



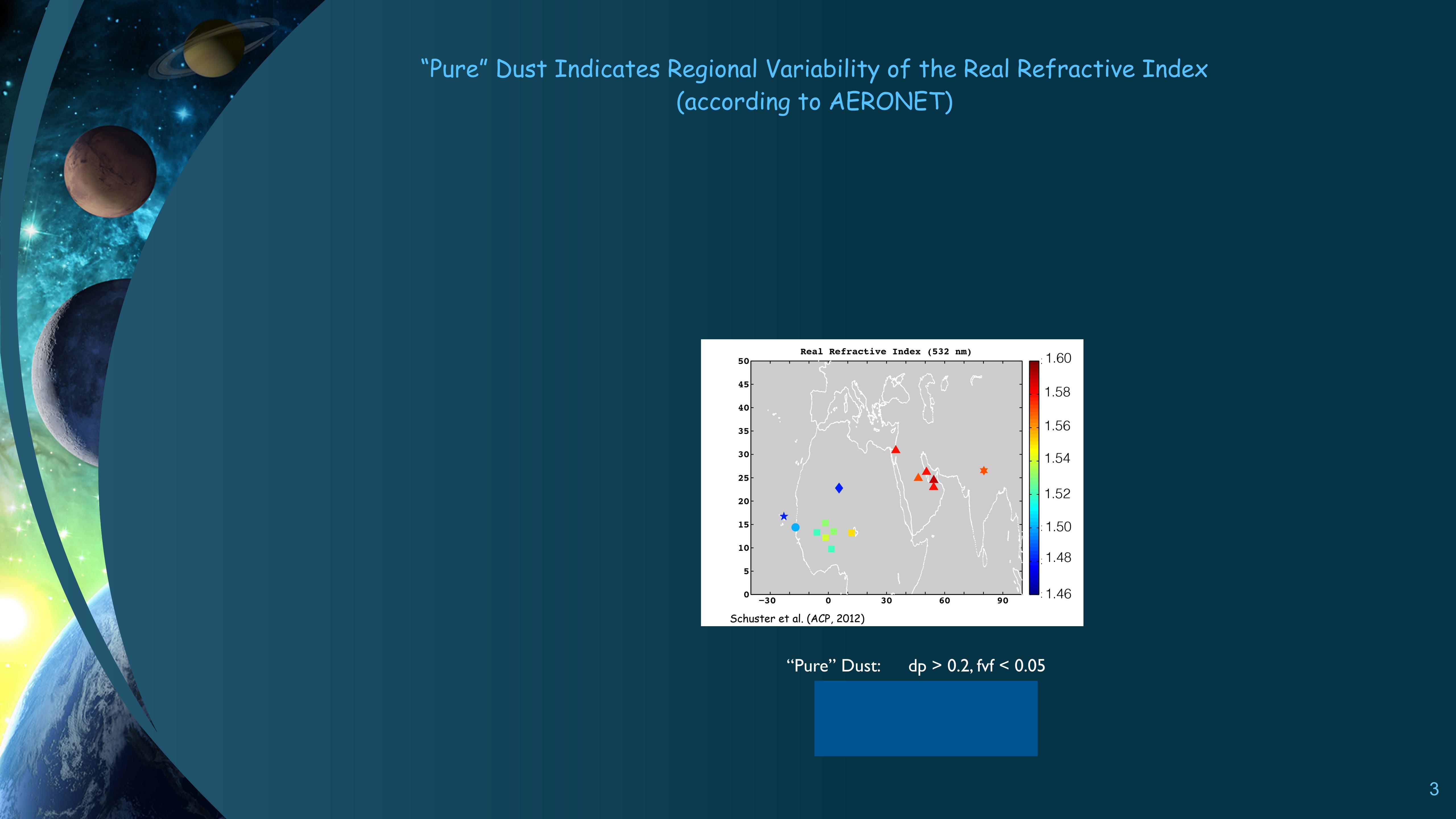
Using a Transport Model for Source-Dependent Lidar Ratios of Dust

Greg Schuster, NASA LaRC; Dongchul Kim, NASA GSFC, USRA;
Zhaoyan Liu, NASA LaRC; Mian Chin, NASA GSFC; Kerstin Schepanski, TROPOS.

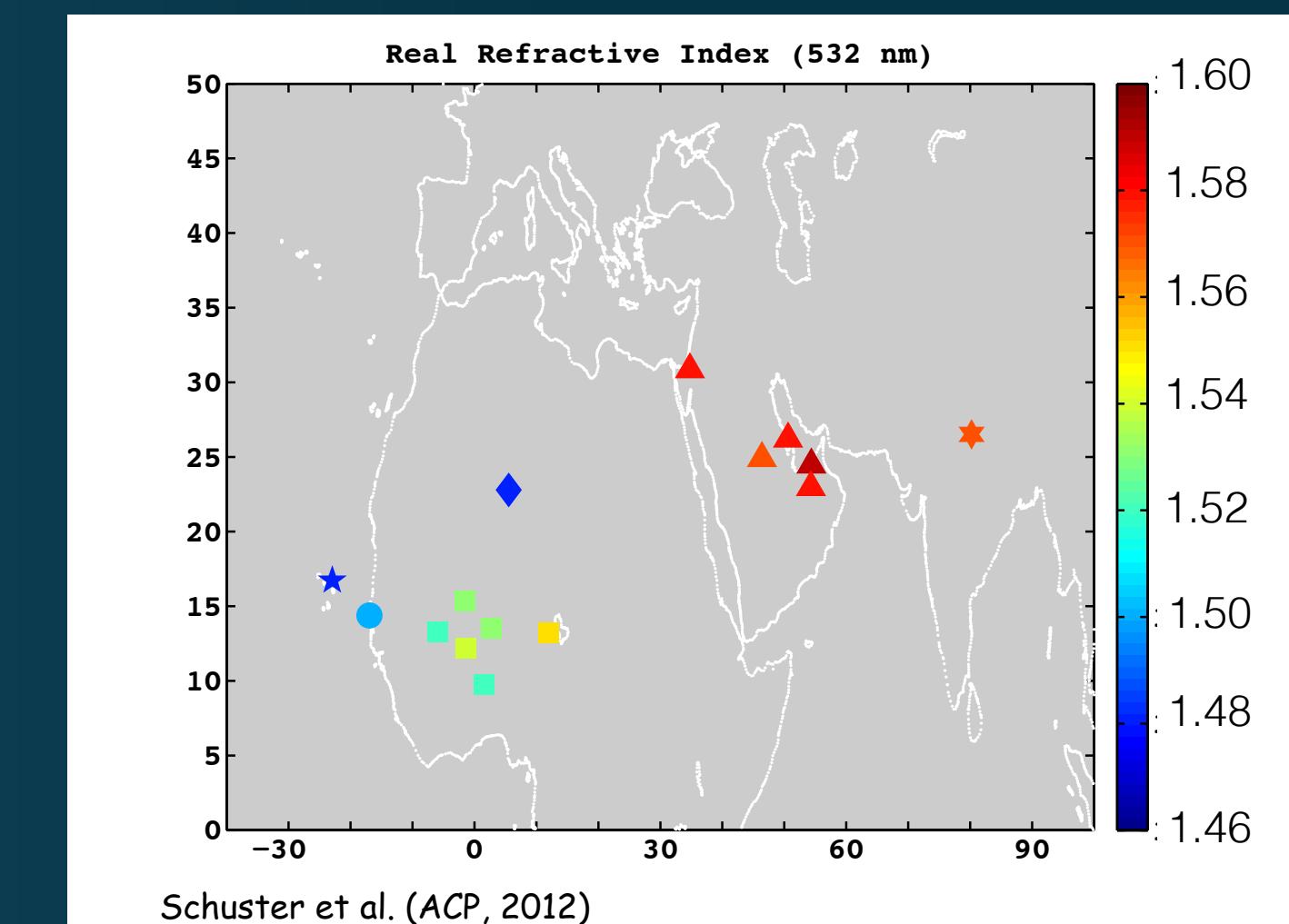


Motivation

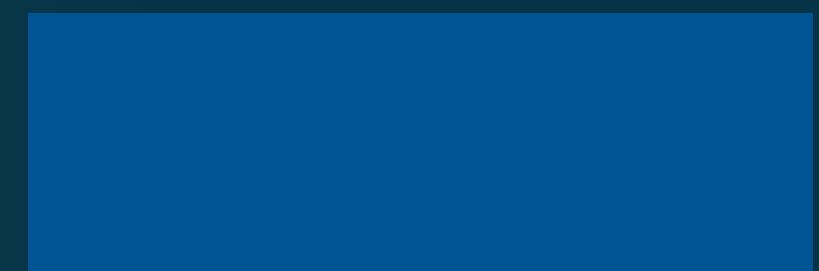
- CALIPSO/CALIOP extinction retrievals require an extinction-to-backscatter ratio assumption (i.e., a lidar ratio assumption).
- Presently, CALIOP uses a single lidar ratio for dust (45 sr).
- However, the lidar ratio of dust is sensitive to mineralogy.
- Dust single-scatter albedo
- Dust aerosol optical depth (AOD)

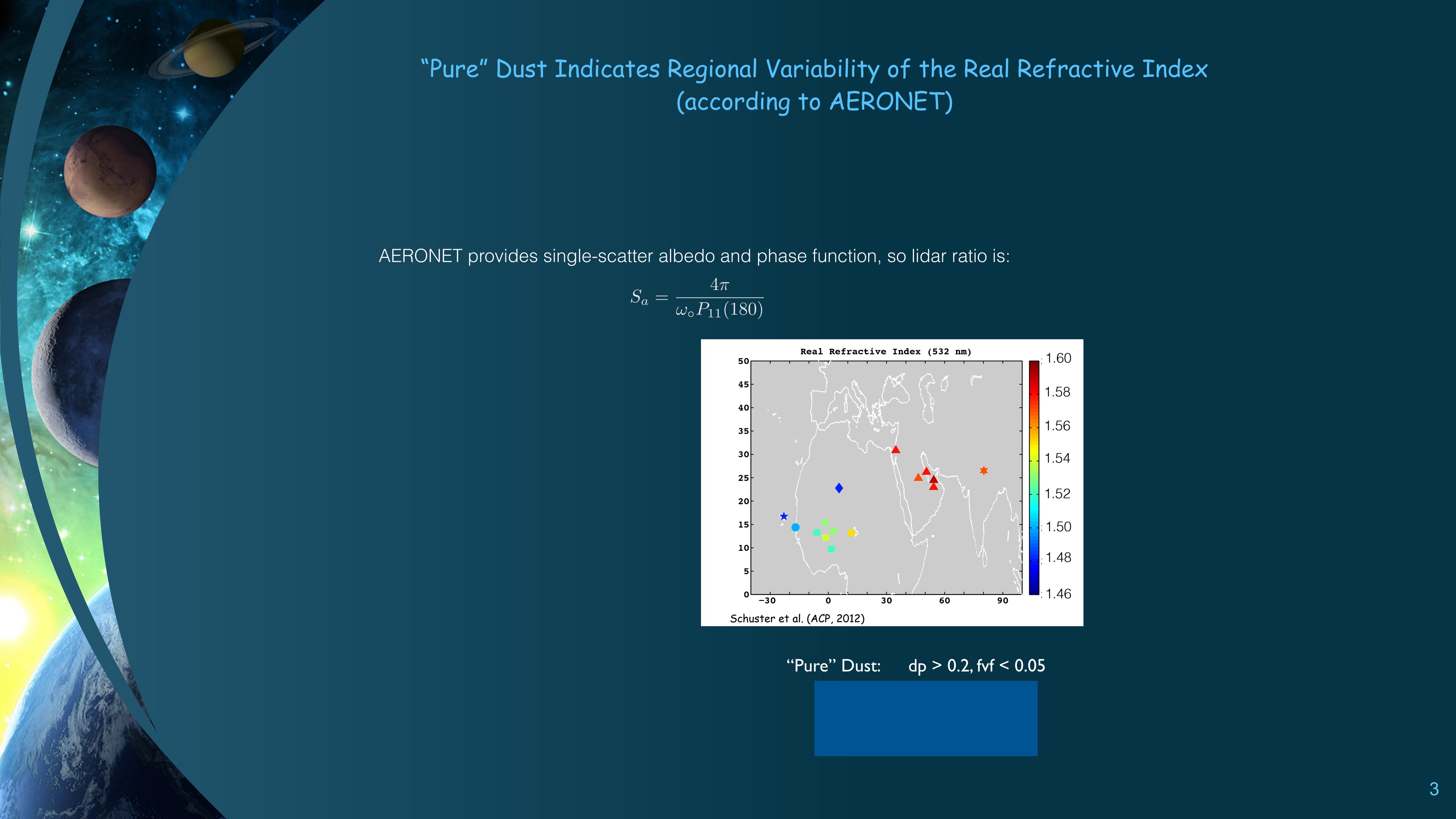


"Pure" Dust Indicates Regional Variability of the Real Refractive Index (according to AERONET)



"Pure" Dust: $d_p > 0.2, fvf < 0.05$

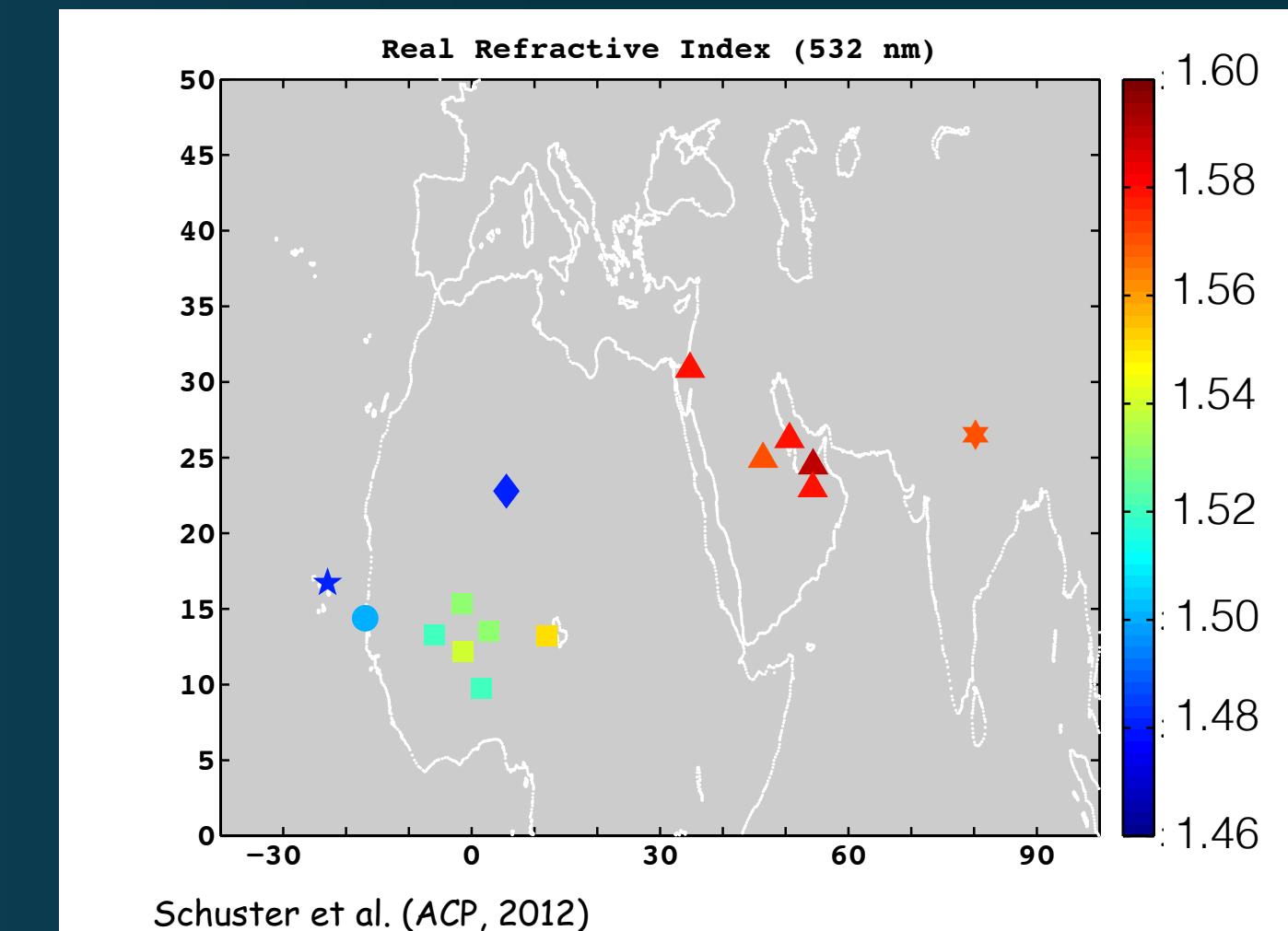




"Pure" Dust Indicates Regional Variability of the Real Refractive Index (according to AERONET)

AERONET provides single-scatter albedo and phase function, so lidar ratio is:

$$S_a = \frac{4\pi}{\omega_o P_{11}(180)}$$



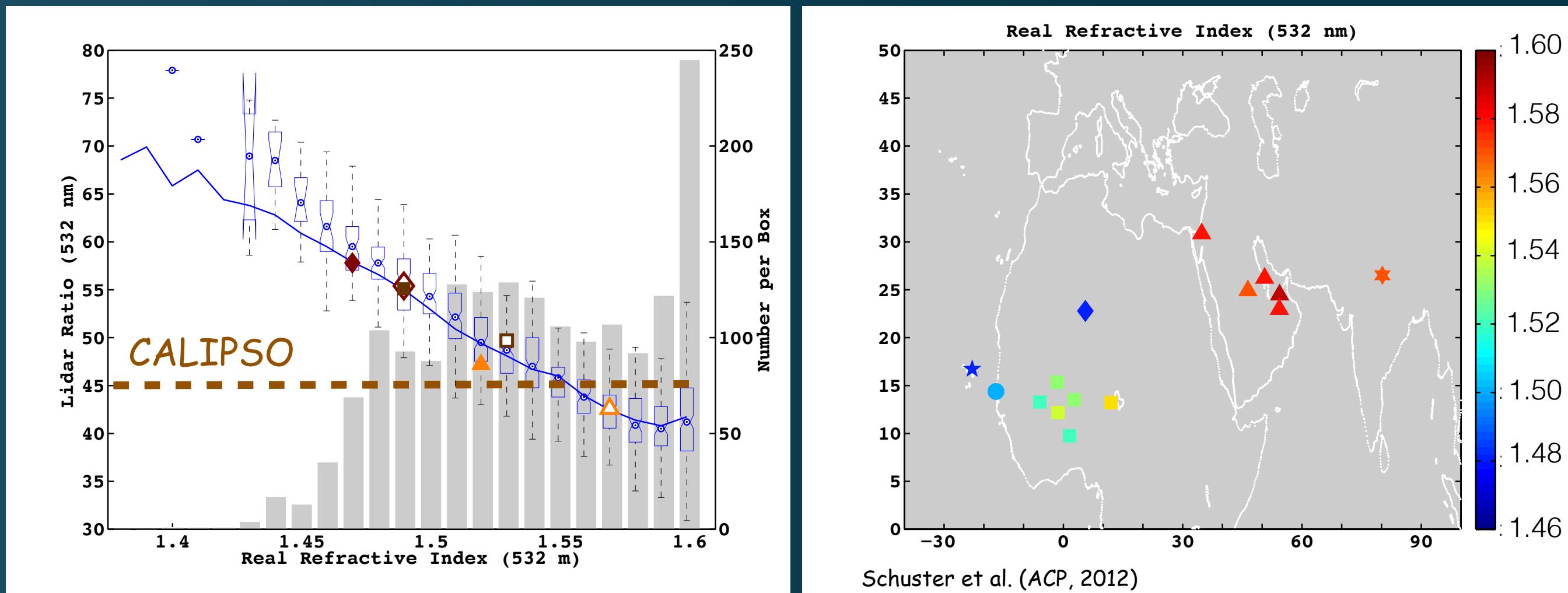
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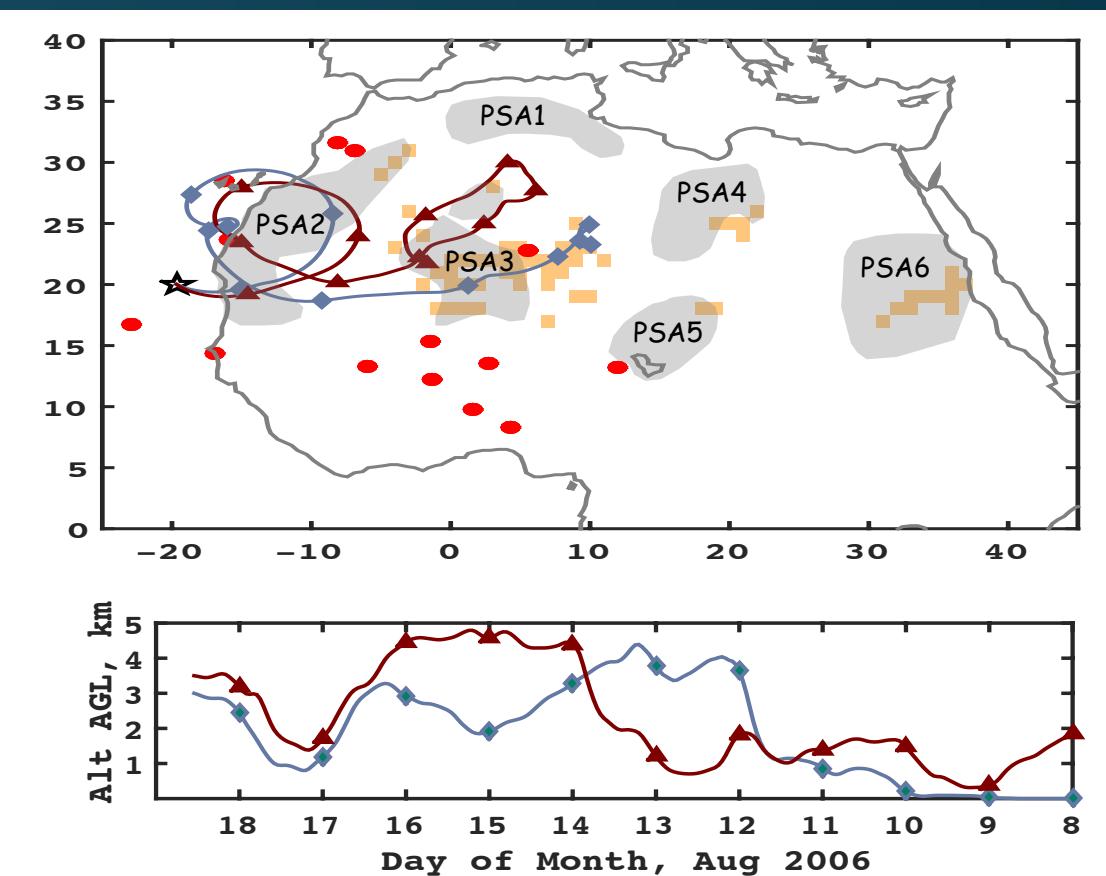


Correlation	“Pure” Dust	Dust
North Africa	-0.85	-0.79
Middle East	-0.68	-0.81

“Pure” Dust: $dp > 0.2, fvf < 0.05$
Dust: $dp > 0.2$

How can we Link Measurements to Potential Source Areas?

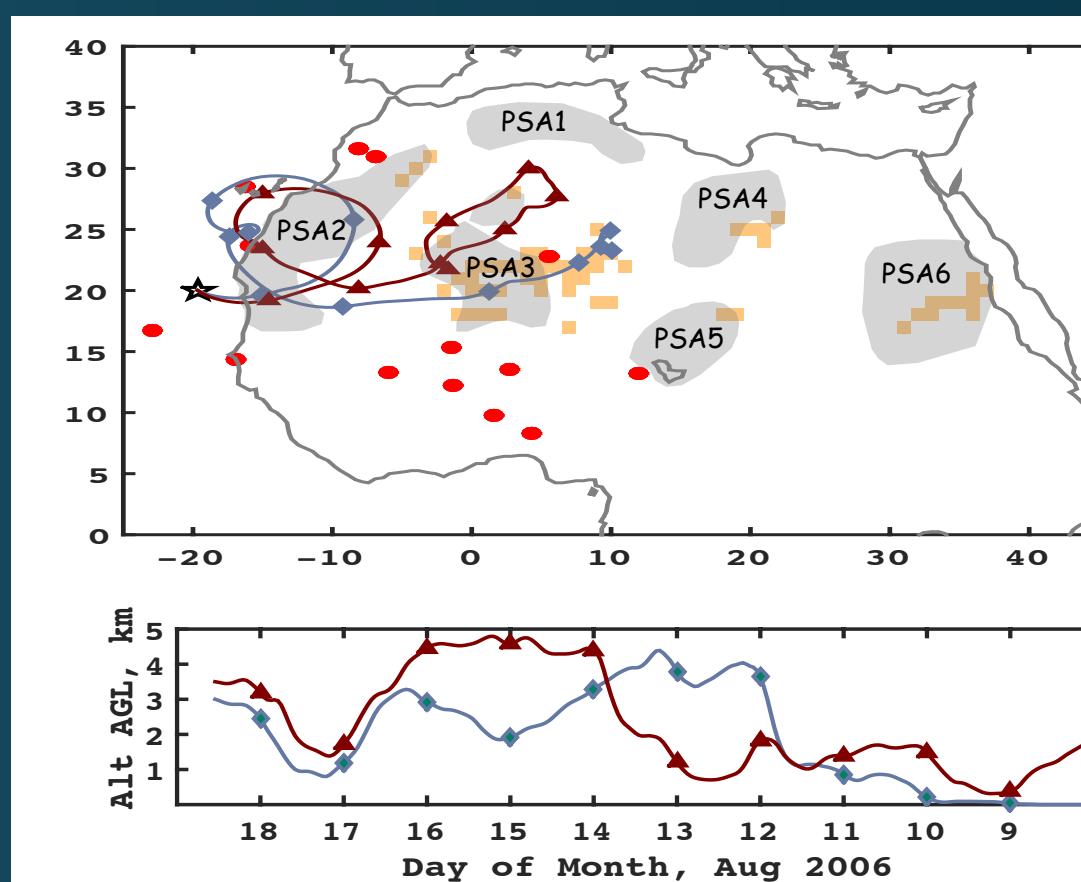
- Backtrajectories can not quantify the proportion of dust originating from different Potential Source Areas (PSAs).



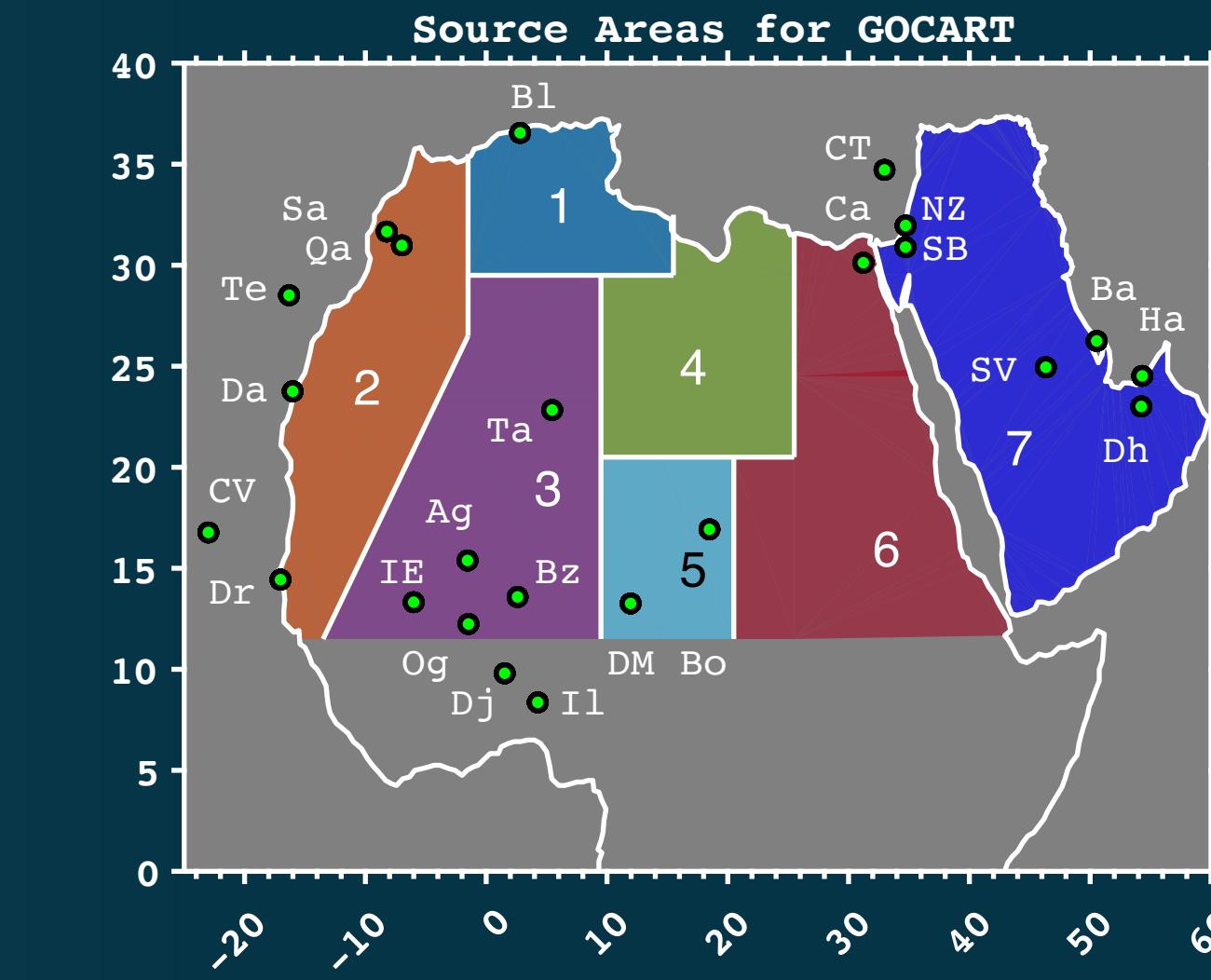
PSA regions from Formenti (ACP, 2011)

How can we Link Measurements to Potential Source Areas?

- Backtrajectories can not quantify the proportion of dust originating from different Potential Source Areas (PSAs).
- Alternative Approach: "Tag" source regions in a transport model (*GOCART*).



PSA regions from Formenti (ACP, 2011)



Quantify the Origin of Aeolian Dust Collected over the Atlantic

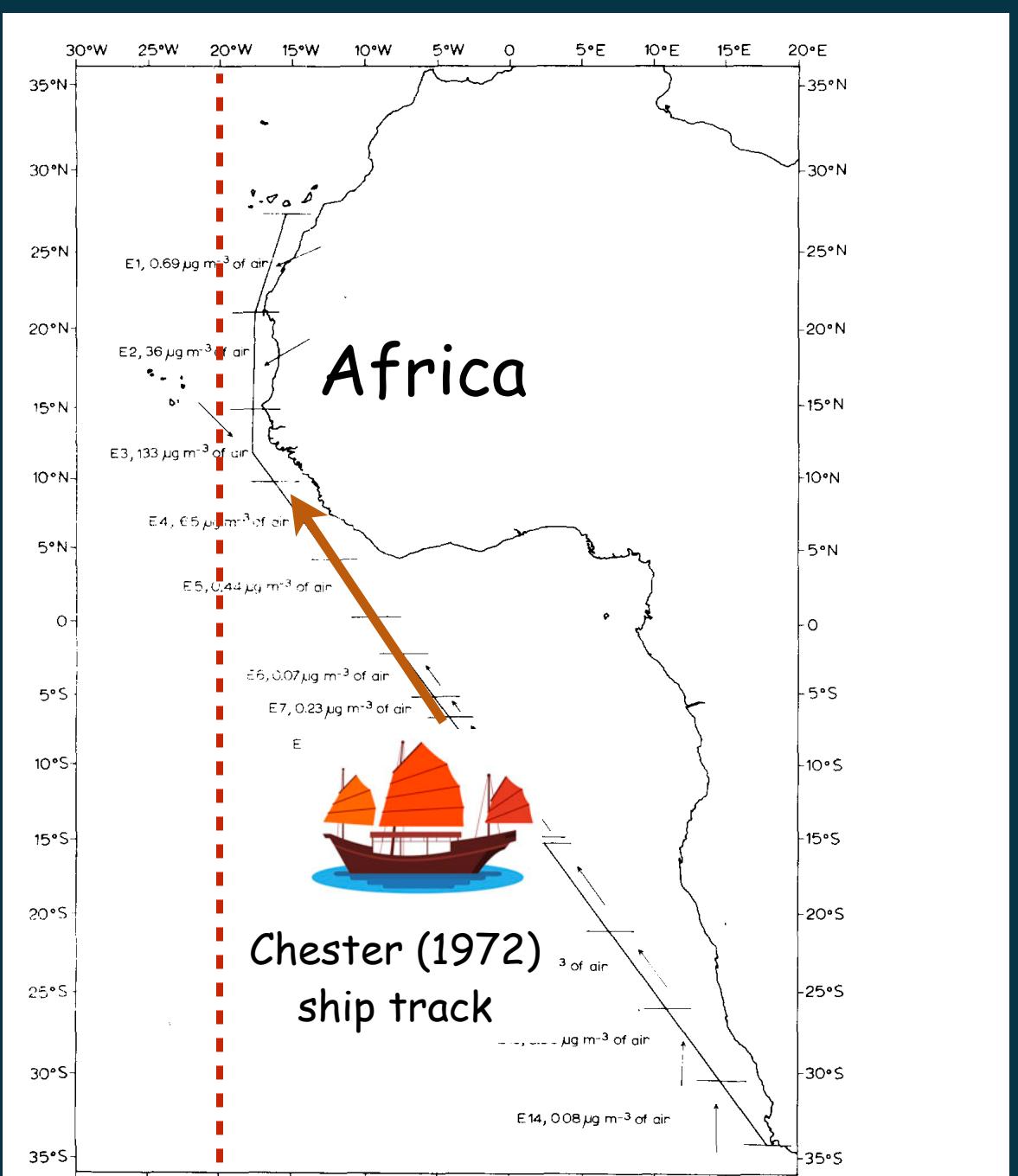
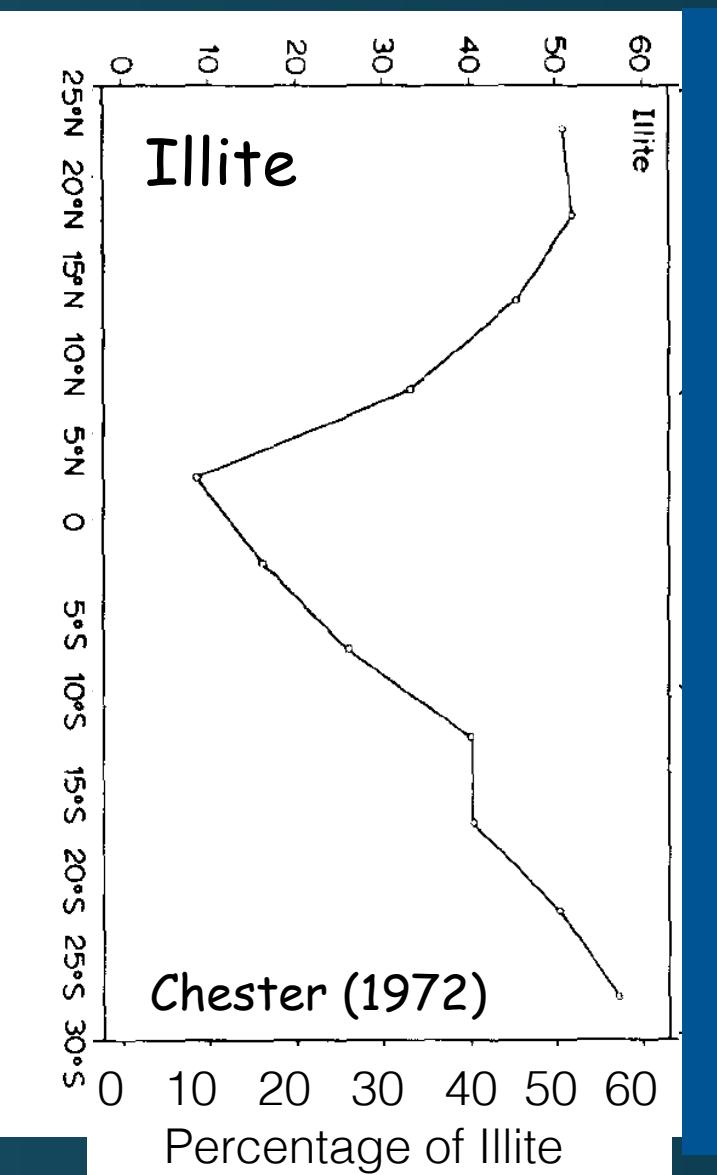
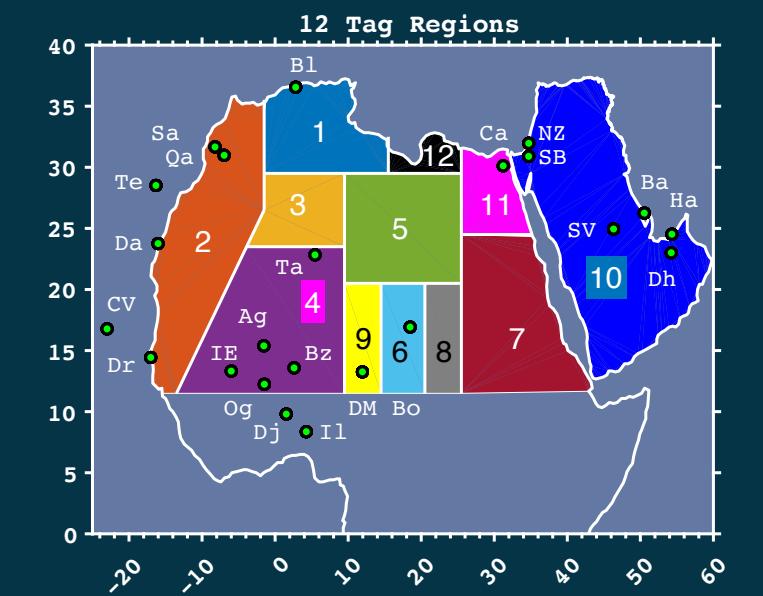
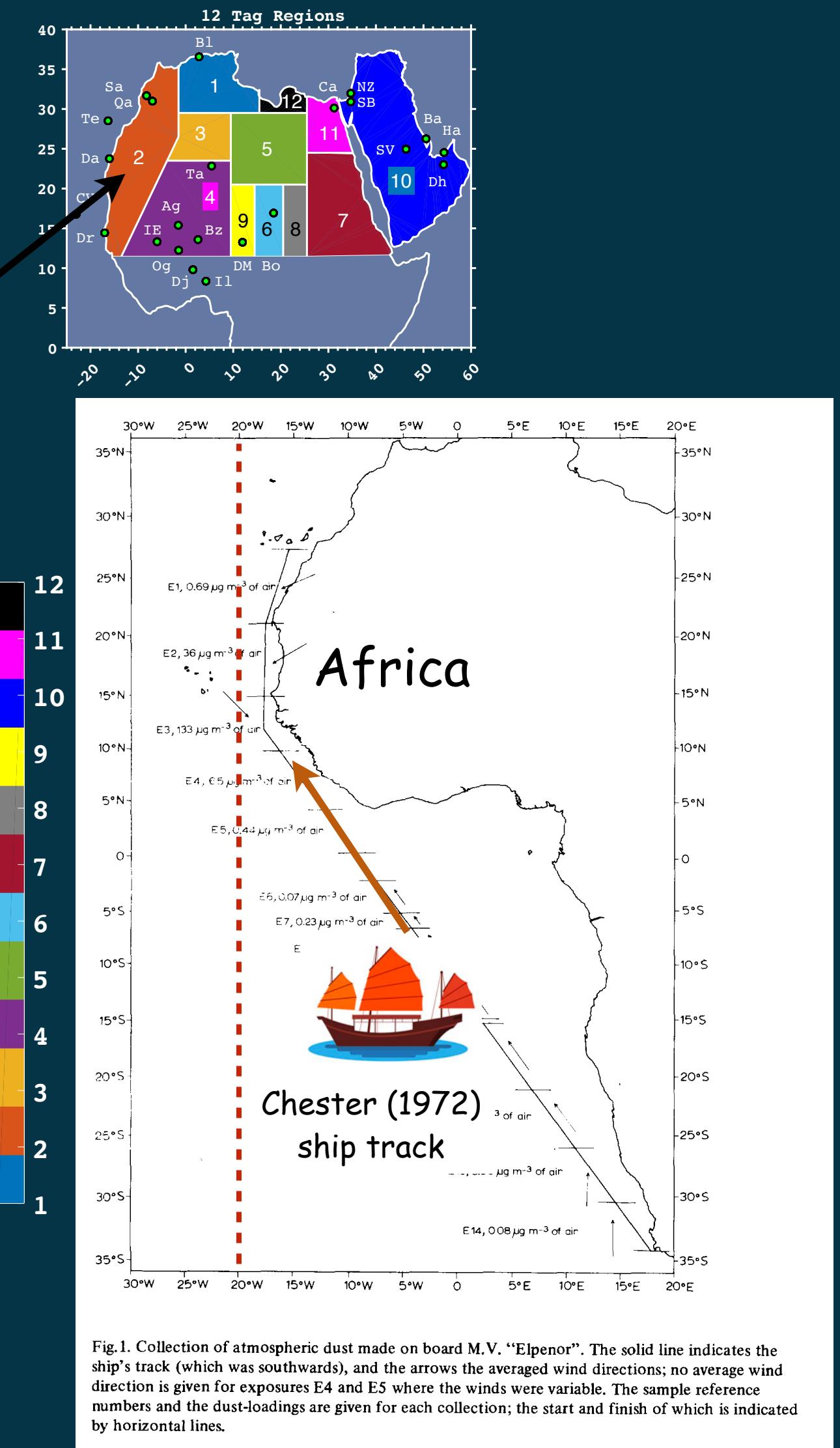
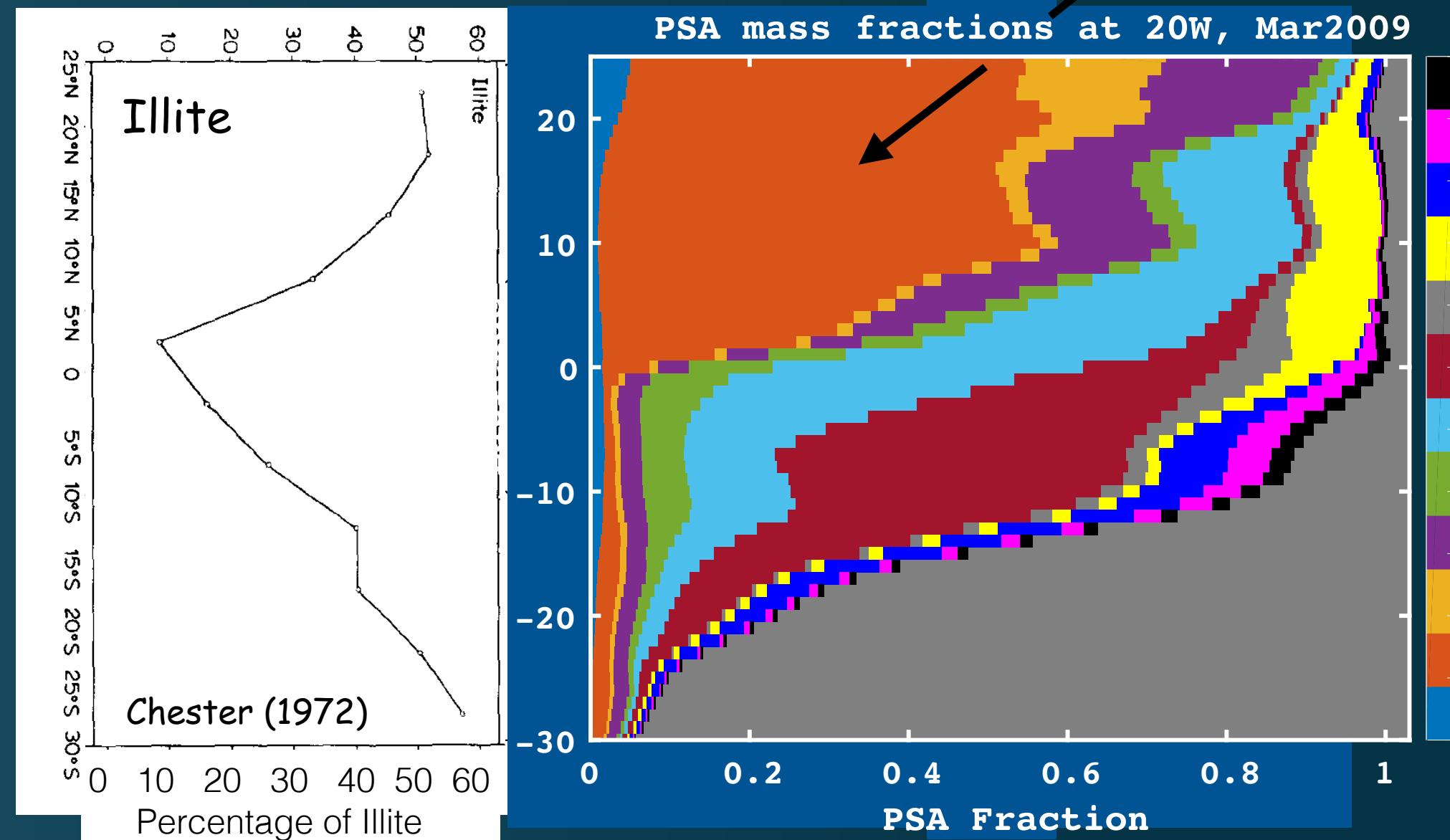


Fig.1. Collection of atmospheric dust made on board M.V. "Elpenor". The solid line indicates the ship's track (which was southwards), and the arrows the averaged wind directions; no average wind direction is given for exposures E4 and E5 where the winds were variable. The sample reference numbers and the dust-loadings are given for each collection; the start and finish of which is indicated by horizontal lines.

Quantify the Origin of Aeolian Dust Collected over the Atlantic

High concentrations of illite in West Africa
(Caquineau, JGR 2002)



Use Generalized Matrix Inverse (GMI) with AERONET retrievals to assign optical properties to each region

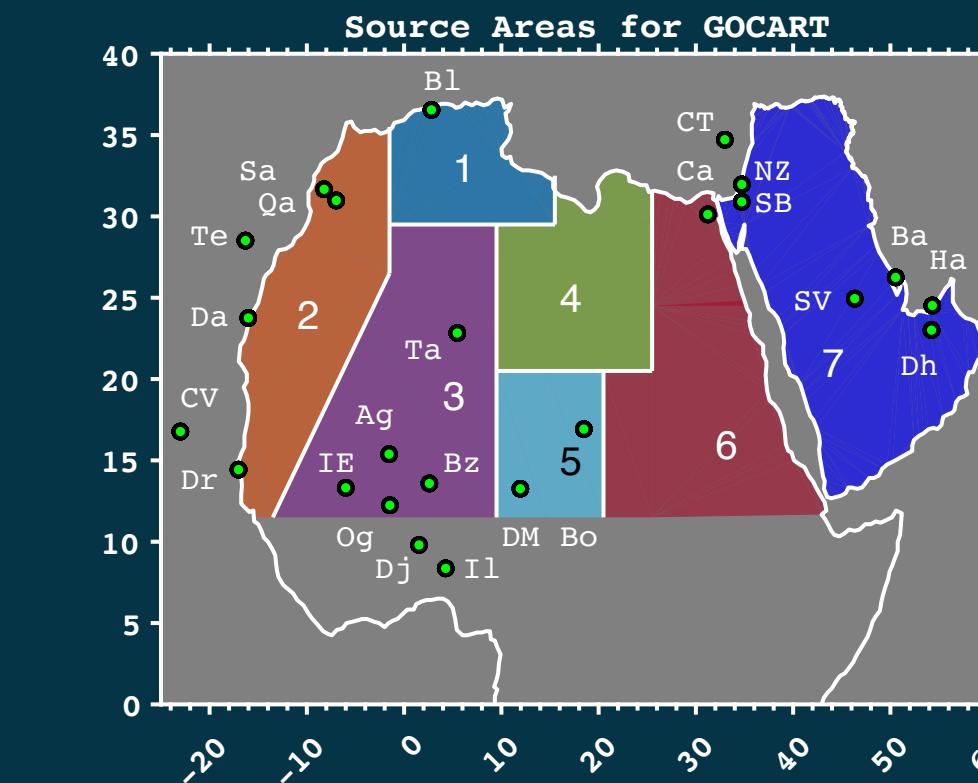
GOCART model runs allow us to write:

$$\vec{n}_a = \mathbf{F} \vec{n}_p$$

↑
↑
Real refr index column vector of
7 + 1 source regions

M X 8 matrix of PSA fractions at AERONETs
(GOCART, 2006-2016)

Real refr index of M AERONET retrievals with
 $AOT(440) \geq 0.4, dp \geq 0.25, fvf < 0.05$



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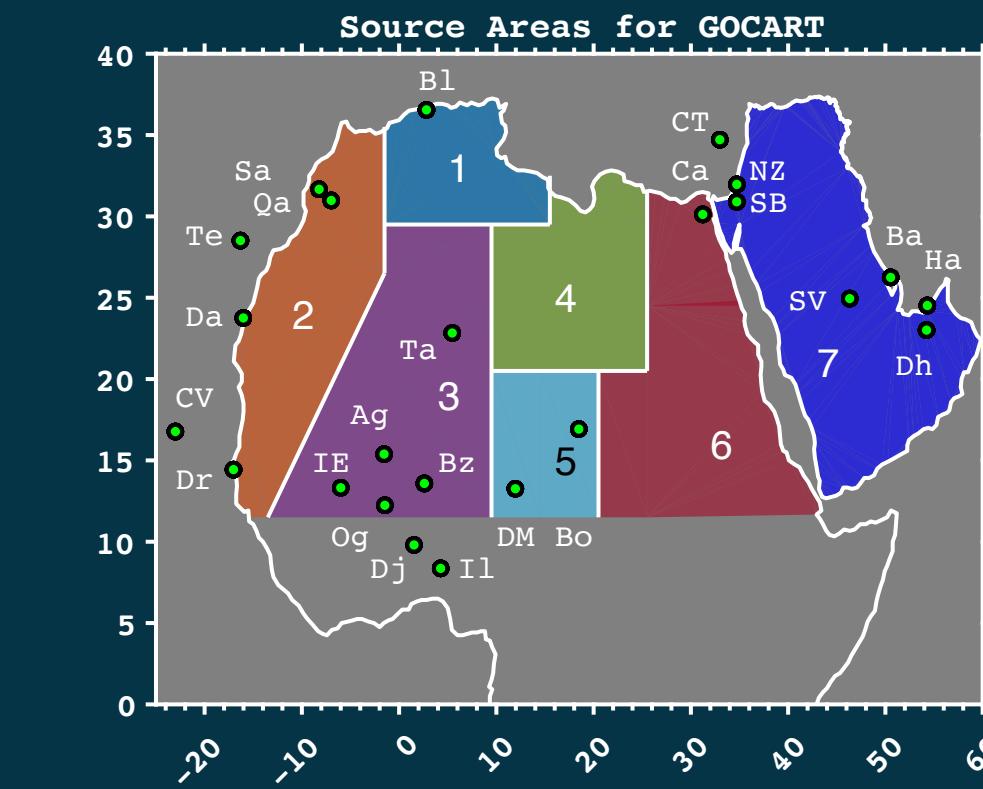
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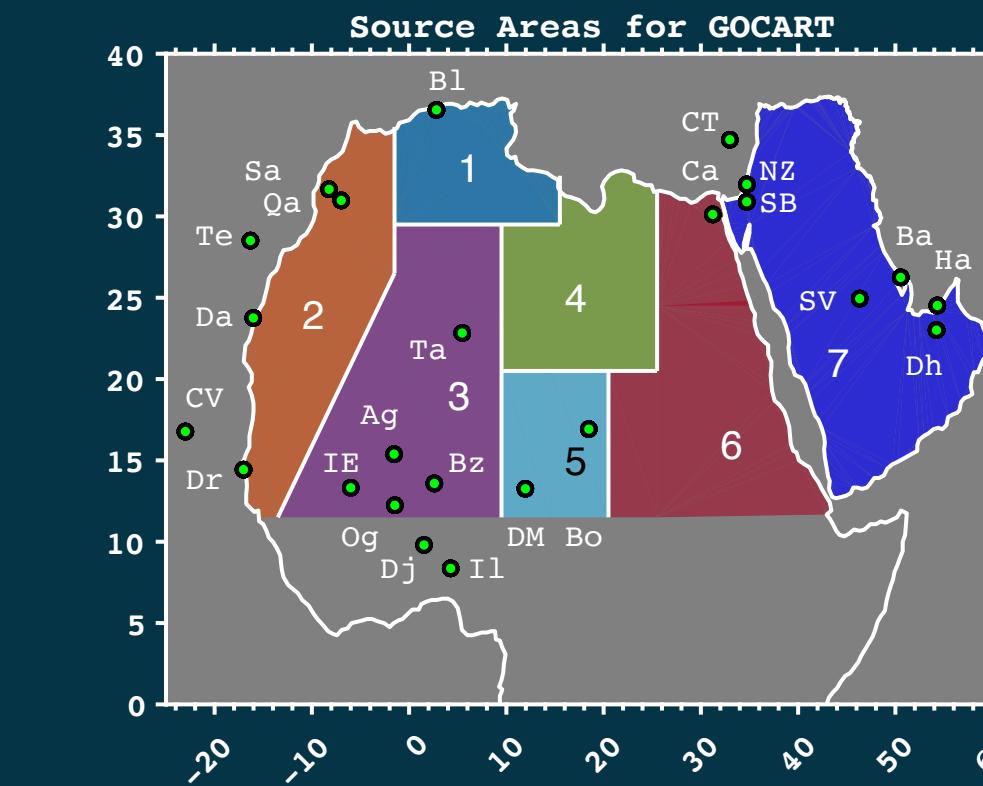
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Have $M = 1388$, so overconstrained (i.e., $M > 8$). Solve for \vec{n}_p

$$\hat{\vec{n}}_p = (\mathbf{F}^t \mathbf{F})^{-1} \mathbf{F}^t \vec{n}_a$$



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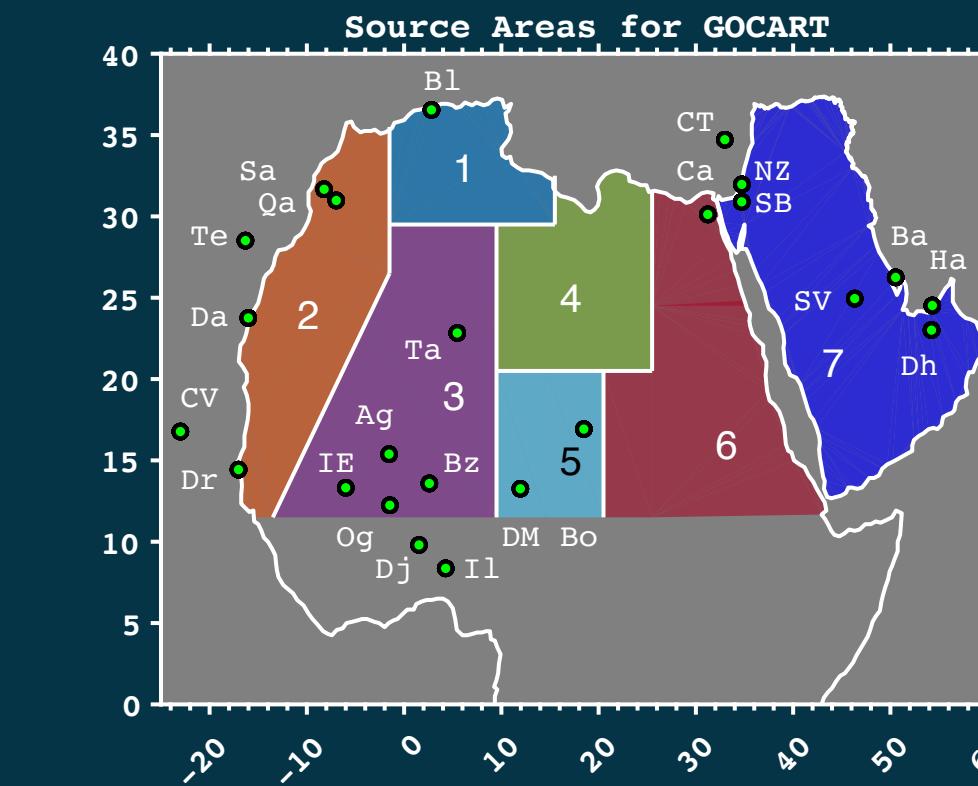
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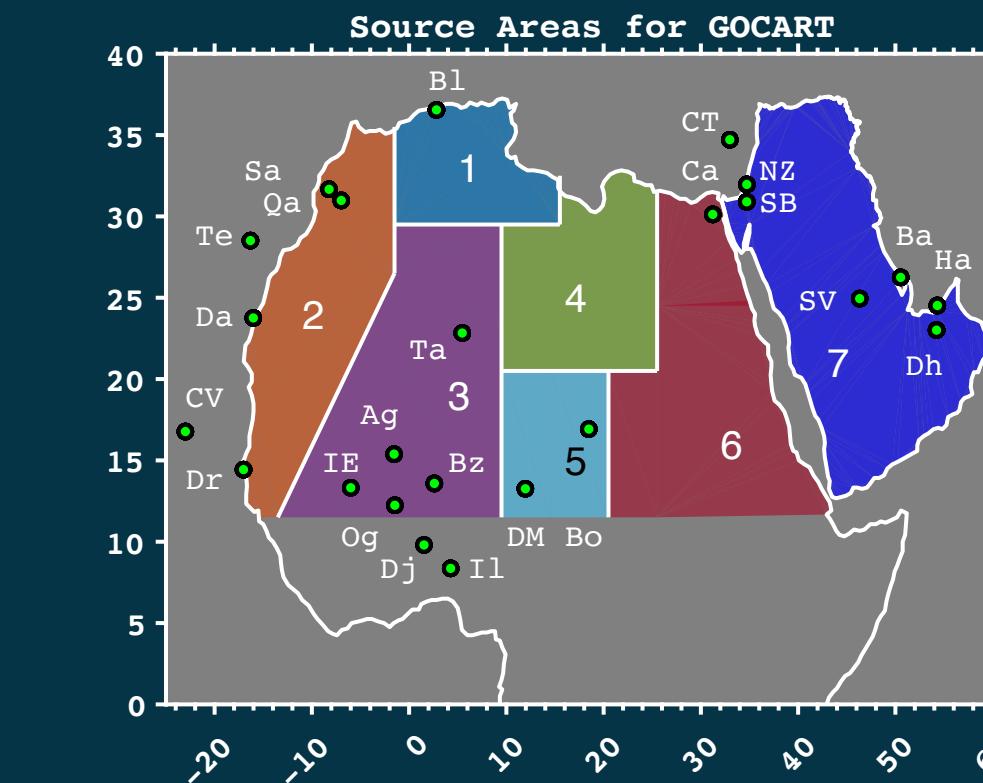
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$$\hat{k}_p = (\mathbf{F}^t \mathbf{F})^{-1} \mathbf{F}^t \vec{k}_a$$

Compute S_a from (\hat{n}_p, \hat{k}_p) and modeled size distributions



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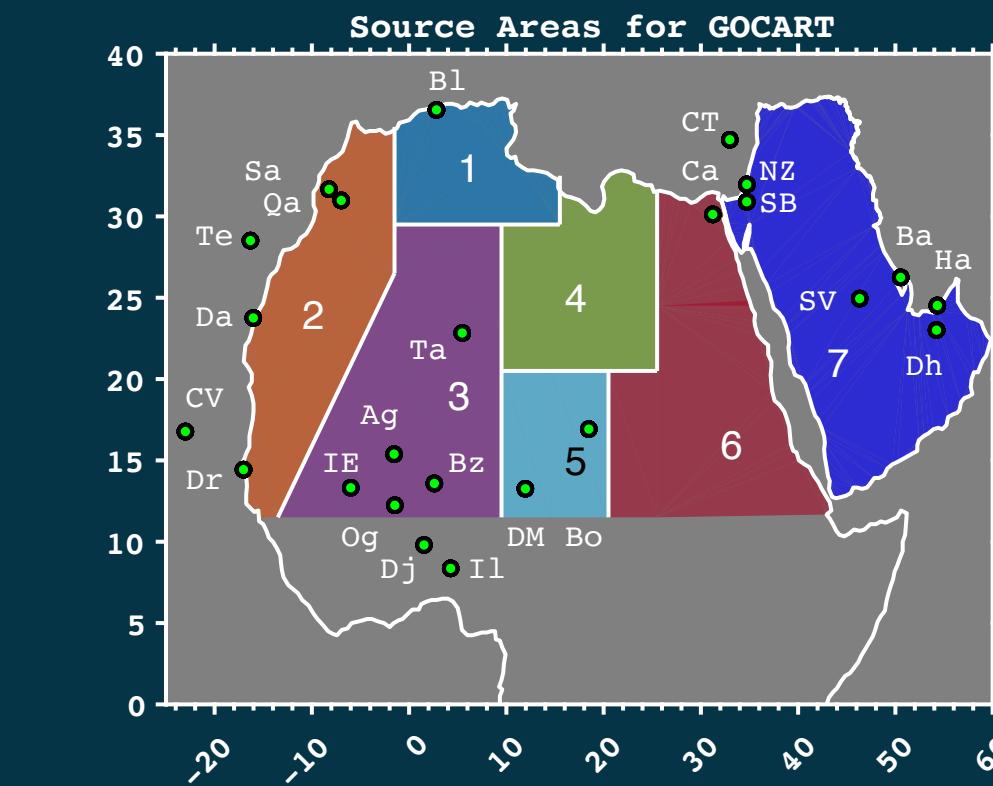
$$\hat{\vec{n}}_p = (\vec{F}^t \vec{F})^{-1} \vec{F}^t \vec{n}_a$$

$$\hat{\vec{k}}_p = (\vec{F}^t \vec{F})^{-1} \vec{F}^t \vec{k}_a$$

$$\hat{\vec{S}}_p = (\vec{F}^t \vec{F})^{-1} \vec{F}^t \vec{S}_a$$

Compute S_a from $(\hat{\vec{n}}_p, \hat{\vec{k}}_p)$ and modeled size distributions

$$S_a = \sum_i F_i(S_p)_i \quad \text{Empirical approach}$$

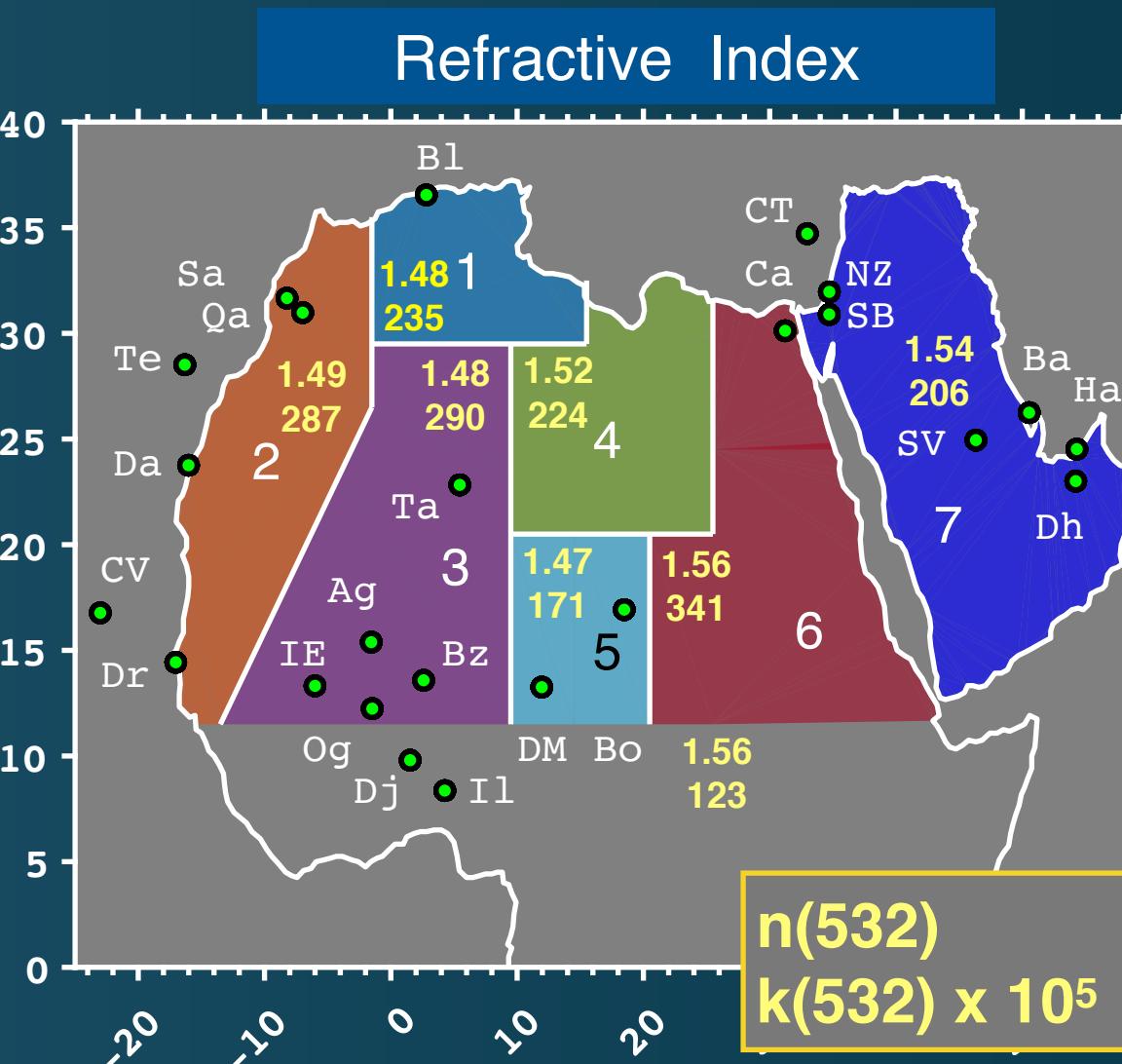


Regional Refractive Indices and Lidar Ratios from GMI

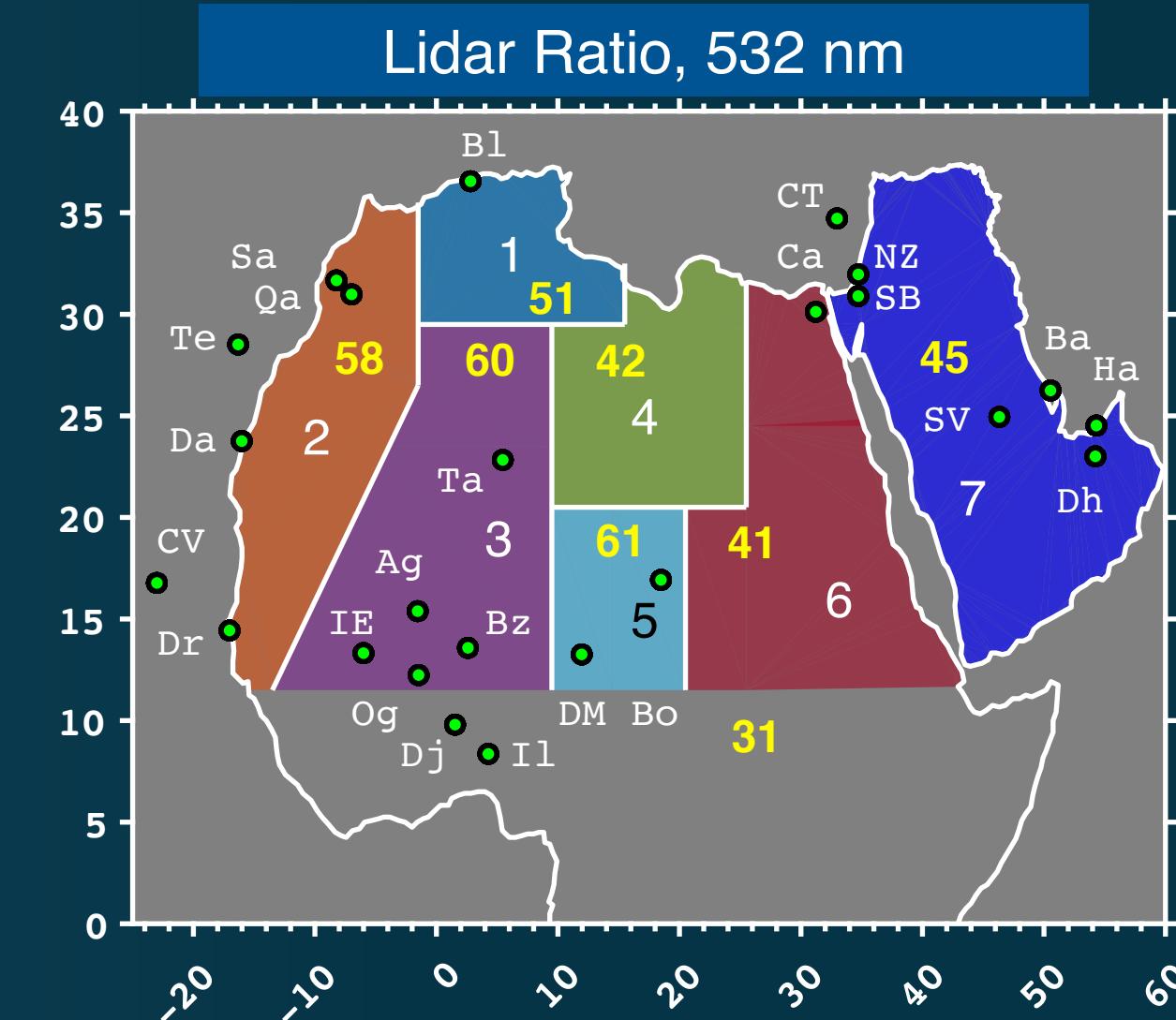
$$\hat{n}_p = (F^t F)^{-1} F^t \vec{n}_a$$

$$\widehat{k}_n = (F^t F)^{-1} F^t \overrightarrow{k}_a$$

$$\widehat{S}_p = (F^t F)^{-1} F^t \overrightarrow{S}_a$$



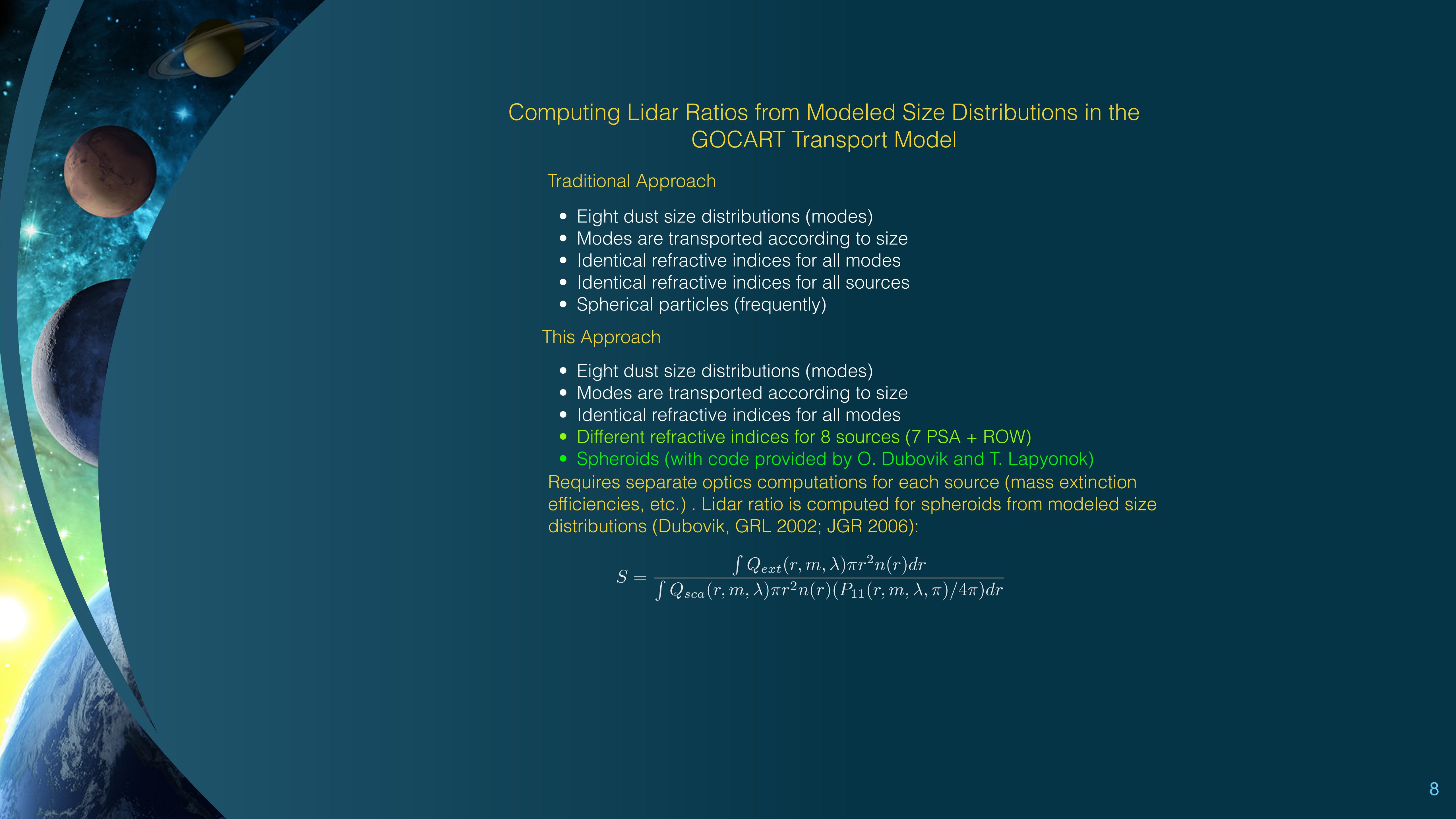
Compute S_a with GOCART: (n_p, k_p) and GOCART $dn/dr \Rightarrow$ Lidar Ratio S_a



Empirical Approach: $S_a = \sum_i F_i(S_p)_i$

Reference k-values:

GOCART $k \times 10^5$:	550
AERONET $k_{min} \times 10^5$:	50



Computing Lidar Ratios from Modeled Size Distributions in the GOCART Transport Model

Traditional Approach

- Eight dust size distributions (modes)
- Modes are transported according to size
- Identical refractive indices for all modes
- Identical refractive indices for all sources
- Spherical particles (frequently)

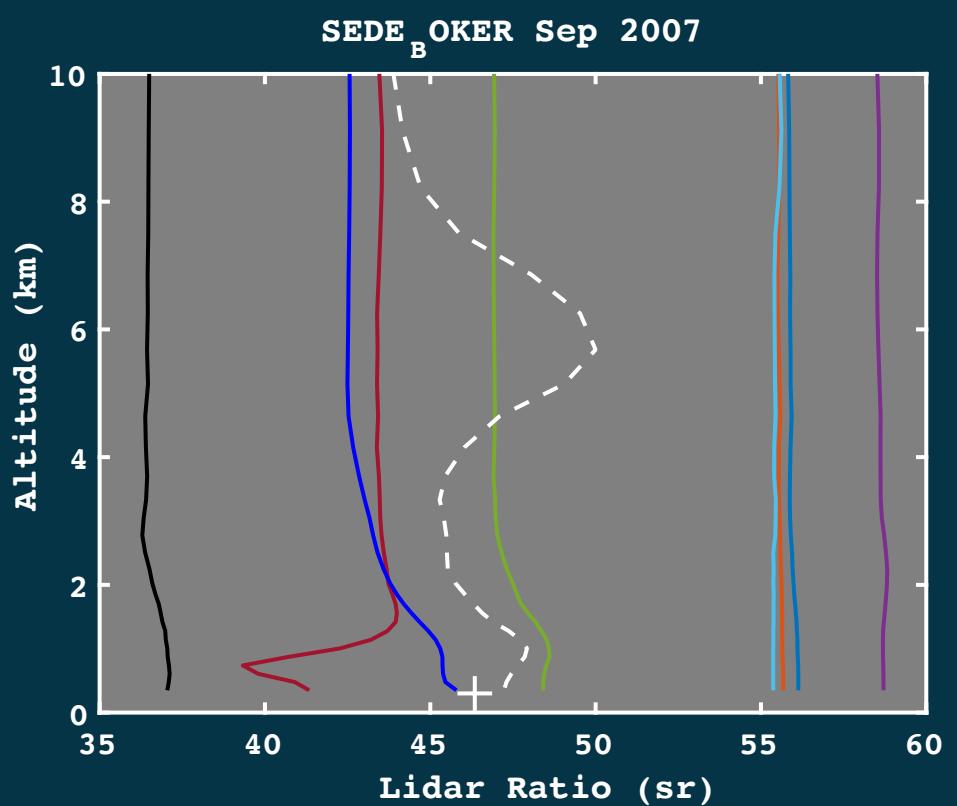
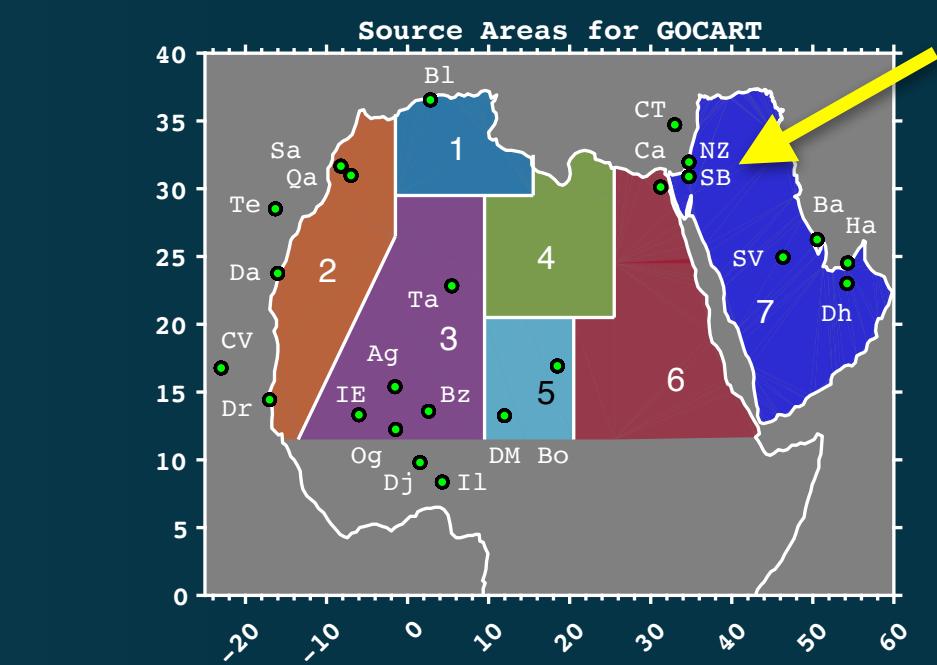
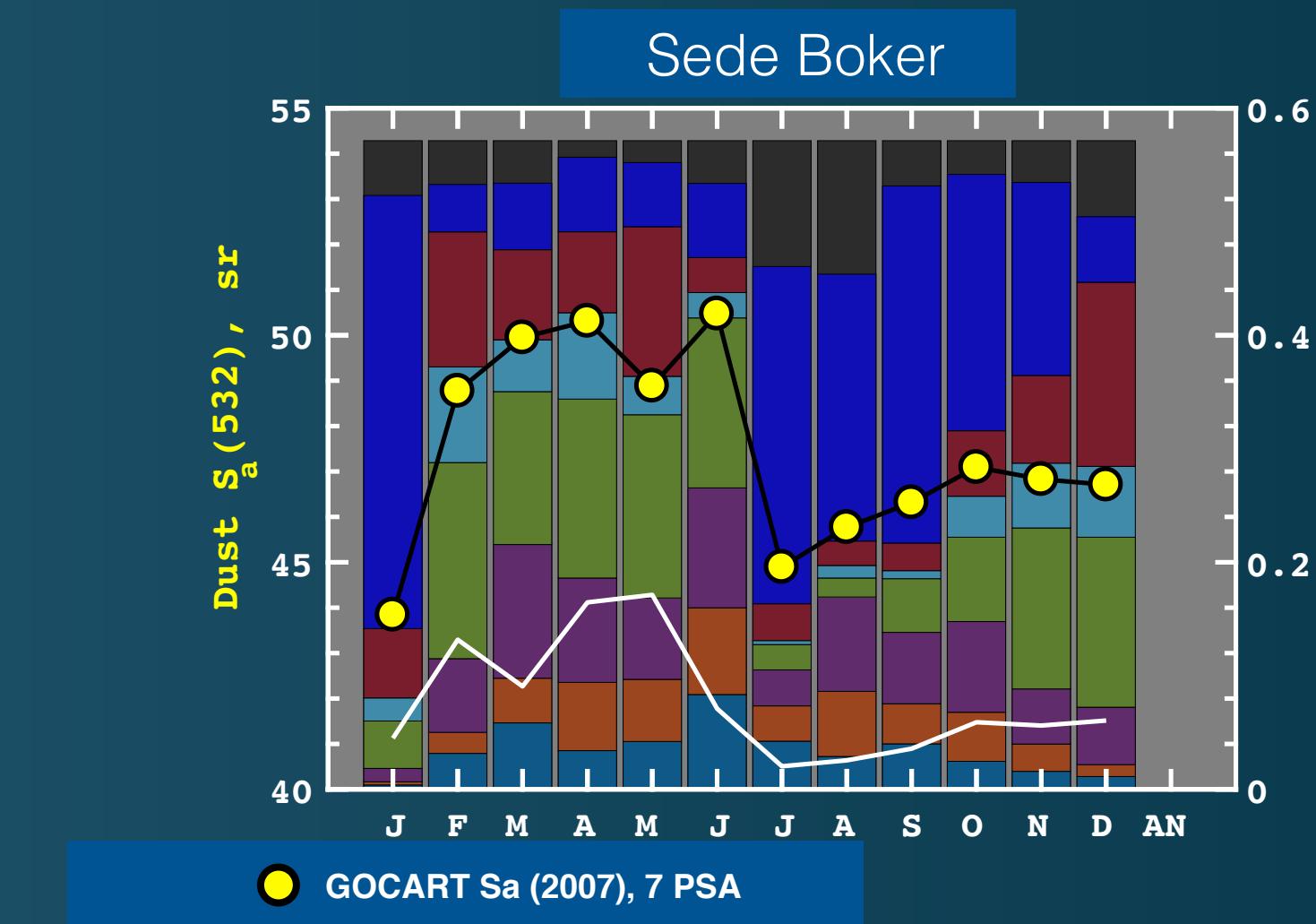
This Approach

- Eight dust size distributions (modes)
- Modes are transported according to size
- Identical refractive indices for all modes
- Different refractive indices for 8 sources (7 PSA + ROW)
- Spheroids (with code provided by O. Dubovik and T. Lapyonok)

Requires separate optics computations for each source (mass extinction efficiencies, etc.) . Lidar ratio is computed for spheroids from modeled size distributions (Dubovik, GRL 2002; JGR 2006):

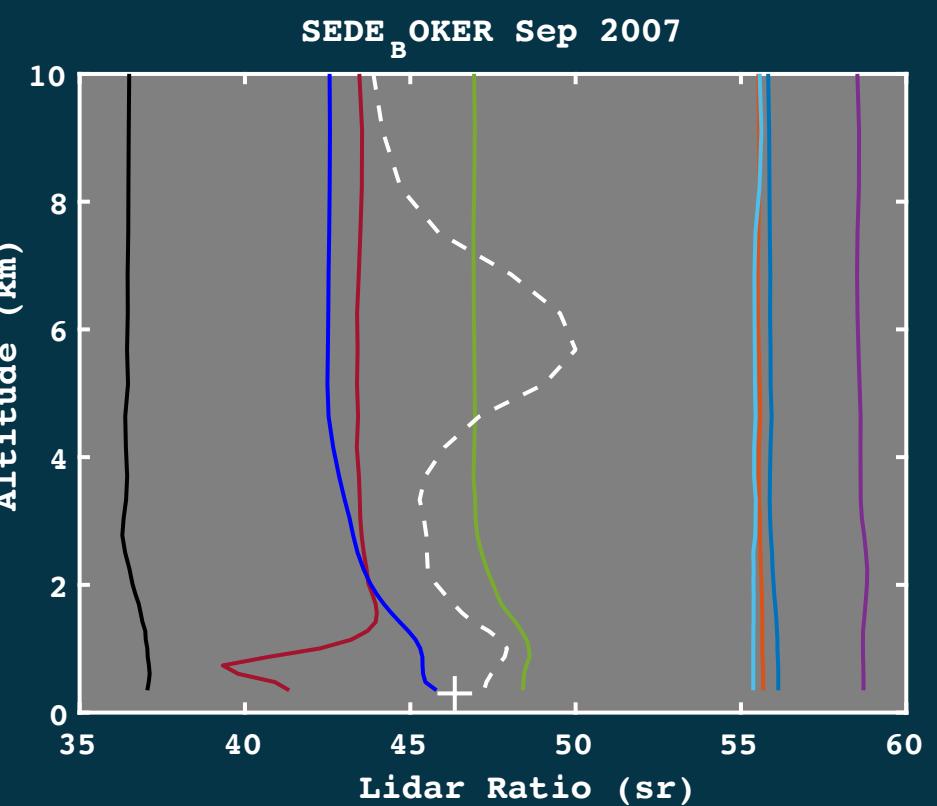
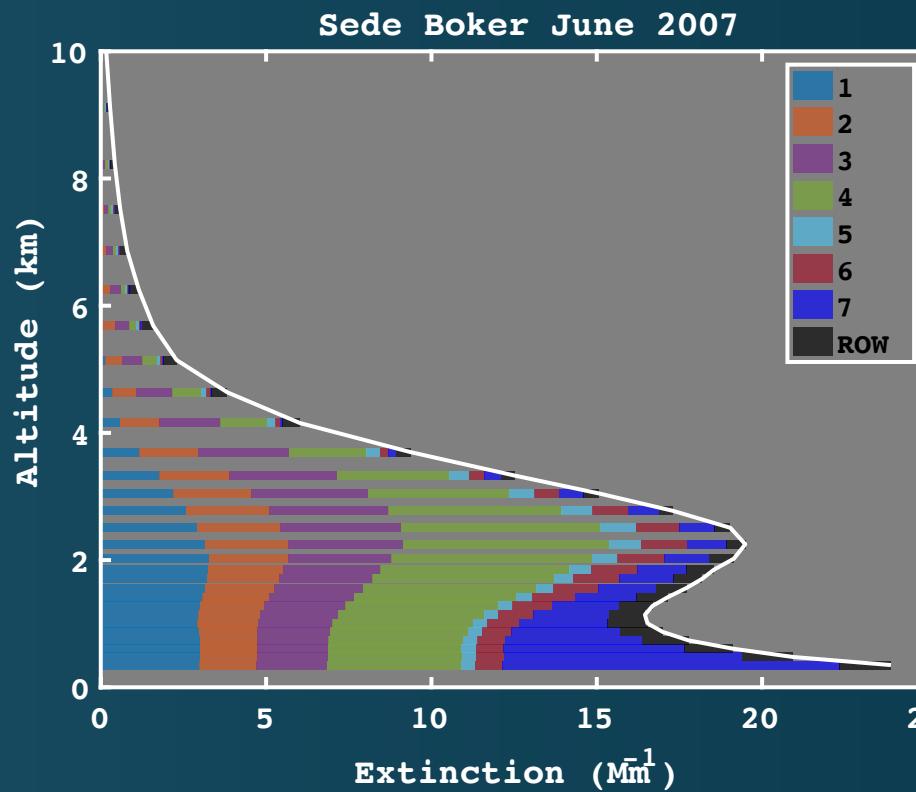
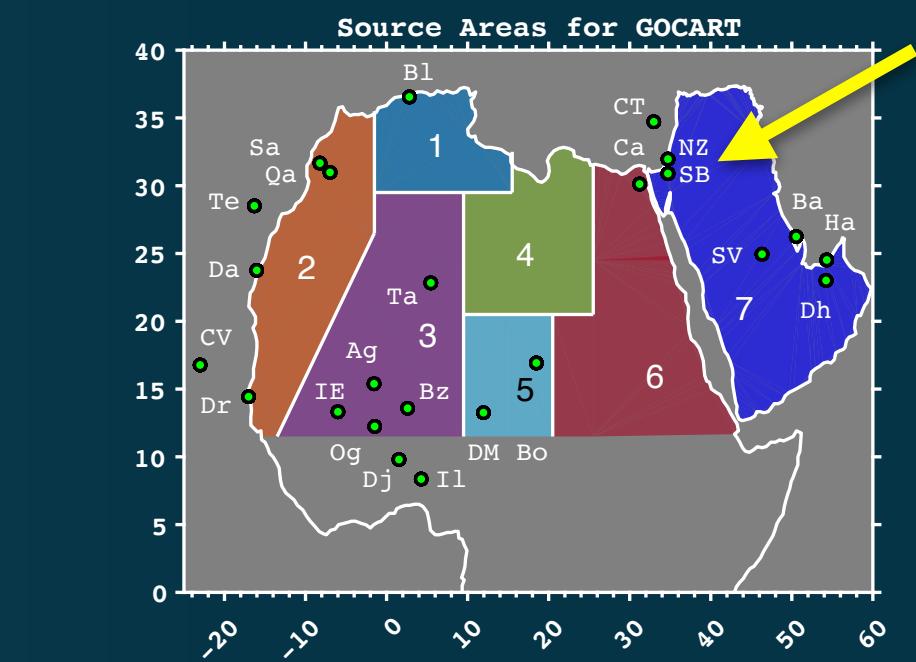
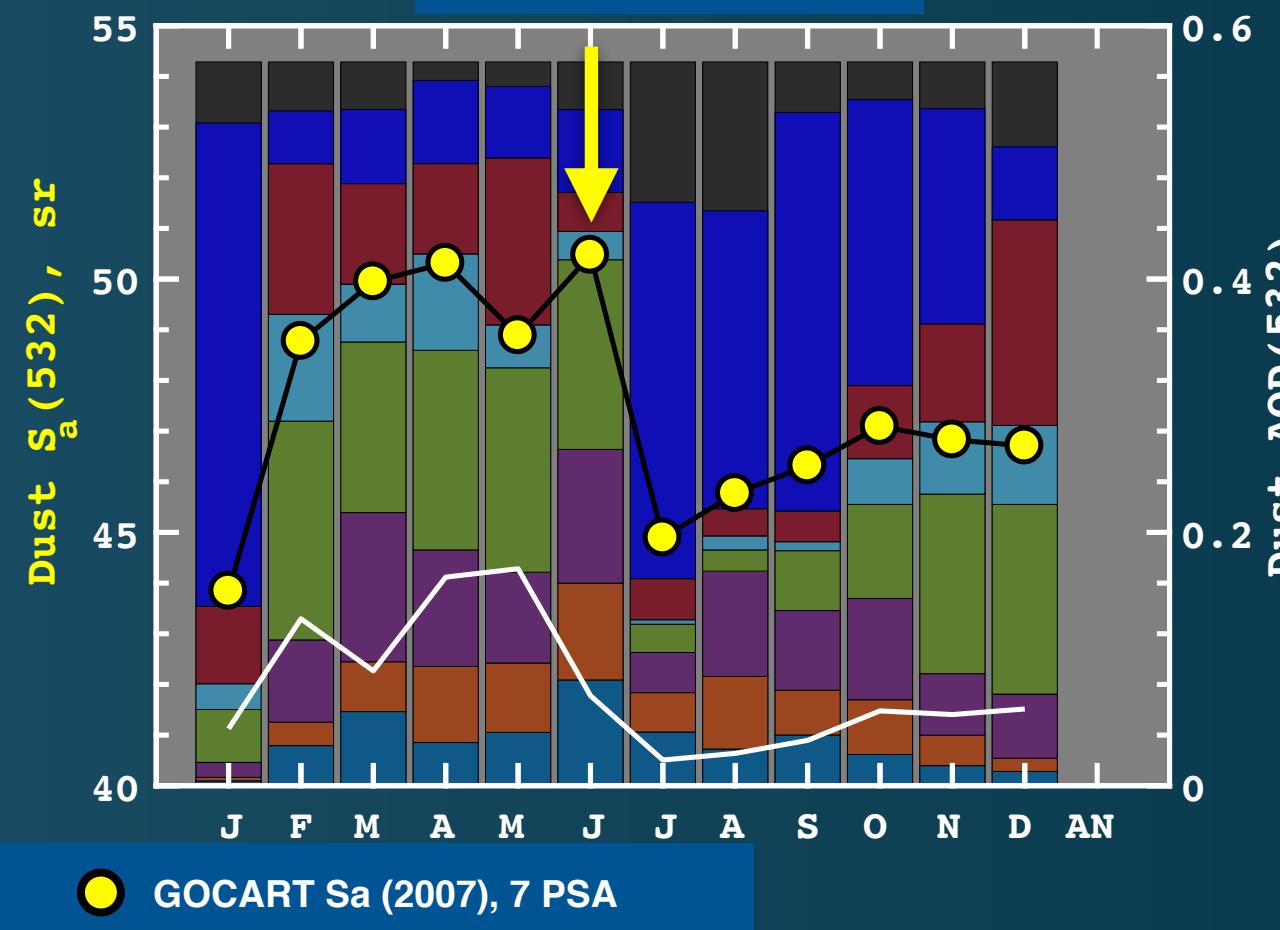
$$S = \frac{\int Q_{ext}(r, m, \lambda) \pi r^2 n(r) dr}{\int Q_{sca}(r, m, \lambda) \pi r^2 n(r) (P_{11}(r, m, \lambda, \pi)/4\pi) dr}$$

Example of Modeled and Empirical lidar ratios

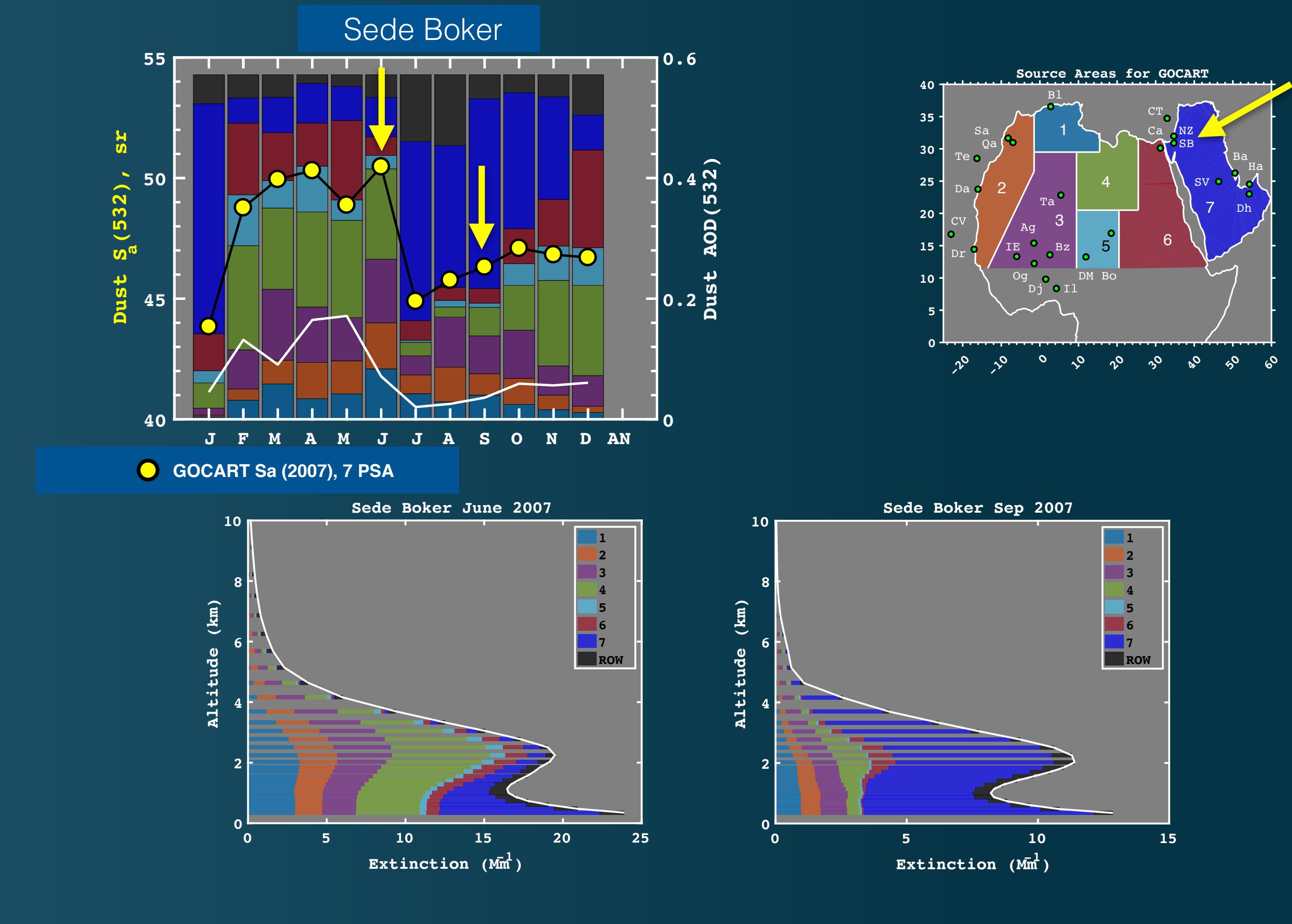


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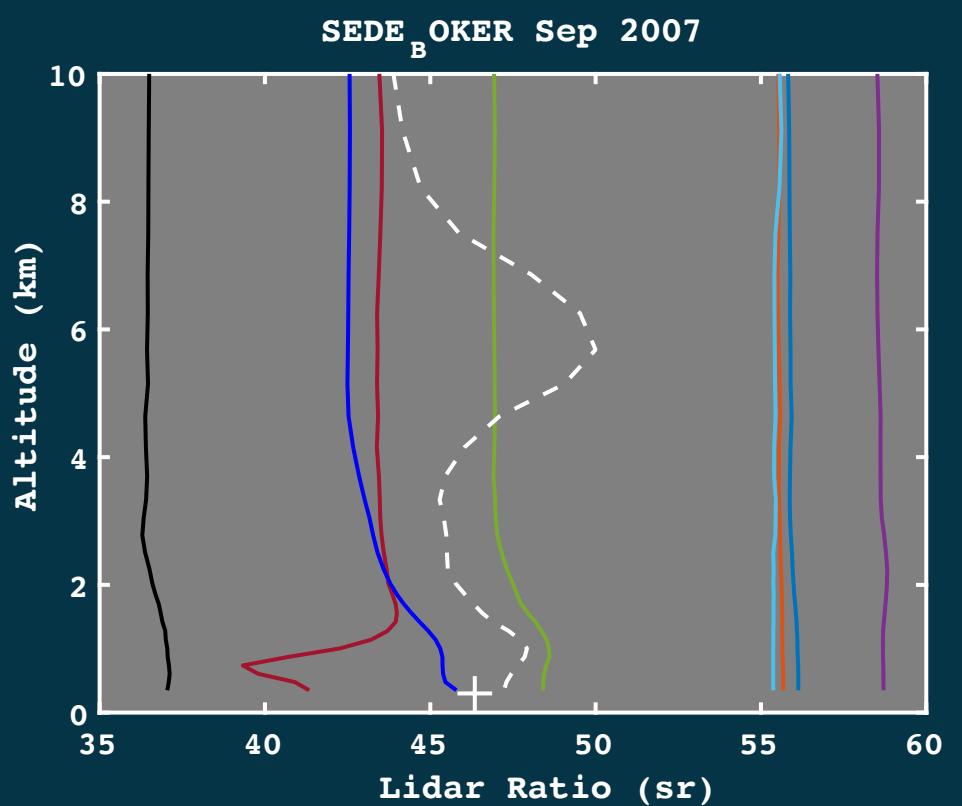
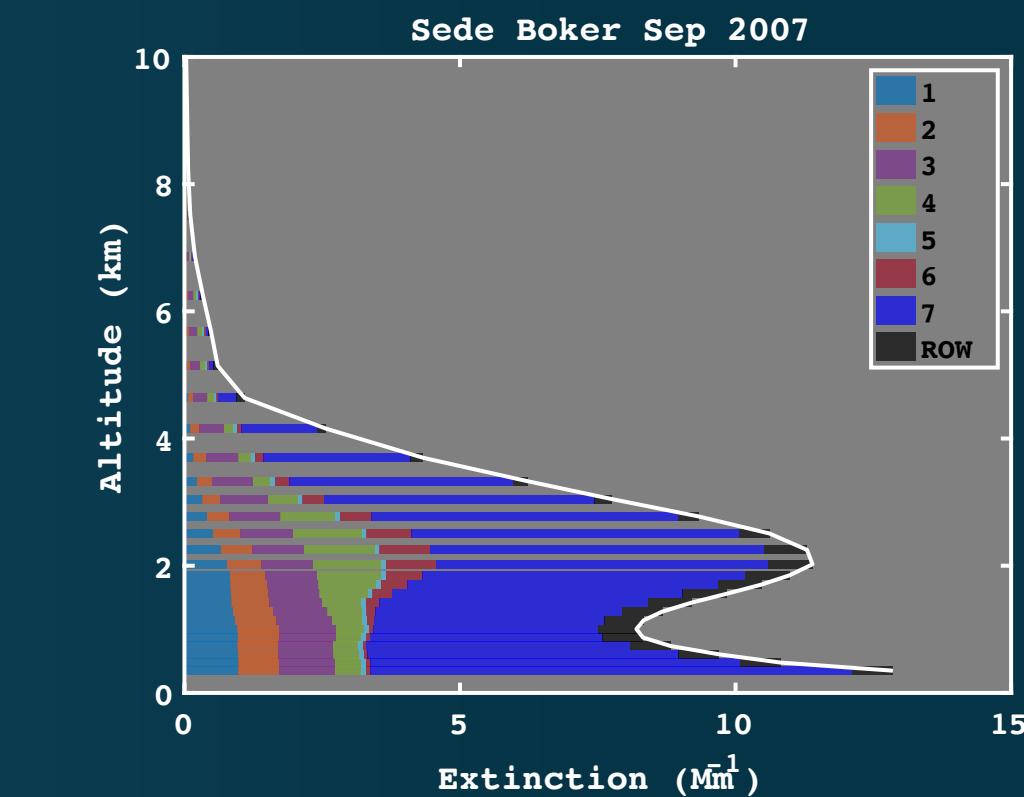
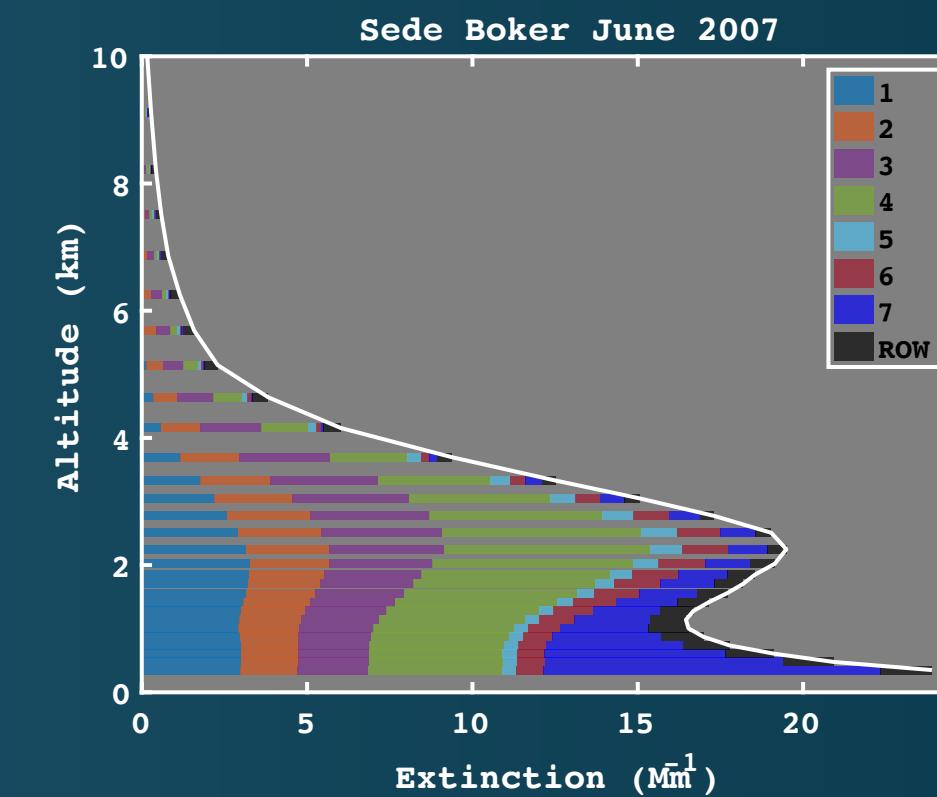
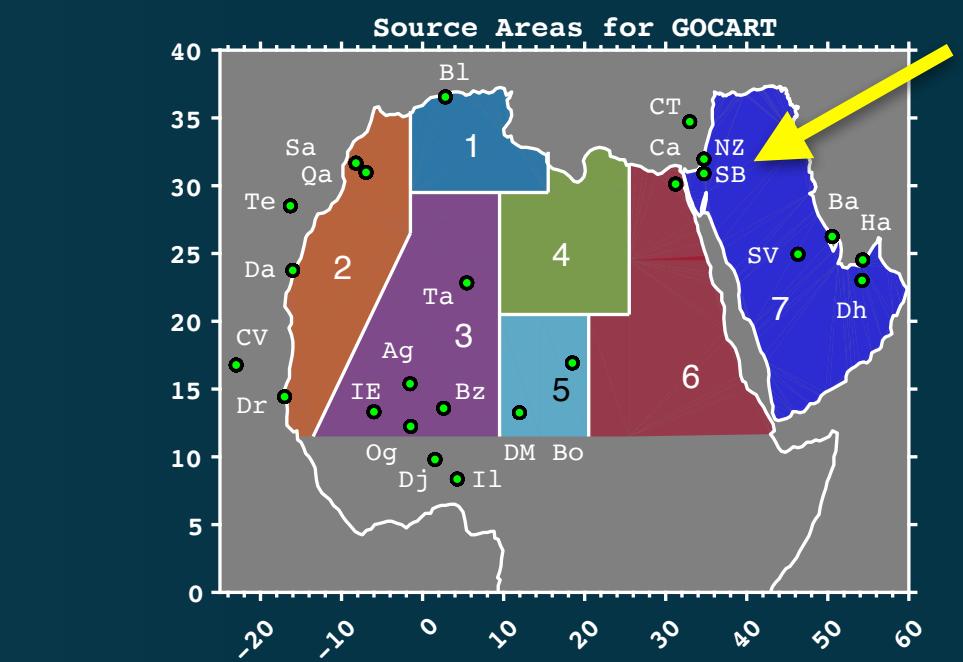
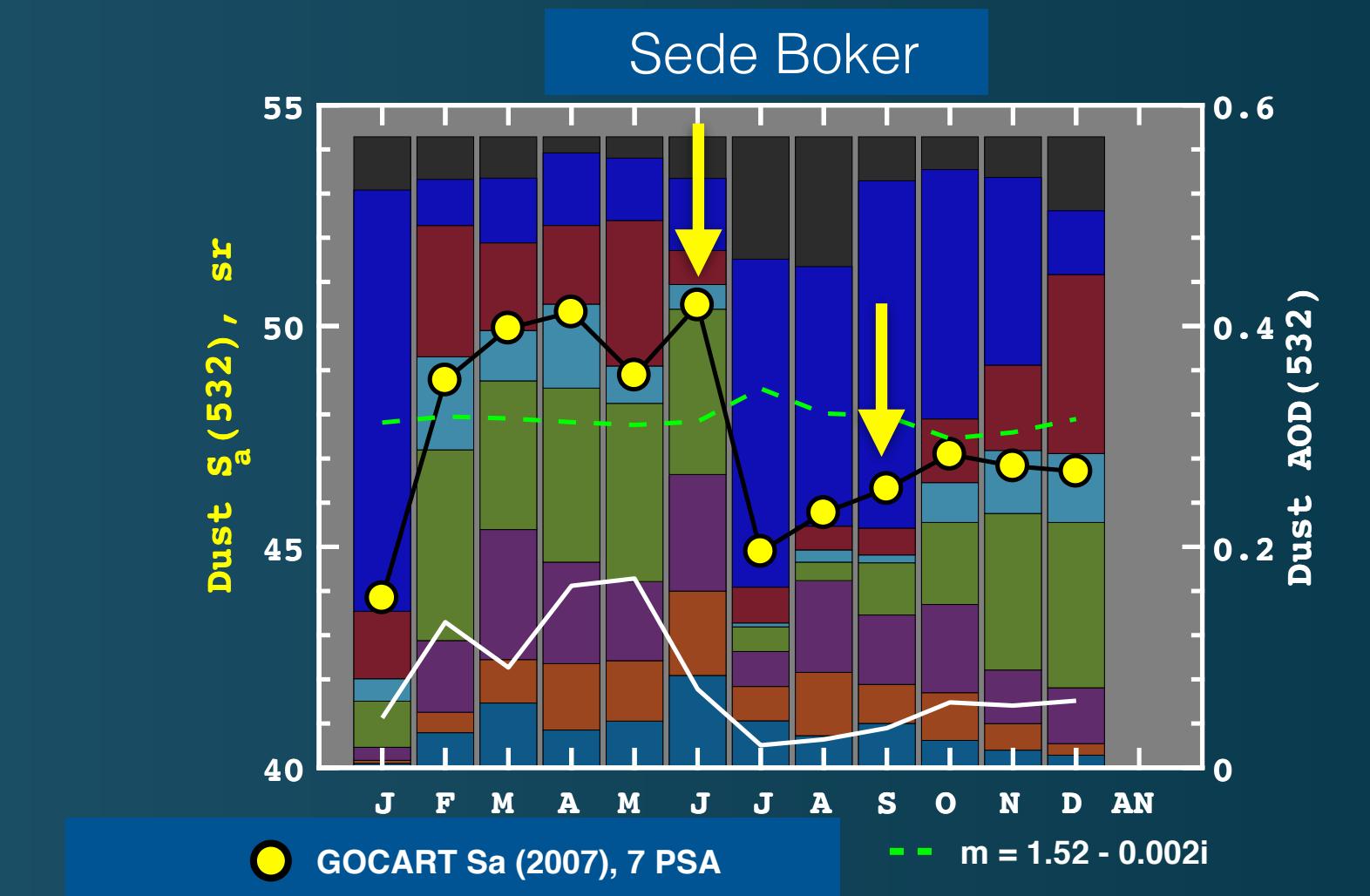
Sede Boker



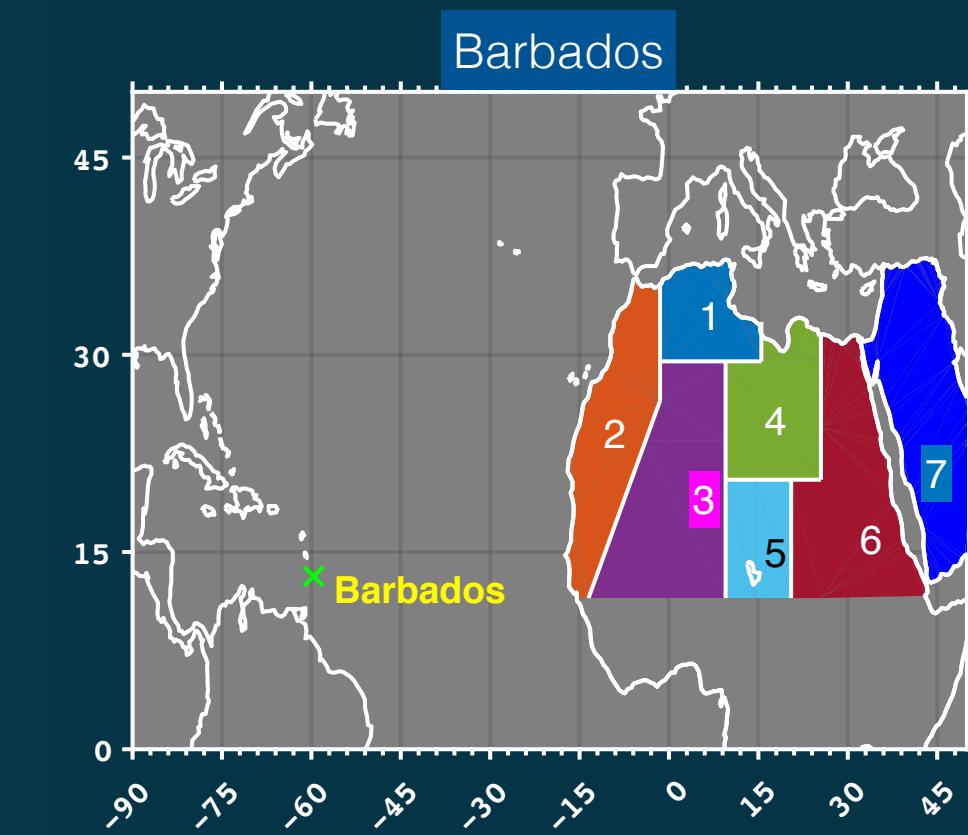
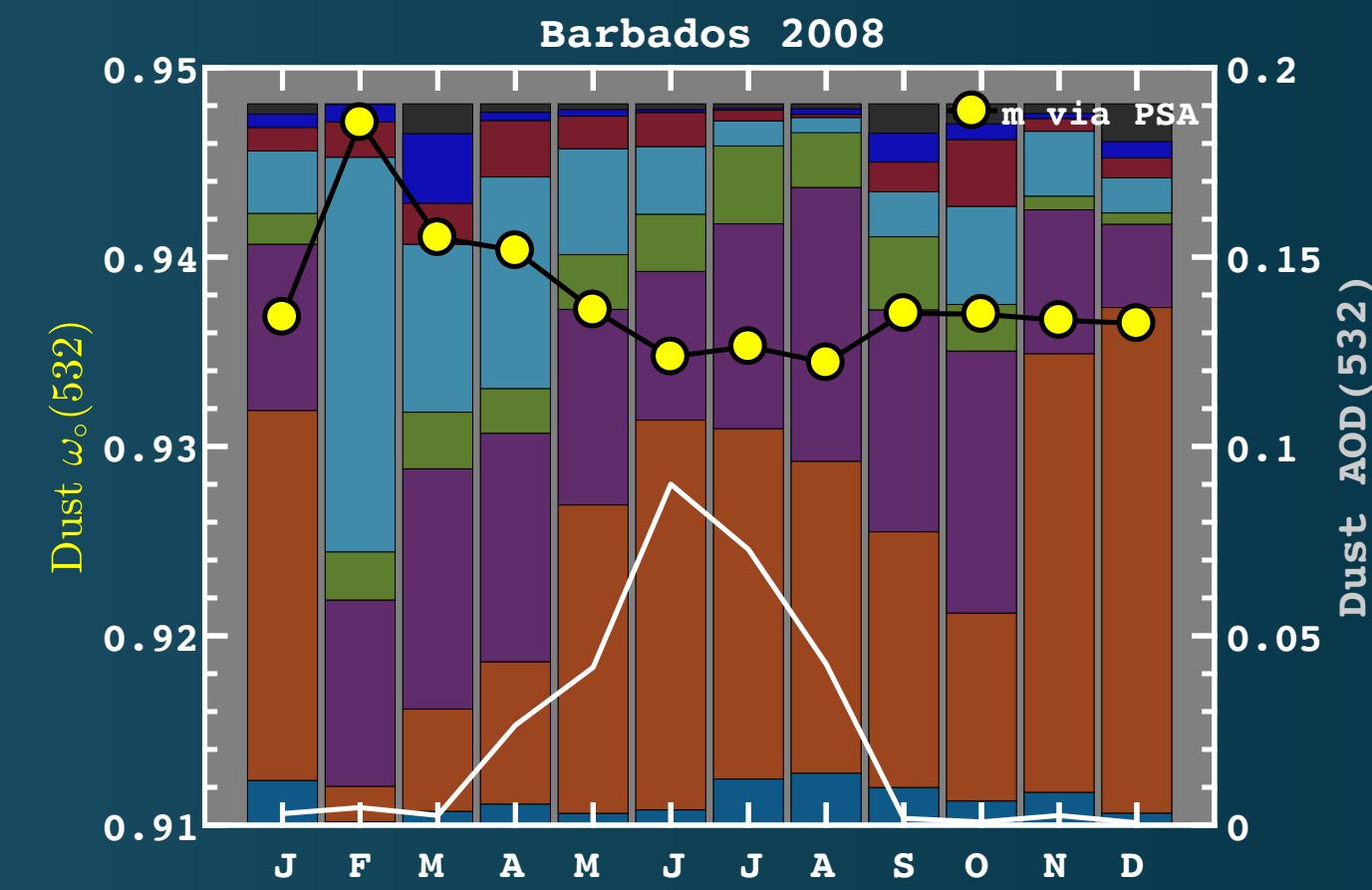
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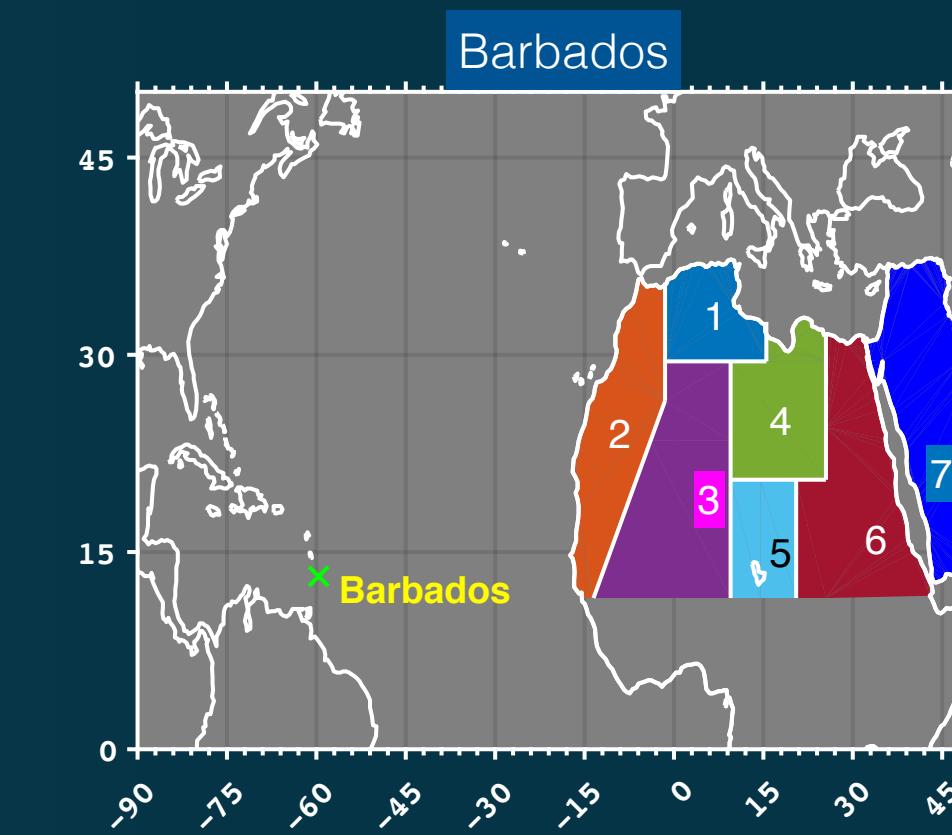
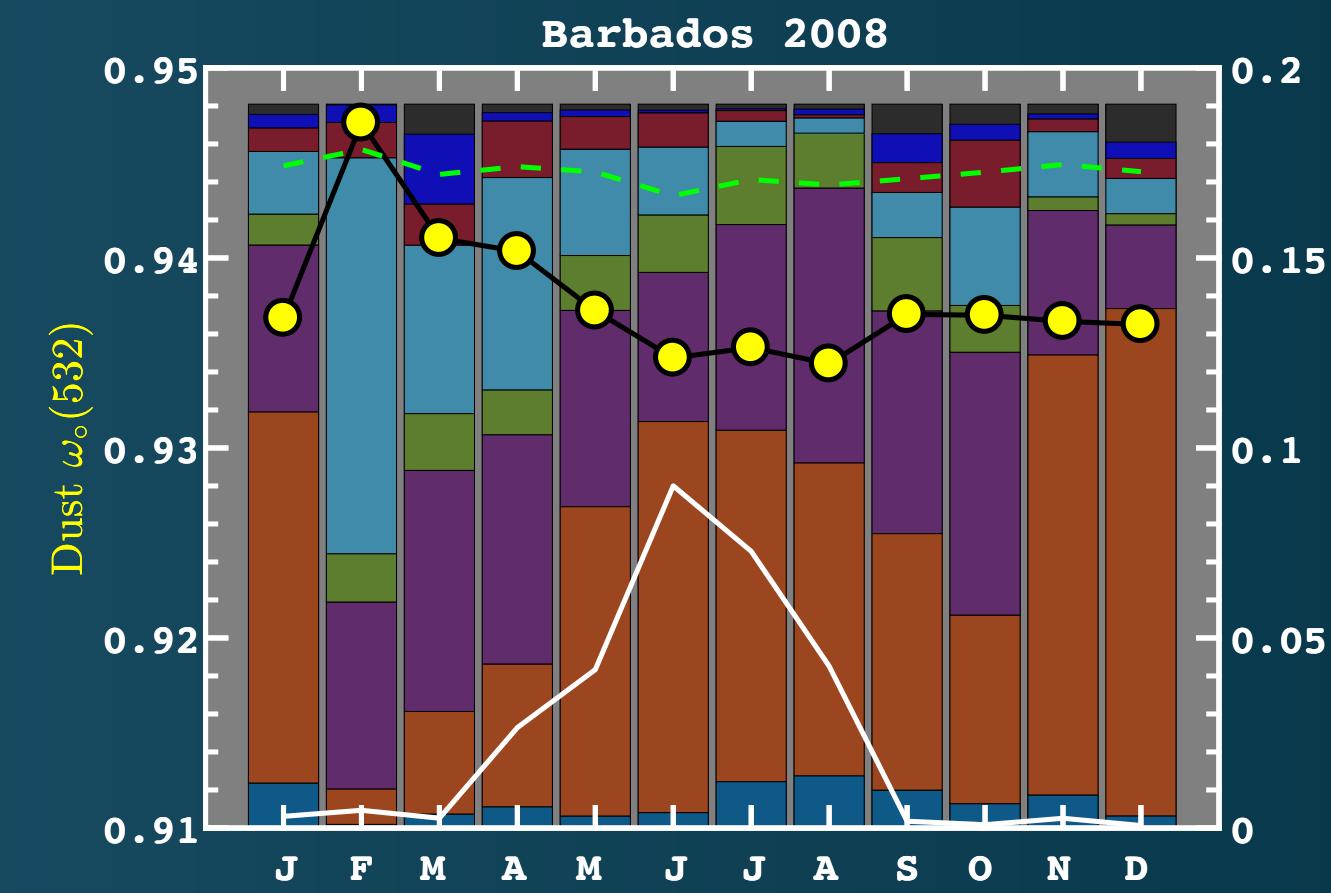
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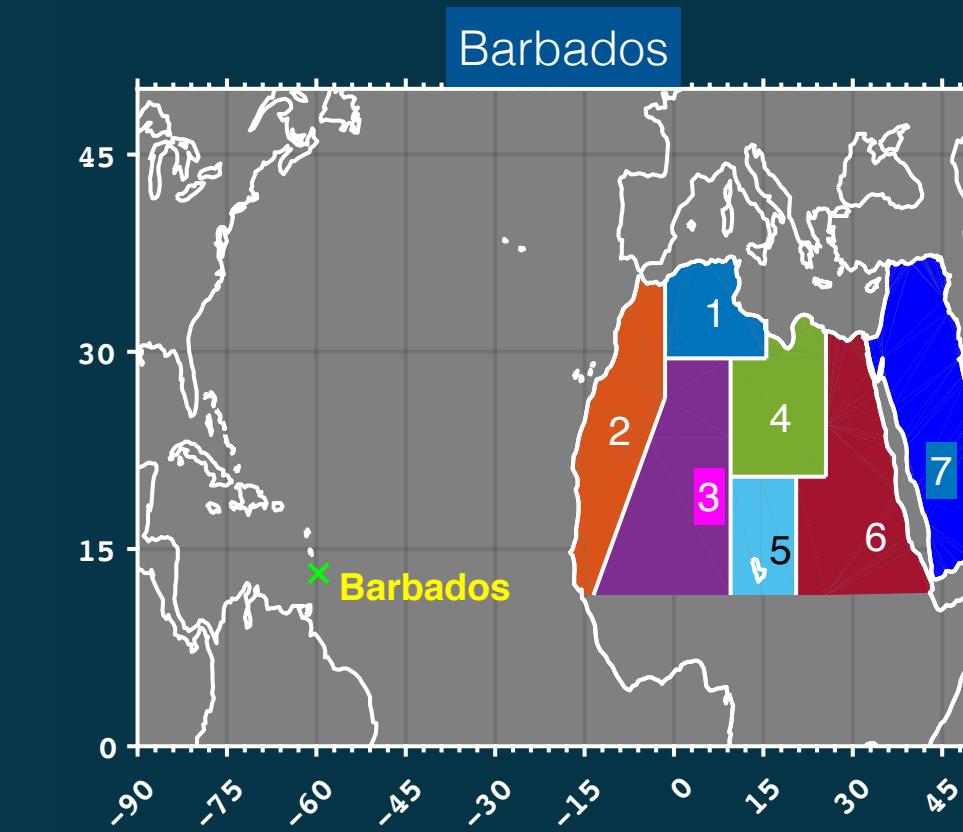
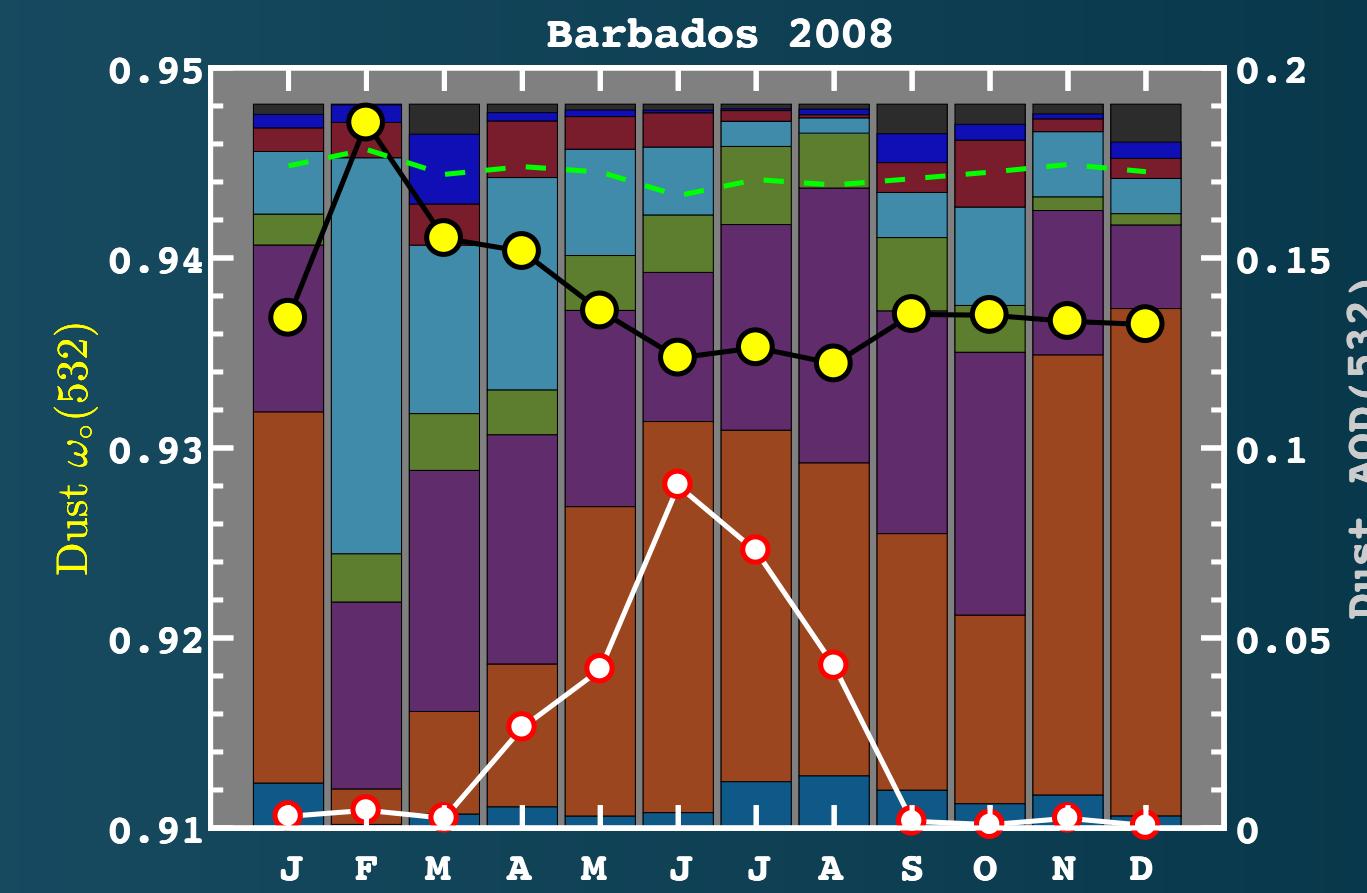
Source region impact on dust single-scatter albedo and AOD



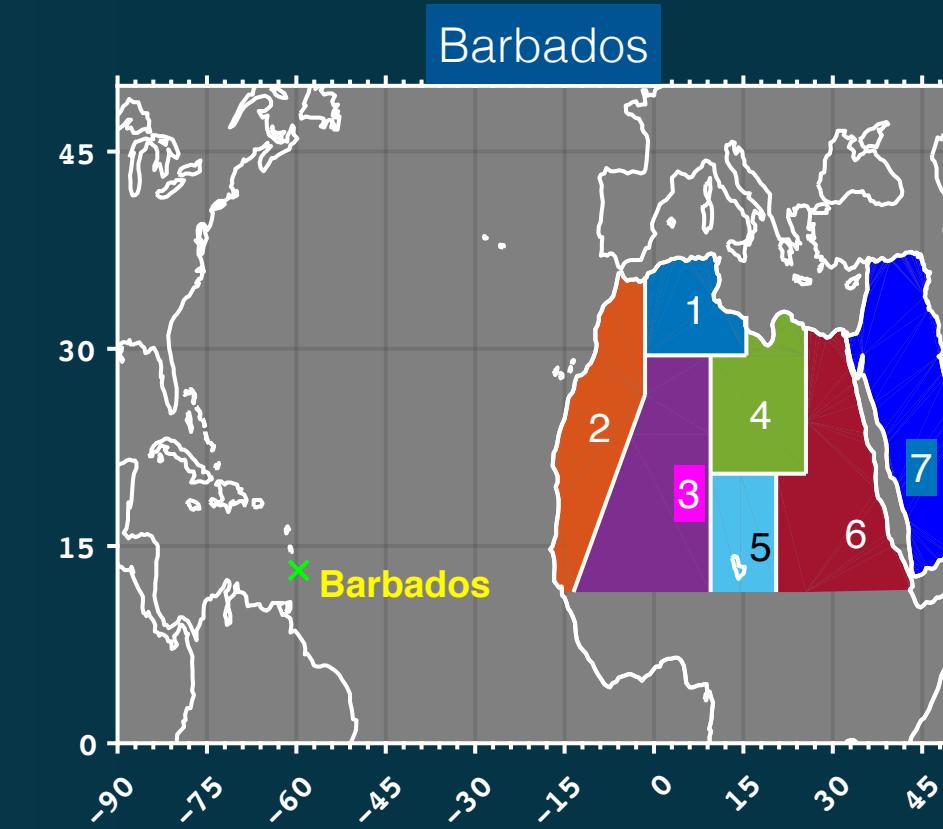
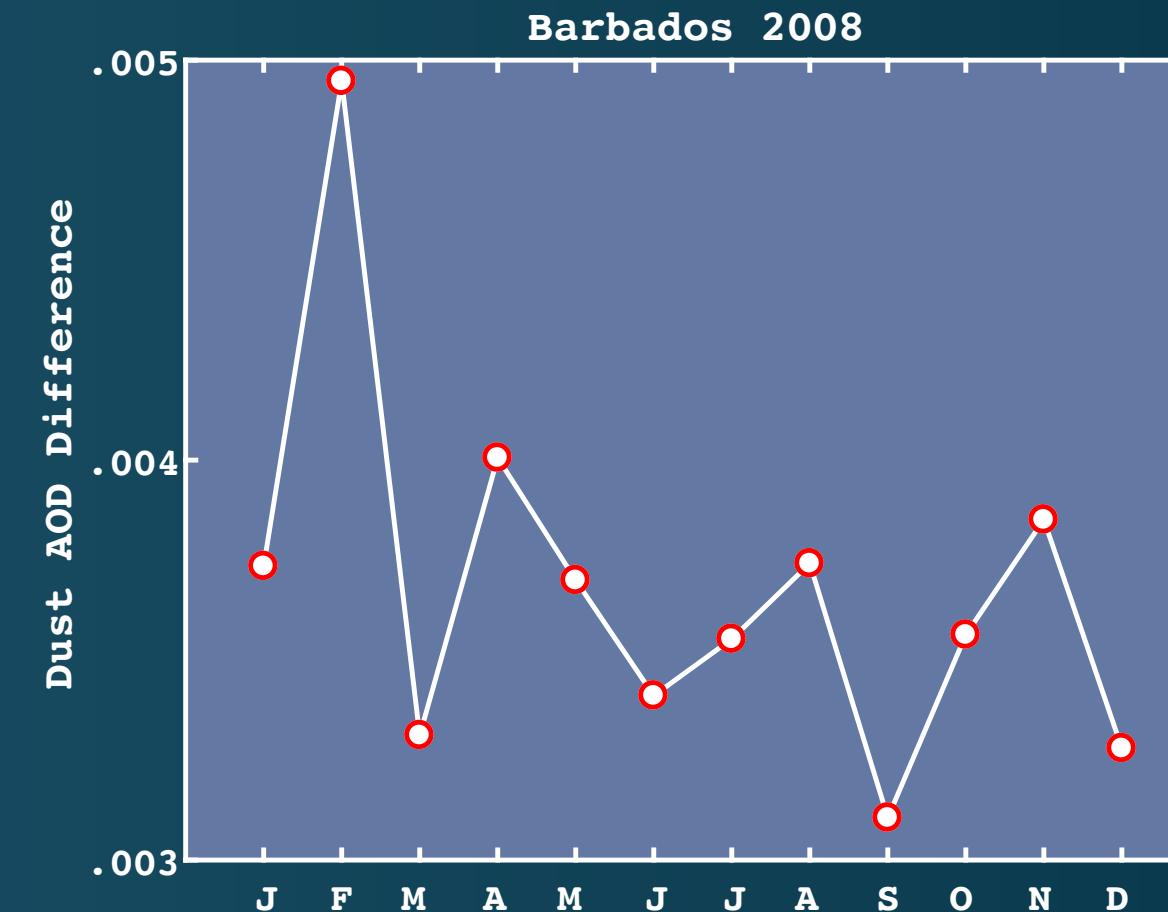
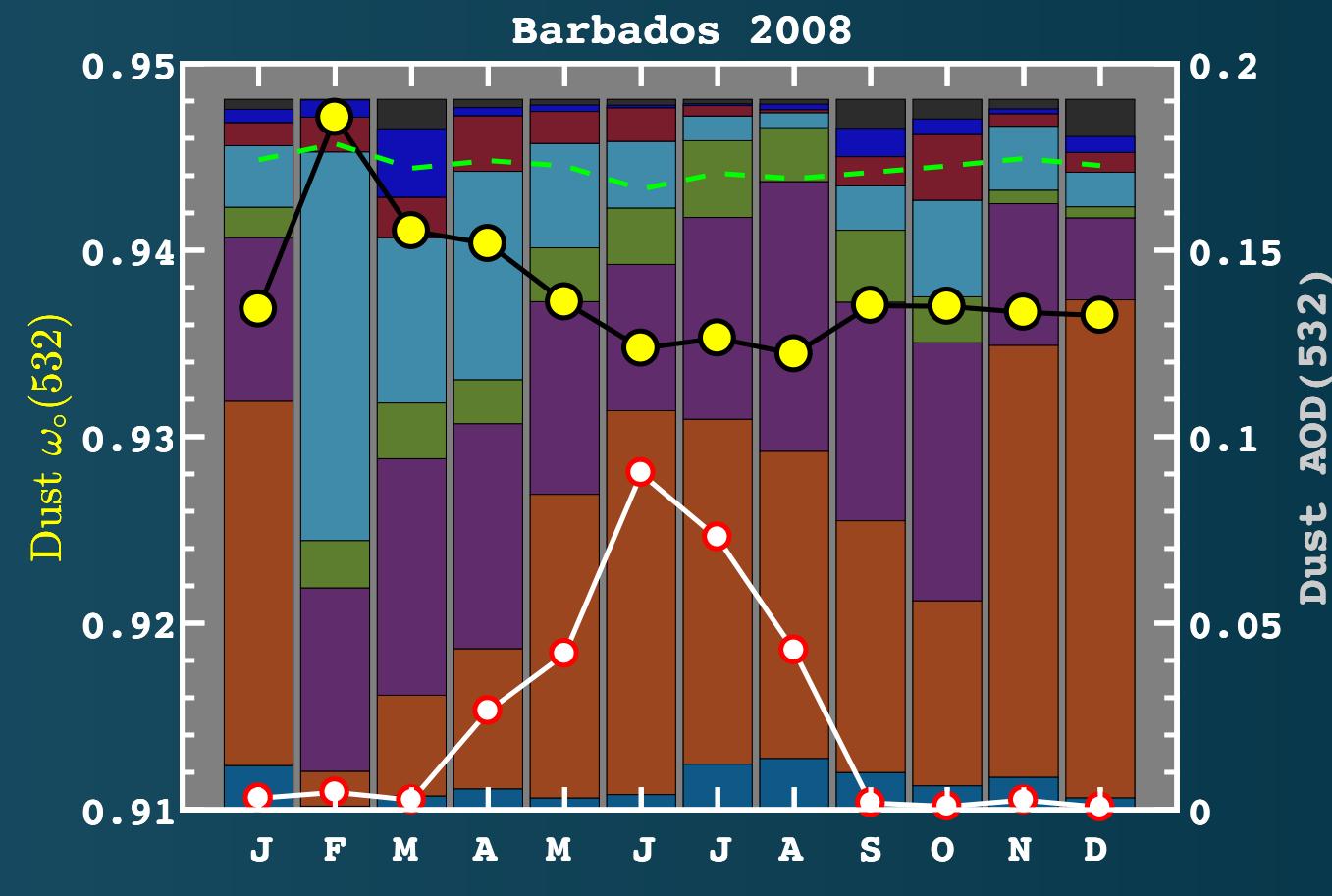
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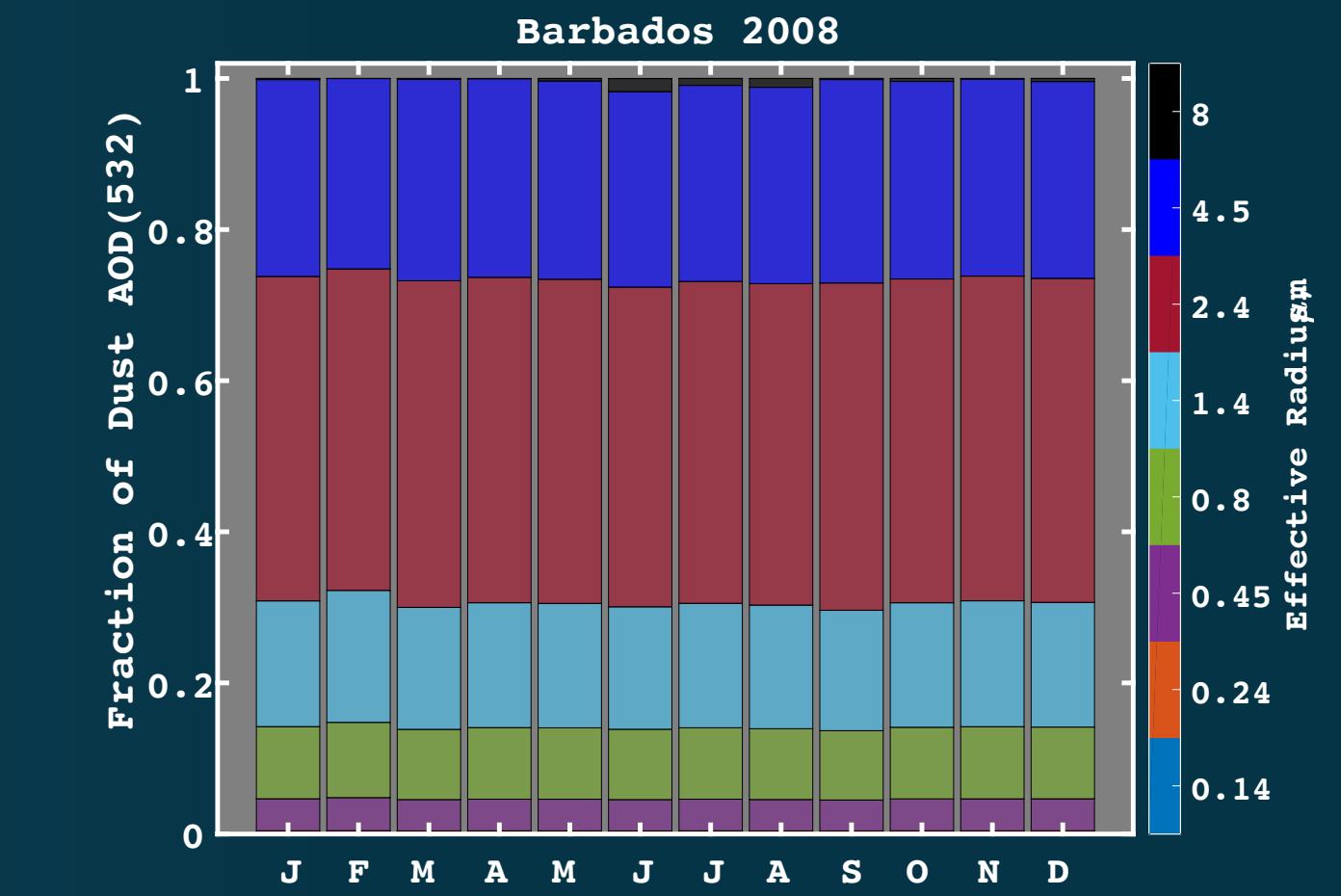
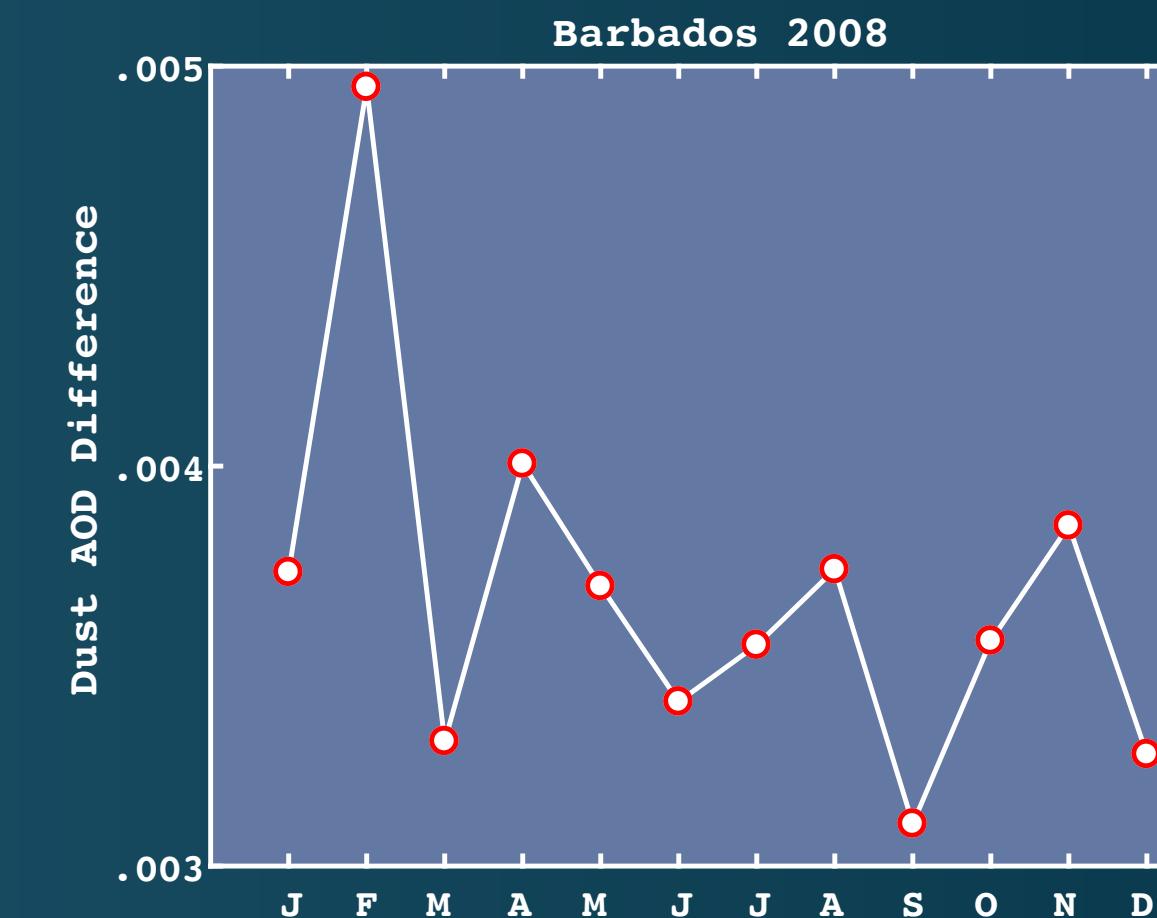
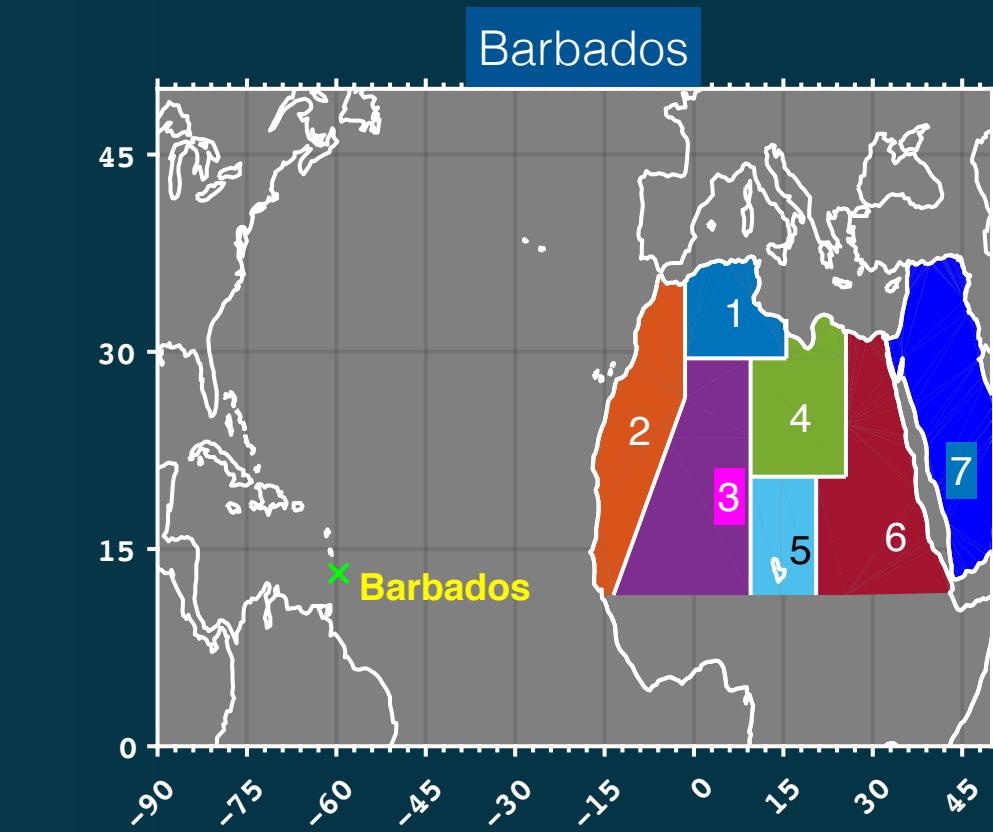
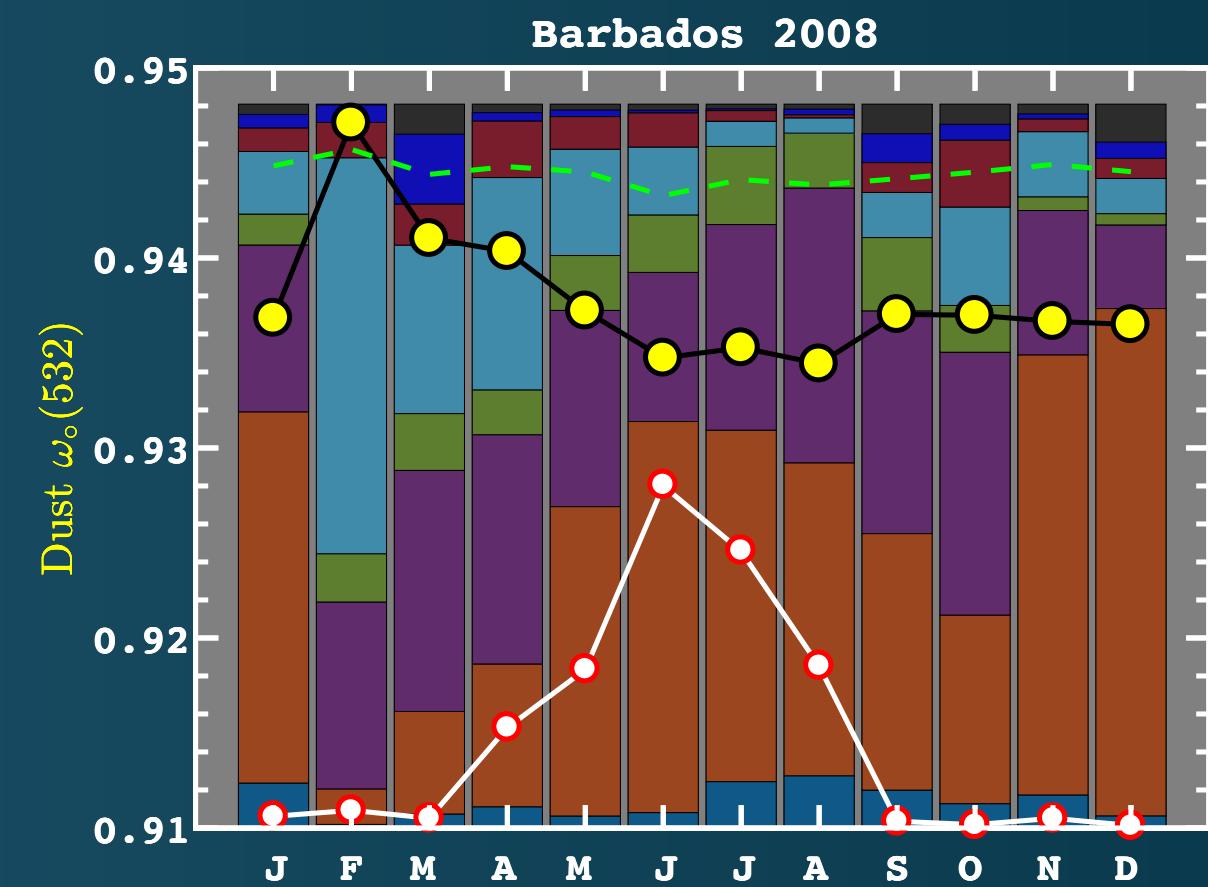
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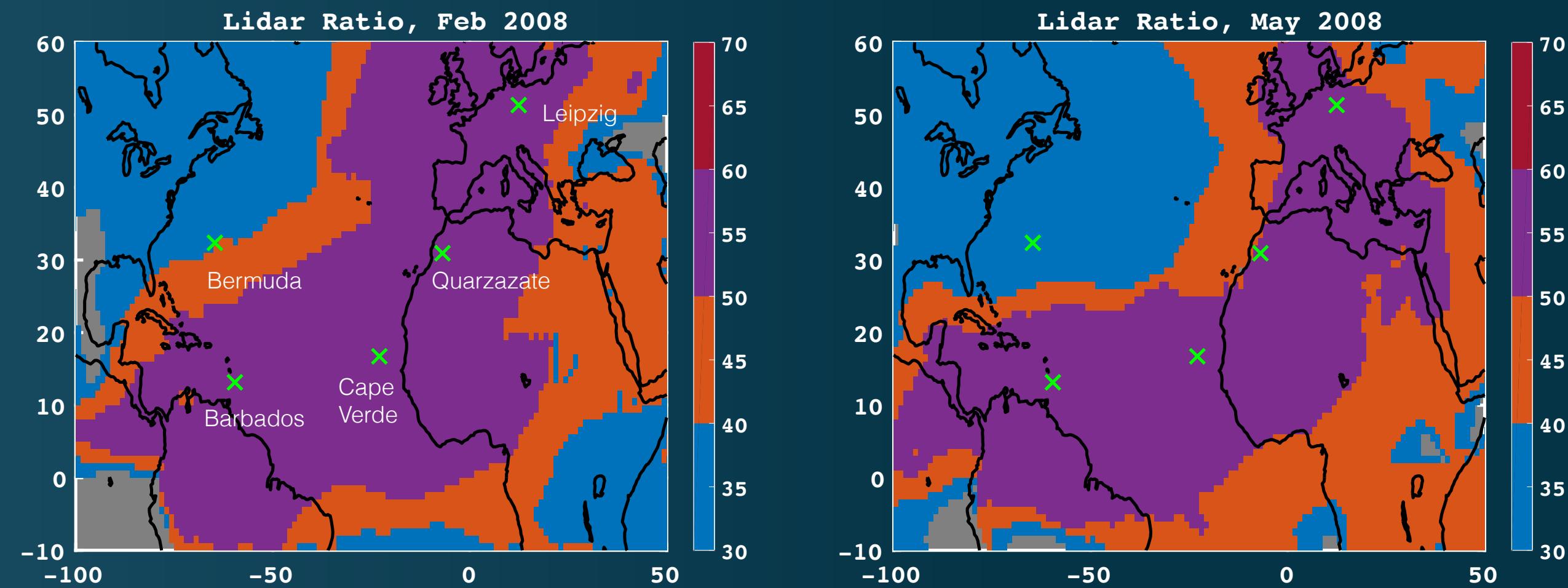
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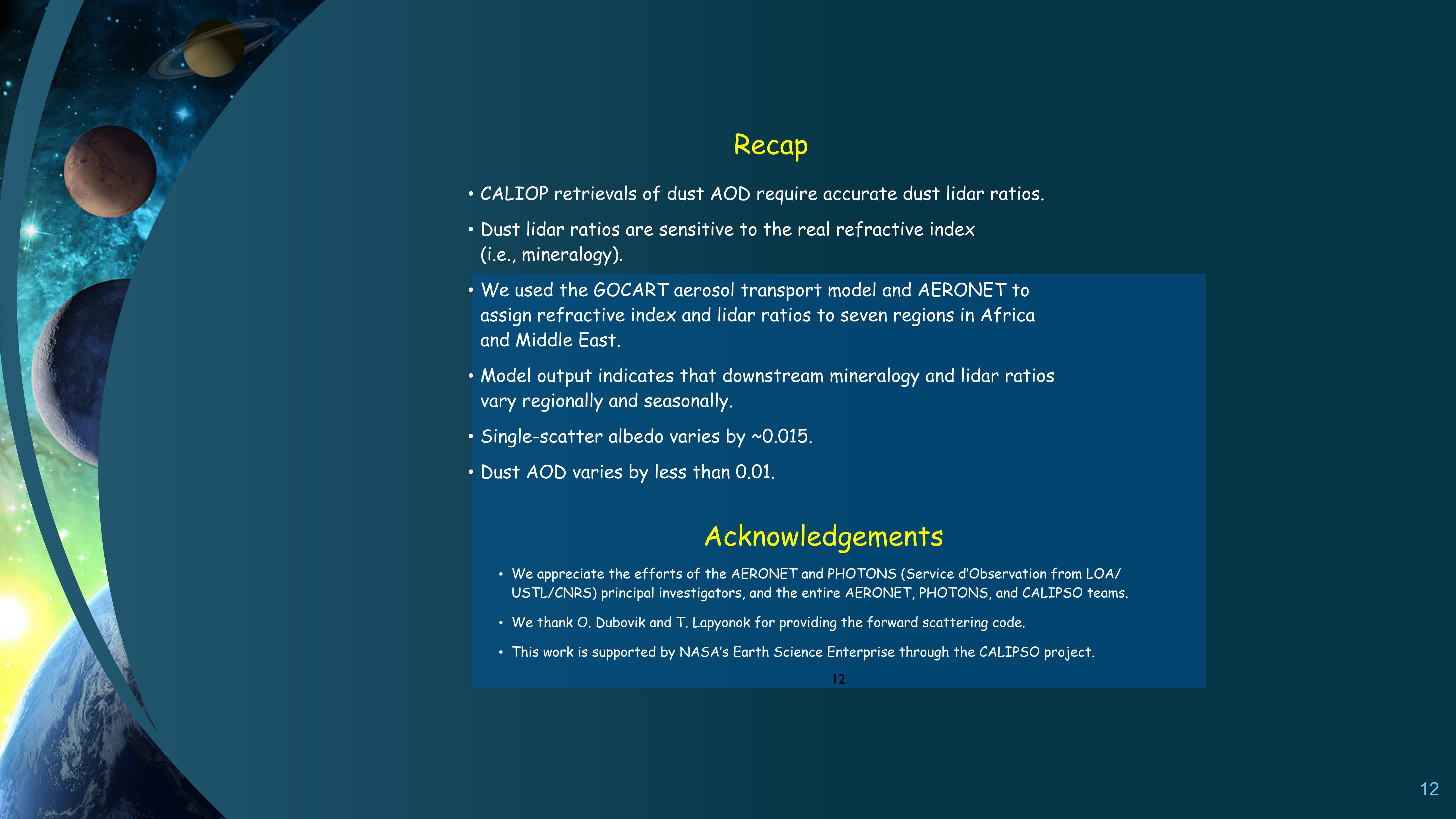
Source region impact on dust single-scatter albedo and AOD



Regional Variability of the Lidar Ratio for Dust



- SAMUM -1 took place near Quarzazate (Morocco) in May/June 2006
- SAMUM-2 took place in Cape Verde in Jan/Feb and May/June 2008
- SALTRACE took place in Barbados in Aug 2013

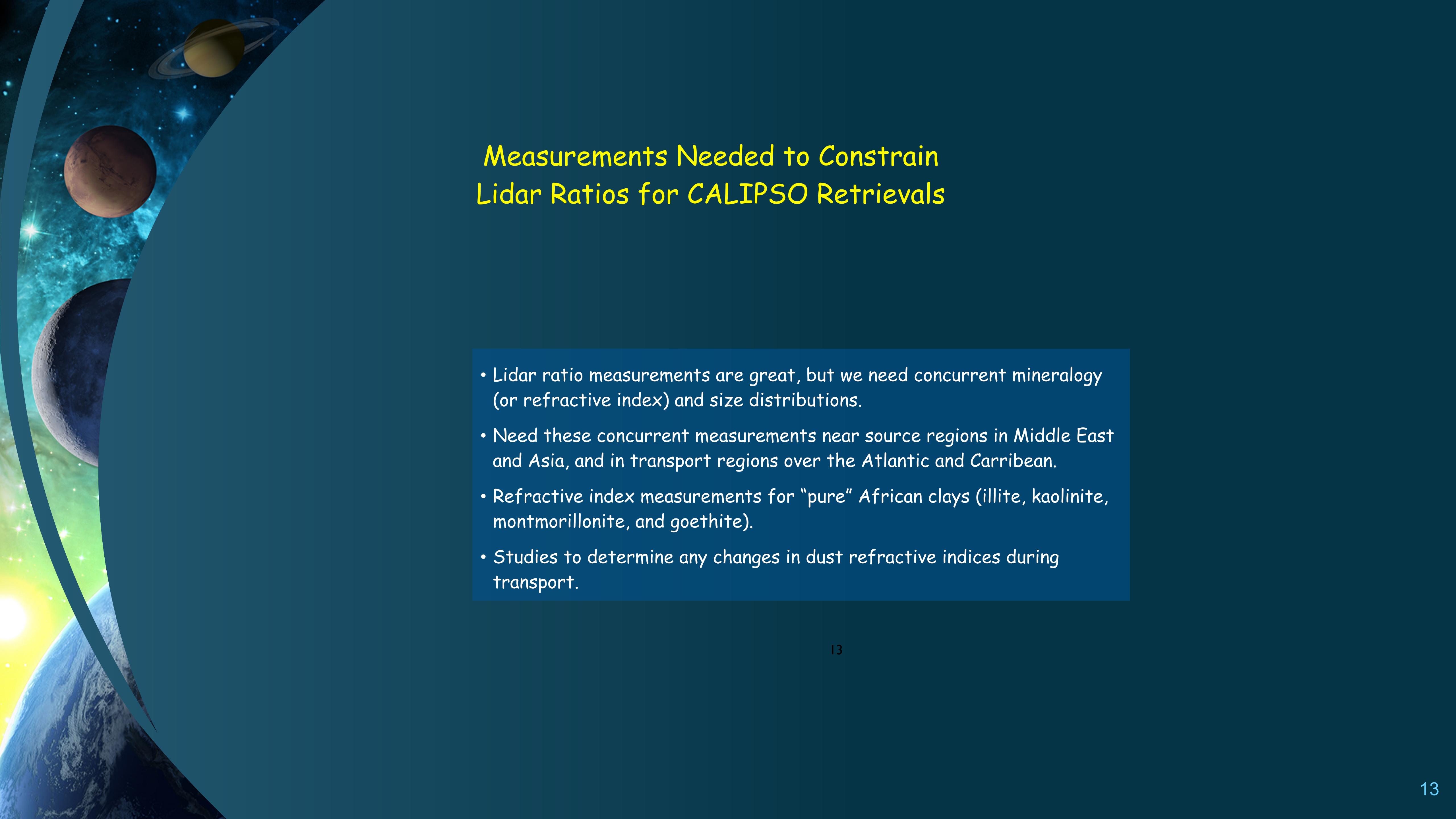


Recap

- CALIOP retrievals of dust AOD require accurate dust lidar ratios.
- Dust lidar ratios are sensitive to the real refractive index (i.e., mineralogy).
- We used the GOCART aerosol transport model and AERONET to assign refractive index and lidar ratios to seven regions in Africa and Middle East.
- Model output indicates that downstream mineralogy and lidar ratios vary regionally and seasonally.
- Single-scatter albedo varies by ~0.015.
- Dust AOD varies by less than 0.01.

Acknowledgements

- We appreciate the efforts of the AERONET and PHOTONS (Service d'Observation from LOA/USTL/CNRS) principal investigators, and the entire AERONET, PHOTONS, and CALIPSO teams.
- We thank O. Dubovik and T. Lapyonok for providing the forward scattering code.
- This work is supported by NASA's Earth Science Enterprise through the CALIPSO project.

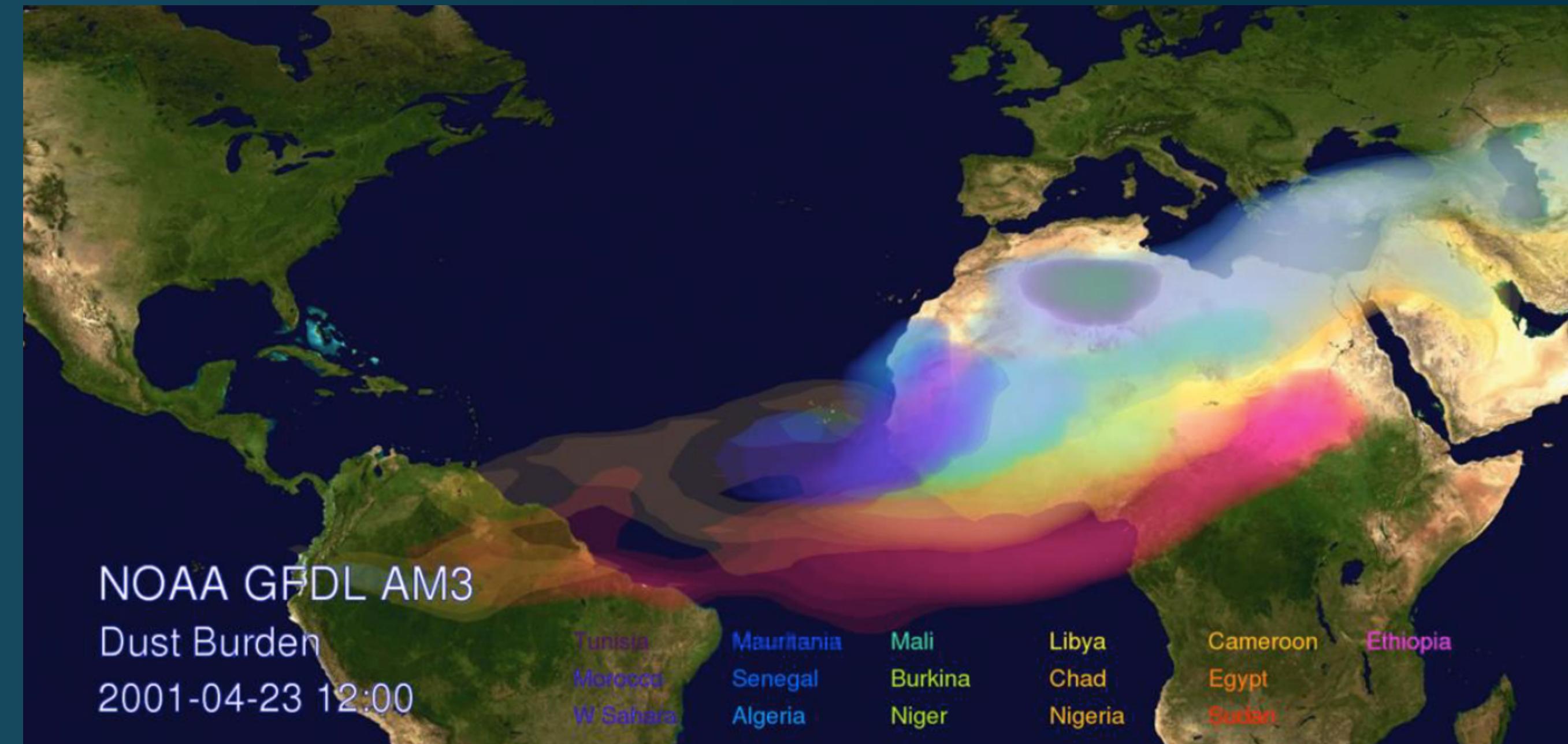


Measurements Needed to Constrain Lidar Ratios for CALIPSO Retrievals

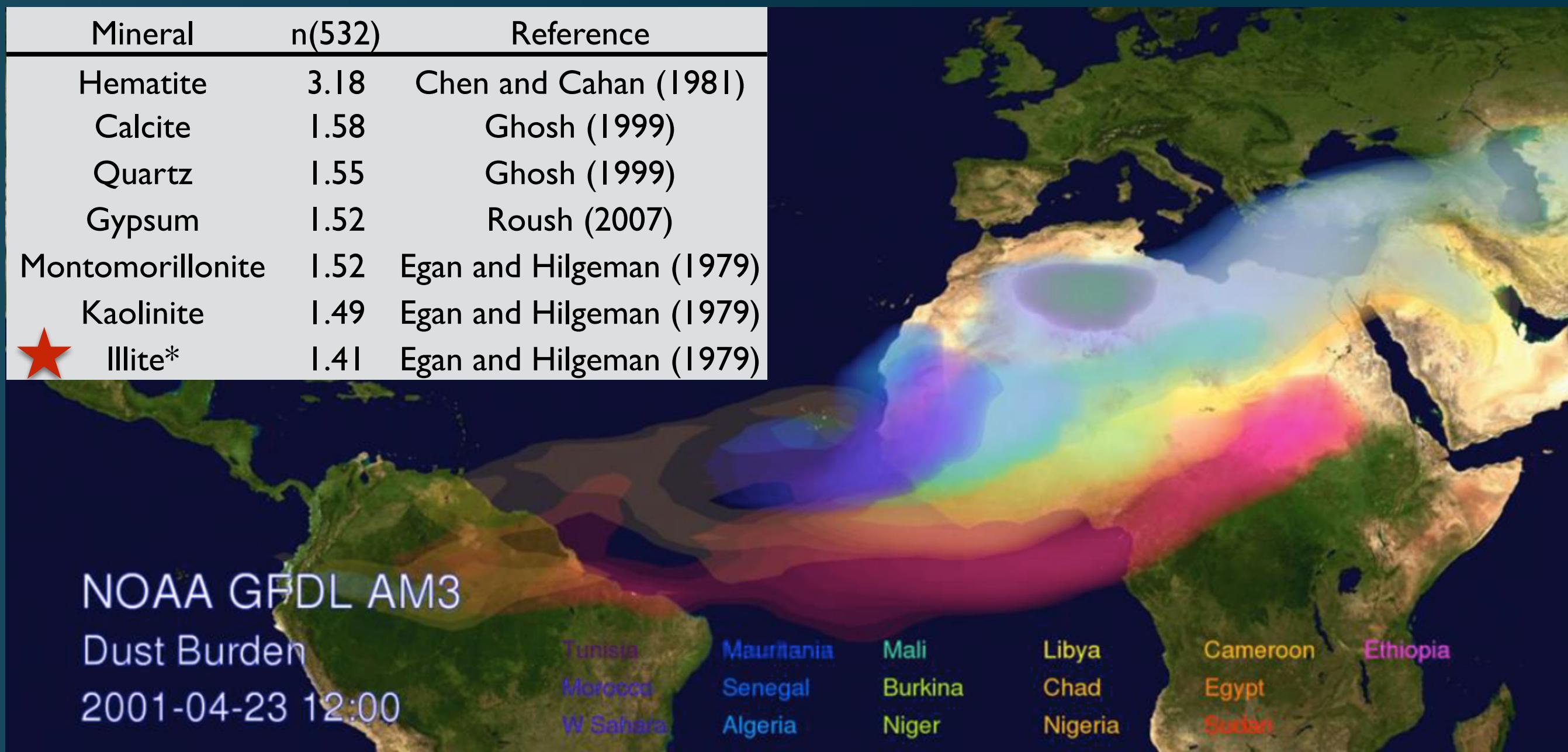
- Lidar ratio measurements are great, but we need concurrent mineralogy (or refractive index) and size distributions.
- Need these concurrent measurements near source regions in Middle East and Asia, and in transport regions over the Atlantic and Caribbean.
- Refractive index measurements for "pure" African clays (illite, kaolinite, montmorillonite, and goethite).
- Studies to determine any changes in dust refractive indices during transport.

Appendix





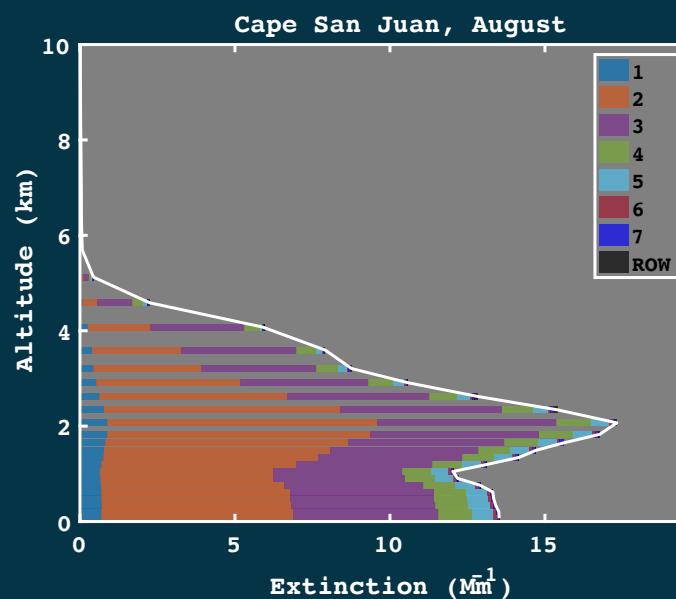
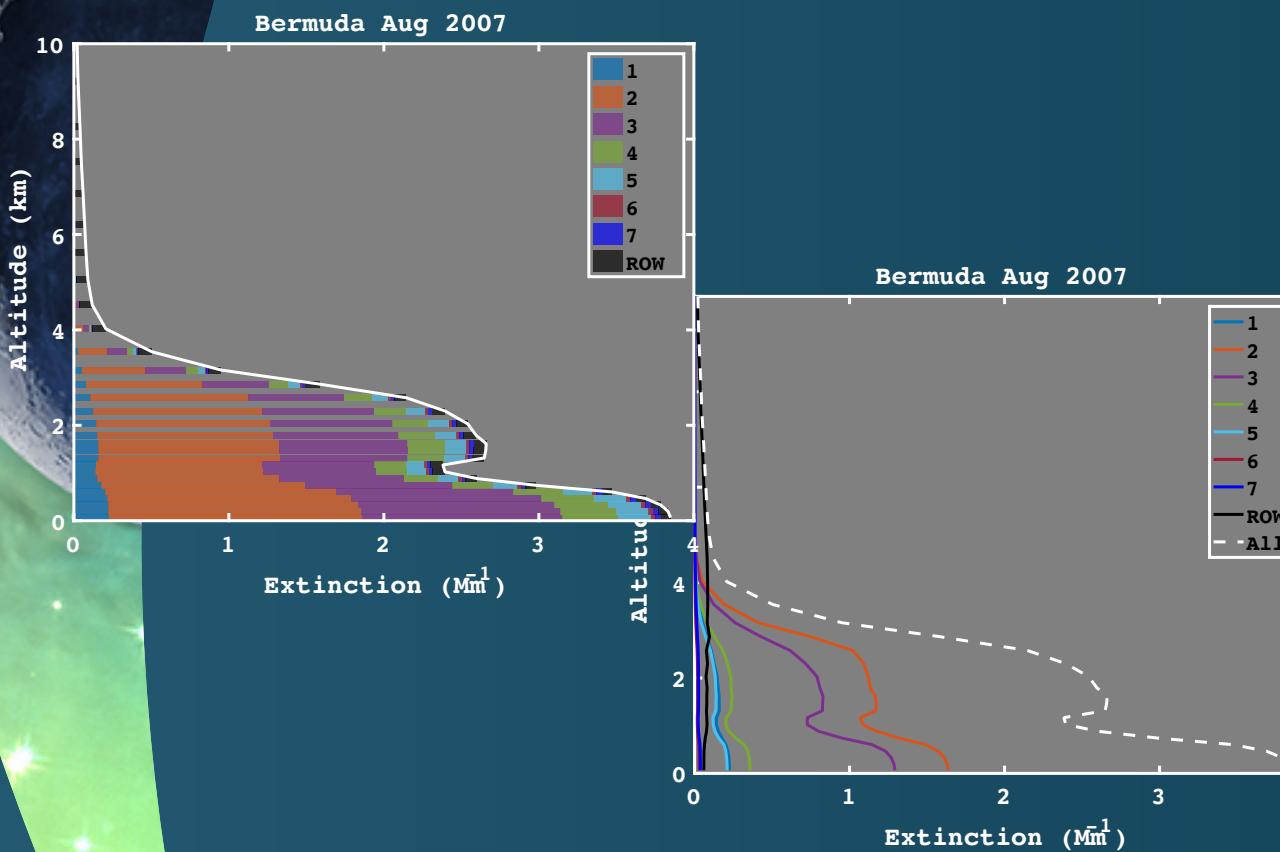
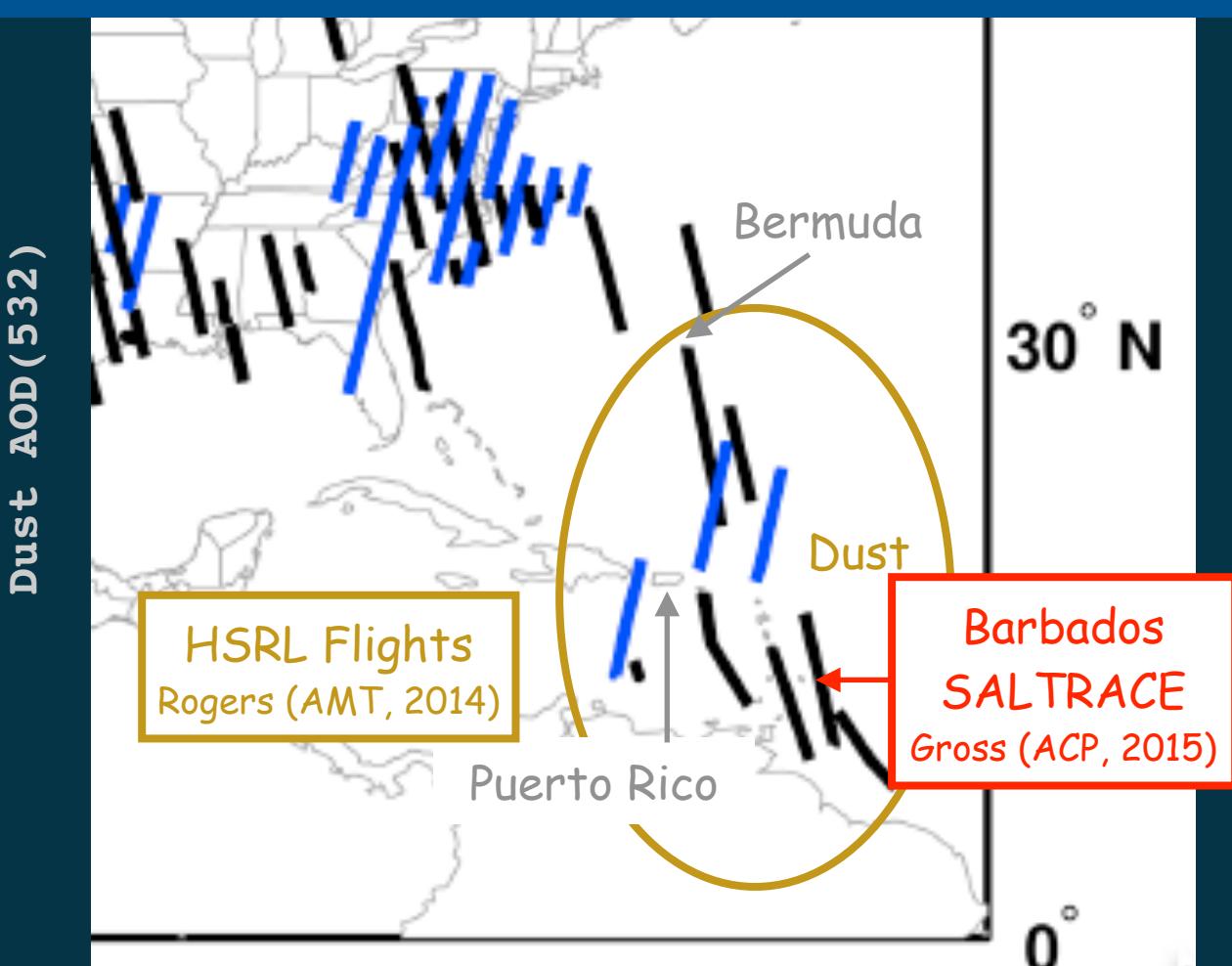
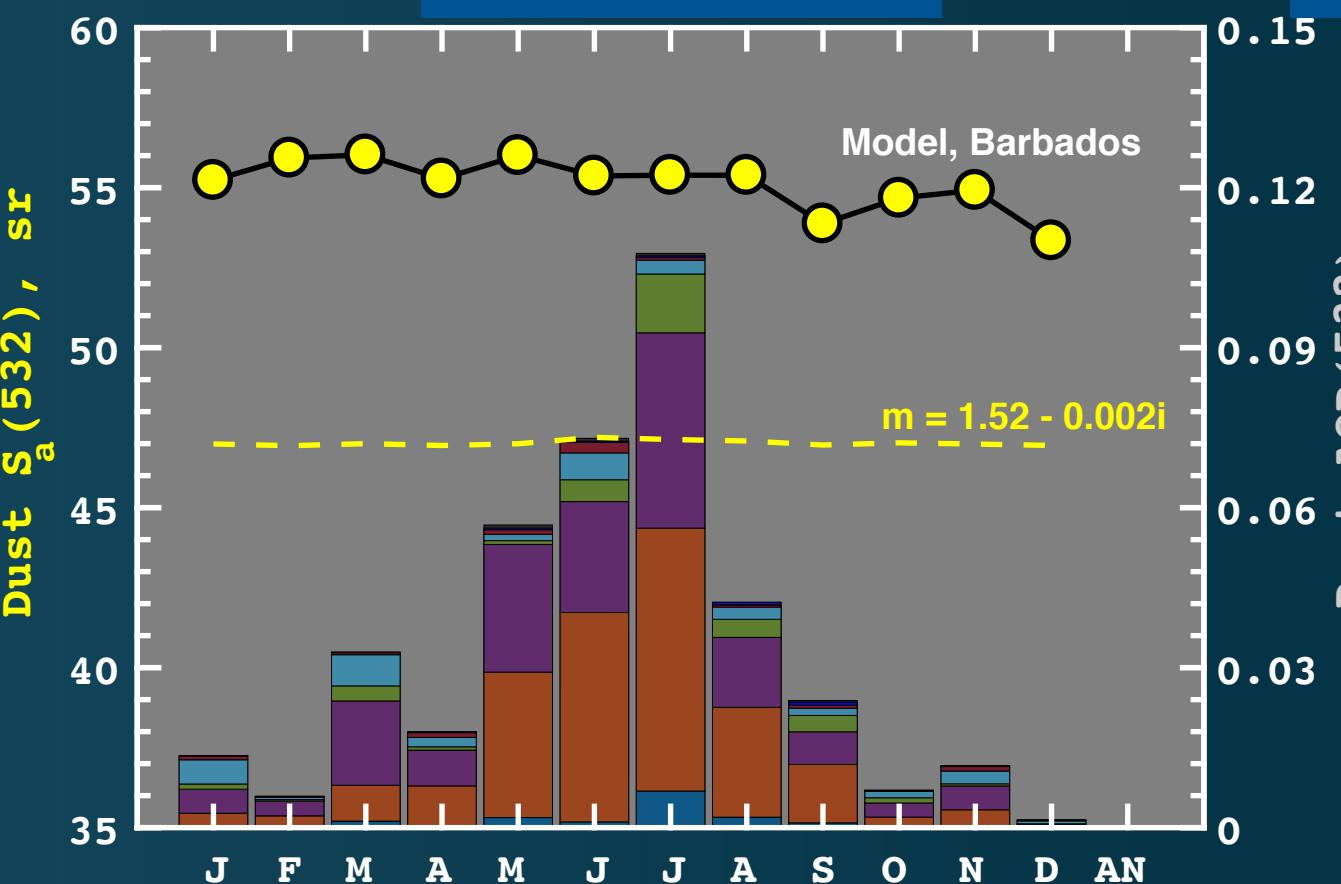
African dust sources presented by Joe Prospero at DUST 2014. Model run by Paul Ginoux



African dust sources presented by Joe Prospero at DUST 2014. Model run by Paul Ginoux

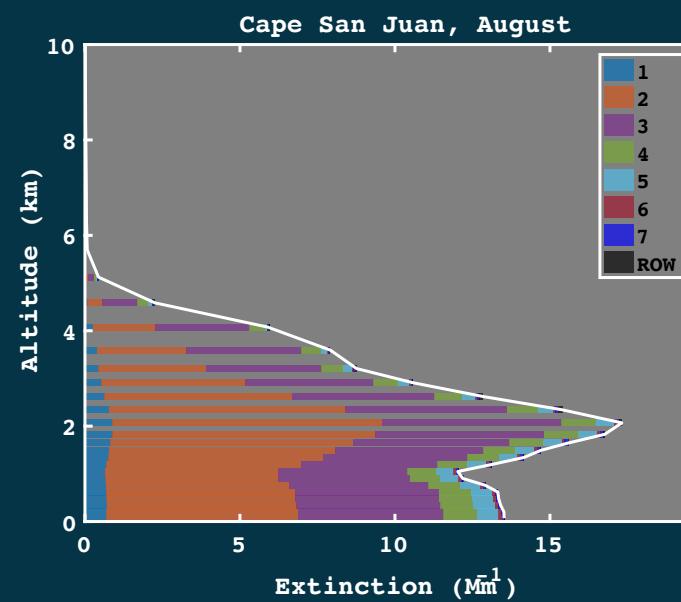
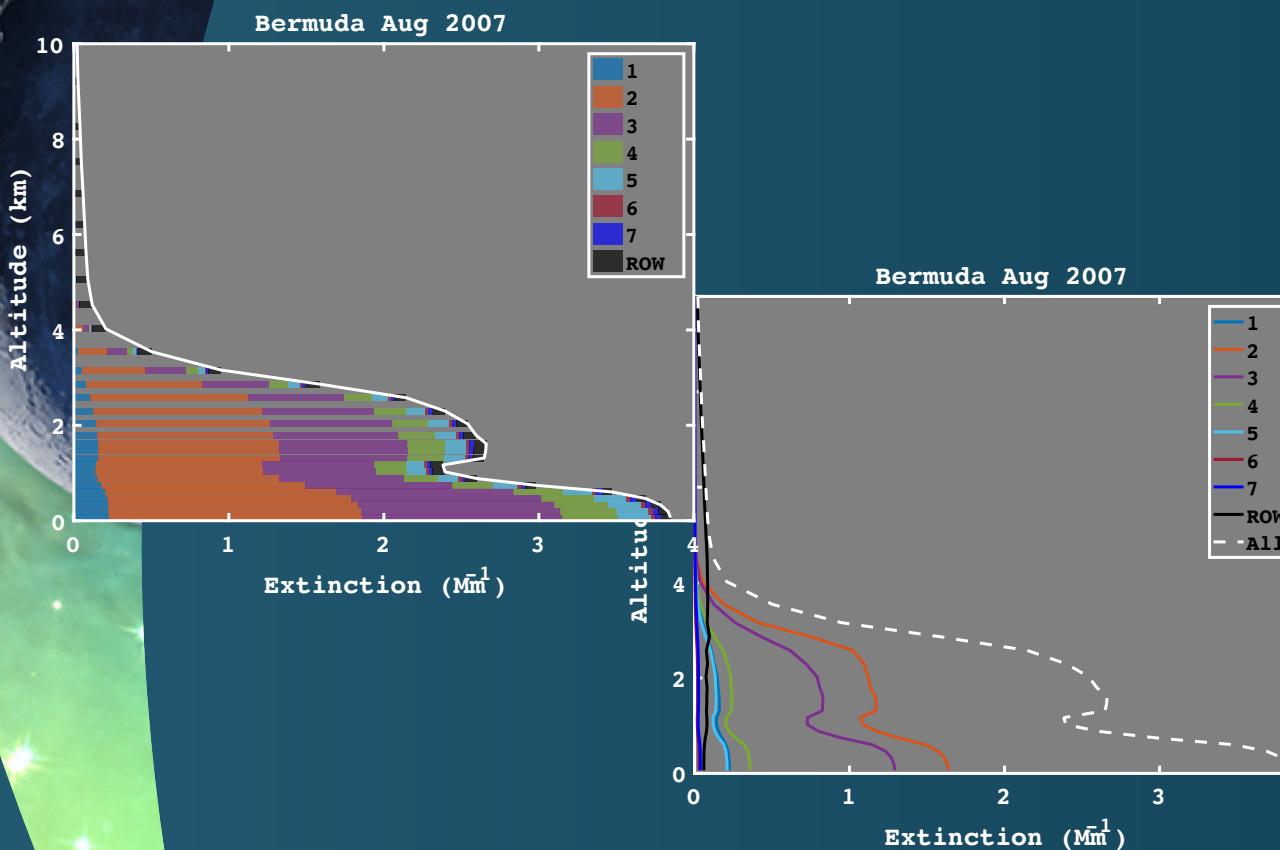
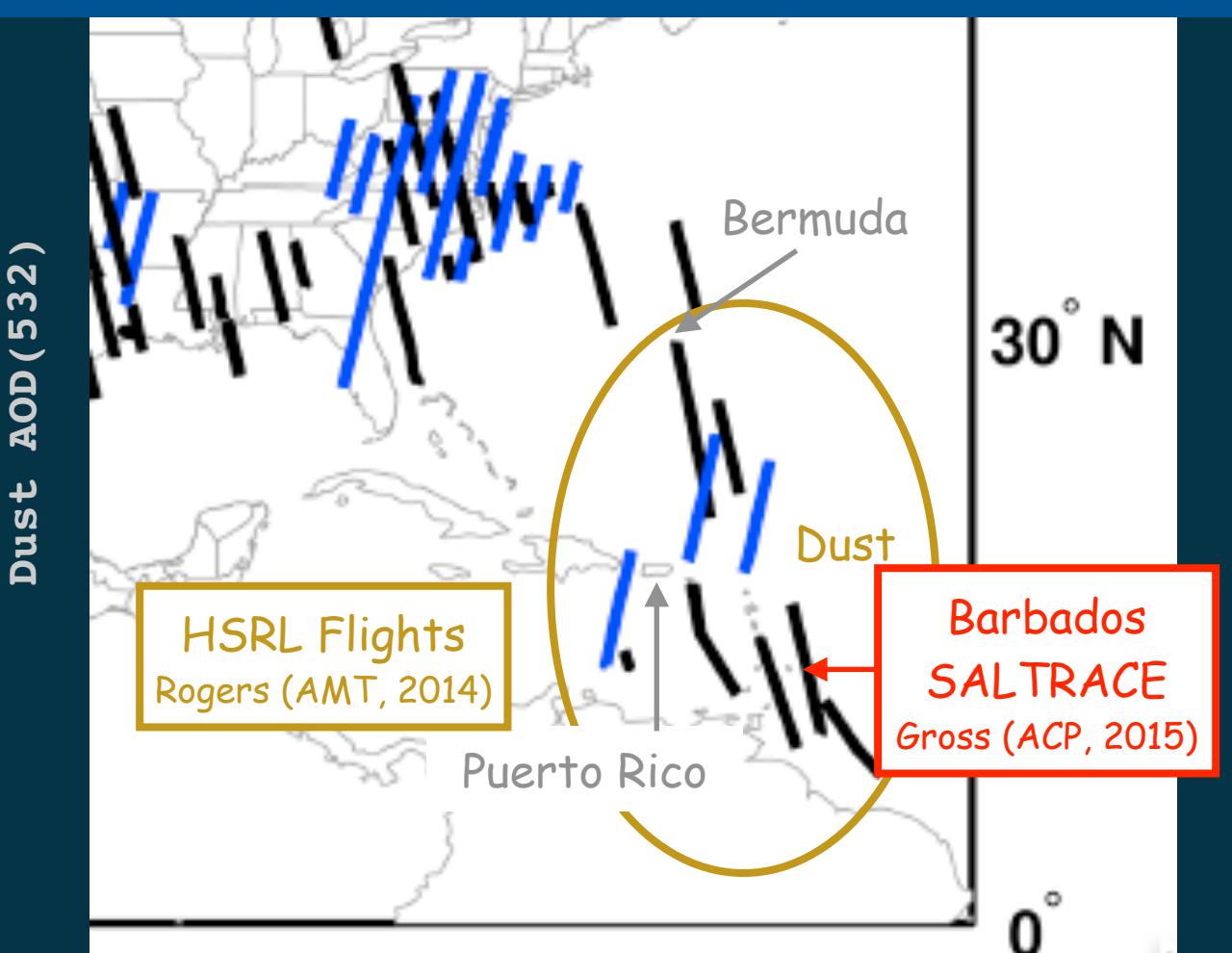
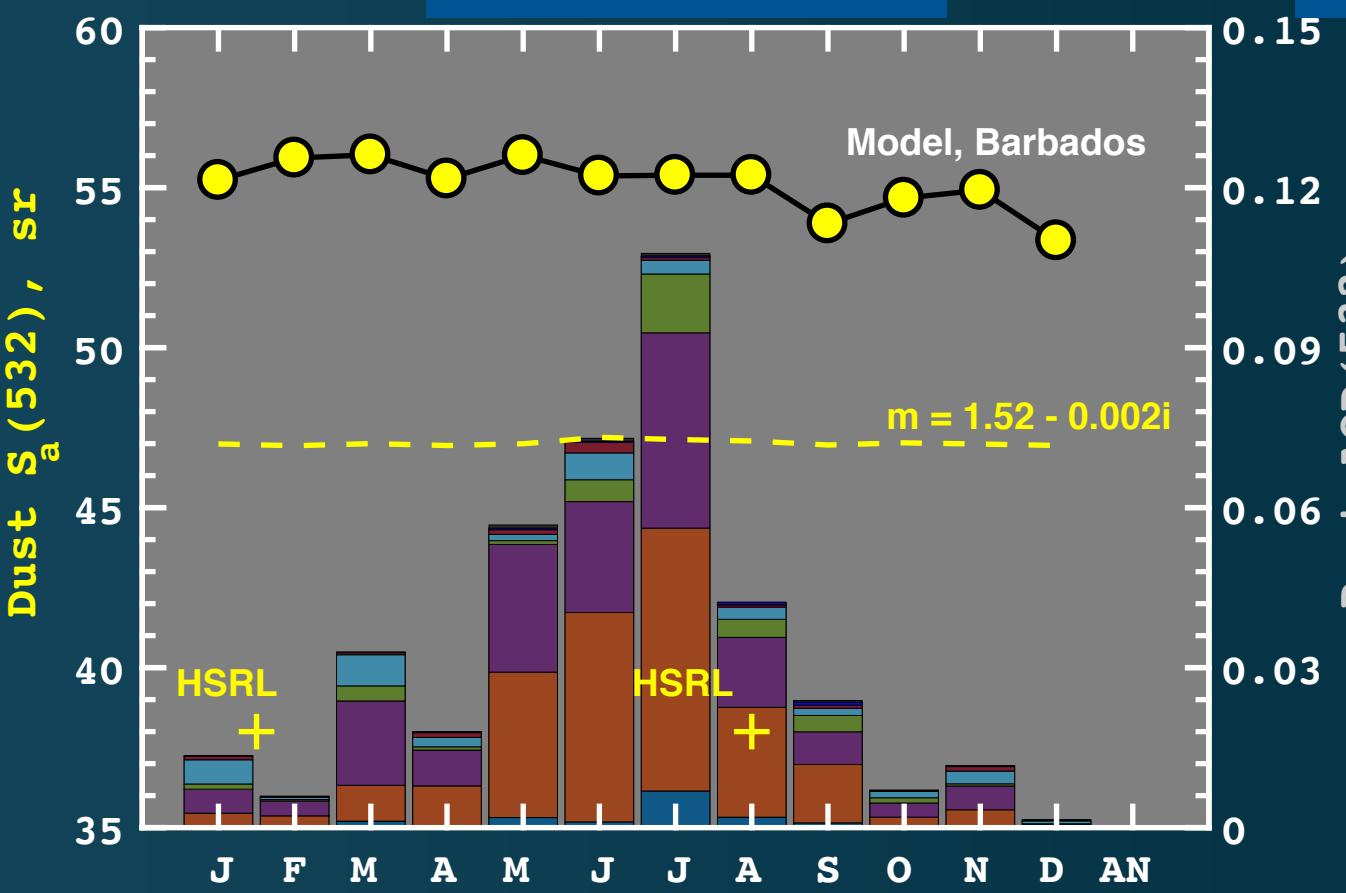
Comparison to HSRL and Raman Lidars in West Atlantic

Model Year 2007



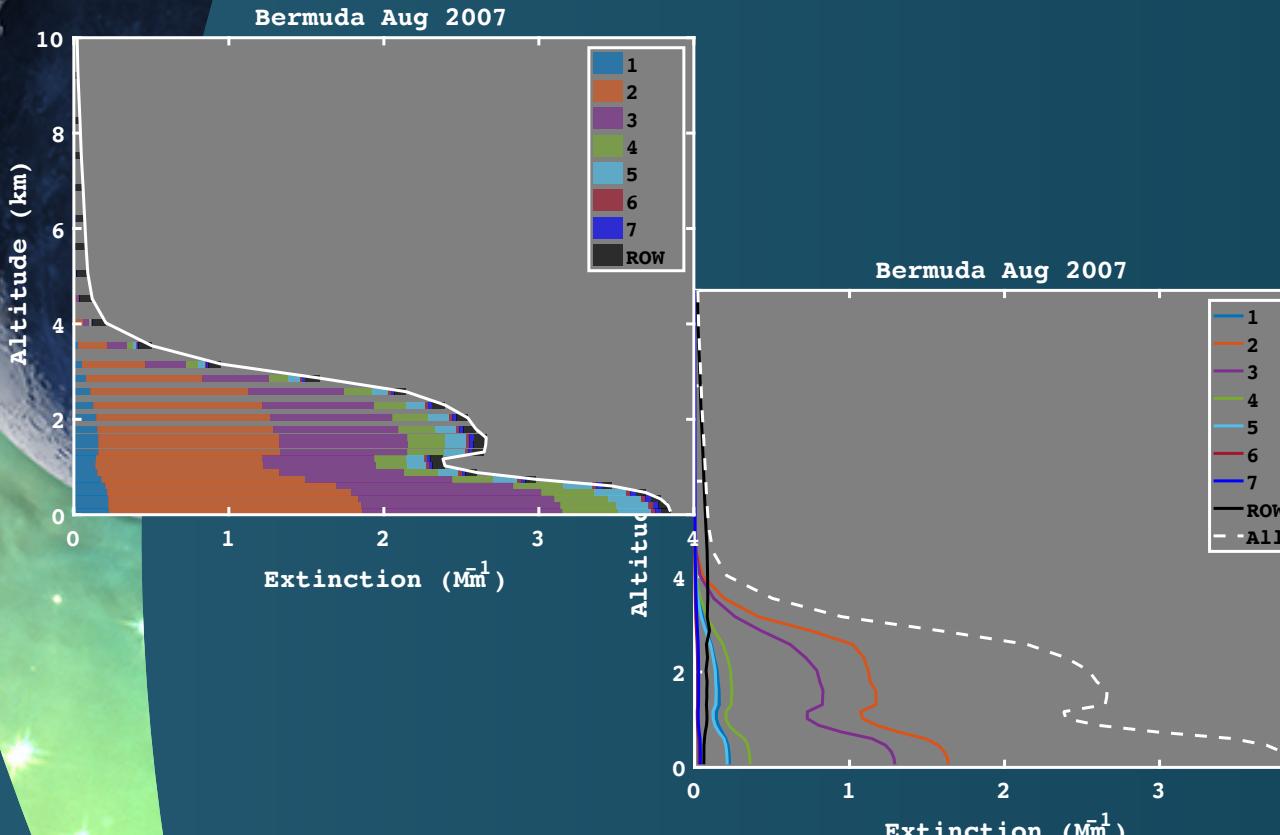
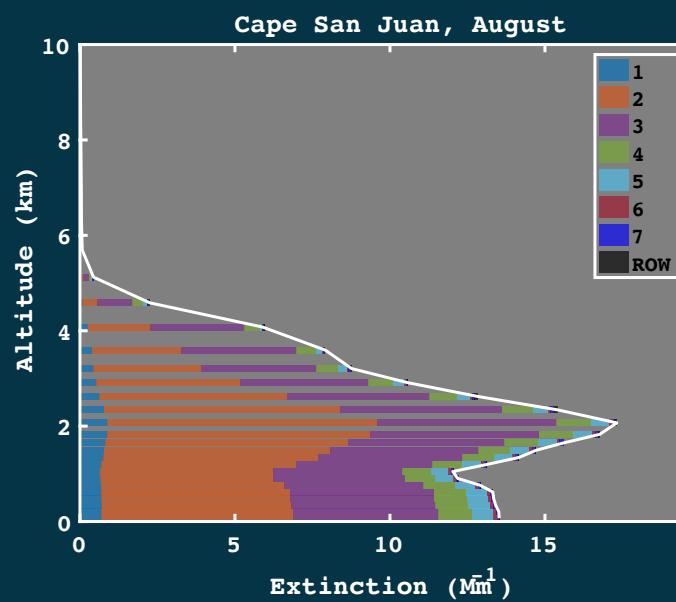
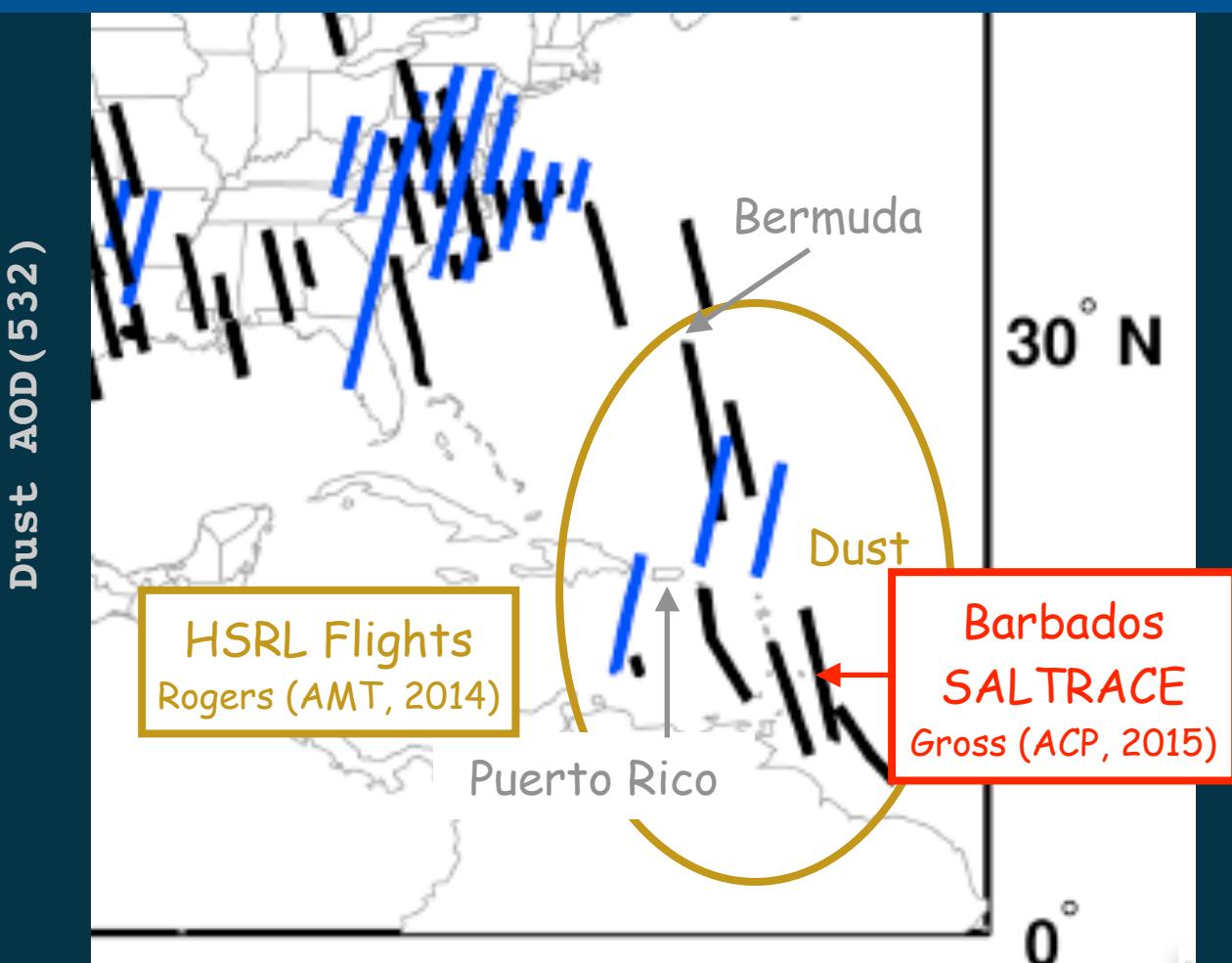
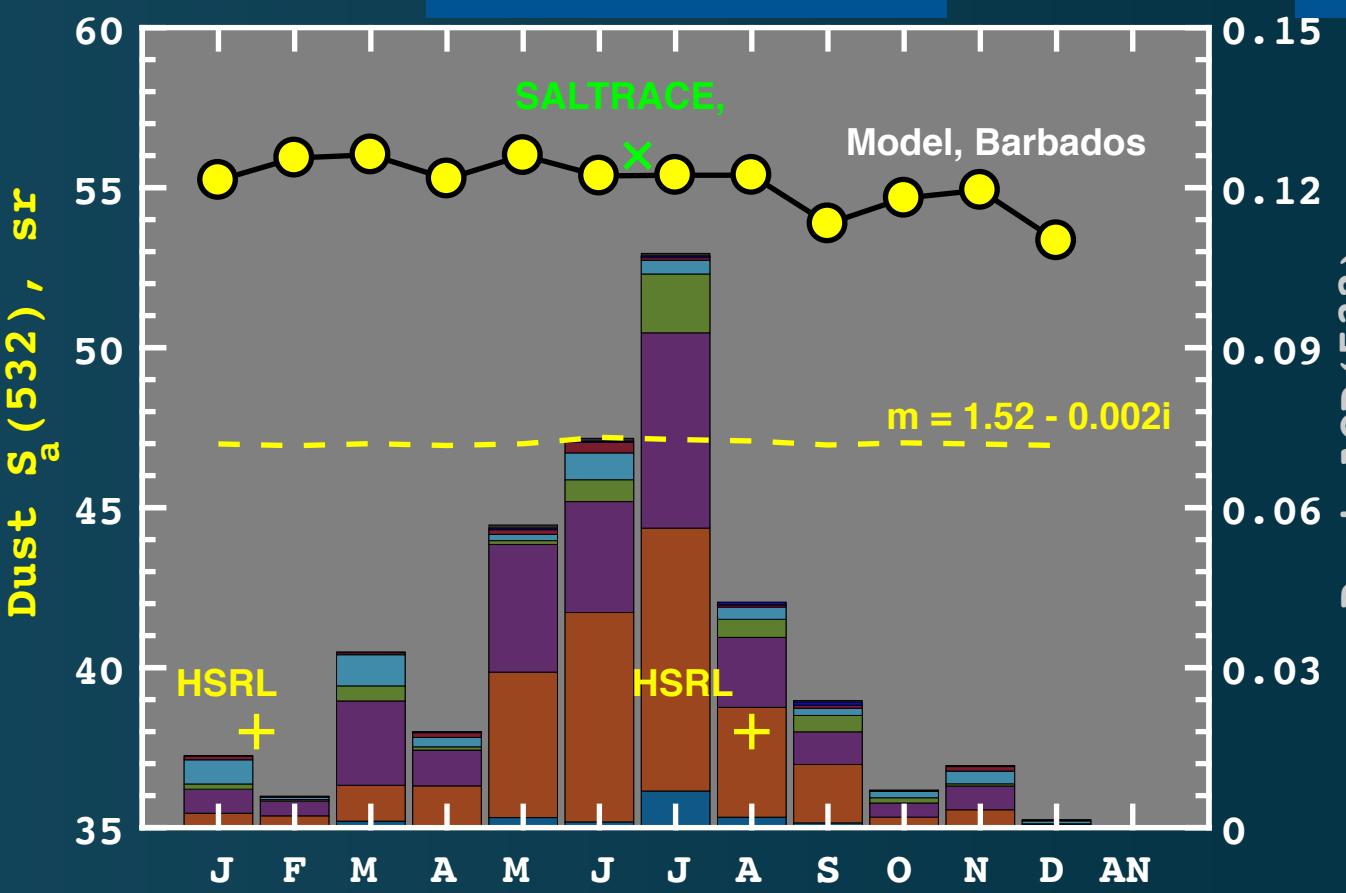
Comparison to HSRL and Raman Lidars in West Atlantic

Model Year 2007



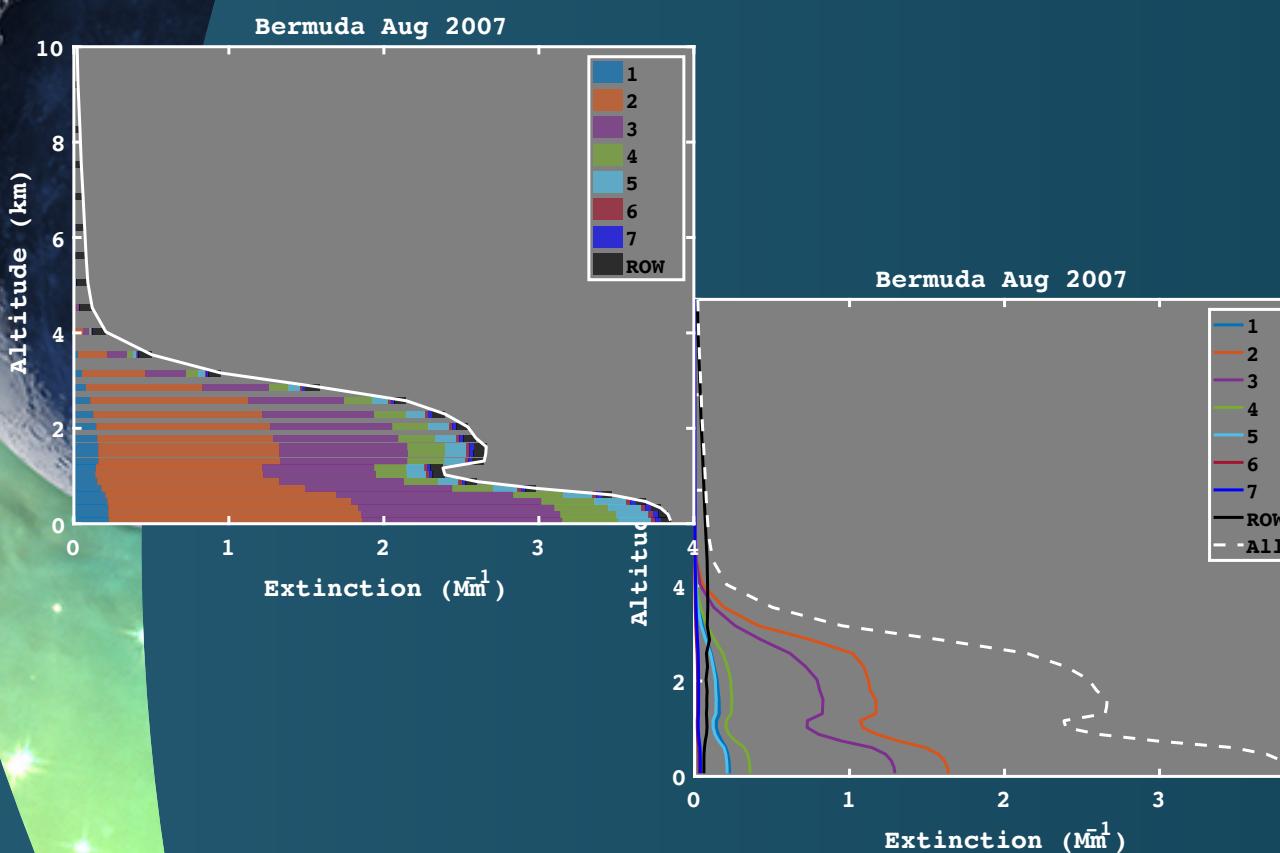
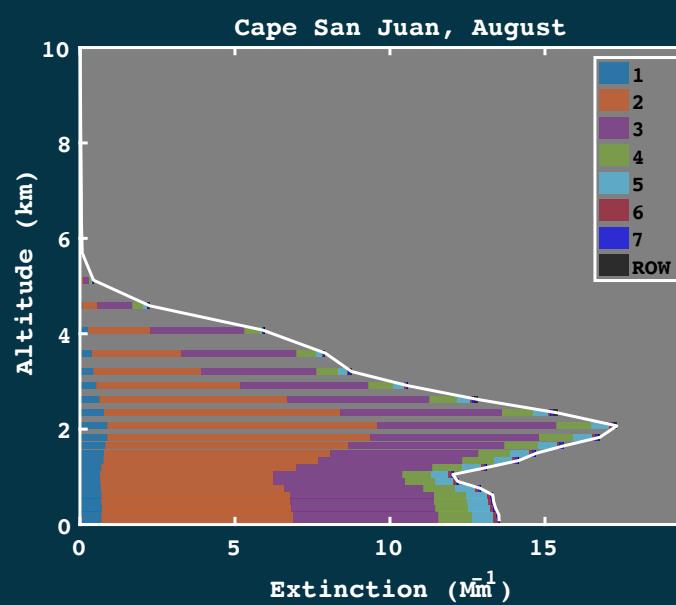
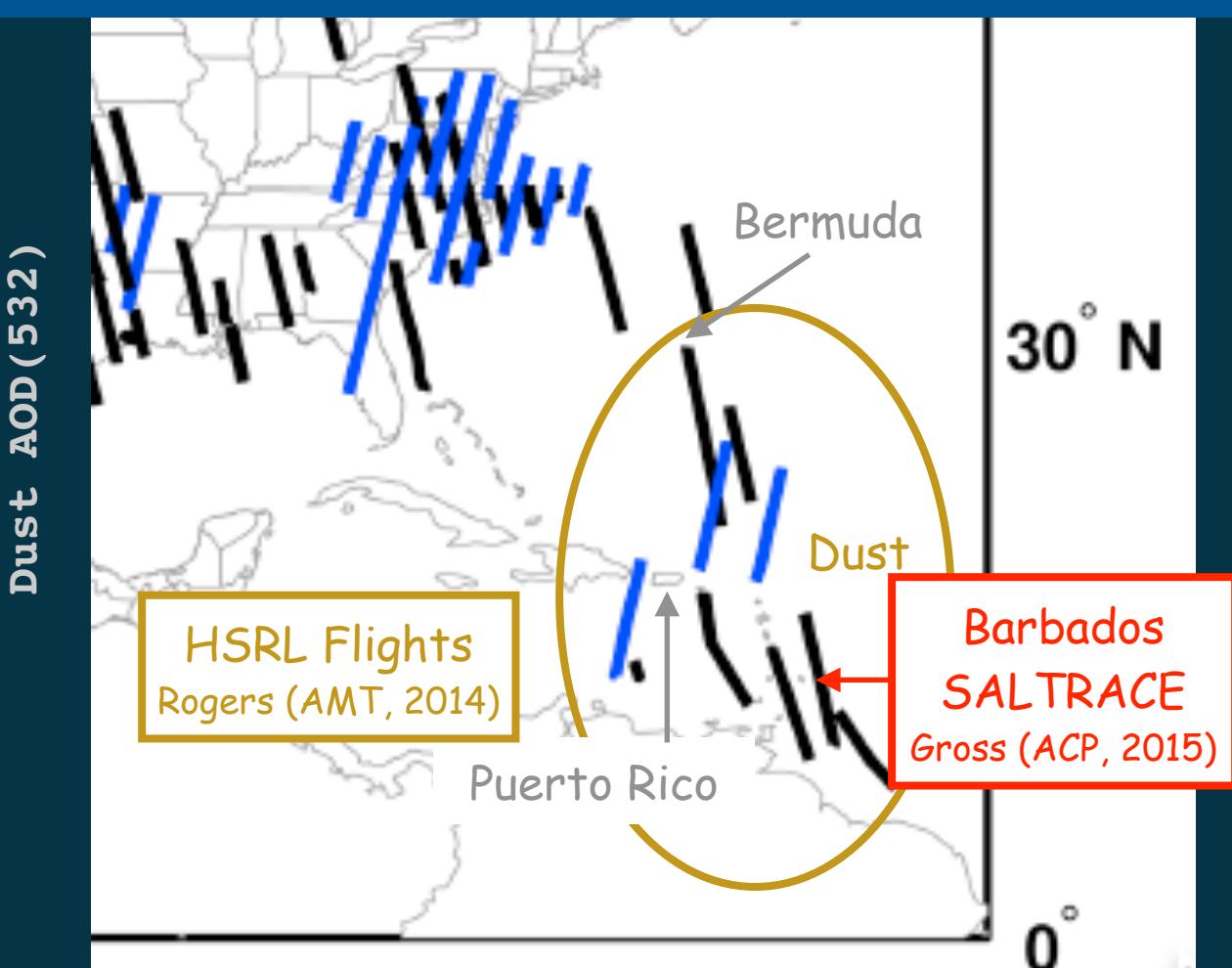
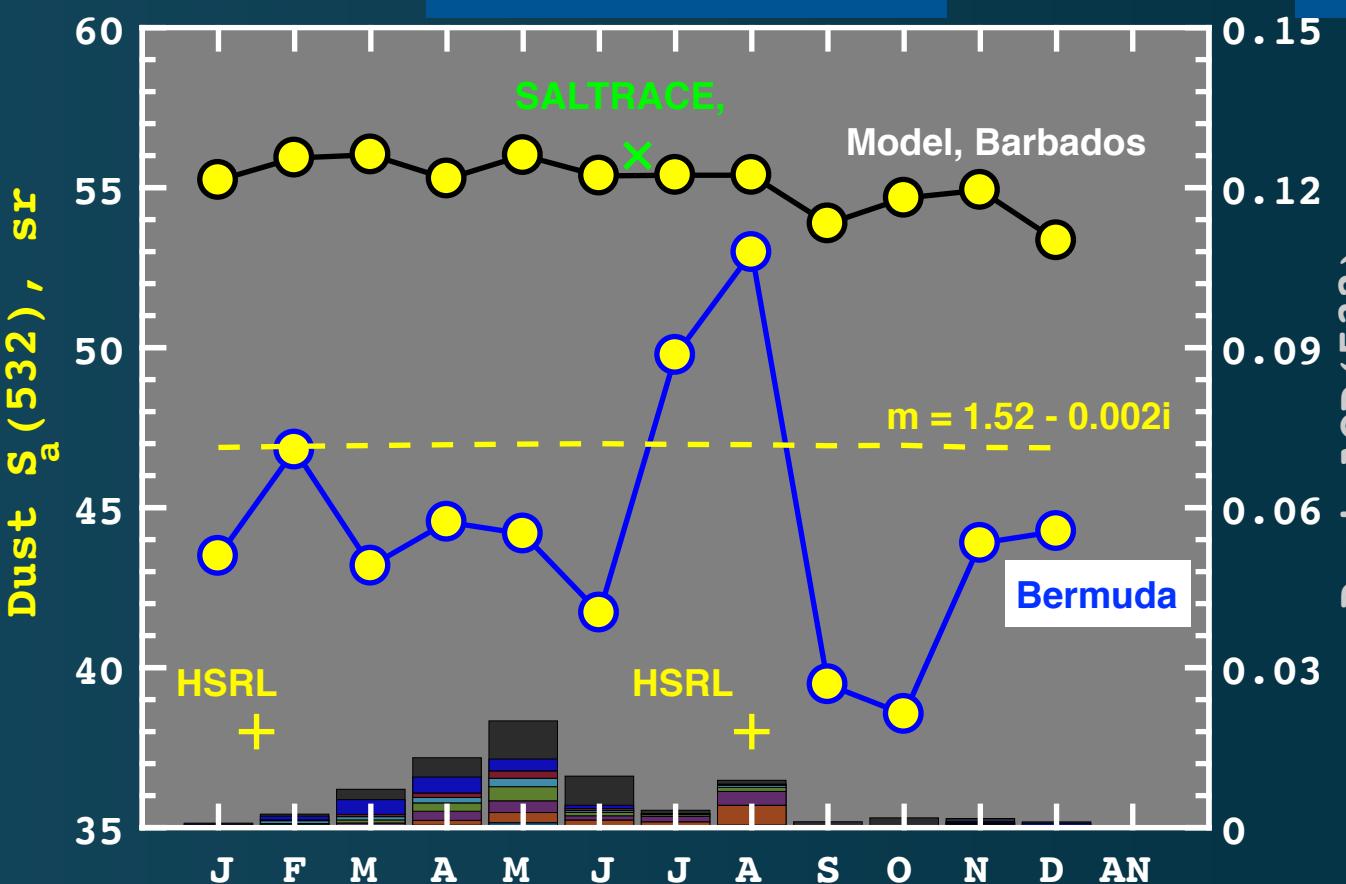
Comparison to HSRL and Raman Lidars in West Atlantic

Model Year 2007



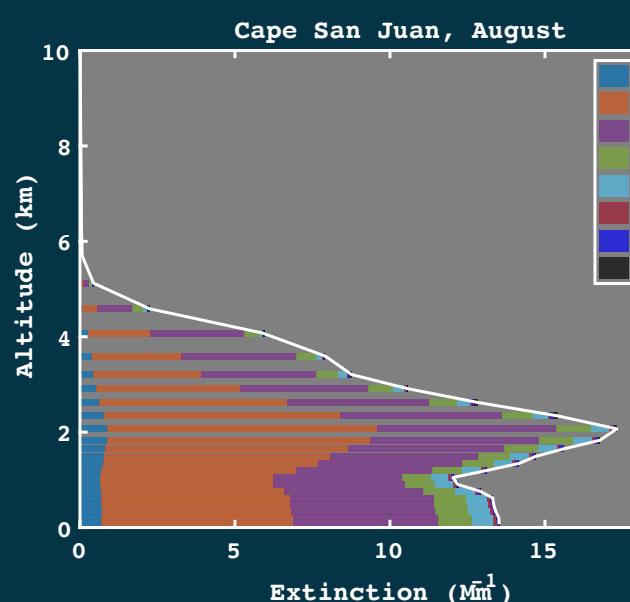
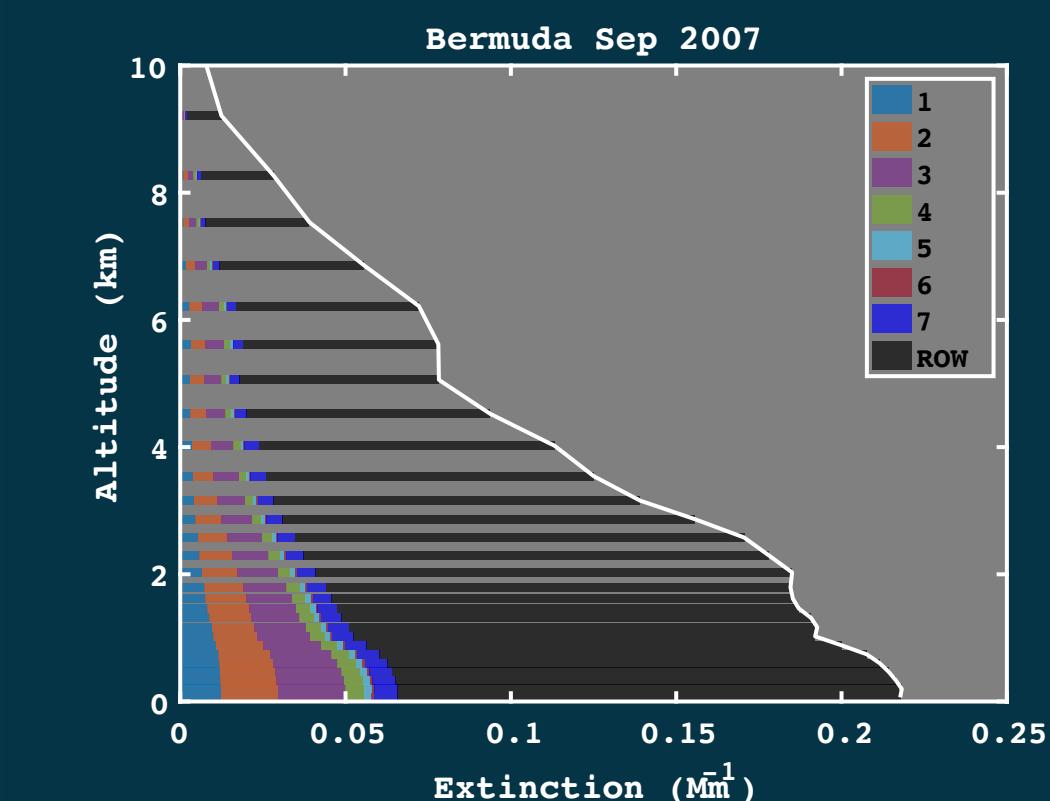
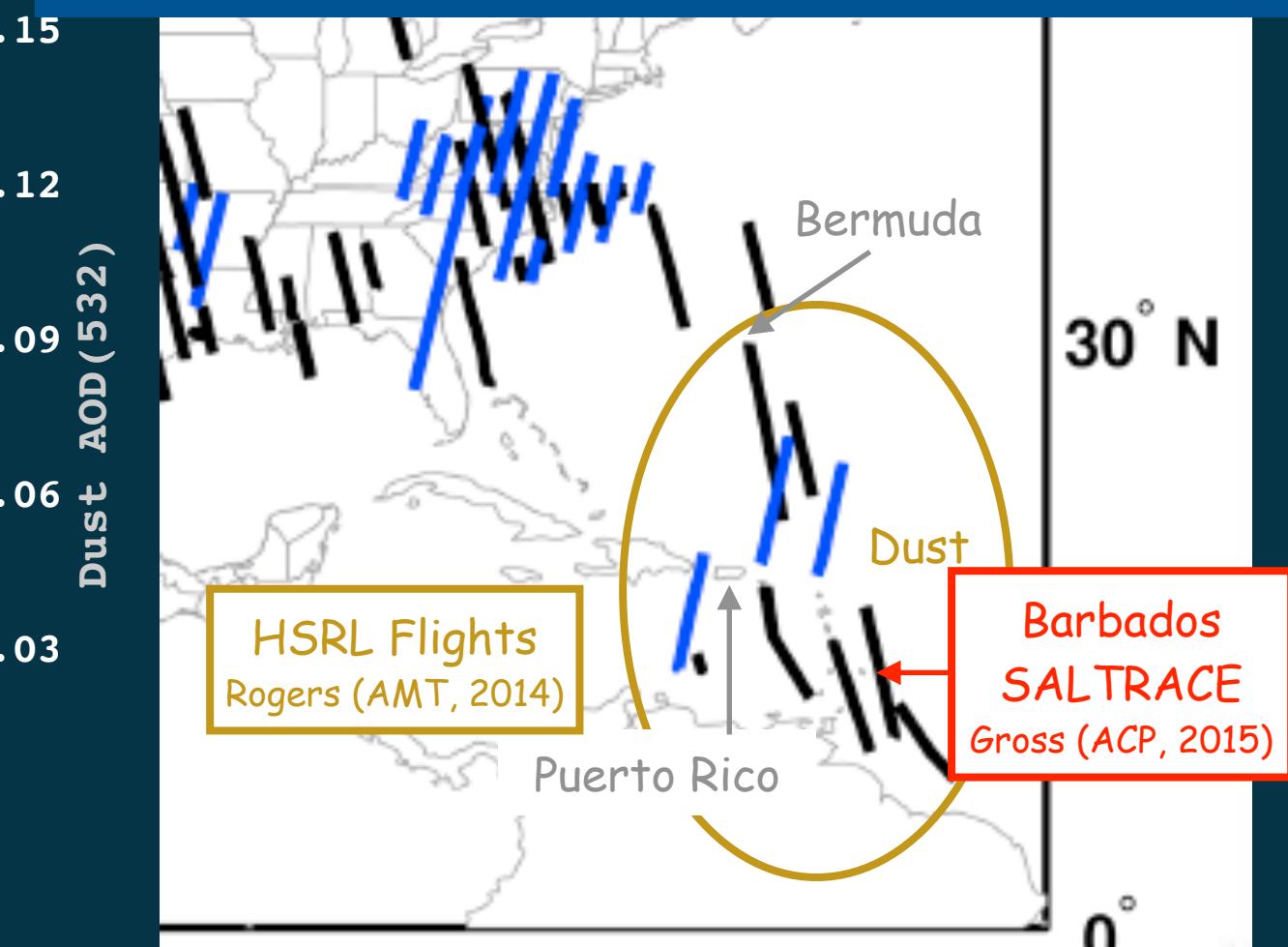
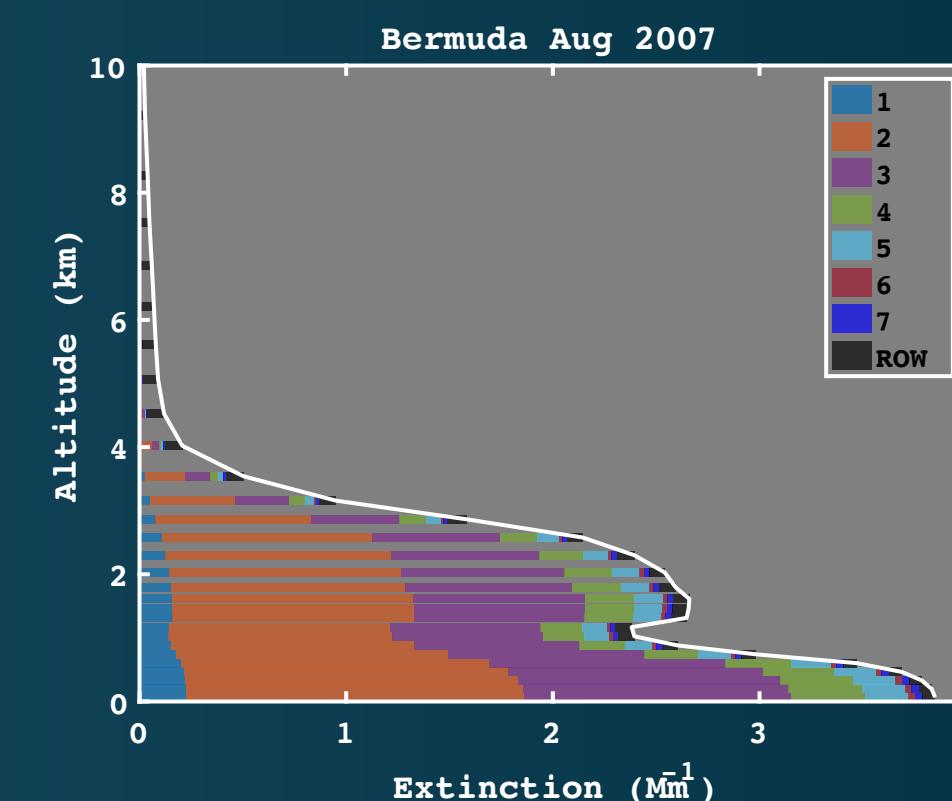
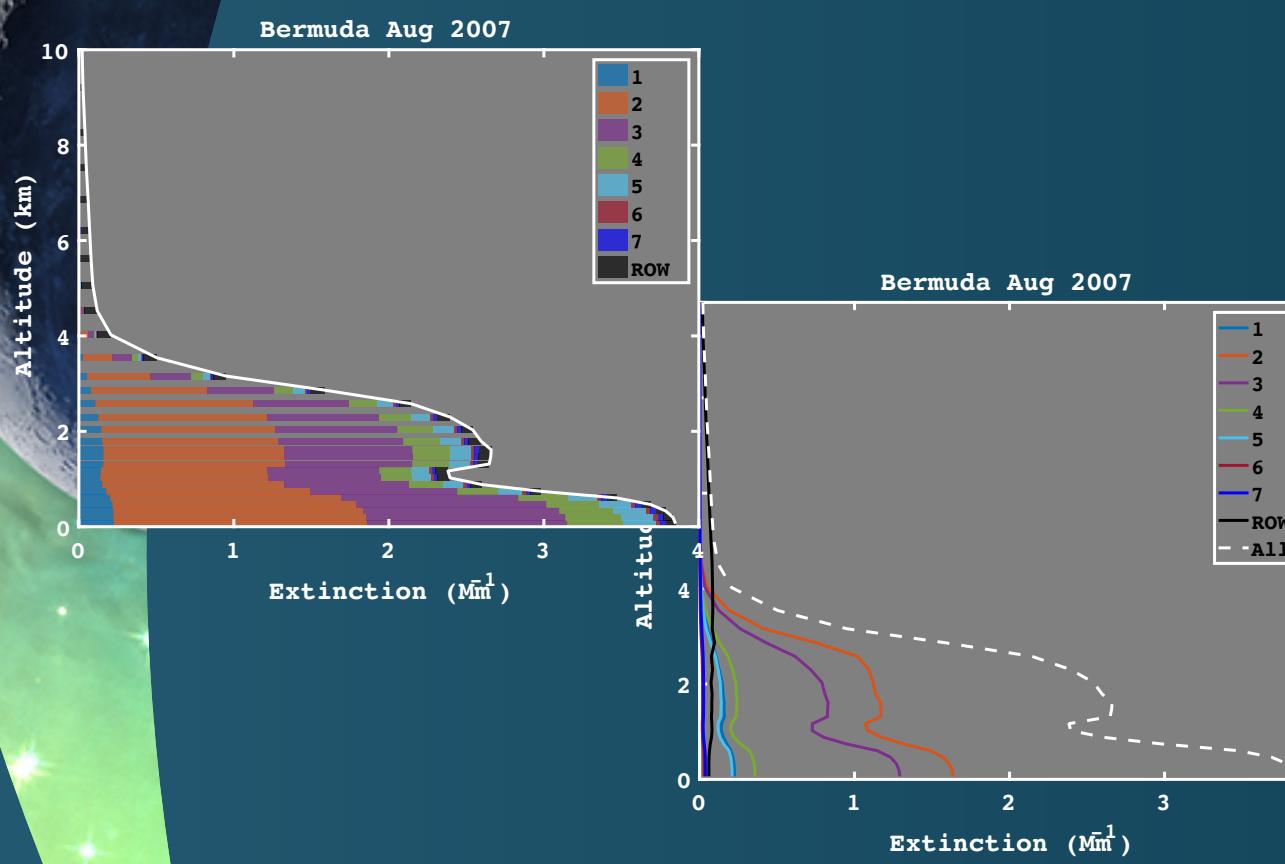
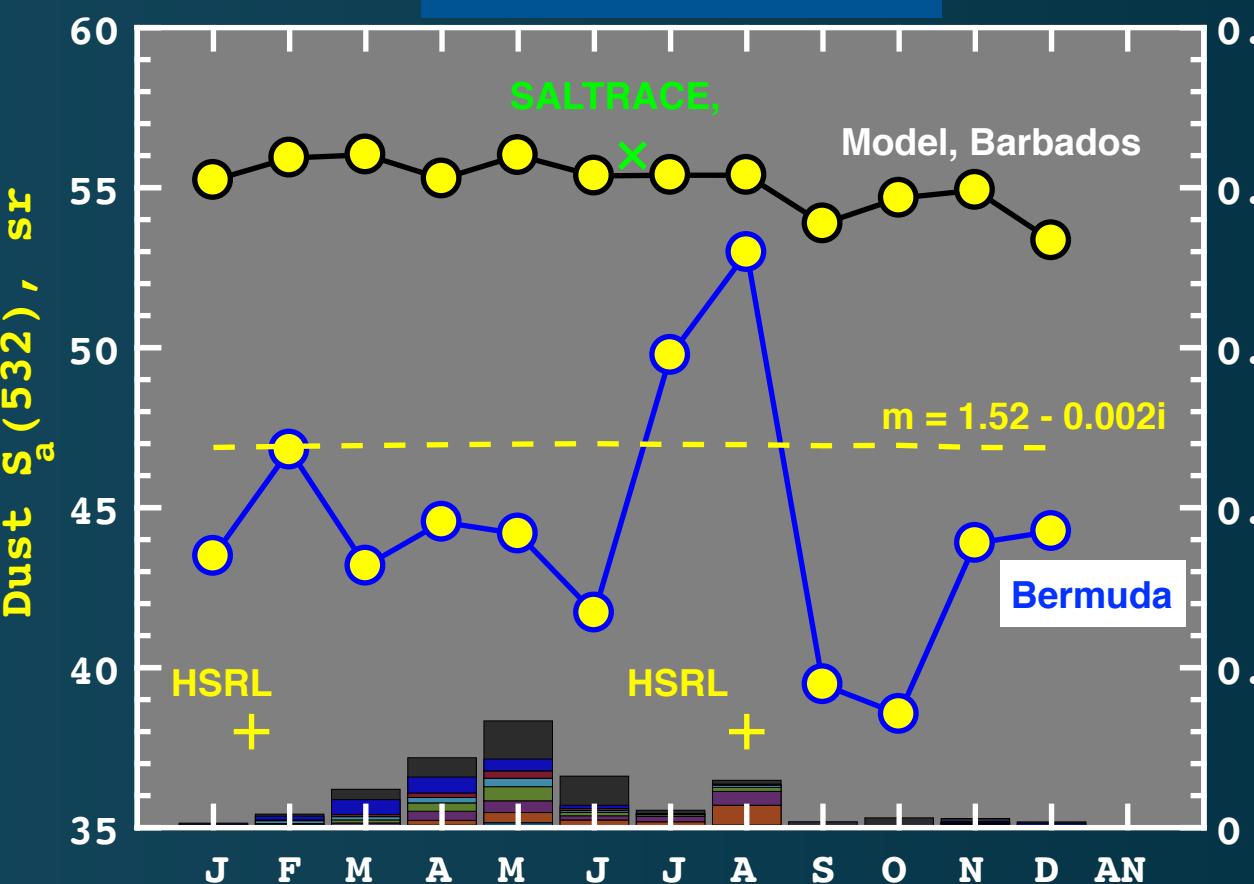
Comparison to HSRL and Raman Lidars in West Atlantic

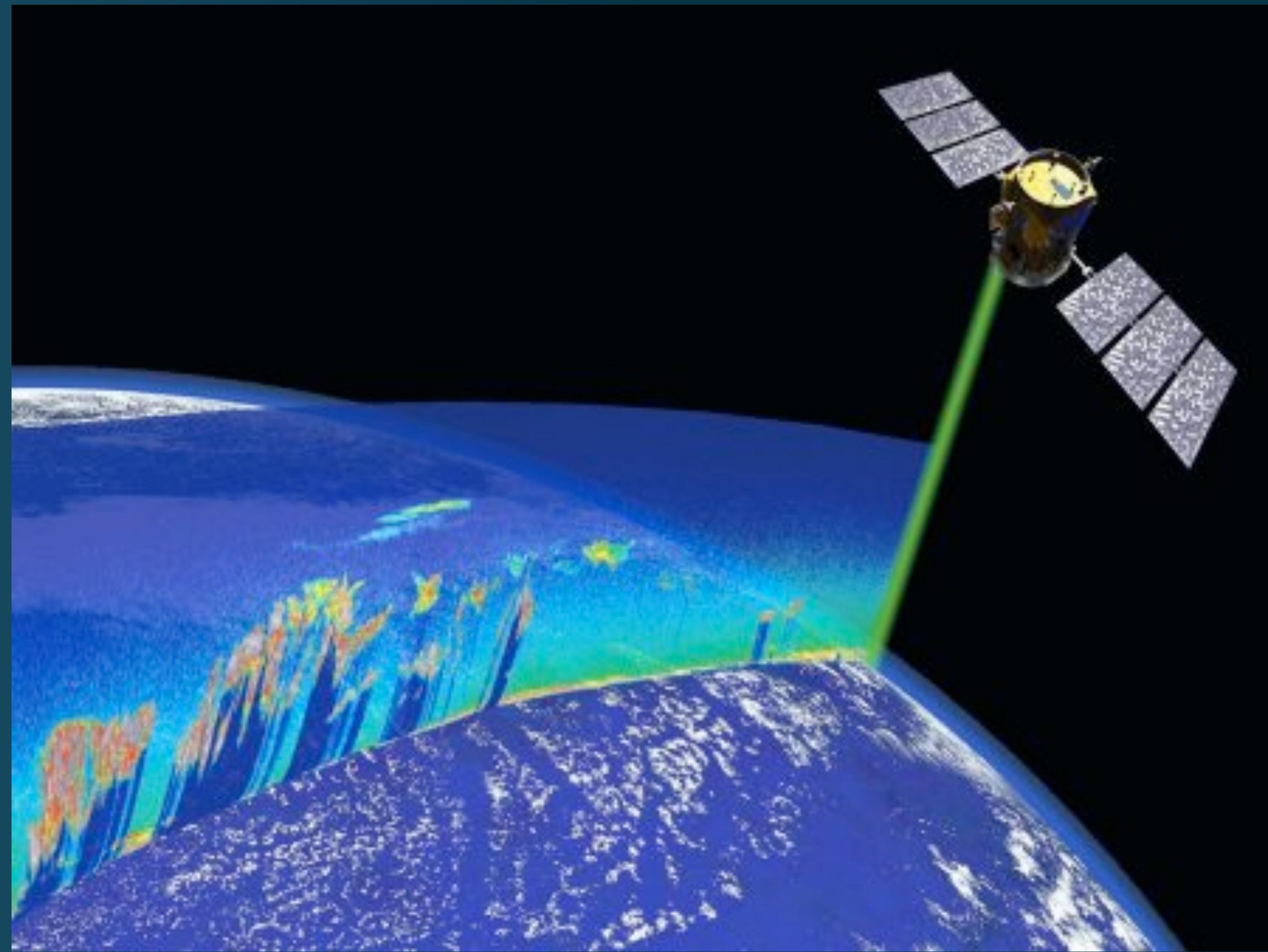
Model Year 2007

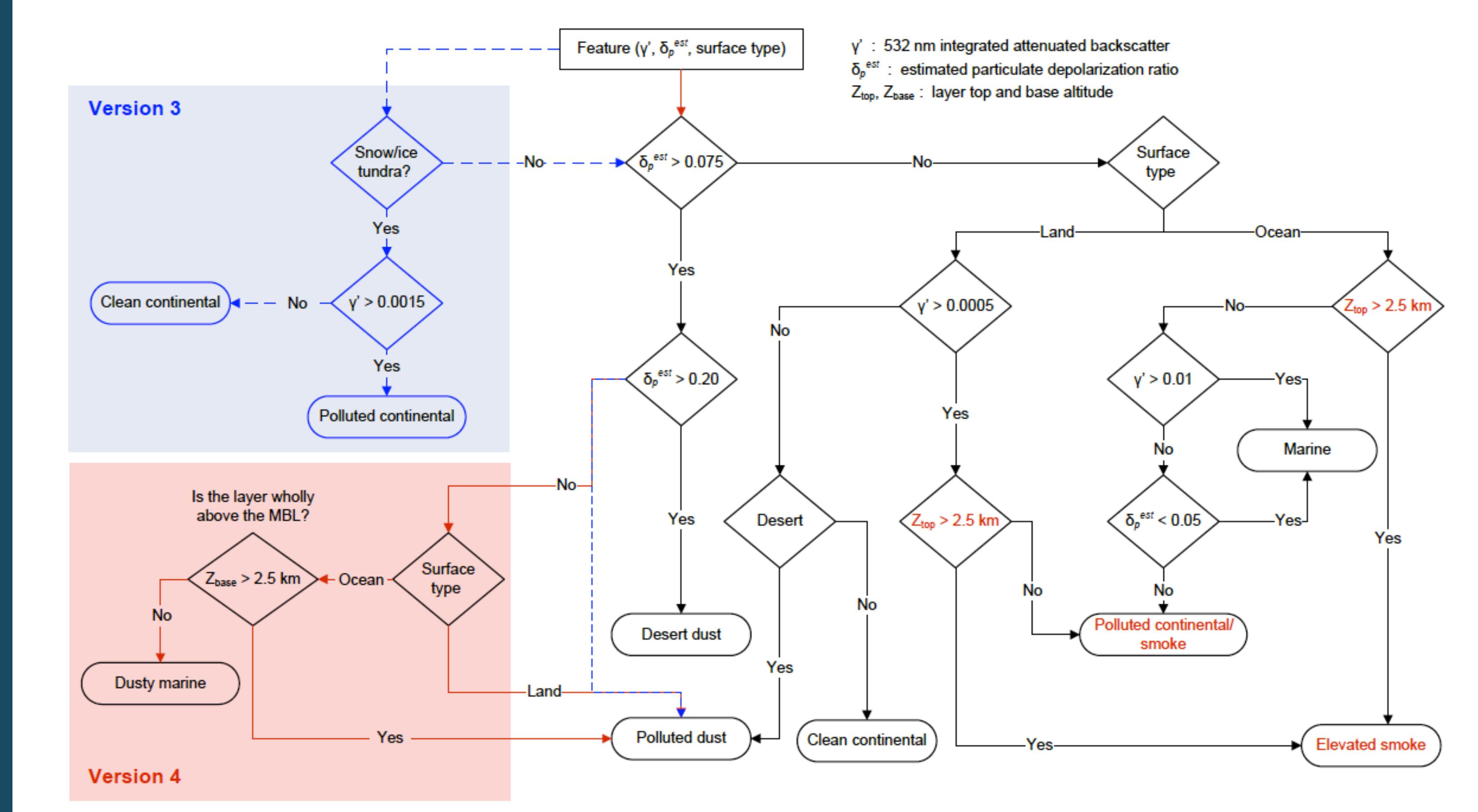


Comparison to HSRL and Raman Lidars in West Atlantic

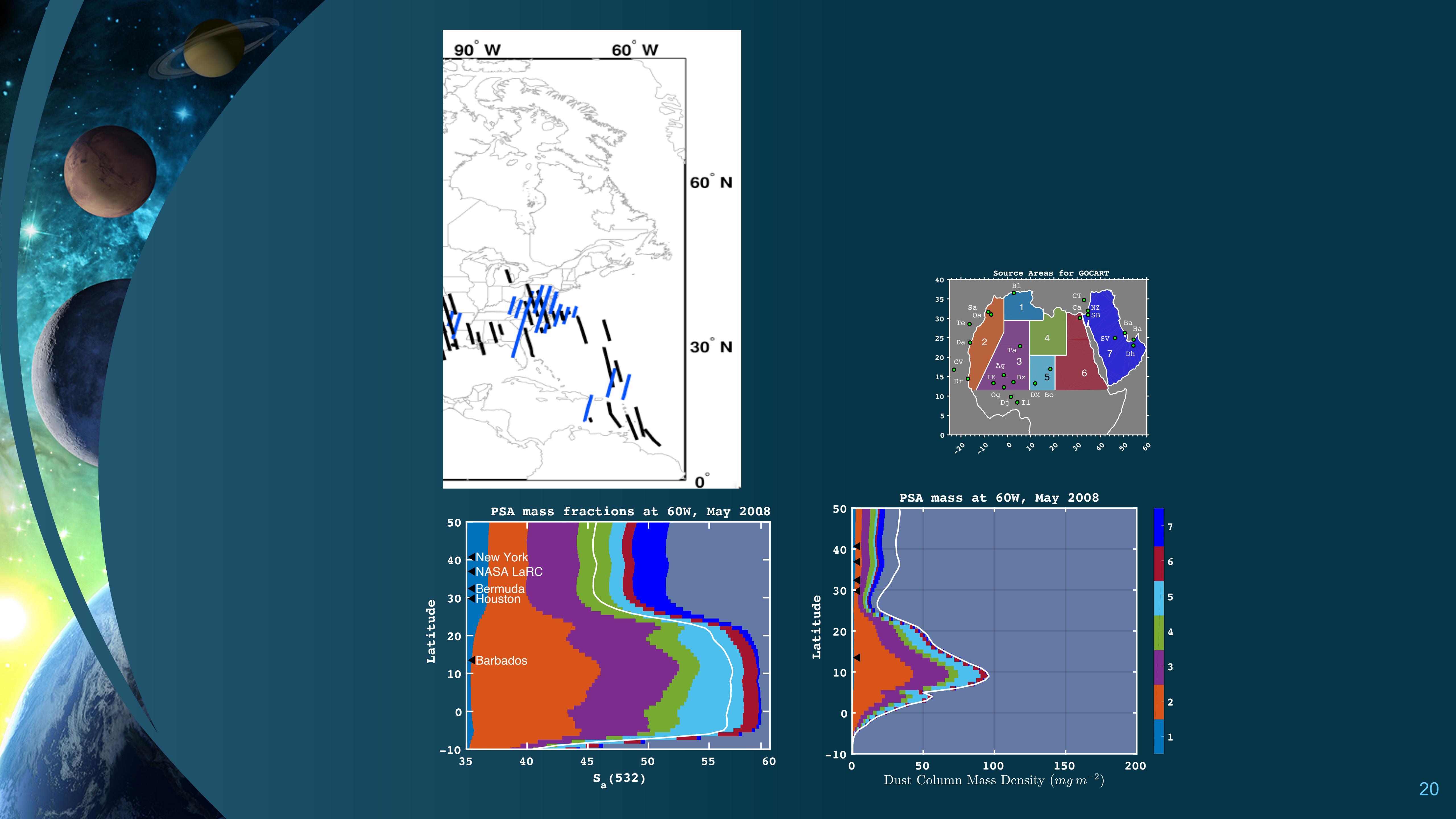
Model Year 2007



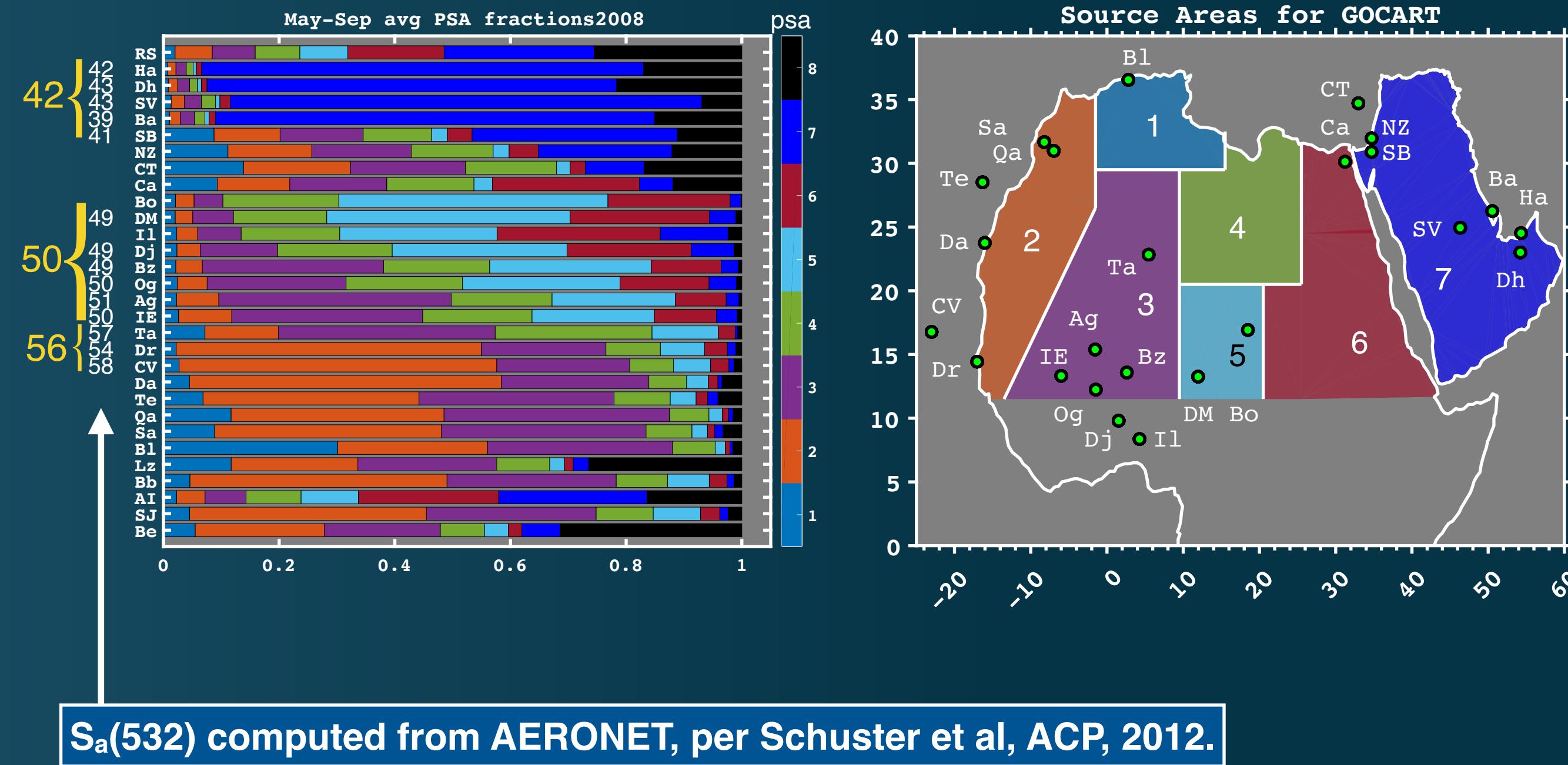




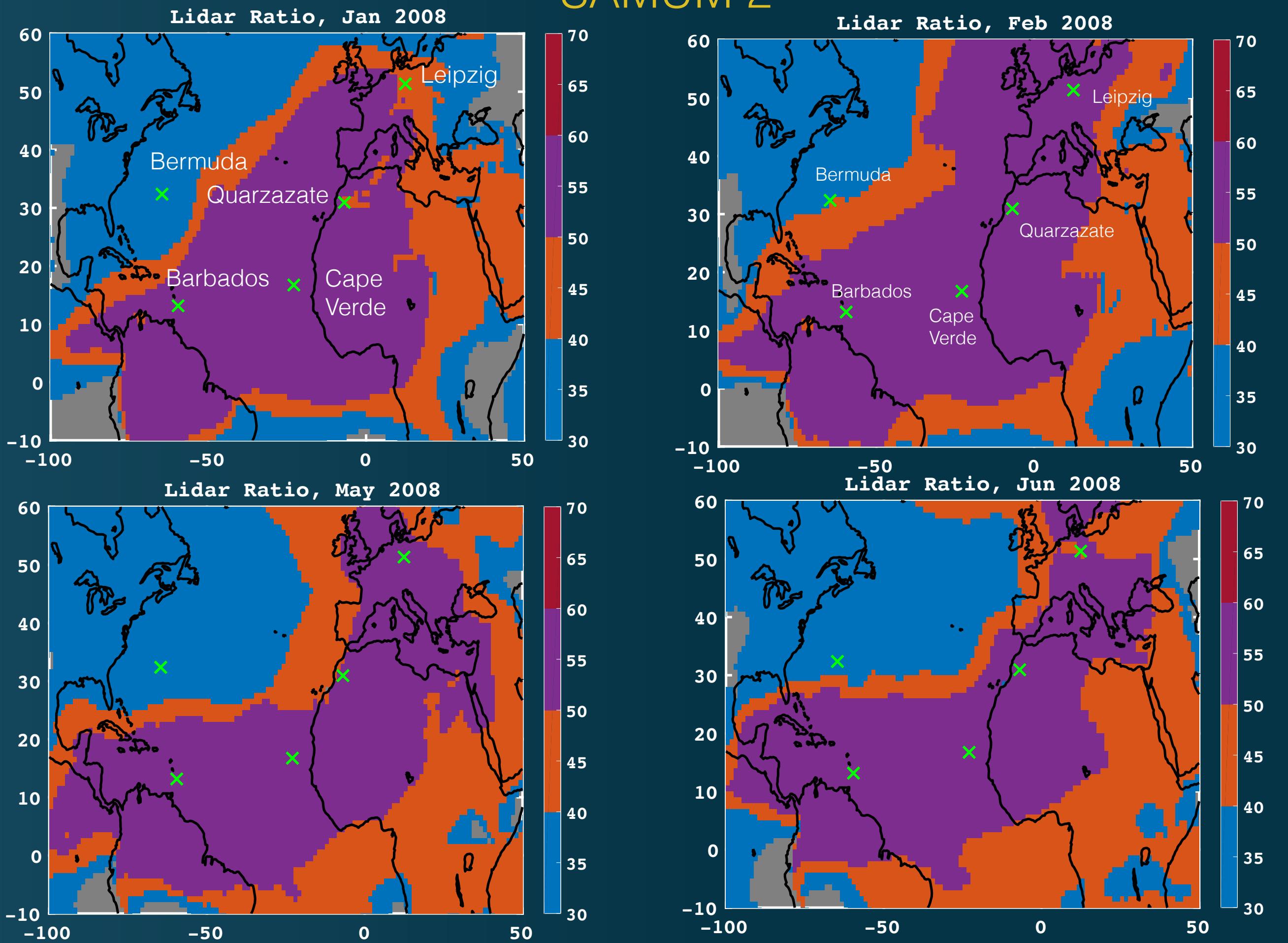
Kim et al (AMT, 2018)

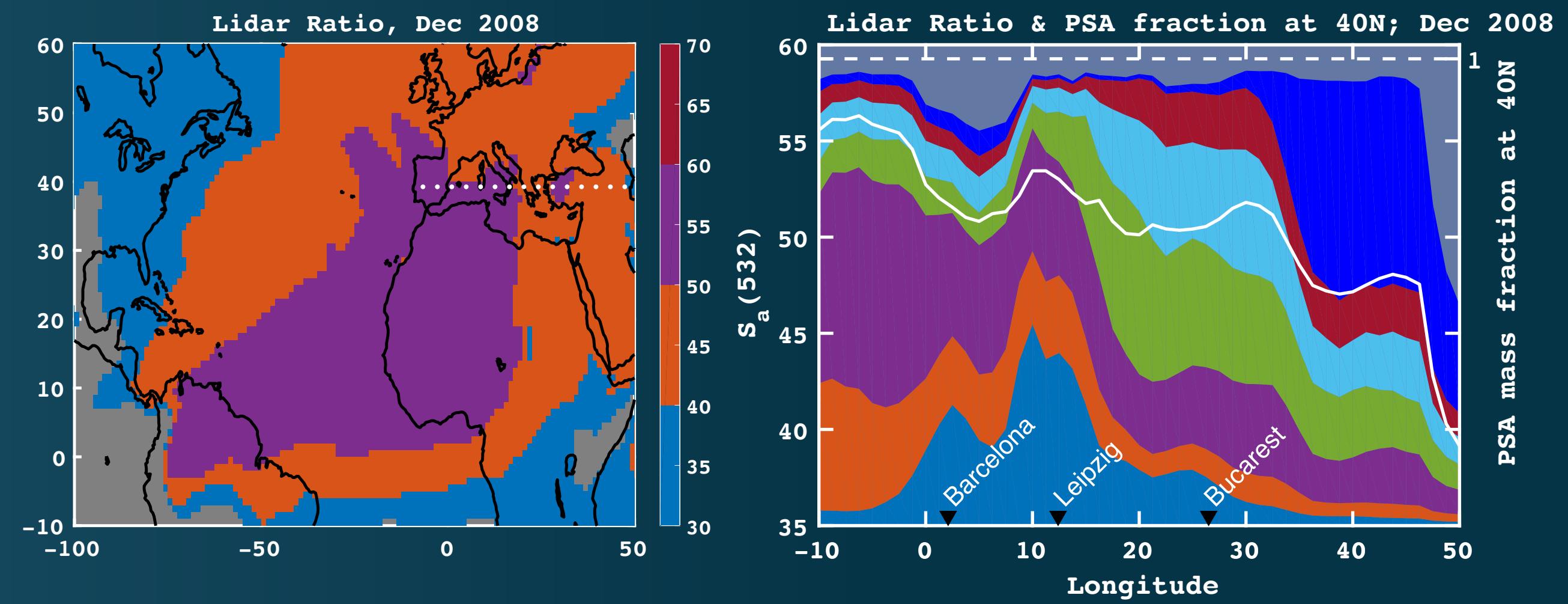


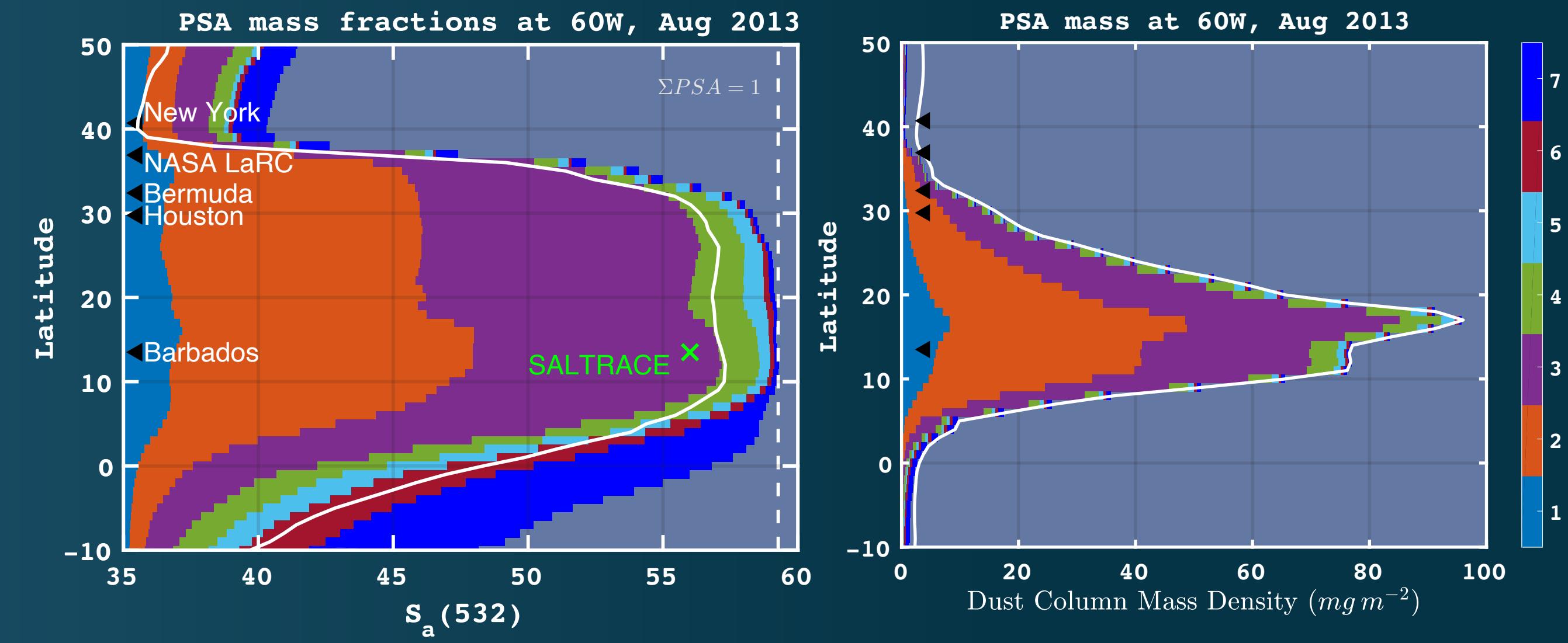
Dust originating from West Africa has highest lidar ratios



SAMUM 2



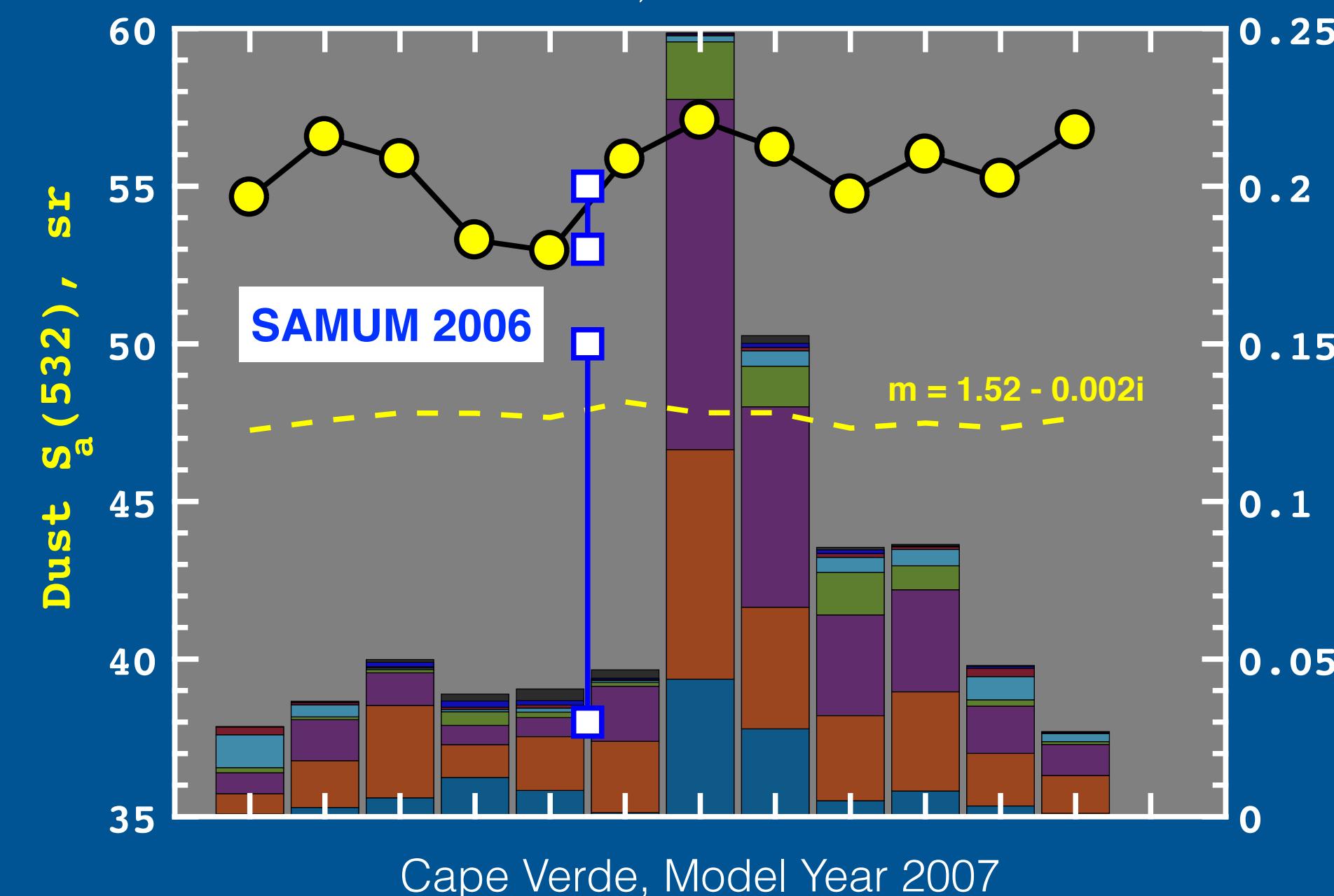




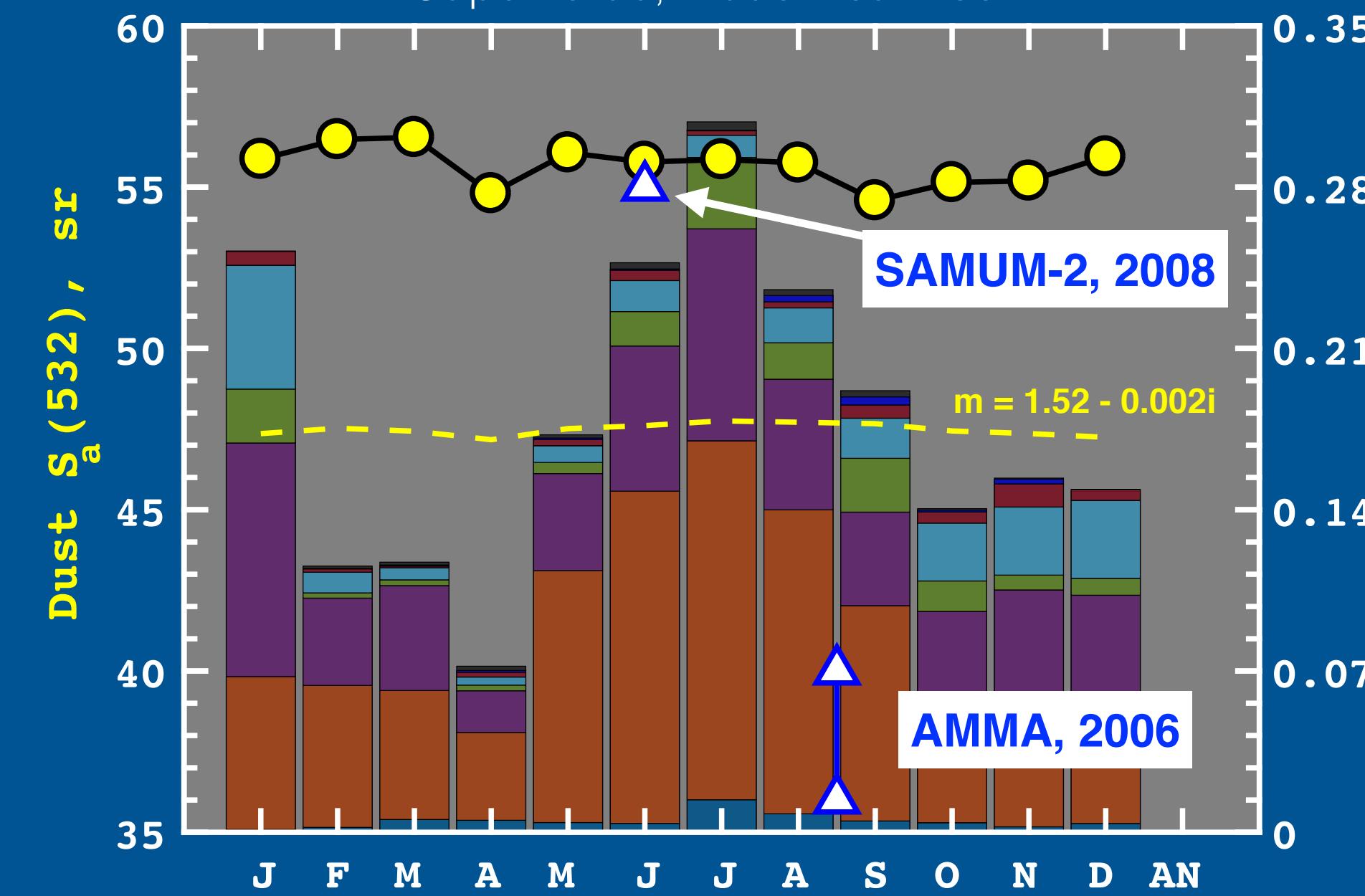
SALTRACE Lidar Ratio from Gross (ACP, 2015)

Other Field Missions with Lidar Ratio Measurements

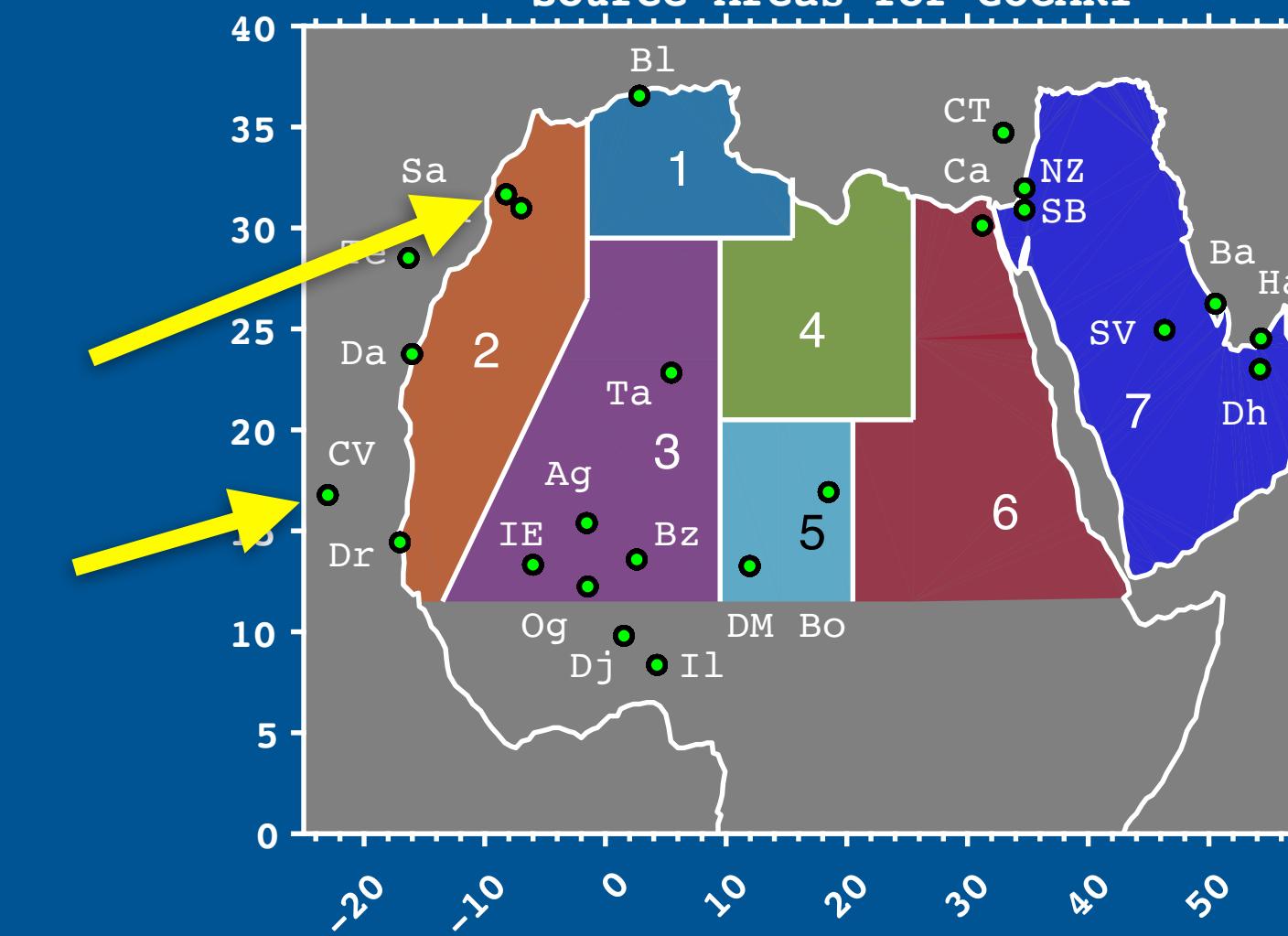
Quarzazate, Model Year 2007



Cape Verde, Model Year 2007



Source Areas for GOCART



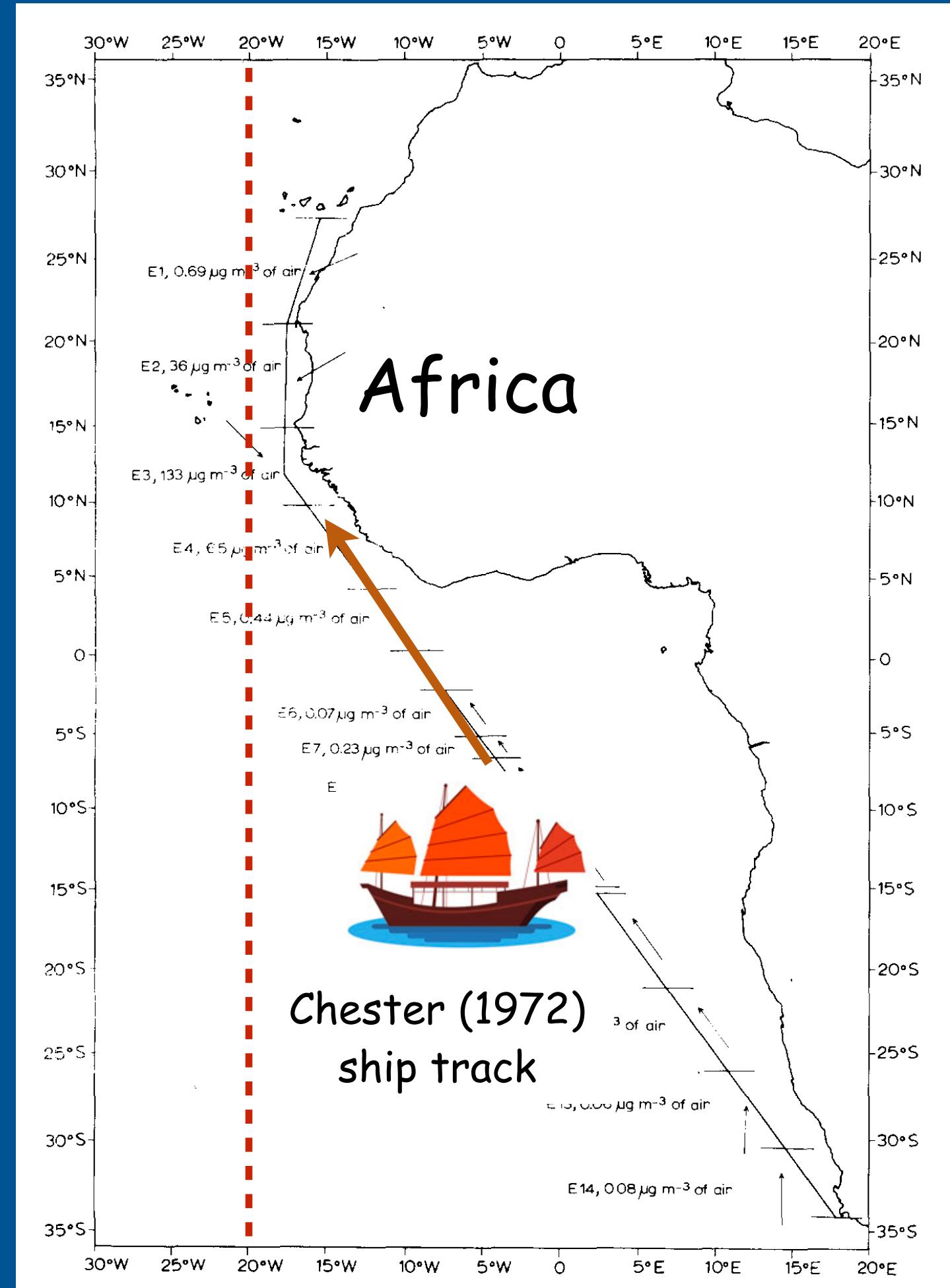
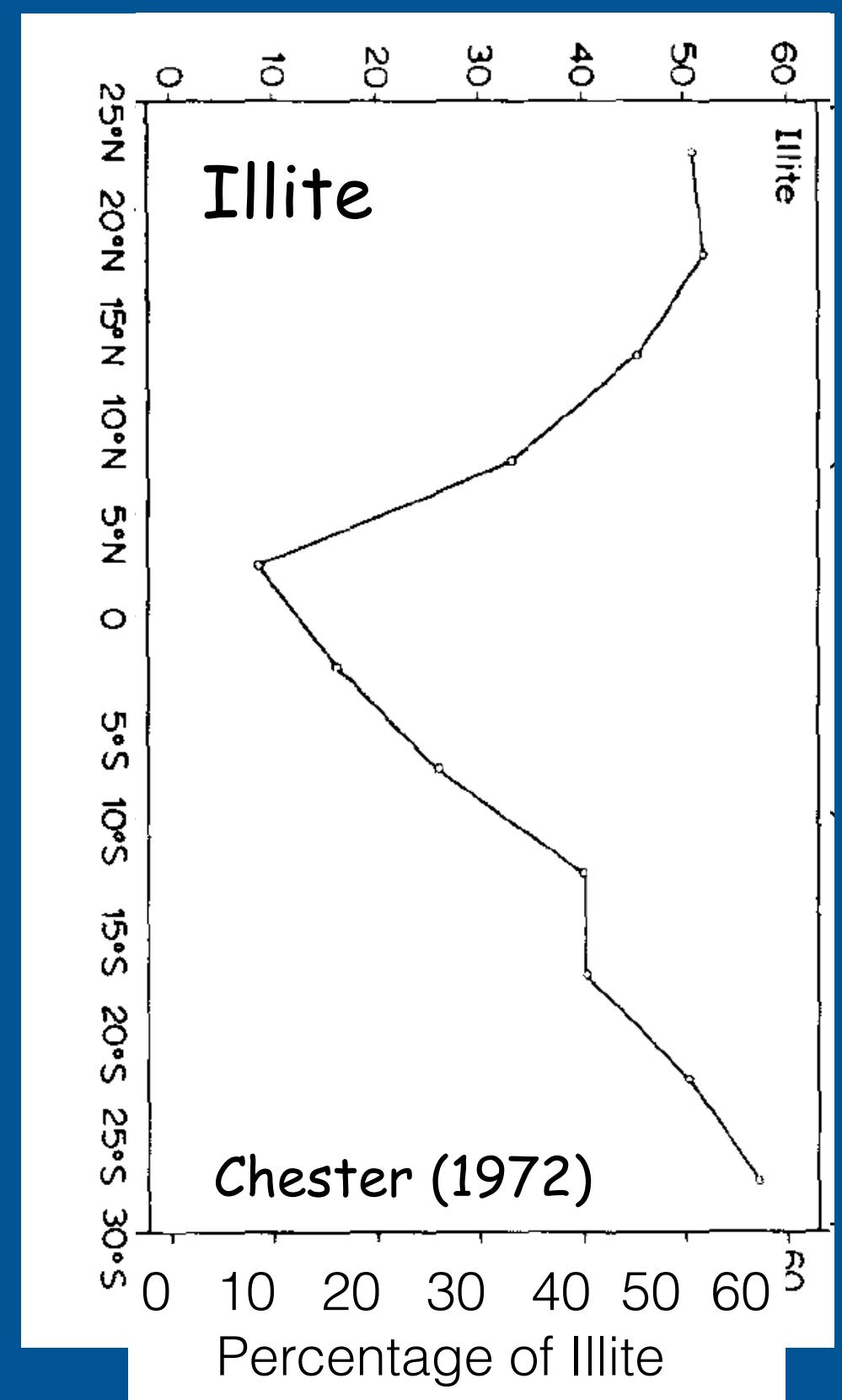
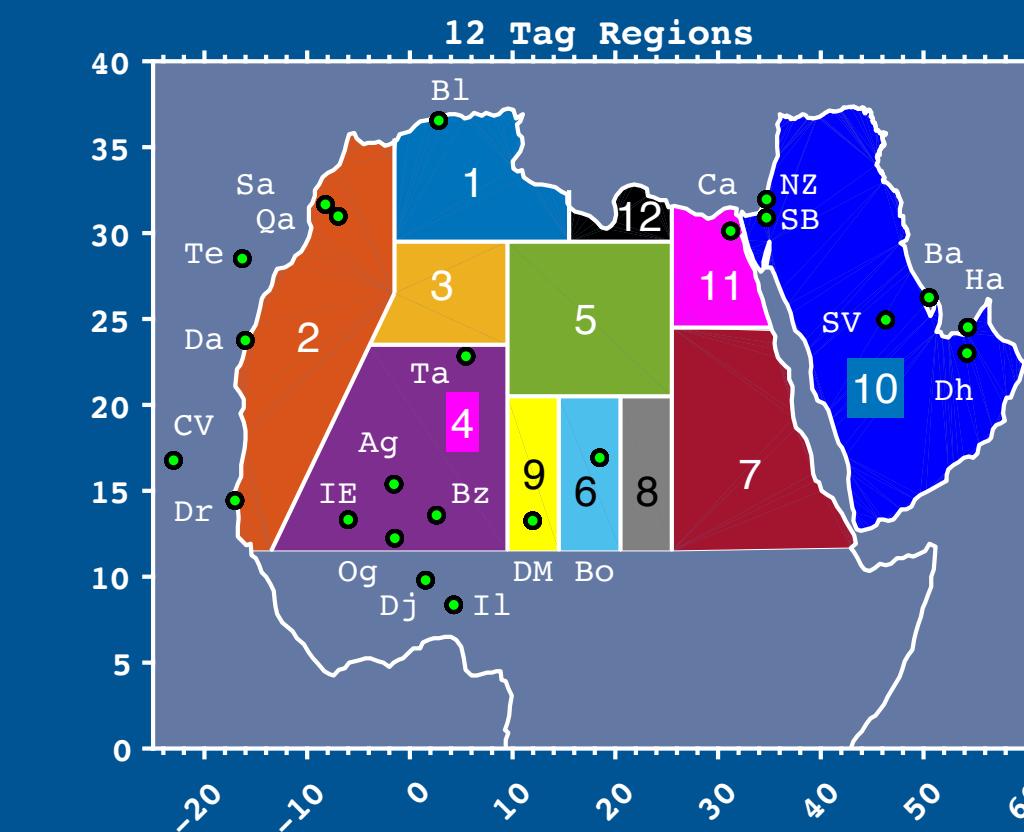
Citations:

AMMA: Omar (JGR, 2010)

SAMUM: Tesche (Tellus, 2008), Esselborn (Tellus, 2008)

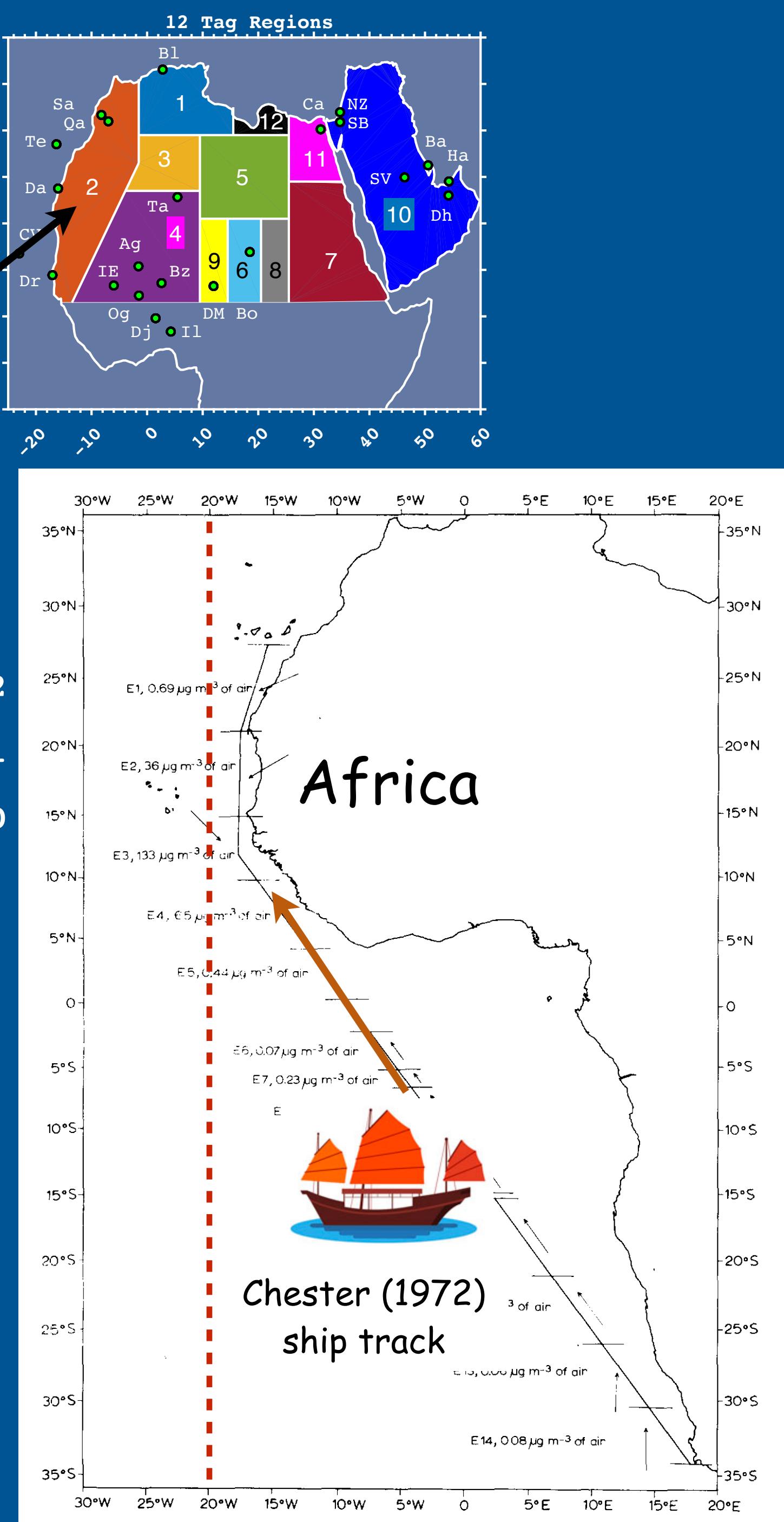
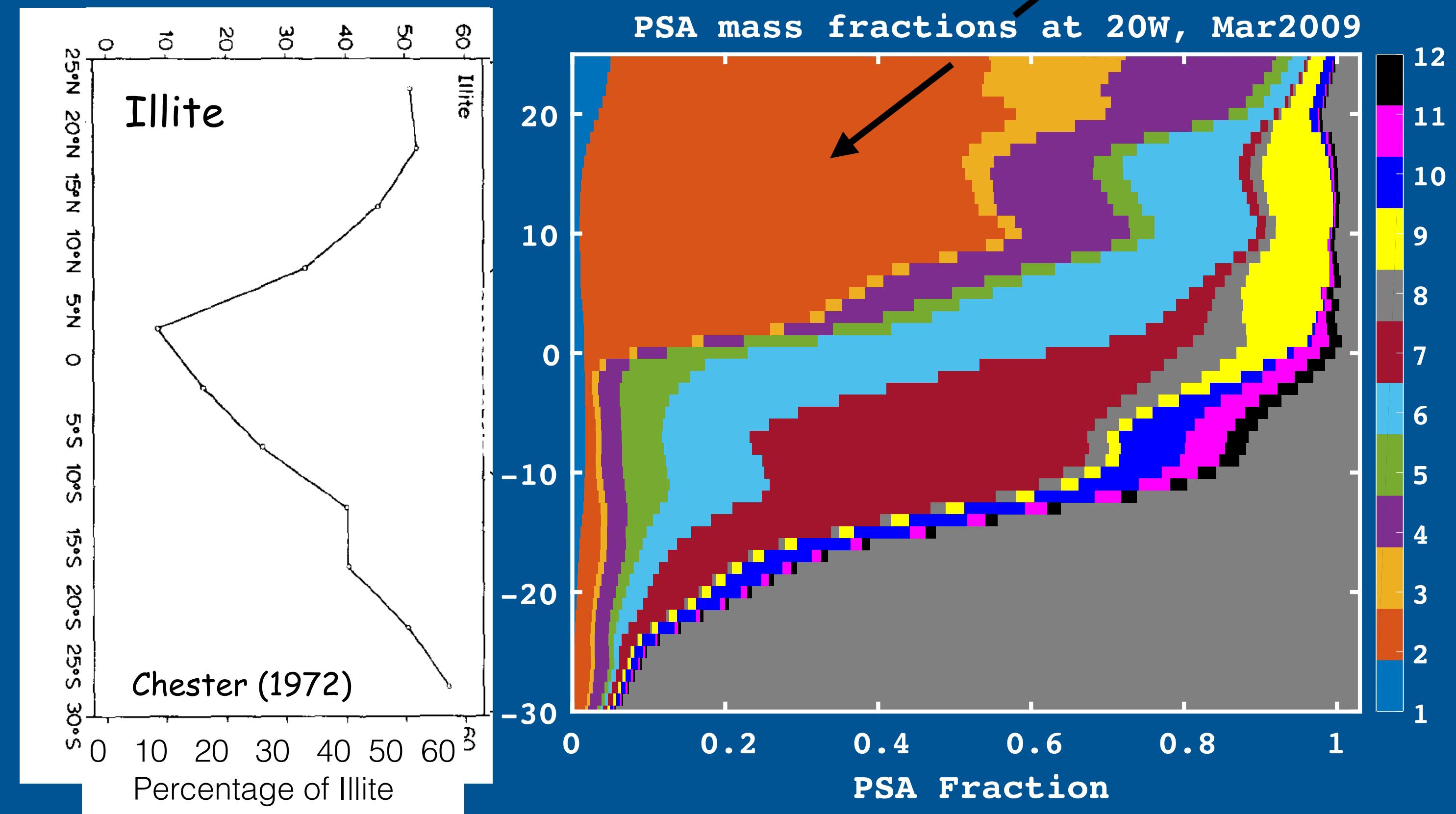
SAMUM-2: Wanginger (GRL, 2010)

Quantify the Origin of Aeolian Dust Collected over the Atlantic

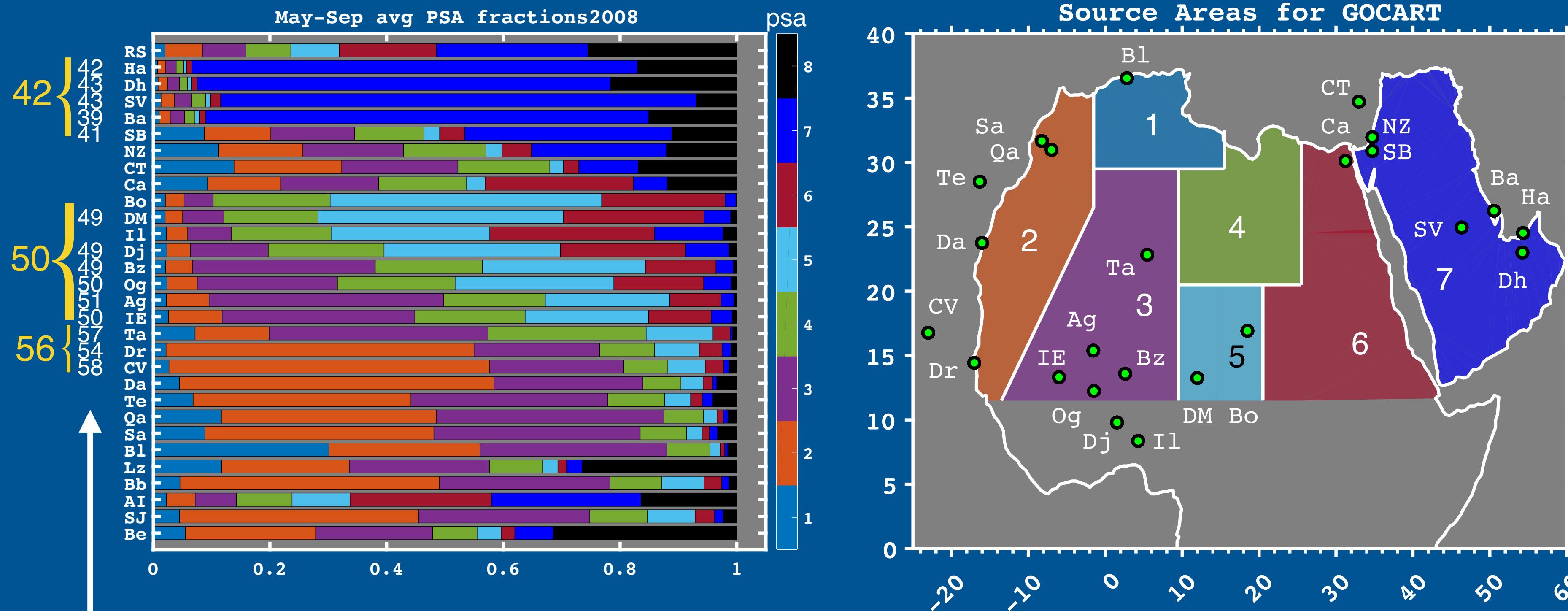


Quantify the Origin of Aeolian Dust Collected over the Atlantic

High concentrations of illite in West Africa
(Caquineau, JGR 2002)



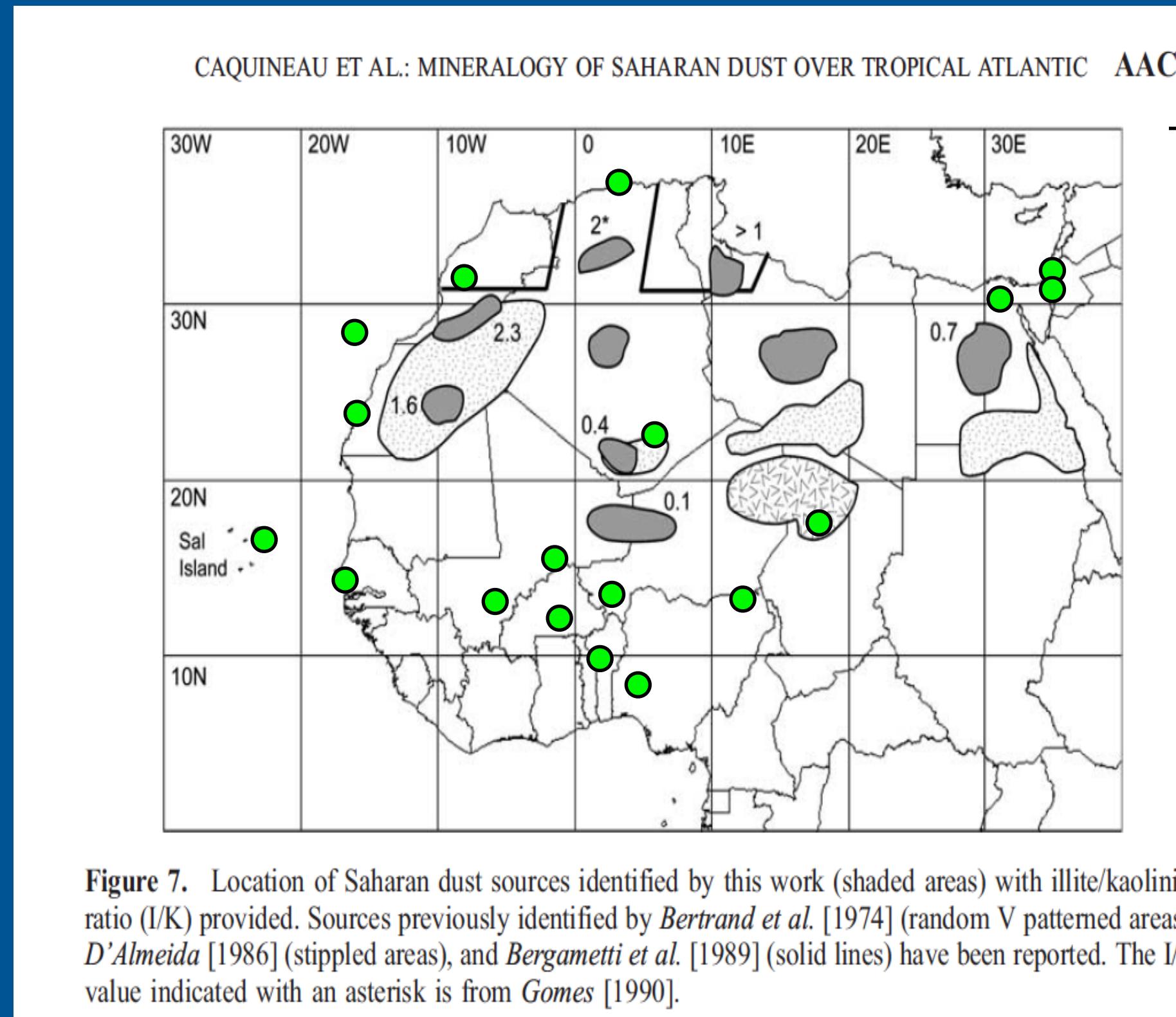
Dust originating from West Africa has highest lidar ratios



$S_a(532)$ computed from AERONET, per Schuster et al, ACP, 2012.

Highest proportion of illite is located in northwest Africa (regions 1 & 2).
 However, some significant illite exists in regions 4 and 11.

Illite/Kaolinite ratios at African PSAs



Illite/Kaolinite ratios at African PSAs

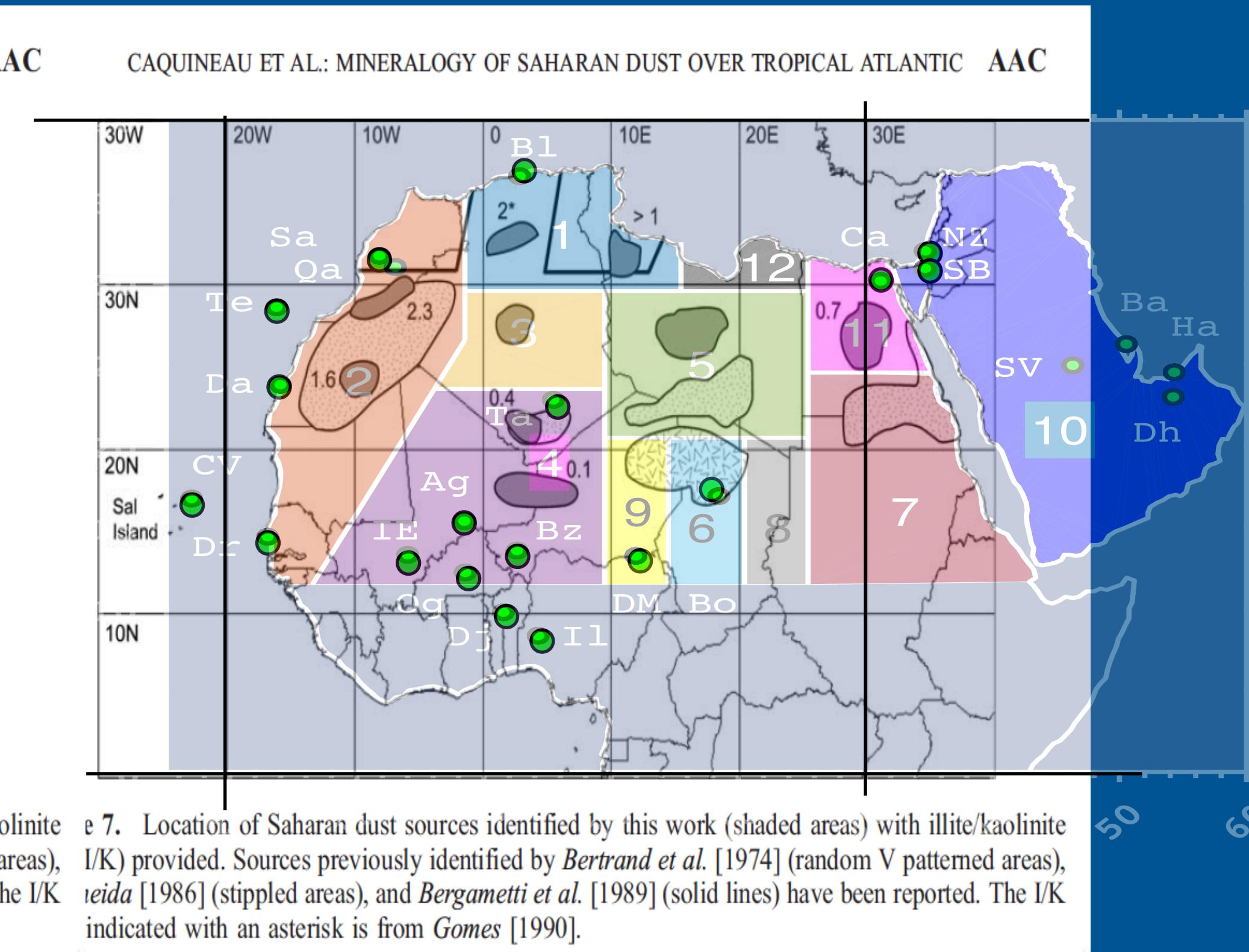


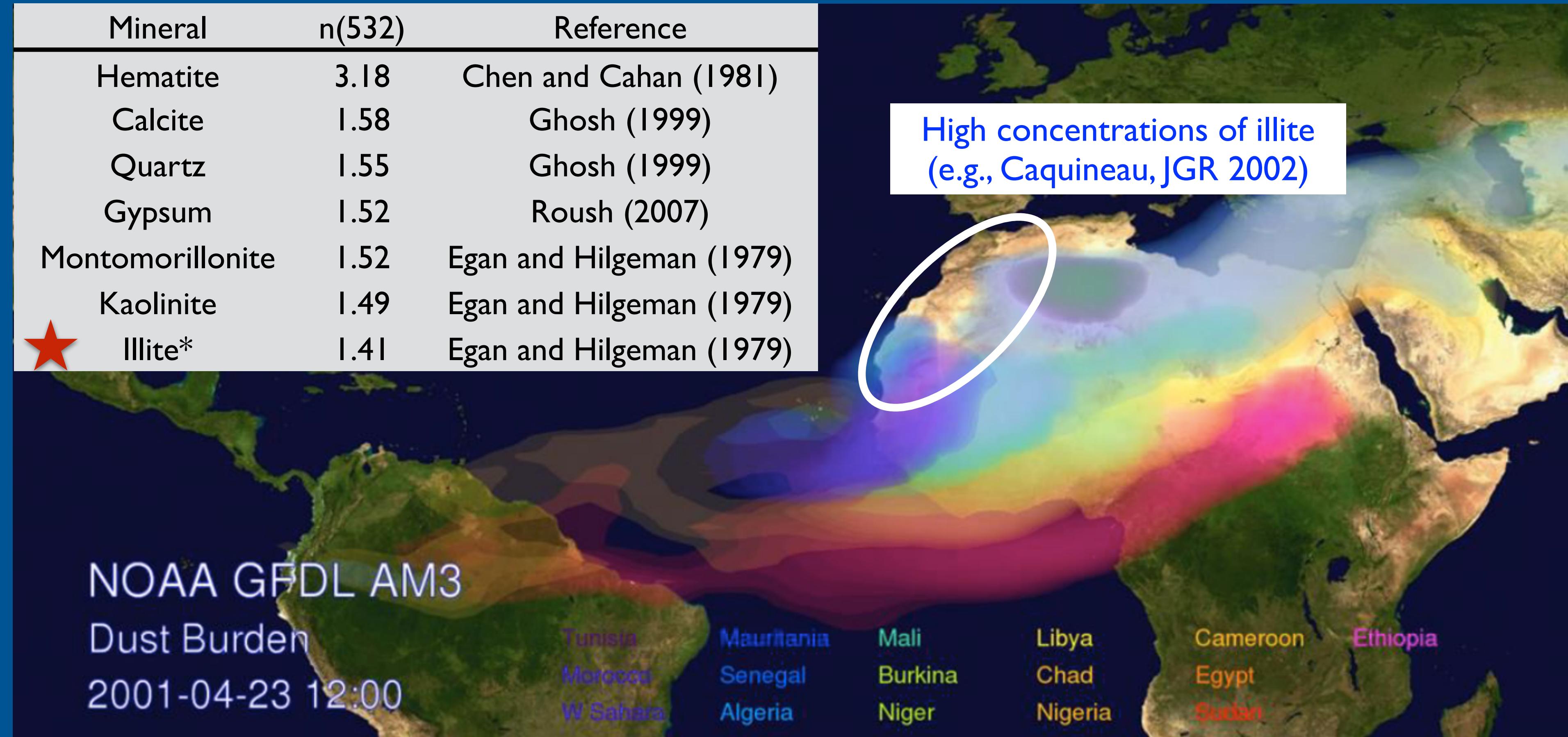
Figure 7. Location of Saharan dust sources identified by this work (shaded areas) with illite/kaolinite ratio (I/K) provided. Sources previously identified by Bertrand et al. [1974] (random V patterned areas), D'Almeida [1986] (stippled areas), and Bergametti et al. [1989] (solid lines) have been reported. The I/K value indicated with an asterisk is from Gomes [1990].

Circles are AERONET sites.

Linking Optical Properties of Dust to Source Regions Over Africa and the Middle East

Greg Schuster, NASA LaRC; Dongchul Kim, NASA GSFC, JCET, UMBC, USRA; Zhaoyan Liu, NASA LaRC; Mian Chin, NASA GSFC; Kerstin Schepanski, TROPOS.

Mineral	n(532)	Reference
Hematite	3.18	Chen and Cahan (1981)
Calcite	1.58	Ghosh (1999)
Quartz	1.55	Ghosh (1999)
Gypsum	1.52	Roush (2007)
Montmorillonite	1.52	Egan and Hilgeman (1979)
Kaolinite	1.49	Egan and Hilgeman (1979)
★ Illite*	1.41	Egan and Hilgeman (1979)



African dust sources presented by Joe Prospero at DUST 2014. Model run by Paul Ginoux

Relative Bias of CALIPSO Lidar Ratio wrt AERONET at Dust Sites

$$Bias = \frac{S_a(CALIPSO)}{S_a(AERONET)} - 1$$

	AERONET		CALIPSO			
	Lidar Ratio Climatology		Lidar Ratio Bias ($S_a = 40$)			
	pure dust	all dust	pure dust	all dust		
Africa	51	56	-0.21	-0.29		
Middle East	43	47	-0.06	-0.15		

depolarization fine volume fraction

pure dust

> 0.2

< 0.05

all dust

> 0.2

no restriction

Relative Bias of CALIPSO Lidar Ratio wrt AERONET at Dust Sites

$$Bias = \frac{S_a(CALIPSO)}{S_a(AERONET)} - 1$$

	AERONET		CALIPSO		Aerosol Optical Depth Bias	
	Lidar Ratio Climatology		Lidar Ratio Bias ($S_a = 40$)			
	pure dust	all dust	pure dust	all dust		
Africa	51	56	-0.21	-0.29	-0.2	
Middle East	43	47	-0.06	-0.15	-0.12	

depolarization fine volume fraction

pure dust

> 0.2

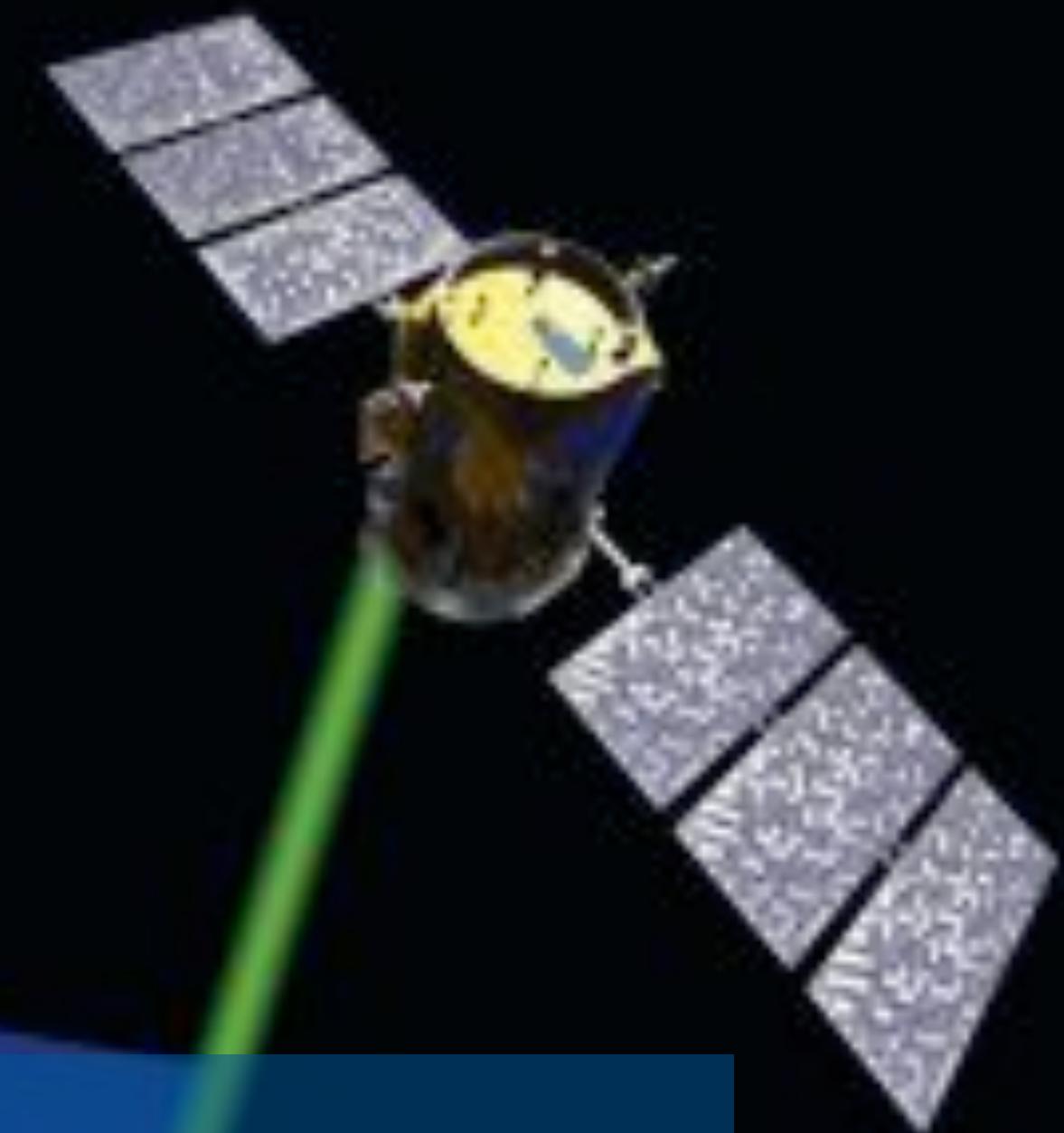
< 0.05

all dust

> 0.2

no restriction

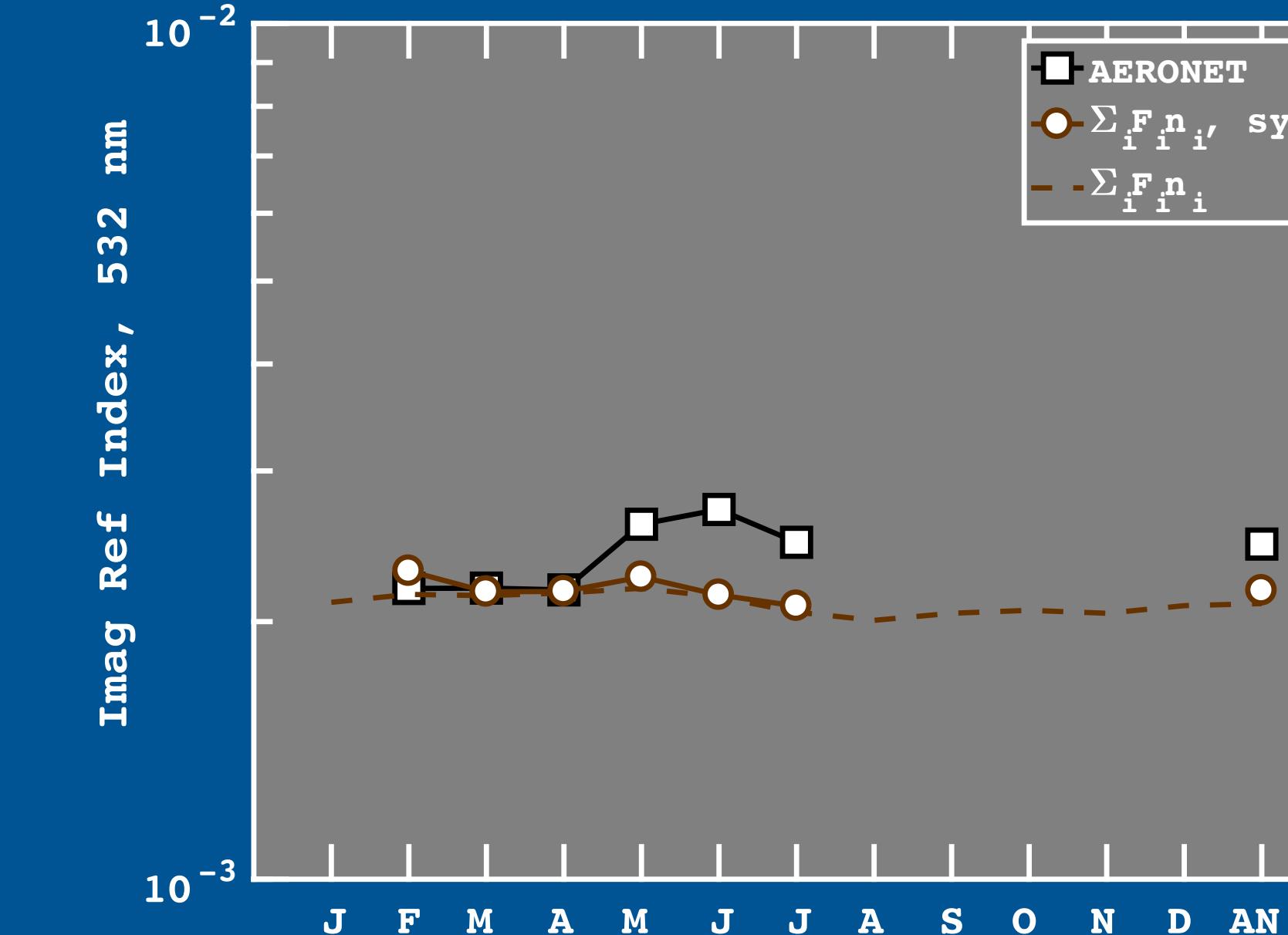
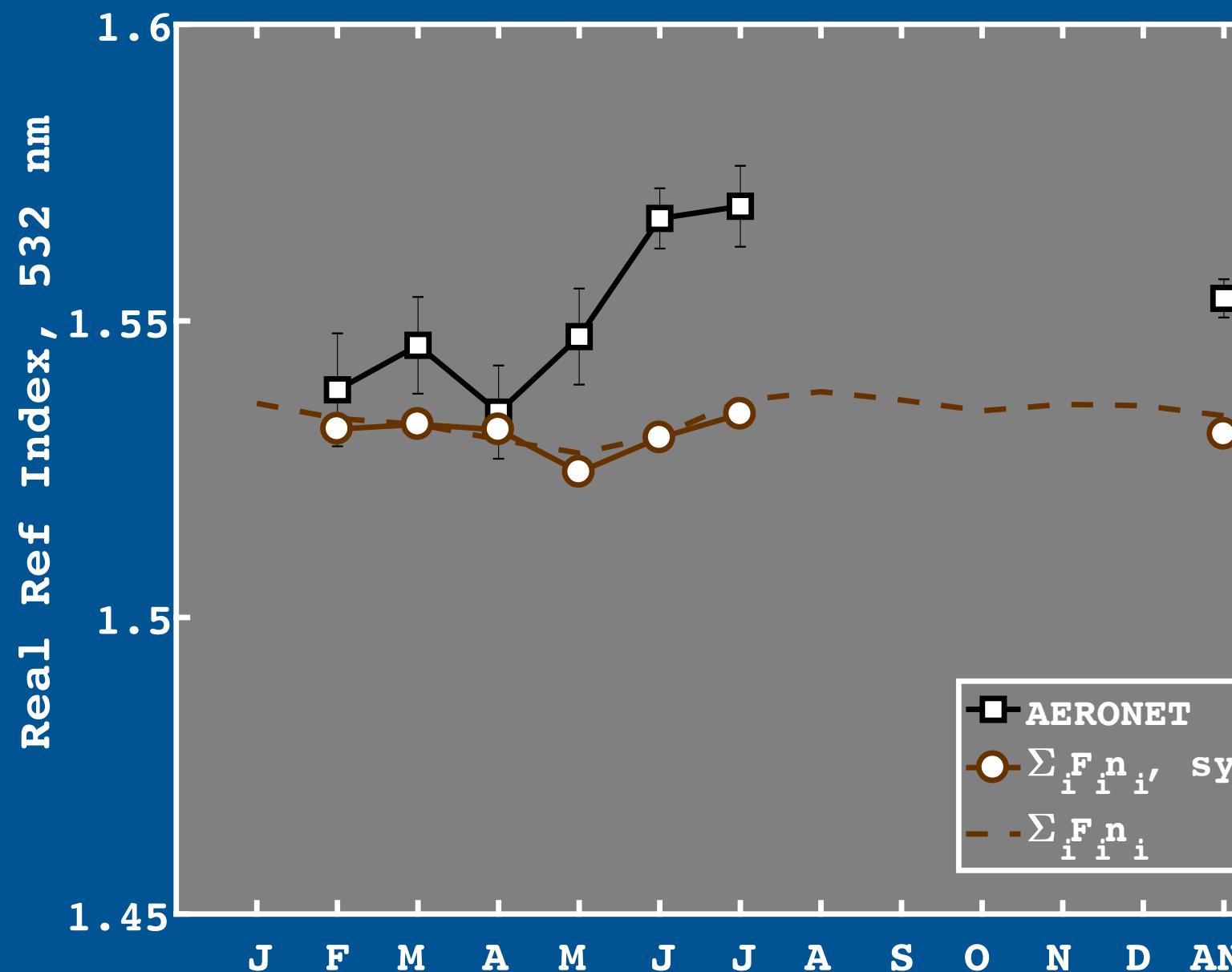
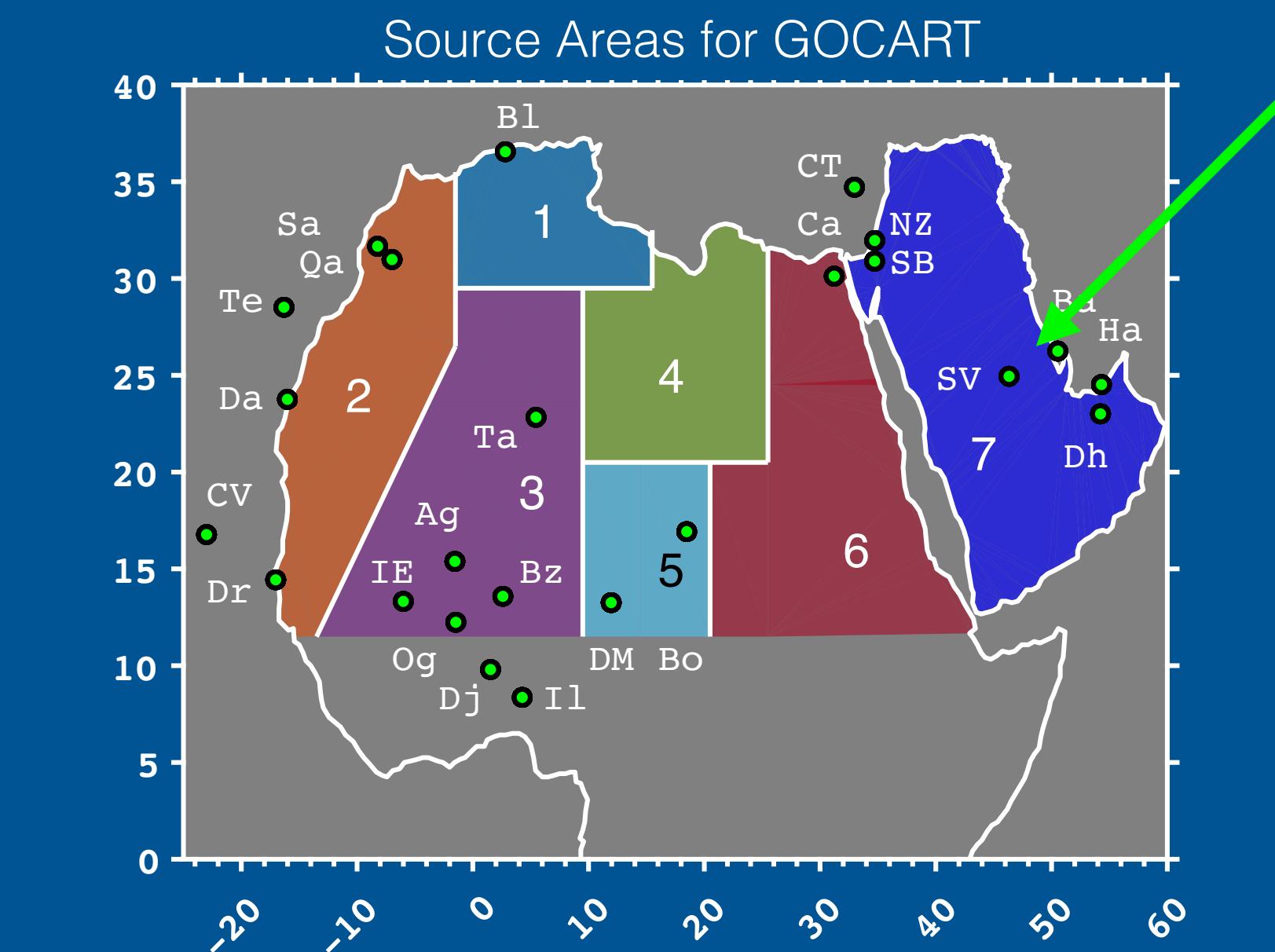
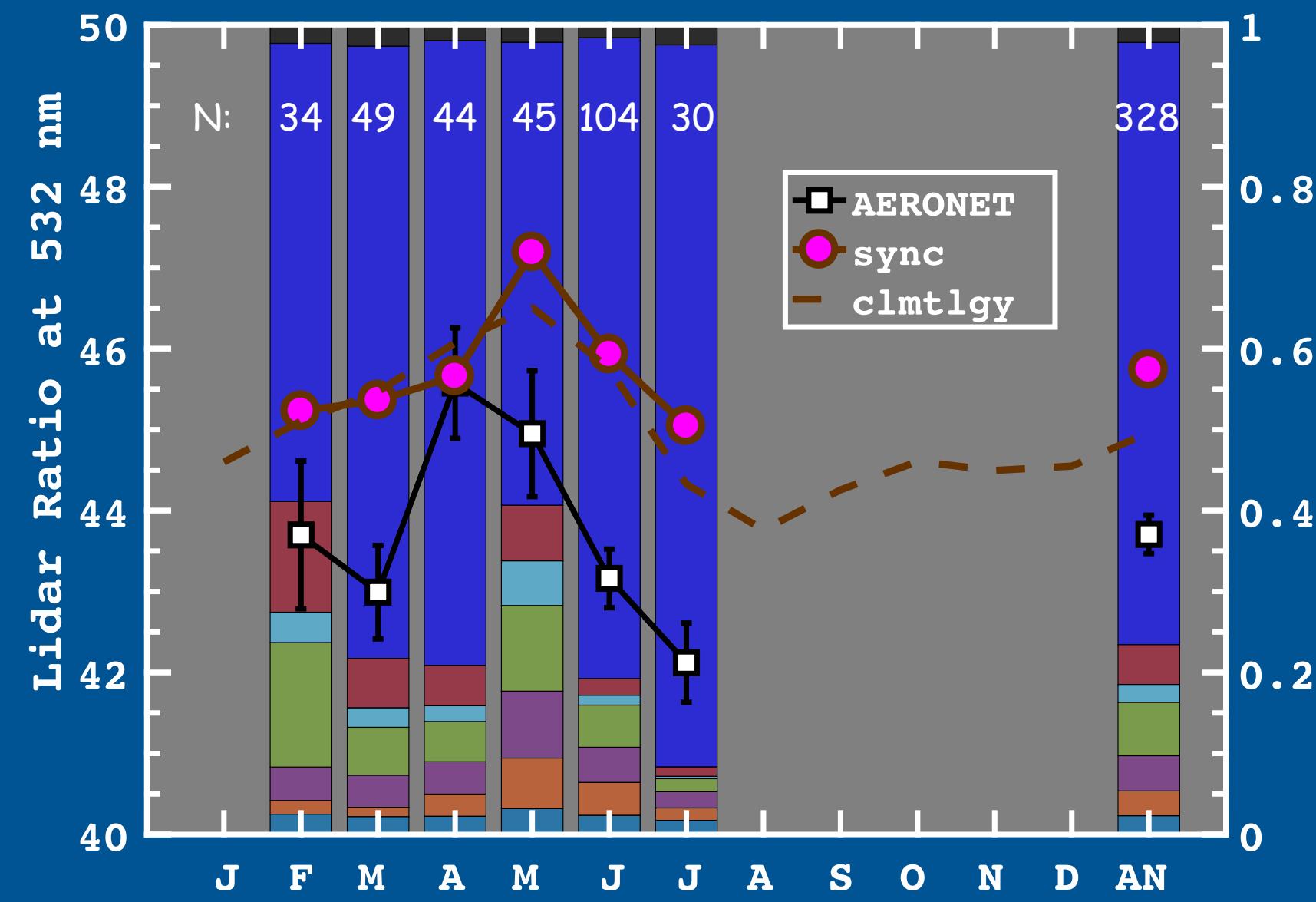
Measurements Needed to Constrain Lidar Ratios for CALIPSO Retrievals



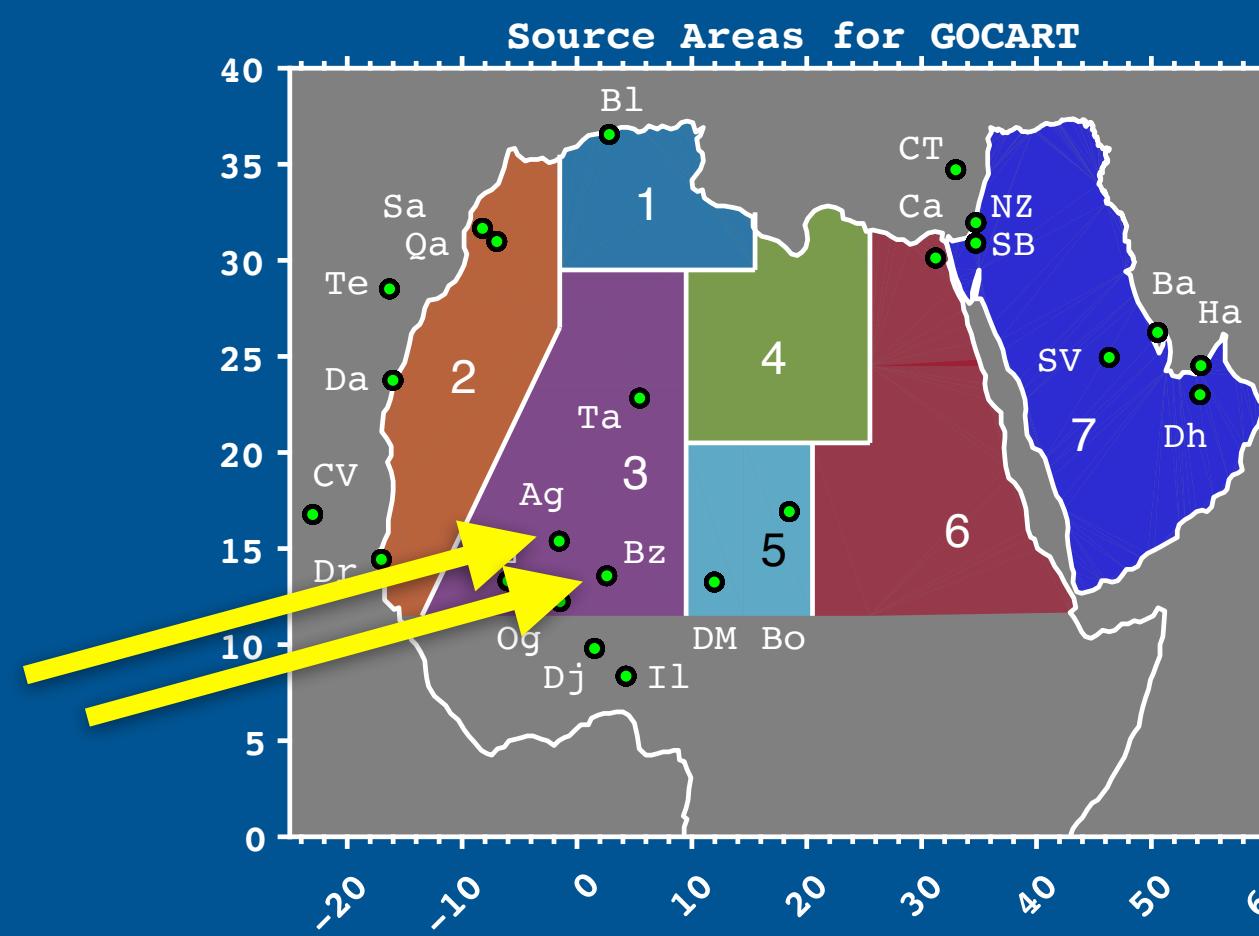
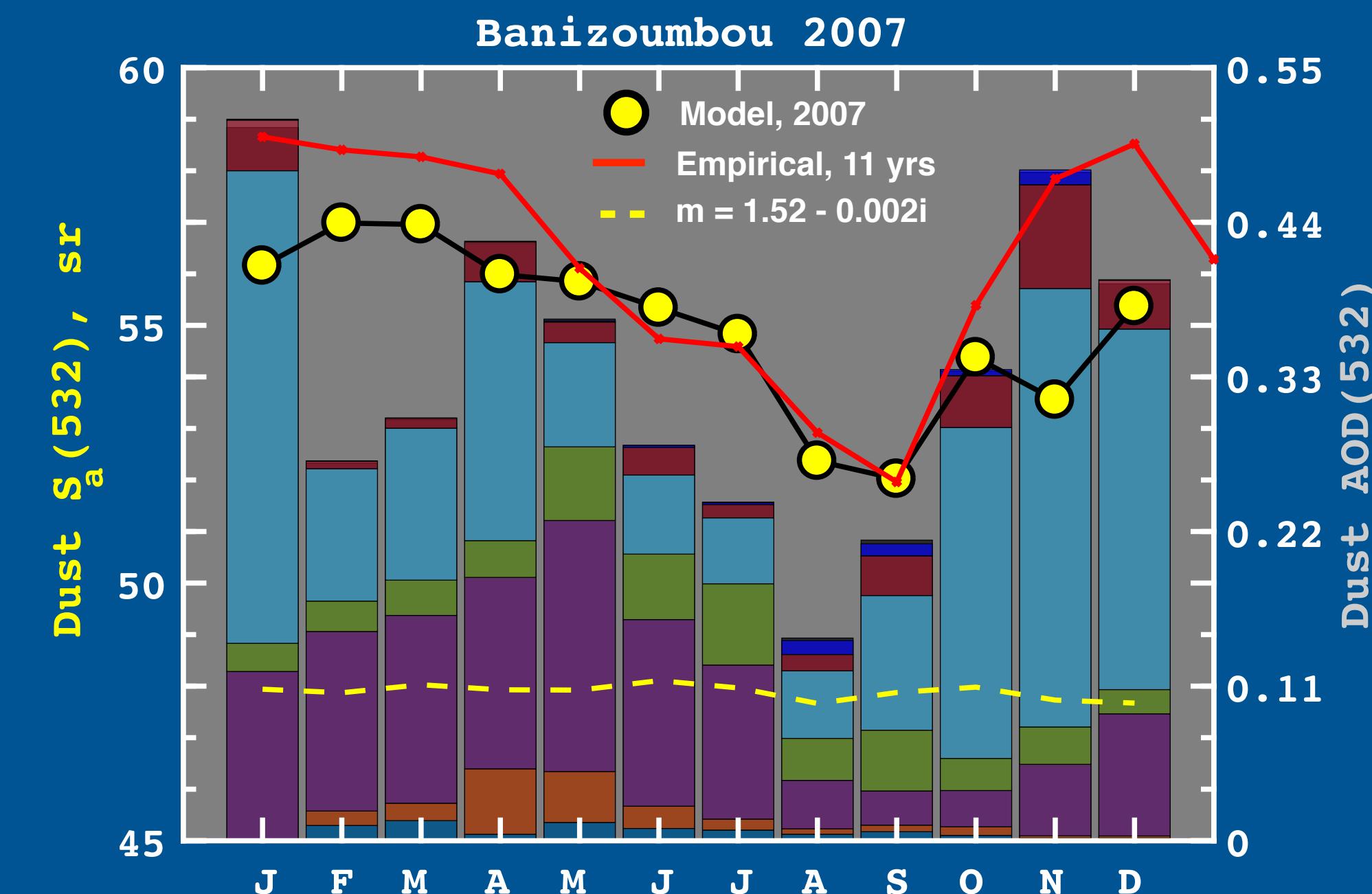
- Lidar ratio measurements are great, but we need concurrent mineralogy (or refractive index) and size distributions.
- Need these concurrent measurements near source regions in Middle East and Asia, and in transport regions over the Atlantic and Caribbean.
- Refractive index measurements for “pure” African clays (illite, kaolinite, montmorillonite, and goethite).
- Studies to determine any changes in dust refractive indices during transport.

Empirical Lidar Ratios and Refr Indices at Solar Village (1/2007 to 12/2016)

Empirical $S_a = \sum_{i=1}^8 F_i (S_{psa})_i$



Seasonal Variability at Sede Boker



Seasonal Variability at Sede Boker

