

# IMA2018 Abstract submission

*Martian mineralogy: observations, experiments, analogues and models*

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## Tracing the evolution of hydrogen in the martian crust through laboratory studies of apatite

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**What is your preferred presentation method?:** Oral presentation

**: Introduction:** Northwest Africa (NWA) 7034 and its pairings represent a regolith breccia of basaltic bulk composition [1], the fine-grained matrix of which bears a strong resemblance to the major and trace element composition estimated for the ancient southern highlands crust on Mars [2]. Therefore, NWA 7034 may represent a key sample for constraining the composition of the martian crust, particularly the ancient highlands. Here we seek to constrain the hydrogen isotopic composition of the martian crust using apatite [Ca<sub>5</sub>(PO<sub>4</sub>)<sub>3</sub>(Cl,F,OH)]. Apatites across all lithologic domains in NWA 7034 have been affected by a Pb-loss event at ~1.5 Ga before present [3] and so they are unlikely to have retained magmatic volatile composition and are more likely to have equilibrated with fluids within the martian crust that may or may not have exchanged with the martian atmosphere.

**Methods & Results:** Electron probe microanalysis (EPMA) of apatite was conducted following the work of McCubbin *et al.*, [4] and Santos *et al.*, [5]. The H-isotopic composition and H<sub>2</sub>O abundance of apatite in two thin sections of NWA 7034 were measured using the Cameca NanoSIMS 50L at The Open University, following protocols described in [6]. Apatites in NWA 7034 were measured in a number of lithologic contexts, including basaltic clasts, Fe, Ti, and P-rich (FTP) clasts, and large crystal clasts within the bulk matrix [5]. All apatites are Cl-rich and contain between 238-1343 ppm H<sub>2</sub>O. In addition, the  $\delta D$  values of apatites ranged from 453‰ to 2564‰.

**Discussion:** In the absence of a magnetosphere, the martian atmosphere has been stripped of <sup>1</sup>H over billions of years by the solar wind [7]. This has produced a heavily fractionated atmosphere ( $\delta D \sim 2500$ - $6100$ ‰,[8,9]) compared to the mantle ( $\delta D < 275$ ‰,[10]). The equilibration of apatite with crustal fluids is supported by the chlorine-rich compositions exhibited by apatites in NWA 7034 in comparison to apatites from other martian meteorites [4]. The apatites in NWA 7034, which likely attained their volatile inventories during a Pb-loss event at 1.5 Ga [3], have  $\delta D$  values  $< 3000$ ‰. All of the H-isotopic compositions that we obtained fall within ~500‰ of the range of values reported for the intermediate H reservoir predicted for the martian crust by Usui *et al.*, [11]. These observations indicate that at least portions of the martian regolith have not exchanged completely with the martian atmosphere. If NWA 7034 records the H-isotopic composition of the crust ca. 1.5 Ga, then Allan Hills 84001 (ALH 84001), the oldest unbrecciated martian meteorite, constrains the H-isotopic composition of the crust ca. 4 Ga [12]. Greenwood *et al.* [13] determined a  $\delta D$  value of ~3000 ‰ for apatite in ALH 84001, which is D-enriched relative to the intermediate reservoir. However, our recent petrologic investigation of ALH 84001 hints that there are at least two generations of apatite within ALH 84001, and it is not clear which generation of apatite was measured by [13]. We will measure the H-isotopic composition of apatite in ALH 84001 to ascertain whether or not ALH 84001 holds new clues to the evolution of hydrogen in the ancient martian crust.

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