

Observation-Based Constraints On Atmospheric And Oceanic Cross-Equatorial Heat Transport

Norman G. Loeb¹, Hailan Wang², Anning Cheng², Seiji Kato¹,
John T. Fasullo³, Kuan-Man Xu¹, Richard P. Allan⁴

¹NASA Langley Research Center, Hampton, VA

²Science Systems and Applications, Inc., Hampton, VA

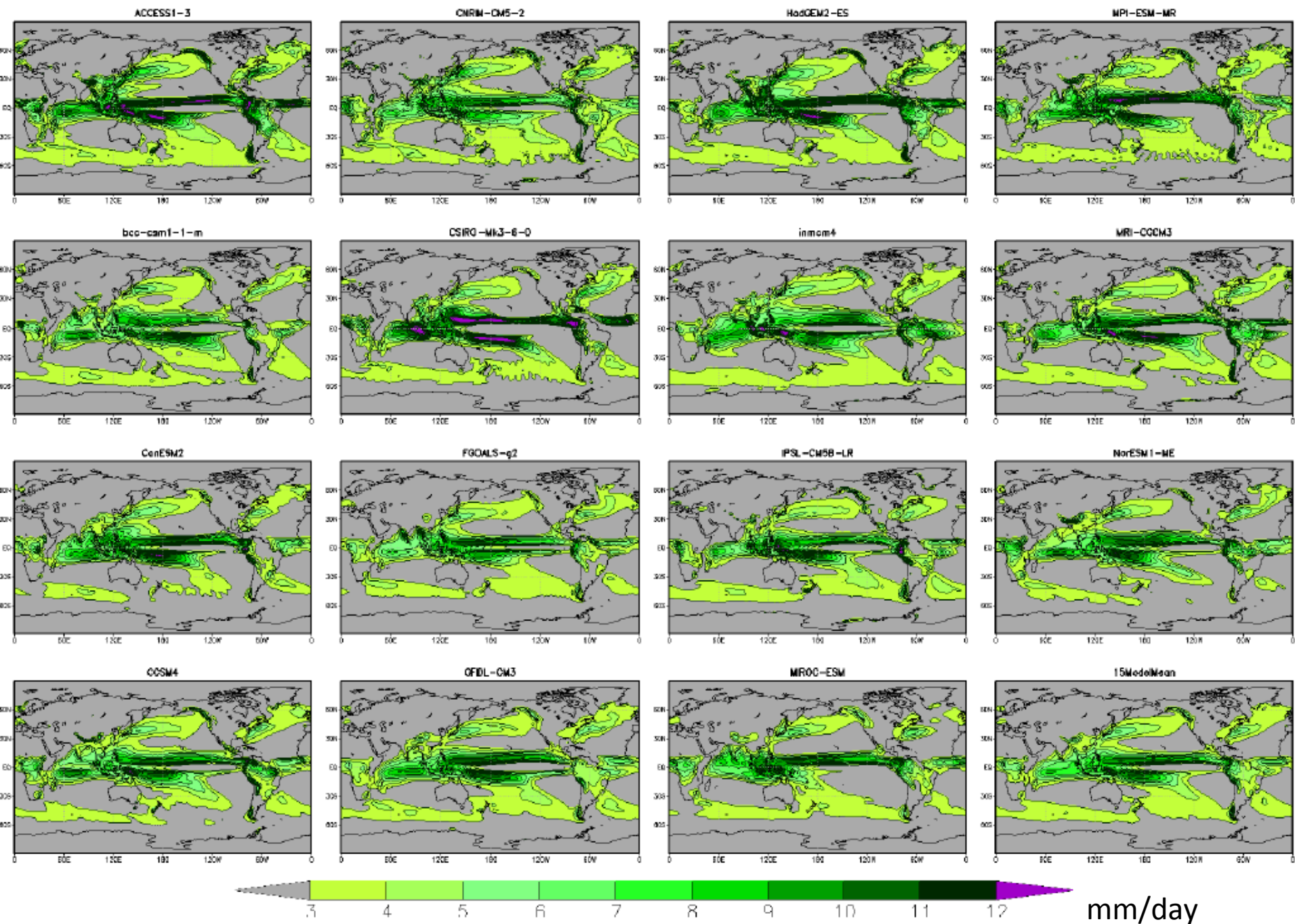
³National Center for Atmospheric Research, Boulder, CO

⁴Department of Meteorology, University of Reading, Reading, UK

Introduction

- Large-scale tropical circulation and precipitation are constrained by the regional distribution of energy.
- The hemispheric asymmetry in energy determines the cross-equatorial heat transport in the atmosphere and ocean.
- This in turn constrains the mean position of the ITCZ.
- ITCZ and associated precipitation is poorly represented in climate models, likely because they do not correctly represent the regional distribution of energy.

CMIP5 Historical Coupled Simulations(1980-2004Mean): Precip



Objective

- Use CERES EBAF (TOA & SFC) Ed 2.8 and ERA-Interim to determine the implied atmospheric and ocean cross-equatorial heat transports.
- Further decompose the implied cross-equatorial heat transport into radiative and non-radiative contributions.
- Evaluate how climate models (CMIP5) represent the cross-equatorial heat transport.

Observations

- CERES EBAF Ed2.8 (TOA and SFC).
- ERA-Interim total energy tendency and column-integrated divergence of total energy ($c_p T + gz + Lq + k$).
 - Version of ERA-Interim used obtained from NCAR: The climate data guide: ERA-Interim: Derived components.
 - In this version, a mass flux correction has been applied to the divergence terms.
- GPCP V2.2
- Time Period: January 2001-December 2012.

CMIP5 Models Considered

Model number	Model name	Country/model group	Resolution (Lon $\times$ Lat)	Rt-Fs
1	ACCESS1.0	Australia/ACCESS	1.875° $\times$ 1.25°	-0.38
2	ACCESS1.3		1.875° $\times$ 1.25°	-0.65
3	CCSM4		1.25° $\times$ 0.9375°	-0.35
4	CESM1-BGC		1.25° $\times$ 0.9375°	-0.34
5	CESM1-FASTCHEM		1.25° $\times$ 0.9375°	-0.35
6	CESM1-WACCM	Australia/CSIRO	2.5° $\times$ 1.89°	-0.24
7	CSIRO-Mk3.6.0		1.875° $\times$ 1.86°	0.33
8	CanESM2		2.8125° $\times$ 2.79°	0.19
9	GFDL-CM3		2.5° $\times$ 2.0°	-0.36
10	GFDL-ESM2G		2.5° $\times$ 2.01°	-0.55
11	GFDL-ESM2 M	US/GISS	2.5° $\times$ 2.01°	-0.56
12	GISS-E2-H		2.5° $\times$ 2.0°	-0.42
13	GISS-E2-H-CC		2.5° $\times$ 2.0°	-0.40
14	GISS-E2-R		2.5° $\times$ 2.0°	-0.39
15	GISS-E2-R-CC		2.5° $\times$ 2.0°	-0.39
16	HadCM3	UK/Met Office	3.75° $\times$ 2.5°	-0.29
17	HadGEM2-CC		1.875° $\times$ 1.25°	-0.48
18	HadGEM2-ES		1.875° $\times$ 1.25°	-0.46
19	IPSL-CM5A-LR	France/IPSL	3.75° $\times$ 1.89°	0.32
20	IPSL-CM5A-MR		2.5° $\times$ 1.27°	0.33
21	IPSL-CM5B-LR		3.75° $\times$ 1.89°	-0.59
22	MIROC4 h	Japan/MIROC	0.5625° $\times$ 0.56°	-0.50
23	MIROC5		1.40625° $\times$ 1.40°	0.24
24	MPI-ESM-LR	Germany/MPI	1.875° $\times$ 1.86°	0.09
25	MPI-ESM-MR		1.875° $\times$ 1.86°	0.21
26	MPI-ESM-P		1.875° $\times$ 1.86°	0.12
27	MRI-CGCM3	Japan/MRI	1.125° $\times$ 1.12°	-0.19
28	MRI-ESM1		1.125° $\times$ 1.12°	-0.20
29	bcc-csm1-1	China/BCC	2.8125° $\times$ 2.79°	-0.97
30	bcc-csm1-1-m		1.125° $\times$ 1.12°	-0.99

Atmospheric & Surface Energy Budgets from CERES and Reanalysis

$$\frac{\partial A_E}{\partial t} = R_T - F_S - \nabla \cdot F_A \quad (1)$$

$$F_S = R_S + LE + S \quad (2)$$

$$F_A = \frac{1}{g} \int_0^{p_s} (h + k) \vec{u} dp$$

$$A_E = c_p T + gz + Lq + k = h + k$$

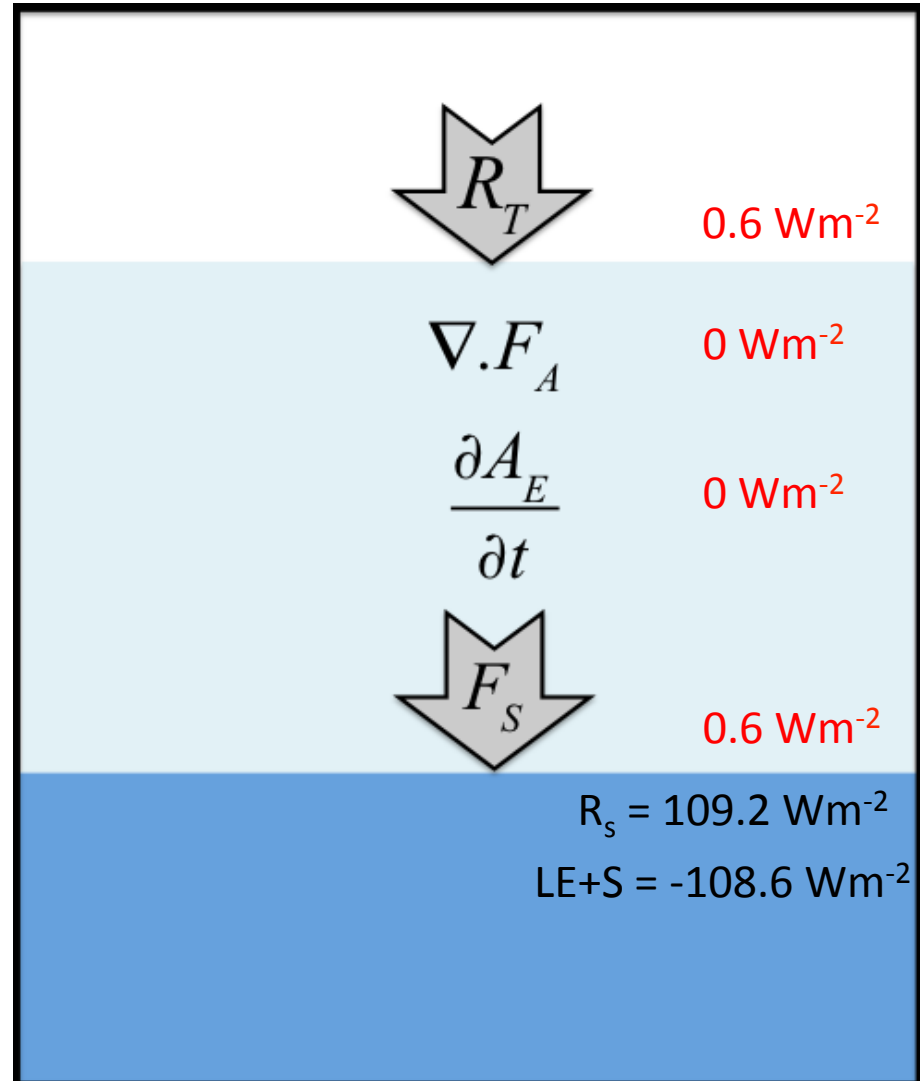
$$\frac{\partial A_E}{\partial t} \text{ \& } \nabla \cdot F_A \Rightarrow \text{ERA-Interim}$$

$$R_T \text{ \& } R_S \Rightarrow \text{CERES EBAF Ed2.8}$$

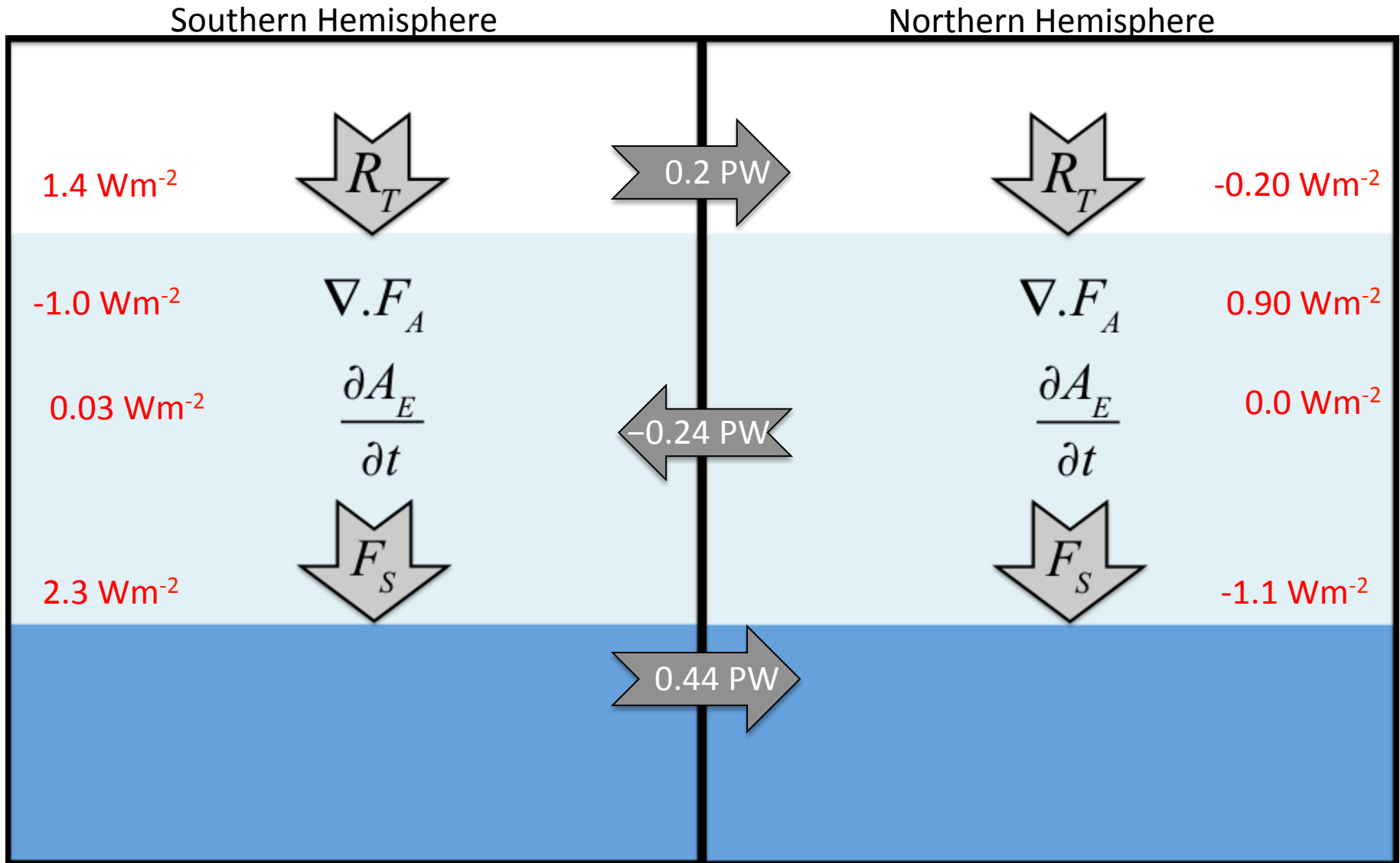
$$F_S \text{ \& } (LE + S) \Rightarrow \text{Residual Terms}$$

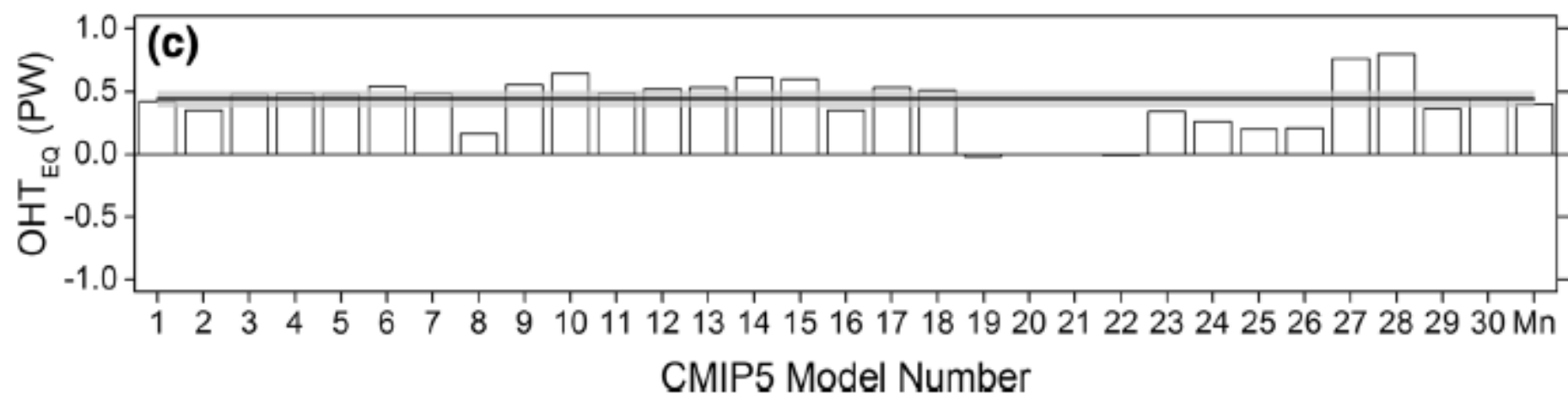
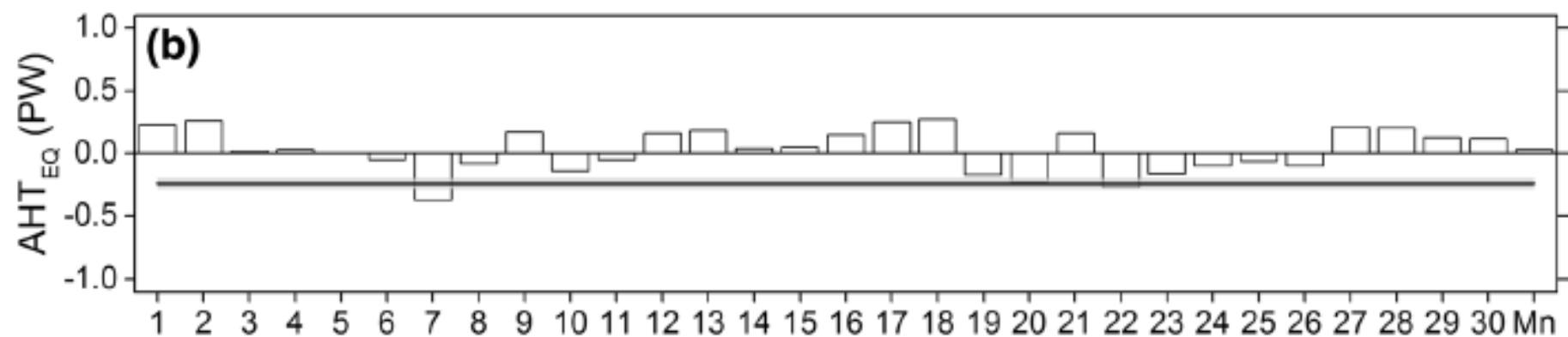
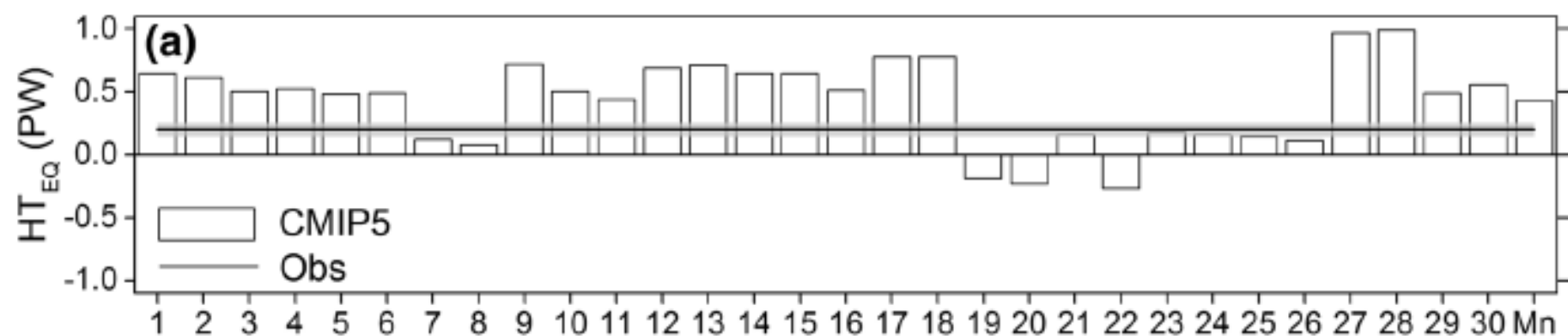
in (1) & (2)

Global Mean



Implied Cross-Eq. Heat Transports in Atmos. & Ocean from Energetic Constraints





Decomposition of Cross-Equatorial Heat Transport into Radiative and Combined Latent and Sensible Heat Flux Components

Atmosphere

$$AHT_{EQ} = \frac{1}{2} \left(\Delta R_T - \Delta F_S - \Delta \frac{\partial A_E}{\partial t} \right)$$

Ocean

$$OHT_{EQ} = \frac{1}{2} \left(\Delta F_S - \Delta \frac{\partial O_E}{\partial t} \right)$$

Δ denotes the SH minus NH difference.

$$F_S = R_S + LE + S$$

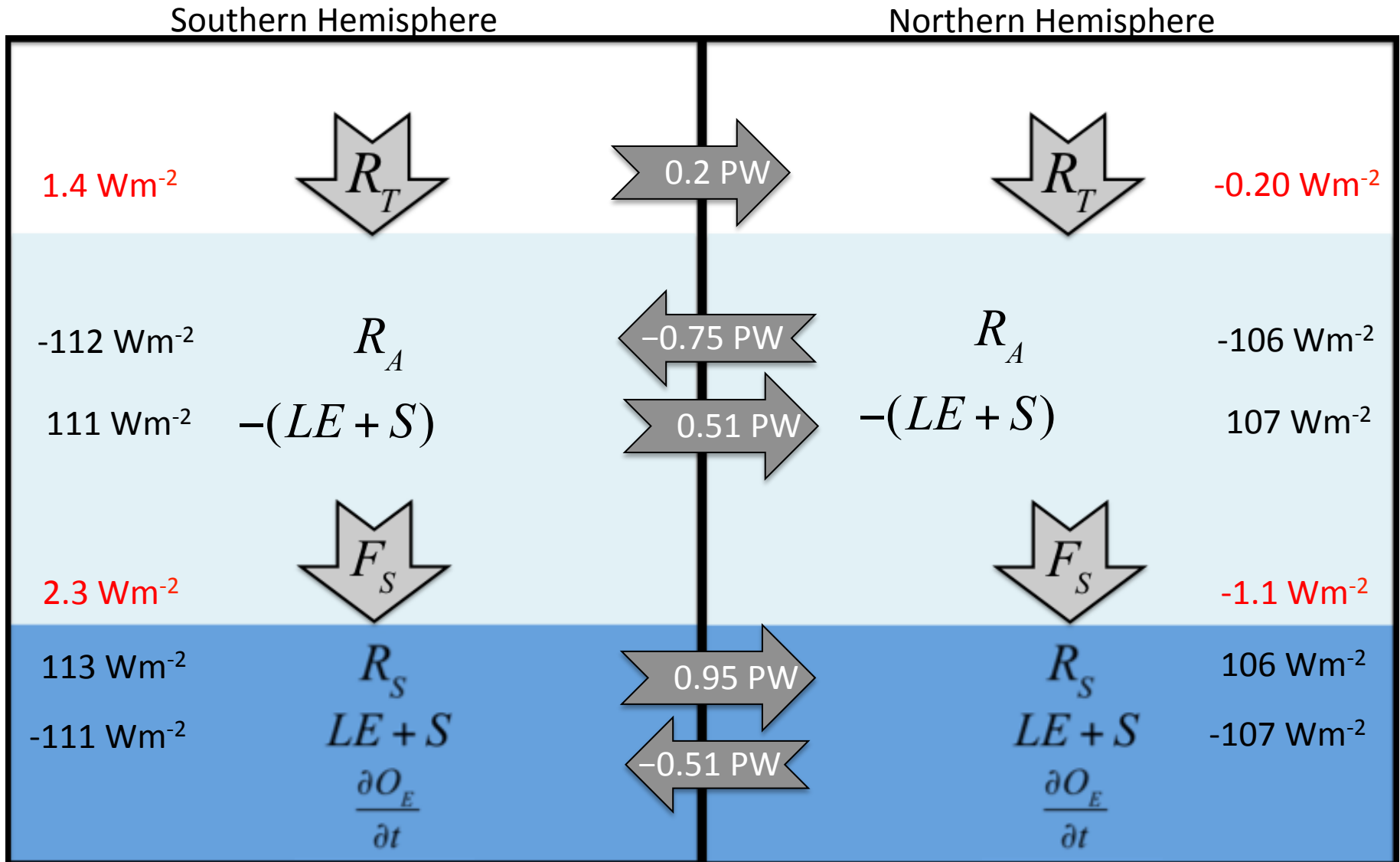
$$AHT_{EQ} = \frac{1}{2} \left(\Delta R_A + \Delta Q_A - \Delta \frac{\partial A_E}{\partial t} \right) \quad OHT_{EQ} = \frac{1}{2} \left(\Delta R_S + \Delta Q_S - \Delta \frac{\partial O_E}{\partial t} \right)$$

$$R_A = R_T - R_S$$

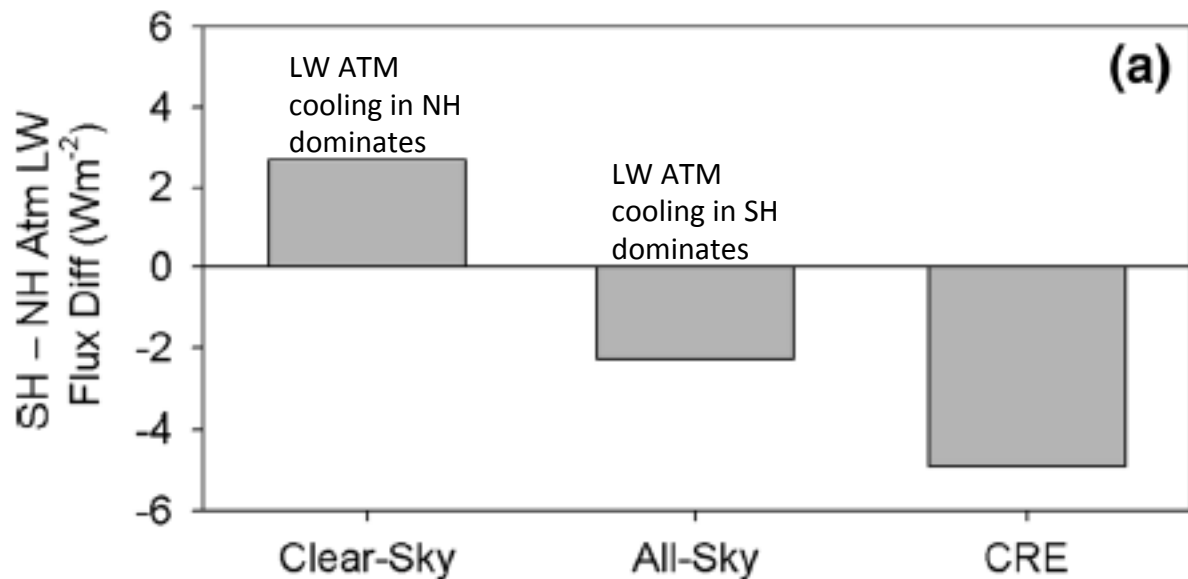
$$Q_S = (LE + S)$$

$$Q_A = -Q_S$$

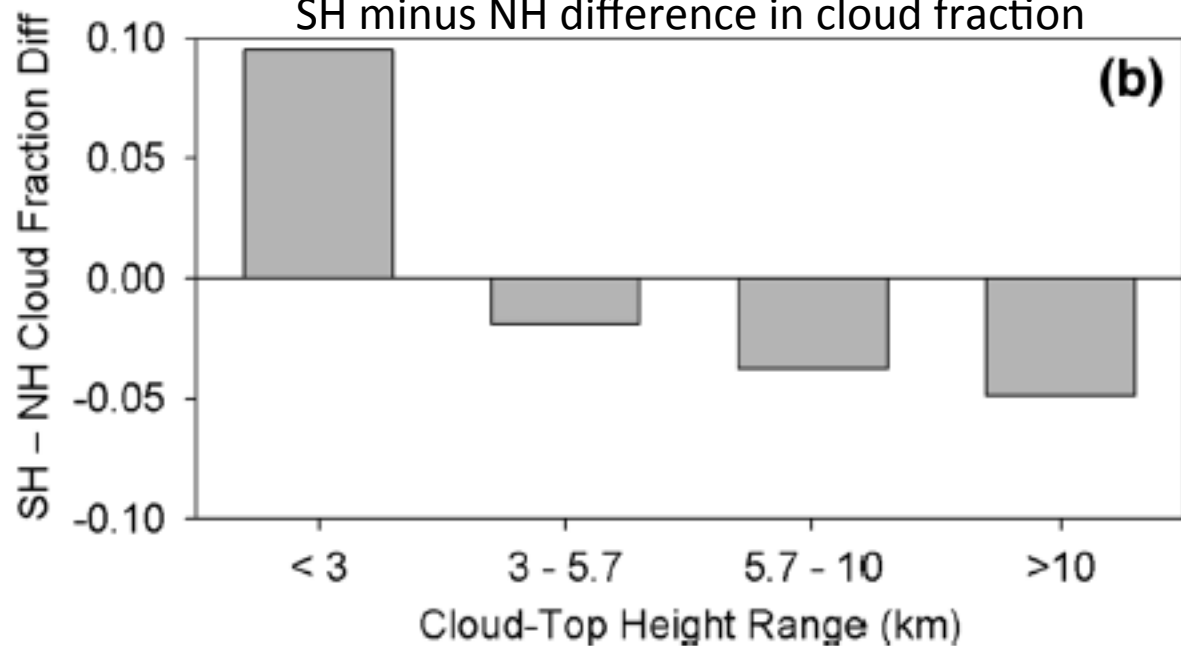
Implied Cross-Eq. Heat Transports in Atmos. & Ocean from Energetic Constraints

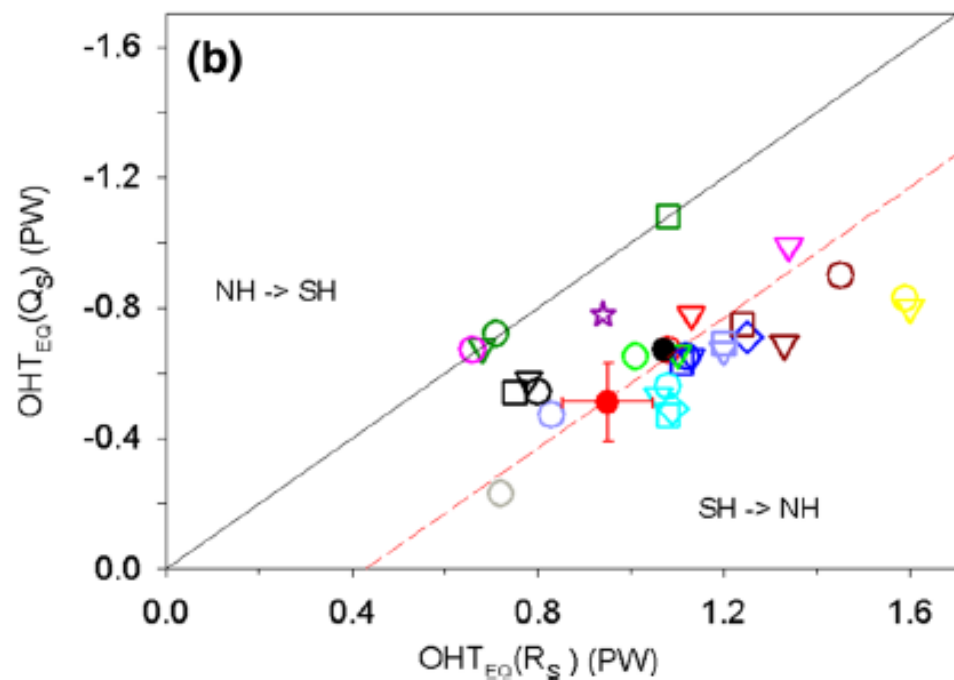
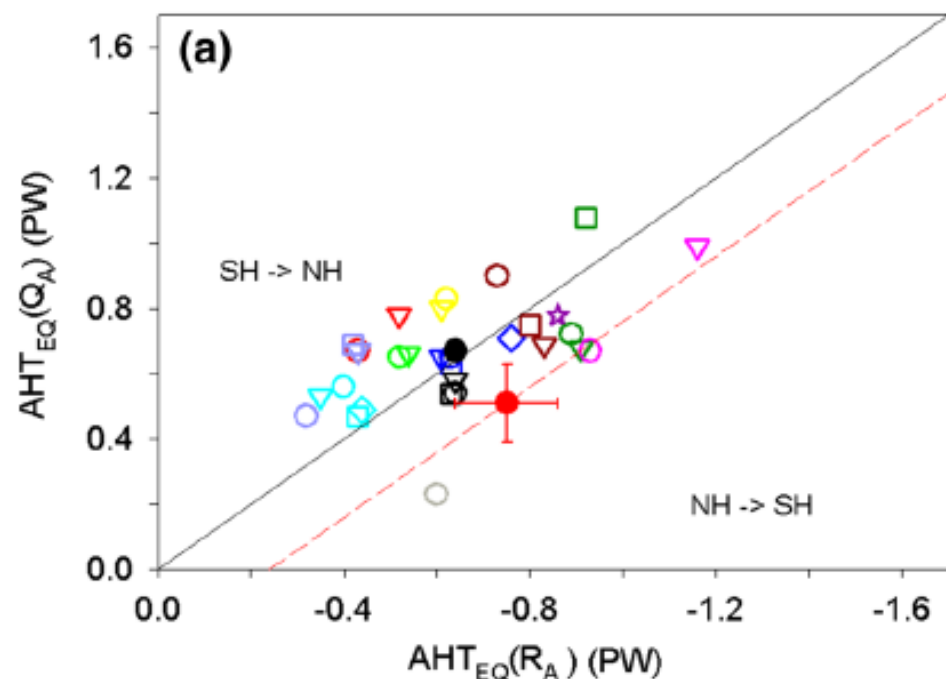


SH minus NH difference in atmospheric LW Flux



SH minus NH difference in cloud fraction





- ACCESS1.0
- ▽ ACCESS1.3
- CCSM4
- ▽ CESM1-BGC
- CESM1-FASTCHEM
- ◇ CESM1-WACCM
- CSIRO-Mk3.6.0
- ☆ CanESM2
- GFDL-CM3
- ▽ GFDL-ESM2G
- GFDL-ESM2M
- GISS-E2-H
- ▽ GISS-E2-H-CC
- GISS-E2-R
- ◇ GISS-E2-R-CC
- HadCM3
- ▽ HadGEM2-CC
- HadGEM2-ES
- IPSL-CM5A-LR
- ▽ IPSL-CM5A-MR
- IPSL-CM5B-LR
- MIROC4h
- ▽ MIROC5
- MPI-ESM-LR
- ▽ MPI-ESM-MR
- MPI-ESM-P
- MRI-CGCM3
- ▽ MRI-ESM1
- bcc-csm1-1
- ▽ bcc-csm1-1-m
- CMIP5 Multi-model Mean
- OBS (CERES/ERA-Interim/GPCP)

CERES/ERA

CMIP5

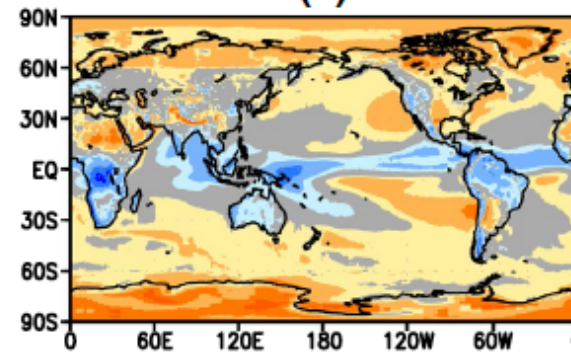
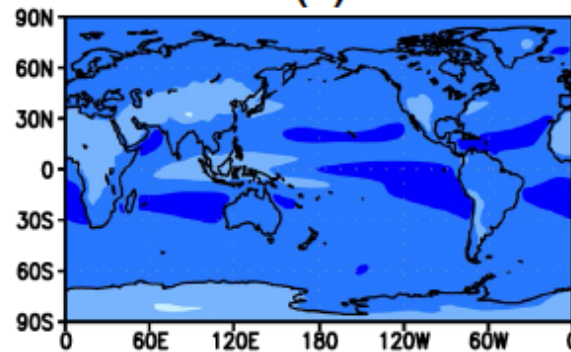
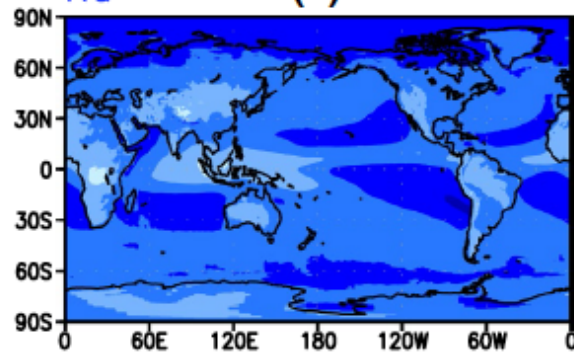
CMIP5-CERES/ERA

 R_a

(a)

(b)

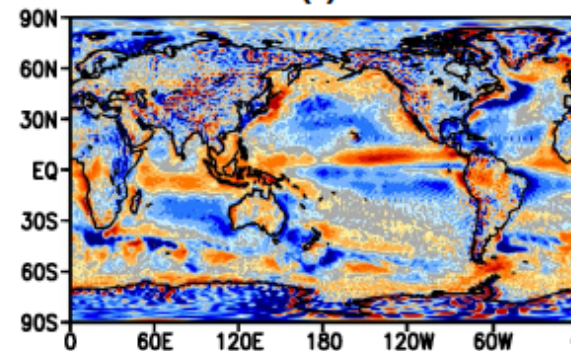
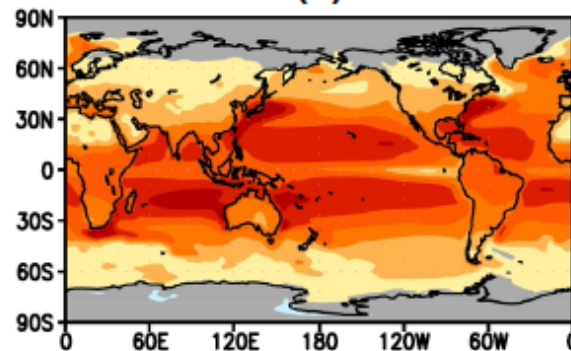
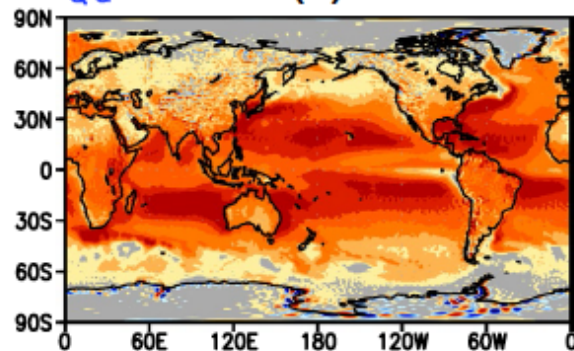
(c)

 Q_a

(d)

(e)

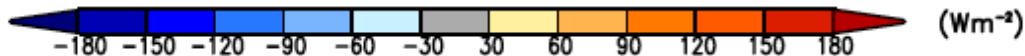
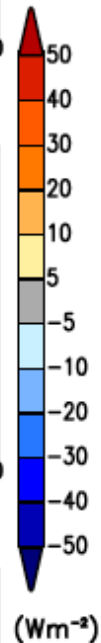
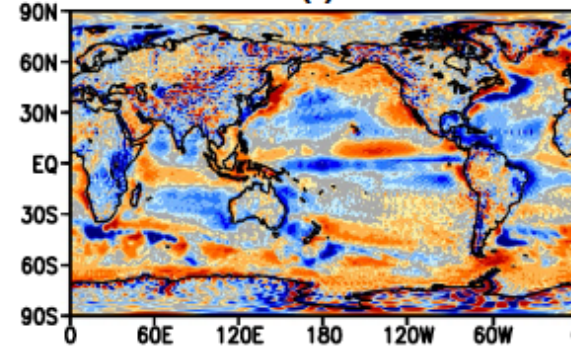
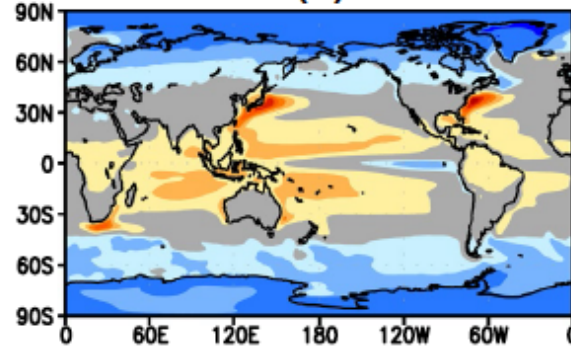
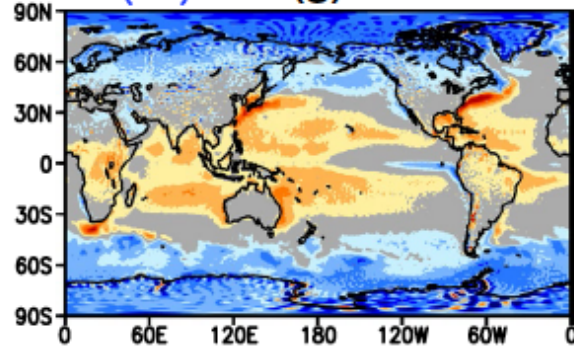
(f)

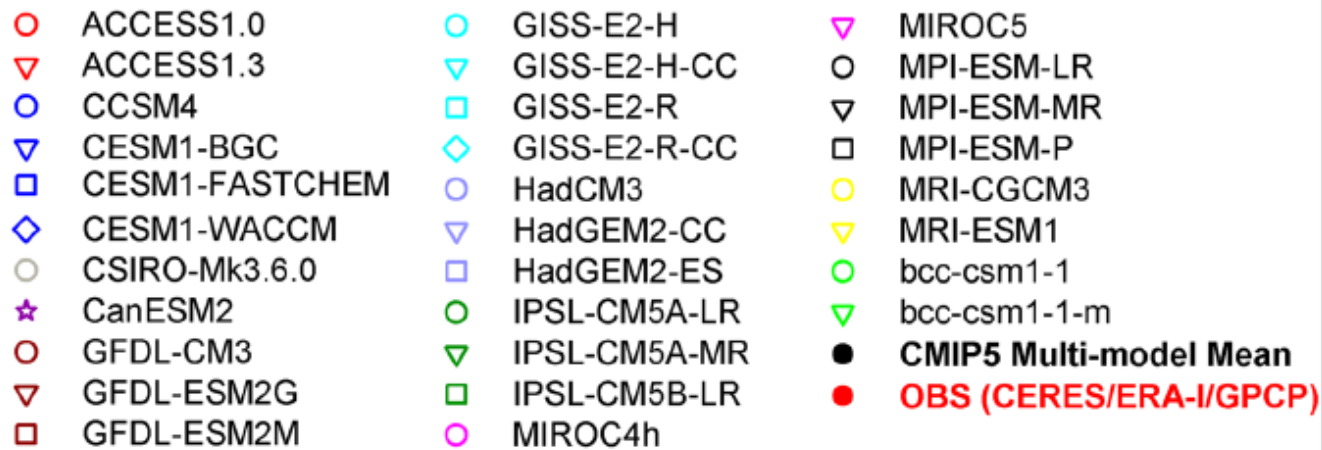
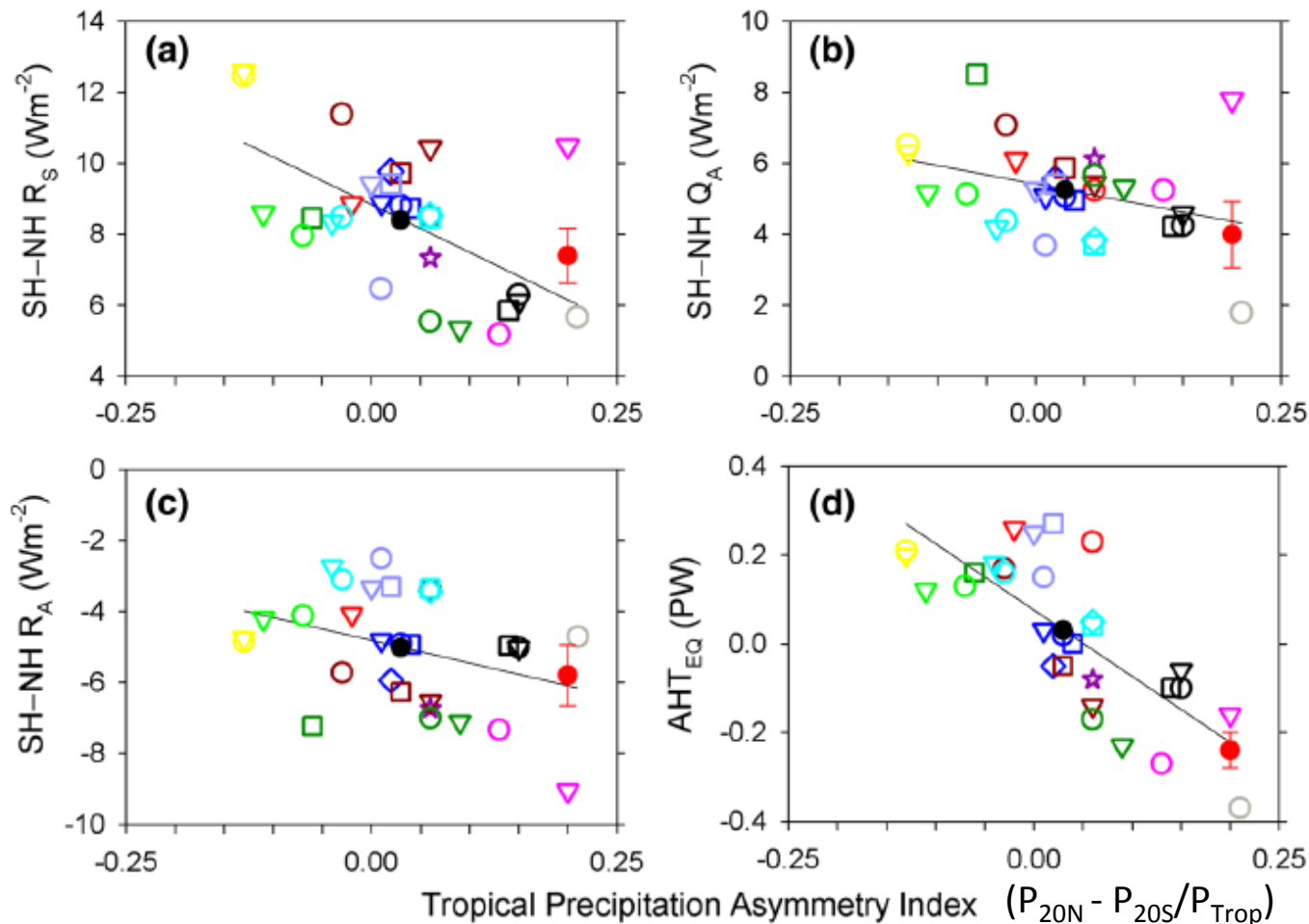
 $Div(F_a)$

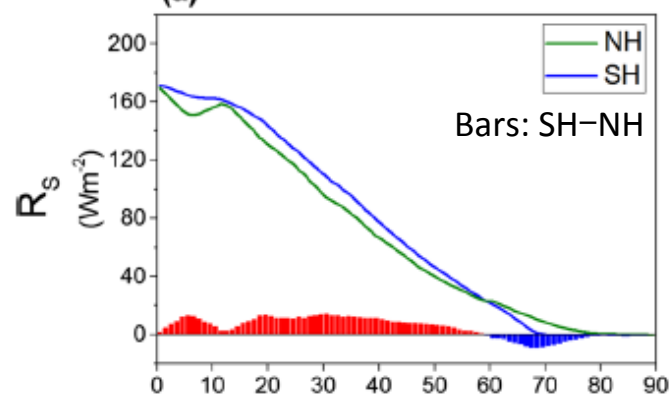
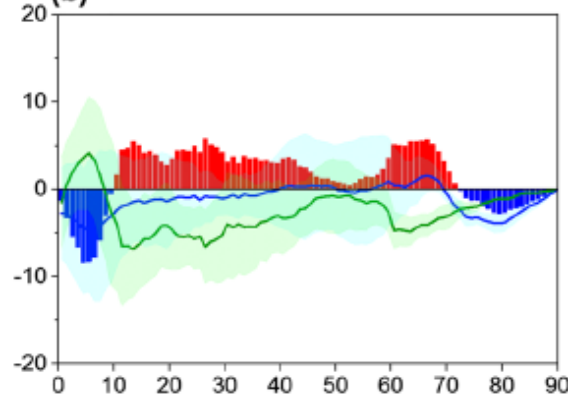
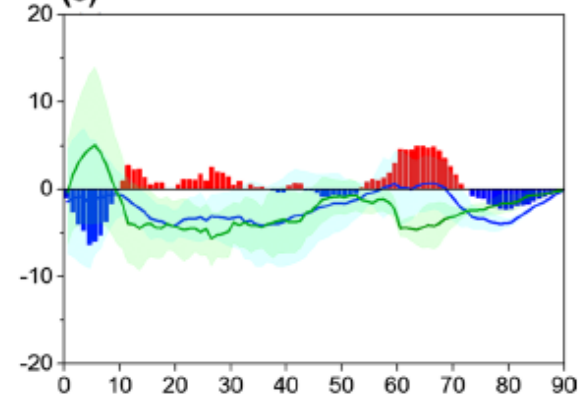
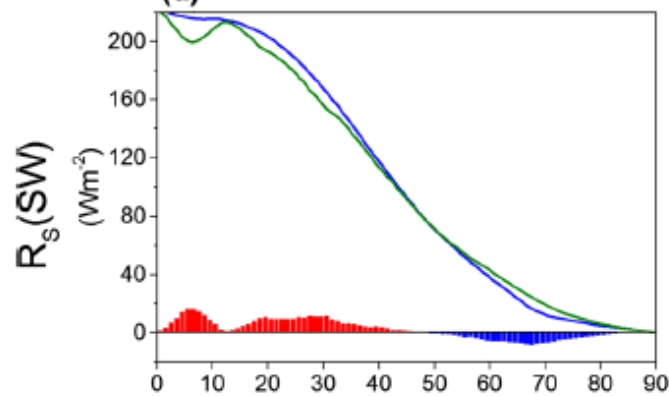
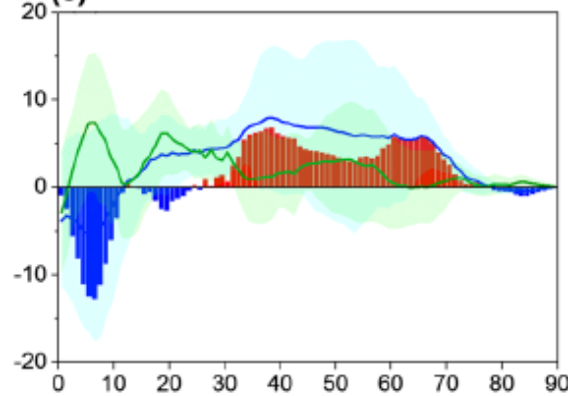
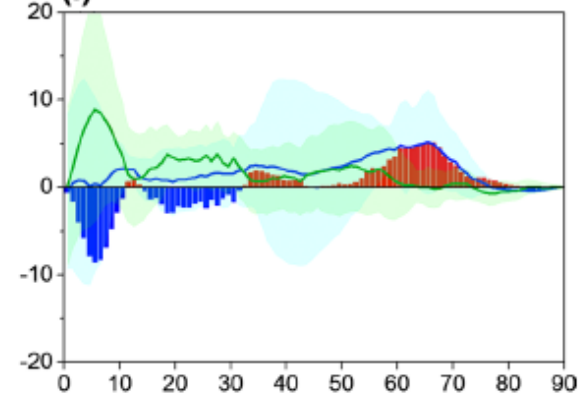
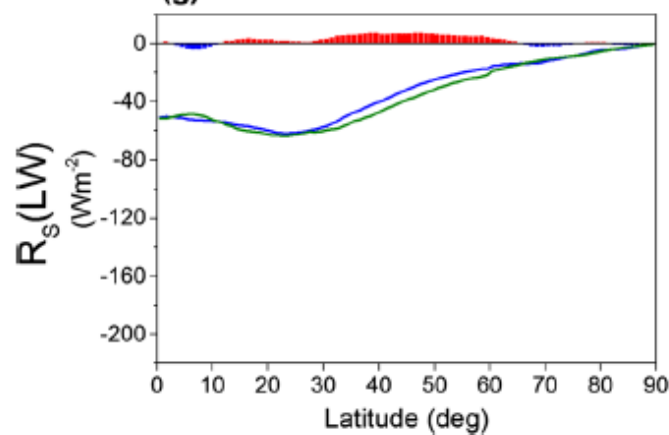
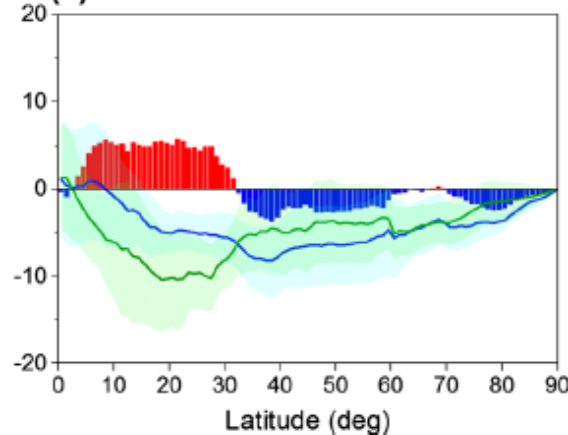
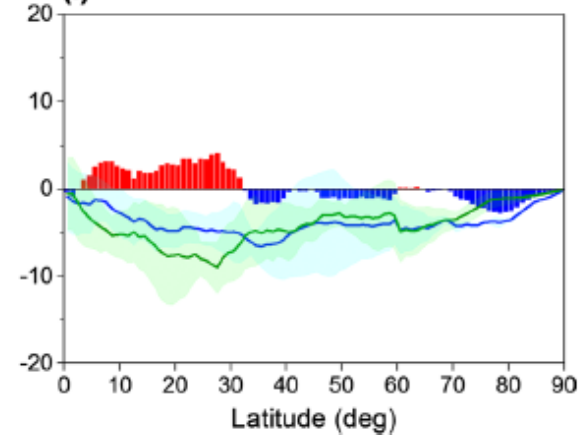
(g)

(h)

(i)





(a) Observations**(b) Model Bias (TPA ≤ 0)****(c) Model Bias (TPA > 0)****(d)****(e)****(f)****(g)****(h)****(i)**

Conclusions

- The large-scale circulation in the tropics and position of the ITCZ are intricately linked with the large-scale distribution of the energy budget.
- CERES EBAF-TOA and SFC combined with ERA-I atmospheric total energy divergence enable a decomposition of cross-equatorial heat transport into radiative and combined latent and sensible heat flux components.
- This decomposition provides a powerful new observational constraint on large-scale energy budget that needs to be satisfied in order to make progress on double ITCZ problem.

Conclusions

- SH has a larger cloud fraction and a greater fraction of low clouds, while the NH has more high clouds. In addition, NH has a higher surface albedo, greater abundance of absorbing aerosols and precipitable water.
 - ⇒ LW radiative cooling is more pronounced in the SH than the NH and SW radiative heating is greater in the NH.
 - ⇒ Net atmospheric radiative effect is more cooling in the SH relative to NH, which implies a NH to SH cross-eq heat transport.
 - ⇒ Surface-to-atmosphere combined latent and sensible heat transport is greater in SH than NH, which compensates somewhat for radiatively driven cross-eq heat transport.

Conclusions

- CMIP5 models that overestimate tropical precipitation in the SH:
 - overestimate net downward surface radiation in SH vs NH
 - overestimate combined latent and sensible heat flux in SH vs NH
 - underestimate atmospheric radiative cooling in the SH vs NH
- ⇒ Excessive heating of the SH atmosphere and anomalous SH to NH cross-equatorial heat transport
- ⇒ Ascending branch of Hadley circulation lies too far to the south (necessary to move excess heat from SH to NH).
- ⇒ Too much SH tropical precipitation.