

# The Climate Response to Emissions Reductions due to COVID-19: Initial Results from CovidMIP

Chris D. Jones<sup>1</sup>, Jonathan E. Hickman<sup>2</sup>, Steven T. Rumbold<sup>3</sup>, Jeremy Walton<sup>1</sup>, Robin D. Lamboll<sup>4</sup>, Ragnhild B. Skeie<sup>5</sup>, Stephanie Fiedler<sup>6,7</sup>, Piers M. Forster<sup>8</sup>, Joeri Rogelj<sup>4,9</sup>, Manabu Abe<sup>10</sup>, Michael Botzet<sup>11</sup>, Katherine Calvin<sup>12,13</sup>, Christophe Cassou<sup>14</sup>, Jason N.S. Cole<sup>15</sup>, Paolo Davini<sup>16</sup>, Makoto Deushi<sup>17</sup>, Martin Dix<sup>18</sup>, John C. Fyfe<sup>15</sup>, Nathan P. Gillett<sup>15</sup>, Tatiana Ilyina<sup>11</sup>, Michio Kawamiya<sup>10</sup>, Maxwell Kelley<sup>19,2</sup>, Slava Kharin<sup>15</sup>, Tsuyoshi Koshiro<sup>17</sup>, Hongmei Li<sup>11</sup>, Chloe Mackallah<sup>18</sup>, Wolfgang A. Müller<sup>11</sup>, Pierre Nabat<sup>20</sup>, Twan van Noije<sup>21</sup>, Paul Nolan<sup>22,23</sup>, Rumi Ohgaito<sup>10</sup>, Dirk Olivié<sup>24</sup>, Naga Oshima<sup>17</sup>, Jose Parodi<sup>25</sup>, Thomas J. Reerink<sup>21</sup>, Lili Ren<sup>26</sup>, Anastasia Romanou<sup>2</sup>, Roland Séférian<sup>20</sup>, Yongming Tang<sup>1</sup>, Claudia Timmreck<sup>11</sup>, Jerry Tjiputra<sup>27</sup>, Etienne Tourigny<sup>28</sup>, Kostas Tsigaridis<sup>29,2</sup>, Hailong Wang<sup>12</sup>, Mingxuan Wu<sup>12</sup>, Klaus Wyser<sup>30</sup>, Shuting Yang<sup>31</sup>, Yang Yang<sup>26</sup>, Tilo Ziehn<sup>18</sup>.

<sup>1</sup>Met Office Hadley Centre, Exeter, UK.

<sup>2</sup>NASA Goddard Institute for Space Studies, New York, NY, USA.

<sup>3</sup>National Centre for Atmospheric Science, University of Reading, UK.

<sup>4</sup>Grantham Institute for Climate Change and the Environment, Imperial College London, London, UK.

<sup>5</sup>CICERO Center for International Climate Research, Oslo, Norway.

<sup>6</sup>University of Cologne, Institute of Geophysics and Meteorology, Cologne, Germany.

<sup>7</sup>Hans-Ertel-Centre for Weather Research, Climate Monitoring and Diagnostics, Bonn/Cologne, Germany.

<sup>8</sup>Priestley International Centre for Climate, University of Leeds, UK.

<sup>9</sup>International Institute for Applied Systems Analysis (IIASA), Laxenburg, Austria.

<sup>10</sup>Japan Agency for Marine-Earth Science and Technology, 3173-25 Showamachi, Kanazawa-ward, Yokohama, 236-0001, Japan.

<sup>11</sup>Max Planck Institute for Meteorology, Hamburg, Germany.

<sup>12</sup>Pacific Northwest National Laboratory, Richland, WA, USA.

<sup>13</sup>Pacific Northwest National Laboratory, College Park, MD, USA.

<sup>14</sup>CECI, Université de Toulouse, CNRS, CERFACS, Toulouse, France.

<sup>15</sup>Canadian Centre for Climate Modelling and Analysis, Environment and Climate Change Canada, Victoria, BC, Canada.

<sup>16</sup>Istituto di Scienze dell'Atmosfera e del Clima, Consiglio Nazionale delle Ricerche (CNR-ISAC), Torino, Italy.

<sup>17</sup>Meteorological Research Institute, Japan Meteorological Agency, 1-1 Nagamine, Tsukuba, Ibaraki, 305-0052, Japan.

<sup>18</sup>CSIRO Oceans and Atmosphere, Aspendale, VIC, Australia.

<sup>19</sup>SciSpace LLC, New York, NY, USA.

38 <sup>20</sup>CNRM, Université de Toulouse, Météo-France, CNRS, Toulouse, France.

39 <sup>21</sup>Royal Netherlands Meteorological Institute (KNMI), 3730 AE De Bilt, the Netherlands.

40 <sup>22</sup>Irish Centre for High-End Computing (ICHEC), 2, 7/F, The Tower, Trinity Technology &  
41 Enterprise Campus, Grand Canal Dock, Dublin 2, Ireland.

42 <sup>23</sup>Research and Applications Division, Met Éireann, Dublin, Ireland.

43 <sup>24</sup>NORCE Norwegian Meteorological Institute, Oslo, Norway.

44 <sup>25</sup>Spanish State Meteorological Agency (AEMET), DT Murcia, Avda de la Libertad 11, 30107  
45 Murcia, Spain.

46 <sup>26</sup>Jiangsu Key Laboratory of Atmospheric Environment Monitoring and Pollution Control,  
47 Jiangsu Collaborative Innovation Center of Atmospheric Environment and Equipment  
48 Technology, School of Environmental Science and Engineering, Nanjing University of  
49 Information Science and Technology, Nanjing, China.

50 <sup>27</sup>Norwegian Research Centre and Bjerknes Centre for Climate Research, Bergen, Norway.

51 <sup>28</sup>Earth Sciences Department, Barcelona Supercomputing Center (BSC), C/ Jordi Girona 29,  
52 08034 Barcelona, Spain.

53 <sup>29</sup>Center for Climate Systems Research, Columbia University, New York, NY, USA.

54 <sup>30</sup>Rosby Centre, Swedish Meteorological and Hydrological Institute (SMHI), 601 76  
55 Norrköping, Sweden.

56 <sup>31</sup>Danish Meteorological Institute (DMI), 2100 Copenhagen, Denmark.

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58 Corresponding author: Chris D. Jones ([chris.d.jones@metoffice.gov.uk](mailto:chris.d.jones@metoffice.gov.uk))

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60 **Key Points:**

- 61 • Lockdown restrictions during COVID-19 have reduced emissions of aerosols and  
62 greenhouse gases
- 63 • 12 CMIP6 Earth system models have performed coordinated experiments to assess the  
64 impact of this on climate
- 65 • Aerosol amounts are reduced over southern and eastern Asia but there is no detectable  
66 change in annually averaged temperature or precipitation  
67

68 **Abstract**

69 Many nations responded to the COVID-19 pandemic by restricting travel and other activities  
70 during 2020, resulting in temporarily reduced emissions of CO<sub>2</sub>, other greenhouse gases and  
71 ozone and aerosol precursors. We present the initial results from a coordinated Intercomparison,  
72 CovidMIP, of Earth system model simulations which assess the impact on climate of these  
73 emissions reductions. Twelve models performed multiple initial-condition ensembles to produce  
74 over 300 simulations spanning both initial condition and model structural uncertainty. We find  
75 model consensus on reduced aerosol amounts (particularly over southern and eastern Asia) and  
76 associated increases in surface shortwave radiation levels. However, any impact on near-surface  
77 temperature or rainfall during 2020-2024 is extremely small and is not detectable in this initial  
78 analysis. Regional analyses on a finer scale, and closer attention to extremes (especially linked to  
79 changes in atmospheric composition and air quality) are required to test the impact of COVID-  
80 19-related emission reductions on near-term climate.

81

82 **Plain Language Summary**

83 Many nations responded to the COVID-19 pandemic by restricting travel and other activities  
84 during 2020. This caused a temporary reduction in emissions of CO<sub>2</sub> and other pollutants. We  
85 compare results from twelve Earth system models to see if the emissions reductions affected  
86 climate. These twelve models performed over 300 experiments using multiple initial-conditions.  
87 We find a consensus that aerosol amounts were reduced, especially over southern and eastern  
88 Asia, during 2020-2024. This led to increases in solar radiation reaching the surface in this  
89 region. However, we could not detect any associated impact on temperature or rainfall. We  
90 recommend more analyses on regional scales. We also suggest that analysis of extreme weather  
91 and air quality would be useful to test the impact on climate of emission reductions due to  
92 COVID-19.

## 93 1 Introduction

### 94 1.1 Impact of COVID-19 lockdown on emissions

95

96 The COVID-19 pandemic led to widespread measures restricting travel, industrial, and  
97 commercial activity during 2020. The effects of these changes in socioeconomic activity on  
98 atmospheric composition have been widely studied including estimates of emissions and  
99 concentrations of species that directly or indirectly affect climate.

100

101 The impacts of COVID-19 measures on long-lived greenhouse gases have been inferred  
102 from both bottom-up estimates using activity data and top-down analysis of atmospheric  
103 observations. Bottom-up estimates using sector activity have estimated global CO<sub>2</sub> emissions  
104 reductions of 8.8% during the first 5 months of 2020 (Liu et al., 2020) and annual reductions  
105 from 4% to 7% (Le Quéré et al., 2020). Top-down assessments have found some indications of a  
106 decrease in CO<sub>2</sub> growth rate during 2020 (Buchwitz et al., 2020), with examples of substantial  
107 local and regional CO<sub>2</sub> and methane (CH<sub>4</sub>) emissions reductions inferred from surface  
108 observations (Tohjima et al., 2020; Turner et al., 2020). However, existing satellite products  
109 could not provide the required coverage to reliably detect changes in CO<sub>2</sub> column densities at the  
110 magnitude expected to be occurring in 2020 (Buchwitz et al., 2020; Chevallier et al., 2020).  
111 Expected growth rates in atmospheric CO<sub>2</sub> fractions vary too much from year to year due to  
112 internal climate variability (Jones and Cox, 2005; Betts et al., 2016) for the effects of emission  
113 reductions on the order of 8% to be clearly detected from observations of CO<sub>2</sub> column densities  
114 (Sussmann and Rettinger, 2020; Tohjima et al., 2020). The long lifetime of CO<sub>2</sub>, and to a lesser  
115 extent CH<sub>4</sub>, means that the small impact of emissions reductions is likely to be long-lived, and  
116 may still exert a non-negligible climate impact on decadal timescales (Forster et al., 2020).

117

118 The largest changes in observed composition attributed to COVID-19 restrictions were  
119 for nitrogen dioxide (NO<sub>2</sub>), with concentration reductions at both national- and city-scales  
120 typically on the order of 20-60% in China, India, Europe, and the United States (Goldberg et al.,  
121 2020; Keller et al., 2020; Menut et al., 2020; Miyazaki et al., 2020; Ordóñez et al., 2020; Venter  
122 et al., 2020; Zhao et al., 2020). The NO<sub>2</sub> decreases have been attributed largely to changes in the  
123 transport sector (Bao and Zhang, 2020; Diamond and Wood, 2020; Lian et al., 2020b; Venter et  
124 al., 2020). The rapid changes in emissions and complex dynamics of short-lived pollutants have  
125 complex and non-uniform implications for climate. In areas where background NO<sub>x</sub>  
126 concentrations were high, reduced NO<sub>x</sub> emissions led to increased tropospheric ozone (O<sub>3</sub>)  
127 concentrations in many regions and cities (Keller et al., 2020; Le et al., 2020; Lian et al., 2020a;  
128 Ordóñez et al., 2020; Sicard et al., 2020; Venter et al., 2020). Elsewhere, tropospheric ozone may  
129 have decreased during lockdowns leading to short-term estimated changes of radiative forcing by  
130 -33 to -78 mWm<sup>-2</sup> (Weber et al., 2020).

131

132 Some studies report substantial decreases in particulate matter (PM) on the order of 10-  
133 30% (Filonchik et al., 2020; Le et al., 2020; Silver et al., 2020; Venter et al., 2020; Xu et al.,  
134 2020), but analyses accounting for long-term trends generally found no lockdown impacts on  
135 aerosol optical depth (AOD) or PM concentrations (Diamond and Wood, 2020; Field et al.,  
136 2020; Zangari et al., 2020). In some regions, PM concentrations increased as a result of altered  
137 dust or biomass burning emissions or as a consequence of changes in emissions and meteorology

138 (Le et al., 2020; Venter et al., 2020). Notably, northern China experienced an increase in haze  
139 during the spring lockdown due to enhanced formation of ozone, which, in combination with  
140 favorable meteorological conditions and changes in heterogeneous chemistry, contributed to  
141 enhanced secondary aerosol formation (Chang et al., 2020; Le et al., 2020; Wang et al., 2020b).  
142

## 143 **1.2 Impact of emissions reductions on climate**

144 The reduction in emissions is expected to have regional impacts on atmospheric  
145 composition, and therefore could have implications for weather and climate. Different species  
146 have very different lifetimes from hours-to-days for aerosols, to decades or longer for long-lived  
147 greenhouse gases, and very different spatial scales, with some being very localized and others  
148 globally well-mixed.

149 For example, Yang et al. (2020) examined climate responses to aerosol emission  
150 reductions during the COVID-19 lockdown, back-to-work and post-lockdown stages throughout  
151 the year 2020 based on CESM1 model simulations. They reported that an anomalous surface  
152 warming appeared over the Northern Hemisphere continents due to the fast climate response to  
153 aerosol reductions. Fyfe et al. (2021) examine a large ensemble of simulations with CanESM5  
154 under an idealized modelling framework of the COVID emission reduction and conclude that  
155 any signal from such short-lived emissions changes is likely to be small or even undetectable.  
156 Forster et al. (2020) developed a two-year COVID-19 emissions reduction scenario for long- and  
157 short-lived species based on mobility data and the bottom-up approach of Le Quéré et al. (2020)  
158 for some sectors and then assumed a recovery over the subsequent two years. Using the FaIR  
159 climate emulator, they simulated the effect of these emissions reductions and found a rapid short-  
160 term warming due to reduced aerosols, which was offset by a slightly slower, but also near-term  
161 cooling due to reduced tropospheric ozone. On longer timescales, well-mixed GHGs, especially  
162 CO<sub>2</sub> became important, and their simulations showed that the net effect of these emissions  
163 changes by 2030 was negligible: a global cooling of about  $0.01 \pm 0.005$  °C.

164 However, because FaIR cannot capture regional climate effects, internal variability or  
165 complex interactions of atmospheric composition and biogeochemistry, there remain unanswered  
166 questions about the possible climatic impact of emissions reductions on regional air quality and  
167 climate. These are beginning to be addressed by single model studies (e.g. Yang et al., 2020  
168 analyse an atmospheric model with prescribed sea surface temperature, and ; Fyfe et al, 2021  
169 analyse a large ensemble of coupled atmosphere-ocean simulations with the CanESM5 model),  
170 but would benefit greatly from being analyzed across an ensemble of Earth system models  
171 (ESMs) run under a common protocol. Hence it was decided that this scenario would form the  
172 basis of a multi-Earth system model intercomparison project (MIP). This paper presents an initial  
173 analysis of the first results coming from this new activity, called CovidMIP. The emissions  
174 estimates and modelling protocol used are described in section 2, results shown in section 3 and  
175 discussed in section 4 in the context of ongoing climate change.

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177

## 178 2 Materials and Methods

### 179 2.1 CovidMIP protocol

180 The emissions estimates assembled by Forster et al. (2020) were collated and gridded,  
181 and made available in Inputs4MIPs data format for use by CMIP Earth system models (Lamboll  
182 et al., 2020). A modelling protocol was agreed, and is incorporated into DAMIP (the Detection  
183 and Attribution MIP; Gillett et al., 2016), which is also described in Lamboll et al. (2020), but  
184 the main points are noted here for convenience.

185 Because any climate signal due to COVID-19-induced emissions reductions was  
186 considered likely to be small, it is advantageous to carry out large initial-condition ensembles  
187 which have been shown to enable detection of even small regional climate signals (e.g. Banerjee  
188 et al., 2020). But cognizant of the computational cost and time required for producing such large  
189 ensembles, a pragmatic recommendation was made that model groups perform at least 10 initial-  
190 condition ensemble members. This was hoped to maximize the number of modelling groups  
191 participating but still produce enough members to enable meaningful analysis.

192 The protocol uses the SSP2-4.5 scenario (O'Neill et al., 2016) as a baseline against which  
193 to apply the emissions reductions. Simulations are run parallel to ssp245, but branching from that  
194 simulation on 1 January 2020 and following the new forcing in line with emissions reductions.  
195 The results will be published on the CMIP6 archive (Earth System Grid Federation) under  
196 experiment name ssp245-covid. Forcing is provided as concentrations of greenhouse gases and  
197 emissions of aerosols and aerosol and ozone precursors. For models with interactive chemistry,  
198 ozone can be simulated otherwise it has been provided as concentrations. Similarly, models can  
199 simulate aerosols or they can be represented with the MACv2-SP parametrisation (Stevens et al.,  
200 2017; Fiedler et al., 2020).

201 In this manuscript we focus on the immediate term impact (from 2020-2024) of the “two  
202 year blip” scenario under which emissions revert to the baseline levels by the end of 2022. In  
203 addition to this, Forster et al. (2020) created a set of scenarios spanning possible future economic  
204 recovery strategies: a reduction in anthropogenic CO<sub>2</sub> emissions post-2020 consistent with  
205 enhanced investment in environmentally friendly technologies (moderate or strong “green  
206 stimulus”), no effect after 2022 (continuation of “two year blip” studied here with emissions  
207 reverting to ssp245) or an increase in anthropogenic CO<sub>2</sub> emissions relative to ssp245 after 2020  
208 consistent with an investment in more traditional fossil-fuel based energy production (or “fossil-  
209 fuelled recovery”). All of these scenarios have become part of the CMIP6 set of experiments,  
210 under the detection and attribution activity (DAMIP: Gillett et al., 2016).

### 211 2.2 Participating Earth system models

212 The protocol is open to all models participating in CMIP6 and to date twelve models  
213 have provided data for analysis (Table 1). A particular value of a multi-model ensemble is being  
214 able to incorporate different levels of process complexity, but this also brings challenges of  
215 interpreting results.

216 Some models prescribe aerosols and ozone, either using their own climatology or  
217 MACv2-SP and/or prescribed ozone 3D concentrations taken from the OsloCTM3 chemical  
218 transport model (Lamboll et al., 2020). Others may simulate either aerosols or ozone  
219 interactively in response to their primary or secondary emissions. The MPI-ESM1-2-LR model  
220 simulated interactive CO<sub>2</sub> while the other models used prescribed CO<sub>2</sub> concentrations. Models  
221 have differing complexity and species richness of aerosols, representing both natural and

222 anthropogenic species such as sulphates, black carbon, organic carbon, sea-salt and mineral-dust,  
223 but many still lack representation of nitrate aerosols.

224 In terms of biogeochemistry many ESMs now represent land and marine ecosystems and  
225 the carbon cycle (Boysen et al., 2020; Séférian et al., 2020; Thornhill et al., 2020). On the near-  
226 term studied here, the carbon cycle is unlikely to have a large effect on climate but impacts of  
227 emissions reductions may show up in terms of changes in carbon fluxes, stores and partitioning  
228 across realms of the Earth system.

229 To generate initial conditions some models (ACCESS-ESM1-5, CanESM5, EC-Earth3,  
230 MIROC-ES2L, MPI-ESM1-2-LR, UKESM1-0-LL) drew on existing ssp245 simulations which  
231 followed on from initial-condition ensembles of the CMIP6 historical simulations. Others  
232 perturbed conditions at the end of the historical period (CESM1, E3SM-1-1, GISS-E2-1-G), or  
233 mixed the two approaches by inflating existing ensembles with additional perturbations applied  
234 (MRI-ESM2-0, CNRM-ESM2-1, NorESM2-LM) or by running on different super-computers  
235 (NorESM2-LM).

236 Future studies will be able to take into account the model complexity and how this affects  
237 the simulated results. For example, are changes in atmospheric circulation or surface climate  
238 affected differently between models with simulated and prescribed ozone and aerosols? How  
239 does the model treatment of interactions between atmospheric composition (such as fraction of  
240 diffuse light or surface ozone) affect vegetation productivity and carbon storage? In this analysis  
241 such considerations are out of scope and we give an overview on each model's results for the  
242 climate response for 2020-2024. The reader is referred to Table 1, which documents the spatial  
243 resolution and the process complexity of each participating model as well as the number of  
244 ensemble members utilized in this study.

245

## 246 **3 Results**

### 247 **3.1 Indicators of global change**

248 Our analysis draws on different sized ensembles from 12 ESMs. Throughout, we base  
249 analysis on ensemble mean anomalies from each model, calculated from a pair-wise difference  
250 between simulations with COVID-19-related emissions reductions (“ssp245-covid”) and  
251 simulations using the standard, baseline SSP2-4.5 scenario (“ssp245”).

252 Globally, for 2020, all models show a reduction in aerosol optical depth (at 550 nm) in  
253 their ensemble mean with 7 out of 11 models which reported this variable having a reduction  
254 greater than 1 standard deviation (Figure 1). In 2021, the AOD anomalies of 10 out of 11 models  
255 remain negative with ACCESS-ESM1-5 showing near-zero deviation. From 2022 onwards there  
256 is no robust global signal in AOD as emissions reductions in this simulation recover to levels in  
257 the baseline scenario and aerosol amounts quickly recover too.

258 This behavior is reflected in the amount of solar radiation reaching the surface, which is  
259 generally simulated to have increased, with all models (of the 11 for whom this variable was  
260 available for this analysis) having a positive anomaly in downwards shortwave (SW) radiation  
261 for both 2020 and 2021 (Figure 1, panel b). Although only MRI-ESM2-0 simulated an ensemble  
262 mean global increase greater than 1 standard deviation. As for AOD, the anomaly quickly

263 recovers and becomes very small from 2022 onwards. We have not yet investigated the extent to  
264 which surface shortwave is directly affected by aerosol absorption or by aerosol-induced changes  
265 in cloud cover. Future studies will also assess impacts and implications of aerosol-cloud  
266 interactions in driving the changes seen here.

267

268 *Figure 1. Annual mean, ensemble average output from ESMs. Each panel shows anomalies from the simulations with COVID-19-*  
269 *related emissions reductions compared to the baseline SSP2-4.5 simulations (“ssp245-covid” minus “ssp245”). (a) Global*  
270 *aerosol optical depth at 550nm; (b) downwards SW radiation at the surface; (c) Global surface air temperature; (d) Global*  
271 *precipitation. Coloured lines show ensemble average results from each model, and paler plumes show ensemble spread for each*  
272 *model calculated here as  $\pm 1$  standard deviation across each model’s ensemble. Vertical bars to the left of each panel show each*  
273 *model spread (mean  $\pm 1$  standard deviation) for the first year, 2020. Each model has performed a different number of ensemble*  
274 *members as listed in Table 1 and shown in square brackets in the caption.*

275 The impact of this, however, on surface climate at a global scale is very small. Figure 1  
276 panels (c) and (d) show globally averaged surface air temperature and precipitation respectively.  
277 No model shows any significant change in either of these quantities at a global level for any year.

278

### 279 **3.2 Patterns of regional changes**

280 Figure 2 shows the regional patterns of the changes in aerosol optical depth for each  
281 model. It is apparent that models agree that the largest response is in Asia, predominantly over  
282 India and China where almost all models show a marked decrease in aerosols as an average over  
283 the 5-year period 2020-2024. Some models also show some patches of aerosol increases, for  
284 example CanESM and E3SM-1-1 over the Himalayan region, and MIROC-ES2L over regions of  
285 North Africa. Reasons for these changes are not explored further here and we do not yet know if  
286 they are caused by changes in anthropogenic or natural sources, such as dust, which can be very  
287 sensitive to variations in windspeed.  
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*Figure 2. Model by model simulated changes in aerosol optical depth (at a wavelength of 550 nm). For each model we plot the ensemble mean response from 2020-2024 inclusive. Blue colours denote a decrease in AOD. Each model has performed a different number of ensemble members as listed in Table 1 and shown in square brackets in the caption. The black box shows the region analysed in Figure 3.*

294 To see if these regional changes in aerosol loading affect regional climate properties, we  
 295 define a region bounded by 60-160°E and 0-50°N which has been chosen subjectively after  
 296 considering all models to cover the main AOD anomalies across models (marked as black boxes  
 297 in Figure 2). We assess annual changes in surface SW radiation, temperature and precipitation in  
 298 this region. Figure 3 shows a similar response to the global metrics shown in Figure 1 but with  
 299 greater magnitudes of average response. Again, there is a strong model agreement of reduced  
 300 aerosols, with all models agreeing on this in their ensemble mean for 2020 and 7 out of 11  
 301 having reductions greater than 1 standard deviation. Averaged across models, global AOD  
 302 reduction in 2020 is  $-0.0027 \pm 0.0012$ , while in southern and eastern Asia it is  $-0.0097 \pm 0.0034$ .  
 303 The associated increase in downwards SW radiation is also apparent, and stronger here: globally  
 304 models show an increase of  $0.21 \pm 0.10 \text{ Wm}^{-2}$  while in southern and eastern Asia it is  $0.69 \pm 0.31$   
 305  $\text{Wm}^{-2}$ .

306 Although most models simulate a slight warming signal in this region in their ensemble  
 307 mean (Figure 3, panel c), the magnitude is very small – less than 0.1 °C, and in all models smaller  
 308 than the standard deviation across ensemble members (typically of the order 0.2°C).

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 310  
 311 *Figure 3. indicators of change in southern and eastern Asia (defined here as 60-160°E and 0-50°N). As for figure 1 results are*  
 312 *plotted as annual mean anomalies, with coloured lines denoting ensemble means from each model and grey shading 1-standard*  
 313 *deviation for each model. (a) Aerosol optical depth; (b) surface downwards shortwave radiation; (c) surface air temperature; (d)*  
 314 *precipitation.*

315  
 316 Outside of this region, models show patchy temperature changes, indicative of random  
 317 changes, and internal variability of modes such as NAO or ENSO. This residual signal of  
 318 internal variability is not eliminated in limited ensemble size and demonstrates the weak signal-  
 319 to-noise ratio (see figure 1 in Supplementary Info). These changes do not appear to be  
 320 systematic, with some regions exhibiting both apparently strong warming and cooling signals in  
 321 different models. The region of northern East Asia often displays a strong temperature signal in  
 322 the model results, with CESM1 displaying a warming as reported in Yang et al. (2020), although  
 323 that study performed simulations with fixed sea-surface temperatures. GISS-E2-1-G and E3SM-  
 324 1-1 also show strong warming patterns here and UKESM1-0-LL, MPI-ESM1-2-LR and  
 325 CanESM5 some warming too. But NorESM2-LM shows a strong cooling and ACCESS-ESM1-  
 326 5, MIROC-ES2L and MRI-ESM2-0 having mixed signals. Models show marked differences  
 327 elsewhere e.g. MPI-ESM1-2-LR and MRI-ESM2-0 have opposite patterns of warming over  
 328 North America while in South America CanESM5 and UKESM1-0-LL show a cooling but  
 329 GISS-E2-1-G and NorESM2-LM show a warming.

330 When looking at regional patterns of precipitation and surface SW radiation (S.I. figures  
 331 2 and 3) there are no robust signals or consistent patterns of change across models. Even the  
 332 increase in surface SW radiation shown in Figure 3 is very hard to see by eye in the patterns of  
 333 change, due to the influence of clouds which can easily mask any signal from changes in  
 334 aerosols. This, and similar incoherent patterns of rainfall change, indicate the substantial  
 335 variability in these quantities and the challenges in detecting robust signals of change under  
 336 conditions of relatively small forcing. Despite a large number of ensembles, it is evident that at  
 337 these smaller regional scales, variability in meteorology prevents robust detection of signals in  
 338 clouds and rainfall.

339  
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#### 341 **4 Discussion and Conclusions**

342 Here we have only begun to scratch the surface of the results becoming available from  
343 the CovidMIP simulations. We stress that this work has been the result of a very rapid response  
344 of the Earth system modelling community. It often takes several years to design and perform  
345 coordinated MIP experiments, and process the data for publication in a community archive. This  
346 activity has taken place in only a matter of months. This paper is just the very first analysis of  
347 initial results and therefore serves only as a first indication of how the climate system has  
348 responded to the perturbations to emissions in response to the COVID-19 pandemic. It is not  
349 possible at this stage to analyze all of the responses, nor the processes responsible for changes  
350 across the whole system. But this study sets the scene and informs priorities for future analysis.

351  
352 We have shown that the imprint of COVID-19-related changes in societal activity is  
353 visible in atmospheric composition – notably aerosol optical depth over southern and eastern  
354 Asia, and in the amount of solar radiation reaching the planet’s surface. Over this most affected  
355 region, the 2-year average effect was more than  $0.5\text{Wm}^{-2}$ . More locally and on shorter timescales  
356 it could be substantially higher. However, despite these changes in the make-up of the  
357 atmosphere, no detectable change in surface temperatures or rainfall could be found. We  
358 conclude that the emissions reductions were too small in magnitude and time to have a  
359 significant effect on global climate, and that larger, sustained changes on a much longer  
360 timescale are required in order to have observable effects (Samset et al., 2020; Tebaldi et al,  
361 2020). The CovidMIP protocol will be extended to include an additional “four year blip”  
362 simulation so that future work can also consider the impact if lockdown restrictions were  
363 prolonged or recovery delayed due to new strains of the Coronavirus.

364  
365 Based on what we have found we recommend further analysis would be fruitful in the  
366 following areas:

- 367 • *Effective radiative forcing (ERF) response to the emissions perturbations.* The  
368 global patterns of downwards SW radiation anomalies are very noisy in these  
369 simulations but the radiation signal would be improved in simulations with fixed-  
370 SSTs which reduce interannual variability in the climate system and allow  
371 quantification of the ERF due to the emission changes (Pincus et al., 2016; Fiedler  
372 et al., 2020). The CovidMIP protocol (Lamboll et al., 2020) defines additional  
373 fixed-SST simulations to isolate the effects of ozone, aerosols and even separate  
374 black carbon, organic carbon and sulphate aerosols. We recommend model groups  
375 perform these complementary simulations to allow the radiative effects of  
376 emissions reductions to be assessed more reliably.
- 377 • *Attribution of drivers of climate signals.* As part of DAMIP, this activity has a  
378 strong interest in performing single-forcing simulations to enable understanding  
379 of different drivers and causes of the climate changes seen. Large ensembles have  
380 been shown to be successful in detecting and attributing changes, e.g., in recent  
381 southern hemisphere circulation changes to stratospheric ozone recovery  
382 (Banerjee et al., 2020). Similar techniques could be used here to separate the  
383 impacts of emissions reductions of GHGs and aerosols as explored in CanESM5  
384 by Fyfe et al. (2021).
- 385 • *Longer term implications of emissions reductions and options for economic  
386 recovery.* Forster et al. (2020) compiled a set of hypothetical recovery scenarios

387 based on moderate or strong green stimulus packages or a fossil-fuel stimulus  
388 rebound. The climate impacts by 2050 showed that how the world's economy  
389 recovers after 2020 can have profound impacts on our ability to meet long-term  
390 climate goals. Multi-model analysis of these simulations will enable clearer  
391 understanding of the threats and opportunities arising from the current situation.

- 392 • *Quantifying changes in extremes.* In addition to annual mean changes, the climate  
393 response in terms of extremes – such as daily maximum or minimum  
394 temperatures or daily precipitation rates – may also show important signals  
395 (Seneviratne and Hauser, 2020).
- 396 • *Influence on atmospheric circulation.* Studies have found a sensitivity of  
397 monsoons to changes in emissions of aerosols (Meehl et al., 2008; Li et al., 2016;  
398 Lau et al., 2017; Zhao et al., 2019). Analysis of these changes in a multi-model  
399 study may be able to detect if COVID-19-related emissions reductions have had a  
400 detectable impact on monsoon circulations, especially over Asia.
- 401 • *Response and impacts of atmospheric composition.* The response of aerosols is  
402 detectable in this ensemble, but we have not yet explored the role of other  
403 chemically active components of the atmospheric composition. Especially, the  
404 role of ozone and its response to changes in emissions of precursors, is a key  
405 components of changes in air quality. Multiple studies have found increases in  
406 ozone in populated urban areas during lockdown (e.g. Keller et al., 2020), in  
407 contrast to a global decrease in tropospheric ozone (Weber et al., 2020). This MIP  
408 provides an opportunity to shed process-level understanding on these changes in a  
409 range of models of varying degrees of complexity with regards to atmospheric  
410 chemistry.
- 411 • *Impact on the global carbon cycle.* There is increasing interest in the ability to  
412 make predictions from one year to the next of changes in atmospheric CO<sub>2</sub> (Betts  
413 et al., 2016; Séférian et al., 2018; Lovenduski et al., 2019; Fransner et al., 2020;  
414 Spring and Ilyina, 2020). These studies require knowledge of natural causes of  
415 interannual variability – notable from ENSO (Watanabe et al., 2020), but they  
416 also require knowledge of up to date estimates of anthropogenic CO<sub>2</sub> emissions.  
417 These are normally expected to vary relatively little from year to year (Le Quéré  
418 et al., 2018) but expected impacts from COVID-19-related emissions reductions  
419 allow us to test out ability to forecast this most important metric of climate  
420 change, and whether external forcing can affect its variability (McKinley et al.,  
421 2020).

422  
423 The SARS-Cov-2 pandemic of 2020 has created one of the biggest health and economic  
424 crises of recent history, but it also presents a remarkable opportunity to study how the climate  
425 system responds to changes in emissions of radiatively active species. From regional air quality  
426 to global climate this database of ESM outputs will enable advances in our understanding of how  
427 the climate system responds to short-term perturbations.  
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482

483 **Conflict of interest**

484 The authors declare no competing interests

485

486 **Data**

487 Model data is published on the CMIP6 archive available via the Earth System Grid Federation.  
488 <https://esgf-index1.ceda.ac.uk/search/cmip6-ceda/>

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720 *Table 1. List of participating models, their main properties and number of ensemble members used in this study.*

Model name	reference	Atmosphere resolution §	Ocean resolution §	ssp245-covid ensemble members	Aerosol processes #	Ozone forcing	Aerosol forcing
ACCESS-ESM1-5	Ziehn et al. (2020)	250km (N96), L38	100km, L50	30	5; CLASSIC	Prescribed <i>ssp245-covid</i> perturbation (Lamboll et al., 2020) *	interactive
CanESM5	Swart et al. (2019)	500km (T63), L49	100km, L45	50	5; Parameterized using a prognostic scheme for bulk concentrations	Prescribed <i>ssp245-covid</i> perturbation (Lamboll et al., 2020) *	interactive
CESM1	Hurrell et al. (2013)	250km (1.9x2.5), L30	100km (gx1v6), L60	10	6; MAM4	Prescribed without ssp245-covid perturbation	interactive
CNRM-ESM2-1	Séférián et al. (2019)	250km (TL127,1.4°), L91	100km (eORCA1), L75	100	5; TACTIC (Michou et al., 2020)	Interactive above 560 hPa, prescribed below ( Michou et al., 2020)	interactive
E3SM-1-1	Burrows et al. (2020)	100km (NE30), L72	60-30 km, L100	10	7; MAM4 (Wang et al., 2020a)	Prescribed without ssp245-covid perturbation	interactive
EC-Earth3		100km (T255), L91	100km (eORCA1), L75	30	n/a	Prescribed <i>ssp245-covid</i> perturbation (Lamboll et al., 2020)	MACv2-SP (Fiedler et al., 2020)
MIROC-ES2L	Hajima et al. (2020); Kawamiya et al. (2020)	500km (T42), L40	100km (360x256), L63	30	5; SPRINTARS	Prescribed <i>ssp245-covid</i> perturbation (Lamboll et al., 2020)	interactive
MPI-ESM1-2-LR	Mauritsen et al. (2019)	250km (T63), L47	150km, L40	10	n/a	Prescribed <i>ssp245-covid</i> perturbation (Lamboll et al., 2020)	MACv2-SP (Fiedler et al., 2020)
MRI-ESM2-0	Yukimoto et al. (2019); Oshima et al. (2020)	100km (TL159, 1.125°), L80	100km (tripolar 1° x 0.3° -0.5°), L61	10	5; MASINGAR mk-2r4c	interactive	interactive
GISS-E2-1-G	Kelley et al., 2020; Ito et al., 2020; Bauer et al., 2020	250km (2x2.5°), L40	100km (1x1.25°), L40	10	8; MATRIX	interactive	interactive

NorESM2-LM	Seland et al. (2020); Tjiputra et al. (2020)	250km (1.9° x 2.5°), L32	100km, L53	10	5; OsloAero6	Prescribed without ssp245-covid perturbation	interactive
UKESM1-0-LL	Sellar et al. (2019)	250km (N96), L85	100km (eORCA1), L75	16	5; UKCA MODE	interactive	interactive

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722 § shown as CMIP “nominal resolution” in km, “L” denotes number of vertical levels. Grid name or information provided if available.

723 # number of aerosol species, and name/description of aerosol sub-model

724 \* These models used the first version of the ozone fields that had a small bug in the vertical interpolation of the ozone perturbation, stretching the ozone  
725 perturbation to too high altitudes. The models were not able to re-run the model simulations with the corrected ozone fields. Radiative kernel calculations  
726 following Skeie et al. (2020) gave 0.6 mWm<sup>-2</sup> stronger total ozone radiative forcing in 2020 for the corrected fields compared to the incorrect ozone fields, that  
727 are small compared to the total ozone radiative forcing of -37 mWm<sup>-2</sup> for *ssp245-covid* relative to *ssp245* in 2020.