The Role of Meltwater in Greenland Ice Sheet Dynamics



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Arctic amplification



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Surface & dynamic ice mass changes

- Surface \rightarrow 70% Dynamic \rightarrow 30%
- Dynamic processes are numerous and are relatively unconstrained, but are critical to determining how much of the ice sheet behaves



Surface melting





850 hPa Temperature Difference from Zonal Mean [°C]

-11 -10 -9 -8 -7 -6 -5 -4 -3 -2 -1 0 1 2 3 4 5 6 7 8 9 10 11 12 20 m s

250 hPa Winds ->

Surface Melt

How do changes in surface melt alter ice dynamics?





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How do changes in surface melt alter ice dynamics?

Surface melting Supraglacial routing Surface-to-bed connections Ice flow variations & englacial storage Subglacial hydrology & pressure variations **Dynamic ice sheet** mass loss Ocean forcings

Processes in the subglacial, englacial, and supraglacial systems, plus changes in the spatiotemporal distribution of melt make the link between melt and ice dynamics complex (and interesting).

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Subglacial hydrology







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Subglacial hydrology



- Formed primarily due to sliding \bullet
- Slow water flow

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- **Operate at high pressure**
- Increase in flux increases lacksquare

subglacial pressure & ice flow



- Formed primarily due to melting
- Fast water flow
- **Operates at low pressure**
- Increase in flux decreases subglacial pressure & ice flow

Subglacial hydrology: inefficient drainage





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Subglacial hydrology: efficient drainage







Subglacial hydrology & ice motion





Subglacial hydrology & ice motion



Andrews et al. (2014)

Monitoring subglacial channels

- Large diurnal variability, specifically, low diurnal minima
- Downstream similarity a system response
- In phase with ice velocity





Hydraulic head (m)

Day of year, 2012



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An evolving relationship between ice velocity and subglacial channel pressure



Andrews et al. (2014)

Velocity evolution without long-term increases efficiency?



This subglacial channel is seemingly mature. We don't observe continuing increases in channel efficiency (i.e. no decreasing pressure observed in moulins). However, ice speed continues to drop, indicating evolution elsewhere in the subglacial system.



Subglacial evolution outside channelized regions

- Low amplitude diurnal changes in boreholes
- Borehole head out of phase with velocity
 - Sampled the (large) 'unconnected or weakly-connected' unchannelized system
 - Flow coupling and 'passive' cavity opening
- Long-term trends in boreholes match trends in ice velocity



Conceptualizing the weakly connected system

- Channels control daily and event (multi-day) variations in velocity
 - Moulin hydraulic head indicates the presence of stable channels during the latter half of the summer
- Late season velocity patterns are likely due to changes occurring within the isolated or weakly-connected regions of the unchannelized system.







Why could weakly connected (inactive regions) of the bed be important? Early melt season Mid melt season

- Extensive and dynamic isolated and ephemerally connected systems under alpine glaciers (e.g., *Hodge*, 1979; *Murray* & Clarke, 1995; Gordon et al., 1998)
- Widespread velocities respond to the integrated basal traction over the entire bed (both connected and weakly-(un)connected regions) (Iken & Truffer, 1997)





Post melt season





pressure

ncreasing

Hoffman et al. (2016)

ncreasing connectivity

Modeling the weakly connected system

Weakly-connected system implemented as as discontinuous reservoir that can exchange water with the surrounding drainage system in each grid cell



Subglacial pressure and ice velocity are coupled through a prescribed basal friction law



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Modeled and observed subglacial pressures

- Model reproduces the basic features of the velocity observation. •
- Control simulation with prescribed, fixed weakly-connected subglacial • pressure fails to produce the observed seasonal evolution. Z





Why is the weakly-connected system important?

- Weakly-connected subglacial evolution previously unconsidered in subglacial hydrology models
- Necessary to explain seasonal hysteresis between moulin and ice velocity records.
- Can explain the observed winter slowdowns (*Hoffman et al.*, 2011; *Sole et al.*, 2013; *Tedstone et al.*, 2013; Tedstone et al., 2015).
- Modified parametrizations, including dynamic hydraulic conductivity better explain observations at multiple locations (Downs et al., 2018; Rada & Schoof, 2018)



Tedstone et al. (2015)

How does meltwater get to the bed?

At higher elevations, moulins tend to form in regions of compression – counter to prevailing assumptions.







Short perturbations may cause widespread basal connections

Lakes tend to drain in clusters and lake drainages are poorly predicted by water volume or surface characteristics, suggesting **ice flow coupling is likely important**.

a. Supraglacial Lake Formation

A supraglacial lake forms on the surface of the Greenland Ice Sheet when surface runoff fills a compressive basin.





b. A Precursor to Rapid Lake Drainage

Tension

Substantial melt-water is routed to the bed via a moulin, causing uplift and tension at the ice sheet surface.

Surface uplift

Basal uplift

Surface

c. Hydro-fracture Opening and Rapid Drainage A hydro-fracture opens through the lake basin, draining

water in the lake to the bed within a few hours.



Stevens et al. (2015)





Using ice speed to model potential subglacial drainage locations

Andrews (2015); Hoffman et al. (2018)



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Ice dynamics model + high resolution ice velocities

- Ice flow models are often forced with ice velocities to obtain basal traction (or slipperiness).
- Instead assimilate high temporal resolution ice velocities to constrain basal traction and ice stresses during periods of rapid velocity change.



Constraints on basal traction and surface stresses







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When might moulins form?



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Transients matter in the glacial system

- Transient stresses driven by ice velocity perturbations likely trigger cascading lake drainage and moulin formation.
- We need to estimate the 'sphere of influence' of lake drainages in order to characterize how englacial connections will evolve.

а

60

40

20

y (km)

6 June

Perturbation

Basal water flux (m³ s⁻¹)

Drain O





What do moulins look like and does their shape matter?

Moulin geometry dictates subglacial pressure and variability – changes in the their shape in space and time will alter the subglacial response to surface meltwater.





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Multiple non-linear processes alter moulin geometry





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Moulin shape matters for subglacial pressure Volume (×10⁴ m³)

Moulin shape can alter the equilibrium response time and modify the magnitude of diurnal variability even when the mean radius is the same.





0 1000 ∟

900

800

Moulin shape evolves diurnally

Moulin geometry can evolve significantly during diurnal fluctuations in melting due to changes in water level (creep closure) and melting (turbulent dissipation).









Still working out the kinks...

- Realistic ice and meltwater discharge characteristics, but potentially unrealistic geometries and water levels suggest missing model parameterizations (elastic deformation)
- Coupling to a subglacial model to more appropriately deal with changes in moulin discharge is in progress.





Understanding moulin geometry is a step toward a stochastic model



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The link between surface melt and ice dynamics is dominated by transient behavior



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