

# Overview of Large-Scale Tropospheric Transport in the Chemistry Climate Model Initiative (CCMI) Simulations

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# Motivation

- Differences in large-scale tropospheric transport among models contribute to differences in aerosol distributions in the Arctic and to the interhemispheric gradients of GHGs and ODSs (e.g. *Shindell et al. (2008)*, *Patra et al. (2011)*, *Monks et al. (2015)*).
- It is not clear, however, whether these uncertainties are driven by large-scale flow biases and/or subgrid-scale processes.
- Few studies have examined how tropospheric transport (e.g. transport to the Arctic, interhemispheric exchange) will change in a warmer climate (e.g. *Holzer and Boer (2001)*, *Doherty et al. (2017)*).

# Motivation

The Chemistry Climate Modeling (CCM) Initiative experiments (*Eyring et al. (2013)*) provide a unique opportunity to examine the relationship between tropospheric transport and large-scale dynamics because:

- Unprecedented number of tropospheric transport diagnostics, including a range of both idealized loss and age tracers (*Waugh et al. (2013), Eyring et al. (2013), Orbe et al. (2016,2017)*)
- Large number of models submitting both “specified-dynamics” and free-running simulations **using the same underlying model code**
- Much more dynamical output, relative to previous composition intercomparisons (e.g. TRANSCOM, ACCMIP).

# Motivation

Here we use the Chemistry Climate Modeling (CCM) Initiative experiments (*Eyring et al. (2013)*), consisting of hindcast simulations over the recent past, performed both in “specified-dynamics” (REF-C1SD) and free-running (REF-C1) modes to evaluate:

*#1 What is the spread in tropospheric transport among CCMs and how is that related to differences in large-scale dynamics and/or (parameterized) convection?*

*#2 Is tropospheric transport better constrained in specified-dynamics (SD)(versus free-running (FR)) simulations?*

# Motivation

Here we use the Chemistry Climate Modeling (CCM) Initiative experiments (*Eyring et al. (2013)*), consisting of hindcast simulations over the recent past, performed both in “specified-dynamics” (REF-C1SD) and free-running (REF-C1) modes, and future (REF-C2) simulations to examine more systematically:

*#1 What is the spread in tropospheric transport among CCMs and how is that related to differences in large-scale dynamics and/or (parameterized) convection?*

*#2 Is tropospheric transport better constrained in specified-dynamics (SD)(versus free-running (FR)) simulations?*

*#3 How is transport to the Arctic and interhemispheric transport projected to change by the end of the 21<sup>st</sup> century?*

# Methods

## A. Experiments:

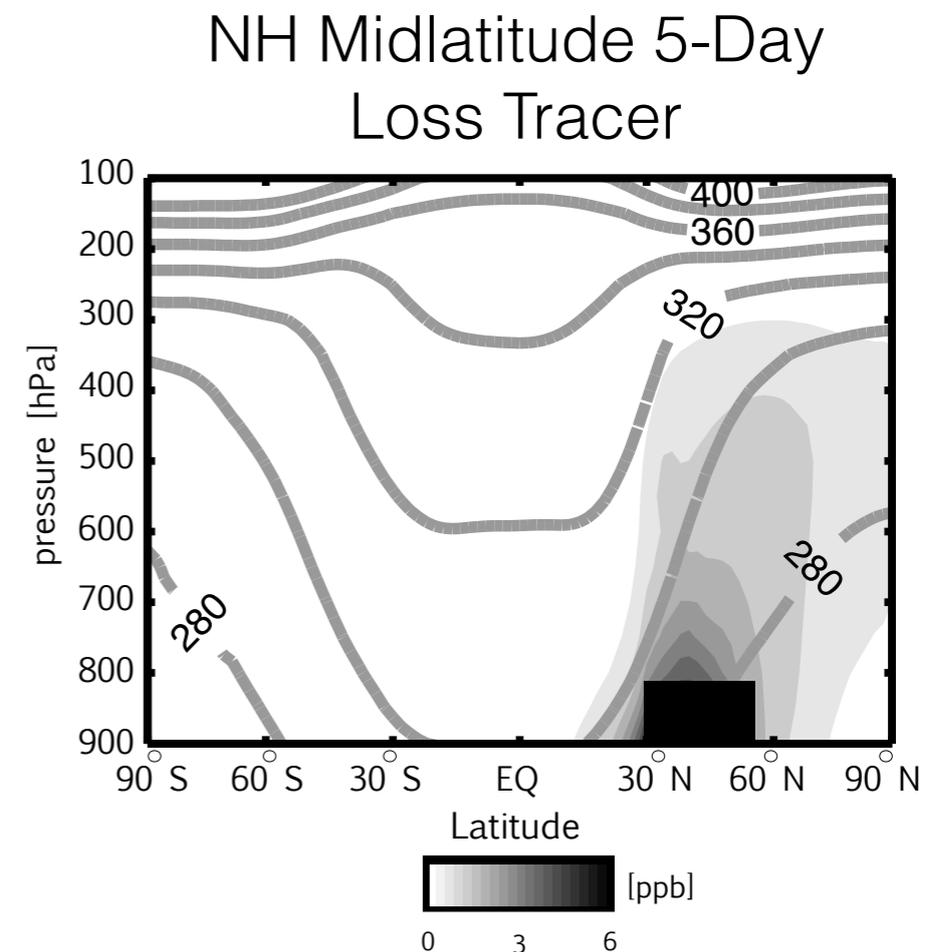
**REF-C1SD** (1980-2010): observed SSTs and SICs, analysis large-scale flow

**REF-C1** (1960-2010): observed SSTs and SICs, free-running

**REF-C2** (1960-2100): modeled SSTs and SICs, free-running, RCP 6.5 scenario

## B. Transport Diagnostics:

Tropospheric transport is inferred from idealized loss tracers with a NH midlatitude source ( $\chi_5$  and  $\chi_{50}$ ) as well as a NH midlatitude mean age tracer ( $\Gamma_{\text{NH}}$ ) (*Waugh et al. (2013)*, *Eyring et al. (2013)*, *Orbe et al. (2016,2017)*).



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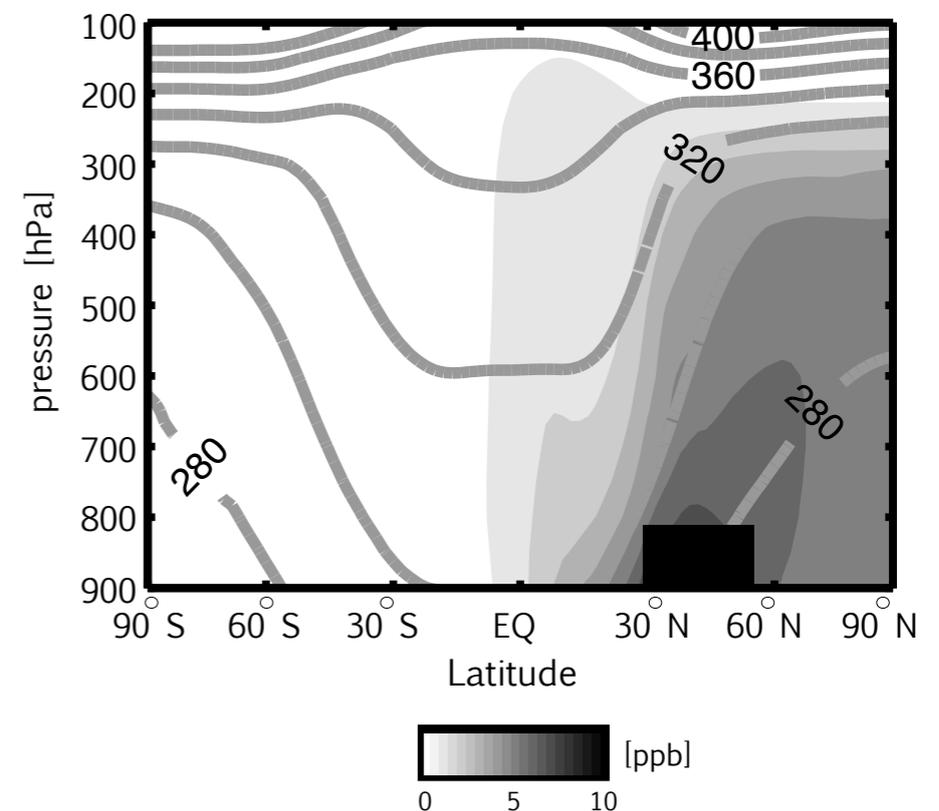
**REF-C1** (1960-2010): observed SSTs and SICs, free-running

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NH Midlatitude 50-Day  
Loss Tracer



# Methods

## A. Experiments:

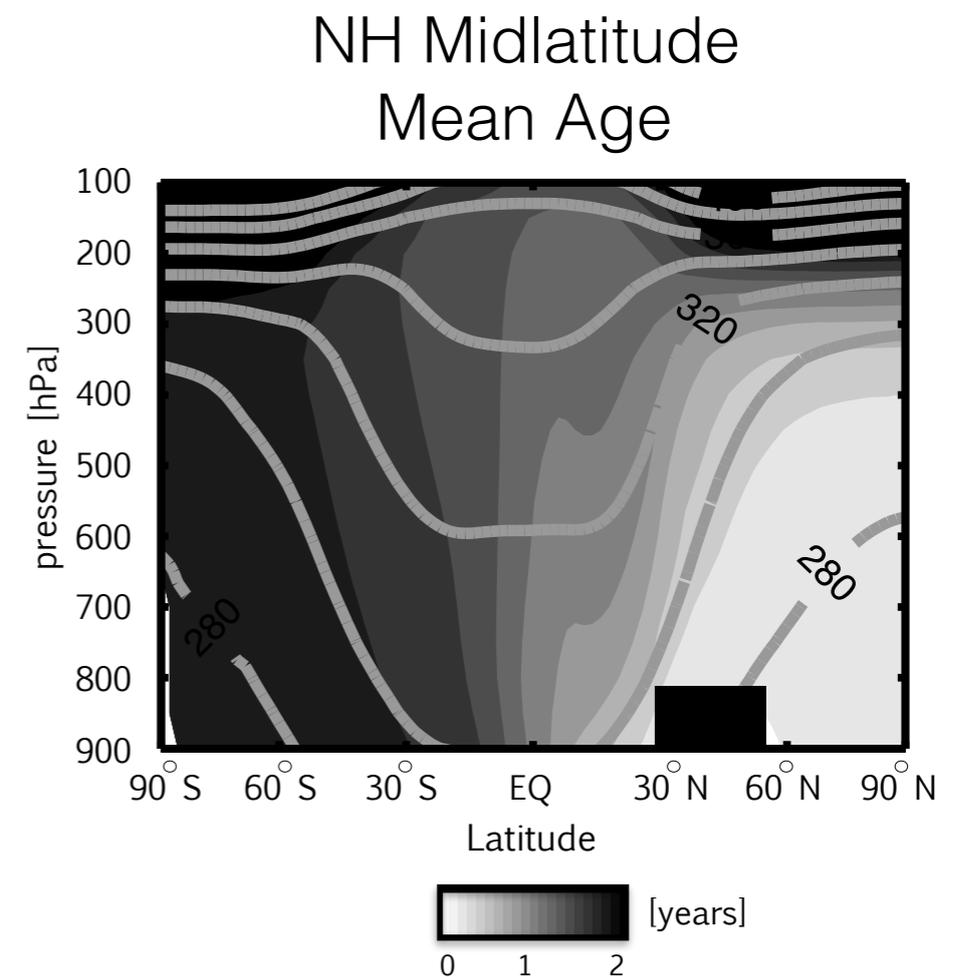
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**REF-C1** (1960-2010): observed SSTs and SICs, free-running

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## B. Transport Diagnostics:

In addition to examining tracers with zonally invariant sources ( $\chi_5$ ,  $\chi_{50}$ ,  $\Gamma_{\text{NH}}$ ) we will also examine more realistic tracers with only land (CO-like) emissions ( $\chi_{\text{CO}50}$ ) (*Shindell et al. (2008), Monks et al. (2015), Doherty et al. (2017), Yang et al. (2018, Under Review)*).

# Models: Hindcast Experiment

Simulation Name	Model (Reference)	Horizontal Resolution	Vertical Levels (Model Top)	Large-Scale Flow (Free/Nudging/CTM)	Convective Parameterization
GEOS-CTM	NASA Global Modeling Initiative Chemical Transport Model Strahan et al., (2013)	2° x 2.5°	72 (0.01 hPa)	MERRA (CTM)	Moorthi and Suarez (1992) Bacmeister et al. (2006)
GEOS-C1SD	Goddard Earth Observing System Version 5 GCM Reinecker et al. (2007); Molod et al. (2015)	" "	" "	MERRA (Nudging)	" "
GEOS-C1	" "	" "	" "	Free-running	" "
WACCM-C1SDV1/V2	Whole Atmosphere Community Climate Model Version 4 (WACCM-4) Marsh et al. (2013); Solomon et al. (2015); Garcia et al. (2016)	1.9° x 2.5°	88 (140 km)	MERRA (Nudging)	Hack (1994) (shallow) Zhang and MacFarlane (1995) (deep)
WACCM-C1	" "	" "	" "	Free-running	" "
CAM-C1SD	Community Atmosphere Model Version 4 (CAM4)-Chem Tilmes et al. (2015)	1.9° x 2.5°	56 (1 Pa)	MERRA (Nudging)	" "
CAM-C1	" "	" "	" "	Free-running	" "
EMAC-L47-C1	ECHAM/ Modular Earth Submodel System (MESSy) Atmospheric Chemistry (EMAC) Jöckel et al. (2010); Jöckel et al. (2016)	T42	47 (0.01 hPa)	Free-running	Tiedtke (1989); Nordeng (1994)
EMAC-L47-C1SD	" "	" "	" "	ERA-Interim (nudging)	" "
EMAC-L90-C1	" "	" "	90 (0.01 hPa)	Free-running	" "
EMAC-L90-C1SD	" "	" "	" "	ERA-Interim (nudging)	" "
MRI-C1SD	Earth System Model MRI-ESM1r1 Yukimoto et al. (2012, 2011); Deushi and Shibata (2011)	TL159	80 (0.01 hPa)	JRA-55 (Nudging)	Yoshimura et al. (2015)
MRI-C1	" "	" "	" "	Free-running	" "
CMAM-C1SD	Canadian Middle Atmosphere Model (CMAM) Jonsson et al. (2004); Scinocca et al. (2008)	T47	71 (0.0008 hPa)	ERA-Interim (Nudging)	Zhang and McFarlane (1995)
CMAM-C1	" "	" "	" "	Free-running	" "
NIWA-C1	National Institute of Water and Atmospheric Research UK Chemistry and Aerosols (NIWA-UKCA) Morgenstern et al. (2009, 2013); Stone et al. (2016)	3.75° x 2.5°	60 (84 km)	Free-running	Hewitt et al. (2011)
SOCOL-C1	Solar-Climate-Ozone Links (SOCOL) v3 Stenke et al. (2013); Revell et al. (2015)	T42	39 (0.01 hPa)	Free-running	Nordeng (1994)
NIES-C1SD	CCSRNIES-MIROC3.2 Imai et al. (2013); Akiyoshi et al. (2016)	T42	34 (0.01 hPa)	ERA-Interim (Nudging)	Arakawa and Schubert (1974)
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MOCAGE-CTM	Modele de Chimie Atmosphérique de Grande Echelle (MOCAGE) Josse et al. (2004); Guth et al. (2016)	2° x 2°	47 (5 hPa)	ERA-Interim (CTM)	Bechtold et al. (2001)
ULAQ-C1	University of L'Aquila (ULAQ)-CCM Pitari et al. (2014)	T21	126 (0.04 hPa)	Free-running	Grewe et al. (2001)
ACCESS-C1	National Institute of Water and Atmospheric Research UK Chemistry and Aerosols (NIWA-UKCA) Morgenstern et al. (2009, 2013); Stone et al. (2016)	3.75° x 2.5°	60 (84 km)	Free-running	Hewitt et al. (2011)

Among the hindcast runs (REF-C1, REF-C1SD) we consider 23 simulations, performed in both specified-dynamics (---) and free-running (—) modes.

# Models: Future Experiment

Simulation Name	Model (Reference)	Horizontal Resolution	Vertical Levels (Model Top)	Large-Scale Flow (Free/Nudging/CTM)	Convective Parameterization
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EMAC-L90-C1	" "	" "	90 (0.01 hPa)	Free-running	" "
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Among the future runs (REF-C2, RCP 6.5) we consider nine simulations that integrated the NH midlatitude idealized tracers.

# Models: Future Experiment

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## ***I. Transport to the Arctic***

— Hindcast  
— Future

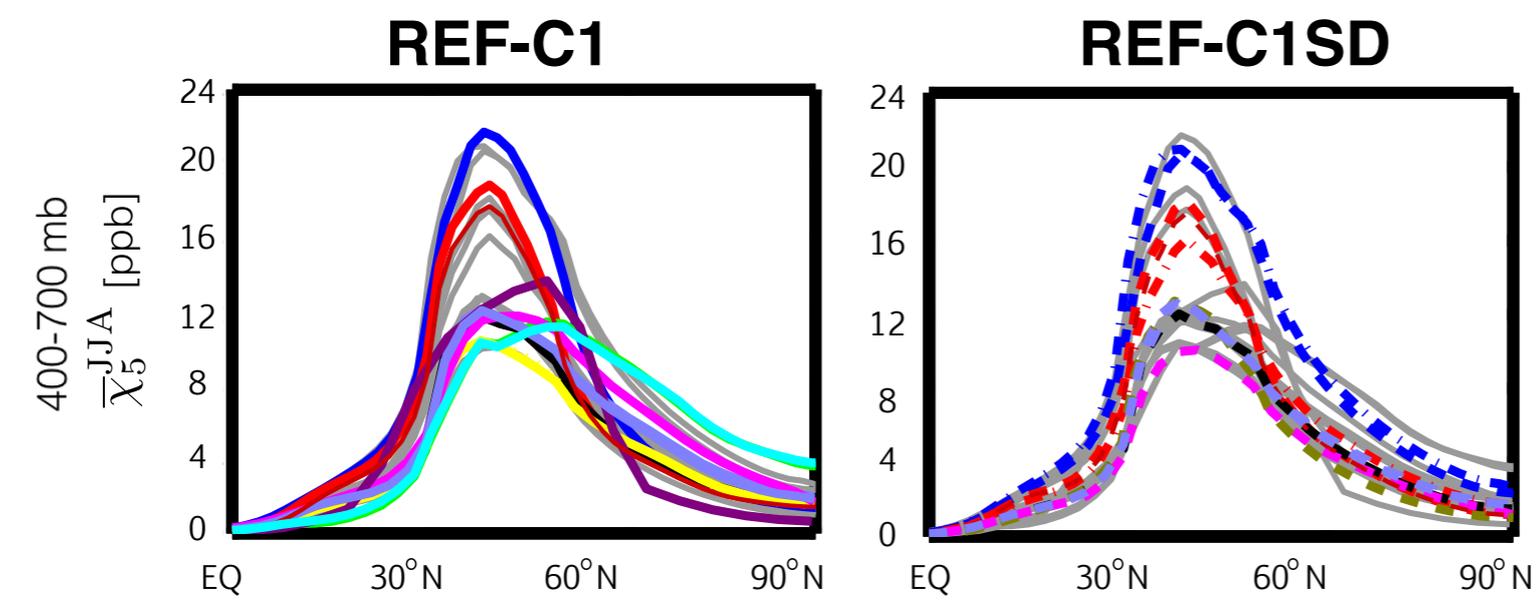
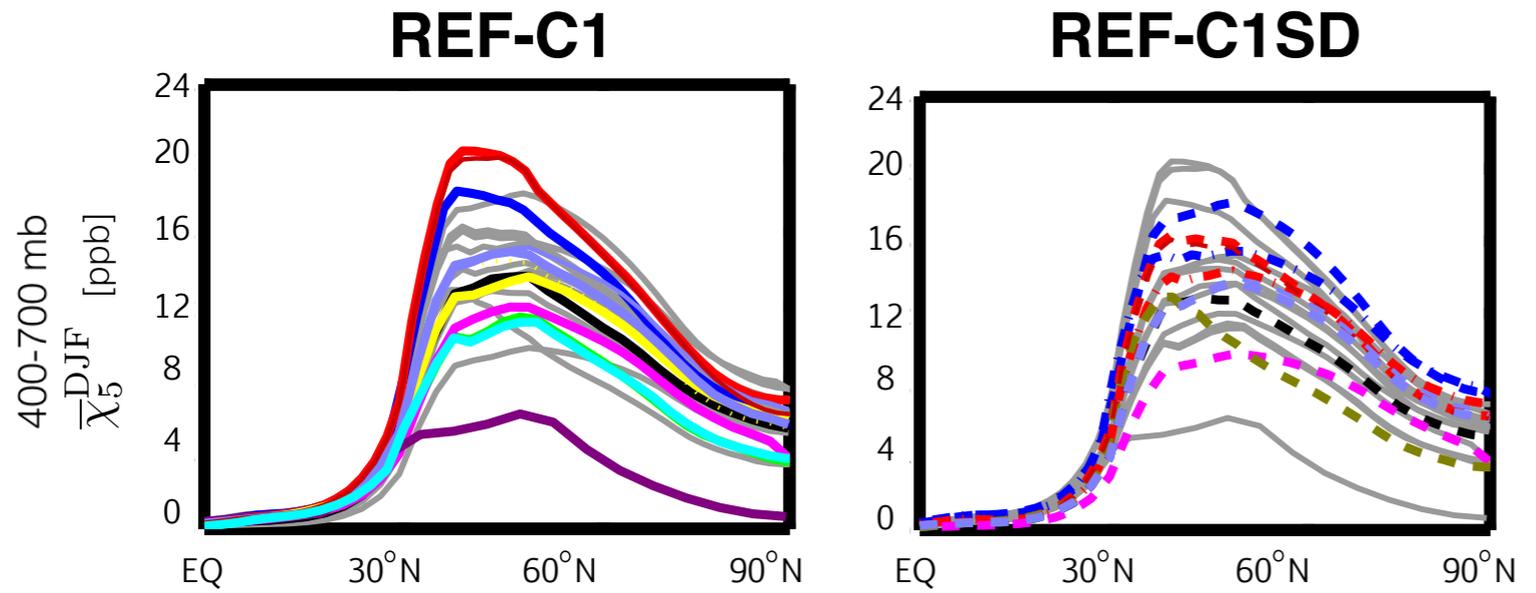
## ***II. Interhemispheric Transport***

— Hindcast  
— Future

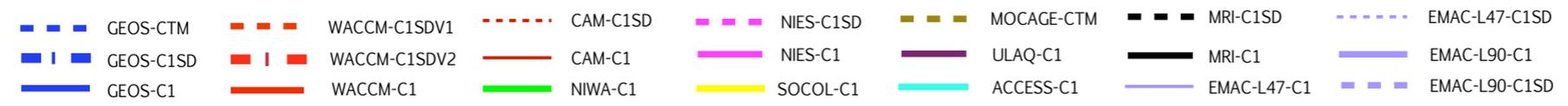
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# Transport to the Arctic

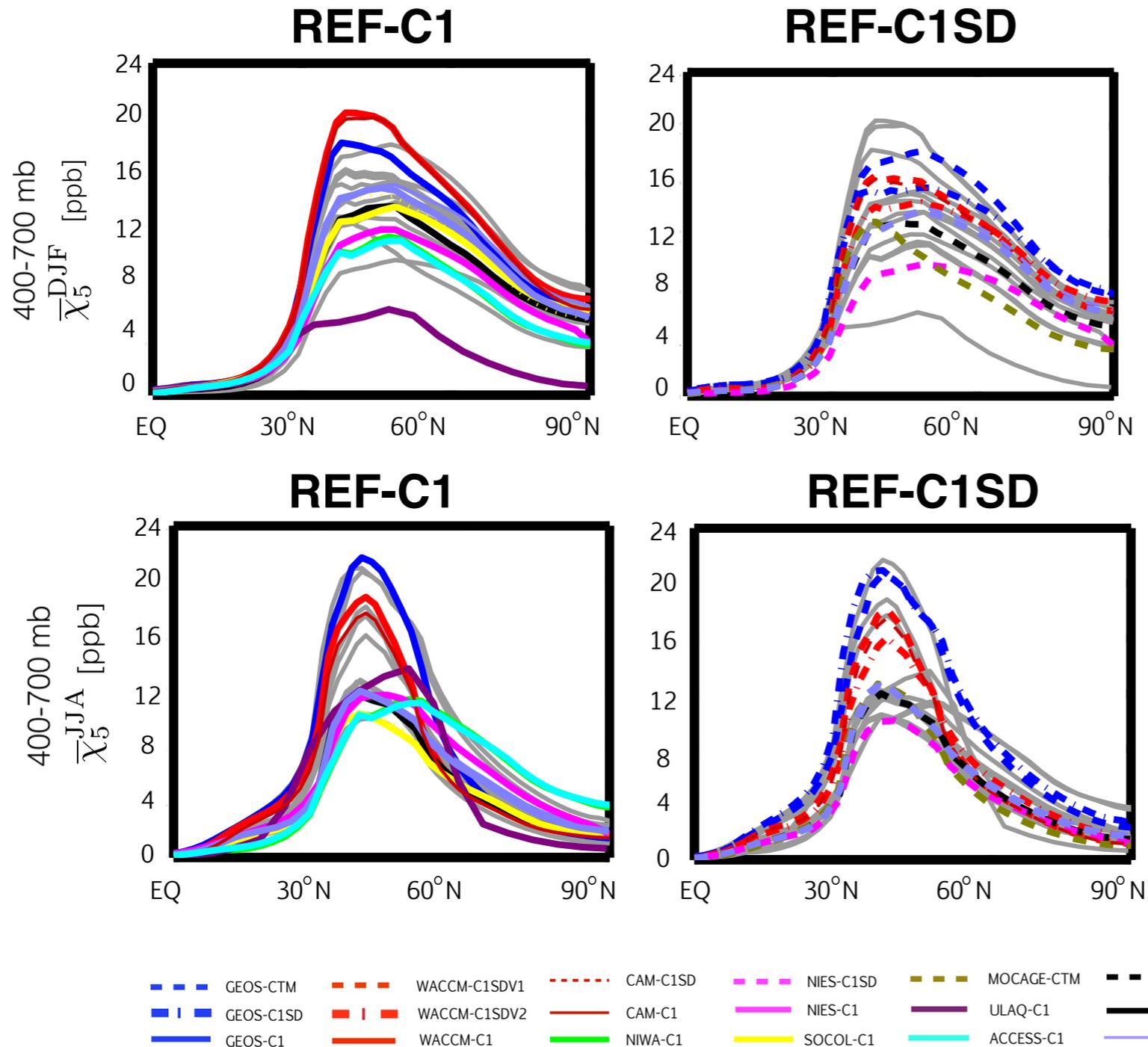


- Large (~30-40%) differences in transport over NH middle and high latitudes.
- The differences among SD simulations are as large (and at places larger) than the differences among FR simulations.



*Orbe et al. (2018, ACP)*

# Transport to the Arctic



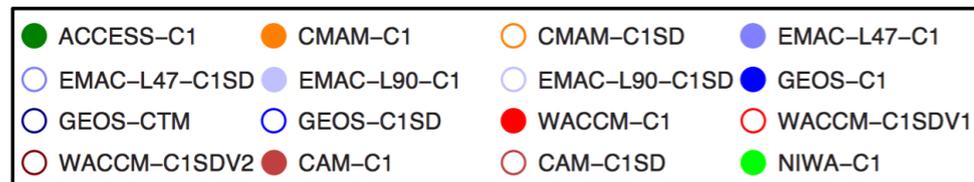
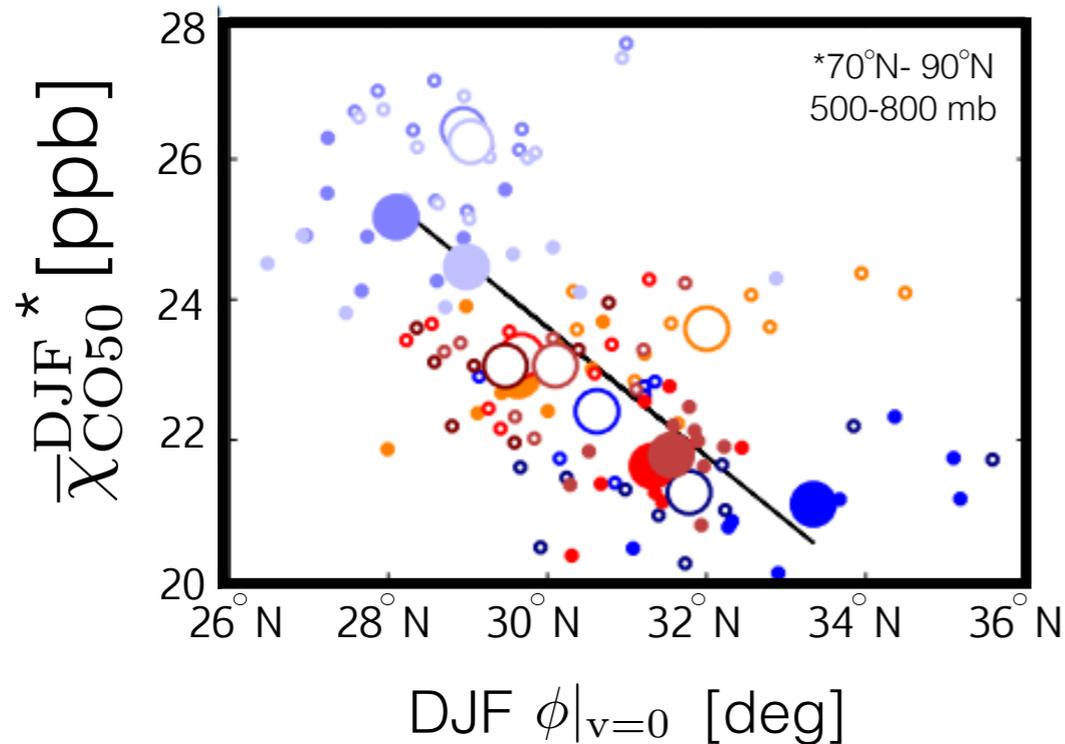
- Similar differences between SD and FR simulations are also exhibited by tracers with only land emissions ( $\chi_{CO50}$ ) (Yang et al. (2018, Under Review in ACPD)).

Orbe et al. (2018, ACP)

# Transport to the Arctic

## Land-Only Sources

Tracer vs. Hadley Cell Edge



#1 For tracers with land-only emissions ( $\overline{\chi_{CO50}^{DJF}}$ ), transport efficiency to the Arctic depends sensitively on the poleward edge of the Hadley Cell (*Yang et al. (2018), Under Review*)

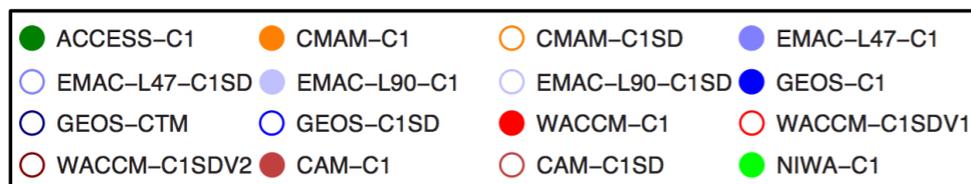
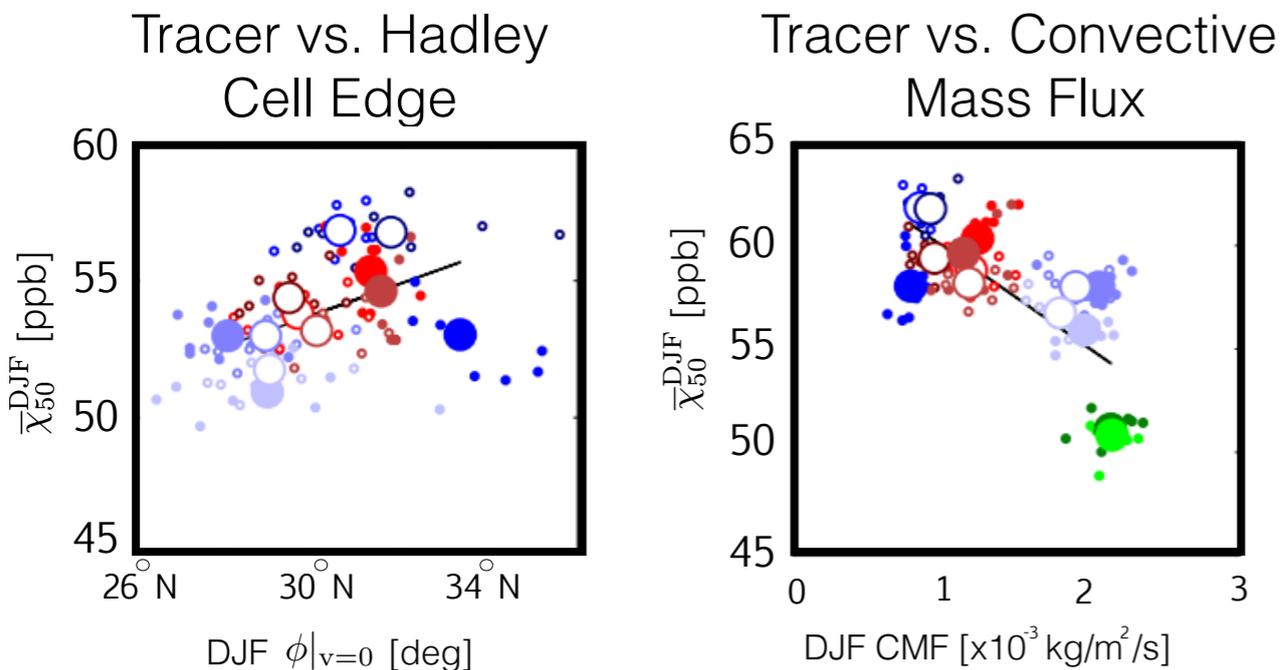
○ Specified-Dynamics

● Free-Running

# Transport to the Arctic

## Land and Ocean Sources

#2 By comparison, tracers with both ocean and land sources ( $\bar{\chi}_{50}^{\text{DJF}}$ ) depend also on convection over oceans (*Orbe et al. (2018), Yang et al. (Under Review)*) and less sensitively on midlatitude jet location and/or Hadley Cell edge.

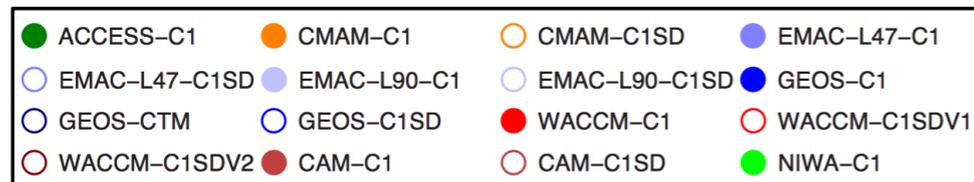
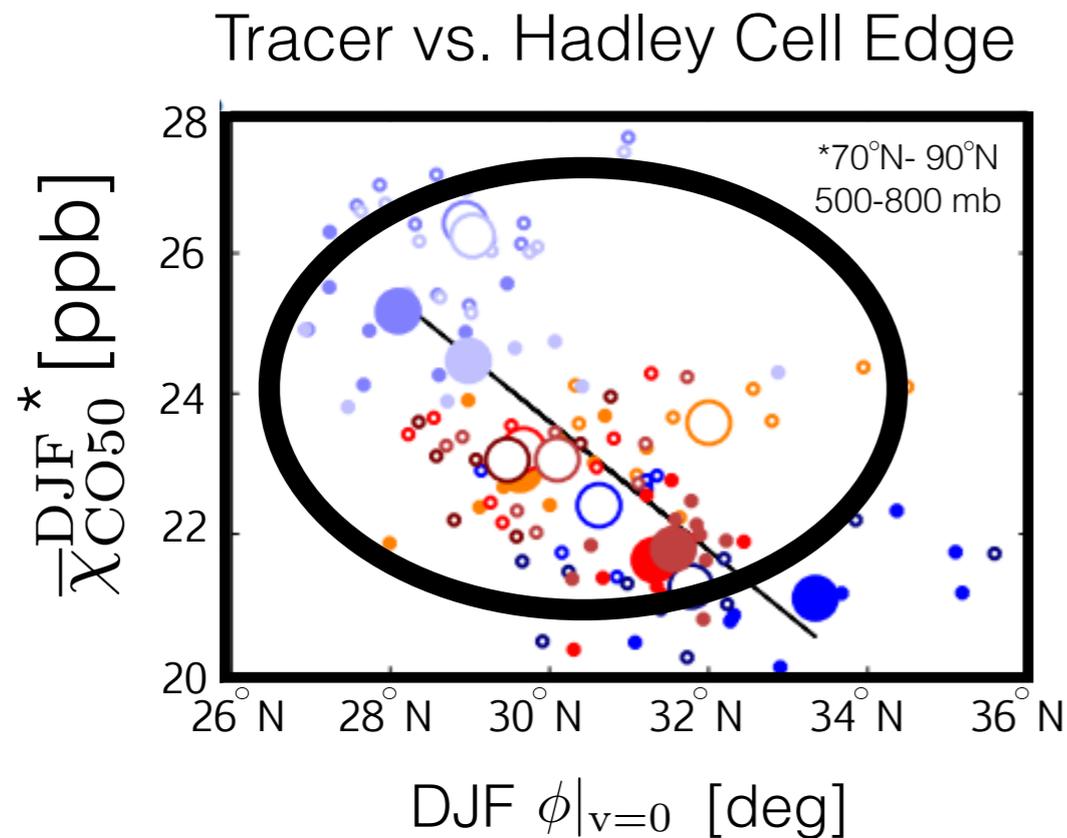


○ Specified-Dynamics

● Free-Running

# Transport to the Arctic

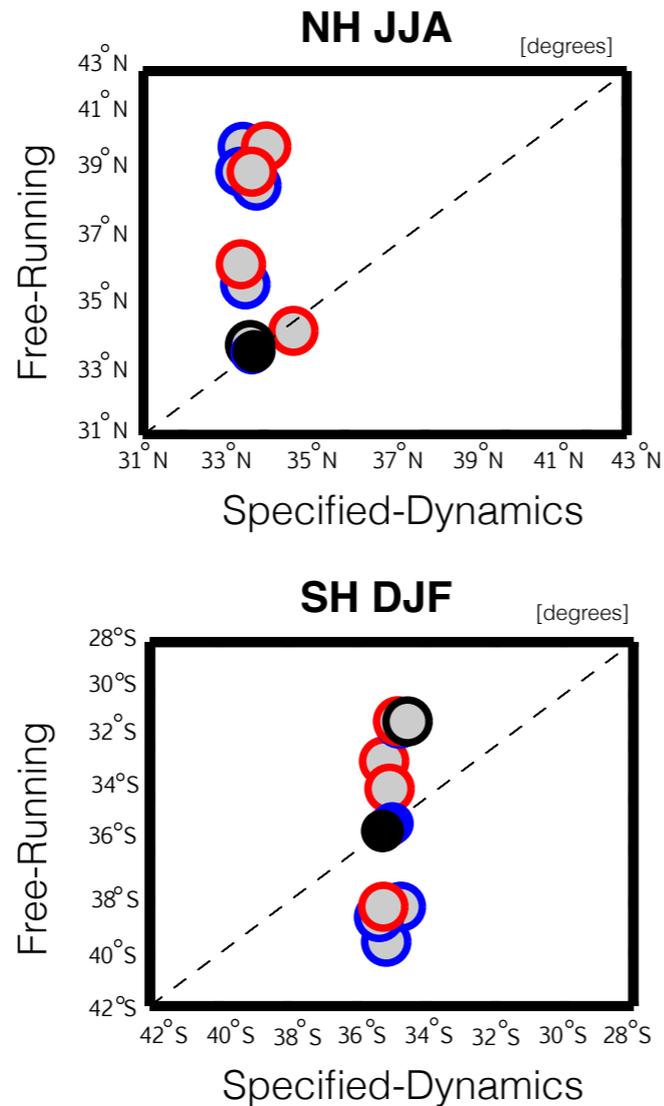
## Land-Only Sources



Note that the differences in Hadley Cell edge among specified-dynamics simulations (○) are as large as the differences among free-running simulations (●). This is somewhat surprising.

# Transport to the Arctic

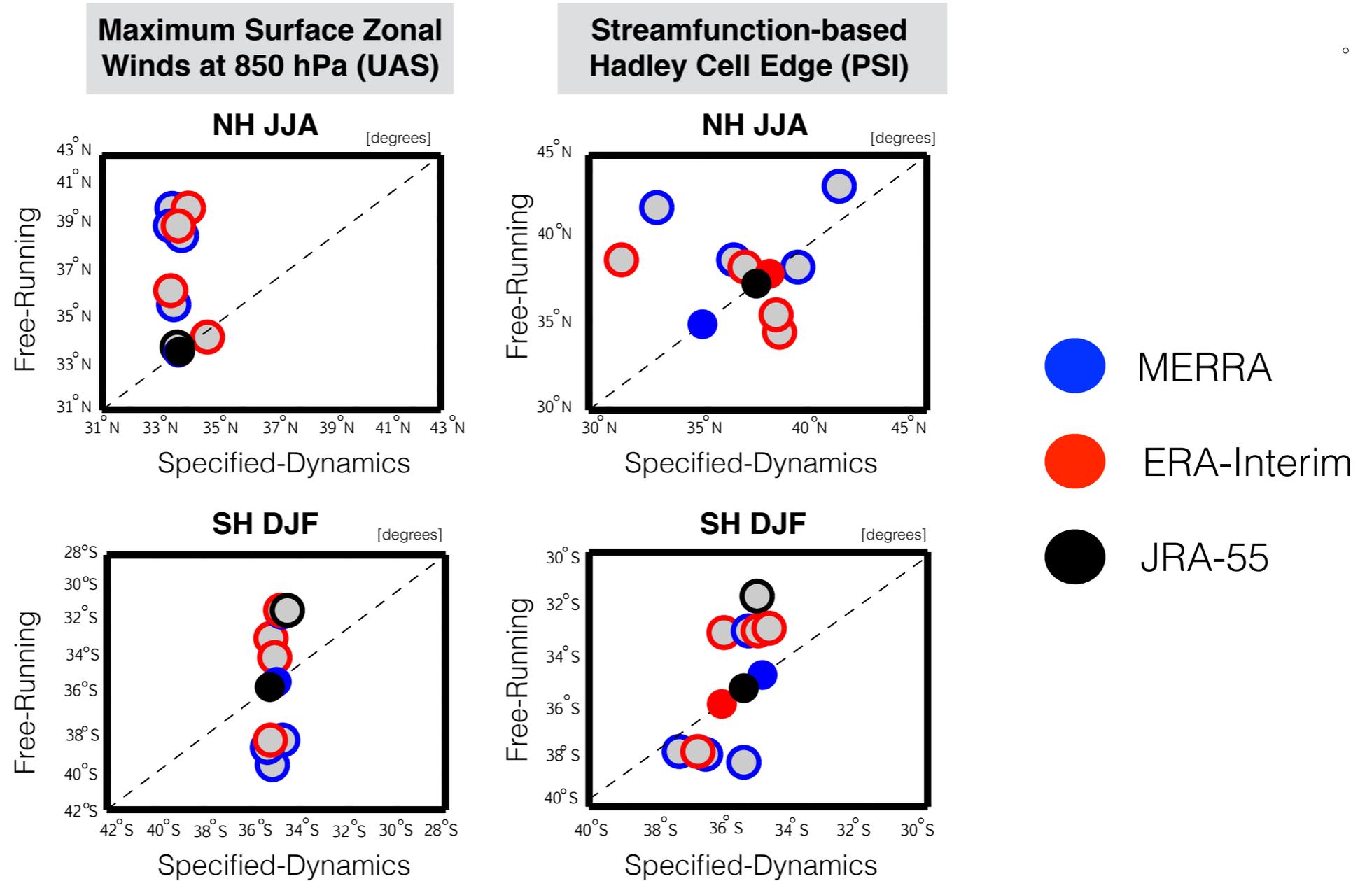
Maximum Surface Zonal Winds at 850 hPa (UAS)



- MERRA
- ERA-Interim
- JRA-55

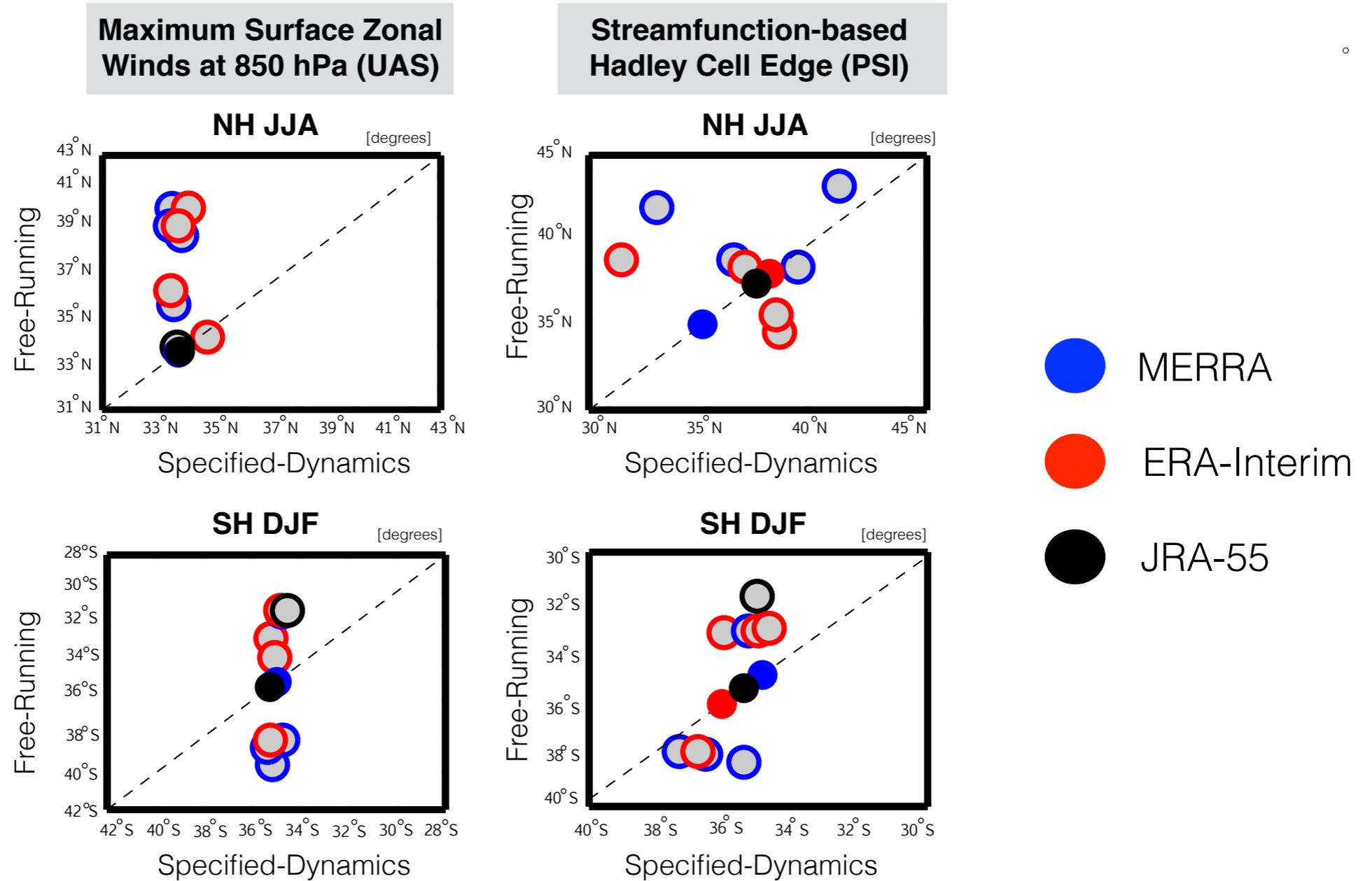
This reflects the fact that, while the zonal winds are well constrained in specified-dynamics simulations...

# Transport to the Arctic



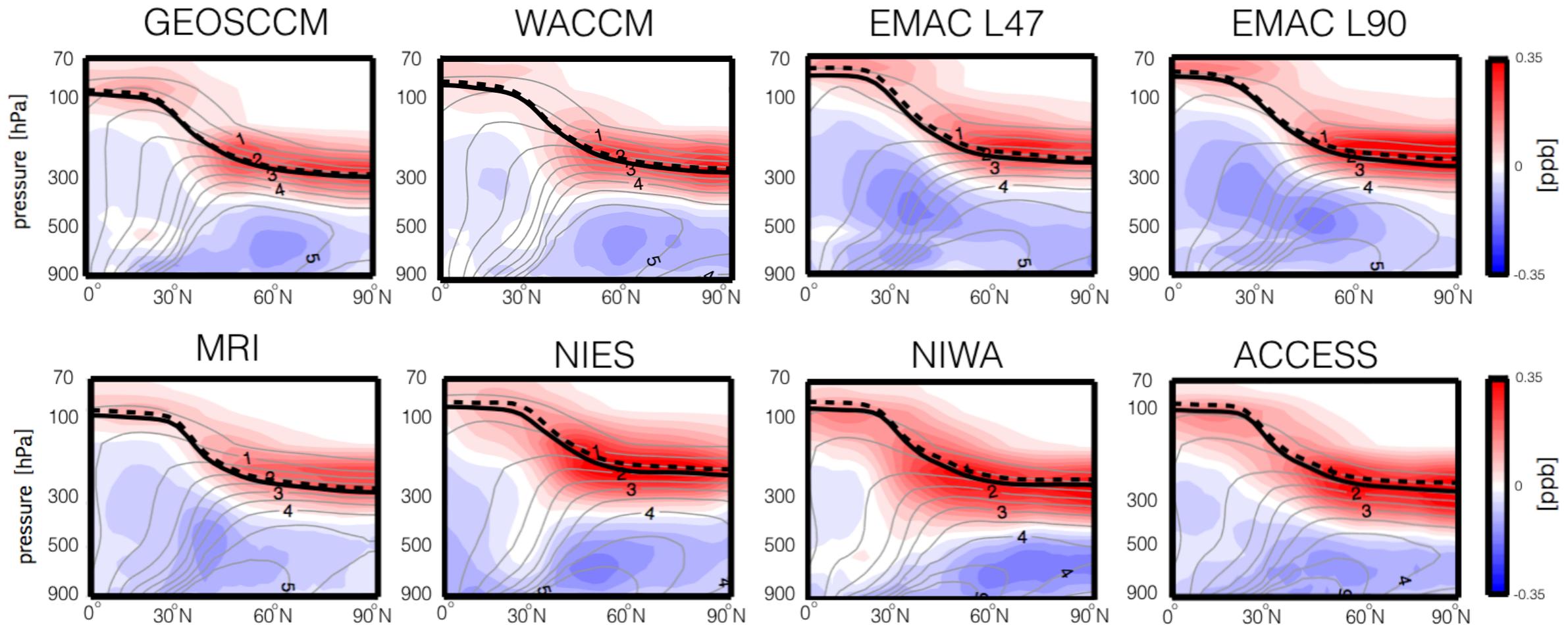
This reflects the fact that, while the zonal winds are well constrained in specified-dynamics simulations, the meridional and vertical component of the flow is not (*Orbe, Plummer et al., In Prep*).

# Transport to the Arctic



Note that differences among specified-dynamics simulations are not obviously related to the use of different analysis products (●, ●, ●), but rather to how the fields are implemented.

# Changes in Transport to the Arctic over the 21<sup>st</sup> Century

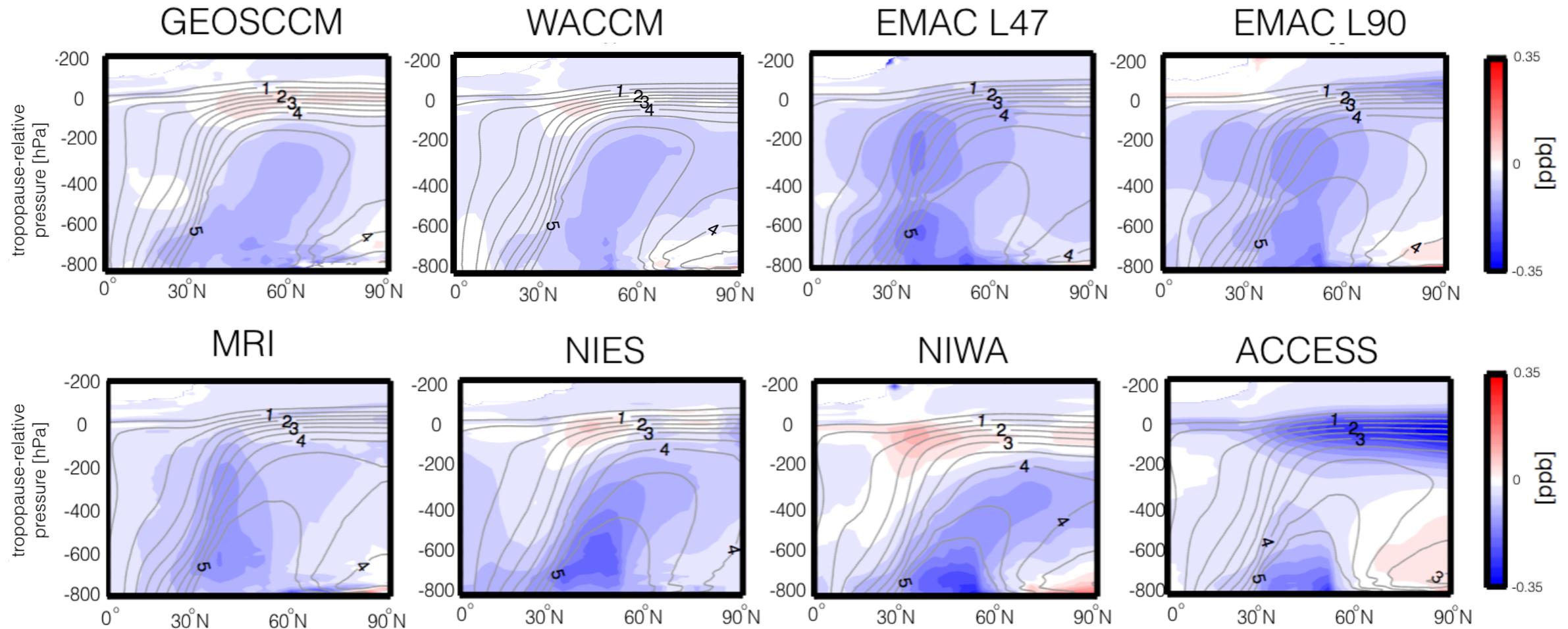


Robust response among CCMI models:

#1 Increased concentrations at the tropopause (—) and UTLS

#2 Reduced concentrations throughout the troposphere

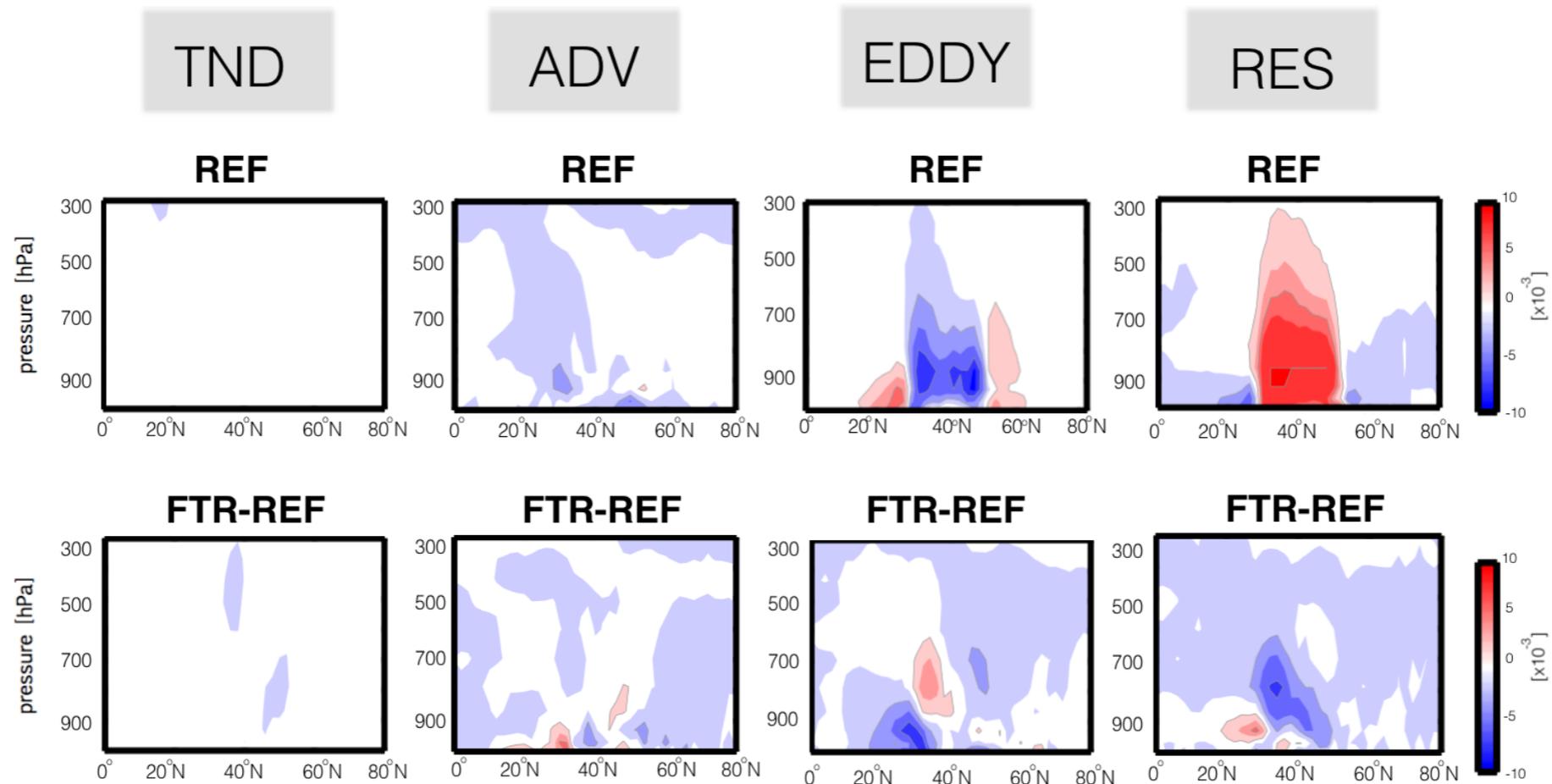
# Changes in Transport to the Arctic over the 21<sup>st</sup> Century



#1 Increased concentrations at the tropopause primarily reflect an increase in tropopause height (*Holzer and Boer (2001), Fang et al. (2011), Doherty et al. (2017), Abalos et al. (2017)*).

#2 Reduced concentrations throughout the troposphere persist.

# Changes in Transport to the Arctic over the 21<sup>st</sup> Century

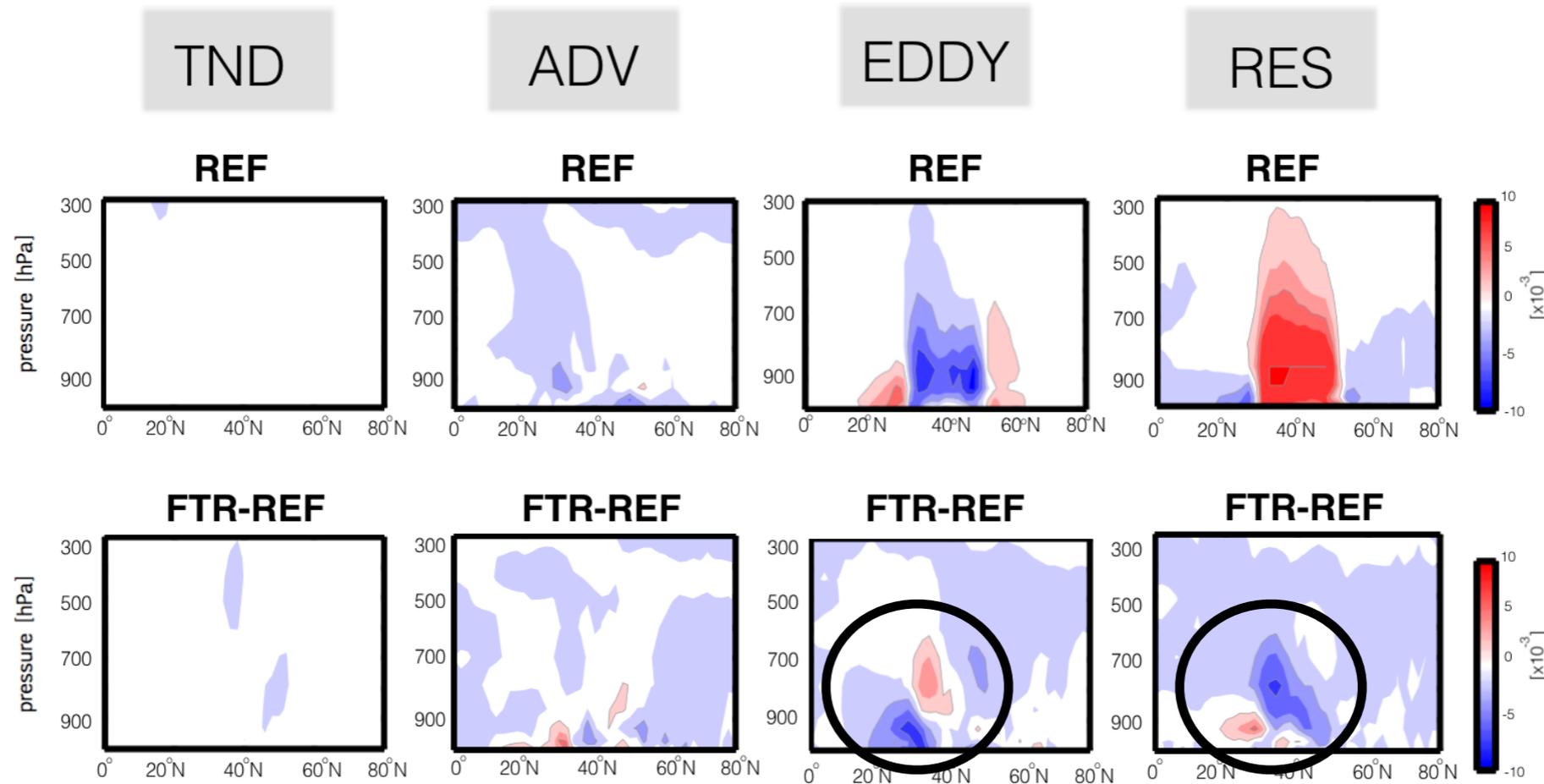


$$\bar{\chi}_t = -\bar{v}^* \bar{\chi}_y - \bar{w}^* \bar{\chi}_z + \nabla \cdot \mathbf{M} + \bar{L} + \bar{X}$$

TND                  ADV                  EDDY                  RES

Tracer budgets, cast in terms of the Transformed Eulerian Mean as in *Abalos et al. (2017)*, indicate that loss tracer concentrations primarily reflect a balance between eddy-induced mixing and transport by (parameterized) convection.

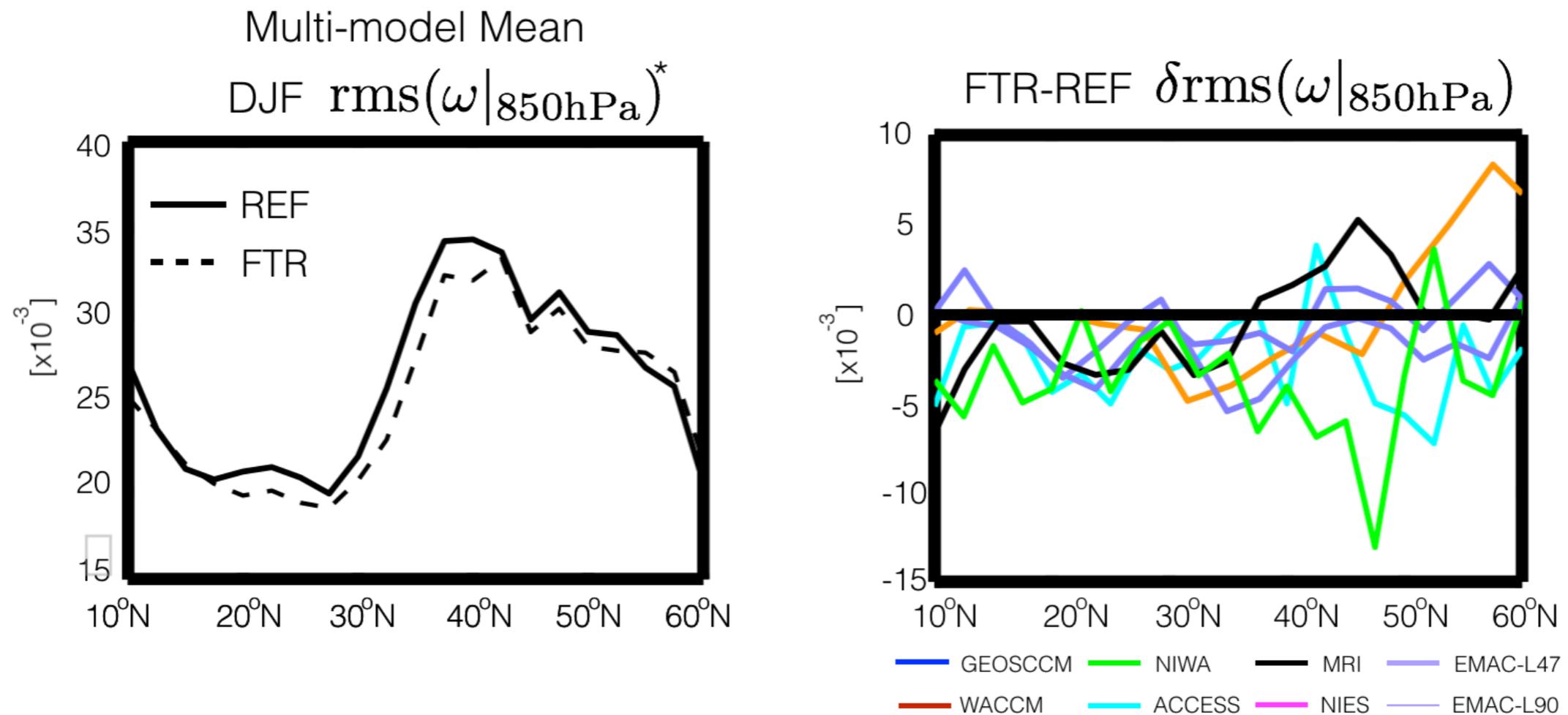
# Changes in Transport to the Arctic over the 21<sup>st</sup> Century



$$\bar{\chi}_t = \underbrace{-\bar{v}^* \bar{\chi}_y}_{\text{TND}} - \underbrace{\bar{w}^* \bar{\chi}_z}_{\text{ADV}} + \underbrace{\nabla \cdot \mathbf{M}}_{\text{EDDY}} + \underbrace{\bar{L} + \bar{X}}_{\text{RES}}$$

Changes in budget terms indicate that reduced concentrations of loss tracers are associated with reduced vertical transport by both eddies and convection, not by changes in the mean circulation.

# Changes in Transport to the Arctic over the 21<sup>st</sup> Century



Consist with both reduced convective mass fluxes in the future (*Held and Soden (2006)*) as well as robust decreases in lower tropospheric vertical motion in stationary eddies (*Wills and Schneider, 2016*).

# ***I. Transport to the Arctic***

## **Hindcast (1960-2010) simulations show:**

-Large differences in transport to high latitudes among *both* specified-dynamics and free-running simulations.

-Poleward extent of the Hadley Cell controls the poleward transport of tracers emitted only over land, whereas ocean convection matters more for tracers with ocean sources.

-Certain measures of the Hadley Cell are poorly constrained in specified-dynamics simulations, consistent with large differences in meridional transport.

**Orbe, C., Yang, H., Waugh, D. W., Zeng, G., Morgenstern, O., Kinnison, D. E. et al. (2018). Large-scale tropospheric transport in the Chemistry-Climate Model Initiative (CCMI) simulations. *Atmospheric Chemistry and Physics*, 18(10), 7217-7235.**

**Yang, H., Waugh, D. W., Orbe, C., Zeng, G., Morgenstern, O., Kinnison, D. E. et al. (2018). Tracer Transport into the Arctic: Relative Roles of the Midlatitude Jet and the Hadley Cell Edge. *Under Review in Atmospheric Chemistry and Physics Discussions*.**

**Orbe, C., D. Plummer., Waugh, D. W., Yang H., and CCMI Co-authors, Description of the Specified-Dynamics Experiment in the Chemistry Climate Model Initiative (CCMI) (*In Prep*)**

# ***I. Transport to the Arctic***

## **Future (1960-2100) simulations show:**

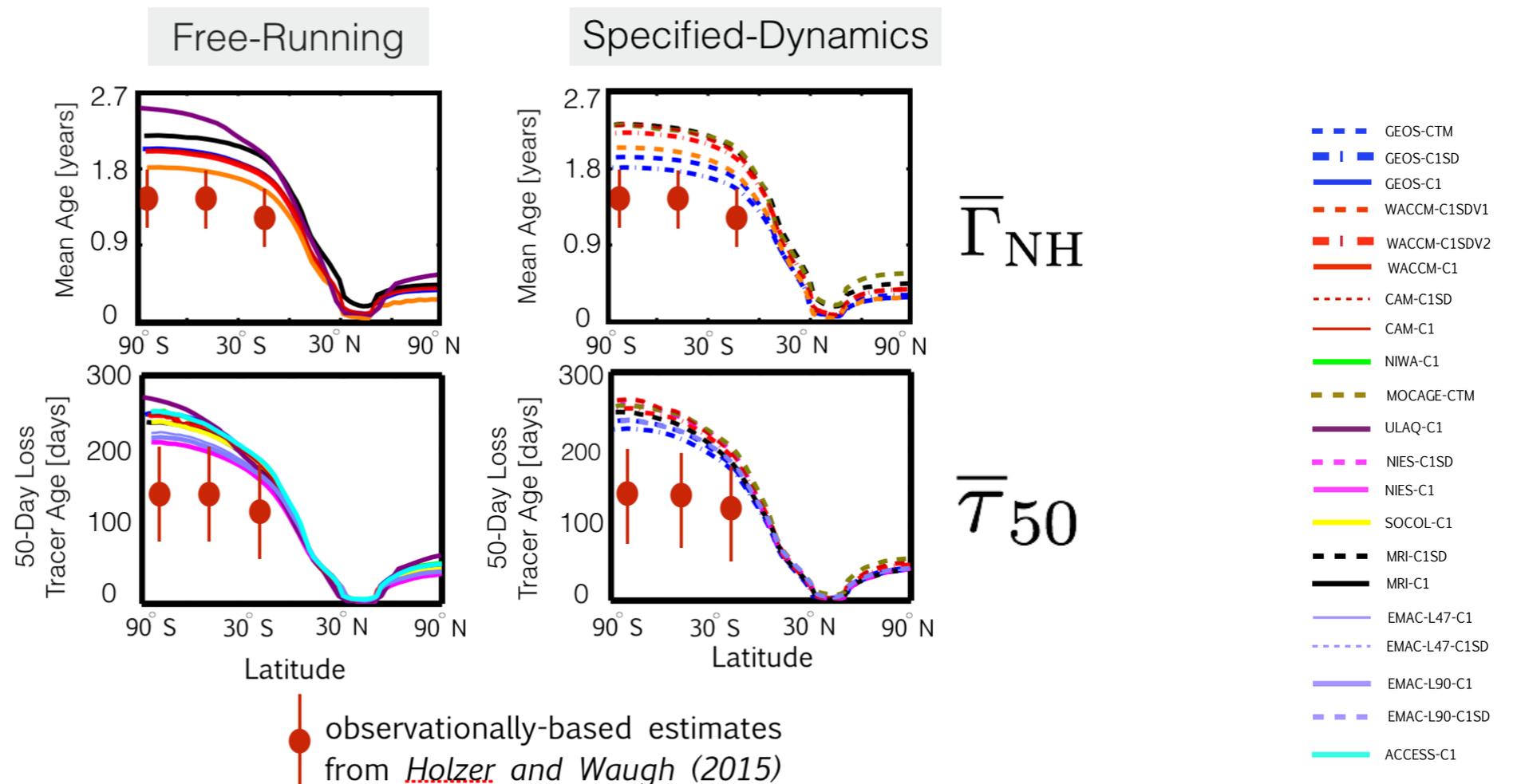
- Consistent increase in transport of NH midlatitude source tracers into the tropopause/lower stratosphere, primarily due to an increase in tropopause height.
- Reduced vertical transport out of the lower troposphere, consistent with weaker vertical eddies and reduced convective mass fluxes.

**Orbe, C., Abalos M., Waugh, D. W., Wang H., et al. Future Projections of Large-scale Tropospheric Transport Changes in the Chemistry-Climate Model Initiative (CCMI) simulations, (*In Prep*).**

# Interhemispheric Transport

- The differences in interhemispheric transport are also large (30-40%) and not better constrained among the SD simulations (versus FR).

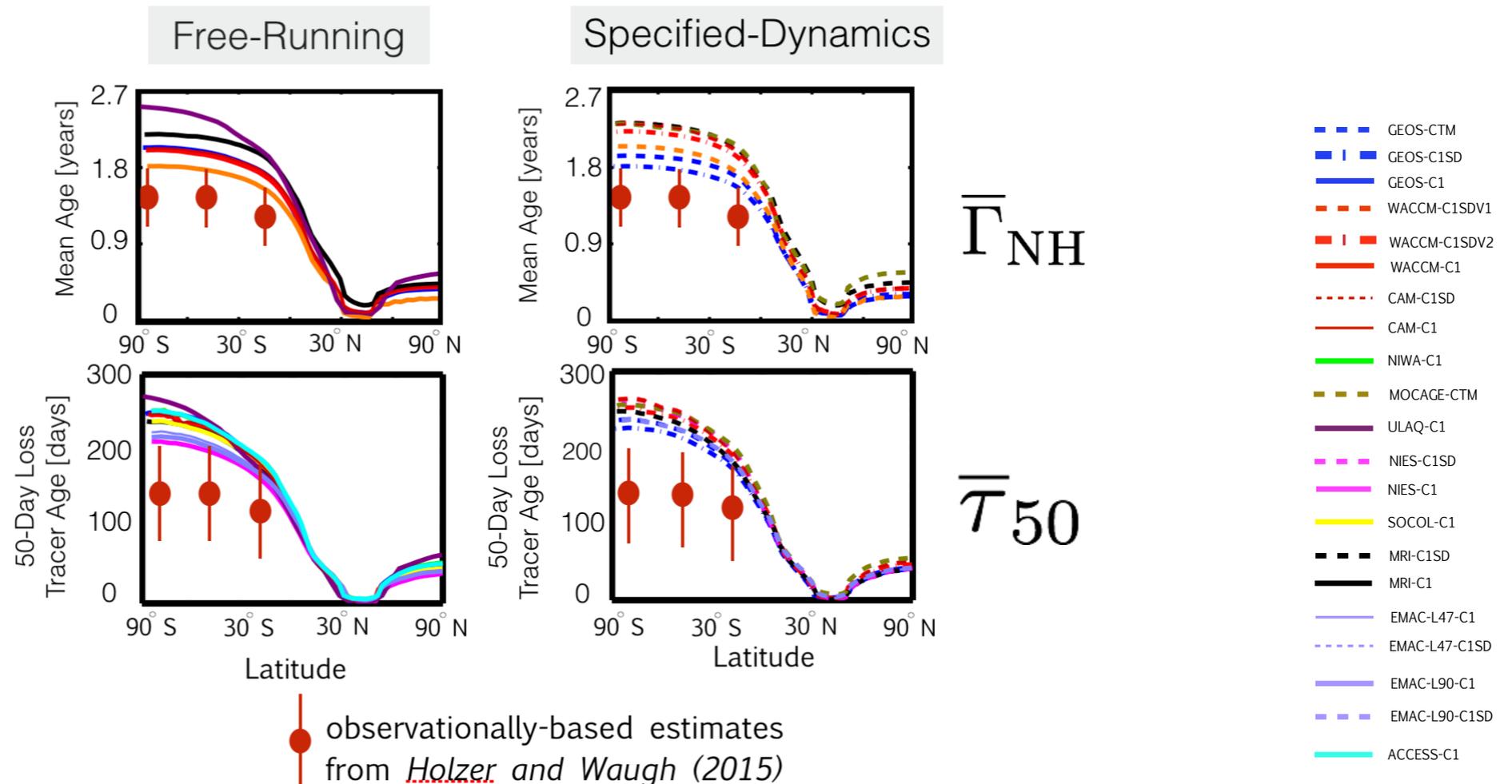
900-1000 mb Annual Mean 50-Day Tracer Ages ( $\bar{\tau}_{50}$ )  
and the NH Midlatitude Mean Age ( $\bar{\Gamma}_{NH}$ )



# Interhemispheric Transport

- Sparse observationally-based estimates of the mean age and 50-day tracer age ( $\bullet$ ) indicate that all models tend to feature too slow transport (*Holzer and Waugh (2015)*).

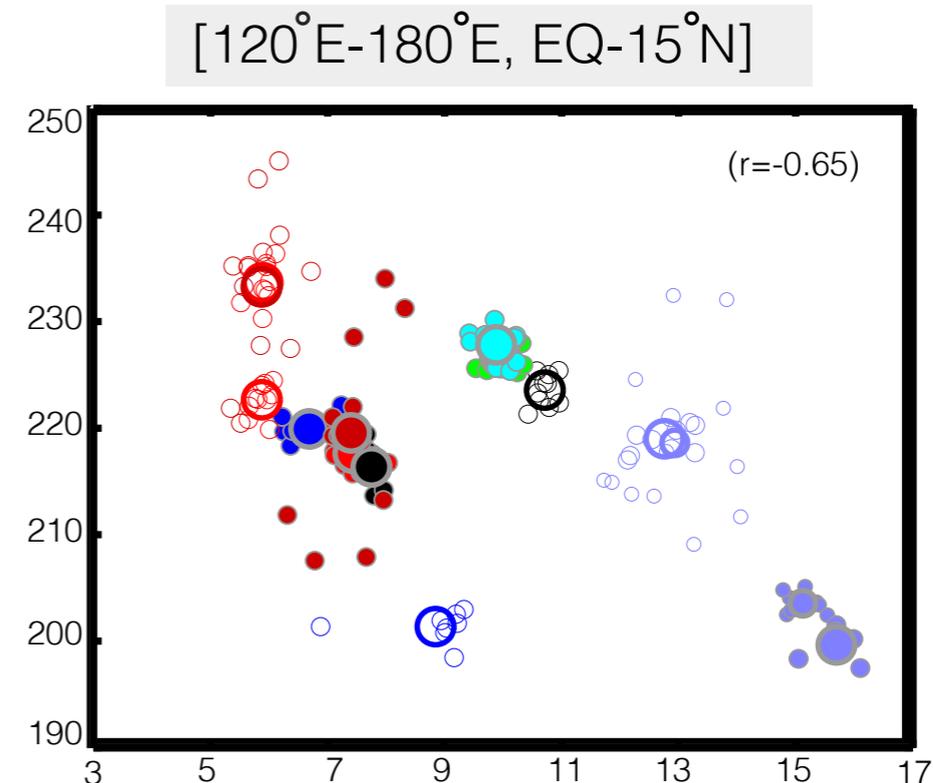
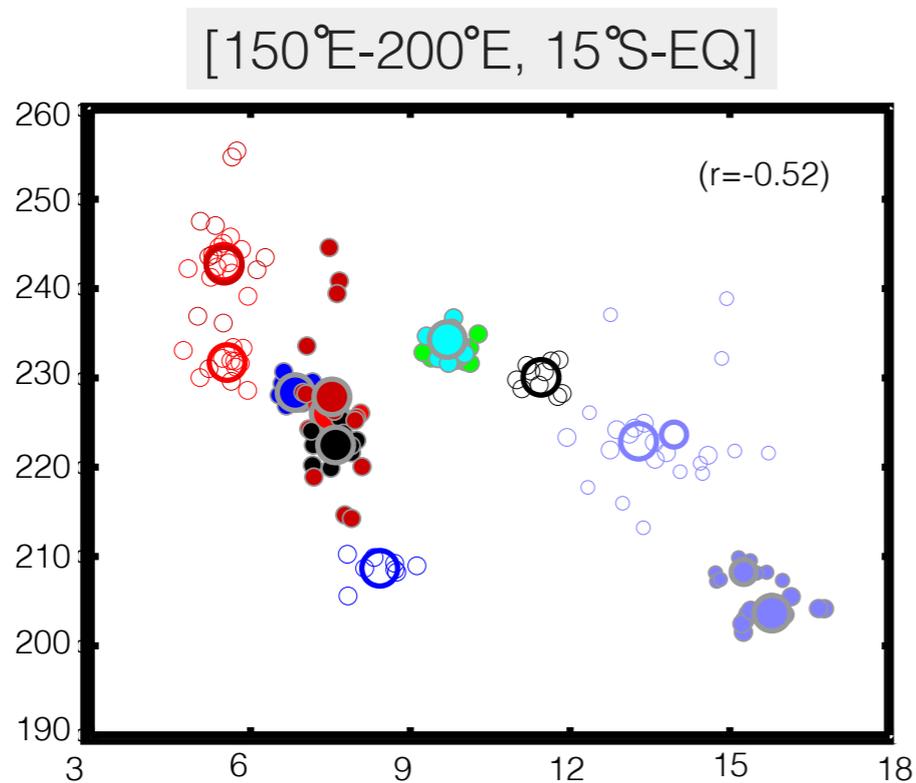
900-1000 mb Annual Mean 50-Day Tracer Ages ( $\bar{\tau}_{50}$ )  
and the NH Midlatitude Mean Age ( $\bar{\Gamma}_{NH}$ )



# Interhemispheric Transport

- In the annual mean, SH tracer age differences correlate best with differences in (parameterized) convection in the tropics and northern subtropics, particularly over the Pacific Ocean.

SH 60°N-90°N 50-Day Age

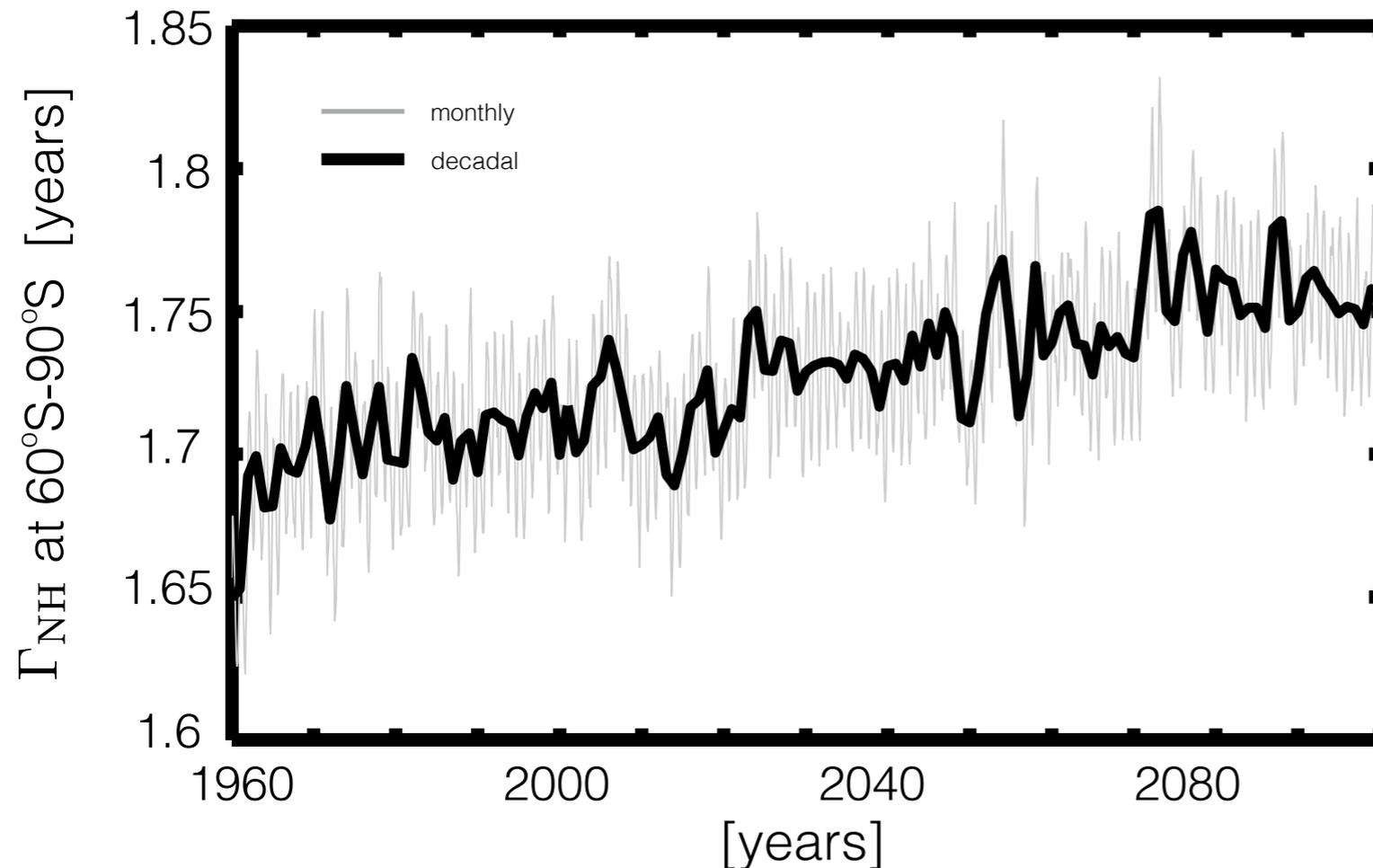


Annual Mean Convective Mass Flux (CMF) [ $\times 10^{-3}$ ] [kg/m<sup>2</sup>/s]



# Changes in Interhemispheric Transport over the 21<sup>st</sup> Century

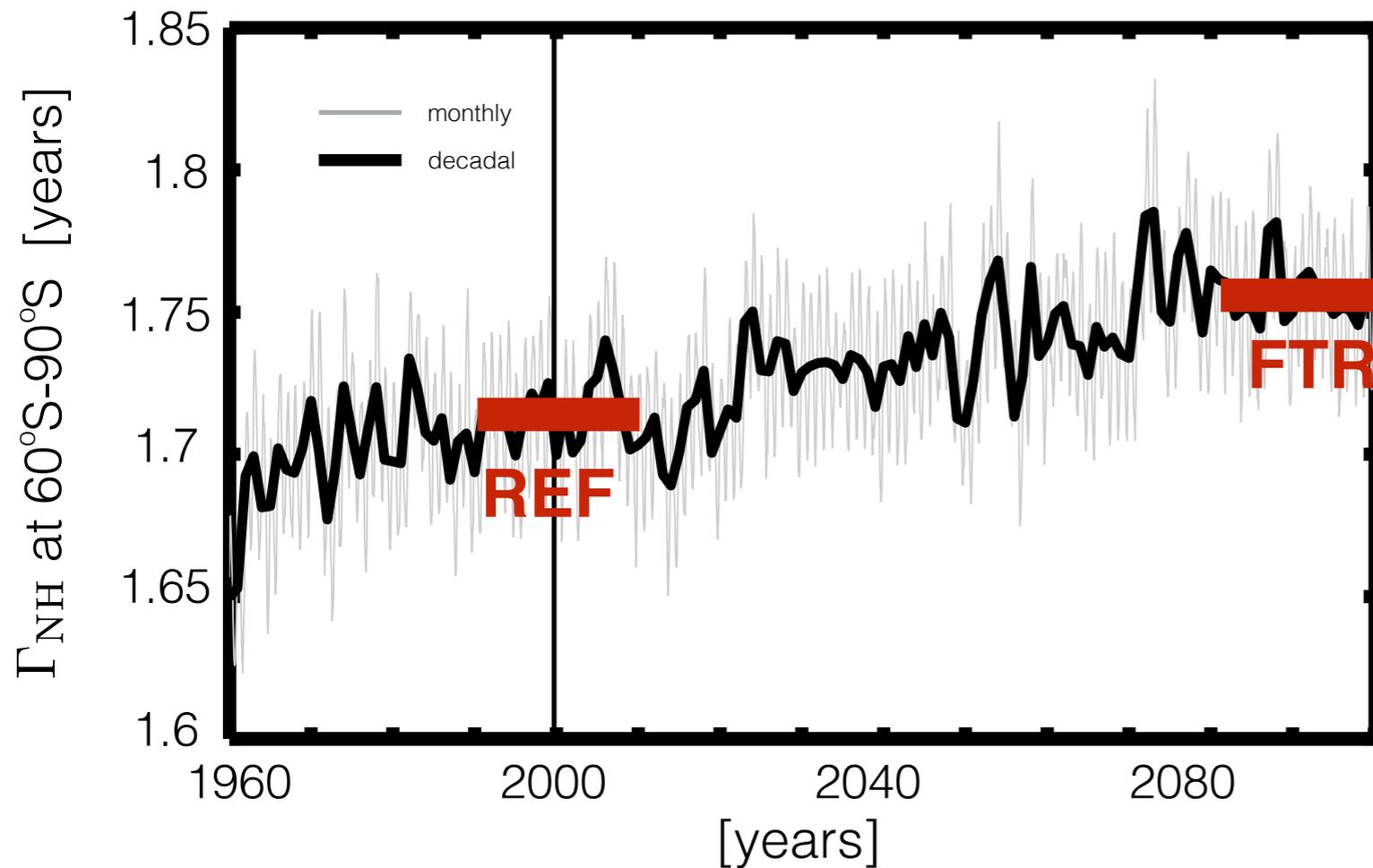
1960-2099 NH Midlatitude Mean Age over the Southern Pole\*



- Some models project slower (5-10%) interhemispheric transport, consistent with previous studies (*Holzer and Boer (2001)*).

# Changes in Interhemispheric Transport over the 21<sup>st</sup> Century

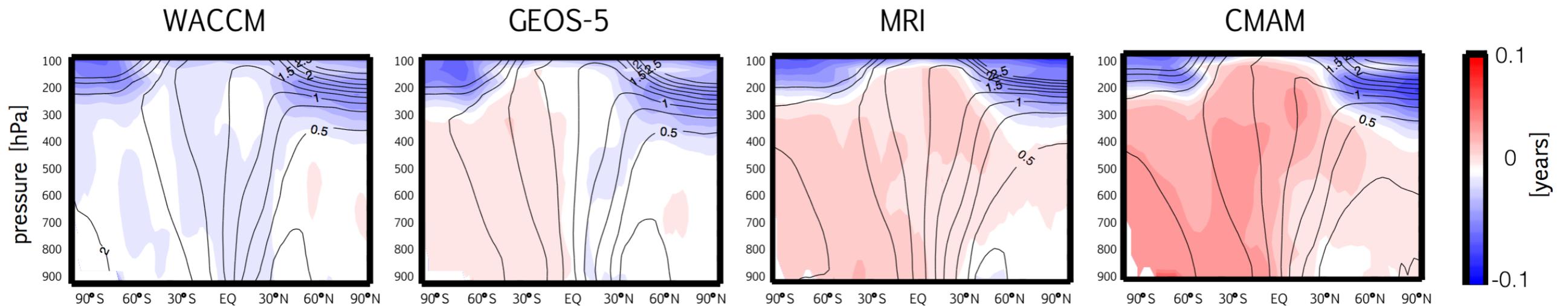
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# Changes in Interhemispheric Transport over the 21<sup>st</sup> Century

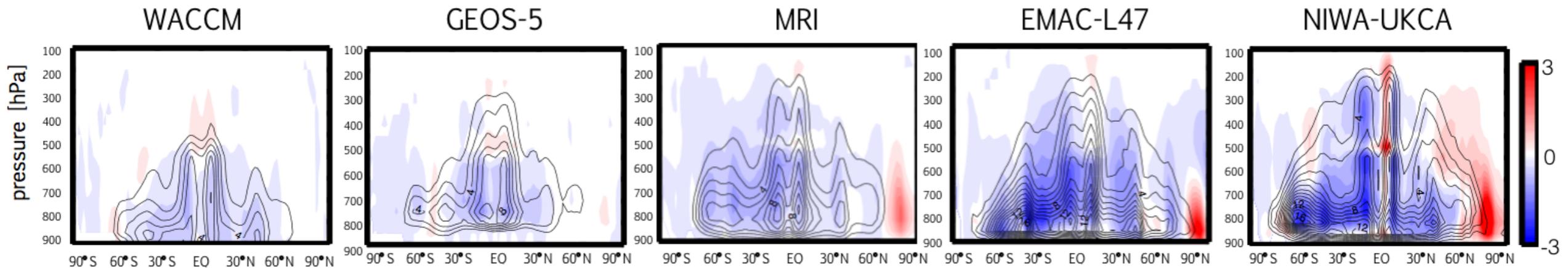
FTR-REF\* Changes in Northern Midlatitude Mean Age ( $\Gamma_{NH}$ )



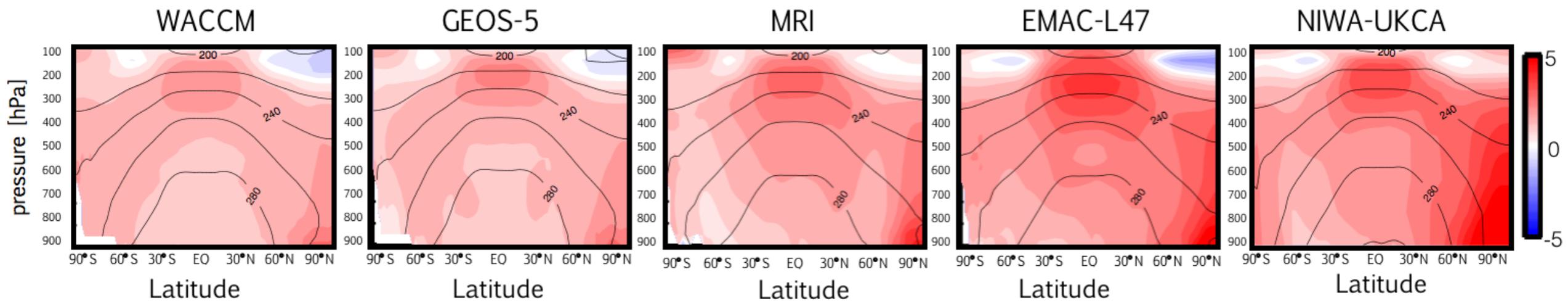
Lack of agreement among models, however, with some showing stronger responses than others.

# Changes in Interhemispheric Transport over the 21<sup>st</sup> Century

FTR-REF\* Convective Mass Flux Changes (CMF:  $10^{-3} \text{ kg/m}^2/\text{s}$ )



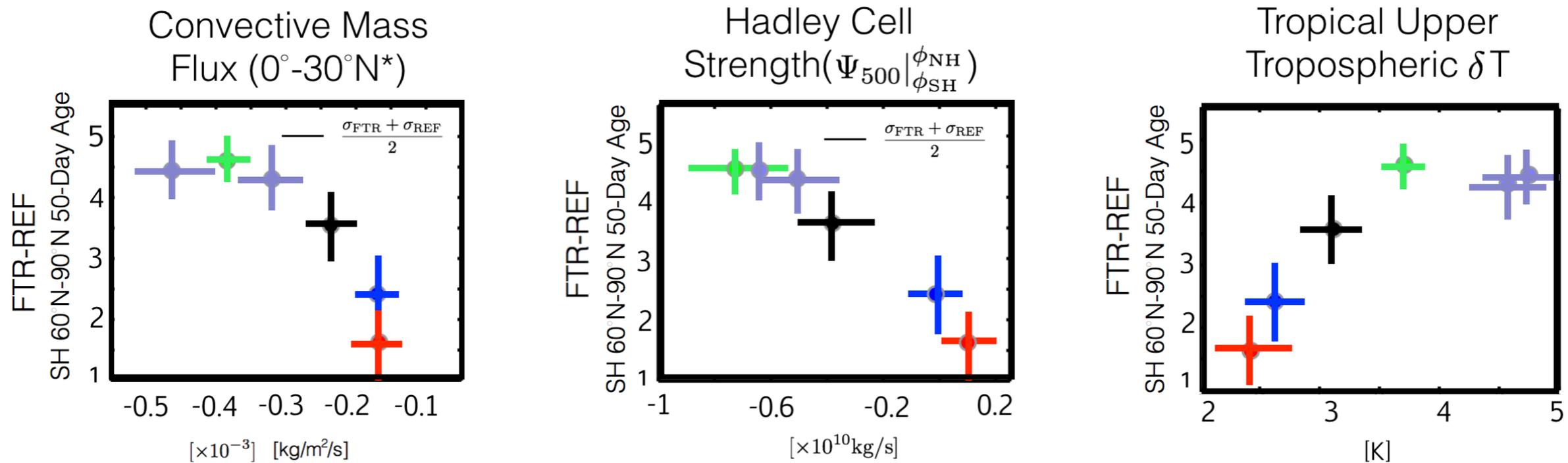
FTR-REF\* Temperature Changes [K]



\*2080-2100 - 1990-2010

Changes in interhemispheric transport are correlated with changes in convective mass fluxes and the amount of upper tropospheric tropical warming.

# Changes in Interhemispheric Transport over the 21<sup>st</sup> Century



Changes in interhemispheric transport are correlated with changes in convective mass fluxes and the amount of upper tropospheric tropical warming.



*(Orbe, Abalos et al. In Prep)*

## ***II. Interhemispheric Transport***

### **Hindcast (1960-2010) simulations show:**

-Large differences in interhemispheric transport among *both* specified-dynamics and free-running simulations.

-Strength of (sub)tropical convection is positively correlated with differences in the efficiency of interhemispheric transport among models.

**Orbe, C., Yang, H., Waugh, D. W., Zeng, G., Morgenstern, O., Kinnison, D. E. et al. (2018). Large-scale tropospheric transport in the Chemistry-Climate Model Initiative (CCMI) simulations. *Atmospheric Chemistry and Physics*, 18(10), 7217-7235.**

**Orbe, C., Plummer D., Waugh, D. W., Yang H., and CCMI Co-authors, Description of the Specified-Dynamics Experiment in the Chemistry Climate Model Initiative (CCMI) (*In Prep*)**

## ***II. Interhemispheric Transport***

### **Future (1960-2100) simulations show:**

-Weaker (~5-10%) interhemispheric transport by the end of the 21<sup>st</sup> century, although some models show no significant changes.

-Interhemispheric transport response is correlated with changes in the strength of lower tropospheric convection and the amount of upper tropospheric tropical warming.

**Orbe, C., Abalos M., Waugh, D. W., Wang H., et al. Future Projections of Large-scale Tropospheric Transport Changes in the Chemistry-Climate Model Initiative (CCMI) simulations. (*In Prep*).**

# Relevant Publications

## Published:

Orbe, C., Yang, H., Waugh, D. W., Zeng, G., Morgenstern, O., Kinnison, D. E. et al. (2018). Large-scale tropospheric transport in the Chemistry-Climate Model Initiative (CCMI) simulations. *Atmospheric Chemistry and Physics*, 18(10), 7217-7235.

Wu, X., Yang, H., Waugh, D.W., Orbe, C., Tilmes, S., and Lamarque J.F., “Spatial and Temporal Variability of Interhemispheric Transport Times,” *Atmospheric Chemistry and Physics* 18.10 (2018): 7439-7452.

## In Prep/Under Review:

Yang, H., Waugh, D. W., Orbe, C., Zeng, G., Morgenstern, O., Kinnison, D. E. et al. (2018). Tracer Transport into the Arctic: Relative Roles of the Midlatitude Jet and the Hadley Cell Edge. (*Under Review in Atmospheric Chemistry and Physics Discussions.*)

Orbe, C., Plummer D., Waugh, D. W., Yang H., et al., Description of the Specified-Dynamics Experiment in the Chemistry Climate Model Initiative (CCMI) (*In Prep*)

Orbe, C., Abalos M., Waugh, D. W., Wang H., et al. Future Projections of Large-scale Tropospheric Transport in the Chemistry-Climate Model Initiative (CCMI) simulations. (*In Prep*).