Mechanical twinning of monazite expels radiogenic Pb 1 2 D. Fougerouse^{1,2*}, S.M. Reddy^{1,2}, A.-M. Seydoux-Guillaume^{3,4}, C.L. Kirkland¹⁺, T.M. Erickson^{1,5}, D.W. Saxey², W.D.A. Rickard², D. Jacob⁶, H. Leroux⁶, C. Clark¹ 3 4 5 ¹School of Earth and Planetary Sciences, Centre for Exploration Targeting – Curtin Node⁺, 6 Curtin University, Perth, WA 6845, Australia 7 ²Geoscience Atom Probe, John de Laeter Centre, Curtin University, Perth, WA 6845, Australia 8 ³CNRS, Université de Lyon, UCBL, ENSL, LGL-TPE, 69622 Villeurbanne, France 9 ⁴Univ Lyon, UJM-Saint-Etienne, F-42023 Saint-Etienne, France 10 ⁵Jacobs – JETS, Astromaterials Research and Exploration Science division, NASA Johnson 11 Space Center, Houston, TX, 77058, USA ⁶Univ Lille, CNRS, INRAE, Centrale Lille, UMR 8207 - UMET - Unité Matériaux et 12 Transformations, F-59000 Lille, France. 13

14 ABSTRACT

Mechanical twins form by the simple shear of the crystal lattice during deformation In order 15 16 to test the potential of narrow twins in monazite to record the timing of their formation, a 17 ~1,700 Ma monazite grain deformed at ~980 Ma was investigated by electron backscattered diffraction (EBSD), transmission electron microscopy (TEM) and atom probe tomography 18 19 (APT). APT ²⁰⁸Pb/²³²Th ages indicate that the twin was entirely reset by radiogenic Pb-loss 20 during its formation at conditions far below the monazite closure temperature. The results are 21 consistent with a model where Pb is liberated during rupture of REE-O bonds in the large 22 REEO₉ polyhedra during twinning. Liberated Pb is likely to migrate along fast diffusion pathways such as crystal defects. The combination of a quantitative microstructural 23 24 investigation and nanogeochronology provides a new approach for understanding the history 25 of accessory phases.

26 INTRODUCTION

The radiogenic decay of U and Th into different isotopes of Pb provides a series of widely used geochronometers. Critical to this application is an understanding of the mobility of radiogenic Pb in minerals and how this underpins our ability to use and develop U-Th-Pb 30 geochronology. Traditional models for Pb mobility assume temperature-dependent volume diffusion of Pb through the crystal lattice of the mineral. More recently, a number of studies 31 have reported a relationship between trace element mobility and deformation (Fougerouse et 32 33 al., 2019; Kirkland et al., 2018; Moser et al., 2009; Piazolo et al., 2012; Reddy et al., 2016; Timms 34 et al., 2011). Pb mobility has been suggested to occur in high closure temperature minerals 35 (zircon and monazite) during crystal-plastic deformation (Moser et al., 2009) and 36 deformation-induced grain boundary migration (Erickson et al., 2015), with radiogenic accumulation continuing after deformation. 37

Deformation twinning is a common, narrow ($\leq 3 \mu m$) microstructural feature in deformed 38 39 minerals and forms by the simple shear of the crystal lattice and a highly-ordered 40 displacement of atoms. Twinning is sensitive to temperature during deformation and strain rate, and typically is more dominant than slip at lower temperatures or at higher strain rates 41 (Christian and Mahajan, 1995). Mechanical twinning is a deformation microstructure that has 42 43 found use as a geothermometer in calcite (Ferrill et al., 2004), as a tool to determine 44 compressive stress direction (Jamison and Spang, 1976) and has been considered a diagnostic microstructure for impact-related deformation in quartz, zircon and monazite (Erickson et al., 45 46 2016b; Goltrant et al., 1992; Moser et al., 2011; Timms et al., 2012).

In this study, we integrate electron backscattered diffraction (EBSD), transmission electron microscopy (TEM) and atom probe tomography (APT) analyses to investigate Pb mobility associated with micrometre-scale deformation twins formed within tectonically deformed monazite. We provide a framework for the analysis of mineral twins in monazite that could be further utilised in the analysis of nanoscale microstructural features in other accessory phases.

53 SAMPLES AND METHODOLOGY

In this contribution, we re-investigate a monazite grain from the granulite-facies rocks of the Sandmata Complex, Rajasthan, India (Fig.1), which was previously studied in detail by Erickson et al. (2015) and Fougerouse et al. (2018). Two high-temperature metamorphic events are recognised in this region, a first event at ~1720 Ma (M1sc; ~7–10 kbar and 800–900 °C), which resulted in growth of the original monazite, and a second, fluid-absent event, at ~1000 Ma (M2sc ; ~5–7 kbar and 600–750 °C; (Buick et al., 2006)). Electron Backscatter Diffraction 60 (EBSD) and SIMS U/Pb analyses reveal partial Pb-loss in deformed domains of monazite and 61 growth of neoblasts in localised deformed domains at 970 \pm 14 Ma (2 σ , n = 6, MSWD = 1.3; 62 Erickson et al., 2015). Due to a spot size larger that the mechanical twins, SIMS analyses of 63 mixed twin and host domains yield imprecise U-Pb discordant data, implying partial Pb loss 64 in the domain analysed. These mixed domains could not be used to successfully constrain the 65 timing of monazite deformation (Erickson et al., 2015).

EBSD analyses were used to investigate crystal lattice orientation variations and were
conducted on a Tescan Mira3 scanning electron microscope equipped with a Nordlys Nano
high resolution detector at Curtin University. Details of EBSD analyses are given in Erickson
et al. (2015).

One atom probe specimen was prepared from a (001) twin (specimen 1) and three specimens from a neoblast (specimen 2 to 4) by focused ion beam – scanning electron microscopy (FIB-SEM; Fig. 1C). A Tescan Lyra3 Ga⁺ FIB-SEM was used to produce needle specimens and precisely select the specimen location (Rickard et al., 2020). A TEM foil was prepared with the same instrument across the boundary between the twin and the monazite host (Fig. 1C) and further thinned using a FEI Helios 600i FIB-SEM (MANUTECH USD, Saint-Etienne, France).

The TEM foil was studied using a FEI TITAN Themis 300 (University of Lille, France)to obtain
high resolution high angle annular dark field (HR-HAADF) images and energy dispersive xray spectroscopy (EDS) with a Super-X windowless, 4 quadrant SDD detector in scanning
transmission electron microscope (STEM) mode.

80 APT analyses were performed on at Cameca LEAP 4000X HR at the Geoscience Atom Probe 81 facility at Curtin University (Reddy et al., 2020). The instrument was operated in laser pulsed mode with a UV laser (λ = 355 nm), laser pulse energy of 100 pJ, base temperature of 50 K and 82 automated detection rate of 0.01 atoms per pulse. Four datasets were collected with ~8, 26, 27 83 and 40 million atoms, respectively. The ²⁰⁸Pb/²³²Th age was measured using the ²⁰⁸Pb⁺⁺ and 84 ²³²ThO⁺⁺ signal and corrected following the protocol defined by Fougerouse et al. (2020). 85 86 Counts were quantified by subtracting the estimated background counts from each peak. The 87 local background level was measured from the counts in a nearby 'peak-free' range and 88 normalised to the width of the range used for quantification (0.1 Da). The isotopic ratio 89 uncertainties are derived from the uncertainty of each species used for calculating the ratio and reported at 95% confidence (2 sigma). The linear correlation (y = $0.381(\pm 0.048)x - 254(\pm 26)$) between the ²⁰⁸Pb/²³²Th fractionation coefficient (y) and the M/ Δ M₁₀ peak parameter (x), where M is the position of the O₂⁺ peak and Δ M₁₀ its full-width-at-tenth-maximum, was used to correct for molecular fractionation and calculate ²⁰⁸Pb/²³²Th ages (Fougerouse et al., 2020). The error propagation of the prediction is derived from the uncertainty of the M/ Δ M₁₀ value, the uncertainty of the measured ²⁰⁸Pb/²³²ThO ratio and the 95% prediction band.

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97 **RESULTS**

The EBSD data show that the twin is crystallographically equivalent to a 180° rotation about (100>, consistent with apparent shearing in the [100] direction on the (001) plane (e.g. Erickson et al., 2016). The observed twin dimensions in the plane of the thin section are approximately 80 μ m long and 2.5 μ m wide. Both the twin domain and the host monazite are relatively undeformed in the region sampled for APT with a maximum of 3.5° cumulative misorientation.

104 The atom probe results of specimen 1 (twin) reveal that the Ca distribution is heterogeneous 105 within two distinct compositional domains separated by the twin boundary (Fig. 2). A proportion of the Ca within the host is clustered along with Si and Pb (~10 at.% Ca, ~1 at.% Si 106 107 and ~0.7 at.% Pb) and is consistent with nanoscale inclusions of apatite in other APT studies 108 of monazite (Fougerouse et al., 2018), whereas the twin appears homogeneous. The trace 109 element compositions of the host and the twin are dissimilar with higher Ca and Pb in the host (1.1 vs 0.8 at.% Ca and 0.09 vs 0.05 at.% Pb, respectively) and lower Si concentration (0.04 110 vs 0.1 at.% Si, respectively). The twin boundary separating these two domains is enriched in 111 Ca (2.3 at.%), Si (0.3 at.%), a weak enrichment in Pb (only visible in 3D atom maps; Fig. 2) and 112 113 depleted in REEs and P.

114 The APT ²⁰⁸Pb⁺⁺/²³²ThO⁺⁺ ratio of the host, including the cluster is 0.1789 ± 0.0058 whereas the 115 twin domain ²⁰⁸Pb⁺⁺/²³²ThO⁺⁺ ratio is 0.0965 ± 0.0057 (2 σ ; Fig. 2). The peak parameter M/ Δ M₁₀ 116 for this specimen is 533.8 ± 2.7. Fractionation corrected APT ²⁰⁸Pb/²³²Th ages for the host 117 domain of specimen 1 is 1,698 ± 293 Ma (2 σ), whereas the twin domain of specimen 1 yields 118 an age of 934 ± 193 Ma. The atom probe results from specimen 2, 3 and 4 from a single neoblast are consistent across all specimens with homogenous distribution of major and trace elements (Fig. 2 & Fig. DR1). In the neoblasts, the APT ${}^{208}Pb^{++}/{}^{232}ThO^{++}$ ratio is 0.1074 ± 0.0048 for specimen 2, 0.1117 ± 0.0056 for specimen 3 and 0.1039 ± 0.0048 for specimen 4 (Fig. 2). The peak parameter M/ Δ M10 for Specimen 2 is 530.2 ± 1.7, 535.8 ± 1.8 for specimen 3 and 532.5 ± 1.8 for specimen 4, which yield corrected ages of 1,008 ± 198 Ma, 1,093 ± 204 Ma, and 994 ± 186 Ma, respectively (Fig. 2). The ages of the twinned domain and neoblast are within uncertainty of each other and have a weighted mean average of 1,005 ± 94 Ma (2 σ).

A HR-HAADF image across the twin boundary indicates the absence of an amorphous film and it is semi-coherent (Fig. 3B; (Ranganathan, 1966)). The STEM-EDS results are consistent with the APT findings, with enrichment in Ca and depletions in LREEs and P close to the boundary (few nm only; Fig. DR2).

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131 AGE RESETTING AND Pb MOBILITY DURING TWIN FORMATION

132 The APT ²⁰⁸Pb/²³²Th ages are in good agreement with previously published SIMS and electron 133 microprobe data for the metamorphic history of the region (Buick et al., 2006; Erickson et al., 134 2015). The age of the monazite host corresponds to the timing of granulite facies 135 metamorphism which affected the region at $1,720 \pm 10$ Ma (Buick et al., 2006). The $1,005 \pm 94$ Ma weighted mean age obtained for recrystallised neoblasts and twinned domain in specimen 136 137 1 is consistent (within uncertainty) with the 970 ± 14 Ma amphibolite facies metamorphism 138 and regional deformation event (Buick et al., 2006; Erickson et al., 2015). Therefore, all radiogenic and initial Pb was expelled from the twinned domain of the monazite crystal 139 140 during the deformation event (Fig. 4).

141 Twin nucleation and growth have been extensively explored in the material sciences literature 142 (Beyerlein et al., 2014; Christian and Mahajan, 1995). Mechanical twins are formed by the 143 simple shear of the crystal lattice and the continuous movement of partial dislocations, 144 however no studies have previously shown a relationship between twin formation and the 145 resetting of the Th-Pb chronometer.

146 Three possible mechanisms can explain the resetting of the twinned domain. These include147 (1) dynamic recrystallization during metamorphism, (2) fluid alteration replacement, and (3)

Pb liberation during crystal shearing. The studied twin is weakly deformed (up to 3.5° of 148 cumulative misorientation), is a single 80 µm long domain and lacks triple junctions. This 149 observation suggests that the twin did not recrystallise after formation. We postulate that 150 151 fluid-assisted recrystallization (i.e. pressure solution) did not operate during deformation of 152 the Sandmata monazite based on the lack of a free, interconnected fluid phase in the lower crust (Yardley and Valley, 1997), coupled with textural evidence including the isolation of 153 strain-free neoblasts within the grain interior and the absence of resetting along the exterior 154 of the monazite (Erickson et al., 2015; Erickson et al., 2016a). In addition, high-resolution TEM 155 data indicate that the host and the twin have a semi coherent boundary, unavailable to fluid 156 infiltration (Fig. 3). The role of small amount of fluids (H_2O , CO_2) at the grain boundary is 157 158 difficult to assess but may have facilitated the reactive transports of ions to/from the crystal surface. Therefore, we favour the third mechanism. 159

160 In monazite, Ca²⁺, Th⁴⁺ and U⁴⁺ reside within the REE sites (Ni et al., 1995) and radiogenic Pb²⁺ is repositioned in the monazite crystal structure on the REE sites after alpha recoil during self-161 162 annealing (Seydoux-Guillaume et al., 2018; Tang et al., 2020). The monazite crystal structure is composed of an arrangement of small PO₄ tetrahedra (P-O bond lengths = 0.1524 to 0.1540 163 nm) and larger REEO₉ polyhedra (REE-O bond lengths = 0.250 to 0.277 nm). These crystal 164 165 lattice parameters suggests that the PO4 tetrahedra are more rigid compared to the REEO9 166 polyhedra and that REE-O bonds are more likely to break than P-O bonds during twinning 167 (Hay and Marshall, 2003). This preferential point of rupture would also affect radiogenic Pb-O and Ca-O bonds as they share the same sites, and potentially release Pb and Ca during this 168 169 process (Fig. 4). The prediction of this model is that the liberation of Ca and Pb would result 170 in a Pb-Ca-free twinned domain and the accumulation of Ca and Pb within the twin boundary. 171 The nanogeochronology data indicate that all of the radiogenic Pb was mobilized during twin 172 formation, while Ca composition changed by ~25%, from 1.1 at.% in the host to 0.8 at.% in the 173 twinned domain. Ca is also highly enriched in the twin boundary (up to 2.3 at. %), whereas 174 Pb is weakly concentrated in the same domain indicating that different mechanisms may affect Pb and Ca mobility after crystal shearing. In nine-fold coordination, the ionic radius of 175 Pb²⁺ (1.35 Å) is larger than Ca²⁺ (1.18 Å; (Shannon, 1976)). During twinning, dislocations are 176 necessary at the point of shearing (Cottrell and Bilby, 1951). Dislocations can trap impurities 177 in the distorted crystal lattice surrounding the dislocation (Cottrell and Bilby, 1949), or in their 178

cores (Johnston and Gilman, 1959). Large ions such as Pb²⁺ are likely hosted in the core of the
dislocation as opposed to smaller ions such as Ca²⁺ more suited to the locally distorted lattice.
Ions in dislocation cores can diffuse faster along the linear defect (Love, 1964), allowing for an
efficient expulsion of Pb compared to Ca. A portion of the Ca in the Cottrell atmosphere may
be reincorporated in the twinned domain as the shearing of the crystal progresses or upon
subsequent thermal annealing.

No experimental data are available to constrain the temperature dependency of Pb diffusion along crystal defects, and the influence of temperature and strain rate on Pb isotopic resetting is uncertain. Nonetheless, this study demonstrates that Pb is mobilized by twin formation during deformation at 600°C. We highlight that this approach offers the potential to temporally constrain larger scale processes including, tectonic and shock driven metamorphism, that hitherto have remained elusive.

191 Acknowledgments

192 The study is supported by the Science and Industry Endowment Fund (SIEF RI13-01). The authors gratefully acknowledge support of Curtin University's Microscopy & Microanalysis 193 Facility and the John de Laeter Centre, whose instrumentation has been supported by 194 University, State and Commonwealth Government funding. DF acknowledges ARC funding 195 DE190101307. We thank Ian Buick for providing the samples used in this study. Stéphanie 196 Reynaud (Université de Saint-Etienne) is thanked for her help with FIB sample preparation of 197 the TEM specimen. HL thanks the electron microscope facility at the University of Lille and 198 the support of the Chevreul Institute, the European FEDER and Région Hauts-de-France. 199 AMSG thanks CNRS INSU (TelluS-SYSTER) for financial support. We are thankful for 200 constructive reviews by Randall Parrish and Fernando Corfu, and editorial handling by 201 202 Dennis Brown.

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204 Figure caption

Fig. 1: A) Boundary map of India; B) Simplified geological map of the Sandmata granulite
complex with location of sample, modified from Buick et al. (2010). AS: Aravalli Supergroup;
BGC: Banded Gneiss Complex; DS: Delhi Supergroup; C) Crystallographic orientation EBSD

map. Grain Reference Orientation Deviation (GROD) up to 15° on a green colour scale; Twins
are colour coded red for misorientation ranging between 0 and 20° from the reference point
"+"; modified from Erickson et al. (2015). The atom probe Tomography (APT) samples were
extracted from a (100) twin (specimen 1) and one neoblast (specimen 2, 3 and 4); D) Lower
hemisphere, equal area projection at site of atom probe specimen 1. The green symbols
represent the monazite host and the red symbols the twin domain.

Fig. 2: A) Reconstructed three-dimensional atom probe image of Ca, Si and ²⁰⁸Pb distribution 214 215 for specimen 1 (twin + host) and specimen 2 (neoblast). Each sphere represents one atom. The 216 twinned domain is marked by a homogeneous Ca, Si and Pb distribution whereas the host is 217 heterogeneous with apatite nano-inclusions. The neoblast Ca, Si and Pb distribution is homogeneous. B) Nanoscale ²⁰⁸Pb/²³²Th age data. The ages are colour coded by grain domains 218 219 (Green: host; Blue: neoblast; Orange: twin). The host monazite is consistent with the granulite 220 metamorphism whereas the neoblast and twin data with deformation associated with 221 amphibolite metamorphism.

Fig. 3: A) HR-HAADF image of the twin boundary showing the absence of amorphous material within the boundary. C) Indexed Fourier Transform of a square area on the right side of image B and the structure model seen along the [010] orientation. Only the heavy REE atoms are seen on the HR-HAADF image.

Fig. 4: Schematic model of the age resetting of the twin during its formation. Diagram not atscale.

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230 References

231	Beyerlein, I. J., Zhang, X., and Misra, A., 2014, Growth twins and deformation twins in
232	metals: Annual Review of Materials Research, v. 44, p. 329-363.
233	Buick, I., Allen, C., Pandit, M., Rubatto, D., and Hermann, J., 2006, The Proterozoic
234	magmatic and metamorphic history of the Banded Gneiss Complex, central
235	Rajasthan, India: LA-ICP-MS U-Pb zircon constraints: Precambrian Research, v. 151,
236	no. 1, p. 119-142.
237	Christian, J. W., and Mahajan, S., 1995, Deformation twinning: Progress in materials science,
238	v. 39, no. 1-2, p. 1-157.
239	Cottrell, A., and Bilby, B., 1951, LX. A mechanism for the growth of deformation twins in
240	crystals: The London, Edinburgh, and Dublin Philosophical Magazine and Journal of
241	Science, v. 42, no. 329, p. 573-581.
242	Cottrell, A. H., and Bilby, B., 1949, Dislocation theory of yielding and strain ageing of iron:
243	Proceedings of the Physical Society. Section A, v. 62, no. 1, p. 49.
244	Erickson, T., Pearce, M., Taylor, R., Timms, N. E., Clark, C., Reddy, S., and Buick, I., 2015,
245	Deformed monazite yields high-temperature tectonic ages: Geology, v. 43, no. 5, p.
246	383-386.
247	Erickson, T., Reddy, S., Timms, N., Pearce, M., Taylor, R., Clark, C., and Buick, I., 2016a,
248	Deformed monazite yields high-temperature tectonic ages: REPLY: Geology, v. 44,
249	no. 1, p. e378-e378.
250	Erickson, T. M., Cavosie, A. J., Pearce, M. A., Timms, N. E., and Reddy, S. M., 2016b,
251	Empirical constraints on shock features in monazite using shocked zircon inclusions:
252	Geology, v. 44, no. 8, p. 635-638.
253	Ferrill, D. A., Morris, A. P., Evans, M. A., Burkhard, M., Groshong Jr, R. H., and Onasch, C.
254	M., 2004, Calcite twin morphology: a low-temperature deformation geothermometer:
255	Journal of structural Geology, v. 26, no. 8, p. 1521-1529.
256	Fougerouse, D., Kirkland, C. L., Saxey, D. W., Seydoux-Guillaume, A. M., Rowles, M. R.,
257	Rickard, W. D. A., and Reddy, S. M., 2020, Nanoscale isotopic dating of monazite:
258	Geostandards and Geoanalytical Research, v. in press.

- Fougerouse, D., Reddy, S. M., Kirkland, C. L., Saxey, D. W., Rickard, W. D., and Hough, R.
 M., 2019, Time-resolved, defect-hosted, trace element mobility in deformed
 Witwatersrand pyrite: Geoscience Frontiers, v. 10, no. 1, p. 55-63.
- Fougerouse, D., Reddy, S. M., Saxey, D. W., Erickson, T. M., Kirkland, C. L., Rickard, W. D.
 A., Seydoux-Guillaume, A. M., Clark, C., and Buick, I. S., 2018, Nanoscale
- 264 distribution of Pb in monazite revealed by atom probe microscopy: Chemical
 265 Geology, v. 479, p. 251-258.
- Goltrant, O., Leroux, H., Doukhan, J.-C., and Cordier, P., 1992, Formation mechanisms of
 planar deformation features in naturally shocked quartz: Physics of the Earth and
 Planetary Interiors, v. 74, no. 3-4, p. 219-240.
- Hay, R., and Marshall, D., 2003, Deformation twinning in monazite: Acta Materialia, v. 51,
 no. 18, p. 5235-5254.
- Jamison, W. R., and Spang, J. H., 1976, Use of calcite twin lamellae to infer differential stress:
 Geological Society of America Bulletin, v. 87, no. 6, p. 868-872.
- Johnston, W., and Gilman, J. J., 1959, Dislocation velocities, dislocation densities, and plastic
 flow in lithium fluoride crystals: Journal of Applied Physics, v. 30, no. 2, p. 129-144.
- Kirkland, C. L., Fougerouse, D., Reddy, S. M., Hollis, J., and Saxey, D. W., 2018, Assessing
 the mechanisms of common Pb incorporation into titanite: Chemical Geology, v. 483,
 p. 558-566.
- 278 Love, G., 1964, Dislocation pipe diffusion: Acta Metallurgica, v. 12, no. 6, p. 731-737.
- Moser, D., Cupelli, C., Barker, I., Flowers, R., Bowman, J., Wooden, J., and Hart, J., 2011,
 New zircon shock phenomena and their use for dating and reconstruction of large
- 281 impact structures revealed by electron nanobeam (EBSD, CL, EDS) and isotopic U–
- 282 Pb and (U–Th)/He analysis of the Vredefort dome: Canadian Journal of Earth
 283 Sciences, v. 48, no. 2, p. 117-139.
- Moser, D., Davis, W., Reddy, S., Flemming, R., and Hart, R., 2009, Zircon U–Pb strain
 chronometry reveals deep impact-triggered flow: Earth and Planetary Science
 Letters, v. 277, no. 1, p. 73-79.
- Ni, Y., Hughes, J. M., and Mariano, A. N., 1995, Crystal chemistry of the monazite and
 xenotime structures: American Mineralogist, v. 80, no. 1, p. 21-26.

Piazolo, S., Austrheim, H., and Whitehouse, M., 2012, Brittle-ductile microfabrics in 289 290 naturally deformed zircon: Deformation mechanisms and consequences for U-Pb dating: American Mineralogist, v. 97, no. 10, p. 1544-1563. 291 292 Ranganathan, S., 1966, On the geometry of coincidence-site lattices: Acta Crystallographica, 293 v. 21, no. 2, p. 197-199. 294 Reddy, S. M., Saxey, D. W., Rickard, W. D. A., Fougerouse, D., Montalvo, S. D., Verberne, R., and van Riessen, A., 2020, Atom Probe Tomography: Development and Application 295 to the Geosciences: Geostandards and Geoanalytical Research, v. 44, no. 1, p. 5-50. 296 Reddy, S. M., van Riessen, A., Saxey, D. W., Johnson, T. E., Rickard, W. D., Fougerouse, D., 297 Fischer, S., Prosa, T. J., Rice, K. P., and Reinhard, D. A., 2016, Mechanisms of 298 299 deformation-induced trace element migration in zircon resolved by atom probe and correlative microscopy: Geochimica et Cosmochimica Acta, v. 195, p. 158-170. 300 Rickard, W. D., Reddy, S. M., Saxey, D. W., Fougerouse, D., Timms, N. E., Daly, L., 301 302 Peterman, E., Cavosie, A. J., and Jourdan, F., 2020, Novel Applications of FIB-SEM-Based ToF-SIMS in Atom Probe Tomography Workflows: Microscopy and 303 304 Microanalysis, p. 1-8. Seydoux-Guillaume, A.-M., Deschanels, X., Baumier, C., Neumeier, S., Weber, W. J., and 305 Peuget, S., 2018, Why natural monazite never becomes amorphous: Experimental 306 307 evidence for alpha self-healing: American Mineralogist, v. 103, no. 5, p. 824-827. 308 Shannon, R. D., 1976, Revised effective ionic radii and systematic studies of interatomic 309 distances in halides and chalcogenides: Acta crystallographica section A: crystal 310 physics, diffraction, theoretical and general crystallography, v. 32, no. 5, p. 751-767. 311 Tang, X., Li, Q.-L., Zhang, B., Wang, P., Gu, L.-X., Ling, X.-X., Fei, C.-H., and Li, J.-H., 2020, 312 The Chemical State and Occupancy of Radiogenic Pb, and Crystallinity of RW-1 313 Monazite Revealed by XPS and TEM: Minerals, v. 10, no. 6, p. 504. Timms, N. E., Kinny, P. D., Reddy, S. M., Evans, K., Clark, C., and Healy, D., 2011, 314 Relationship among titanium, rare earth elements, U-Pb ages and deformation 315 microstructures in zircon: Implications for Ti-in-zircon thermometry: Chemical 316 Geology, v. 280, no. 1-2, p. 33-46. 317 Timms, N. E., Reddy, S. M., Healy, D., Nemchin, A. A., Grange, M. L., Pidgeon, R. T., and 318 319 Hart, R., 2012, Resolution of impact-related microstructures in lunar zircon: A shock-

- deformation mechanism map: Meteoritics & Planetary Science, v. 47, no. 1, p. 120-
- **321** 141.
- Yardley, B. W., and Valley, J. W., 1997, The petrologic case for a dry lower crust: Journal of
 Geophysical Research: Solid Earth, v. 102, no. B6, p. 12173-12185.

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