

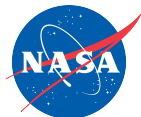
Assimilation of GMI and ATMS observations in the rainbands of hurricanes

Isaac Moradi^{1,2} Frank Evans³, William McCarty¹,
Frank Marks⁴

1. GMAO/GSFC/NASA, 2. ESSIC, University of Maryland,
3. U. of Colorado, 4. AOML/NOAA

33rd Conference on Hurricanes and Tropical Meteorology
16 – 20 April 2018 Ponte Vedra, FL.

April 17, 2018



**This work is supported by the NASA Data
for Operation and Assessment program,
grant# NNX17AE89G.
Many thanks to NDOA Program Manager,
Dr. Tsengdar Lee.**



Overview

- Due to limitations in directly assimilating microwave cloudy observations in the rain-bands of hurricanes, a new technique named Bayesian Monte Carlo Integration (BMCI) is introduced.
- The BMCI technique is used to retrieve T , q , wind speed and several other parameters from microwave radiances.
- These retrievals can be either directly used by forecasters to evaluate the structure of hurricanes or be assimilated into NWP models.

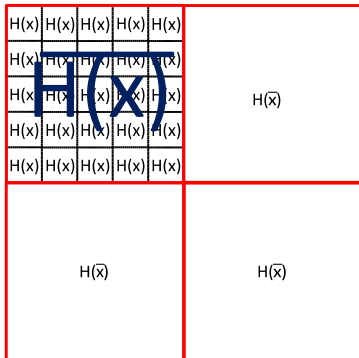


The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$
 \vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$

\vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

Non-linearity in the forward model: \vec{x} is the mean value of the model variables within grid-box and because H is a non-linear: $\overline{H(\vec{x})} \neq H(\vec{x})$.



The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$

\vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

Non-linearity in the forward model: \vec{x} is the mean value of the model variables within grid-box snf because H is a non-linear: $\overline{H(\vec{x})} \neq H(\overline{\vec{x}})$.

Inaccuracy in the first-guess: the models do not provide a close first guess for cloud parameters or clouds are often displaced.

The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$

\vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

Non-linearity in the forward model: \vec{x} is the mean value of the model variables within grid-box snf because H is a non-linear: $\overline{H(\vec{x})} \neq H(\overline{\vec{x}})$.

Inaccuracy in the first-guess: the models do not provide a close first guess for cloud parameters or clouds are often displaced.

Lack of required RT inputs: \vec{p}_s are neither provided by the model nor fully measurable in real world thus are estimated from limited in-situ and aircraft measurements.

The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$

\vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

Non-linearity in the forward model: \vec{x} is the mean value of the model variables within grid-box snf because H is a non-linear: $\overline{H(\vec{x})} \neq H(\overline{\vec{x}})$.

Inaccuracy in the first-guess: the models do not provide a close first guess for cloud parameters or clouds are often displaced.

Lack of required RT inputs: \vec{p}_s are neither provided by the model nor fully measurable in real world thus are estimated from limited in-situ and aircraft measurements.

Simplified RT models: Operational RT models that use a simplified RT framework, such as spherical hydrometeors, which is not appropriate at higher microwave frequencies where ice scattering is important.

The relation between the observations (O) and the forward operator (H) can be expressed as: $O = H(\vec{x}, \vec{p}_b, \vec{p}_s) + \epsilon$

\vec{x} state vector, \vec{p}_b parameters such as shape and size distribution of hydrometers, \vec{p}_s indicates the scattering parameters (e.g., phase function)

Non-linearity in the forward model: \vec{x} is the mean value of the model variables within grid-box snf because H is a non-linear: $\overline{H(\vec{x})} \neq H(\overline{\vec{x}})$.

Inaccuracy in the first-guess: the models do not provide a close first guess for cloud parameters or clouds are often displaced.

Lack of required RT inputs: \vec{p}_s are neither provided by the model nor fully measurable in real world thus are estimated from limited in-situ and aircraft measurements.

Simplified RT models: Operational RT models that use a simplified RT framework, such as spherical hydrometeors, which is not appropriate at higher microwave frequencies where ice scattering is important.

Assuming Gaussian Errors: DA systems assume Gaussian error statistics, examined using the departures, but in the case of cloudy radiances the departures are likely to be non-Gaussian

The BMCI technique

The BMCI technique can be summarized in three steps:

- generation of a retrieval database of atmospheric state and cloud variables using a-priori information. The database should also include extreme cases as the extrapolation is not allowed.

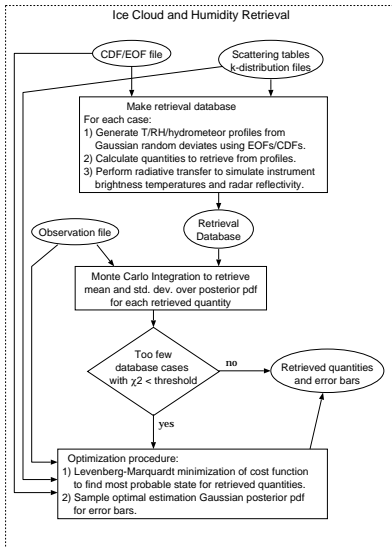
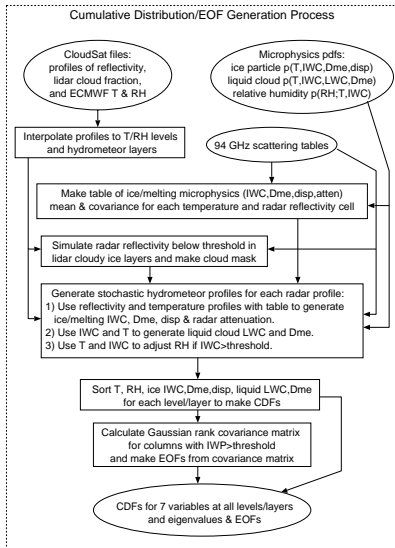
The BMCI technique

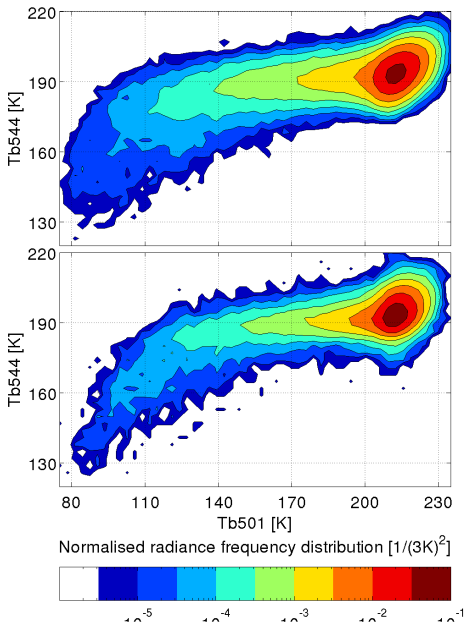
The BMCI technique can be summarized in three steps:

- generation of a retrieval database of atmospheric state and cloud variables using a-priori information. The database should also include extreme cases as the extrapolation is not allowed.
- the atmospheric state and cloud variables are fed into the RT model to generate the synthetic observations. In addition to the state variables such as temperature, water vapor, and cloud profiles, cloud microphysics and parameterization such as particles' shape and size distribution are also utilized as input.

The BMCI technique can be summarized in three steps:

- generation of a retrieval database of atmospheric state and cloud variables using a-priori information. The database should also include extreme cases as the extrapolation is not allowed.
- the atmospheric state and cloud variables are fed into the RT model to generate the synthetic observations. In addition to the state variables such as temperature, water vapor, and cloud profiles, cloud microphysics and parameterization such as particles' shape and size distribution are also utilized as input.
- real measurements along with the generated database are given to the retrieval package, then the retrieval package will select the cases which are close to the real measurements and integrate them according to the Bayes' theorem to give the estimate of the mean and uncertainty of the state and cloud variables.





Rydberg et al., 2009

$$p_{post}(\vec{x}|\vec{y}) = \frac{p_f(\vec{y}|\vec{x})p_p(\vec{x})}{\int p_f(\vec{y}|\vec{x}')p_p(\vec{x}')d\vec{x}'} \Rightarrow \text{Posterior} = \frac{\text{Likelihood} \times \text{Prior}}{\text{Marginal Likelihood}}$$

The retrieved values for atmospheric or cloud state (\hat{x}) can be computed by integrating over the posterior pdf:

$$\hat{x} = \int \vec{x} p_{post}(\vec{x}|\vec{y}) d\vec{x}$$

$$\hat{x} = \frac{\sum_i x_i p_f(\vec{y}|\vec{x}_i)}{\sum_i p_f(\vec{y}|\vec{x}_i)}$$

The variance (error) of the posterior pdf is calculated as:

$$\sigma_x^2 = \int (x - \hat{x})^2 p_{post}(\vec{x}|\vec{y}) d\vec{x}$$

The conditional pdf can be defined using the probability density of the measured vectors for any given atmospheric state (j channel number):

$$P(\vec{y}|\vec{x}) = \prod_{j=1}^m \frac{1}{\sqrt{2\pi\sigma_j^2}} \exp\left(-\frac{[\vec{y}_j - H_j(\vec{x})]^2}{2\sigma_j^2}\right)$$

σ is the noise in the measurements and \hat{x} now can be calculated as:

$$\hat{x} = \frac{\sum_i \vec{x}_i \exp(-\frac{1}{2}\chi^2)}{\sum_i \exp(-\frac{1}{2}\chi^2)}$$

$$\chi^2 = \sum_{j=1}^M \frac{[\vec{y}_j - H_j(\vec{x})]^2}{\sigma_j^2}$$

Improvements to the BMCI Retrievals

We have made a significant progress on enhancing the BMCI retrieval system and adding new functionalities to the code. Some of the major enhancements to the code are as follows:

- Adding temperature profile retrieval capability as well as the ocean skin temperature and near surface wind speed

Improvements to the BMCI Retrievals

We have made a significant progress on enhancing the BMCI retrieval system and adding new functionalities to the code. Some of the major enhancements to the code are as follows:

- Adding temperature profile retrieval capability as well as the ocean skin temperature and near surface wind speed
- Computing ice particle scattering properties at new frequencies and generating new scattering tables

Improvements to the BMCI Retrievals

We have made a significant progress on enhancing the BMCI retrieval system and adding new functionalities to the code. Some of the major enhancements to the code are as follows:

- Adding temperature profile retrieval capability as well as the ocean skin temperature and near surface wind speed
- Computing ice particle scattering properties at new frequencies and generating new scattering tables
- Implementing the FASTEM microwave ocean surface emissivity model, both forward and adjoint, in the BMCI code

We have made a significant progress on enhancing the BMCI retrieval system and adding new functionalities to the code. Some of the major enhancements to the code are as follows:

- Adding temperature profile retrieval capability as well as the ocean skin temperature and near surface wind speed
- Computing ice particle scattering properties at new frequencies and generating new scattering tables
- Implementing the FASTEM microwave ocean surface emissivity model, both forward and adjoint, in the BMCI code
- Modifying the original CloudSat reflectivity profile based CDF/EOF program to also use GPM Dual-frequency Precipitation Radar (DPR) reflectivity profiles

Improvements to the BMCI Retrievals

We have made a significant progress on enhancing the BMCI retrieval system and adding new functionalities to the code. Some of the major enhancements to the code are as follows:

- Adding temperature profile retrieval capability as well as the ocean skin temperature and near surface wind speed
- Computing ice particle scattering properties at new frequencies and generating new scattering tables
- Implementing the FASTEM microwave ocean surface emissivity model, both forward and adjoint, in the BMCI code
- Modifying the original CloudSat reflectivity profile based CDF/EOF program to also use GPM Dual-frequency Precipitation Radar (DPR) reflectivity profiles



- Analyzing in situ warm cloud and rain microphysical data from the Hurricane Research Division (HRD) and generating stochastic profiles of warm liquid cloud profiles



- Analyzing in situ warm cloud and rain microphysical data from the Hurricane Research Division (HRD) and generating stochastic profiles of warm liquid cloud profiles
- Adding ERA-Interim profiles of stratospheric temperature and water vapor matched to CloudSat times and locations to complement the CloudSat ECMWF-AUX profiles (which only reach 24 km).

Improvements to the BMCI Retrievals

- Analyzing in situ warm cloud and rain microphysical data from the Hurricane Research Division (HRD) and generating stochastic profiles of warm liquid cloud profiles
- Adding ERA-Interim profiles of stratospheric temperature and water vapor matched to CloudSat times and locations to complement the CloudSat ECMWF-AUX profiles (which only reach 24 km).
- Modifying the CDF-EOF algorithm to allow for clear layers using a hydrometeor masking procedure for ice, rain, and liquid cloud



- Analyzing in situ warm cloud and rain microphysical data from the Hurricane Research Division (HRD) and generating stochastic profiles of warm liquid cloud profiles
- Adding ERA-Interim profiles of stratospheric temperature and water vapor matched to CloudSat times and locations to complement the CloudSat ECMWF-AUX profiles (which only reach 24 km).
- Modifying the CDF-EOF algorithm to allow for clear layers using a hydrometeor masking procedure for ice, rain, and liquid cloud
- Modifying the 1D Bayesian retrieval program to input the new CDF-EOF a priori file and generate consistent profiles of temperature, relative humidity, and ice particle, raindrop, and cloud droplet size distribution parameters to use in the Bayesian profile retrievals.

Hurricane Sandy CloudSat overpassed it on October 27, 2012

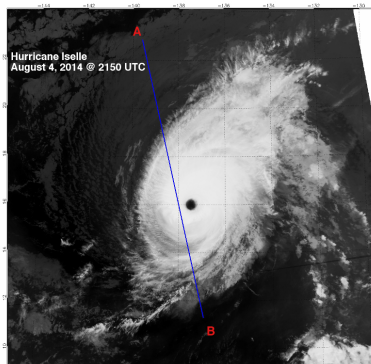
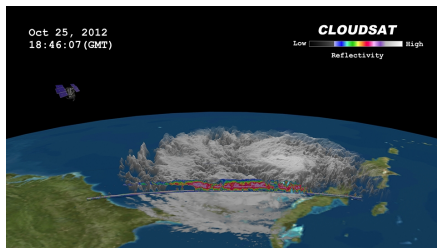
Hurricane Iselle CloudSat overpassed the hurricane on August 4, 2014

Hurricane Amanda CloudSat overpassed Hurricane Amanda on May 25, 2014

Hurricane Joaquin In an early stage of the formation of Hurricane Joaquin, on September 29, 2015, CloudSat passed over the center of the hurricane in the Caribbean.

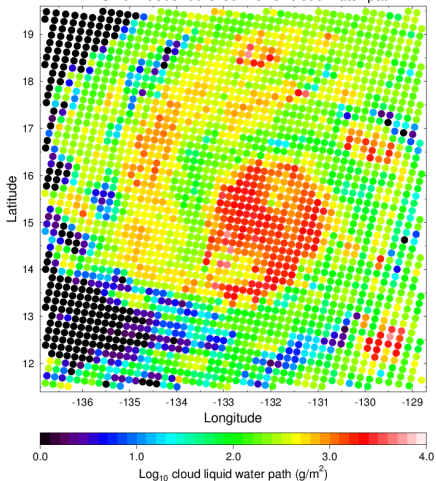
Image credit:

cloudsat.atmos.colostate.edu

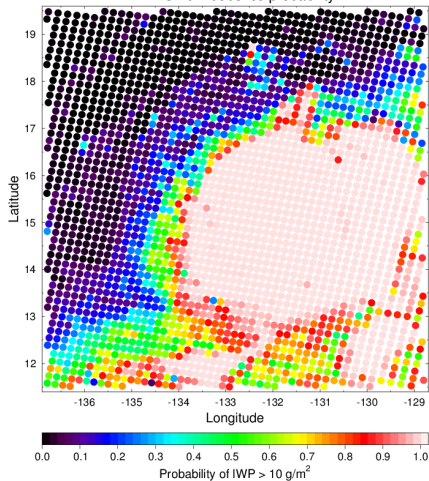


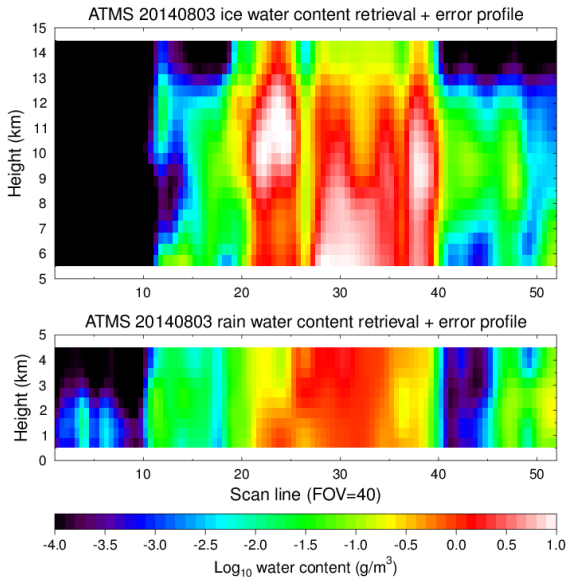


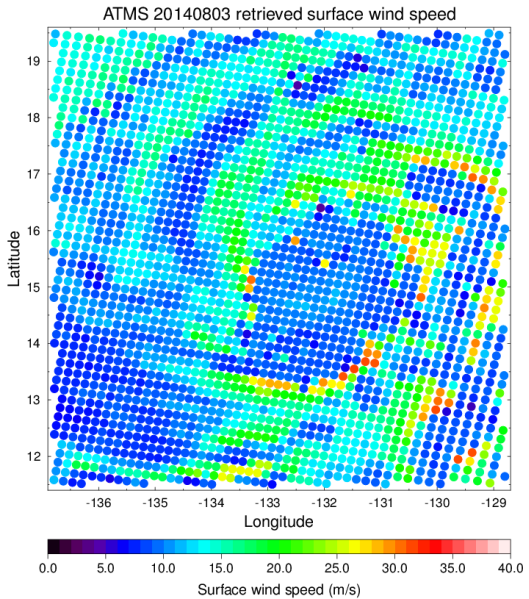
ATMS 20140803 retrieved + error cloud water path



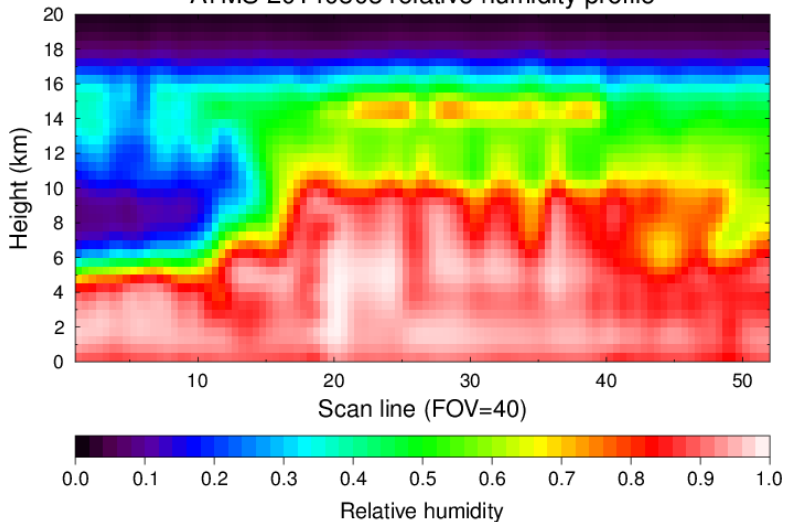
ATMS 20140803 ice probability







ATMS 20140803 relative humidity profile



Top: SkinTemp (left), IWP (right), Bottom: Rain WP (left), Surface Wind Speed (right)

- Conventional data assimilation schemes cannot properly assimilate satellite radiances in the rainband of tropical cyclones due to inaccuracy in RT scattering parameters as well as inaccuracy in the first guess provided by NWP models
- A new technique is proposed that does not depend on the minimization of the cost function.
- Preliminary results from BMCI technique are encouraging but require extensive validation, though validation itself is challenging
- These retrieved profiles are valuable for both analyzing the structure of the hurricanes as well as to provide more accurate initial conditions for the NWP models

Thank you for
your attention!

**This work is supported by the NASA Data
for Operation and Assessment program,
grant# NNX17AE89G.
Many thanks to NDOA Program Manager,
Dr. Tsengdar Lee.**