

1 **Continental patterns of bird migration linked to climate variability**

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12 **Abstract**

13 For ~100 years, the continental patterns of avian migration in North America have been described in
14 the context of three or four primary flyways. This spatial compartmentalization often fails to
15 adequately reflect a critical characterization of migration — phenology. This shortcoming has been
16 partly due to the lack of reliable continental-scale data, a gap filled by our current study. Here, we
17 leveraged unique radar-based data quantifying migration phenology and used an objective
18 regionalization approach to introduce a new spatial framework that reflects interannual variability.
19 Therefore, the resulting spatial classification is intrinsically different from the “flyway concept”. We
20 identified *two* regions with distinct interannual variability of spring migration across the contiguous
21 U.S. This data-driven framework enabled us to explore the climatic cues affecting the interannual
22 variability of migration phenology, “specific to each region” across North America. For example, our
23 “two-region” approach allowed us to identify an east-west dipole pattern in migratory behavior linked
24 to atmospheric Rossby waves. Also, we revealed that migration movements over the western U.S. was
25 inversely related to interannual and low-frequency variability of regional temperature. A similar link
26 but weaker and only for interannual variability was evident for the eastern region. However, this region
27 was more strongly tied to climate teleconnections, particularly to the East Pacific-North Pacific (EP-
28 NP) pattern. The results suggest that oceanic forcing in the tropical Pacific—through a chain of
29 processes including Rossby wave trains—controls the climatic conditions, associated with bird
30 migration over the eastern U.S. Our spatial platform would facilitate better understanding of the
31 mechanisms responsible for broad-scale migration phenology and its potential future changes.

32 **Capsule summary**

33 The contiguous U.S. is objectively divided into two regions based on bird migration phenology. We
34 explore the climatic cues associated with this new spatial framework.

35 **1. Motivation**

36 The seasonal migration of birds is a prominent feature of the natural world. Every spring, migratory
37 birds arrive from south and central America to the contiguous U.S. (CONUS) and Canada for breeding
38 (Gauthreaux 1971; Lowery 1945; Dokter et al. 2018; Lane et al. 2012). Exogenous forces, such as
39 climate and changes in primary productivity, influence migration speed and phenology, defined as the
40 seasonal timing of life cycle events (La Sorte et al. 2014a; Zuckerberg et al. 2020; Gordo 2007; Smith
41 and Deppe 2008). Endogenous forces, such as circadian cycles and site fidelity, also play a role
42 (Gwinner 1996; Cohen et al. 2012; Alerstam et al. 2003). Together, these forces suggest that migratory
43 pathways should be stable over time, but also reflect broad-scale and regular patterns in climate
44 variability. Traditionally, spatial classification of bird migration in CONUS is viewed in the context of
45 “flyways”, and the region is commonly divided into four principal routes (Pacific, Central, Mississippi,
46 and Atlantic), largely derived from waterfowl ecology (Hawkins 1984; Lincoln 1935; Waller et al.
47 2018). An alternative representation is three routes, western, central and eastern (La Sorte et al. 2014b;
48 Horton et al. 2020), although some similarities have been identified between the latter two routes that
49 may be indicative of a larger migration system (La Sorte et al. 2014b).

50
51 However, such a large-scale characterization of migratory routes has not been fully understood, and the
52 common spatial classification approaches are either subjective or based on the time-averaged
53 migratory behavior and therefore neglect year-to-year variability (Hawkins 1984; La Sorte et al.
54 2014b; Olsen et al. 2006). Those studies that consider interannual variability are limited to
55 observations from individual sites (Van Buskirk et al. 2009; Oliver et al. 2020; Ballard et al. 2003). To
56 fill these voids, we have proposed a new geographic framework, which would reflect the interannual
57 variability of bird migration at the continental-scale. This approach would be essential for better

58 understanding how patterns in climate variability influence broad-scale animal movements and
59 migration phenology (Strong et al. 2015; Zuckerberg et al. 2020).

60

61 The main obstacle for spatio-temporal analysis of bird migration has been the lack of reliable data over
62 a sufficiently long period and with broad spatial coverage across CONUS (Horton et al. 2020). This
63 data limitation has hampered the proper assessments of spatial properties and annual timing events.
64 Recently-published data of migration phenology, derived from weather radar observations (Horton et
65 al. 2020), provides a unique opportunity to perform such analysis at the continental-scale. Leveraging
66 these data, we have revisited the traditional spatial framework to test whether there is coherent
67 interannual and low-frequency variability in migration timing across the continent, and whether that
68 exhibits spatial variability that could be used to improve our knowledge of the drivers of year-to-year
69 variability of bird migration. In other words, we aim to identify sub-regions based on similarity of
70 interannual variability of bird migration and consequently explore regional and remote climatic drivers
71 specific to each region.

72

73 **2. Bird migration data**

74 We used nocturnal migration data that has been recently compiled from the NOAA's Next Generation
75 Radar (NEXRAD) system (Horton et al. 2020). NEXRAD is a network of 143 stations across the
76 contiguous U.S. and provides exceptional spatial and temporal coverage for continental-scale analysis
77 (Ansari et al. 2018; Dokter et al. 2019; Rosenberg et al. 2019). Real-time and archived NEXRAD data
78 are shared on Amazon Web Service (AWS) and can be accessed via simple application program
79 interfaces (APIs). The AWS cloud has facilitated data access and created new research opportunities,
80 including analysis of avian migration.

81

82 The bird migration data has been developed using a convolutional neural network (CNN) to exclude
83 precipitation contamination and subsequently quantify the phenology of migratory movements (Lin et
84 al. 2019). This approach employs a neural network trained using per-pixel labels (biology or weather)
85 derived from a polarimetric variable, specifically correlation coefficient (ρ_{HV}). Correlation coefficient
86 quantifies the consistency of the shapes and sizes of targets within the radar beam and is used to
87 distinguish between meteorological and non-meteorological objects. If the correlation coefficient
88 exceeds 0.95, reflectivity is classified as precipitation, otherwise it is classified as biological. Their
89 algorithm also removes stationary clutter. Following these filtering steps, vertical profiles of radar
90 reflectivity are constructed to quantify migration activity from 100-3000 m layer above ground level
91 (AGL), from spring (1 March to 15 June) 1995 to spring 2018. Some sites have data over a shorter
92 period. To analyze the timing of migration, consistent with Horton et al. (2020), we used median
93 migration date (q50), defined as the date by which 50% of the cumulative passage occurred at each
94 radar station.

95

96 **3. Weather and climate data**

97 Monthly meteorological data are obtained from NASA Modern-Era Retrospective Analysis for
98 Research and Applications, version 2 (MERRA-2, Gelaro et al. 2017). That includes upper-level
99 geopotential heights and meridional (north-south) wind as well as 2-meter air temperature (T-2m
100 above the surface), available at $0.5^\circ \times 0.625^\circ$ regular latitude by longitude grids.

101

102 We used the 300 hPa pressure level for geopotential heights and winds to capture the quasi-stationary
103 Rossby waves, although similar wave patterns were also apparent at 500 hPa (Holton et al. 2003).
104 These waves appear as a series of troughs and ridges looping around the globe with typical zonal wave
105 numbers of 4-6. Rossby waves, and in particular, tropically forced Rossby waves (Hoskins and

106 Karoly, 1981), play an important role in modulating mid-latitude weather at subseasonal to seasonal
107 time scales. Since these waves tend to be barotropic (do not vary in the vertical) in middle latitudes,
108 their impacts extend down to the near-surface meteorological fields, including T-2m (e.g., Schubert et
109 al. 2011), which we have used as a proxy for temperature variability within the layer that most of the
110 migration occurs (up to ~1500 m AGL). Pressure level corresponding to the top of this layer varies
111 across CONUS due to the east-west topographic contrast. For this reason, we have verified for the
112 western and eastern CONUS separately, that temperature patterns remain vertically uniform in the bird
113 migration layer, so that we would be able to use T-2m to represent that layer.

114

115 For composite analysis—focusing on select extreme years—the meteorological variables were
116 averaged over April and May to represent the peak cumulative flow of migratory birds. For correlation
117 analysis, we used the entire spring migration season (March-April-May, MAM). A regional mean
118 time-series was generated for the western and eastern sectors of the U.S., separated at 102 W
119 longitude, using an objective regionalization approach discussed in Section 4. Anomaly time-series
120 (subtracting the mean) were used for comparing the data that have the same unit such as in Fig. 1b.
121 When data with different units were compared, such as in Fig. 3, each time-series was standardized by
122 subtracting the mean and dividing by the standard deviation. In either case the time-series were linearly
123 detrended to focus on interannual variability. Note that because of sporadic data coverage in space and
124 time, the time-series are normalized (or standardized) based on the mean and standard deviation of the
125 period on which they are presented, e.g., 2004-2018 for Fig. 1b.

126

127 In addition to regional air temperature, seasonal time-series of the normalized difference vegetation
128 index (NDVI) and various climate indices were correlated with q50. The p-value for each correlation
129 coefficient was then adjusted using the false discovery rate method (Benjamini and Hochberg 1995).

130 These indices that represent different modes of climate variability over the Pacific and Atlantic Oceans
131 include: Niño 3.4, Pacific North American index (PNA), East Pacific/North Pacific Oscillation
132 (EP/NP), North Pacific pattern (NP), Pacific Decadal Oscillation (PDO), North Atlantic Oscillation
133 (NAO), Arctic Oscillation (AO), North Tropical Atlantic index (NTA), and Atlantic Meridional Mode
134 (AMM). The climate indices data were obtained from NOAA’s Physical Science Laboratory
135 (<https://psl.noaa.gov/data/climateindices/list/>). Monthly NDVI was used from the Moderate Resolution
136 Imaging Spectroradiometer (MODIS) collection 6 product (MOD13C2), available at
137 <https://modis.gsfc.nasa.gov/data/dataproduct/mod13.php>.

138

139 **4. Regionalization based on bird migration**

140 Regionalization is a common practice for climate variability analysis (Fovell and Fovell 1993; Comrie
141 and Glenn 1998; Dezfuli 2011; Dezfuli and Nicholson 2013). However, to the best of our knowledge,
142 this is the first study to perform an objective regionalization based on interannual variability of bird
143 migratory phenology at the continental-scale. The process involved multiple steps and quality control
144 measures to ensure the robustness of the spatio-temporal patterns and properly address the issues
145 arising from the gaps and intrinsic noise in migratory data. Those efforts have resulted in excluding
146 stations with a large number of missing data as well as those with noisy behavior that are most likely
147 dominated by local characteristics.

148

149 In preparing the data, we first identified and at this initial stage eliminated the years in which more
150 than half the stations had missing data. A second filter was applied to keep only stations that had q50
151 observations over all those years. These restricting criteria were imposed to meet the minimum
152 requirements for a first estimate of regionalization and provided a 35 (stations) by 21 (years) matrix
153 used in the regionalization model, *HiClimR* (Badr et al. 2015). This is an open-source tool that uses

154 hierarchical clustering to regionalize any number of spatial points such as radar stations into
155 homogeneous regions with respect to similarity of their temporal variability. Note that the 21 years
156 used in the initial stage may not necessarily represent a continuous period. This step of the analysis
157 tried to maximize the number of years, so that the temporal similarity between stations would be
158 meaningful. It aimed to simultaneously maintain a minimum number of stations that would provide a
159 reasonable representation of the spatial variability. This effort would inform us about the optimum
160 number of regions and the longitudes at which they should be separated, therefore the 35 by 21 matrix
161 was not used to generate regional time-series. The results at this stage are used as a guideline and
162 suggested an optimum number of two regions, separated at about 102 W longitude. Using these two
163 pieces of information, we modified the preliminary results in order to address the known intrinsic
164 shortcoming of hierarchical algorithms that may result in removing or reassigning inconsistent
165 members. In addition, applying those assumptions to the q50 data allowed for larger spatial coverage
166 and maintained temporal continuity of the regional mean q50.

167

168 Consequently, 2004-2018 was chosen as a period over which most stations (121 of 143) had
169 continuous observations. Two regional time-series were created, by averaging standardized q50
170 anomalies of all stations located to the west and east of 102 W, respectively. Pearson correlation
171 coefficient between each regional time-series and all its individual members were calculated. The
172 stations with a correlation coefficient less than 0.4 (an arbitrary value corresponding to $p < .14$) were
173 flagged as noise. Modified regional time-series were calculated after removing those stations, so that
174 they would represent the large-scale spatial signal in bird migration phenology. The regional time-
175 series were then detrended to focus on the interannual variability. The western and eastern regions
176 consisted of 28 and 38 stations, respectively.

177

178 We evaluated the regionalization performance using intra-regional and inter-regional correlations
179 (Dezfuli 2011; Badr et al. 2015; Badr et al. 2016). A high value of “intra-regional”, defined as the
180 mean correlation between each regional time-series and its members, assures homogeneity of the
181 regions. A low value of “inter-regional”, defined as the correlation between regional mean time-series,
182 satisfies separability of the regions. Both these criteria were simultaneously met in our regionalization
183 (Fig. 1a), as shown in the high intra-regional correlations for western ($R_w = 0.57$) and eastern ($R_E =$
184 0.62) regions as well as in the low value for inter-regional correlation ($R_{W-E} = -0.04$).

185

186 It is important to emphasize that we have used anomalies rather than absolute values of q50 because
187 we are interested in regional interannual variability. Using anomalies would allow equal contribution
188 from all stations to the regional means. Therefore, areal average time-series would represent the entire
189 region and are not biased toward stations with higher q50 values located in the northern latitudes. To
190 further elaborate on this approach, we have compared two arbitrary stations in the western region
191 (KMTX, 41.3N & 112.4W and KNKX, 32.9N & 117W). The correlation coefficient between their
192 time-series was 0.75 ($p < .005$), though they are ~1000 km apart and the mean q50 of the northern
193 stations is ~13 days higher. Similarly, the time-series of KGRB (44.5N & 88.1W) and KTLH (30.4N &
194 84.3W) in the eastern region—nearly 1600 km apart—are highly correlated ($R = 0.78$, $p < .0001$). It
195 is interesting that some stations in the western region (e.g., KDAX, 38.5N & 121.6W) are strongly
196 negatively correlated ($R = -.69$, $p < .005$) with other stations in the eastern region such as KTLH,
197 located ~3500 km away. However, this dipole does not seem to be a continental-scale characteristic
198 since R_{W-E} is nearly zero and therefore is not further investigated in this study.

199

200 We also tried the regionalization for 3 and 4 regions, but both were rejected as the separability criterion
201 was not achieved. At this stage the regionalization process is complete, and we next explored the

202 differences between the temporal characteristics of the two regions such as their interannual variability.
203 The standard deviation of regional mean time-series of q50 anomalies over the period 2004-2018
204 shows a relatively higher variability in the western region (2.4 days) than in its eastern counterpart (1.7
205 days). Using a two-tailed F-test, the difference between variance of the two regional time-series takes a
206 *p-value* less than 0.22.

207

208 Our two-region compartmentalization is intrinsically different from the previously used classifications,
209 which are based on three- and four-flyway strategies, both in how it has been achieved and its
210 applications. Our approach reflects the interannual variability in timing of bird arrival and therefore is
211 distinct from migratory corridors. We utilized an objective statistical approach to define the regions.
212 This work relies on the fact that variability of bird migration phenology can be divided into two
213 components, “noise” and “signal”. The noise part may be determined by factors such as local
214 environmental conditions, local geographical features and species-specific characteristics (Vardanis et
215 al. 2011; Somveille et al. 2019; Deppe et al. 2015; Youngflesh et al. 2021). On the other hand,
216 common behavioral factors among species as well as large-scale climatic phenomena would
217 collectively result in a spatio-temporal “signal” in interannual variability. We argue that our
218 regionalization approach, reflecting this coherent “spatial signal”, enables us to better identify the
219 drivers of interannual and potentially decadal variability of migration timing at the continental scale.
220 Here, we provide examples of large-scale impacts of climate conditions on bird migration, facilitated
221 by our regionalization. It is worth noting that the three-year running averages are only used to
222 qualitatively discuss the low-frequency variability in data. All quantitative analysis, including
223 regionalization, significance tests, and correlations incorporate unsmoothed time-series.

224

225 **5. Climate-migration association for the two regions**

226 Comparing the mean time-series of the two regions (Fig. 1b) allowed us to identify years with notably
227 east-west contrast in median passage date anomalies. That contrast was most evident in 2005 and 2010,
228 when the western and eastern sectors experienced considerably different median passage dates, with
229 the western region exhibiting an earlier date in 2005 and a later date in 2010. We attribute this zonal
230 (east-west) dipole pattern in q50, in part, to the near surface air temperature (Fig. 2a) and, to a lesser
231 degree, the meridional winds (Fig. 2b) during the peak migration months, April and May. The warmer
232 than normal temperatures and southerly anomalies over the western region in 2005 favor an earlier
233 arrival than in 2010. The opposite pattern is apparent for the eastern region. The strong linkage with
234 temperature is likely due to the fact that temperature serves as a surrogate for resources (Studds and
235 Marra 2011; Van Doren and Horton 2018). We speculate that the winds at the height of migrating birds
236 that are linked to the gradient of surface temperature via thermal wind balance may play a secondary
237 role. This zonal configuration of temperature and meridional winds resembles a pattern that is
238 consistent with that of a quasi-stationary atmospheric Rossby wave. The spatial structure of
239 geopotential heights captures the areas of high- and low-pressure anomalies, associated with the wave
240 (Fig. 2a). This anomaly pattern over the U.S. is part of a wave train that extends from the central North
241 Pacific into the North Atlantic (Fig. 2b); it was especially prominent during 2010. The effect of the
242 waves—likely triggered by sea surface temperature (SST) anomalies over the Pacific Ocean—is
243 reflected at the lower troposphere through downward penetration of potential vorticity.

244

245 Another capability of our regionalization approach is that it enables us to identify variability patterns
246 specific to each region and their associated controlling factors. One advantage of objective
247 regionalization is that once the borders are determined, the regions are assumed homogeneous and
248 therefore the time-series can be extended over the years that were excluded from the original
249 regionalization due to the low number of sites with data available. This advantage allowed us to extend

250 the time-series of q50 over 1996-2018, recognizing the potential uncertainties and errors, arising from
251 using a smaller number of stations for the years prior to 2004. However, we have computed correlation
252 coefficient between q50 and various climate indices for both periods (Table 1).

253

254 The western region shows a significant negative correlation with T-2m averaged over the same area
255 ($R = -0.79$) for 1996-2018. One noticeable pattern in q50 of this region is its low-frequency
256 variability that is also apparent in the regional T-2m (Fig. 3) and PDO (not shown), where
257 positive/negative phases of PDO are coincident with early/late arrival dates. However, we recognize
258 that the period of this analysis is not sufficiently long to confidently support this link, which can be
259 considered as a viable hypothesis for further investigation when data becomes available. In contrast, a
260 low-frequency pattern is not evident over the eastern region, and q50 over this area shows a weaker
261 interannual association, though statistically significant, with its regional mean temperature ($R =$
262 -0.56). This different magnitude of response to temperature is intriguing because CONUS can be
263 divided into two homogeneous regions with respect to interannual variability of spring temperature
264 (MAM), and the separating longitude is roughly the same as that of the regions based on q50 (Fig. 4).
265 The regionalization was objectively performed with *HiClimR* package, using seasonal T-2m gridded
266 data from MERRA-2. The two-region classification was obtained from simultaneous minimization of
267 inter-regional and maximization of intra-regional correlations. In addition, this division closely
268 corresponds to differences in patterns of greenness and habitat between eastern and western CONUS
269 (White et al. 2005). Interannual variability of q50 in the western region presents a strong negative
270 correlation ($R = -0.50$) with NDVI—unlike the eastern region ($R = -0.12$, Table 1). The latter low
271 correlation may be attributed to several factors including heterogeneity of interannual variability of
272 greenness within the eastern region and species and latitudinal dependencies on vegetation patterns
273 (Mayor et al. 2017; Youngflesh et al. 2021).

274

275 Although the eastern region shows relatively weaker association with regional variability, its link to
276 teleconnection patterns is stronger than that of the western region (Table 1). The highest correlations
277 are with EP/NP ($R = 0.58$), NTA ($R = 0.52$) and AO ($R = -0.50$) indices. To assess the extent to
278 which these climate phenomena manifest the impact of ocean variability on bird migration, we
279 evaluated the spatial correlation between q50 of the eastern region and large-scale SSTs (Fig. 5a). The
280 spatial patterns of EP/NP and NTA can be particularly identified from the regions with significant
281 correlations, although the North Pacific correlations may also resemble the PDO structure. Analysis of
282 spatial correlation between 300-hPa geopotential heights and q50 shows that the impact of SST is
283 likely reflected through Rossby waves that are excited over the tropical Pacific (Fig. 5b). These waves
284 are often associated with the North American ridge-trough dipole that controls the temperature over the
285 eastern CONUS. Although the dipole is commonly known for its influence on boreal winter
286 temperature (Wang et al. 2014; Singh et al. 2016; Schulte et al. 2018), it is also present during spring
287 (Schulte and Lee 2017).

288

289 The western region, on the other hand, shows strong correlation only with geopotential heights over
290 the same area and SSTs along the West Coast of North America (Fig. 6). The negative correlations
291 with SST imply that the adjacent waters likely affect the region through temperature advection. A
292 Rossby wave train originating from the tropical Pacific is also apparent (Fig. 6b), but it is much weaker
293 and more spatially limited than the one shown for the eastern region. Additional climate modes were
294 examined but the results were not included in Table 1 because they were either not statistically
295 significant (e.g., PNA) or considered redundant due to high co-variability with indices already
296 presented in the table. For example, NTA was highly correlated with AMM ($R = 0.9$), so was AO with
297 NAO ($R = 0.74$) and PDO with NP ($R = -0.7$). However, the climate modes shown in Table 1 would

298 adequately represent variability over both tropical and extratropical parts of the Pacific and Atlantic
299 Oceans.

300

301 **6. Discussions and future work**

302 Our analysis approach is different from previous studies of long-term changes, which have largely
303 focused on the trends of migration phenology; many do not consider year-to-year variability in these
304 dynamics. In contrast, our approach has incorporated detrended data to facilitate the study of
305 interannual variability and its drivers. As a byproduct, this strategy detects the years during which the
306 western and eastern U.S. present an opposite migratory behavior and attempts to explore climatic
307 processes responsible for such a dipole pattern.

308

309 Some differences were noticed between drivers of interannual variability of the western and eastern
310 regions. While the western region shows a strong link to the regional temperature, the eastern region
311 presents statistically significant relationships with several climate modes of variability including
312 atmospheric Rossby waves, which appear to be excited in the tropical Pacific Ocean. While some co-
313 variability may exist between these modes, some of them can act quite independently, suggesting that
314 bird migration is likely controlled by combined effect of these teleconnections. Such complex
315 interactions require further investigation. Also, we speculate that spatial variability of species
316 composition may partly contribute to different responses of the western and eastern regions to regional
317 climate conditions (La Sorte et al. 2014b; Horton et al. 2020). However, NEXRAD data is agnostic to
318 species composition, therefore long-term species-specific observations with high spatial resolution, for
319 example, from citizen-science would be crucial to address this question. Other potential future work
320 could focus on future projection of spring temperature variability mainly for the western region,
321 changes in teleconnections affecting the eastern region, and seasonal prediction skill of atmospheric

322 phenomena, such as Rossby waves, that influence the migration system. The new spatial framework
323 presented here would facilitate such follow-up studies.

324

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335

336 **Data availability**

337 The bird migration data sets are available at <https://doi.org/10.6084/m9.figshare.10062239.v1>.
338 MERRA-2 products are publicly available at <https://disc.gsfc.nasa.gov> with
339 <https://doi.org/10.5067/KVIMOMCUO83U> for T-2m and <https://doi.org/10.5067/2E096JV59PK7> for
340 upper levels. PDO time-series can be found at <http://research.jisao.washington.edu/pdo/PDO.latest>.

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470 **Table 1.** Pearson correlation coefficient between regional mean q50 and seasonal (March-April-May)
 471 mean of various climate indices. Calculations are made for both 2004-2018 (n= 15) with minimum
 472 missing data and the extended period 1996-2018 (n= 23). Corresponding p-values, adjusted with the
 473 false discovery rate method (Benjamini and Hochberg 1995) are also provided (in parentheses). Only
 474 adjusted p-values close to or less than 0.1 are shown (in bold).

Climate Index	West/East	2004-2018	1996-2018 ⁴⁷⁵
T-2m	W	-.83 (.0009)	-.79 (.00005)
	E	-.55 (.08)	-.56 (.02)
NDVI	W	-.63 (.04)	-.50 (.1)*
	E	-.17	-.12
Nino3.4	W	-.33	-.24
	E	.28	.30
EP/NP	W	-.37	-.17
	E	.55 (.08)	.58 (.02)
PDO	W	-.35	-.29
	E	.31	.42 (.06)
AO	W	-.02	.00
	E	-.60 (.08)	-.50 (.03)
NTA	W	.34	.21
	E	.49 (.11)	.52 (.03)

476 * For 2000-2018.

477

478 **Figure Captions**

479 **Fig. 1** (a) Two regions identified based on interannual variability of peak bird migration date (q50) in
480 spring. Circles show the location of NEXRAD stations in each region. (b) Regional mean time-series
481 of the two regions. Time-series are detrended anomalies. Years with notably west-east contrast in q50
482 anomalies are marked with open circles.

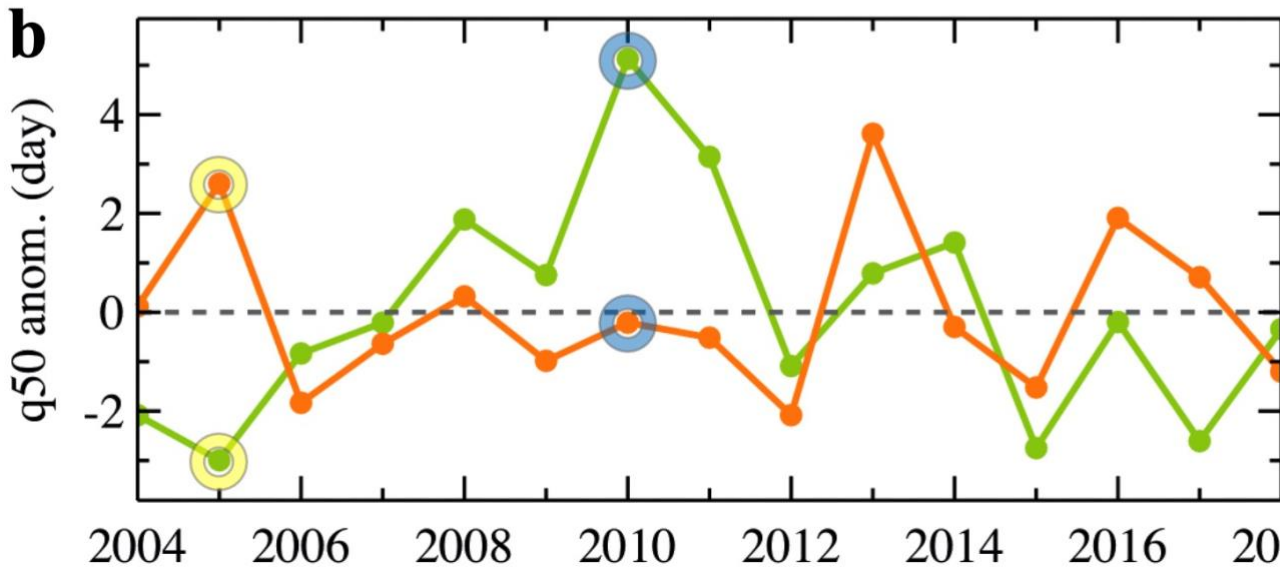
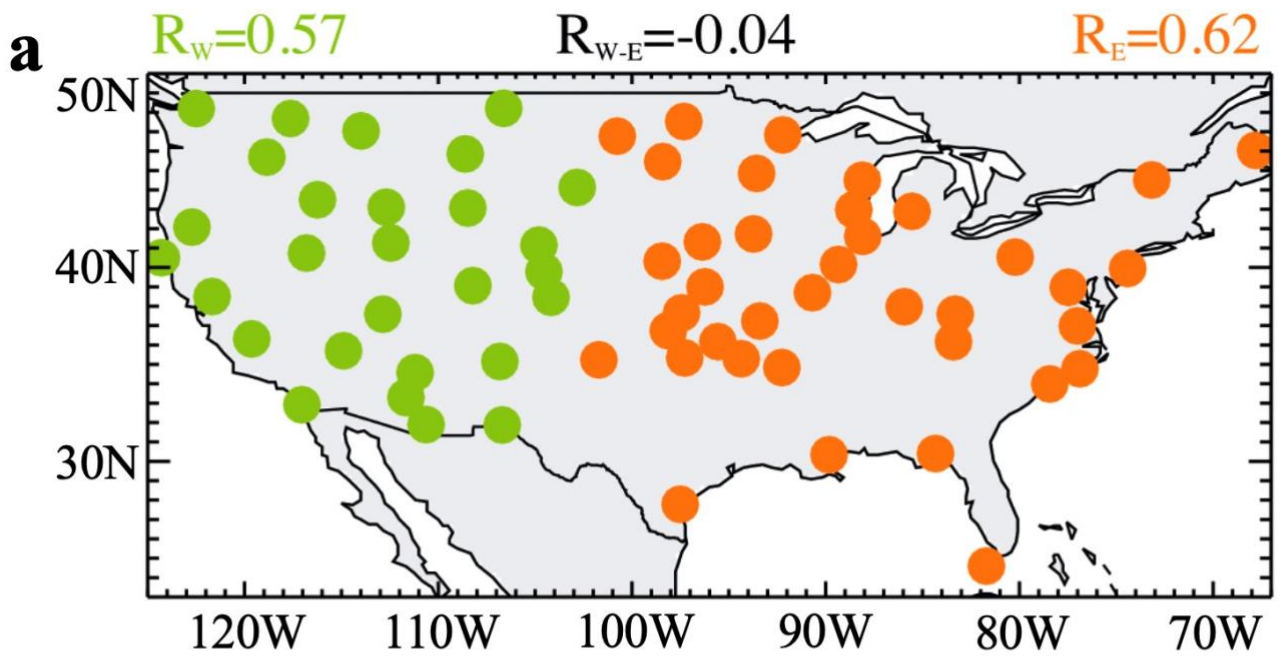
483 **Fig. 2** (a) T-2m (shading) and geopotential heights at 300 hPa level during April/May (blue and red
484 contour lines) for 2005 minus 2010. (b) The same as (a) but for 300-hPa meridional wind (shading)
485 over a longitudinally extended area to capture the Rossby wave train. Regions with high and low
486 pressure anomalies are labeled with H and L, respectively.

487 **Fig. 3** Three-year running average of spring q50 and T-2m seasonal mean (Mar-Apr-May) over the
488 western region. Time-series are standardized and detrended for better comparison of variables with
489 different units.

490 **Fig. 4** Climate regions obtained objectively based on similarity in interannual variability of Mar-Apr-
491 May T-2m (shading). Location of the stations for the two regions identified based on interannual
492 variability of peak migration date (q50) are superimposed for comparison.

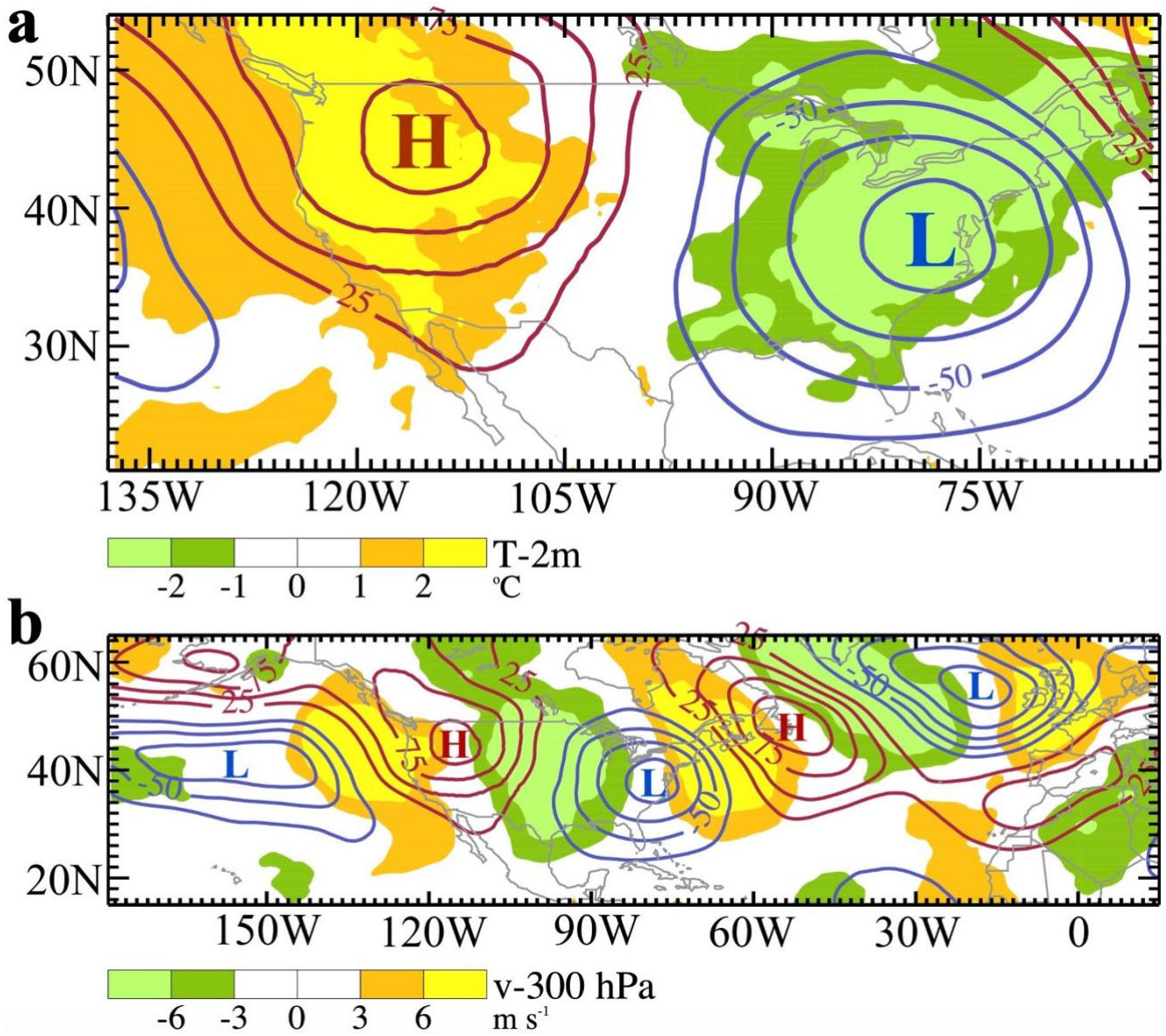
493 **Fig. 5** Correlation patterns between regional q50 time-series of the eastern region and the large-scale
494 (a) SST and (b) 300-hPa geopotential heights for the eastern region. All time-series are seasonal means
495 (Mar-Apr-May) for 1996-2018. Black dots show areas with correlation coefficient significant at 10%
496 level.

497 **Fig. 6** The same as Fig. 5 but for the western region.



498

499 **Fig. 1** (a) Two regions identified based on interannual variability of peak bird migration date (q50) in
 500 spring. Circles show the location of NEXRAD stations in each region. (b) Regional mean time-series
 501 of the two regions. Time-series are detrended anomalies. Years with notably west-east contrast in q50
 502 anomalies are marked with open circles.



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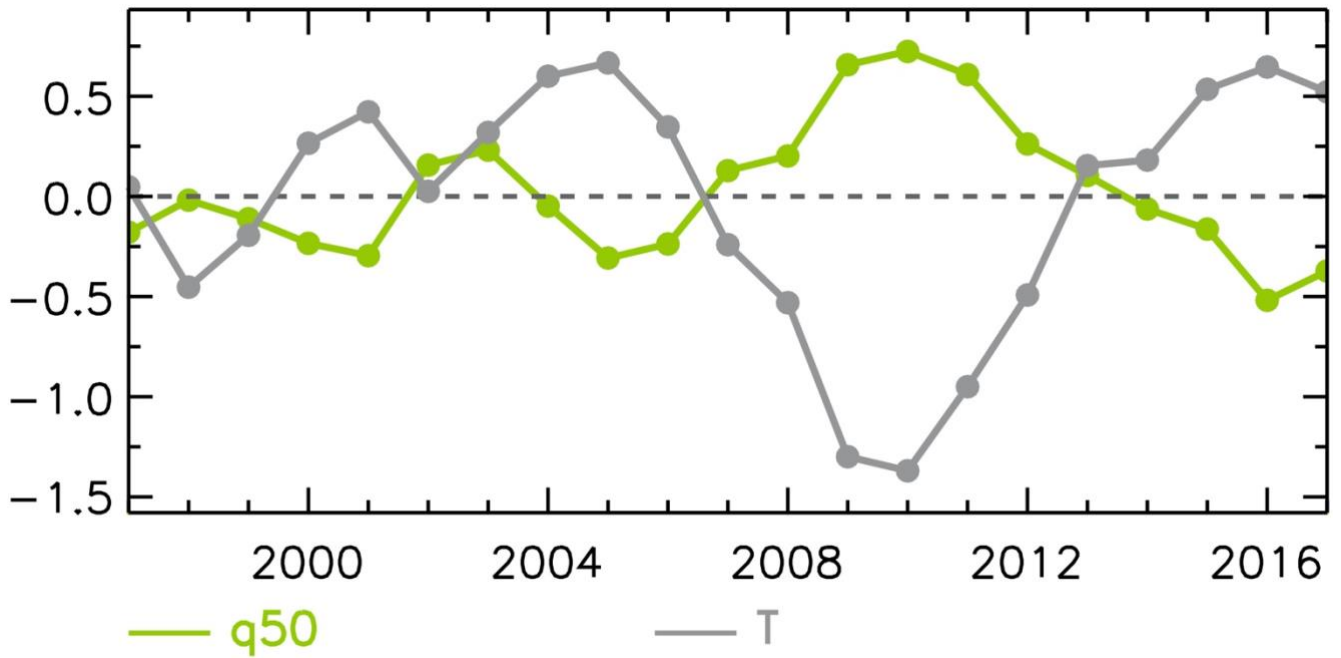
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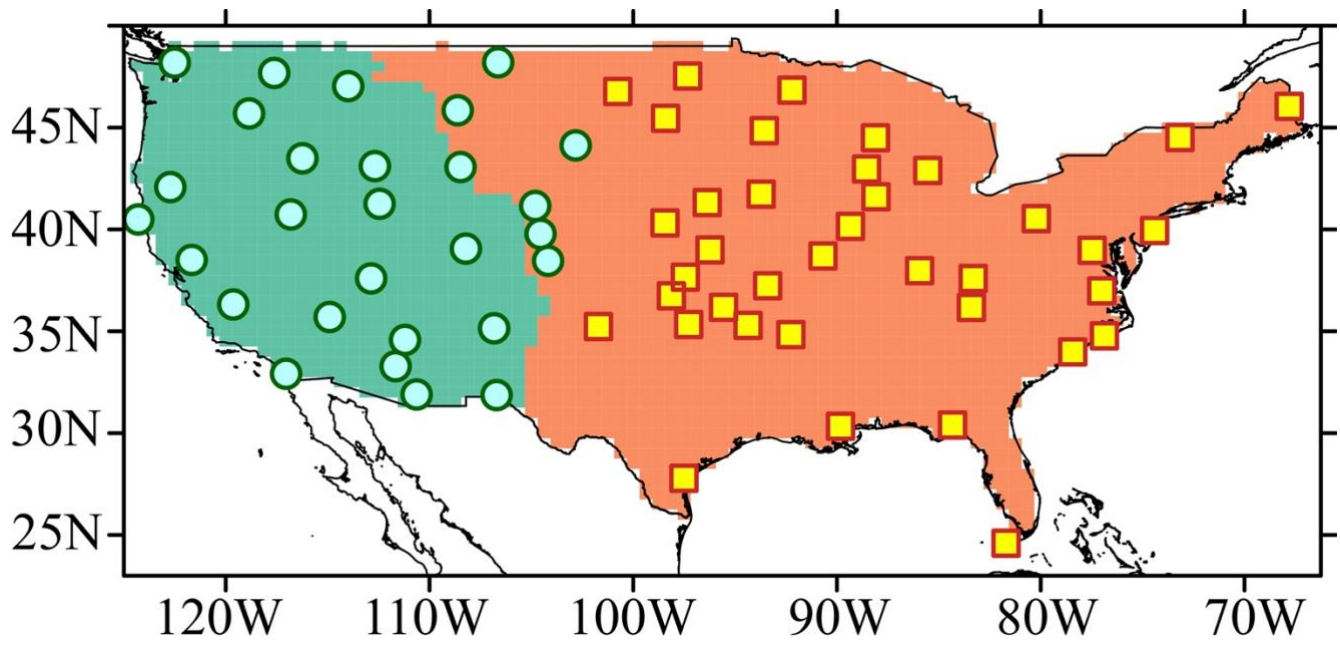
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509

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 511 western region. Time-series are standardized and detrended for better comparison of variables with
 512 different units.



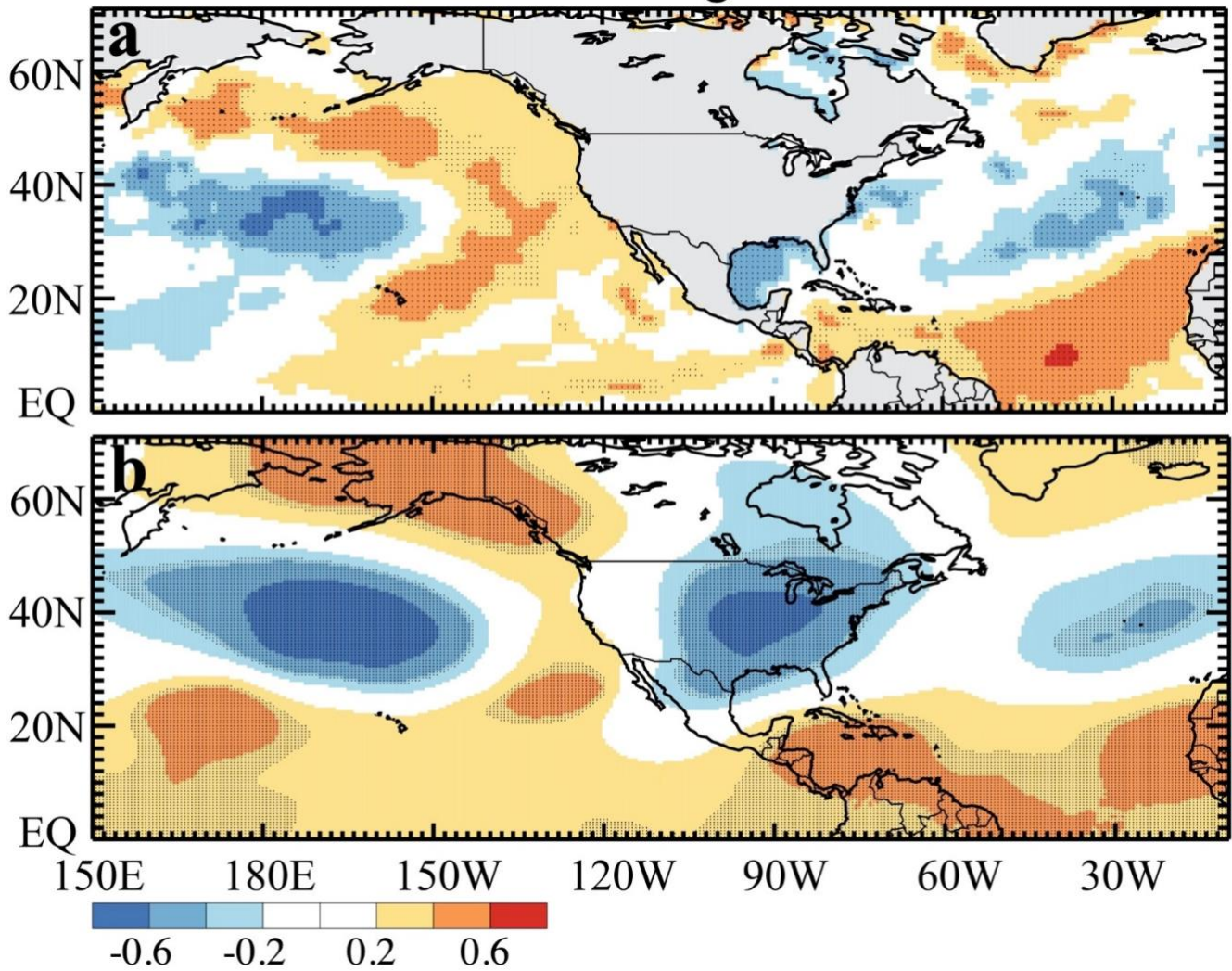
513

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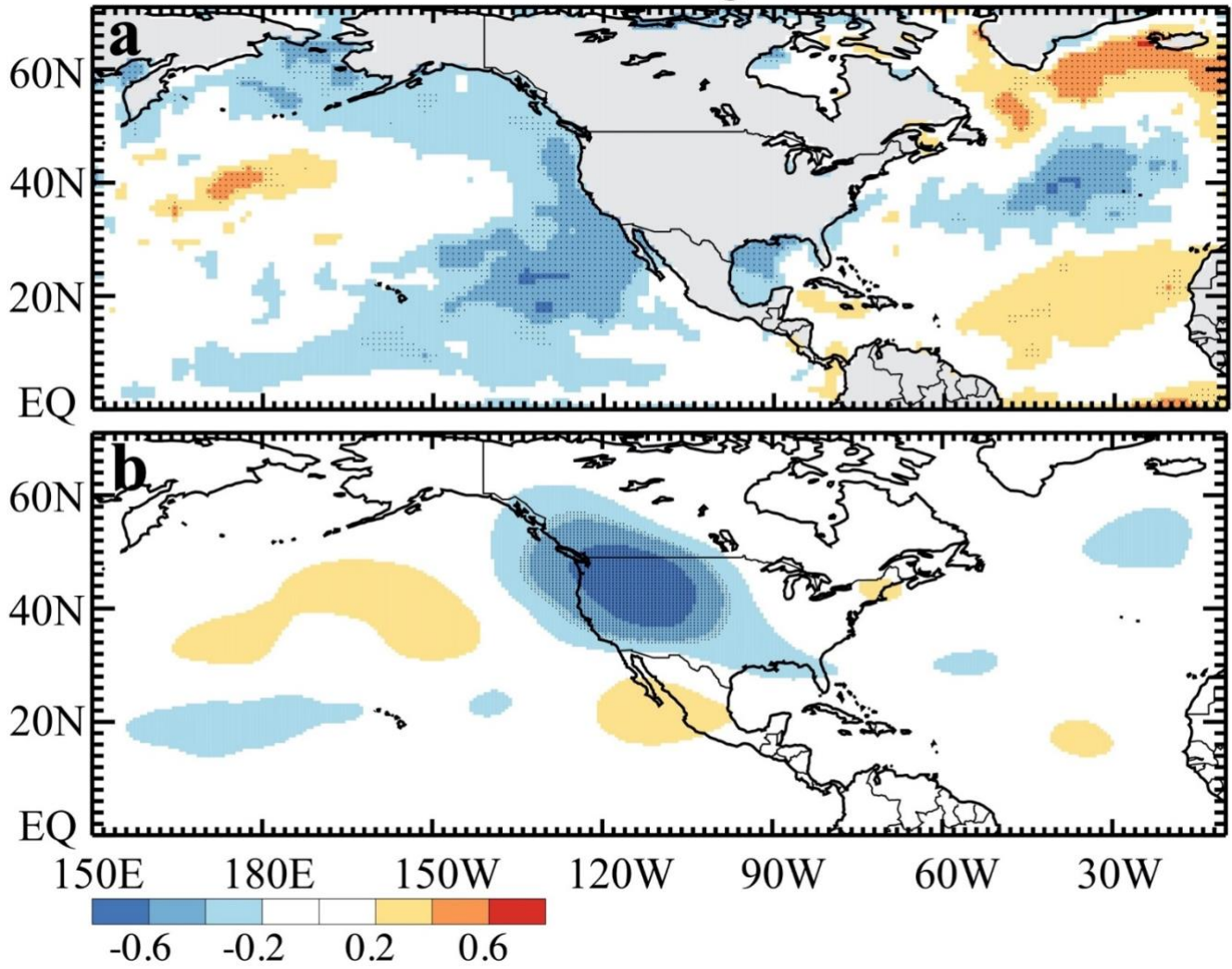
Eastern Region



517

518 **Fig. 5** Correlation patterns between regional q50 time-series of the eastern region and the large-scale
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520 (Mar-Apr-May) for 1996-2018. Black dots show areas with correlation coefficient significant at 10%
521 level.

Western Region



522

523 **Fig. 6** The same as Fig. 5 but for the western region.