Cross-Instrument Comparison of MapCam and OVIRS on OSIRIS-REx

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18 Abstract

- 19 Two of the instruments onboard the OSIRIS-REx spacecraft, the MapCam color imager and the OVIRS visible and
- 20 infrared spectrometer, observed the surface of asteroid (101955) Bennu in partially overlapping wavelengths.
- 21 Significant scientific advances have been enabled by using data from these two instruments in tandem, but a robust
- 22 statistical understanding of their relationship is needed for future analyses to cross-compare their data as accurately
- and sensitively as possible. Here we present a cross-instrument comparison of data acquired by MapCam and
- 24 OVIRS, including methods and results for all global and site-specific observation campaigns in which both
- 25 instruments were active. In our analysis, we consider both the absolute radiometric offset and the relative
- 26 (normalized) variation between the two instruments; we find that both depend strongly on the photometric and
- instrumental conditions during the observation. The two instruments have a large absolute offset (>15%) due to their
- 28 independent radiometric calibrations. However, they are very consistent (relative offset as low as 1%) when each
- 29 instrument's response is normalized at a single wavelength, particularly at low phase angles where shadows on
- 30 Bennu's rough surface are minimized. We recommend using the global datasets acquired at 12:30 pm local solar
- 31 time for cross-comparisons; data acquired at higher phase angles have larger uncertainties.

32 **1** Introduction

33 The Origins, Spectral Interpretation, Resource Identification, and Security–Regolith 34 Explorer (OSIRIS-REx) spacecraft (Lauretta et al. 2021; Lauretta et al. 2017) observed the 35 surface of asteroid (101955) Bennu, a B-type near-Earth asteroid (Clark et al. 2011; 36 Hergenrother et al. 2013; Lauretta et al. 2019a), for approximately two years before sampling 37 regolith from its surface on 2020 October 20. Though the observations taken during those two 38 years were primarily driven by the need to identify a safe and sampleable surface location, they 39 also provided a tremendous dataset for scientific analysis of the asteroid and scientific context 40 for the sample. These data revealed Bennu to have a diverse surface with macroscopic 41 heterogeneity in albedo (Golish et al. 2021c; Lauretta et al. 2019a), color (DellaGiustina et al. 42 2020), composition (Simon et al. 2020b), photometric response (Golish et al. 2021b; Li et al. 43 2021; Zou et al. 2021), physical structure (Rozitis et al. 2020; Scheeres et al. 2020), and texture 44 (Bennett et al. 2021; Walsh et al. 2019). Some of these characteristics are enhanced by Bennu's 45 dynamic history, which includes relatively recent surface changes due to mass movement (Jawin 46 et al. 2020), thermal fracturing (Molaro et al. 2020), impacts (Ballouz et al. 2020), and ongoing 47 particle ejections (Hergenrother et al. 2020; Lauretta et al. 2019b). The MapCam imager of the 48 OSIRIS-REx Camera Suite (OCAMS; Rizk et al. 2018) and the OSIRIS-REx Visible and 49 Infrared Spectrometer (OVIRS; Reuter et al. 2018) provided the data underlying many of these 50 discoveries.

51 MapCam imaged Bennu's surface in visible (VIS) wavelengths also observed by OVIRS, providing the opportunity for a direct comparison. The two instruments were designed as 52 53 complements to each other, with MapCam providing broadband spectrophotometric data at high 54 spatial resolution and OVIRS providing high spectral resolution, that extends into the near-55 infrared (NIR), with coarse spatial scales. Several studies performed during Bennu's proximity 56 operations used the instruments' complementary designs to strengthen their analyses (e.g., 57 DellaGiustina et al. 2021; Kaplan et al. 2020). Though both instruments went through extensive 58 ground and in-flight calibration campaigns (Golish et al. 2020; Rizk et al. 2018; Simon et al. 59 2018; Simon et al. 2021), those calibrations have independent uncertainties, and no formal 60 attempt has previously cross-calibrated the instruments. Moreover, both instruments have 61 idiosyncrasies that are documented in their individual calibrations, but not with respect to each 2

other. Here, we take advantage of concurrent observations by MapCam and OVIRS to perform a
comparison of the datasets. A quantitative comparison enables more in-depth studies of Bennu's
surface variation, taking advantage of the corresponding capabilities of the two instruments to
perform high spatial and spectral resolution analyses.

66

1.1 THE OSIRIS-REX CAMERA SUITE

67 OCAMS is a suite of three scientific imagers designed with individual and overlapping 68 capabilities (Golish et al. 2020; Rizk et al. 2018). PolyCam is a narrow-angle panchromatic 69 camera used to create high-resolution global and regional maps of Bennu's surface. MapCam is a 70 medium-angle camera with a series of optical filters used to make color maps of Bennu's 71 surface. SamCam is a moderately wide-angle, panchromatic camera used during and after the 72 sampling event. For the purposes of the comparison with OVIRS, we considered only MapCam. 73 Though the wavelengths imaged by the panchromatic filters in PolyCam and SamCam overlap 74 with OVIRS's spectral sensitivity, the bandwidth of those filters (~0.300 µm) is sufficiently 75 broad that a consistent radiometric calibration is challenging for either instrument. More 76 importantly, the color radiometric comparison is more relevant for most analyses that might combine data from both instruments to achieve high spatial and spectral resolution. 77 78 MapCam has four narrowband color filters and one wideband panchromatic (pan) filter. The 79 effective wavelengths of the filters are 0.473, 0.550, 0.698, 0.847 and 0.646 µm for the b', v, w, 80 x, and pan filters, respectively (Golish et al. 2020). The filter cut-on/off wavelengths are 0.439– 81 0.500, 0.521–0.578, 0.671–0.731, 0.815–0.893, and 0.489–0.815, respectively. These filters are 82 comparable to the Eight Color Asteroid Survey bands (Zellner et al. 1985) and were selected to 83 capture spatially resolved variations in Bennu's spectral slope and band ratios in the visible 84 wavelengths (DellaGiustina et al. 2018).

Additional effort was put into radiometric calibration of MapCam during ground and inflight calibration due to the sensitivity of color and color ratio mapping of planetary surfaces
(DellaGiustina et al. 2020). The calibration effort (Golish et al. 2020) utilized images of Earth's
Moon acquired during the OSIRIS-REx Earth gravity assist (Lauretta et al. 2018) and a Robotic
Lunar Observatory (ROLO; Buratti et al. 2011) model of lunar albedo and photometry.
Unfortunately, the Moon presented a small target in MapCam's field of view (~40 pixels across)

and did not provide strong statistics for the calibration. Moreover, MapCam imaged the Moon at a very different sub-spacecraft latitude and longitude than ROLO (which observes from Earth). We applied photometric corrections to the ROLO data to match the conditions of MapCam's observation, but that process is also very sensitive to the resolution of the image. As a result, the lunar calibration predicted a moderate absolute radiometric uncertainty (\pm 5%, 1 σ). However, MapCam's four filters share that absolute uncertainty, such that the calibration estimated a low relative (filter-to-filter) radiometric uncertainty of <2%.

98 The OCAMS imagers have a number of second-order effects that can increase the 99 uncertainty of the radiometric measurements, depending on the conditions of the observations. 100 The OCAMS calibration pipeline does not correct detector non-linearity. The OCAMS detectors 101 are >99.5% linear over most of their dynamic range, but become increasingly non-linear when 102 measuring very high or very low signals (Golish et al. 2020). Nearly all OCAMS images were 103 acquired with exposure times that captured the bulk of the surface within the linear regime. 104 However, extremely bright exogenic material (DellaGiustina et al. 2021) and deep shadows were 105 sometimes imaged with non-linearity greater than 2%.

All OCAMS detectors experience artifacts referred to as *icicles* in images acquired with extremely low exposure times (<3 ms; Golish et al. 2020). OCAMS only acquired images with these exposure times when longer exposures would overexpose portions of the surface. This occurred only for the panchromatic filters of MapCam and PolyCam at low phase angles. For the purposes of this study, icicles were only present for images acquired at 12:30 pm local solar time (Section 1.4) with MapCam's pan filter.

112 MapCam also has some out-of-field stray light that couples to the detector (Rizk et al. 2018). 113 The stray light is primarily noticeable when there is a bright source just outside MapCam's field 114 of view, such as when Bennu is larger than the field of view. The noise due to stray light is <1%115 and is not significant in single-filter images and mosaics, which typically have a signal-to-noise 116 ratio of <1% (DellaGiustina et al. 2020). However, the amount of stray light is wavelength-117 dependent. Therefore, 0.5% variations due to stray light can add significant noise when 118 calculating color ratios (DellaGiustina et al. 2020), which measure variations on the order of a 119 few percent.

120 **1.2** THE OSIRIS-REX VISIBLE AND INFRARED SPECTROMETER

121 OVIRS is a point spectrometer with a field of view of 4 mrad and a spectral range of 0.4 to 122 4.3 um; the full spectrum is obtained simultaneously for each 4 mrad spot (Reuter et al. 2018). 123 OVIRS achieves this spectral range with a series of wedged filters that split five overlapping 124 segments of the full spectral range onto different regions of a Teledyne H1RG infrared detector. 125 The detector is cooled with a passive radiator to reduce dark current and the optics are thermally 126 isolated from the spacecraft deck (Reuter et al. 2018). The first two segments (1a from 0.392-127 0.670 and 1b from 0.652–1.090 µm) overlap MapCam's color filters. Importantly, the OVIRS 128 segments image to different locations on the detector in the following order: 1a, 2, 3, 4, 1b. As a 129 result, the two short-wavelength segments are on opposite ends of the detector and may image 130 slightly different regions on the surface when the spacecraft is slewing (Simon et al. 2021). In 131 locations with a sharp discontinuity on the surface (e.g., a deep shadow), the two segments can 132 measure substantially different signals. The boundary between the two segments is 133 approximately at the low-wavelength cutoff of MapCam's w filter, making segment-related 134 artifacts manifest differently when comparing the b' and v filters with the w and x filters. 135 The main science objective of the OVIRS instrument was to detect spectral features and 136 spectral variability of the surface (Kaplan et al. 2020; Lauretta et al. 2021; Simon et al. 2020b), 137 requiring high relative (channel to channel) accuracy (2%) and moderate absolute accuracy (5%). 138 OVIRS's wavelength range was selected to capture Bennu's overall VIS-NIR spectral slope and 139 detect absorption features due to hydrated minerals (e.g., 0.7 and 2.7 µm) and organic molecules 140 (e.g., between 3.3–3.6 µm) (Reuter et al. 2018). The OVIRS ground calibration was performed 141 during environmental testing with NIST-traceable sources and showed excellent relative 142 (channel to channel) radiometric accuracy and precision (<1%; Simon et al. 2018; Simon et al. 143 2021). However, the ground equipment did not cover all wavelengths, and post-testing issues 144 were found with the short wavelength source (Simon et al. 2018; Simon et al. 2021). Data of the 145 Earth were used to adjust the wavelength and radiometric calibration in flight; however, the 146 available dark ocean views were not ideal for cross-calibration with Earth-viewing satellites 147 (Simon et al. 2018). Final adjustments to the radiometric calibration were made using the 148 asteroid itself, based on Earth-based reflectance data, improving calibration in the 2 to 2.5-149 micron region, but leaving the absolute radiometric accuracy less well defined (>5%).

150 Additionally, the OVIRS radiometric uncertainty increases when the OVIRS detector is 151 outside its nominal temperature design range (90–105 K), because the detector loses long 152 wavelength sensitivity at higher temperatures, making out-of-band filter effects at all 153 wavelengths more difficult to characterize. This thermal effect was a minor issue in global 154 imaging campaigns, where the detector maintained a temperature around 105 K, primarily due to 155 parasitic heat from the spacecraft itself (Kaplan et al. 2020; Simon et al. 2020b). When the 156 spacecraft was closer to Bennu, however, radiator views of the hot surface caused an increase in 157 the OVIRS detector temperature, increasing the radiometric uncertainty (Simon et al. 2021).

158 **1.3 SCIENTIFIC ADVANCES MADE POSSIBLE BY INSTRUMENT COMPARISON**

159 Much scientific progress has already been enabled by using MapCam and OVIRS data 160 together. The high spatial resolution of MapCam color images provides a guide for interpreting 161 the geologic context of OVIRS data, whose spectrometer spot size is $\sim 60 \times$ larger than the 162 MapCam pixel scale. Additionally, the broader wavelength range provided by OVIRS can be 163 used to definitively link MapCam color signatures to compositional units, thereby extending the 164 spatial scale where we can discern composition on Bennu. Concurrent observations by MapCam 165 and OVIRS that reveal the same phenomena independently confirm one another. Because of 166 these complementary aspects, examining MapCam and OVIRS in concert can result in 167 substantially more robust scientific interpretations. Below we highlight some major findings 168 made by analyzing data from both instruments in tandem.

169 The earliest resolved low-phase angle ($\sim 5^{\circ}$) MapCam images of Bennu revealed that Roc 170 Saxum - the largest and darkest exposed boulder on Bennu's surface - had a shallow absorption 171 feature in the v-band (0.55 µm), consistent with the iron-oxide magnetite (Lauretta et al. 2019a). 172 However, the low spectral resolution of the MapCam colors rendered this interpretation 173 ambiguous. In later MapCam images acquired at higher phase angles (~8–11°), this absorption 174 feature appeared more shallow, further complicating this interpretation. It was unclear if this 175 change was related to instrumental artifacts or known phase angle effects that can decrease 176 absorption feature depths (e.g., Takir et al. 2015). However, later OVIRS data confirmed the 177 presence of a broad feature centered near $0.55 \,\mu\text{m}$ in spectra that are redder than average; we 178 also found two minor lines at 0.50 µm and 0.59 µm (Simon et al. 2020a). Features in this region

179 are usually attributed to an iron transition band (Izawa et al. 2019) and are consistent with the 180 iron oxides magnetite, goethite, and some Fe-bearing phyllosilicates (Cloutis et al. 2011b; 181 Cloutis et al. 2011a; Sherman and Waite 1985). Of the minerals typically found in aqueously 182 altered carbonaceous meteorites, magnetite is the best spectral match for a 0.55-µm feature with 183 more minor features at 0.50 and 0.59 µm (Simon et al. 2020a). Collectively, the detection of 184 magnetite in MapCam color and OVIRS spectra indicates that Bennu's parent body underwent 185 extensive aqueous alteration. Examining MapCam data at finer spatial scales (~25 cm/pixel) 186 indicates that magnetite may be concentrated in dark boulders and freshly exposed surfaces 187 (DellaGiustina et al. 2020).

188 One of the more surprising discoveries at Bennu was the detection of meter-scale, bright 189 pyroxene boulders on the surface of the asteroid (DellaGiustina et al. 2021). These boulders 190 showed a downturn in the x-band (0.847 μ m), the longest wavelength MapCam filter. This 191 downturn is consistent with an absorption feature found in mafic minerals, such as pyroxene or 192 olivine. Since MapCam only captured one shoulder of this presumed absorption, we could make 193 no further inferences on the composition of these boulders. However, spectra collected by 194 OVIRS showed that these bright boulders contained pyroxene and not olivine, as indicated by a 195 second absorption near 2 µm (DellaGiustina et al. 2021). Although Bennu's blue slope dominated 196 the OVIRS data of these boulders (which occupied $\sim 1\%$ of the instrument spot size), a pyroxene 197 signature was detected when their spectra were divided by the global average spectrum. These 198 normalized spectra have clear absorption bands at 1 and 2 µm, consistent with calcium-poor 199 pyroxenes. Band centers of the pyroxene absorption bands closely match those in the howardite-200 eucrite-diogenite meteorites from Vesta and resulted in the conclusion that pyroxene-bearing 201 boulders on Bennu are exogenous (DellaGiustina et al. 2021). This finding has been applied to 202 higher-resolution MapCam data to track the overall distribution of exogenous material on 203 Bennu's surface (Le Corre et al. 2021; Tatsumi et al. 2021).

Though these studies have examined OVIRS and MapCam data in tandem, the comparisons have mainly been qualitative. In this paper, we summarize the datasets collected by the two instruments and outline recommendations for more accurate, potentially more sensitive comparisons and assessments of uncertainty. Future VIS-NIR studies of Bennu's mineralogy 208 should use data from both instruments to provide a unified description of the surface at both high 209 spatial and high spectral resolution, following the recommendations we present.

210

1.4 OSIRIS-REX OBSERVATION CAMPAIGNS

211 The OSIRIS-REx mission carried out a series of global and regional imaging campaigns to 212 characterize the surface and potential sample collection sites (Lauretta et al. 2021). OVIRS 213 acquired data in almost every observation campaign; MapCam acquired images in the subset 214 dedicated to color imaging. Table 1 lists the observations used in this work, which are described 215 in detail below.

216 The Detailed Survey global imaging mission phase was comprised of the Baseball Diamond 217 and Equatorial Stations campaigns (Lauretta et al. 2021). The Equatorial Stations (EQ) campaign 218 was designed to acquire spectrometer and MapCam data at a series of stations with phase angles 219 ranging from 7° to 130° (Golish et al. 2021a; Lauretta et al. 2021). MapCam and OVIRS 220 acquired all EQ data from the equatorial plane of the asteroid, with a range to surface of 221 approximately 5 km. The spacecraft slewed north/south for at least a full Bennu rotation. For two 222 of the high-phase-angle stations — 6 am and 3:20 am (90° and 130° phase, respectively) — the 223 instruments observed for an additional quarter Bennu turn with the spacecraft pointed toward the 224 lit side of the asteroid. OVIRS acquired data in an identical way during all spacecraft slews. 225 MapCam alternated filters every slew, rotating through the full set (pan, b', v, w, x), such that 226 every fifth slew was imaged with the same filter.

227 In the Baseball Diamond campaign, OVIRS and MapCam were used concurrently in Flybys 228 2a (FB2a) and 2b (FB2b). These flybys were designed to acquire MapCam data for color maps 229 of Bennu (DellaGiustina et al. 2020) with a range to surface of ~3.6 km. FB2b is a re-fly of 230 FB2a, which had large pointing offset to the south caused by a missed spacecraft ephemeris 231 update (Lauretta et al. 2021). Both flybys utilized a point-and-stare observation pattern where 232 MapCam's pointing was held fixed for all five filters. For FB2b, MapCam acquired images with 233 southern, equatorial, and northern pointings. FB2a had only two pointings and, owing to the 234 missed ephemeris update, the nominally southern and northern looks were pointed off-body and 235 at Bennu's southern hemisphere, respectively. OVIRS acquired data during the point-and-stares,

236 during the transition between pointings, and from the end of one slew (northern look) to the start

- of the next (southern look).
- 238
- 239 Table 1: OVIRS and MapCam observations used in this comparison.

	Date of observation	Average phase angle (°)	Local solar time	Range to surface (km)	Surface coverage
Baseball Diamond					
FB2a	2019 Mar 14	8	12:30 pm	3.6	Global
FB2b	2019 Sep 26	8	12:30 pm	3.6	Global
Equatorial Stations					
EQ1	2019 Apr 25	45	3 pm	5	Global
EQ2	2019 May 02	130	3:20 am	5	Global
EQ3	2019 May 09	8	12:30 pm	5	Global
EQ4	2019 May 16	30	10 am	5	Global
EQ5	2019 May 23	90	6 am	5	Global
EQ6	2019 May 30	130	8:40 pm	5	Global
EQ7	2019 Jun 06	90	6 pm	5	Global
Reconnaissance A					
Sandpiper	2019 Oct 05	35	12:30 pm	1	Regional
Osprey	2019 Oct 12	40	1 pm	0.9	Regional
Kingfisher	2019 Oct 19	40	1:30 pm	1	Regional
Nightingale	2019 Oct 26	30	11:30 am	1	Regional
Reconnaissance B					
Nightingale	2020 Jan 21	65	4 pm	0.65	Regional
Osprey	2020 Feb 11	15	7:30 am	0.7	Regional

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241 After four potential sample sites were selected in the summer of 2019 (Sandpiper, Osprey,

242 Kingfisher, and Nightingale), OSIRIS-REx carried out a series of reconnaissance flybys that

imaged the surface at closer ranges (Lauretta et al. 2021). These flybys are referred to as Recon

A (~1 km), Recon B (~0.62 km), and Recon C (~0.25 km). Both instruments observed the four

candidate sample sites in Recon A; only the final two candidate sample sites (Nightingale and
Osprey) were observed in Recon B and Recon C. MapCam acquired images in the Recon A and
Recon B campaigns between large PolyCam mosaics, whereas OVIRS acquired data throughout
the flyby. As a result, similarly to Baseball Diamond, OVIRS acquired data concurrent with and
between groups of MapCam images. Unlike Baseball Diamond, the MapCam data were minimal
(sometimes limited to a single set of 10 color images), which limited the time ranges over which
comparable OVIRS data were acquired.

252 2 Cross-instrument Comparison Approach

253 2.1 COMPARISON PHILOSOPHY

Instrument and observation conditions affected the quality of the acquired data. Both instruments' calibration pipelines mitigated these effects, but some residual errors were unavoidable without hand-tuned adjustment of individual spectra and images. As such, we approached the comparison of the instruments on a per-dataset basis. That is, we analyzed the relative calibration of the instruments for each set of instrumental and observational conditions independently (e.g., a single Equatorial Station or a single Reconnaissance flyby).

260 Both instruments have independent absolute radiometric calibrations with moderate 261 uncertainties (Sections 1.1 and 1.2; Golish et al. 2020; Simon et al. 2018; Simon et al. 2021). The 262 data archived in the Planetary Data System (PDS; Reuter et al. 2019; Rizk et al. 2019) have been 263 calibrated by these published methods, therefore we find it most appropriate to compare the 264 archived calibrated data, rather than attempt to implement an absolute correction. To the notable 265 extent that the absolute radiometric calibrations were different, we did not attempt to determine 266 which instrument was more correct. We established the difference in a rigorous way, and across 267 multiple datasets, to provide future users of these data with context and uncertainties for their 268 analyses.

We performed this analysis using the SPICE kernels (Acton et al. 2018) produced by the OSIRIS-REx navigation team and archived with the Navigation and Ancillary Information

271 Facility (NAIF). Though multiple other analyses, particularly for OCAMS data, have updated the

pointing and/or position of MapCam during an observation (e.g., DellaGiustina et al. 2020;

273 Golish et al. 2021b), those updates do not necessarily apply to OVIRS. Registration of the data 274 with Bennu's shape model has no impact on our results, as the comparison is between 275 instruments. The only impact such alignment had was for creating maps of the comparison 276 (Section 4.2), but that impact is less than an OVIRS footprint. Moreover, future users of these 277 data are most likely to characterize them with the kernels available from NAIF. Therefore, it is 278 most broadly applicable to compare the data using the publicly accessible kernels. Nonetheless, 279 using the NAIF kernels for both instruments obfuscated some geometric offset between the two. 280 The SPICE frame and instrument kernels that define the boresights of the instruments were 281 designed in ground testing and updated after launch, but have some residual error. We estimate 282 that the pointing offset between the two instruments was less than an OVIRS footprint in the 283 global imaging campaigns, but likely introduced some error into this analysis (Section 4.3). 284 For a given OVIRS spectrum, we used the MapCam image acquired closest in time for 285 comparison. This minimized the photometric variation that occurs between data acquired at 286 different times, owing either to a change in spacecraft position or to Bennu rotation. OVIRS 287 observations typically started before, and ended after, MapCam imaging. To avoid unbound 288 photometric changes, we limited the OVIRS spectra to those taken between the first and last 289 MapCam images acquired. Even with this constraint, some photometric variation was 290 unavoidable between the OVIRS and MapCam data. In the Equatorial Stations data, MapCam 291 switched filters every slew, repeating every five slews. This results in a *slew aliasing* effect, 292 wherein a given OVIRS spectrum was between zero and two spacecraft slews away from the 293 closest MapCam image with a given filter. The spacecraft completed a slew every 2.7-3° of 294 Bennu rotation, such that the alignment between OVIRS and MapCam data varies between 0 and 295 6° of Bennu rotation. This had minimal impact at low phase angles (e.g., EQ3), but increasingly 296 large impact at higher phase angles, making these data less reliable.

For the Baseball Diamond flybys, OVIRS data acquired during MapCam's point-and-stare are temporally aligned, but OVIRS data taken in between MapCam imaging sets and between slews have an offset due to the time gap.

MapCam acquired only sparse data during the Reconnaissance phases, typically only taking one set of images. Thus, we expanded the time window for the Reconnaissance data to include a full scan of the site with OVIRS before and after MapCam imaging. This relaxation increased the amount of data available, but also increased the photometric variation between the data fromeach instrument significantly.

305 We further determined data validity by a number of observational factors. OVIRS spectra 306 that were acquired above 50° N/S latitude were excluded, because the high emission angles 307 cause increased uncertainty in the OVIRS radiometric calibration. We removed this limitation 308 for data acquired at the Nightingale site (which is at ~56°N) in the Reconnaissance phases. We 309 excluded OVIRS spectra with segment discontinuities greater than 2% (Section 1.2, 4.1). We 310 excluded panchromatic MapCam images acquired with very short exposures times in EQ3, 311 FB2a, and FB2b (which have icicle artifacts), as well as off-body or calibration MapCam 312 images. Pixels within a MapCam image that were outside the detector's linear regime (Golish et 313 al. 2020) were also excluded.

314 2.2 OVIRS SPATIAL FOOTPRINT

315 For a single OVIRS spectrum, we identified the five images, one for each MapCam filter, 316 acquired closest in time. OSIRIS-REx typically acquired data while the spacecraft slewed and 317 always while Bennu was rotating. For MapCam, the exposure times are short enough that motion 318 blur is << 1 MapCam pixel. OVIRS's exposure times, however, typically smeared the OVIRS 319 observation by $\sim 1/2$ of an OVIRS footprint. To account for the changing surface, we calculated 320 the location of the footprint throughout the observation (Figure 1). To start, we calculated the 321 Bennu latitude and longitude intersected by the OVIRS boresight at the start of the observation, 322 using SPICE kernels and a global shape model (Barnouin et al. 2020; Daly et al. 2020). We 323 calculated latitude and longitude backplanes for every MapCam image and found the pixel in the 324 nearest MapCam image that corresponded to the latitude and longitude of the OVIRS footprint. 325 OVIRS's field of view is 4 mrad; MapCam's instantaneous field of view (iFOV; the angle 326 subtended by a single pixel) is 0.067 µrad. Therefore, the OVIRS footprint encompassed pixels 327 within a 59-pixel diameter of the center point. We then translated the OVIRS footprint from the 328 start to the end of that OVIRS observation. At 100 points along the track, we repeated the 329 footprint calculation, building a weighted OVIRS mask (Figure 1(c)). The center of the track was 330 more heavily weighted because OVIRS observed it throughout the integration, whereas it 331 observed the edges only at the beginning or end. We applied this mask to the MapCam image to

- 332 calculate a weighted average of MapCam pixels corresponding to this OVIRS spectrum. We then
- took the mean of those pixels, because a mean represents OVIRS's physical averaging of
- 334 photons from multiple surface locations. We repeated this for each MapCam filter to produce a
- 335 five-point MapCam spectrum corresponding to the OVIRS spectrum.
- 336



Figure 1: An OVIRS spatial footprint on a MapCam image acquired at 17:53:34 on 2019 May 16. We compared OVIRS spectra to the closest (in time) MapCam image (a). The surface locations observed by OVIRS were identified with a weighted mask (b,c). Taking the mean of MapCam pixels weighted by the mask (d; location indicated by the blue rectangle in a), for each of MapCam's filters, produces an equivalent MapCam measurement.

- 342 2.3 MAPCAM SPECTRAL FOOTPRINT
- 343 Similarly, we extracted a five-point OVIRS spectrum from an OVIRS observation by
- 344 imparting a MapCam spectral footprint. MapCam's spectral responsivity was characterized in
- 345 extensive ground testing and documented in Golish et al. (2020). The per-filter spectral
- 346 responsivity included filter transmission, optics throughput, and detector sensitivity. We
- 347 multiplied an example OVIRS spectrum (Figure 2) by the normalized MapCam responsivities to
- 348 produce a five-point OVIRS spectrum.



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Figure 2: MapCam spectral footprints, from ground-based responsivity testing, plotted with an example OVIRS spectrum
 acquired at 17:53:46 on 2019 May 16.

353 2.4 SPECTRAL COMPARISON

354 For each OVIRS spectrum, we calculated the ratio of the five-point spectra in both absolute 355 and relative terms. The absolute five-point ratio (Figure 3(a)) gives the absolute radiometric 356 offset between the two instruments for each of MapCam's filters. The relative five-point ratio 357 (Figure 3(b)), which we normalized to the v filter (0.55 μ m), expresses the filter-to-filter offset 358 of the two instruments. We are primarily interested in how the four narrowband MapCam filters 359 compare with OVIRS. Though the pan filter also overlaps OVIRS wavelengths, it is not as useful 360 for spectral comparison because of the width of the filter. Nonetheless, we included it in the 361 analysis for completeness.

362 We repeated this comparison for every valid OVIRS spectrum. We depict the mean and variation of each filter by plotting the relative ratios on a scatter plot – with small, random 363 364 perturbations in wavelength for visualization (Figure 4(a)). Here the spectra are normalized to 365 the v filter; therefore, all v-filter data have a mean of exactly 1 with no variation. We also plotted 366 the reduced I/F from both instruments (Figure 4(b); v filter). If the instruments were perfectly 367 calibrated, the data would fall on the 1:1 dashed line. To the extent that their absolute radiometric 368 calibration differs (Sections 1.1 and 1.2; Figure 3(a)), the data would fall along a line with a 369 different slope. Because the data were noisy, they populate a scatter envelope around the line. 14

370 These results, for each OVIRS spectrum, MapCam filter, and OSIRIS-REx observation

371 campaign, were compiled to produce a per-filter, per-dataset comparison of the two instruments.



374 375 Figure 3: Five-point spectral ratios for the example shown in Sections 2.2/2.3, in absolute (a) and relative (b) terms, compares the response of each instrument at the same location ($\sim 23S$, 272E) on Bennu's surface.



377

378 Figure 4: Five-point spectra normalized to MapCam's v filter (0.55 µm) illustrate the mean and variation of the offset between 379 380 the instruments (a). Individual points are colored arbitrarily and randomly spread over 50 µm, around the filter's center wavelength, to help distinguish individual points among the cluster. Plotting the measured I/Fs against each other (b) further 381 382 illustrates the comparison, where the dashed line has a slope of 0.82 (equivalent to the absolute radiometric offset between the



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384 3 Cross-instrument Comparison Results

385 **3.1 GLOBAL OBSERVATION CAMPAIGNS**

386 The global observation campaigns provided the best opportunity to compare the instruments, 387 particularly those acquired at low phase angle. The FB2a, EO3, and FB2b datasets were all 388 acquired at 12:30 pm local time (~8° phase angle). These low-phase-angle observations had 389 minimal shadows, which are otherwise prevalent on Bennu's rough surface. Following the 390 procedure described above, we calculated the per-filter median and standard deviation of all five-391 point absolute and relative spectral ratios (Figure 5). MapCam's pan filter had a $\sim 12\%$ absolute 392 radiometric offset from OVIRS in these data, despite the fact that the pan filter essentially 393 encompasses MapCam's v and w filters. The difference was a result of the MapCam radiometric 394 calibration (Golish et al. 2020), which noted a higher response by the pan filter than the other 395 filters. The narrowband filters all had absolute offsets between 17 and 18%, suggesting a large 396 discrepancy between the MapCam and OVIRS (Table 2). However, when normalizing to the v 397 filter at 550 nm (i.e., removing the absolute offset), the four narrowband filters compared very 398 well to OVIRS (<1% residual offset; Table 3). This suggests that any comparative analysis that 399 uses color ratios (OVIRS to MapCam or MapCam filter to filter) would have radiometric 400 uncertainty of <1%.

401 Data from FB2a and FB2b had similar absolute and relative offsets (Figure 5(a,b,e,f), Table 402 2), though with slightly higher standard deviations (represented as error bars). As noted in 403 Section 1.4, the Equatorial Stations and Baseball Diamond flybys had different types of slew 404 aliasing, which likely introduced photometric differences between the observations that 405 increased the noise. Moreover, OVIRS did not have complete surface coverage in Baseball 406 Diamond (see maps in Section 4.2) owing to the MapCam-driven observation strategy. FB2a's 407 absolute offset was slightly less (~14%), despite having nearly identical imaging geometry to 408 FB2b. However, OVIRS data in FB2a were slightly saturated over some brighter portions of the 409 surface. This would result in depressing the reflectance that OVIRS measured, thereby 410 decreasing the offset with respect to OCAMS. Nonetheless, the relative offset remained <1% for 411 all three low-phase global datasets (Table 3).



Figure 5: Median five-point absolute (left; Table 2) and relative (right; Table 3) spectral ratios for FB2a, EQ3, and FB2b data (12:30 pm LST, ~8° phase angle).

	Median absolute MapCam/OVIRS ratio (± 1σ)					
	b'	V	W	Х	pan	
EQ3 (12:30 pm)	0.826 ± 0.016	0.830 ± 0.016	0.832 ± 0.016	0.828 ± 0.015	0.874 ± 0.021	
FB2a (12:30 pm)	0.855 ± 0.032	0.857 ± 0.031	0.859 ± 0.030	0.856 ± 0.029	0.882 ± 0.036	
FB2b (12:30 pm)	0.823 ± 0.023	0.831 ± 0.023	0.835 ± 0.023	0.833 ± 0.022	0.869 ± 0.024	
EQ4 (10 am)	0.801 ± 0.046	0.818 ± 0.045	0.829 ± 0.043	0.837 ± 0.041	0.885 ± 0.043	
EQ1 (3 pm)	0.854 ± 0.082	0.828 ± 0.083	0.796 ± 0.083	0.765 ± 0.085	0.766 ± 0.094	
EQ5 (6 am)	0.766 ± 0.111	0.794 ± 0.115	0.810 ± 0.119	0.830 ± 0.129	0.876 ± 0.151	
EQ2 (3:20 am)	0.648 ± 0.481	0.754 ± 0.522	0.852 ± 0.499	0.905 ± 0.496	0.990 ± 0.594	

420 Table 2: Absolute I/F ratios from cross-instrument global comparisons

422 Table 3: Normalized I/F ratios from cross-instrument global comparisons

	Median v-normalized MapCam/OVIRS ratio (± 1σ)					
	b'	V	W	Х	pan	
EQ3 (12:30 pm)	0.995 ± 0.003	1	1.002 ± 0.005	0.997 ± 0.005	1.052 ± 0.015	
FB2a (12:30 pm)	0.999 ± 0.005	1	1.003 ± 0.004	0.999 ± 0.005	1.033 ± 0.029	
FB2b (12:30 pm)	0.991 ± 0.002	1	1.004 ± 0.003	1.003 ± 0.005	1.046 ± 0.007	
EQ4 (10 am)	0.979 ± 0.007	1	1.014 ± 0.010	1.024 ± 0.014	1.083 ± 0.018	
EQ1 (3 pm)	1.033 ± 0.014	1	0.959 ± 0.017	0.921 ± 0.027	0.920 ± 0.042	
EQ5 (6 am)	0.962 ± 0.032	1	1.022 ± 0.036	1.050 ± 0.066	1.113 ± 0.106	
EQ2 (3:20 am)	0.799 ± 0.195	1	1.221 ± 0.293	1.390 ± 0.623	1.653 ± 1.080	

423

424 As phase angle increased in the other Equatorial Stations, so did the shadows on the surface, 425 which in turn increased the offset and noise between the two instruments. The data collected 426 during EQ4 (10 am, ~30° phase), for example, still compared well (~2% variation between instruments in the v-normalized spectrum; Table 3), with a slightly higher absolute offset (~20%; 427 428 Table 2). However, the higher-phase stations became increasingly variable. Figure 6 plots the 429 absolute and relative ratios for each of the global datasets on the same axes. If the instruments 430 were perfectly calibrated, with respect to each other, these ratio spectra would be horizontal lines 431 with a ratio value of 1. Deviations from a value of 1 indicate a calibration offset between OVIRS

- 432 and MapCam for that dataset. These ratio spectra highlight that the low-phase-angle data have
- 433 smaller offsets and compare better with each other than with higher-phase-angle data.



Figure 6: Absolute (a,c) and relative (b,d) ratios comparing OVIRS with MapCam's narrowband filters for the global datasets.
Figures (c) and (d) plot the same data as (a) and (b), respectively, with a cropped y-axis to better visualize the comparison of the best datasets.

438 OSIRIS-REx was pointed toward the terminator and the night side of Bennu during the 6 pm 439 and 8:40 pm equatorial stations, respectively. As a result, both the OCAMS and OVIRS five-440 point spectra measure primarily noise and are not included here. The same was true for most of 441 the 6 am and 3:20 am equatorial stations, but OSIRIS-REx acquired data for a quarter-Bennu-442 turn with the spacecraft pointed toward the lit side of the asteroid. The data from just the quarter-443 turn were included here, but were quite noisy, leading to larger offsets, particularly for 3:20 am. 444 The 3 pm station had not only larger offsets, but also a different spectral trend than the other 445 stations. Although at these phase angles it is difficult to assign a cause definitively, the direction 446 of shadows likely played a role. Because we excluded the 6 pm and 8:40 pm stations, the 3 pm 447 station was the only one analyzed here with eastward shadows (the other stations were in the 448 morning or close to noon). Shifting the shadows may change the instruments' relative response 449 to the surface, considering any pointing offset between them and the OVIRS segment read-out

order mentioned in Section 1.2. We tabulate the median and standard deviation, per dataset, of
the absolute MapCam/OVIRS ratios in Table 2 and of the relative spectra in Table 3.

452

2 3.2 REGIONAL OBSERVATION CAMPAIGNS

453 Comparing the instruments during the OSIRIS-REx Reconnaissance campaigns was more 454 challenging due to the closer range to the surface, which increased the OVIRS detector 455 temperature, noise, and radiometric uncertainty. This environment directly affected the absolute 456 radiometric ratio of the two instruments, but was less impactful on the relative ratio. Increased 457 OVIRS detector temperature decreases the long-wavelength sensitivity (Simon et al. 2018; 458 Simon et al. 2021), which is outside MapCam's spectral coverage. Higher temperatures can 459 affect the correction of out-of-band leaks at short wavelengths (Simon et al. 2021), but we 460 mitigated this by excluding spectra with large discontinuities (Section 2.1, 4.1).

461 The closer range also amplified the effect of the instruments' angular pointing inaccuracies. 462 These inaccuracies were much less than an OVIRS footprint when the spacecraft was 3.6-5 km 463 from the surface, but the inaccuracies increased linearly with decreased distance. At ranges of ~ 1 464 km (Recon A) and ~0.62 km (Recon B), the pointing offset was a significant fraction of an OVIRS 465 footprint. This caused increased differences between individual OVIRS spectra and their 466 corresponding OCAMS footprint. On the other hand, the regional nature of the data decreased the 467 scatter induced by varying albedo on Bennu. As a result, the relative radiometric ratios for the 468 regional datasets were only 2–3% (Figure 7), but had standard deviations several times larger 469 (Table 4). Again, for perfectly calibrated instruments, these median ratios would be 1 at all 470 wavelengths. However, these regional datasets emphasize that both the median ratio, and standard 471 deviation around that median, are needed to represent the fidelity of the cross-instrument 472 comparison.



474 Figure 7: Absolute (a,c) and relative (b) ratios comparing OVIRS with MapCam's narrowband filters for each of the regional
475 datasets (Recon A and B, abbreviated RA and RB), shown with the EQ3 dataset for comparison. Figure (c) plots the same data as
476 (a), with a cropped y-axis to better visualize the comparison of the best datasets.

478 Table 4: Absolute I/F ratios from cross-instrument Recon A (RA) and Recon B (RB) comparisons.

	Median absolute MapCam/OVIRS ratio (± 1σ)					
	b'	V	W	Х	pan	
Sandpiper (RA)	0.849 ± 0.070	0.858 ± 0.071	0.862 ± 0.069	0.862 ± 0.069	0.882 ± 0.071	
Osprey (RA)	0.863 ± 0.121	0.872 ± 0.124	0.876 ± 0.120	0.875 ± 0.121	0.892 ± 0.126	
Kingfisher (RA)	0.853 ± 0.074	0.861 ± 0.075	0.866 ± 0.073	0.865 ± 0.072	0.883 ± 0.077	
Nightingale (RA)	0.863 ± 0.297	0.874 ± 0.296	0.880 ± 0.291	0.882 ± 0.287	0.901 ± 0.302	
Nightingale (RB)	0.642 ± 0.594	0.646 ± 0.593	0.639 ± 0.582	0.628 ± 0.551	0.657 ± 0.584	
Osprey (RB)	0.963 ± 0.269	0.984 ± 0.276	0.988 ± 0.276	0.992 ± 0.283	1.054 ± 0.307	

-

	Median v-normalized MapCam/OVIRS ratio (± 1σ)				
	b'	V	W	Х	pan
Sandpiper (RA)	0.989 ± 0.003	1	1.004 ± 0.006	1.005 ± 0.008	1.028 ± 0.010
Osprey (RA)	0.990 ± 0.005	1	1.004 ± 0.009	1.004 ± 0.012	1.024 ± 0.013
Kingfisher (RA)	0.991 ± 0.004	1	1.005 ± 0.009	1.005 ± 0.014	1.025 ± 0.016
Nightingale (RA)	0.987 ± 0.004	1	1.008 ± 0.009	1.011 ± 0.011	1.031 ± 0.008
Nightingale (RB)	0.988 ± 0.027	1	1.002 ± 0.039	1.005 ± 0.064	1.026 ± 0.031
Osprey (RB)	0.977 ± 0.019	1	1.004 ± 0.018	1.010 ± 0.031	1.071 ± 0.056

486 **4 Comparison of Individual Spectra**

487 The combined results from the previous sections demonstrate reasonably good agreement 488 between the two instruments when averaged over entire datasets from discrete observational 489 campaigns. However, the standard deviations attached to those averages (which are performed 490 over thousands of spectra) indicate significant spectrum-to-spectrum variation. In general, the 491 variations (both filter-to-filter and as represented by 1σ error bars) listed in the previous section 492 should be used as uncertainties for any cross-instrument comparison that uses individual spectra 493 (Table 2 and Table 4 for absolute comparisons, Table 3 and Table 5 for filter-relative 494 comparisons). The differences between the instruments discussed in Section 1 have a direct impact 495 on the comparison of individual spectra.

496

4.1 SEGMENT DISCONTINUITIES

497 As discussed in Section 1.2, OVIRS has a segment boundary at approximately the short-498 wavelength end of the OCAMS w filter. Because the two segments that compose this boundary 499 are on opposite sides of the OVIRS detector, and read out at slightly different times, they imaged 500 slightly different portions of the surface as the spacecraft was slewing during an OVIRS 501 integration. If these portions of the surface were not spatially uniform, the two segments could 502 have measured signals that were different in proportion to that heterogeneity. Figure 8 depicts an 503 example of this where OVIRS observed a large shadow in two subsequent integrations. In the

504 first spectrum (Figure 8(a-c)), the shadow strongly influenced segment 1a and suppressed the 505 signal below $\sim 0.68 \,\mu\text{m}$. In the second spectrum (Figure 8(d-f)), the spacecraft had slewed such 506 that segment 1b was most affected by the shadow, suppressing the longer wavelengths. In the 507 analyses described above, we rejected any OVIRS spectrum with a segment discontinuity larger 508 than 2%. Continuity was calculated by taking the median of the spectra over the wavelengths 509 0.040 μ m before and after the boundary (i.e., 0.64 – 0.68 μ m and 0.68 – 0.72 μ m). However, 510 such discontinuities can influence individual spectra for analyses of specific surface features that 511 were much brighter or darker than their surroundings.





Figure 8:Large shadows on the surface observed by MapCam (a,d; 17:53:42 on 2019 May 16) – such as one located at ~43°S,
272°E (b,e; location indicated by the blue rectangle in a,d) – are more susceptible to segment discontinuities in OVIRS spectra
(c,f, dashed black lines; 17:53:38 and 17:53:40 on 2019 May 16), due to the 1a and 1b segments imaging portions of the surface
with different brightness. MapCam filters (c,f, color lines) are sensitive to wavelengths on either side of the discontinuity.

519 4.2 MAPS

520 To better visualize the spatial distribution of differences between the two instruments, we 521 produced maps of the OVIRS and MapCam comparisons. We constructed these maps by 522 averaging OVIRS footprints into latitude/longitude bins. As such, the maps are at OVIRS's 523 approximate spatial resolution. Though this approach sacrificed MapCam's much finer spatial 524 resolution, it maintained a 1:1 spatial match between the two datasets (as opposed to comparing 525 the OVIRS map with a native resolution MapCam mosaic such as those in DellaGiustina et al. 526 (2020)). These maps facilitate the comparisons of various albedo or spectral parameter maps 527 derived from the two instruments (e.g., DellaGiustina et al. 2020; Fornasier et al. 2020; Golish et 528 al. 2021b; Kaplan et al. 2020; Li et al. 2021; Simon et al. 2020; Zou et al. 2021). 529 Figure 9 shows OVIRS, MapCam, and ratio maps for the EQ3 (12:30 pm) dataset. The 530 Bennu albedo map (Golish et al. 2021c) is also shown for reference; it is not used in this 531 analysis. The albedo map values are not directly comparable because the albedo map has been 532 photometrically corrected and the EQ3 I/F maps have not, but there are qualitative spatial 533 correlations between the maps. As shown in the previous section (Figure 6 and Table 2), OVIRS 534 measured the mean I/F as ~17.5% larger than MapCam did. As expected, many of the regions 535 that deviated from this mean (higher or lower) corresponded to large features on Bennu's surface that cast shadows, even at low phase angles. These features often have a 'bright' side and a 536 537 'dark' side in the ratio map, presumably due to slightly different photometric conditions between 538 MapCam and OVIRS as the spacecraft slewed over the features.

539 The remaining structure is not random and is likely driven by the photometric variation 540 induced by slew aliasing between OVIRS and MapCam data, coupled with Bennu's terrain. 541 Regardless of its source, this structure will interfere with any individual spectrum comparison. 542 Though the v-normalized spectra, on average, agree within 1% between the two instruments, and 543 with a standard deviation < 1% (Figure 5), comparing individual spectra can have differences as 544 high as 10% around large surface features.



Figure 9: Comparisons of OVIRS (a) and MapCam (b) EQ3 I/F maps illustrate the absolute offset between the instruments. A ratio of the two I/F maps shows the terrain- and slew-dependent noise in the comparison. The Bennu normal albedo map (d; Golish et al. 2021c) is included for visual reference, it was not used in the analysis.

549 At higher phase angles (and therefore larger shadows), differences induced by terrain and 550 slew aliasing become more pronounced. Even at 10 am (Figure 10(a)), vertical artifacts resulting 551 from slew aliasing become qualitatively obvious. At 3 pm (Figure 10(b)), as discussed in Section 552 3.1, shadows were larger and in the opposite direction. At 6 am and 3:20 am (Figure 10(c,d)), 553 only the lit quarter turn provided usable data, which covered a small portion of the surface and 554 did so with large shadows and resulting noise. 555 The Baseball Diamond data (FB2a and FB2b) compared well between the two instruments, 556 as we would expect for low phase angles. However, the ratio maps (Figure 10(e,f)) illustrate the 557 sparse OVIRS coverage during these MapCam-focused observations. As such, the EQ3 data are

558 generally preferred.



Figure 10: Maps of MapCam/OVIRS ratios where their coverage overlaps in the global campaigns, when compared with EQ3 at 12:30 pm (Figure 9), illustrate decreasing utility with increasing phase angle. The color scale represents the MapCam/OVIRS ratio. Data with ratios around the instruments' radiometric offset (~0.82) compare well between the instruments; data far from that offset indicate poor cross-comparison. The FB2a and 2b maps illustrate the sparse OVIRS coverage during the MapCam-focused flybys.

The Recon A data had much more variation in ratios, as we would expect for the mid-phase angles and closer range to surface. As shown in Table 4, though the median MapCam/OVIRS ratio is similar to the low phase global campaigns, the standard deviation and the variation in the ratio maps (Figure 11(a-d)), are much higher. This emphasizes that the instruments were spectrally similar on average but have significant spectrum-to-spectrum variation. Any analysis that includes individual spectrum comparisons should acknowledge this variation. Finally, the maps for Recon 571 B Figure 11(e,f) illustrate the lack of utility of these data. The data that passed even our relaxed 572 validity constraints (Section 2.1) were noisy and did not cover the bulk of the sampling sites. We 573 include the Recon B results not as reliable statistics for future analyses, but as caution against using 574 them without further calibration and analysis.



575

Figure 11: Maps of overlapping MapCam/OVIRS ratios (indicated by the color scale) for the regional campaigns illustrate large variability in the Recon A data when comparing the instruments. Recon B data, which rarely met our data validity requirements, sparsely covered the site and were not reliable in an instrument-to-instrument comparison without further spectrum-specific calibration.

580 4.3 SPATIAL CO-REGISTRATION

581 As described in Section 2.1, we made no attempt in this analysis to align the OVIRS and

582 MapCam data. The pointing of the data from both instruments is described in the mission kernels

583 derived by the OSIRIS-REx navigation team and archived with NAIF. Nonetheless, the ratio

584 maps shown in the previous section depict reasonable spatial co-registration. We further evaluate

585 that registration by plotting OVIRS and MapCam spectra as the instruments slew over notable 586 surface features. Roc Saxum (~25°E, 25°S), in Bennu's southern hemisphere, is ~20% darker 587 than average Bennu and ~100 m long, making it the most prominent albedo feature on Bennu's 588 surface. Figure 12(a) plots the OVIRS observation track as it slewed over Roc Saxum six times 589 during the EQ3 (12:30 pm) station. The absolute I/F tracks (Figure 12(b)) are indicative of the 590 absolute radiometric offset between the instruments. However, when we normalized the spectral 591 tracks to Bennu's average I/F (as measured by each instrument and filter), they reveal that the 592 MapCam data undergo a deeper drop in Bennu-normalized I/F in the first slew (Figure 12(c)). 593 This slew was along the eastern edge of Roc Saxum, which was the shadowed edge because 594 these data were acquired slightly past noon (local solar time). We have seen throughout the 595 analysis that shadows were the biggest driver for differences between the instruments, which 596 seems to be confirmed here. In addition, because the first slew was along a relatively sharp 597 albedo transition, any east-west misregistration between the two instruments would manifest as a 598 difference here.

599 Figure 12(d) plots the v-normalized spectral ratios for the slews over Roc Saxum, showing 600 several deviations, particularly in the w and x filters. These deviations are most prominent when 601 the instruments slewed on and off Roc Saxum (i.e., coincident with a rapid change in albedo). As 602 described in Section 1.2, the w and x filters correspond to the OVIRS segment 1b, which was 603 imaging a slightly different part of the surface than segment 1a (MapCam filters b' and v). This 604 rolling shutter effect likely results in w- and x-filter deviations. The width of these deviations are 605 a few OVIRS integrations, giving a rough sense of the spatial offset (~ 10 m) in this dataset. 606 However, even these outlier spectrum-to-spectrum deviations are less than 2%, while most 607 deviations are less than 1%, indicating reliable comparison between the instruments.



609

Figure 12: Spectral tracks from OVIRS and MapCam as the instruments slewed over Roc Saxum (a). In the full observation, the data continued toward the northern part of Bennu before slewing back over Roc Saxum; we show only a subset of the slews here.
Plots of I/F ratios (b), Bennu-normalized I/F ratios (c), and v-normalized I/F ratios (d) track the response of the instruments throughout the slews. Vertical dotted lines indicate the beginning and end of the slews that imaged Roc Saxum.

614 4.4 SAMPLE SPECTRA

615 Despite the qualifications and uncertainties detailed throughout this analysis, meaningful comparative work can be and has been performed (DellaGiustina et al. 2021; Kaplan et al. 2020) 616 617 by cross-referencing data from the two instruments. Using two notable surface features — Roc 618 Saxum and another large boulder, Benben Saxum — Figure 13 plots the absolute and relative 619 spectra acquired by both instruments during EQ3. The relative spectrum was normalized to 620 Bennu's average spectrum (calculated using the EQ3 data). We selected spectra from the middle 621 of the boulders because data from the edges can lead to artifacts. The spectra from the two 622 instruments have the 17–18% absolute radiometric offset established in Section 3.1. However, 623 the Bennu-relative spectra agree within 1% and confirm the colors (DellaGiustina et al. 2020; 624 Simon et al. 2020b) of these features: Roc Saxum and Benben Saxum are redder and bluer, 625 respectively, relative to average Bennu. A filter-relative analysis using these data, or any EQ3 626 data, should carry the uncertainties identified in Table 3 ($\pm 1\%$).





Figure 13: Individual spectra of Roc (a) and Benben (d) Saxa indicate the validity and uncertainty of OVIRS and MapCam comparisons. The absolute I/F offset (b,e) is larger than the \sim 5% radiometric uncertainty predicted by both instruments (indicated by the error bars). However, when compared filter-to-filter (c,f), OVIRS and MapCam agree to <1%.

633 **5 Conclusions**

This work provides a complete summary of concurrent OVIRS and MapCam datasets
 acquired during OSIRIS-REx proximity operations and recommendations for how to most
 accurately compare them.

637 The instruments have a large absolute radiometric offset ($\sim 15-20\%$) that stems from 638 independent calibration processes with independent sources of error. However, the offset is 639 consistent among all four MapCam color filters for the low-phase-angle datasets. In low-phase-640 angle observations, when shadows and instrument effects have minimal impact on the data 641 quality, the OVIRS-to-MapCam and MapCam filter-to-filter relative calibration are very good 642 (<1% uncertainty). The EQ3 dataset (acquired at ~12:30 pm local solar time) provides the most 643 thorough surface coverage and highest-quality cross-instrument comparison, with a < 2%644 spectrum-to-spectrum 1σ absolute radiometric uncertainty.

We strongly recommend using this dataset whenever possible when comparing data from these two instruments, because higher-phase-angle data require larger uncertainties to be applied. Even with the EQ3 dataset, we advise some caution when analyzing individual OVIRS spectra, due to imperfect instrument co-registration and OVIRS segment discontinuities. Nonetheless, this cross-instrument comparison allows future analyses to apply realistic uncertainties to overlays, ratios, and other quantitative comparisons of OVIRS and MapCam data acquired at Bennu, and perhaps to identify subtler signals that have been previously discernable.

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661 **7 Data Availability**

662 The OVIRS (Reuter et al. 2019) and OCAMS (Rizk et al. 2019) data used in this analysis
663 are available at the Planetary Data System Small Bodies Node

664 (https://sbn.psi.edu/pds/resource/orex/). The results of this analysis — tabulated data from Tables

2 and 3 and raster images representing the ratio maps in Figures 9, 10, and 11 — are archived in

666 Golish et al. (2021d).

667 8 References

- 668 C. Acton, N. Bachman, B. Semenov, and E. Wright, P&SS 150, 9 (2018).
- 669 R.-L. Ballouz, K. J. Walsh, O. S. Barnouin, D. N. DellaGiustina, M. Al Asad, E. R. Jawin,
- 670 M. G. Daly, W. F. Bottke, P. Michel, C. Avdellidou, M. Delbo, R. T. Daly, E. Asphaug, C. A.
- 671 Bennett, E. B. Bierhaus, H. C. Connolly, D. R. Golish, J. L. Molaro, M. C. Nolan, M. Pajola, B.
- 672 Rizk, S. R. Schwartz, D. Trang, C. W. V. Wolner, and D. S. Lauretta, Nature 587, 205 (2020).
- O. S. Barnouin, M. G. Daly, E. E. Palmer, C. L. Johnson, R. W. Gaskell, M. Al Asad, E. B.
- 674 Bierhaus, K. L. Craft, C. M. Ernst, R. C. Espiritu, H. Nair, G. A. Neumann, L. Nguyen, M. C.
- Nolan, E. Mazarico, M. E. Perry, L. C. Philpott, J. H. Roberts, R. J. Steele, J. Seabrook, H. C. M.

676 Susorney, J. R. Weirich, and D. S. Lauretta, P&SS 180, 104764 (2020).

- 677 C. A. Bennett, D. N. DellaGiustina, K. J. Becker, T. L. Becker, K. L. Edmundson, D. R.
- 678 Golish, R. J. Bennett, K. N. Burke, C. N. U. Cue, B. E. Clark, J. Contreras, J. D. P. Deshapriya,
- 679 C. D. d'Aubigny, G. Fitzgibbon, E. R. Jawin, T. Q. Nolan, N. A. Porter, M. M. Riehl, H. L.
- 680 Roper, B. Rizk, Y. Tang, Z. Zeszut, R. W. Gaskell, E. E. Palmer, J. R. Weirich, M. M. Al Asad,
- L. Philpott, M. G. Daly, O. S. Barnouin, H. L. Enos, and D. S. Lauretta, Icarus **357**, 113690
- 682 (2021).
- B. J. Buratti, M. D. Hicks, J. Nettles, M. Staid, C. M. Pieters, J. Sunshine, J. Boardman, and
 T. C. Stone, JGR 116, E00G03 (2011).
- 685 B. E. Clark, R. P. Binzel, E. S. Howell, E. A. Cloutis, M. Ockert-Bell, P. Christensen, M. A.
- 686 Barucci, F. DeMeo, D. S. Lauretta, H. Connolly, A. Soderberg, C. Hergenrother, L. Lim, J.
- 687 Emery, and M. Mueller, Icarus **216**, 462 (2011).
- E. A. Cloutis, T. Hiroi, M. J. Gaffey, C. M. O. D. Alexander, and P. Mann, Icarus 212, 180
 - 32

689 (2011a).

- 690 E. A. Cloutis, P. Hudon, T. Hiroi, M. J. Gaffey, and P. Mann, Icarus 216, 309 (2011b).
- 691 L. Le Corre, V. Reddy, W. F. Bottke, D. N. DellaGiustina, K. N. Burke, J. Nolau, R. B. Van
- Auken, D. R. Golish, J. A. Sanchez, J. Y. Li, C. Y. Drouet d'Aubigny, B. Rizk, and D. S.
- 693 Lauretta, PSJ **2**, 114 (2021).
- M. G. Daly, O. S. Barnouin, J. A. Seabrook, J. Roberts, C. Dickinson, K. J. Walsh, E. R.
- Jawin, E. E. Palmer, R. Gaskell, J. Weirich, T. Haltigin, D. Gaudreau, C. Brunet, G.
- 696 Cunningham, P. Michel, Y. Zhang, R.-L. Ballouz, G. Neumann, M. E. Perry, L. Philpott, M. M.
- 697 Al Asad, C. L. Johnson, C. D. Adam, J. M. Leonard, J. L. Geeraert, K. Getzandanner, M. C.
- Nolan, R. T. Daly, E. B. Bierhaus, E. Mazarico, B. Rozitis, A. J. Ryan, D. N. Dellaguistina, B.
- 699 Rizk, H. C. M. Susorney, H. L. Enos, and D. S. Lauretta, SciA 6, eabd3649 (2020).
- D. N. DellaGiustina, C. A. Bennett, K. Becker, D. R. Golish, L. Le Corre, D. A. Cook, K.
- 701 L. Edmundson, M. Chojnacki, S. S. Sutton, M. P. Milazzo, B. Carcich, M. C. Nolan, N. Habib,
- 702 K. N. Burke, T. Becker, P. H. Smith, K. J. Walsh, K. Getzandanner, D. R. Wibben, J. M.
- 703 Leonard, M. M. Westermann, A. T. Polit, J. N. Kidd, C. W. Hergenrother, W. V. Boynton, J.
- 704 Backer, S. Sides, J. Mapel, K. Berry, H. Roper, C. Drouet d'Aubigny, B. Rizk, M. K. Crombie,
- E. K. Kinney-Spano, J. de León, J. L. Rizos, J. Licandro, H. C. Campins, B. E. Clark, H. L.
- 706 Enos, and D. S. Lauretta, E&SS 5, 929 (2018).
- 707 D. N. DellaGiustina, K. N. Burke, K. J. Walsh, P. H. Smith, D. R. Golish, E. B. Bierhaus,
- 708 R.-L. Ballouz, T. L. Becker, H. Campins, E. Tatsumi, K. Yumoto, S. Sugita, J. D. P. Deshapriya,
- E. A. Cloutis, B. E. Clark, A. R. Hendrix, A. Sen, M. M. Al Asad, M. G. Daly, D. M. Applin, C.
- 710 Avdellidou, M. A. Barucci, K. J. Becker, C. A. Bennett, W. F. Bottke, J. I. Brodbeck, H. C.
- 711 Connolly, M. Delbo, J. de Leon, C. Y. Drouet d'Aubigny, K. L. Edmundson, S. Fornasier, V. E.
- 712 Hamilton, P. H. Hasselmann, C. W. Hergenrother, E. S. Howell, E. R. Jawin, H. H. Kaplan, L.
- 713 Le Corre, L. F. Lim, J. Y. Li, P. Michel, J. L. Molaro, M. C. Nolan, J. Nolau, M. Pajola, A.
- 714 Parkinson, M. Popescu, N. A. Porter, B. Rizk, J. L. Rizos, A. J. Ryan, B. Rozitis, N. K. Shultz,
- 715 A. A. Simon, D. Trang, R. B. Van Auken, C. W. V. Wolner, and D. S. Lauretta, Science 370,
- 716 eabc3660 (2020).
- 717 D. N. DellaGiustina, H. H. Kaplan, A. A. Simon, W. F. Bottke, C. Avdellidou, M. Delbo,
- 718 R.-L. Ballouz, D. R. Golish, K. J. Walsh, M. Popescu, H. Campins, M. A. Barucci, G. Poggiali,

- 719 R. T. Daly, L. Le Corre, V. E. Hamilton, N. Porter, E. R. Jawin, T. J. McCoy, H. C. Connolly, J.
- L. R. Garcia, E. Tatsumi, J. de Leon, J. Licandro, S. Fornasier, M. G. Daly, M. M. Al Asad, L.
- 721 Philpott, J. Seabrook, O. S. Barnouin, B. E. Clark, M. C. Nolan, E. S. Howell, R. P. Binzel, B.
- 722 Rizk, D. C. Reuter, and D. S. Lauretta, NatAs 5, 31 (2021).
- 723 S. Fornasier, P. H. Hasselmann, J. D. P. Deshapriya, M. A. Barucci, B. E. Clark, A. Praet,
- V. E. Hamilton, A. Simon, J.-Y. Li, E. A. Cloutis, F. Merlin, X.-D. Zou, and D. S. Lauretta,
- 725 A&A **644**, A142 (2020).
- D. R. Golish, D. N. DellaGiustina, J.-Y. Li, B. E. Clark, X.-D. Zou, P. H. Smith, J. L.
- 727 Rizos, P. H. Hasselmann, C. A. Bennett, S. Fornasier, R.-L. Ballouz, C. Drouet d'Aubigny, B.
- 728 Rizk, M. G. Daly, O. S. Barnouin, L. Philpott, M. M. Al Asad, J. A. Seabrook, C. L. Johnson,
- 729 and D. S. Lauretta, Icarus **357**, 113724 (2021a).
- 730 D. R. Golish, C. Drouet d'Aubigny, B. Rizk, D. N. DellaGiustina, P. H. Smith, K. Becker,
- 731 N. Shultz, T. Stone, M. K. Barker, E. Mazarico, E. Tatsumi, R. W. Gaskell, L. Harrison, C.
- 732 Merrill, C. Fellows, B. Williams, S. O'Dougherty, M. Whiteley, J. Hancock, B. E. Clark, C. W.
- 733 Hergenrother, and D. S. Lauretta, SSRv **216**, 12 (2020).
- D. R. Golish, J. Y. Li, B. E. Clark, D. N. DellaGiustina, X. D. Zou, J. L. Rizos, P. H.
- 735 Hasselmann, C. A. Bennett, S. Fornasier, C. Drouet d'Aubigny, B. Rizk, M. G. Daly, O. S.
- 736 Barnouin, J. A. Seabrook, L. Philpott, M. M. Al Asad, C. L. Johnson, B. Rozitis, A. Ryan, J. P.
- 737 Emery, and D. S. Lauretta, Planetary Science Journal, (2021b).
- 738 D. R. Golish, N. K. Shultz, T. L. Becker, K. J. Becker, K. L. Edmundson, D. N.
- 739 DellaGiustina, C. Drouet d'Aubigny, C. A. Bennett, B. Rizk, O. S. Barnouin, M. G. Daly, J. A.
- 740 Seabrook, L. Philpott, M. M. Al Asad, C. L. Johnson, J.-Y. Li, R.-L. Ballouz, E. R. Jawin, and
- 741 D. S. Lauretta, Icarus **355**, 114133 (2021c).
- 742 D. R. Golish, A. A. Simon, D. C. Reuter, S. Ferrone, B. E. Clark, J.-Y. Li, D. N.
- 743 DellaGiustina, C. D. D'Aubigny, B. Rizk, and D. S. Lauretta, Figshare (2021d).
- 744 C. W. Hergenrother, C. D. Adam, S. R. Chesley, and D. S. Lauretta, JGRE 125, 1 (2020).
- 745 C. W. Hergenrother, M. C. Nolan, R. P. Binzel, E. A. Cloutis, M. A. Barucci, P. Michel, D.
- 746 J. Scheeres, C. D. D'Aubigny, D. Lazzaro, N. Pinilla-Alonso, H. Campins, J. Licandro, B. E.
- 747 Clark, B. Rizk, E. C. Beshore, and D. S. Lauretta, Icarus 226, 663 (2013).
- M. R. M. Izawa, E. A. Cloutis, T. Rhind, S. A. Mertzman, D. M. Applin, J. M. Stromberg,34

- 749 and D. M. Sherman, Icarus **319**, 525 (2019).
- 750 E. R. Jawin, K. J. Walsh, O. S. Barnouin, T. J. McCoy, R. -L. Ballouz, D. N. DellaGiustina,
- H. C. Connolly, J. Marshall, C. Beddingfield, M. C. Nolan, J. L. Molaro, C. A. Bennett, D. J.
- 752 Scheeres, M. G. Daly, M. Al Asad, R. T. Daly, E. B. Bierhaus, H. C. M. Susorney, H. H. Kaplan,
- 753 H. L. Enos, and D. S. Lauretta, JGRE **125**, e2020JE006475 (2020).
- H. H. Kaplan, D. S. Lauretta, A. A. Simon, V. E. Hamilton, D. N. DellaGiustina, D. R.
- 755 Golish, D. C. Reuter, C. A. Bennett, K. N. Burke, H. Campins, H. C. Connolly, J. P. Dworkin, J.
- P. Emery, D. P. Glavin, T. D. Glotch, R. Hanna, K. Ishimaru, E. R. Jawin, T. J. McCoy, N.
- 757 Porter, S. A. Sandford, S. Ferrone, B. E. Clark, J.-Y. Li, X.-D. Zou, M. G. Daly, O. S. Barnouin,
- 758 J. A. Seabrook, and H. L. Enos, Science **370**, eabc3557 (2020).
- D. S. Lauretta, S. S. Balram-Knutson, C. A. Bennett, B. J. Bos, C. D. D'Aubigny, P. R.
- 760 Christensen, E. C. A. Church, D. N. DellaGiustina, H. L. Enos, D. R. Golish, V. E. Hamilton, C.
- 761 W. Hergenrother, E. S. Howell, J. N. K. Jr., M. C. Nolan, D. C. Reuter, B. Rizk, A. A. Simon,
- and O.-Re. Team, Abstract 1240 Presented at 2018 Lunar and Planetary Science Conference,
- 763 The Woodlands, Texas, 19-23 Mar (2018).
- D. S. Lauretta, S. S. Balram-Knutson, E. Beshore, W. V. Boynton, C. Drouet d'Aubigny, D.
- N. DellaGiustina, H. L. Enos, D. R. Golish, C. W. Hergenrother, E. S. Howell, C. A. Bennett, E.
- 766 T. Morton, M. C. Nolan, B. Rizk, H. L. Roper, A. E. Bartels, B. J. Bos, J. P. Dworkin, D. E.
- 767 Highsmith, D. A. Lorenz, L. F. Lim, R. Mink, M. C. Moreau, J. A. Nuth, D. C. Reuter, A. A.
- 768 Simon, E. B. Bierhaus, B. H. Bryan, R. Ballouz, O. S. Barnouin, R. P. Binzel, W. F. Bottke, V.
- 769 E. Hamilton, K. J. Walsh, S. R. Chesley, P. R. Christensen, B. E. Clark, H. C. Connolly, M. K.
- 770 Crombie, M. G. Daly, J. P. Emery, T. J. McCoy, J. W. McMahon, D. J. Scheeres, S. Messenger,
- 771 K. Nakamura-Messenger, K. Righter, and S. A. Sandford, SSRv 212, 925 (2017).
- D. S. Lauretta, D. N. DellaGiustina, C. A. Bennett, D. R. Golish, K. J. Becker, S. S.
- 773 Balram-Knutson, O. S. Barnouin, T. L. Becker, W. F. Bottke, W. V. Boynton, H. Campins, B. E.
- 774 Clark, H. C. Connolly, C. Y. Drouet d'Aubigny, J. P. Dworkin, J. P. Emery, H. L. Enos, V. E.
- Hamilton, C. W. Hergenrother, E. S. Howell, M. R. M. Izawa, H. H. Kaplan, M. C. Nolan, B.
- 776 Rizk, H. L. Roper, D. J. Scheeres, P. H. Smith, K. J. Walsh, and C. W. V. Wolner, Nature 568,
- 777 55 (2019a).
- D. S. Lauretta, H. Enos, A. T. Polit, H. L. Roper, and C. W. V. Wolner, in *Sample Return*

- 779 *Missions*, edited by A. Longobardo (Elsevier, Amsterdam, Netherlands, 2021), pp. 163–194.
- 780 D. S. Lauretta, C. W. Hergenrother, S. R. Chesley, J. M. Leonard, J. Y. Pelgrift, C. D.
- Adam, M. Al Asad, P. G. Antreasian, R.-L. Ballouz, K. J. Becker, C. A. Bennett, B. J. Bos, W.
- 782 F. Bottke, M. Brozović, H. Campins, H. C. Connolly, M. G. Daly, A. B. Davis, J. de León, D. N.
- 783 DellaGiustina, C. Y. Drouet d'Aubigny, J. P. Dworkin, J. P. Emery, D. Farnocchia, D. P. Glavin,
- D. R. Golish, C. M. Hartzell, R. A. Jacobson, E. R. Jawin, P. Jenniskens, J. N. Kidd, E. J.
- 785 Lessac-Chenen, J.-Y. Li, G. Libourel, J. Licandro, A. J. Liounis, C. K. Maleszewski, C.
- 786 Manzoni, B. May, L. K. McCarthy, J. W. McMahon, P. Michel, J. L. Molaro, M. C. Moreau, D.
- 787 S. Nelson, W. M. Owen, B. Rizk, H. L. Roper, B. Rozitis, E. M. Sahr, D. J. Scheeres, J. A.
- 788 Seabrook, S. H. Selznick, Y. Takahashi, F. Thuillet, P. Tricarico, D. Vokrouhlický, and C. W. V.
- 789 Wolner, Science **366**, eaay3544 (2019b).
- 790 J.-Y. Li, X.-D. Zou, D. R. Golish, B. E. Clark, S. Ferrone, S. Fornasier, P. H. Hasselmann,
- 791 A. J. Ryan, B. Rozitis, J. P. Emery, M. A. Siegler, A. A. Simon, D. C. Reuter, V. E. Hamilton,
- 792 and D. S. Lauretta, PSJ (2021).
- J. L. Molaro, K. J. Walsh, E. R. Jawin, R. L. Ballouz, C. A. Bennett, D. N. DellaGiustina,
- D. R. Golish, C. Drouet d'Aubigny, B. Rizk, S. R. Schwartz, R. D. Hanna, S. J. Martel, M.
- Pajola, H. Campins, A. J. Ryan, W. F. Bottke, and D. S. Lauretta, NatCo 11, 2913 (2020).
- 796 D. C. Reuter, A. A. Simon, J. Hair, A. Lunsford, S. Manthripragada, V. Bly, B. Bos, C.
- 797 Brambora, E. Caldwell, G. Casto, Z. Dolch, P. Finneran, D. Jennings, M. Jhabvala, E. Matson,
- M. McLelland, W. Roher, T. Sullivan, E. Weigle, Y. Wen, D. Wilson, and D. S. Lauretta, SSRv
 214, 54 (2018).
- 800 D. C. Reuter, A. A. Simon, A. Lunsford, R. Cosentino, N. Gorius, and D. S. Lauretta,
- 801 NASA Planetary Data System (2019).
- 802 B. Rizk, C. Drouet d'Aubigny, D. Golish, C. Fellows, C. Merrill, P. Smith, M. S. Walker, J.
- 803 E. Hendershot, J. Hancock, S. H. Bailey, D. N. DellaGiustina, D. S. Lauretta, R. Tanner, M.
- 804 Williams, K. Harshman, M. Fitzgibbon, W. Verts, J. Chen, T. Connors, D. Hamara, A. Dowd, A.
- Lowman, M. Dubin, R. Burt, M. Whiteley, M. Watson, T. McMahon, M. Ward, D. Booher, M.
- 806 Read, B. Williams, M. Hunten, E. Little, T. Saltzman, D. Alfred, S. O'Dougherty, M. Walthall,
- 807 K. Kenagy, S. Peterson, B. Crowther, M. L. Perry, C. See, S. Selznick, C. Sauve, M. Beiser, W.
- 808 Black, R. N. Pfisterer, A. Lancaster, S. Oliver, C. Oquest, D. Crowley, C. Morgan, C. Castle, R.

- 809 Dominguez, and M. Sullivan, SSRv 214, 26 (2018).
- 810 B. Rizk, C. Drouet d'Aubigny, D. R. Golish, D. N. DellaGiustina, and D. S. Lauretta,
- 811 NASA Planetary Data System (2019).
- B. Rozitis, A. J. Ryan, J. P. Emery, P. R. Christensen, V. E. Hamilton, A. A. Simon, D. C.
- 813 Reuter, M. Al Asad, R.-L. Ballouz, J. L. Bandfield, O. S. Barnouin, C. A. Bennett, M. Bernacki,
- 814 K. N. Burke, S. Cambioni, B. E. Clark, M. G. Daly, M. Delbo, D. N. DellaGiustina, C. M. Elder,
- 815 R. D. Hanna, C. W. Haberle, E. S. Howell, D. R. Golish, E. R. Jawin, H. H. Kaplan, L. F. Lim, J.
- 816 L. Molaro, D. P. Munoz, M. C. Nolan, B. Rizk, M. A. Siegler, H. C. M. Susorney, K. J. Walsh,
- 817 and D. S. Lauretta, SciA 6, eabc3699 (2020).
- 818 D. J. Scheeres, A. S. French, P. Tricarico, S. R. Chesley, Y. Takahashi, D. Farnocchia, J.
- 819 W. McMahon, D. N. Brack, A. B. Davis, R.-L. Ballouz, E. R. Jawin, B. Rozitis, J. P. Emery, A.
- J. Ryan, R. S. Park, B. P. Rush, N. Mastrodemos, B. M. Kennedy, J. Bellerose, D. P. Lubey, D.
- 821 Velez, A. T. Vaughan, J. M. Leonard, J. Geeraert, B. Page, P. Antreasian, E. Mazarico, K.
- 822 Getzandanner, D. Rowlands, M. C. Moreau, J. Small, D. E. Highsmith, S. Goossens, E. E.
- 823 Palmer, J. R. Weirich, R. W. Gaskell, O. S. Barnouin, M. G. Daly, J. A. Seabrook, M. M. Al
- Asad, L. C. Philpott, C. L. Johnson, C. M. Hartzell, V. E. Hamilton, P. Michel, K. J. Walsh, M.
- 825 C. Nolan, and D. S. Lauretta, SciA 6, eabc3350 (2020).
- D. M. Sherman and T. D. Waite, American Mineralogist 70, 1262 (1985).
- 827 A. Simon, D. Reuter, N. Gorius, A. Lunsford, R. Cosentino, G. Wind, and D. Lauretta,
- 828 Remote Sensing **10**, 1486 (2018).
- A. A. Simon, H. H. Kaplan, E. Cloutis, V. E. Hamilton, C. Lantz, D. C. Reuter, D. Trang, S.
- 830 Fornasier, B. E. Clark, and D. S. Lauretta, A&A 644, 1 (2020a).
- A. A. Simon, H. H. Kaplan, V. E. Hamilton, D. S. Lauretta, H. Campins, J. P. Emery, M. A.
- 832 Barucci, D. N. DellaGiustina, D. C. Reuter, S. A. Sandford, D. R. Golish, L. F. Lim, A. Ryan, B.
- 833 Rozitis, and C. A. Bennett, Science **370**, eabc3522 (2020b).
- A. A. Simon, D. C. Reuter, and D. S. Lauretta, Journal of Astronomical Telescopes,
- 835 Instruments, and Systems 7, 1 (2021).
- 836 D. Takir, V. Reddy, J. A. Sanchez, L. Le Corre, P. S. Hardersen, and A. Nathues,
- 837 Astrophysical Journal Letters **804**, L13 (2015).
- E. Tatsumi, M. Popescu, H. Campins, J. de León, J. L. R. García, J. Licandro, A. A. Simon,
 37

- 839 H. H. Kaplan, D. N. DellaGiustina, D. R. Golish, and D. S. Lauretta, Monthly Notices of the
- 840 Royal Astronomical Society **508**, 2053 (2021).
- K. J. Walsh, E. R. Jawin, R.-L. Ballouz, O. S. Barnouin, E. B. Bierhaus, H. C. Connolly, J.
- L. Molaro, T. J. McCoy, M. Delbo', C. M. Hartzell, M. Pajola, S. R. Schwartz, D. Trang, E.
- 843 Asphaug, K. J. Becker, C. B. Beddingfield, C. A. Bennett, W. F. Bottke, K. N. Burke, B. C.
- 844 Clark, M. G. Daly, D. N. DellaGiustina, J. P. Dworkin, C. M. Elder, D. R. Golish, A. R.
- 845 Hildebrand, R. Malhotra, J. Marshall, P. Michel, M. C. Nolan, M. E. Perry, B. Rizk, A. Ryan, S.
- A. Sandford, D. J. Scheeres, H. C. M. Susorney, F. Thuillet, and D. S. Lauretta, NatGe 12, 242
- 847 (2019).
- 848 B. Zellner, D. J. Tholen, and E. F. Tedesco, Icarus **61**, 355 (1985).
- X.-D. Zou, J.-Y. Li, B. E. Clark, D. R. Golish, S. Ferrone, A. A. Simon, D. C. Reuter, D. L.
- 850 Domingue, H. Kaplan, M. A. Barucci, S. Fornasier, A. Praet, P. H. Hasselmann, C. A. Bennett,
- E. A. Cloutis, E. Tatsumi, D. N. DellaGiustina, and D. S. Lauretta, Icarus 358, 114183 (2021).