The Global Satellite Precipitation Constellation: current status and future requirements

Chris Kidd[1,2]
George Huffman[2]
Viviana Maggioni[3]
Philippe Chambon[4]
Riko Oki[5]

[1] *Earth System Science Interdisciplinary Center, University of Maryland, College Park, MD. USA*
[2] *NASA/Goddard Space Flight Center, Greenbelt, MD. USA*
[3] *Sid and Reva Dewberry Department of Civil, Environmental, and Infrastructure Engineering, George Mason University, Fairfax, VA, USA*
[4] *CNRM, Université de Toulouse, Météo-France, CNRS, Toulouse, France*
[5] *Earth Observation Research Center, Japan Aerospace Exploration Agency, Ibaraki, Japan.*

Capsule

A robust constellation of precipitation-capable satellite sensors is necessary to capture and map the spatial and temporal variability of precipitation across the Earth's surface.

Abstract

To address the need to map precipitation on a global scale a collection of satellites carrying passive microwave (PMW) radiometers has grown over the last 20 years to form a constellation of about 10-12 sensors at any one time. Over the same period, a broad range of science and user communities has become increasingly dependent on the precipitation products provided by these sensors. The constellation presently consists of both conical and cross-track scanning precipitation-capable multi-channel instruments, many of which are beyond their operational and design lifetime but continue to operate through the cooperation of the responsible agencies. The Group on Earth Observations and the Coordinating Group for Meteorological Satellites (CGMS), among other groups, have raised the issue of how a robust, future precipitation constellation should be constructed. The key issues of current and future requirements for the mapping of global precipitation from satellite sensors can be summarised as providing: 1) sufficiently fine spatial resolutions to capture precipitation-scale systems and reduce the beam-filling effects of the observations; 2) a wide channel diversity for each sensor to cover the range of precipitation types, characteristics and intensities observed across the globe; 3) an observation interval that provides temporal sampling commensurate with the variability of precipitation; and 4) precipitation radars and radiometers in low inclination orbit to provide a consistent calibration source, as demonstrated by the first two spaceborne radar/radiometer combinations on the Tropical Rainfall Measuring Mission (TRMM) and Global Precipitation Measurement (GPM) mission Core Observatory (CO). These issues are critical in determining the direction of future

constellation requirements, while preserving the continuity of the existing constellation necessary for long-term climate-scale studies.

1 Introduction

Water is not only a fundamental element of the Earth system, but also vital to all life on Earth. Consequently, the observation and measurement of precipitation (rainfall and snowfall) on a global scale is crucial to our understanding of the Earth system while impacting society across many levels (Kirschbaum et al. 2017; Skofronick-Jackson et al. 2017). Precipitation provides a direct link between the global cycles of energy and water through constraining and enabling the exchange of energy (Trenberth et al. 2009) and is the primary control of many natural hazards such as droughts and floods (Vicente-Serrano et al. 2010; Kundzewicz et al. 2014). However, the variability of precipitation makes it difficult to fully capture the characteristics of precipitation (Sidebar 1). Conventional measurements made by rain (and snow) gauges are generally representative of a very small area close to each gauge (Kyriakidis et al. 2001; Lundquist et al. 2019). While surface-based weather radar observations have limitations (Harrison et al. 2000; Ciach and Krajewski 1999), they provide valuable spatial measurements over large areas, complementing the global coverage provided by gauge data. Crucially, over the oceans few or no observations are available (or possible) through conventional means (Kidd et al. 2017). Even fundamental questions, such as addressing the amount and occurrence of precipitation, are generally limited to land regions (e.g. Sun et al. 2006; Herold et al. 2015). Observations provided by satellite sensors are therefore key in the measurement of precipitation on a global scale (e.g. Adler et al. 2003). Since no single satellite can possibly achieve this coverage alone, a stable and robust constellation is critical for providing sufficient temporal sampling to capture the vagaries of precipitation, particularly in surface-data sparse regions such as over the Poles or oceans, to provide consistent global precipitation products to the user community (Huffman et al. 2007). Furthermore, long-term climate studies into changes in precipitation across the Earth's surface can only be achieved through maintaining such a constellation (Adler et al. 2017; Levizzani et al. 2018).

2 Conventional precipitation measurements

2.1 Gauge measurements

The *de facto* measuring device for precipitation is the rain (snow) gauge, typically consisting of a funnel and a collection vessel, but with a wide range of designs (Strangeways 2004; Sevruk and Klemm 1989). The amount of water collected is usually measured daily, but less often in more remote regions. Automatic recording of the rainfall is possible with, for example, tipping bucket gauges which record the time/date that a known quantity of rainfall tips a small bucket (Sevruk 2005). Gauges are however subject to errors and uncertainties in their measurements (Ciach 2003; Villarini et al. 2008). The primary error of gauge measurements results from turbulence around the gauge orifice (Duchon and Biddle 2010) which may cause significant under-catch in light rain and/or strong winds (Kochendorfer et al. 2017). Measuring snowfall is more problematic due to the lower fall speeds of snowflakes

(Thériault et al 2012), resulting in the use of 'shields' around snow gauges to reduce the wind speed around the gauge orifice (Duchon and Essenberg 2001). Despite issues relating to their accuracy, gauges remain the mainstay of conventional global precipitation measurements, particularly when corrections are applied (Michelson 2004). Even so, their coverage across the globe is extremely variable. Over land, some regions have what could be considered adequate coverage, while other regions have none. Measurements over oceans are only available from a few atoll or island locations and a very small number of moored or drifting buoys. Overall, the global surface represented by gauge measurements is pitifully small (Kidd et al. 2017).

2.2 Weather radar measurements

The technological development of weather radar systems has created an important source of information on precipitation at local to regional scales (Whiton et al 1998a, 1998b). Unlike the point-measurements of gauges, radars can provide frequent, 3-dimensional observations of precipitation up to about 250 km from the radar location (Zhang et al. 2011; Zhang et al. 2016). However, radars rely upon several assumptions to convert the backscatter signal from rain and/or snow to an equivalent rain intensity (Campos and Zawadzki 2000; Uijlenhoet 2001; Uijlenhoet et al 2003). To avoid surface clutter the radar beam is usually elevated resulting in the altitude of the beam increasing with range, thereby no longer measuring surface precipitation and making near-surface phase (rainfall vs snowfall) detection difficult (Mimikou and Baltas 1996). Furthermore, quantitative radar measurements of precipitation are usually calibrated against gauge data where available, and due to the expense of installation, operation, and maintenance, weather radars tend to be clustered in developed countries, while their data may not be freely available.

2.3 Emerging systems

Over the last few years, a number of new surface-based instruments and techniques have shown merit for measuring precipitation and augmenting conventional measurements, particularly in regions with few or no surface observations. Over the oceans precipitation data are particularly scarce, being limited island-gauges, buoys, or ships, and are typically not be representative of the open ocean precipitation. Estimating rainfall through the use of underwater hydrophones is presented by Pumphrey et al. (1989), Medwin et al. (1992) and Forster (1994). More recently Ma and Nystuen (2005) found an excellent agreement between acoustic, gauge and TRMM satellite measurements, particular at higher rain rates. While very promising, the expense for deployment and maintenance of such instrumentation is significant, although the rewards are likely to be great. Over land, the attenuation of microwave communication signals by rainfall has been studied since the late 1960's (see Semplak and Turrin 1969). This path attenuation can now be fully exploited due to the current widespread mobile phone infrastructure, as exemplified by the work by Leijnse at al. (2007) over The Netherlands. Since the microwave paths are close to the surface, they tend to be more representative of surface rainfall than weather radars. A more comprehensive study by Overeem et al. (2011) showed very good correlations between the link-derived estimates

and those from the weather radar. Further development would complement and augment existing precipitation measurements in regions where few precipitation observations exist.

3 Satellite Precipitation Measurements

3.1 Satellite systems and retrievals

Precipitation-capable satellites may be classified by their orbit and sensing frequencies. Low Earth Orbiting (LEO) satellites orbit between 400-800 km altitude, provide about 14-16 orbits per day and typically operate in a Sun-synchronous orbit. Some low-inclination, non-Sun-synchronous satellite orbits provide observations at different times of day as the orbits precess over periods of weeks or months. The LEO satellites sensors generally provide a broad sub-satellite swath of data, although some only provide data over a narrow swath or only at nadir. The Cloud Profiling Radar (CPR) on CloudSat (Stephens et al. 2002; Stephens et al. 2018), although not designed to retrieve precipitation, has shown significant merit in observing and estimating light precipitation. Despite the nadir-only observations of the CPR, its increased sensitivity has proved invaluable at providing an additional calibration/validation data source for precipitation retrieval schemes, particularly for light precipitation and snowfall (Battaglia et al. 2020). LEO satellites most relevant for precipitation studies carry visible (VIS), Infrared (IR), passive microwave (PMW) and active microwave (AMW) sensors (Kidd and Levizzani 2011). Geostationary (GEO) satellites occupy a much higher orbit of about 35,800 km and are synchronised with the Earth's rotation so appear stationary over a fixed location at the Equator, allowing frequent and regular observations to be made. However, due to their altitude and requirements for sufficiently fine-resolution measurements, these observations are currently restricted to VIS/IR sensors. Mission concepts have been developed for geostationary PMW sensors (e.g. Tanner et al. 2007; Lambrigtsen et al. 2007; Duruisseau et al. 2017; Lambrigsten 2019) but these typically would only provide limited coverage at any one time and with limited frequencies. Table 1 summarises the range of satellite observing systems.

Satellite observations for precipitation estimation extend back over 40 years, with the longest data records based upon VIS and/or IR imagery. Early precipitation estimates relied upon empirical relationships between the cloud top (brightness and/or temperature) characteristics and surface rainfall (Kidd and Levizzani 2011). However, this relationship is generally poor and inconsistent over time and space (Kidd and Muller 2010; Kidd and Levizzani 2019). PMW radiometers, developed in the mid-1970s, rely upon the upwelling radiation from the Earth's surface that is largely unaffected by the presence of cloud, particularly at the lower frequencies (below 37 GHz). Sufficiently large liquid and ice particles (as in the case of precipitation) affect the upwelling radiation, resulting in increased radiation at the lower frequencies due to emission from liquid droplets, while at the higher frequencies ice particles cause a decrease in the upwelling radiation (Kummerow 2020). Although PMW observations are more direct than those of the VIS/IR, many assumptions are necessary to convert the satellite observations into precipitation estimates, particularly over land, where low frequency observations are impractical for retrieving precipitation. The most direct measure of

precipitation from space is obtained from precipitation radars and relies upon the backscatter from the precipitation sized particles to estimate precipitation intensity in the same way as their ground-based counterparts (see Battaglia et al. 2020). Two precipitation-specific radars have flown; the first, the Tropical Rainfall Measuring Mission (TRMM) Precipitation Radar (PR) operated at 13.6 GHz (Kummerow et al. 1998) and the Global Precipitation Measurement (GPM) mission Dual-frequency Precipitation Radar (DPR) operating at 13.6 and 35.5 GHz (Hou et al. 2014), while the cloud-orientated CPR operates at 94 GHz. As with ground-based radars, the backscatter-to-precipitation relationship is not consistent, and they cannot retrieve precipitation within about 500-2000m of the surface due to clutter (Skofronick-Jackson et al. 2019).

On a global scale the satellite-based observations largely address the shortcomings of conventional measurements for observing precipitation while also providing data for studying the characteristics and mechanisms of precipitation (Tapiador et al. 2011). However, to measure and map precipitation correctly, satellite systems must provide sufficient sampling to capture the temporal and spatial variability of precipitation.

3.2 Current precipitation constellation

The evolution of the constellation of precipitation-capable sensors is shown in Figure 1. Conically-scanning PMW sensors began in the late 1980's, with cross-track scanning PMW sounding instruments from the late 1990's onwards. The latter, although primarily designed for retrieving temperature and humidity (Mo 1995), have proved valuable in increasing the temporal sampling necessary for precipitation measurements (Kidd et al. 2016, 2021; Bagaglini et al. 2021). The launch of TRMM (Simpson et al. 1988; Kummerow et al. 1998) in 1997 facilitated multi-sensor retrievals through the inter-calibration of the then-available PMW sensors with the TRMM instruments. This inter-calibration concept was expanded with GPM (Hou et al. 2014; Skofronick-Jackson et al. 2017), launched in 2014.

The GPM-CO carries the DPR and the GPM Microwave Imager (GMI), both of which have been shown to be very well calibrated (Wentz and Draper 2016). Alongside the GPM-CO, a constellation of about 10-12 PMW-based precipitation-capable satellites is provided by several international agencies (see Table 2). Amongst these are the operational missions of NOAA (with the MHS and ATMS sensors), those of EUMETSAT (MetOp MHS) and the U.S. Department of Defense (DoD) SSMIS sensors (Kidd et al. 2020). In addition, JAXA contributes the AMSR-2, while the Centre National d'Etudes Spatiales (CNES) and the Indian Space Research Organisation (ISRO) contribute SAPHIR from the Megha-Tropiques mission (Roca et al 2015). Data archives of the observations from all of these and previous PMW missions provide the necessary input for routine global estimates throughout the entire record. Additional operational precipitation capable PMW missions exist (e.g. Li et al. 2018), but due to data access/usage arrangements these cannot yet be fully exploited by the wider precipitation community. The multi-agency co-ordination of the orbital crossing times of these missions affects the temporal sampling. For example, at present the EUMETSAT,

NOAA and JAXA missions generally use station-keeping to ensure consistent overpass times, while the crossing times of other PMW satellites drift over the course of 14-15 years between the extremes of ca.13:30 to 22:30 (ascending node). A number of missions (i.e. TRMM, GPM and Megha-Tropiques) had, or have, non-sun-synchronous precessing orbits and therefore observe the full diurnal cycle at any one location over the period of a few months, albeit with highly intermittent sampling (Roca et al. 2018).

In addition to the PMW sensors, a ring of GEO satellites provides frequent and regular vis/IR observations (see Table 3) that are used to augment the LEO PMW observations. These satellites' IR data are used as input for the NOAA Climate Precipitation Center (CPC) 4 km 30-minute global IR composite (Janowiak et al. 2001) and the GridSat collection at 10 km 3-hour subsampled data (Knapp and Wilkins 2018) that are often used for precipitation retrievals in conjunction with the LEO PMW observations. Merged PMW and IR satellite schemes such as the CMORPH (Joyce and Xie 2011), GSMaP (Kubota et al 2020) and IMERG (Huffman et al. 2020) allow precipitation products to be generated at resolutions of 30-minutes/10-km or better, which critically relies upon sufficient high-quality PMW instantaneous retrievals of precipitation.

One of the pressing issues of the current constellation is the age of the satellite missions. It is fortuitous that many of the precipitation-capable missions have lasted beyond their designed operational lifetime with their respective agencies providing support to keep them operating. However, before 2016 the mean age of these satellites was 5-7 years; after 2016 the age has slowly risen and now is over 9 years (see Figure 2). Crucially, satellites fail and sensors fail, often unexpectedly. A practical consequence of the gains and losses of these sensors is the direct impact upon the temporal sampling by the constellation. For example, a failure of the three oldest sensors in the constellation (the SSMISs on the DMSP F16 and F17 and the SAPHIR on Megha-Tropiques) would lead to a loss in the temporal sampling, extending the gaps between successive PMW observations, including the critical longest gaps, as illustrated in Figure 3. A concerted program of new satellites/sensors is therefore necessary to ensure the adequate continuation of precipitation measurements in a controlled and planned manner to support the user requirements discussed below.

3.3 Planned precipitation-capable missions

Going forward, there are several planned satellite missions that are capable of providing observations from which precipitation may be retrieved(see Figure 1). The precipitation-capable missions planned for launch over the next decade include:

EUMETSAT: The European Polar-orbiting System (EPS) Second Generation (SG) will provide continuity to the current MetOp series of satellites, and so the orbital characteristics are likely to be commensurate with the current missions, including station keeping (Accadia et al. 2020). The SG-A satellites will carry the cross-track MicroWave Sounder (MWS; 24 bands, 23-229 GHz, 40-17 km resolution), while the SG-B satellites will carry the conically

scanning MicroWave Imager (MWI; 18-bands, 18.7-183 GHz, 50-10 km resolution; Accadia et al. 2020) and the Ice Cloud Imager (ICI; 11 bands, 183-664 GHz, 16 km resolution; Eriksson et al. 2020). The SG-A and SG-B satellites are likely to be in the same orbital plane but ½ orbit offset allowing adjacent swath coverage. Precise launch dates are yet to be determined, but likely to be begin in the 2022/2023 timeframe.

NOAA: Additional members of the current Joint Polar System Satellite (JPSS) series (JPSS-2/-3/-4) through the 2030's are likely to occupy the same orbit as NPP and NOAA-20 (Goldberg et al. 2013). Each will carry the ATMS sensor (22 band, 23-183 GHz), with resolutions on the order of 16 km at nadir for the higher frequency channels, although the resolution at 88.2 GHz channel is 32 km.

US Department of Defense: The DoD has a long history of precipitation-capable missions (SSMI and SSM/I), and DoD has commissioned Ball Aerospace to build a new PMW imager: the Weather Satellite Follow-on – Microwave (WSF-M) (see Newell et al. 2020). The planned sensor is a 6 frequency, 17 channel radiometer covering frequencies between 10-89 GHz with a finest spatial resolution of 15x10 km. The contractual launch date is set as October 2023.

JAXA: The third generation of the Advanced Microwave Scanning Radiometer (AMSR-3) sensor (Kasahara et al. 2020) is being built for installation on JAXA's Global Observation SATellite for Greenhouse gases and Water cycle (GOSAT-GW) with a scheduled launch date in 2023. The AMSR-3 sensor provides similar channel selection as the other AMSR sensors (12-channel, 6.7-89 GHz), but with additional higher-frequency channels at 166 and 183 GHz.

NASA: Time-Resolved Observation of Precipitation structure and storm Intensity with a Constellation of small Satellites (TROPICS) is a NASA Earth Venture Mission, providing 6 (plus one pathfinder) cubesats, primarily focused on the evolution of weather systems across the Tropics (Blackwell et al. 2018). Each cubesat carries a small microwave radiometer operating at frequencies between 90 and 204 GHz with spatial resolutions similar to the MHS/ATMS sensors, and thus should be capable of providing observations for precipitation retrievals. An initial pathfinder mission is scheduled for launch in June 2021.

Chinese Meteorological Agency (CMA): The FY series of satellites carrying the MWHS-2 and the MWRI instruments (Guo et al. 2015; Lawrence et al. 2018) has a proven record of precipitation-capable missions. The planned Chinese Rain Mapping missions (FY-3I and FY-3J) are expected to have capabilities similar to the GPM core observatory, hosting the PMW MWRI/MWHS-2 sensors and a rain radar providing observations at 13.6 and 35.5 GHz, and are currently planned to launch in 2022/2023.

Other missions of interest being developed include the European Copernicus Imaging Microwave Radiometer (CIMR; Accadia et al. 2020), the ESA Arctic Weather Satellite (AWS) and EarthCARE, and the NASA-led Aerosols, Clouds, Convection, Precipitation mission (ACCP), which should be available for launch within the next decade. The Russian Meteor-N series also host both PMW imagers and sounders although with limited dataaccess. Many other small satellite missions (cubesats and smallsats) are proposed that could provide precipitation-relevant data (see Stephens et al. 2020). However, it is vital that such innovations support the spatial, temporal, channel, and quality requirements to address the needs of the user community, as well as facilitate the integration of their observations into the near real-time production of satellite precipitation products.

4 Defining future mission requirements

4.1 User requirements

The main drivers for all new missions are a necessary compromise between the requirements of the user communities, the engineering/physical constraints, and the available budget. For precipitation, the temporal and spatial resolutions for the observations (and hence derived products) is determined by the capabilities of the observing system and the needs of the user communities and their applications. Polls of the user community found significant variations in the requirements of temporal and spatial sampling, as well as the latency (observation-to-delivery delay) (see Table 4 and Figure 2, Friedl 2014) that depend on a huge range of research and application topics. Perhaps the most stringent requirements were associated with emergency managers, who require good spatial resolution (20 km or better), frequent (every few minutes) and immediate (within a few minutes) information. For comparison, products from the current constellation using combined PMW and IR observations provide products at ca. 10 km resolution, every 30 minutes, within about 4 hours, although very near real time IR-only products are available with reduced veracity. Importantly, the requirements within each user community are very diverse, and moreover, the scales may vary within a particular application, as demonstrated by Reed et al. (2015), who investigated the temporal sampling of precipitation datasets for hydrological modelling to meet the necessary flood-forecasting requirements. In that study, catchments with fast(slow) run-off characteristics require information more(less) frequently. Consequently, it is important to address the most stringent requirements as these are often those which have the greatest impact on both the physical and human environment, although in reality, the available scales and sampling are often coarser than the physical scales of precipitation and are further constrained by the physics and engineering of the sensors.

4.2 Addressing the temporal sampling and spatial resolution.

4.2.1 Spatial resolution at microwave frequencies is essentially limited by the size of the antenna (dish) that can be deployed, which is driven by current-generation engineering, launch vehicle and budget limitations. For a given antenna size, the best available resolution is inversely related to the observation frequency, meaning high-frequency channels have finer resolution than low-frequency channels for the same sized antenna. At present, the finest

PMW resolution is about 3x5 km at 89, 166, and 183 GHz using a 2 m dish (e.g. the GMI). Utilization of a larger dish is extremely problematic, both in terms of launch, and in terms of operation since the dish is continually rotating. Deployments of large (5-6m) mesh antennas are planned, such as the CIMR mission (Accadia et al. 2020), but are limited to the lower frequency channels (<37 GHz) by the current-generation limits to dish surface conformance. While GEO-based Synthetic Aperture systems have been studied for precipitation missions (e.g. Lambrigsten 2019), they require significant additional processing, are not truly global and envisage having only high frequency channels. Thus they presently do not necessarily provide significant advantage over LEO-based radiometers. Since the spatial variability of precipitation is on the order of a few km, resolving this variability at 1 km or less would be ideal for satisfying the more-stringent user needs, and remains a significant and unmet challenge.

4.2.2 Temporal sampling from LEO satellites is constrained not only by the physical number of satellite sensors, but also by the swath over which each sensor collects data. For example, the AMSR2 sensor (Imaoka et al. 2010) collects data across a 1450 km swath, while the GMI has a swath of 885 km. Temporal sampling of 3 hours is often quoted as a minimum requirement. This was originally based upon the number of samples necessary to reduce ambiguities when observing the diurnal cycle of precipitation (by providing at least the first three sinusoidal components), as well as being inherited from intermediate and Synoptic hours (00 UTC, 03 UTC, etc), which has been the basis for the WMO-agreed distribution of near-real-time GEO IR imagery. However, 3-hourly observations cannot adequately capture the true precipitation accumulation at the daily scale, as illustrated in Figure 4. At 1 km an accumulation of the 3-hourly sampled data attains a correlation of less than 0.5 against the daily precipitation total and, while coarser resolutions provide higher correlations, even at 25 km resolution this does not exceed 0.8. This issue is further emphasised in Figure 5, which shows the correlation of instantaneous PMW and IR estimates of precipitation against surface radar over the course of 1 day (25 July 2015) over the central UK. Each peak in the correlation represents the retrieval at a PMW sensor overpass. While the correlations between the satellite and surface radar products are good across all the satellites at the time of overpass, the correlation (in common with other measures) quickly deteriorates at times away from the time of the overpass. Even when the IR data are directly calibrated with the surface radar, as is done for Figure 5, the PMW products are much better. Despite a reasonable number of PMW overpasses being available (as is currently the situation), there are clear gaps (around 05:00 and 23:00 UTC) on this day where no PMW overpasses occur and satellite precipitation estimates typically have to rely more heavily upon IR data.

5 Strategies for maintaining a robust constellation

The current precipitation constellation provides a significant number of observations to generate precipitation estimates, but the continuity of these observations is very precarious. Historically, the precipitation community has become very adept at collecting observations from a diverse range of satellite missions and sensors, often not originally designed for the

retrieval of precipitation, as well as incorporating information from other water-related missions to help close the water-cycle loop (Brocca et al. 2014; Behrangi 2020), but there is presently no adequate substitute for a steady supply of precipitation-relevant observations. To ensure a continuation of precipitation measurements from satellite systems, a number of strategies need to be considered.

5.1 Maintaining/strengthening the constellation through new missions

This is undoubtedly the largest driver for maintaining the capabilities of the precipitation constellation. The current constellation is continuously aging, with many of the current sensors now more than 10 years old, well beyond their anticipated mission lifetime (see Figure 2). Given the length of time required to design, build, test and launch new sensors it is imperative that a long-term strategy for the constellation be devised and implemented. As discussed in Section 3.3, several operational missions are planned, such as the NOAA JPSS series, and the EUMETSAT EPS-SG series, but there are fewer dedicated long-term precipitation-specific missions with mapping capabilities. While many operational missions provide valuable observations for the precipitation community, such satellite sensors are not necessarily optimised for precipitation retrievals, nor do the orbital characteristics provide the frequent or uniform sampling necessary to capture the variability of precipitation. Observing System Simulation Experiments (OSSE) could be used (e.g. Chambon et al. 2014) to improve coordination between the satellite agencies is necessary to provide an optimal sampling strategy. Studies have been undertaken to assess the likely impact of the loss of one or more sensors within the constellation (Chambon et al. 2012).

5.2 Robustness through redundancy

Larger satellites are generally more robust at dealing with failures due to built-in redundancy, allowing multi-decadal records of observations to be collected, as in the case of TRMM. The long-term reliability of precipitation-capable cubesats and smallsats has yet to be fully evaluated, but it is anticipated that the orbital characteristics would determine the mission lifetime rather than system failure. Many operational meteorological satellites, such as the GEO missions (GOES, Meteosat, etc.) and the MetOp and NOAA polar-orbiting missions have on-orbit backup sensors. However, when the backup satellites continue to collect data while the primary satellite is functioning, they generally do so over the same space/time domains of the primary missions and therefore contribute relatively modestly to the overall temporal sampling.

5.3 Extended mission lifetimes

Allowing missions to continue past their design lives and to continue contributing to the precipitation constellation is a proven concept. However, the extension of missions is often fraught with obstacles, not least being the need to comply with the modern standards for end-of-life disposal of satellites to avoid space junk (Crowther 2002; Witze 2018). On a positive note, the precipitation community has, in part, been successful in persuading agencies to keep such missions operating after the end of their designed lifetime (e.g. TRMM), with partial

failures (e.g. Megha-Tropiques), or after station-keeping fuel was exhausted (e.g. MetOp-A). The extension of mission lifetimes has only been, and likely will be, an issue for the larger satellite systems.

5.4 Retrieval scheme resilience

Many current retrieval schemes require a full set of observations to generate a precipitation estimate. However, many schemes could still contribute useful information with fewer channels. Figure 6 shows a range of retrieval scenarios using different channel combinations. The loss of a single channel on a diverse-channel sensor, such as the GMI, only degrades the retrieved precipitation marginally. The better performance of precipitation retrievals from the observations gathered over a wide range of frequencies has been shown by Kidd et al. (2018). Additionally, the flexible utilisation of channels in the retrievals has particular merit when dealing with surface-based Radio-Frequency Interference (RFI; Wu and Weng 2011), which necessitates the exclusion of certain channels from retrieval schemes at certain times/locations. Furthermore, the calculation of the errors and uncertainties associated with the retrieved precipitation is urgently needed to allow users to assess the usefulness of different retrievals. New techniques should also be investigated and developed that merge observational data before the retrieval stage, rather than merging precipitation estimates post-retrieval. The last point might be a longer-term goal, but it is possible to envisage a scenario where two satellites in very similar orbits, both experiencing channel degradation could jointly provide the capabilities of a single sensor.

5.5 Data availability and access

While most precipitation-related satellite observations are freely available both in terms of being available and accessible to any particular user, and for that user to share more widely, there are many data sets that are more restrictive and may not be accessible to all potential users. Furthermore, the access to such data sets in very near real time is of great importance to many user applications, such as flood forecasting, to ensure timely integration into their processing systems. Crucially, science works best when such data are accessible to the community, as shown by the open release of the DMSP SSMI data in 1987 by the US Department of Defense, which enabled the careers many of the current generation of precipitation scientists and the development of their retrieval schemes.

6 Recommendations

Based upon the current precipitation constellation and planned missions, the following course of action is necessary to ensure the long-term continuity of global satellite precipitation observations:

i) reaffirm commitment and support for current and planned precipitation-capable missions with free and open data sharing by the appropriate agencies and organisations;
ii) develop a long-term strategy for a viable constellation of precipitation-capable sensors that meet the necessary scientific and user requirements. Specifically,

a) PMW sensors with diverse channels covering the primary precipitation-sensitive frequencies with good spatial resolution as exemplified by the AMSR/GMI class of sensors, and
b) operational AMW sensors in a non-Sun-synchronous orbit for cross-calibration and reference standard for PMW (and IR) precipitation estimates, as exemplified by the mapping capabilities of the PR/DPR and the sensitivity of the CPR;

iii) support the continuation of precipitation-capable missions beyond nominal mission lifetime operations, with due regard for the limitations imposed by deorbiting/sensor degradation considerations;

iv) integrate new technologies, such as smallsats and cubesats, with access to new data sets where these address the necessary scientific and user requirements; and

v) implement mitigation strategies within the precipitation retrieval schemes to maximise use of sub-optimal observations, including failed/denied channels, to help ensure continuity in adequate sampling.

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787

Sidebar 1:

Precipitation Characteristics

The requirements for observing precipitation are driven by the very specific scale and characteristics that precipitation exhibits, creating a challenging statistical problem.

At the micro-physical scale water is the only element that co-exists in all three phases, namely vapour, liquid, and solid, a situation that is common in many precipitation systems. Liquid-phase water starts with the formation of water droplets (about 10 μm), usually on cloud condensation nuclei (c.0.1 μm), through the growth of cloud droplets, to precipitation-sized particles (c.100 μm, up to 4-5 mm; Ma et al. 2019). Smaller droplets are spherical, but large droplets become flatter or even umbrella-shaped due to air resistance prior to breakup. Pristine ice-phase water particles have similar growth, except physical aggregation of particles and interactions with liquid (riming) also occur. The resulting icy particles exhibit a vast range of shapes and sizes, with significant implications for the retrieval and estimation of (frozen) precipitation.

At the precipitation system scale the mechanisms driving the microphysical processes that form clouds and falling precipitation result in variations in precipitation that range from a few metres to 1000 km or more, and from a few seconds through to days, weeks, and longer (Trenberth et al. 2009). Thus, the observation of precipitation is very much affected by interaction between the time/space scales of the rain being observed and the resolution/sampling of the observing system (Luini and Capsoni 2012).

The precipitation statistics in space/time are unusual in that the normal/modal value is zero: for most of the time, and for the majority of the globe it is not raining/snowing. Furthermore, when precipitation does occur, it is heavily skewed towards light precipitation intensities, while the accumulation of precipitation (being a function of occurrence × intensity) is more log-normally distributed (see Figure SB1). As instantaneous samples are accumulated over time and space, the distribution shifts towards a more normal distribution. This complicates any statistical evaluation and requires extreme care when analysing or evaluating precipitation data sets. Specifically, the distribution of precipitation intensities is very much dependent upon the spatial and temporal scales being considered (Luini and Capsoni 2012), such that observing a precipitation system at the same time with sensors with different spatial resolution will yield different results. Similarly, comparing instantaneous precipitation with precipitation accumulated over a few minutes, days, or months will reveal very different characteristics. The practical implication is that instantaneous point measurements cannot be directly compared with those collected over an area and/or over time. This issue in compounded by the fact that precipitation events tend to last for periods ranging from a few minutes to hours.

Table 1: Characteristics of present-day LEO and GEO satellites and their observational capabilities.

	Orbit	**LEO**[1]	**GEO**[2]
Band	**Characteristics**	400-800 km altitude orbits c.14-16 orbits/day Polar orbiting Sun-synchronous (up to 2 overpasses/day), or Low inclination precessing orbits	c.36,000 km altitude orbits Geosynchronous orbit – satellite remains stationary relative to sub-satellite point.
Vis/IR	Cloud top properties Reflection, emission (T) Texture & particle sizes	Multi-spectral Orbital swaths <1 m to 1 km	Multi-spectral Frequent/regular samples over sectors or full disk coverage <1-4 km footprints at sub-satellite point
PMW	Hydrometeor column Liquid, ice and water vapour	Multi-channel Orbital swath c.5 – 70 km footprints	*Not possible at present* *Several feasibility studies*
AMW	Vertical profiles of hydrometeors Backscatter Liquid, ice	Single/dual frequency (13.6/35/94 GHz) Single beam (CloudSat with 1.4/5.4 km footprints) or narrow swath (DPR 245 km swath and 5.4x5.4 km footprints)	*Not possible at present* *Some feasibility studies*

[1] LEO: Vis/IR observations provided by a large group of sensors ranging from land surface monitoring missions (e.g. Landsat) through to meteorological missions (e.g. AVHRR) with resolutions typically matched to user requirements. PMW channels used for the retrieval of geophysical and meteorological parameters, including precipitation.

[2] GEO: constellation of GEO satellites provides global coverage with frequent/regular observations, disseminated at a nominal 3 hourly interval, but with sensors typically providing operational 10-15 min. data collection, and rapid scans <1 minute.

Table 2: Microwave sensors contributing to the GPM precipitation constellation. The current precipitation constellation missions are highlighted in bold. *(*retrieval resolution is that of the NASA GPROF scheme).*

Satellite	Agency	Sensor	number	Channels	Retrieval resolution*
AMW instruments					
GPM	**NASA/JAXA**	**DPR**	**x1**	**13.6/35.5 GHz**	**5.4 x 5.4 km**
TRMM	*NASA/JAXA*	*PR*		*13.6 GHz*	*4.3 x 4.3 km*
PMW imagers					
GPM	**NASA/JAXA**	**GMI**	**x1**	**10.7-183.31 GHz**	**10.9 x 18.1 km***
DMSP F16,17,18,*19*	**US DoD**	**SSMIS**	**x3**	**19.35-183.31 GHz**	**45 x 74 km***
GCOM-W1	**JAXA**	**AMSR2**	**x1**	**6.7-89.0 GHz**	**14 x 22 km***
TRMM	*NASA/JAXA*	*TMI*		*10.7-89.0 GHz*	*20.9 x 34.6 km**
PMW sounders					
NOAA-*18***,19 METOP-***A***,B,C**	**NOAA/EUMETSAT**	**MHS**	**x3**	**89.0-183.31 GHz**	**17.12 x 21.64 km***
NPP, NOAA-20	**NOAA**	**ATMS**	**x2**	**23.0-183.31 GHz**	**16.51 x 16.22 km***
MeghaTropiques	**ISRO/CNES**	**SAPHIR**	**x1**	**183.31 GHz (x6)**	**7.34 x 7.27 km**

850 Table 3: Current GEO VIS/IR sensors: those that actively contribute to the global 30-
851 minute 4 km IR imagery are highlighted in bold. *(All of these provide multi-channel*
852 *VIS/IR observations with temporal sampling of 15 minutes or better, and < 1km / 4km*
853 *for VIS/IR at sub-satellite point).*
854

Satellite	Agency	Sensor	Longitude	Channels/number		Sub-satellite resolution
GOES-13 (Storage)	NOAA	IMAGER	60°W	VIS/IR	x5	1 km / 4 km
GOES-14 (backup)	NOAA	IMAGER	105°W	VIS/IR	x5	1 km / 4 km
GOES-15 (West backup)	NOAA	IMAGER	128°W	VIS/IR	x5	1 km / 4 km
GOES-16 (GOES-East)	**NOAA**	**ABI**	**75.2°W**	**VIS/IR**	**x16**	**0.5 km / 2 km**
GOES-17 (GOES-West)	**NOAA**	**ABI**	**137.2°W**	**VIS/IR**	**x16**	**0.5 km / 2 km**
Meteosat-8 (IODC)	**EUMETSAT**	**SEVIRI**	**41.5°E**	**VIS/IR**	**x12**	**1 km / 3 km**
Meteosat-9 (rapid scan)	EUMETSAT	SEVIRI	3.5°E	VIS/IR	x12	1 km / 3 km
Meteosat-10 (rapid scan)	EUMETSAT	SEVIRI	9.5°E	VIS/IR	x12	1 km / 3 km
Meteosat-11	**EUMETSAT**	**SEVIRI**	**0°**	**VIS/IR**	**x12**	**1 km / 3 km**
Himawari-8	**JMA**	**AHI**	**140.7°E**	**VIS/IR**	**x16**	**0.5 km / 2 km**
Himawari-9 (standby)	JMA	AHI	140.7°E	VIS/IR	x16	0.5 km / 2 km

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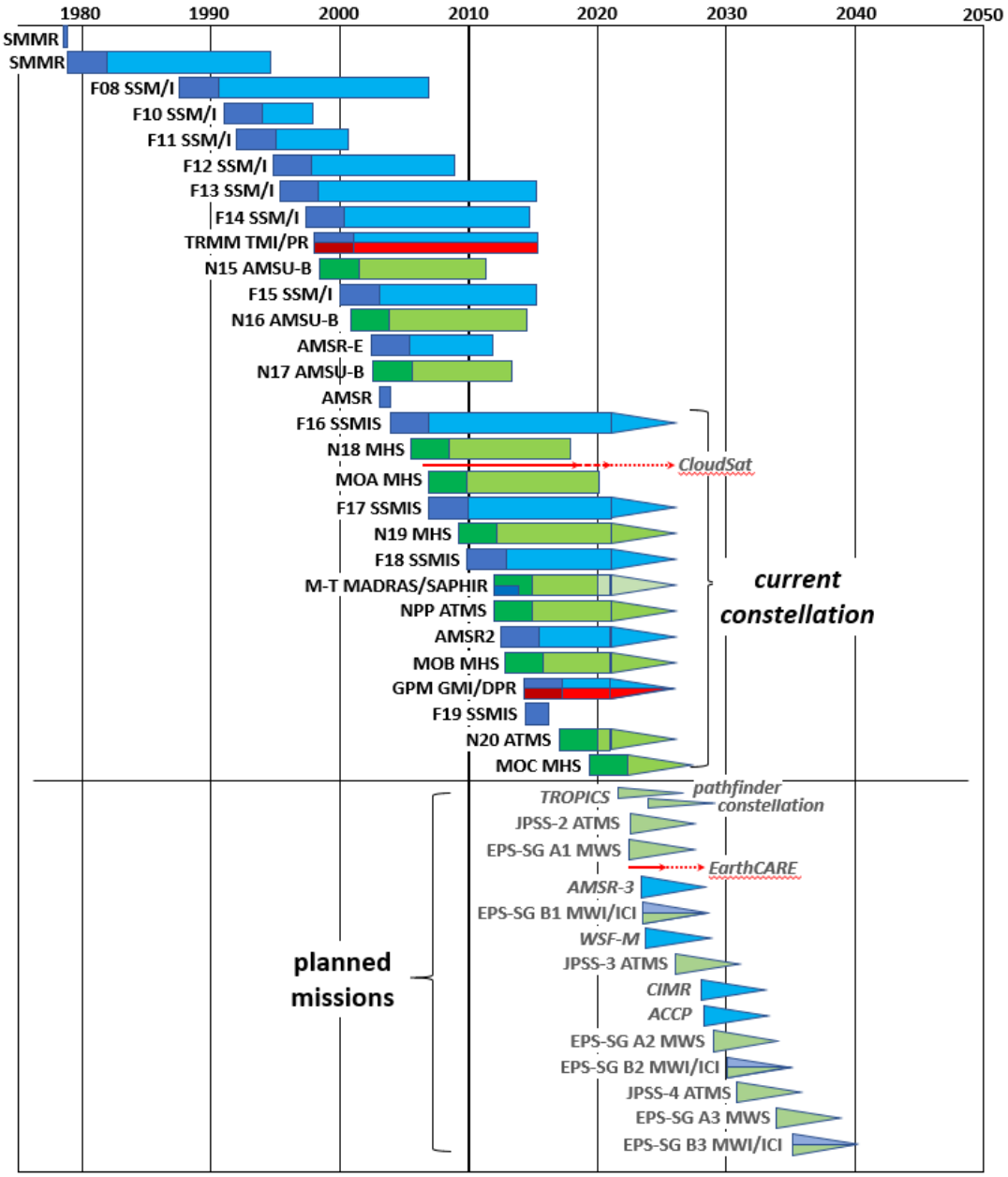

Figure 1: Timeline of PMW satellites and sensors providing analysis-ready data. The wide bars and arrows indicate swath-based or nadir-only observations respectively. Blue bars represent imaging/conical scanning radiometers, green bars represent sounding/cross-track radiometers and red bars indicate AMW (radar) sensors. Triangles indicate those sensors that currently provide data (as of 2021-02-09) and may continue to do so, together with future missions. *(Data source: based upon World Meteorological Office (WMO) Observing Systems Capability Analysis and Review Tool (OSCAR) database and EUMETSAT)*

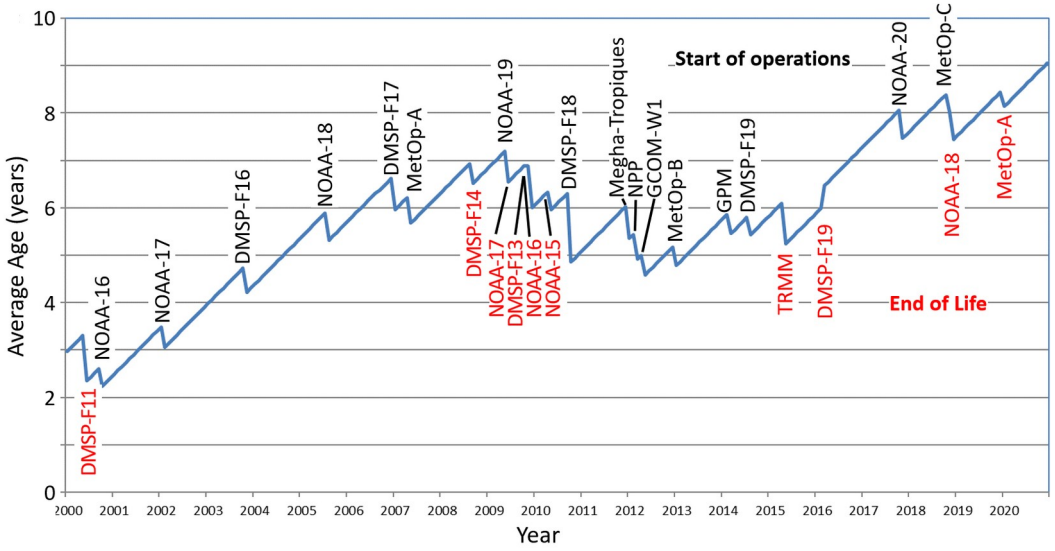

Figure 2: Mean age of the PMW satellites up to 1st January 2021. From 2005-2016 the mean age was around 5-7 years, but now currently exceeds 9 years. Note that the constellation becomes younger both with a new launch and if a long-term mission reaches its end of life.

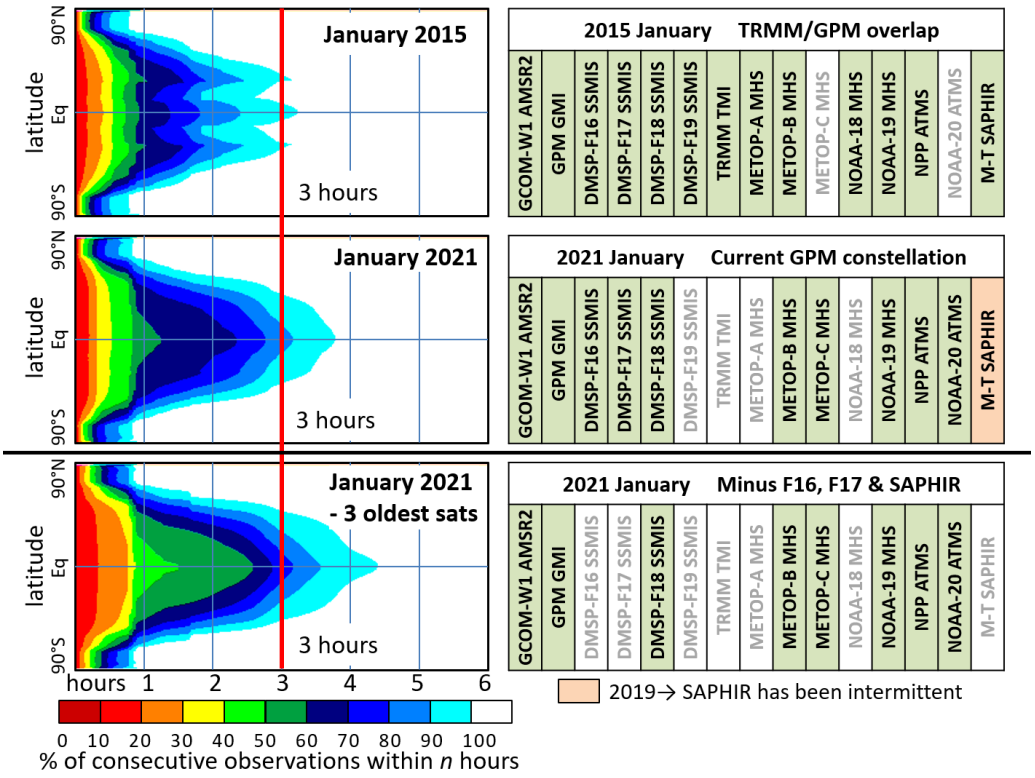

Figure 3: Comparison of the revisit times by latitude for three selected dates: the baseline sampling of the GPM mission is shown in the top row for January 2015, while the current sampling is shown in the middle row for January 2021. The bottom row shows a possible scenario if data from the three oldest sensors (SSMIS F16 and F17, and SAPHIR) are not included. The red vertical line represents the widely accepted 3-hour minimum revisit time necessary to adequately capture the accumulation of precipitation at daily, 0.25° scales. While the 3-hour revisit time was attained more than 90% of the time in January 2015, the reduction in the constellation numbers has reduced this to about 80% in January 2021.

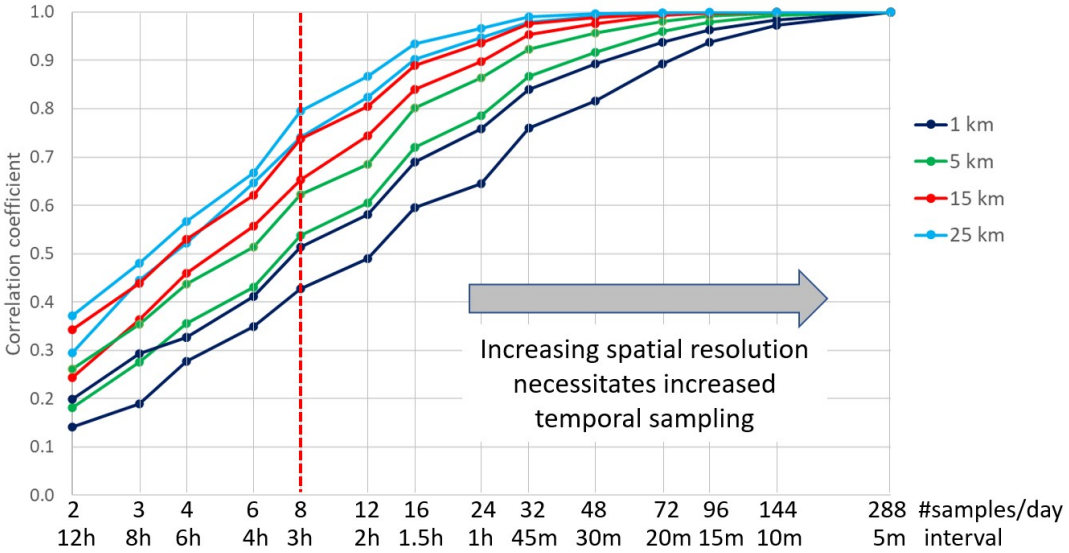

Figure 4: correlation between temporally sub-sampled 5-minute 1-km radar data and the daily total (accumulating all 5-min. samples) for two 300x300 km areas (for each resolution) over the central UK during 2019. Despite the different relationships within each resolution category due to regional variations in precipitation characteristics, the general trend is that finer spatial resolutions require more samples. Note that for the nominal 3-hourly sampling (8 samples/day) the correlation is only between 0.4-0.5 at 1 km resolution, and no better than 0.8 at 25 km resolution.

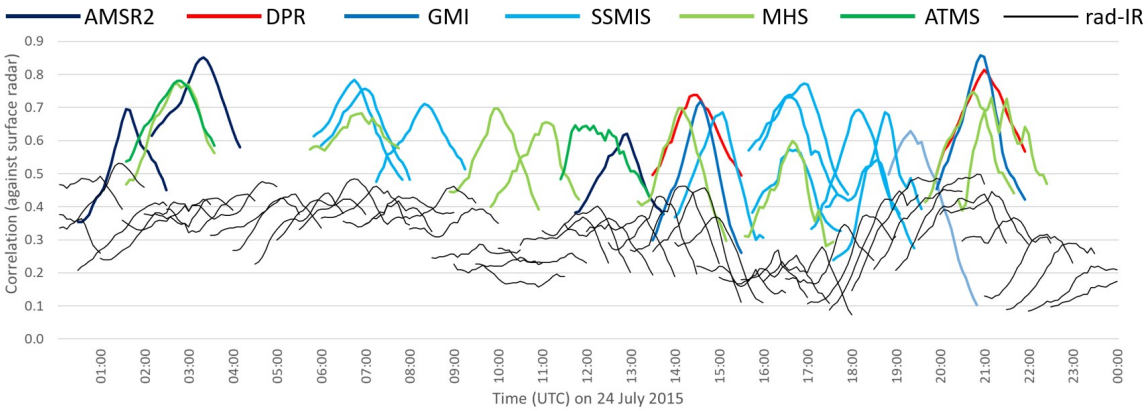

Figure 5: Correlation of satellite precipitation retrievals from IR, passive, and AMW sensors against surface radar for 24 July 2015 over the central UK. Each line segment covers a window over which the satellite estimate is compared with radar data up to ±1 hour centred on the observation time of the satellite sensor. The microwave measurements generally have a good correlation at overpass time, but this falls quite quickly on either side of the observation time. The estimates derived from the radar-calibrated 30-minute IR data are much poorer (rad-IR), as shown by the thin black lines.

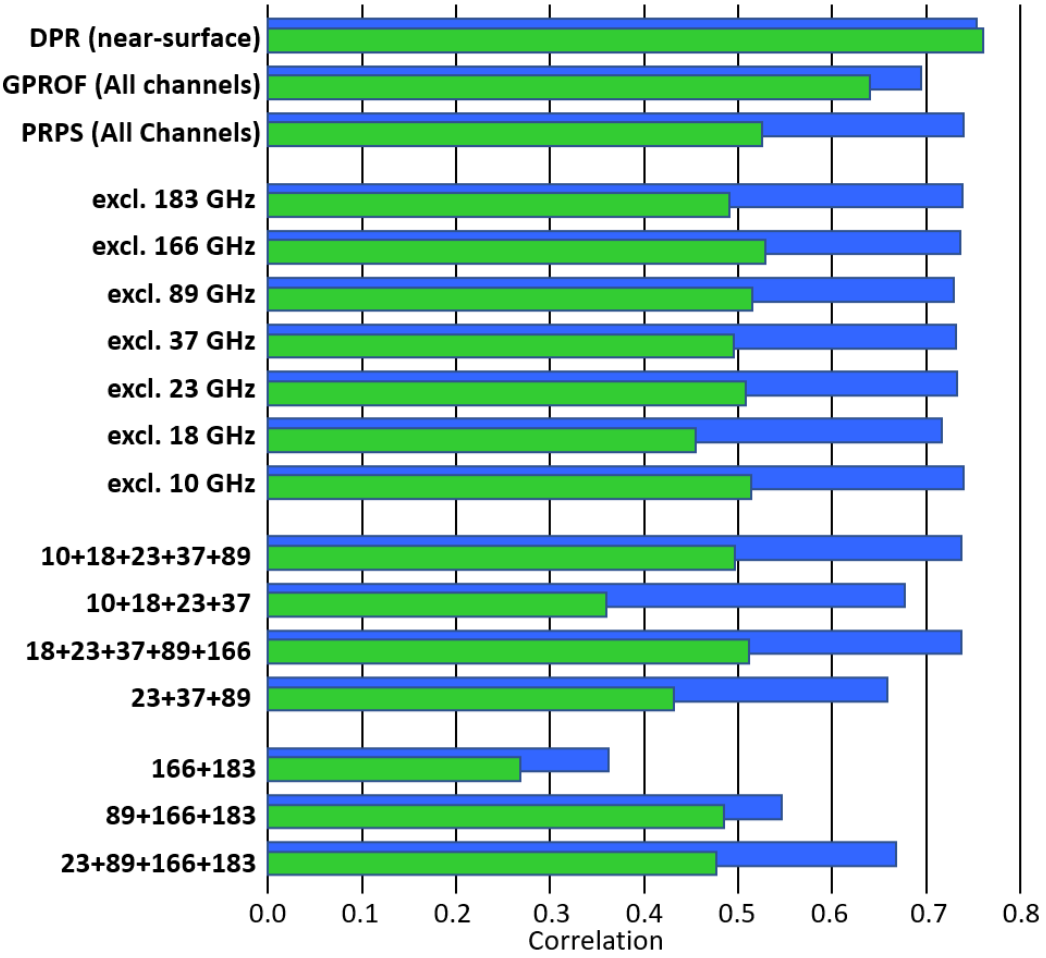

Figure 6: Scenarios used to study the effects of channel loss and channel combinations, based upon instantaneous PRPS retrievals from the 13-channel, 7-frequency GMI sensor, compared with surface radar data over the United Kingdom for 2017. The blue bars are for over ocean retrievals, while the green bars are for over land. The top three retrievals (DPR, GPROF and PRPS, where the last two are computed from GMI) can be used as benchmarks. The exclusion of a single channel has relatively little effect, except for the 18 GHz primarily over land. Generally, the more channels that are available, the better the retrieval as seen in the multi-frequency 10-89 or 18-166 GHz retrieval. A narrow frequency range results in poorer performance, particularly at high frequency channels, although the inclusion of a low frequency water vapour channel does significantly improve the performance over the ocean.

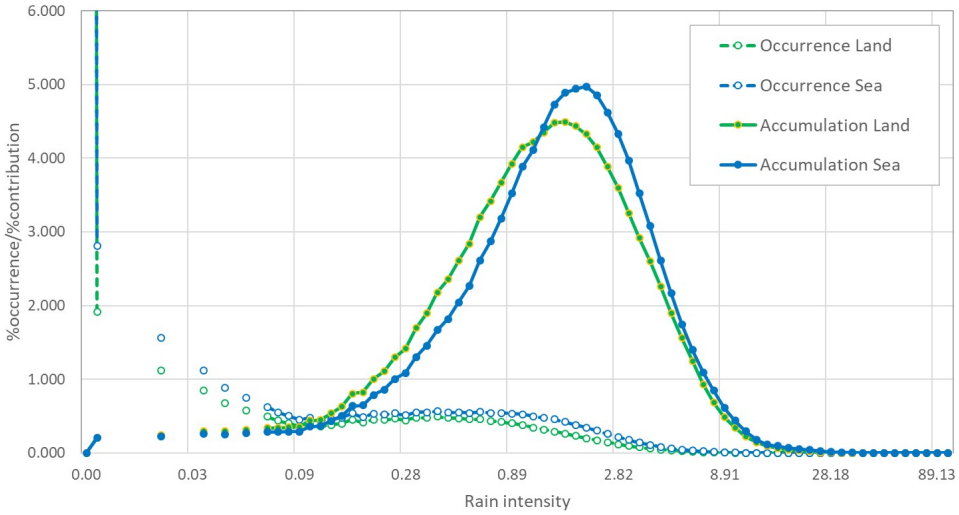

Figure SB1: Distribution of the occurrence/contribution of precipitation by intensity based upon instantaneous surface radar data over the United Kingdom (2014-2019) at 15 km x 15 km resolution. Note that the precipitation intensity is plotted on a log scale; the occurrence of precipitation is highly skewed towards zero.