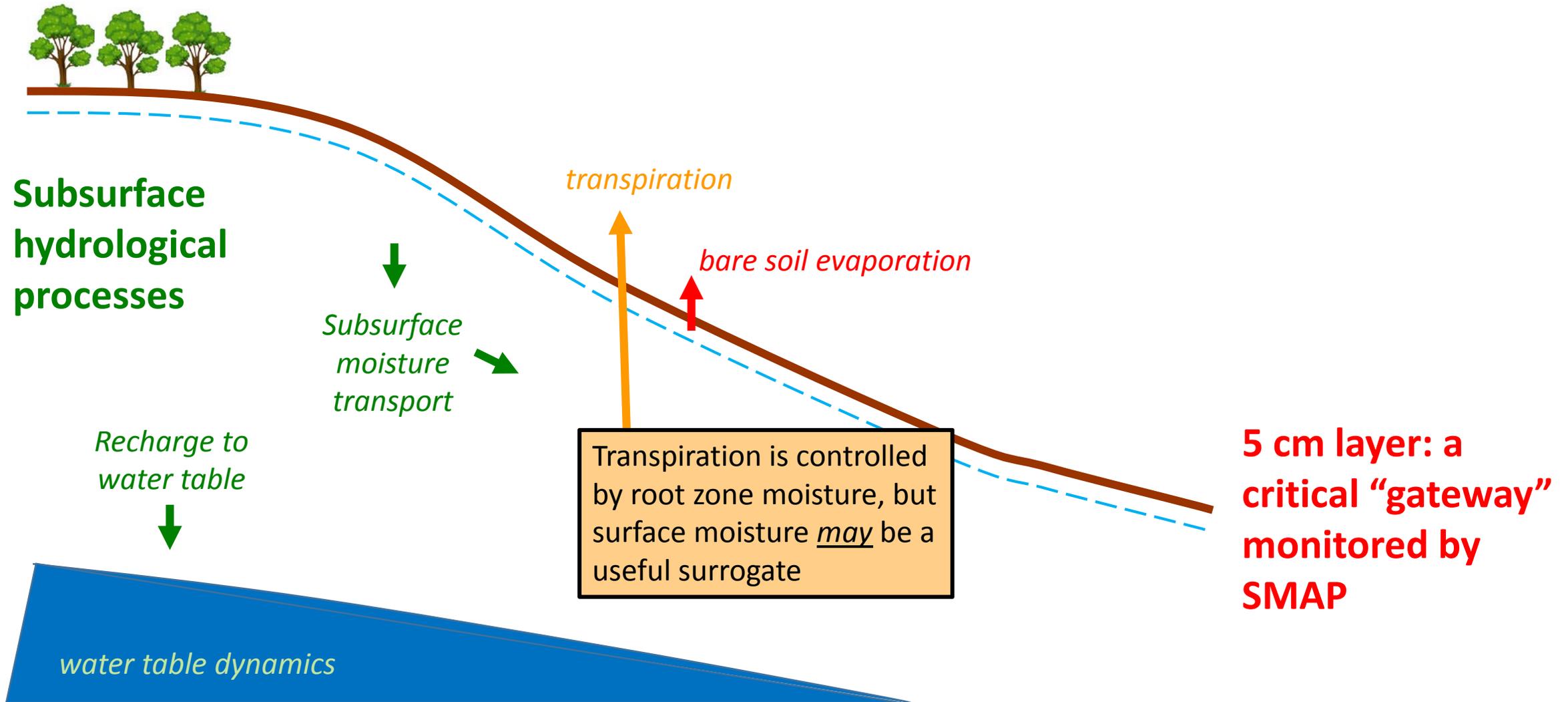
**Atmospheric  
hydrological  
processes**

**Subsurface  
hydrological  
processes**

**5 cm layer: a  
critical “gateway”  
monitored by  
SMAP**

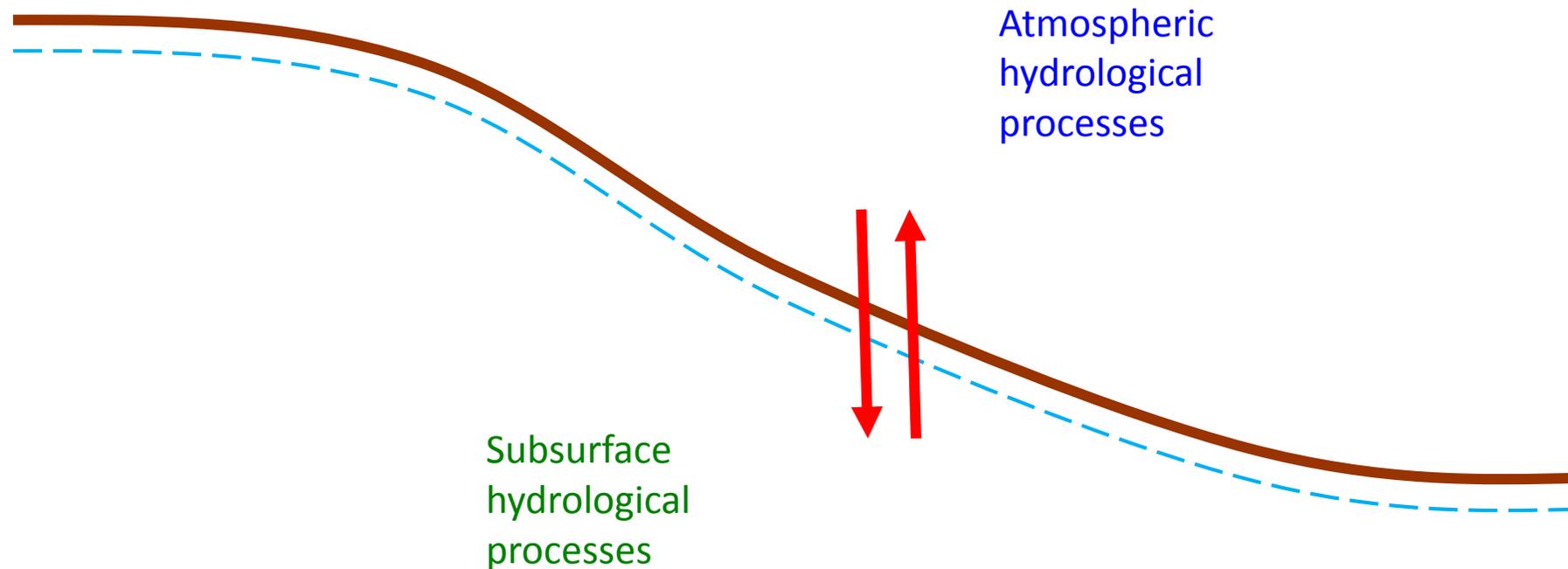




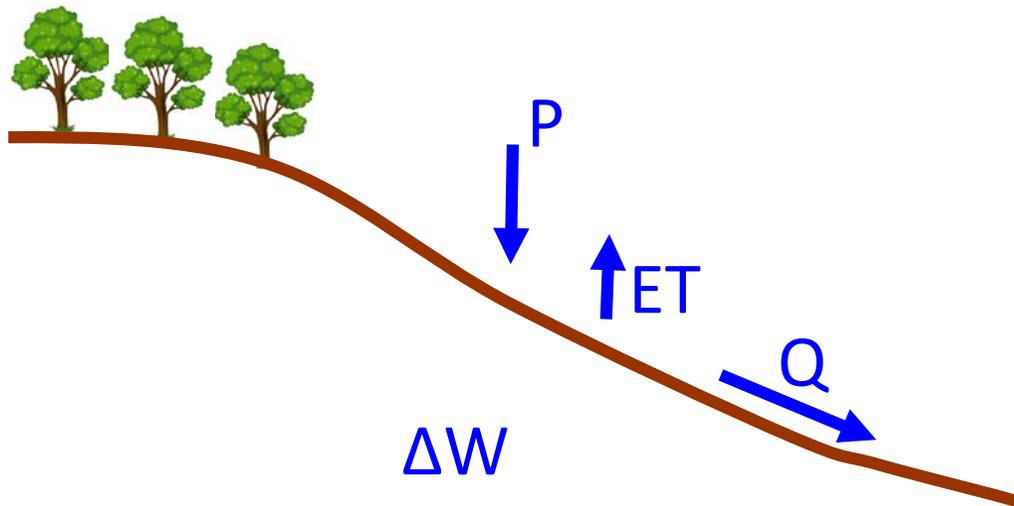


**Key point:** Because SMAP data monitor the “critical gateway”, they contain information on many important hydrological processes.

⇒ *In essence, SMAP does more than just measure soil moisture...*



Sample Hydrological Application: Estimating terms in a basin's water balance.



$$P = ET + Q + \Delta W$$

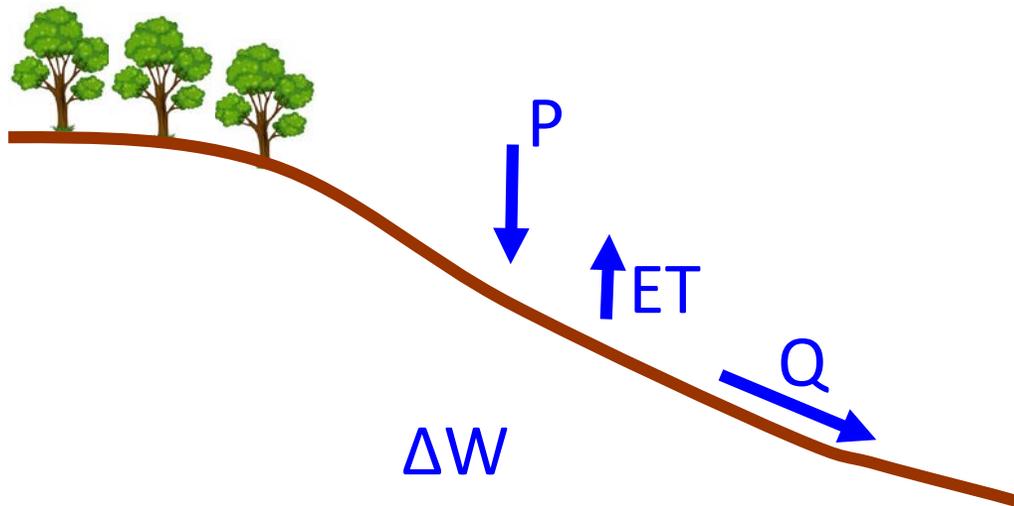
precipitation

runoff (streamflow)

evapotranspiration

change in storage

**Sample Hydrological Application:** Estimating terms in a basin's water balance.



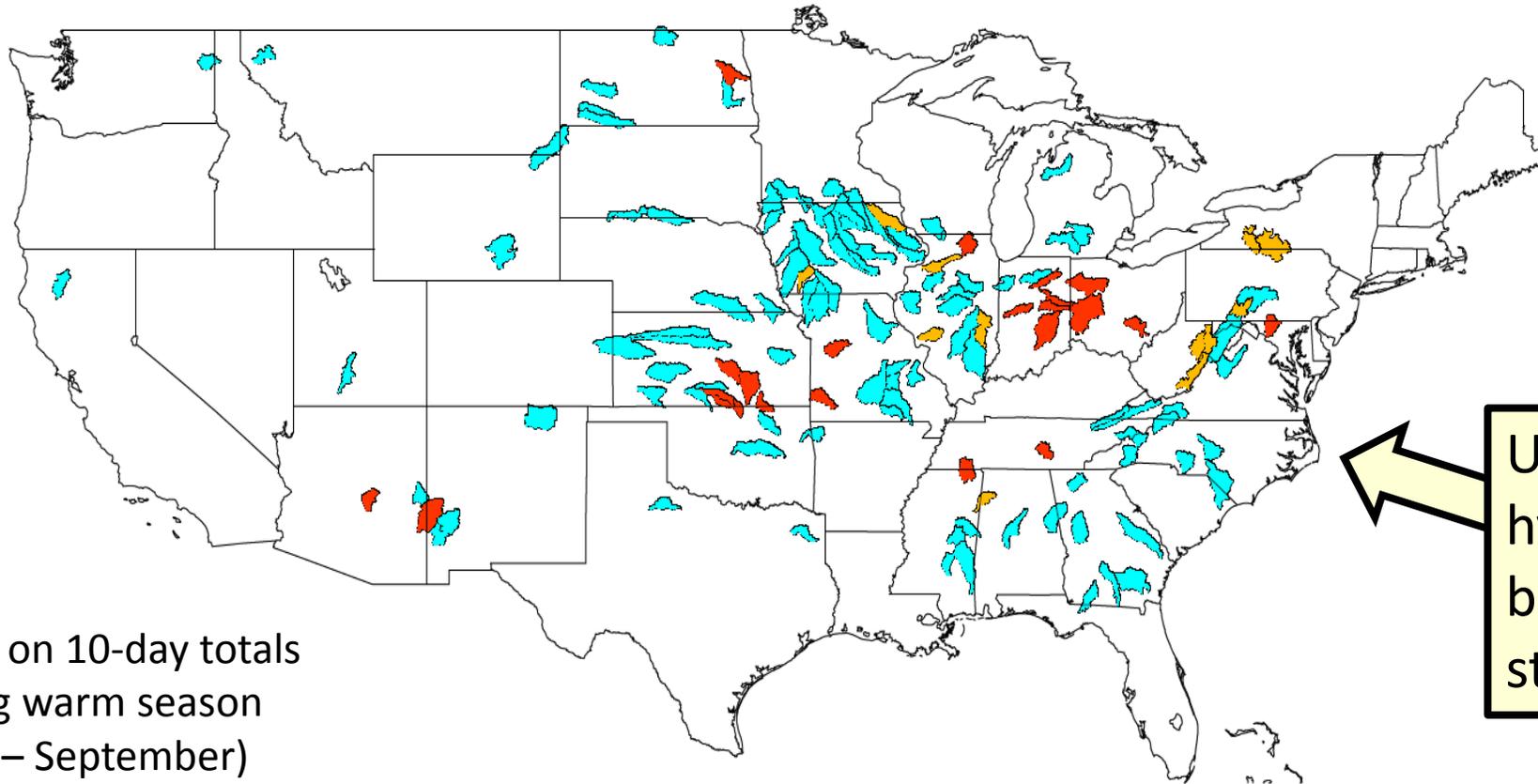
$$P = ET + Q + \Delta W$$

Today: a demonstration of how these terms can be estimated with SMAP soil moisture data

(not perfectly, but with some skill)

Basin level analysis (to allow for joint calculation of P and Q)

$$\text{P} = \text{ET} + \text{Q} + \Delta\text{W}$$



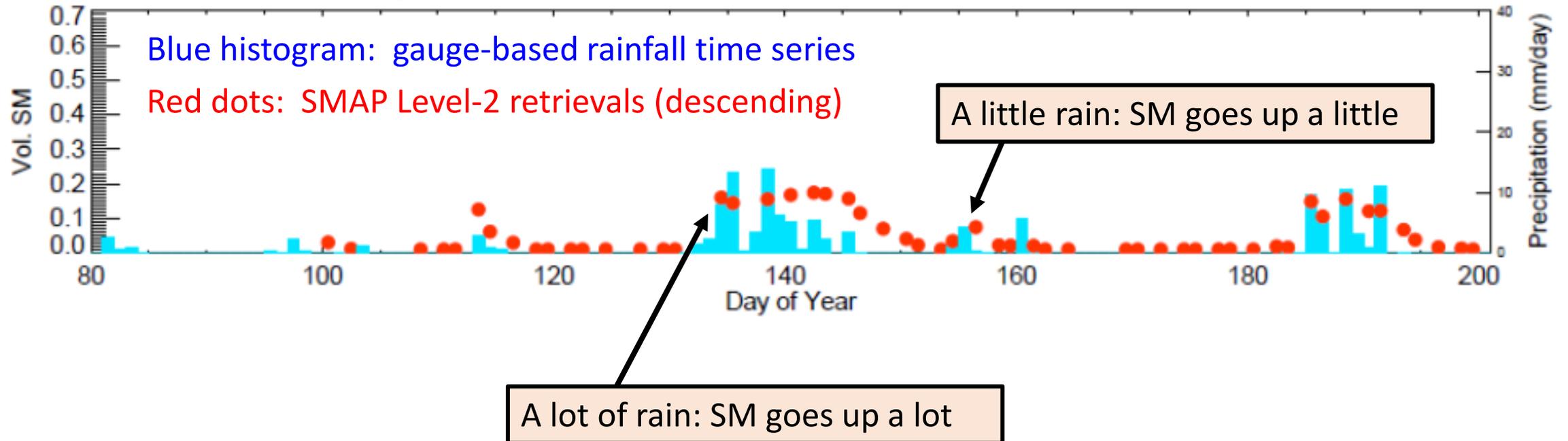
Unregulated hydrological basins with USGS stream gauges

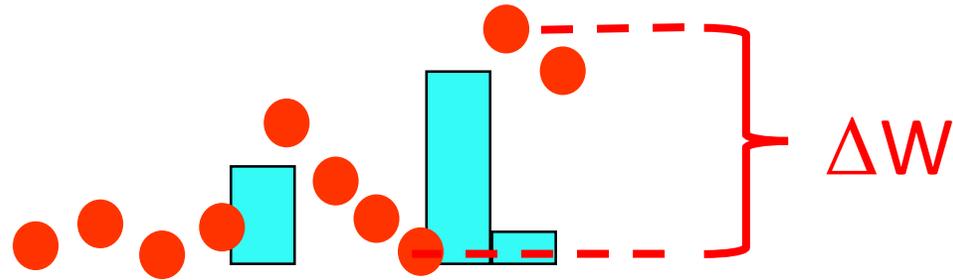
Focus on 10-day totals during warm season (June – September)

# Step 1: Precipitation Estimation

Notice: Rain gauge data and SMAP radiometer data generally look nicely consistent.

a. Location A: 119.3W, 41.80N (Western U.S.)

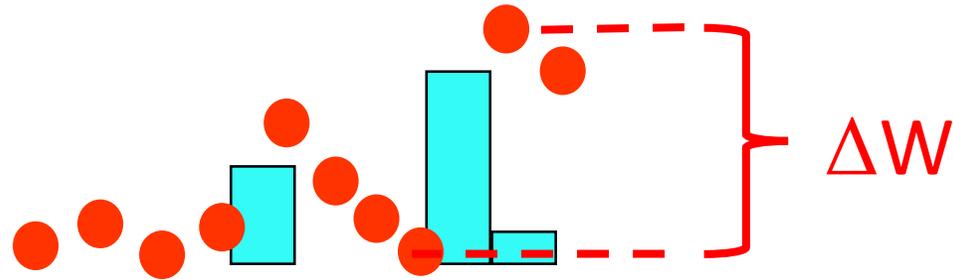




Precipitation estimation approach from Brocca et al. (2013):

$$P \sim \text{Max} ( \Delta W - a W_{\text{ave}}^b , 0. )$$

$W_{\text{ave}}$  = average of the two consecutive retrievals



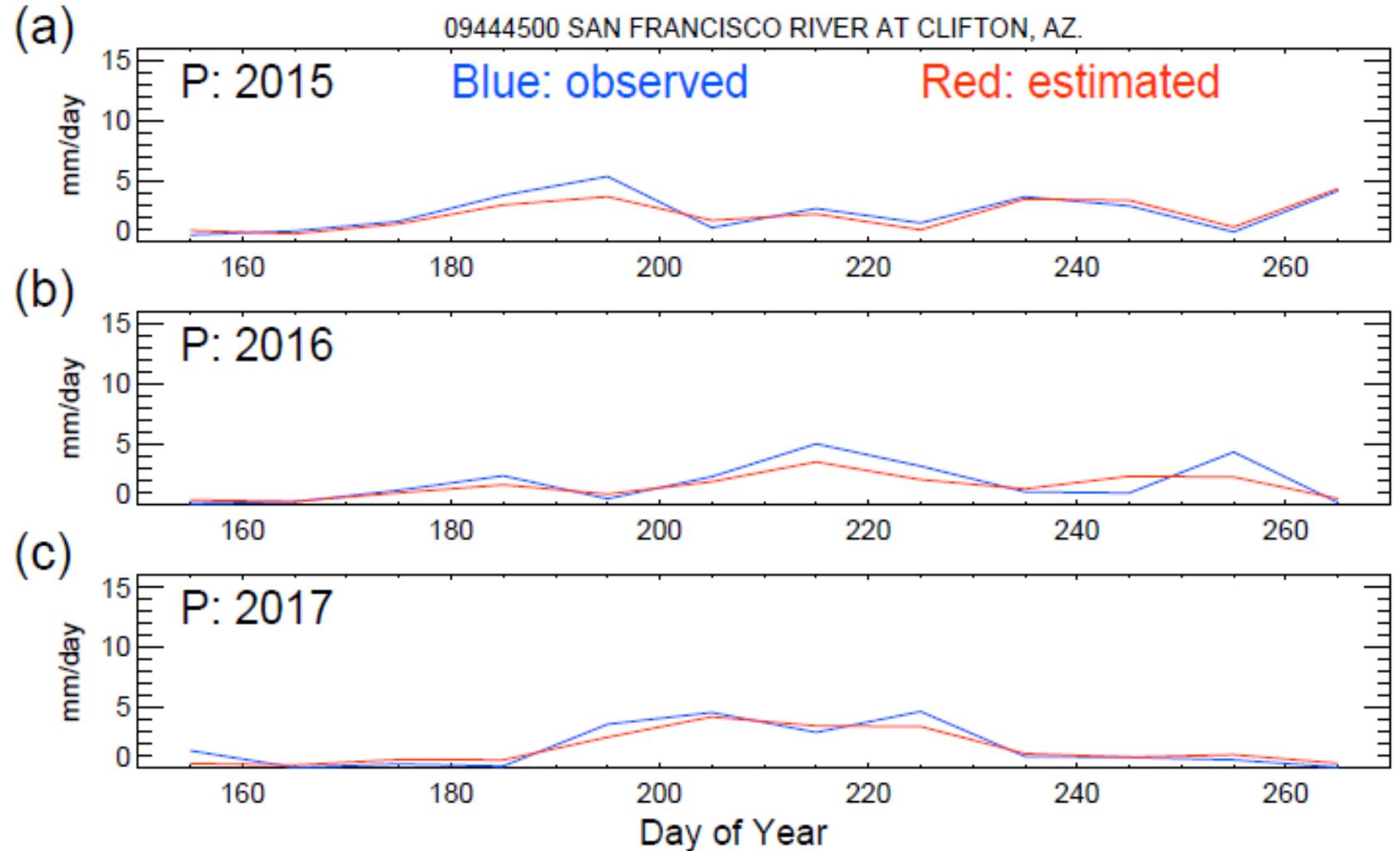
Updated approach:  
This term replaced with  
calibrated loss function

Precipitation estimation approach from Brocca et al. (2013):

$$P \sim \text{Max} ( \Delta W - a W_{\text{ave}}^b, 0. )$$

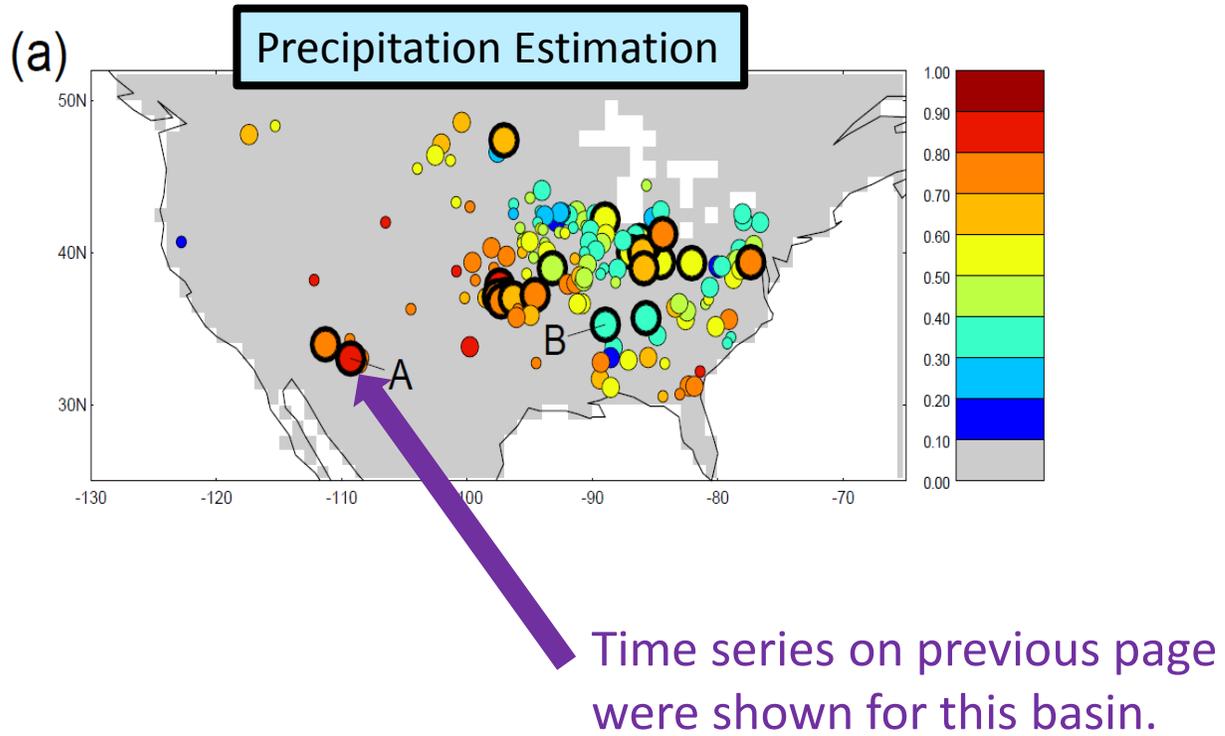
$W_{\text{ave}}$  = average of the two consecutive retrievals

## Some results! (One of the better estimations):

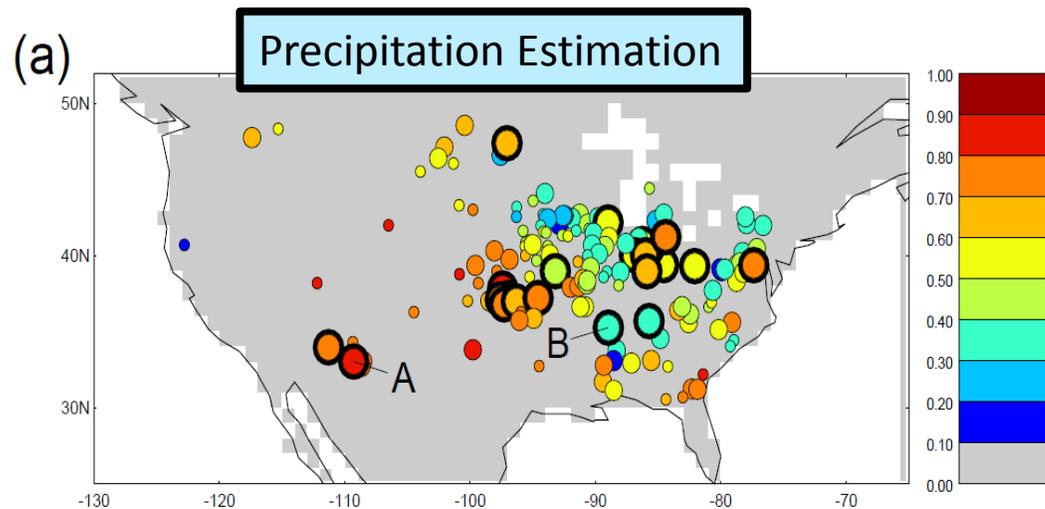


*We can characterize the agreement in these time series with the square of the correlation coefficient,  $r^2$ .*

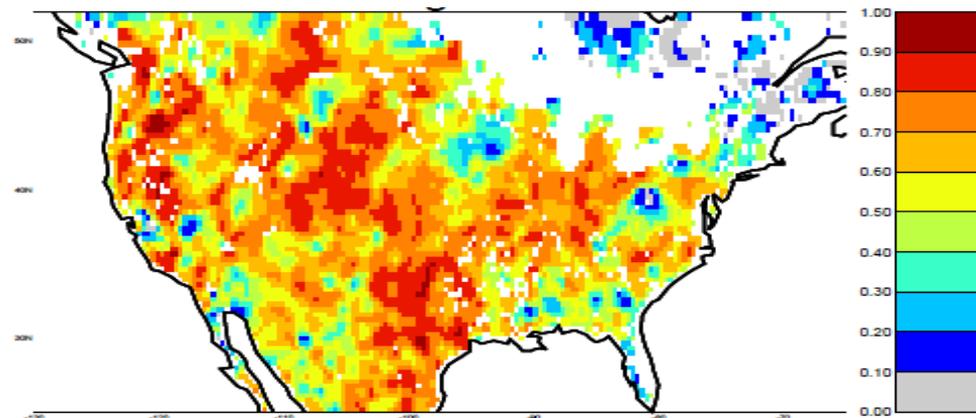
# Basin level skill scores (time series of 10-day precipitation totals: $r^2$ vs observations)



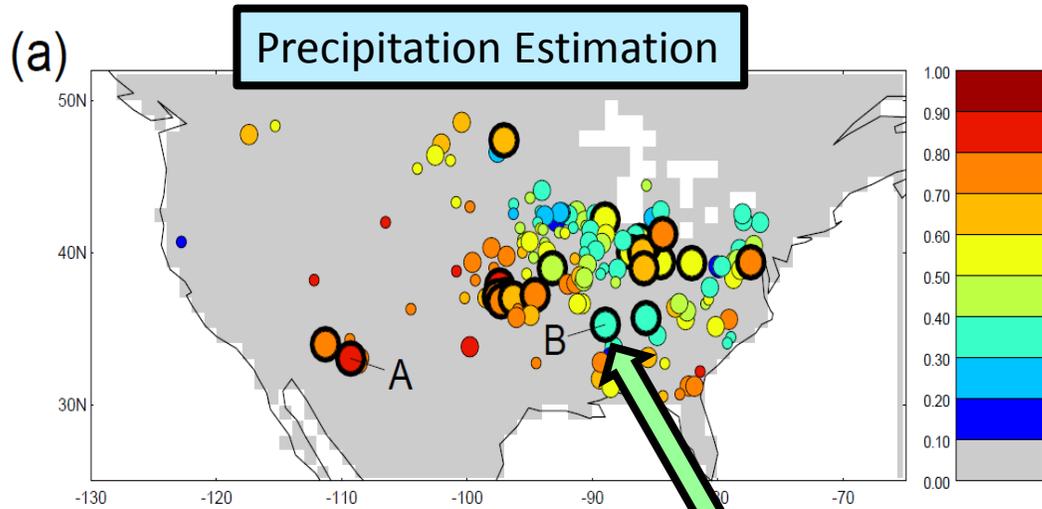
# Basin level skill scores (time series of 10-day precipitation totals: $r^2$ vs observations)



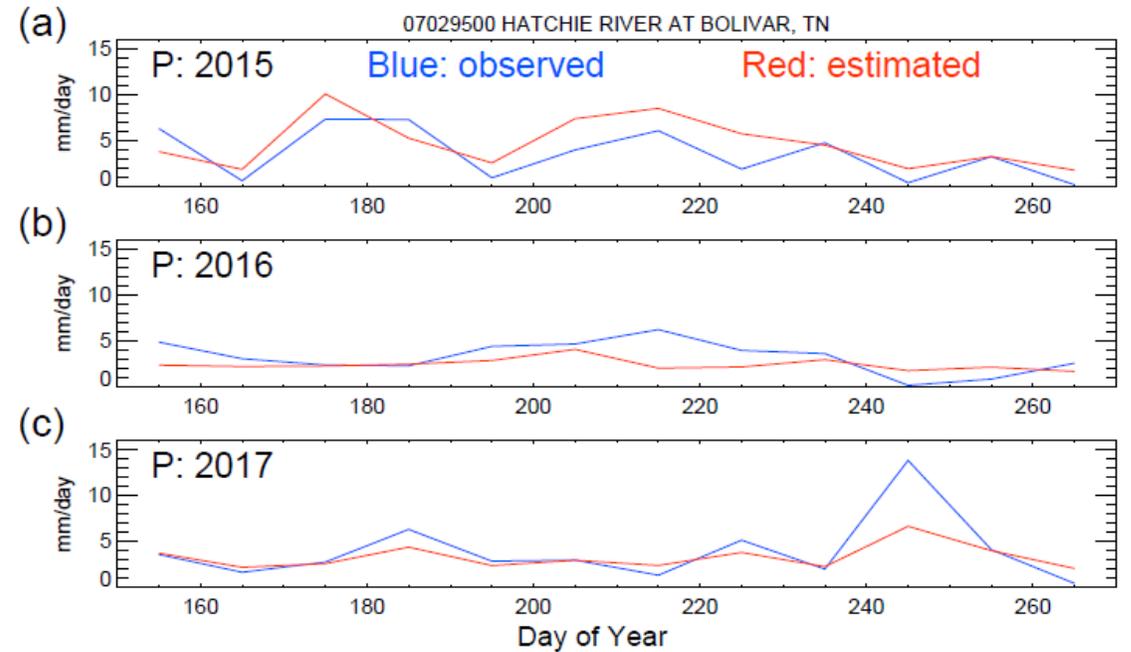
Aside: extending this analysis across the US, beyond “basins”, indicates high skill throughout the west.



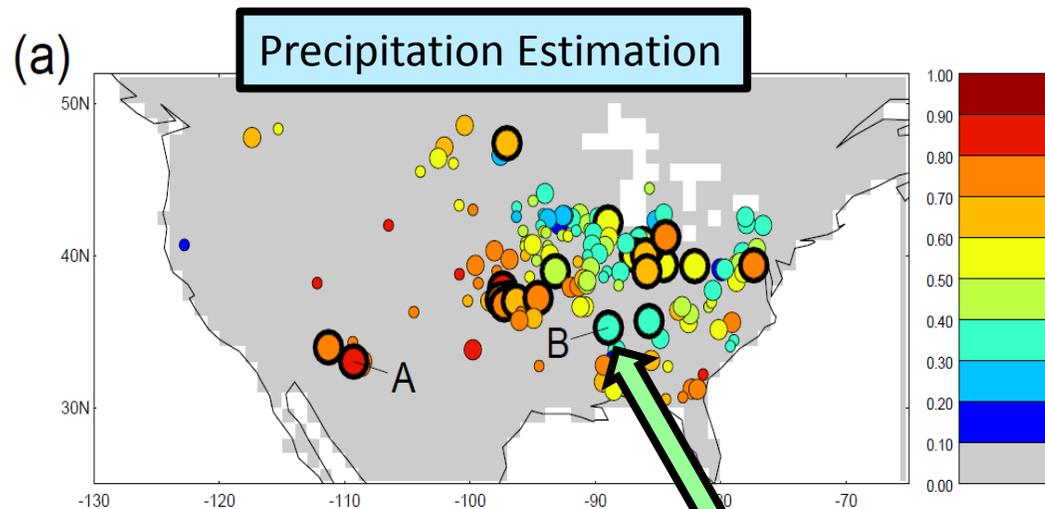
# Basin level skill scores (time series of 10-day precipitation totals: $r^2$ vs observations)



The algorithm works relatively poorly in this basin. Still, there is some valid information in the estimates

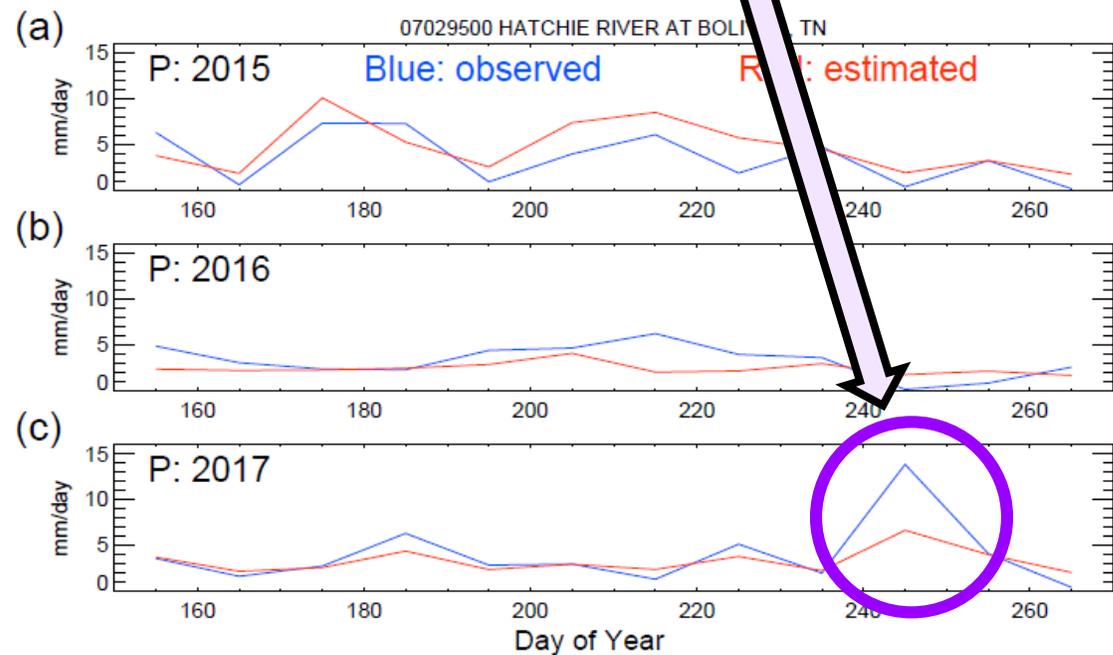


# Basin level skill scores (time series of 10-day precipitation totals: $r^2$ vs observations)



The algorithm works relatively poorly in this basin. Still, there is some valid information in the estimates

*Pardon the pun:*  
High rainfall saturates the soil moisture signal, making rainfall estimation difficult.



## Step 2: Streamflow Estimation

### Consider:

- The fraction of rainfall,  $P$ , that is converted to surface runoff,  $Q_{\text{fast}}$ , increases with surface soil moisture,  $W_{\text{surf}}$  :

$$Q_{\text{fast}} / P = f_1 (W_{\text{surf}})$$

- Drainage of moisture,  $Q_{\text{slow}}$ , to the water table (and eventually into streams) increases with increased soil moisture,  $W$ :

$$Q_{\text{slow}} = f_2 (W)$$

Apply multiple regression:

$$Q = Q_{\text{fast}} + Q_{\text{slow}} = a P W + b W + c$$

Diagram annotations:  
- A red arrow points from the text "basin-scale P" to the variable  $P$  in the equation.  
- Two green arrows point from the text "basin-scale W" to the two occurrences of the variable  $W$  in the equation.

Note – in practice, more complex and accurate approaches would be used. This simple approach has the advantage, though, of demonstrating unequivocally that SMAP data hold relevant information.

# Cross-validate!

Calibrate (i.e., find a, b, and c) using two years of observations:

obs Q from 2 SMAP years

obs P from 2 SMAP years

SMAP W from 2 SMAP years

$$Q = a P W + b W + c$$

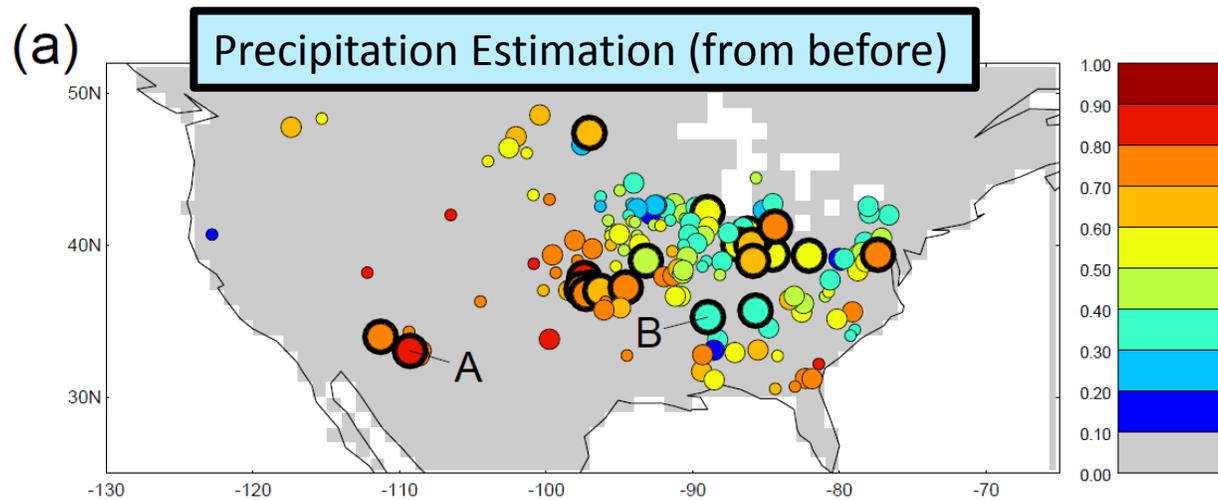
Validate against data from third year:

SMAP-based P estimate for 3<sup>rd</sup> year

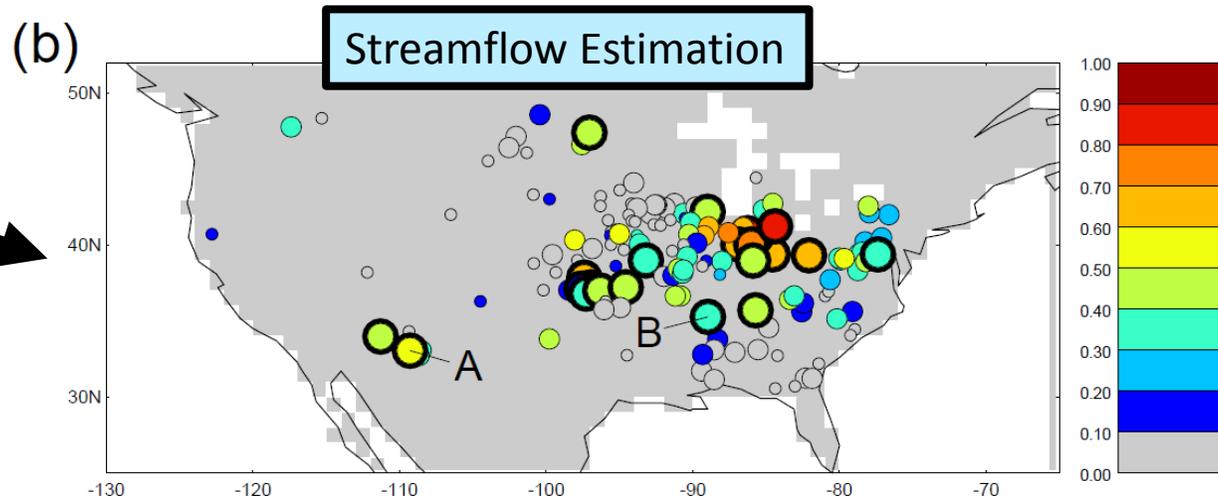
SMAP W in 3<sup>rd</sup> year

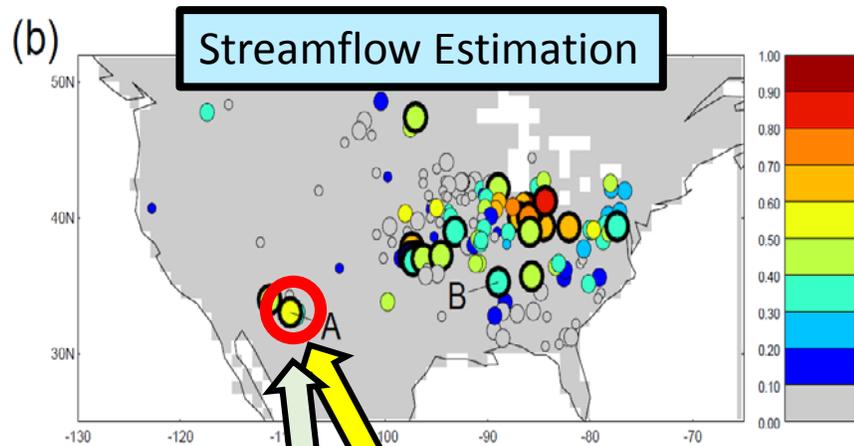
$$Q = a P W + b W + c$$

estimate of Q in 3<sup>rd</sup> year derived solely from SMAP data



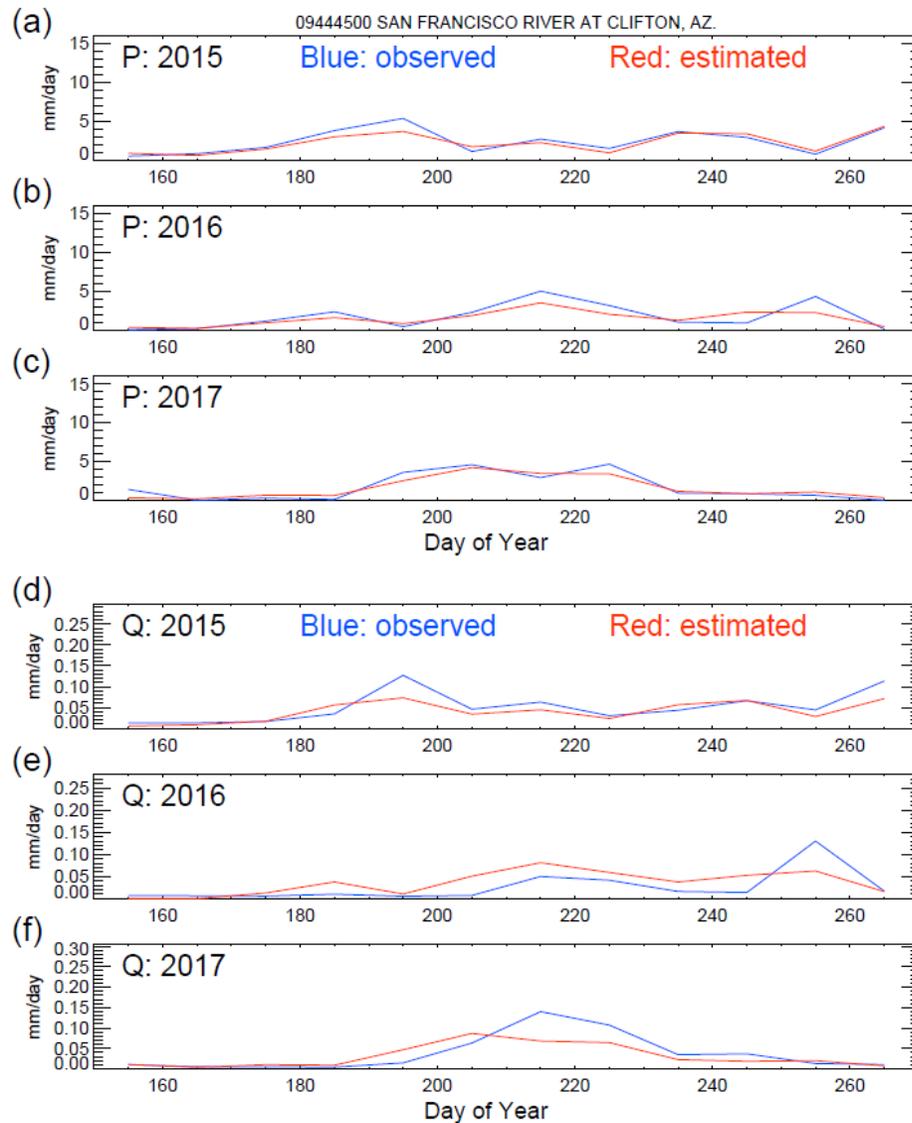
**Estimation skill  
(time series of 10-day  
streamflow totals:  
 $r^2$  vs observations)**

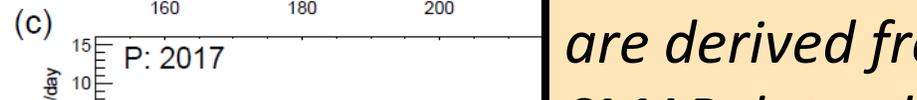
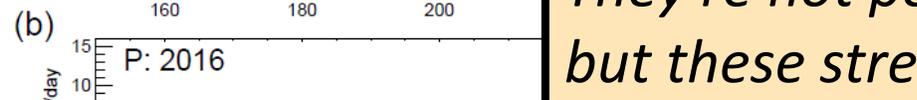
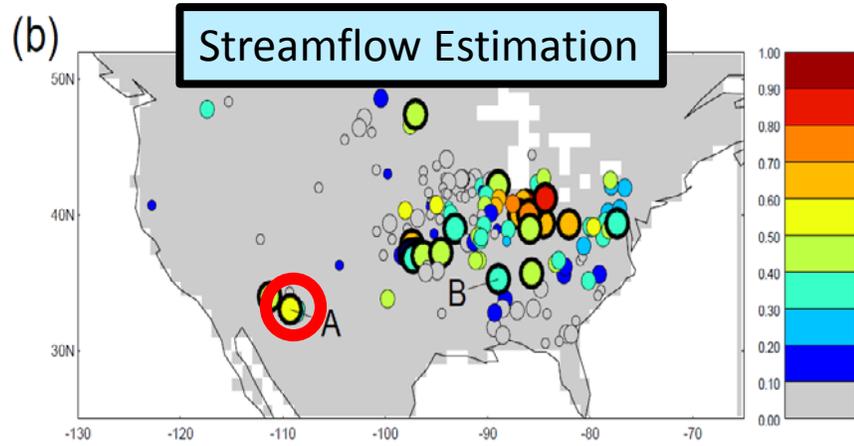




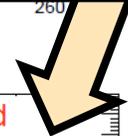
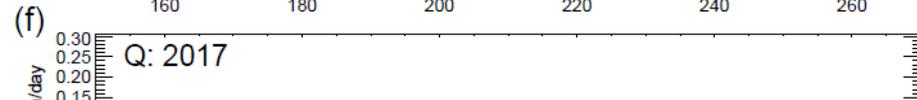
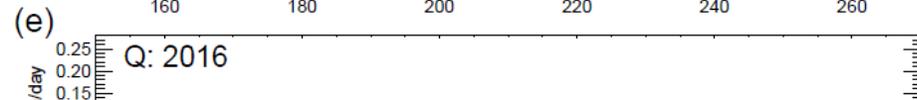
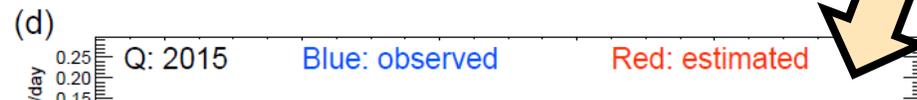
Precipitation estimates (from before)

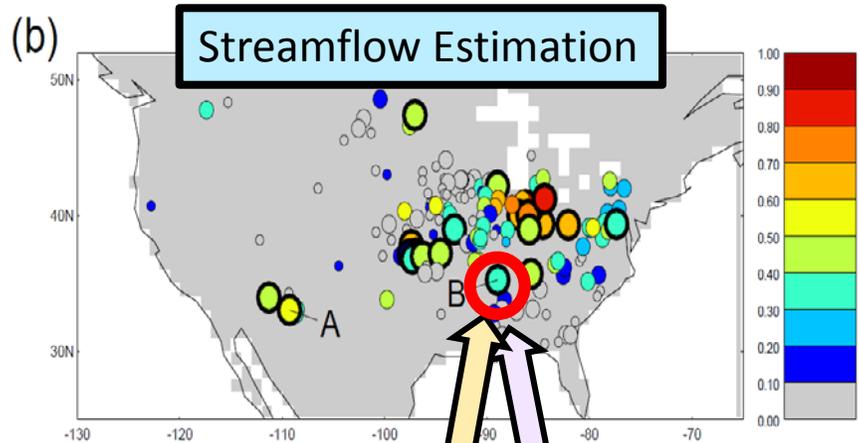
Streamflow estimates capture some of the observed behavior.





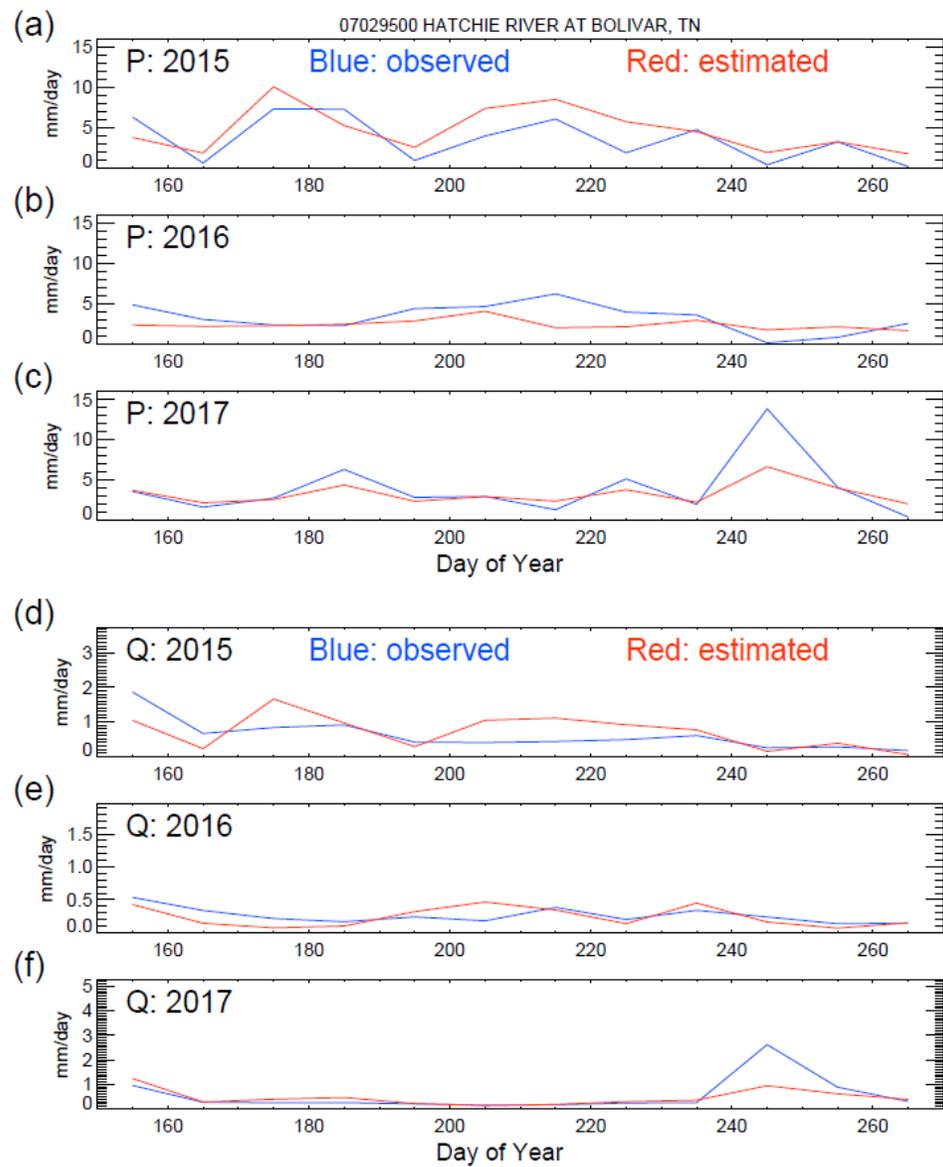
*They're not perfect, but these streamflow estimates (red lines) are derived from SMAP data alone!*

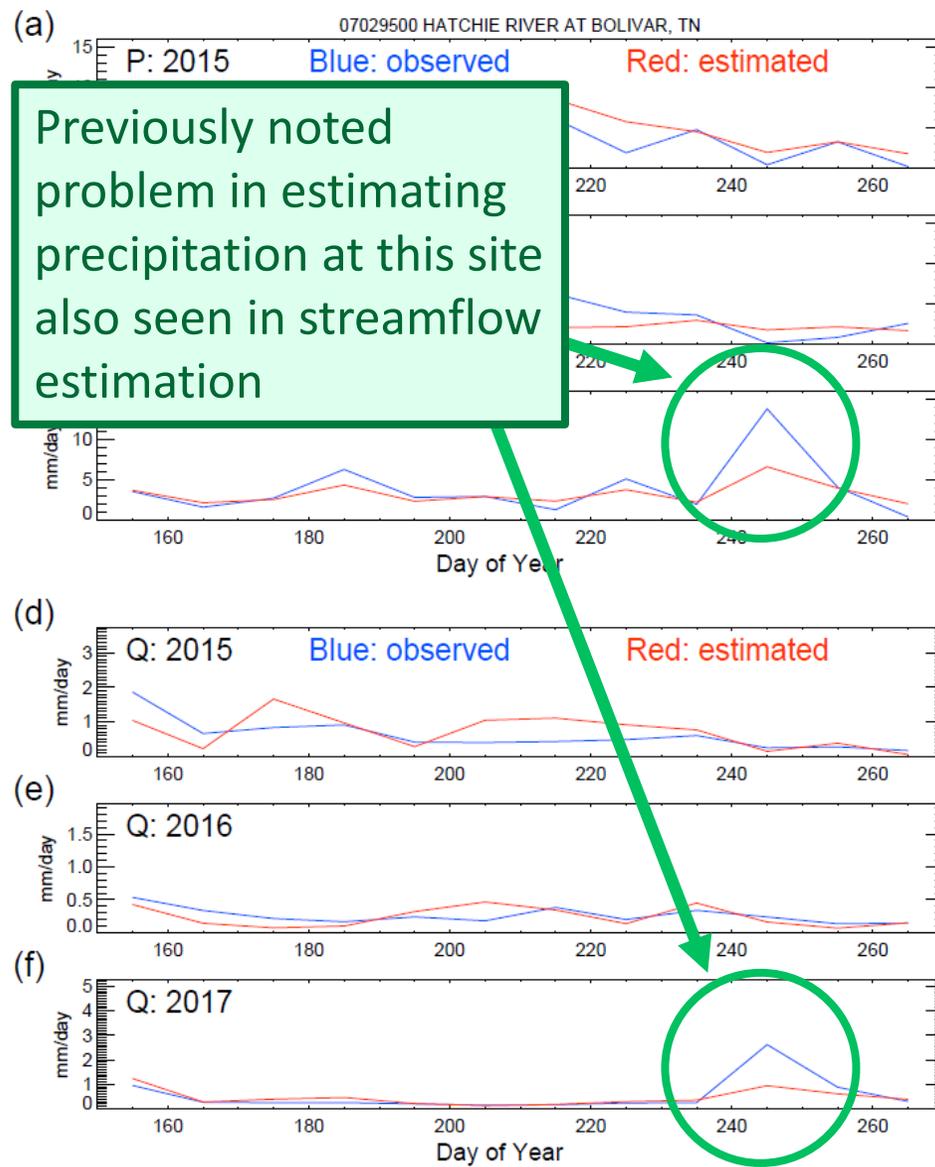
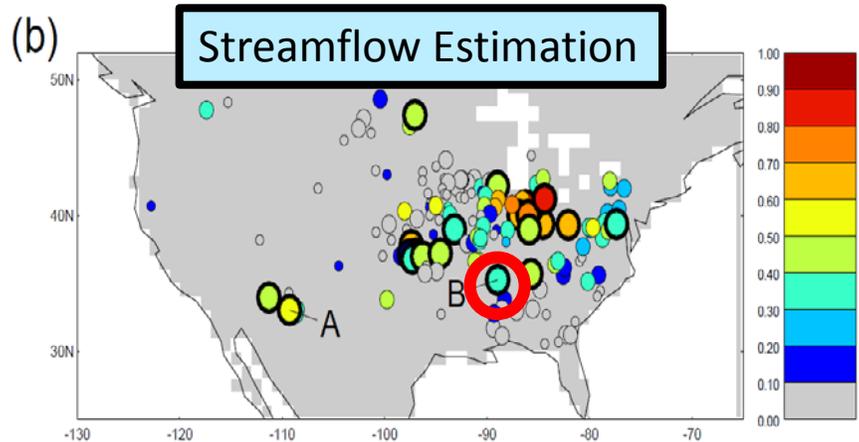




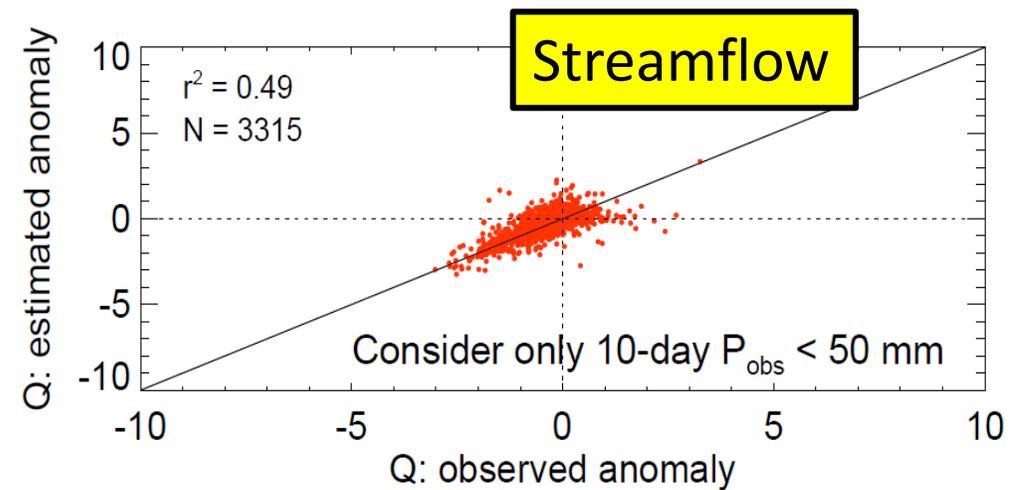
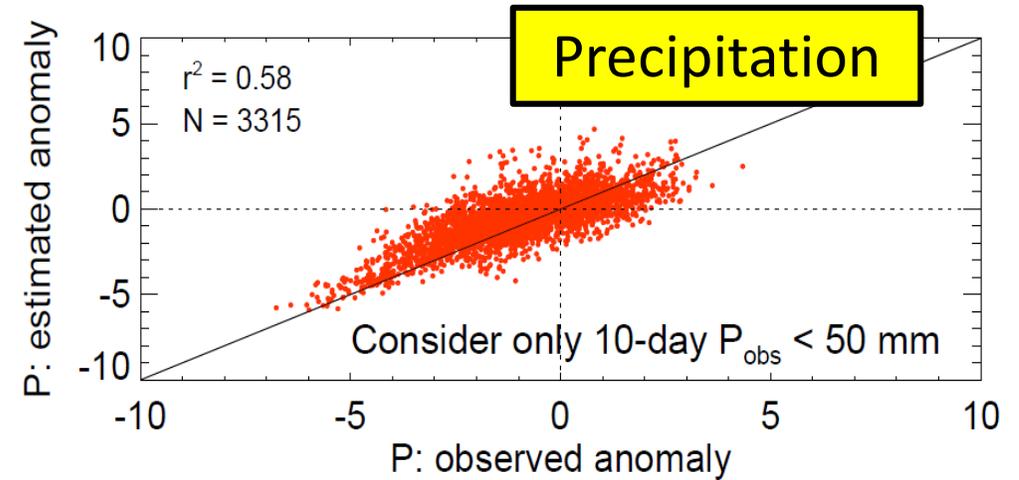
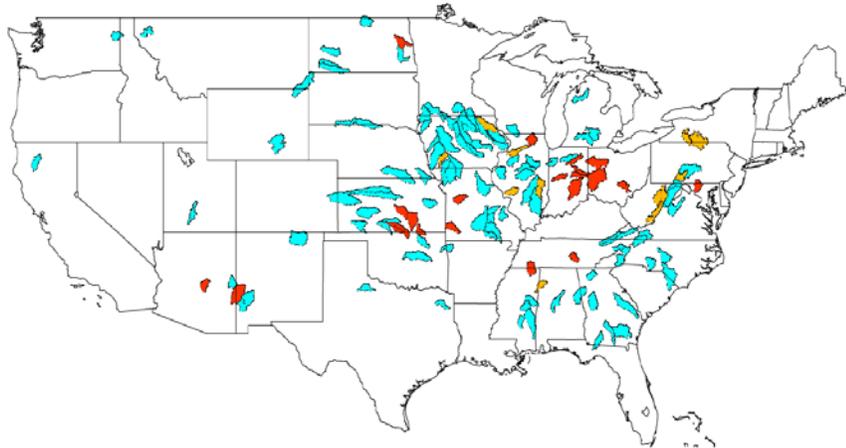
Precipitation estimates in "poorer basin" (from before)

Streamflow estimates capture less of the observed behavior.





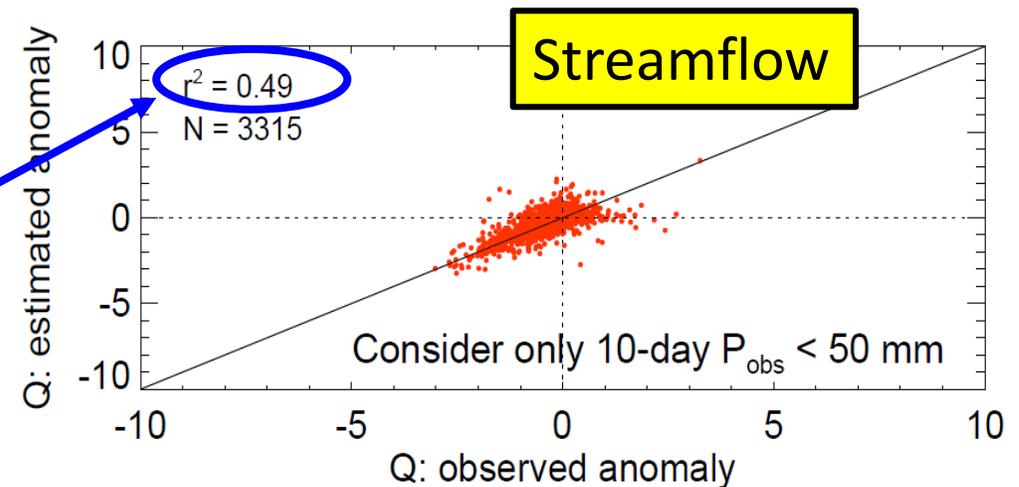
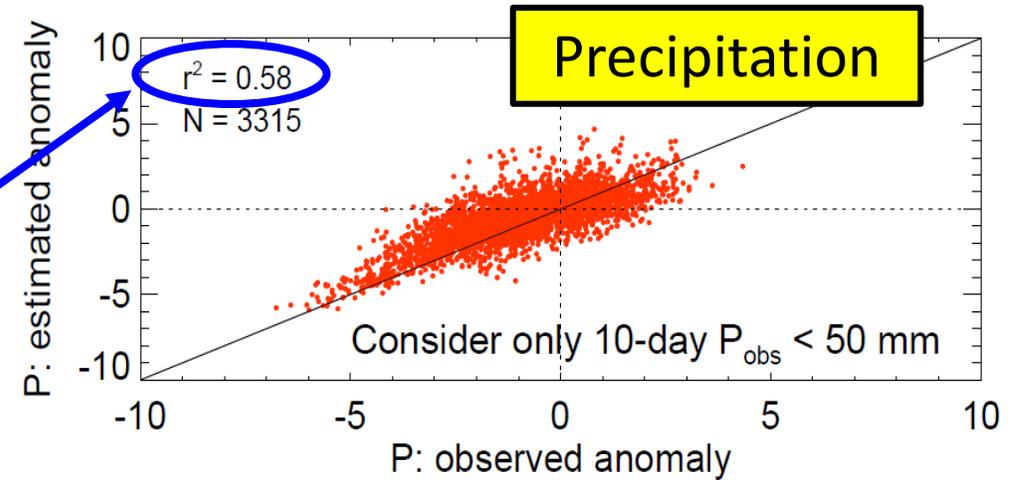
# Inter-basin analysis (throw results for all considered basins onto the same scatterplot)

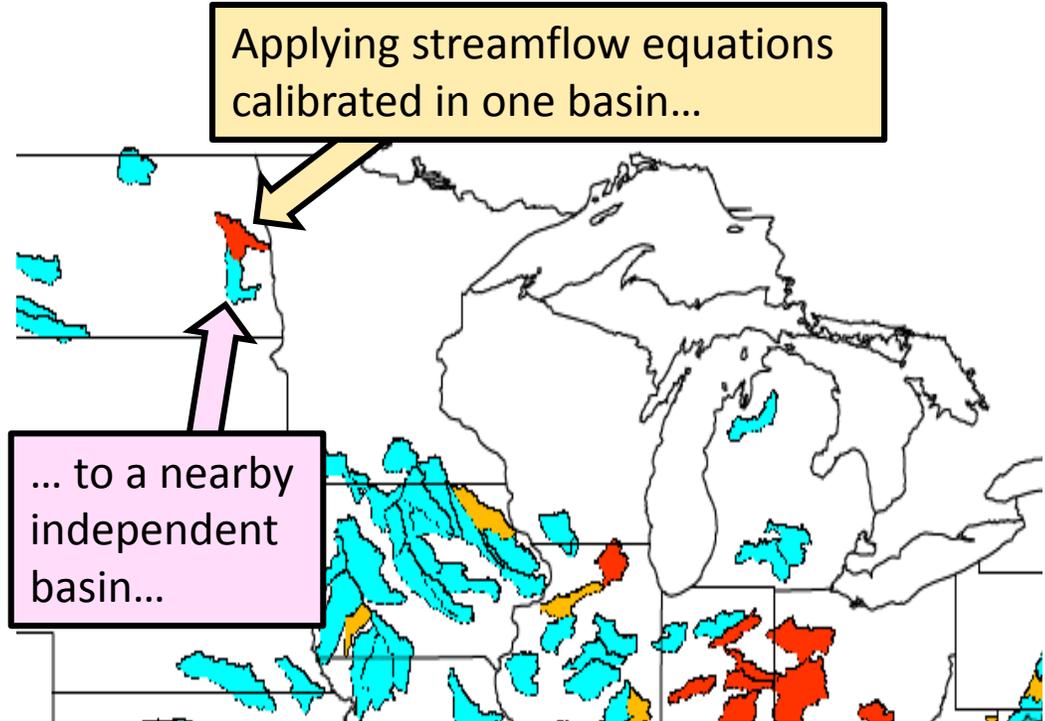


# Inter-basin analysis (throw results for all considered basins onto the same scatterplot)

For 10-day periods in which rainfall does not exceed 50 mm), SMAP data “explain” 58% of precipitation variance...

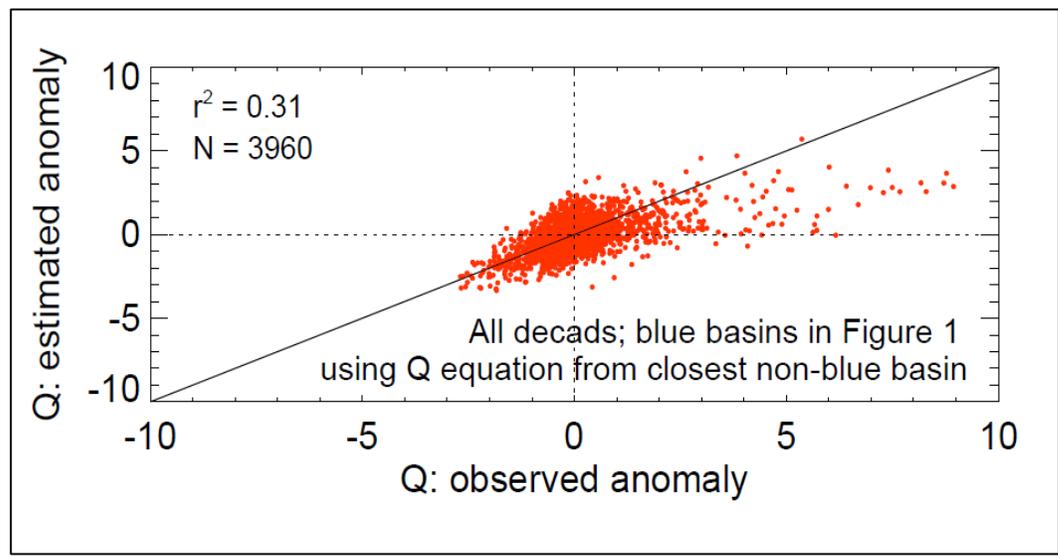
... and 49% of the streamflow variance





# Transferability (streamflow)

... produces streamflow estimates with some skill  
⇒ *there's hope for estimating streamflow in basins that never had a streamflow gauge.*



## Main Finding:

The SMAP estimates of rainfall and streamflow are not perfect, but they do contain relevant information.

*⇒ At the very least, they should prove useful for constraining, or otherwise contributing to, rainfall and streamflow estimates obtained with more conventional approaches.*

## A final comment...

Obvious question: What is the potential for examining other basin water budget components?

$$P - E - Q - \Delta\text{storage} = 0$$

We know that evapotranspiration is a strong function of soil moisture  
 ⇒ SMAP data could, in theory, be used to estimate it.

SMAP actually measures directly some of the storage change.

⇒ *The potential for estimating the other components as well is indeed there.*



Extra Slides

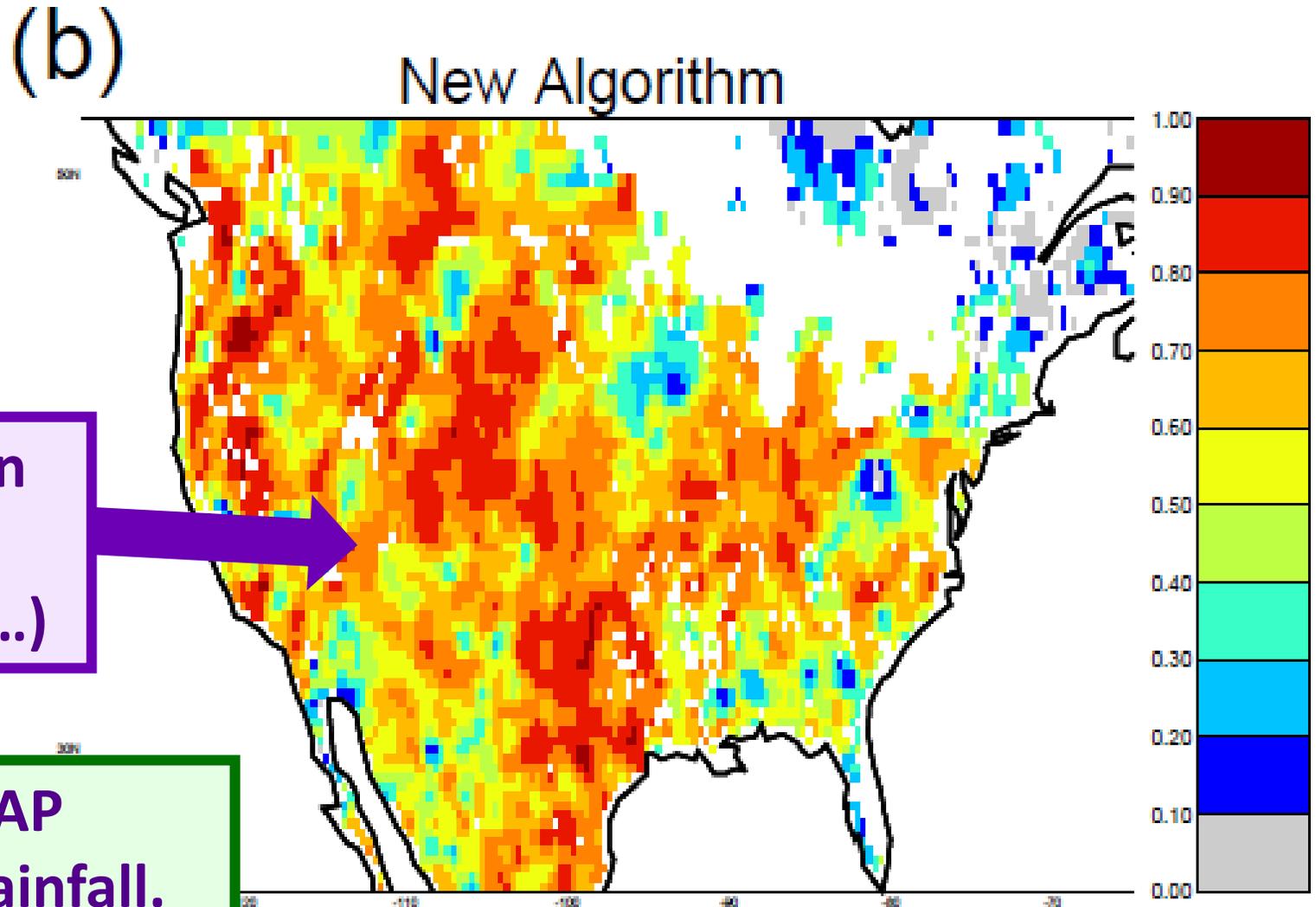
Skill ( $r^2$ , vs rain gauge observations) of ~100 km, 5-day precipitation estimates

Calibrate: using P observations in June - September of 2015-2016

Validate (map on right): June - September of 2017

**SMAP-based precipitation estimates are accurate!  
(At least to some degree...)**

**⇒ We can use SMAP data to estimate rainfall.**



Aggregate 10-day gridded precipitation estimates  
across hydrological basin of interest

⇒ hydrological basin precipitation estimates.

(Cross-validate:

- Use 2015,2016 data to calibrate model for 2017 estimates.
- Use 2015,2017 data to calibrate model for 2016 estimates.
- Use 2016,2017 data to calibrate model for 2015 estimates.)

## Basic idea:

We have already shown that precipitation rates,  $P$ , can be estimated with SMAP Level 2 data (Koster et al., 2016):

$$P = F_1(W_{SMAP})$$

Wade has also shown that SMAP data contain information on the ratio of runoff,  $Q$ , to precipitation (Crow et al., 2017):

$$Q/P = F_2(W_{SMAP})$$

Logically, then, runoff (streamflow) itself should be extractable from SMAP data:

$$Q = F_3(W_{SMAP})$$