

Perspective on Satellite-Based Land Data Assimilation to Estimate Water Cycle Components in an Era of Advanced Data Availability and Model Sophistication

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39 **Abstract**

40 The beginning of the 21st century is marked by a rapid growth of land surface satellite data and model
 41 sophistication. This offers new opportunities to estimate multiple components of the water cycle via
 42 satellite-based land data assimilation (DA) across multiple scales. By resolving more processes in land
 43 surface models and by coupling the land, the atmosphere, and other Earth system compartments, the
 44 observed information can be propagated to constrain additional unobserved variables. Furthermore,
 45 access to more satellite observations enables the direct constraint of more and more components of the
 46 water cycle that are of interest to end users. However, the finer level of detail in models and data is
 47 also often accompanied by an increase in dimensions, with more state variables, parameters, or
 48 boundary conditions to estimate, and more observations to assimilate. This requires advanced DA
 49 methods and efficient solutions. One solution is to target specific observations for assimilation based
 50 on a sensitivity study or coupling strength analysis, because not all observations are equally effective
 51 in improving subsequent forecasts of hydrological variables, weather, agricultural production, or
 52 hazards through DA. This paper offers a perspective on current and future land DA development, and
 53 suggestions to optimally exploit advances in observing and modeling systems.

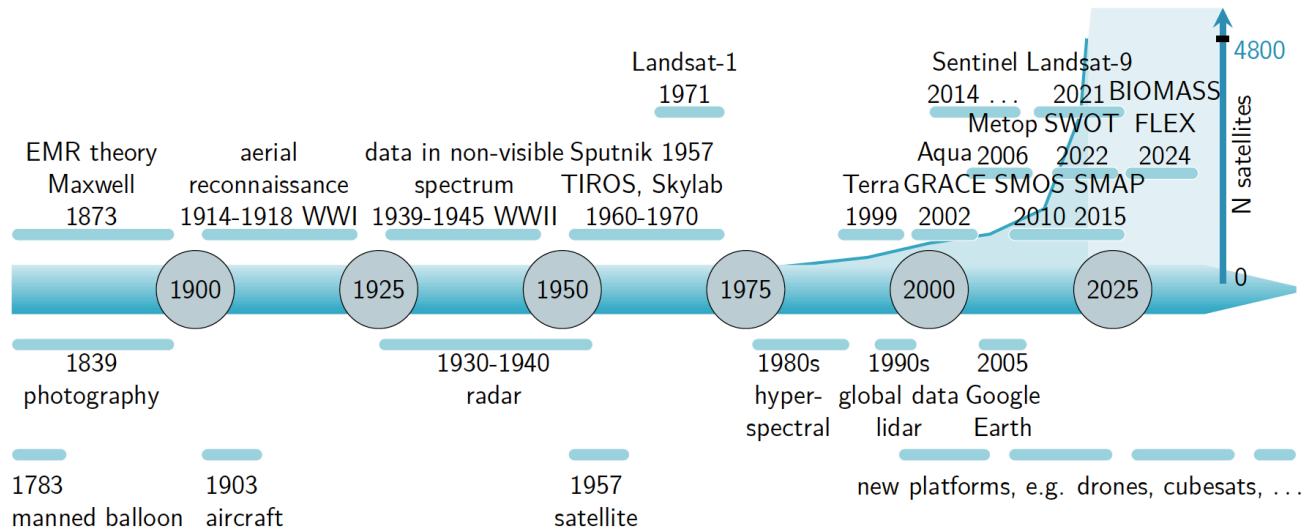
54 **1 Introduction**

55 The distribution of water on Earth determines human livelihoods and is itself influenced by human
 56 activities. Estimating the water availability in various terrestrial compartments is essential for water
 57 resources management, agricultural monitoring, natural hazards and disaster risk assessment,
 58 biodiversity and planet health protection, numerical weather prediction (NWP), seasonal prediction,
 59 and climate change mitigation and adaptation. Currently, the most complete regional- to global-scale
 60 estimates of water-related variables are obtained by merging satellite data records into numerical
 61 models of Earth system processes through data assimilation (DA) (Asch et al., 2016). DA can combine
 62 the unprecedented amounts of satellite data with the steadily acquired process understanding of the
 63 past decades. Specifically, DA uses the satellite observations to correct errors in model simulations,
 64 including errors in unobserved variables. Thereby, DA adds value to the observations by inferring
 65 unobserved information, filling gaps and/or enhancing the spatial resolution of satellite data. In the
 66 geosciences, DA mostly refers to state estimation theory, but it more generally covers any technique
 67 that uses data to estimate the most accurate possible system state (Carrassi et al., 2018) and associated
 68 fluxes. Therefore, DA also encompasses model parameter optimization and the correction of boundary
 69 conditions, including meteorological forcings.

70 Land DA developments have been reviewed earlier (Reichle, 2008, Lahoz and De Lannoy, 2014, De
 71 Lannoy et al., 2016, Jin et al., 2018, Huang et al., 2019, Xia et al., 2019, Giroto et al., 2020, Durand
 72 et al., 2021, Baatz et al., 2021). In parallel to our paper, Kumar et al. (2022, in review) review and
 73 identify current community-agreed gaps and priorities for the future of state estimation via land DA.
 74 In this paper, we reflect on advances in observing, modeling and DA techniques, the associated
 75 opportunities and complexities of future land DA systems, and solutions to keep land DA efficient and
 76 effective, in the presence of rapid data growth and model sophistication in the first half of the 21st

77 century. First, we summarize the state of the art of land DA for the estimation of water cycle variables
 78 (**Section 2**). Next, we offer a perspective on current observing, modeling and DA systems (**Section 3**)
 79 and on the future goals of land DA (**Section 4**). The focus will be on soil moisture, snow and vegetation
 80 estimation and how to extend the impact of satellite-based land DA to improved dynamic estimates of
 81 the atmosphere, vegetation, hydrological and biogeochemical cycles, as well as of natural hazards.

82 **2 State of the Art**



83
 84 **Figure 1.** The 21st century within the history of remote sensing, which started with the understanding
 85 of electromagnetic radiation (EMR). Select satellite missions used for land DA are marked and
 86 discussed in the text.

87 The beginning of the 21st century has seen a sustained increase in remotely sensed data of the Earth
 88 system. **Figure 1** shows the exponential growth in satellite missions, with about 4,800 active satellite
 89 platforms orbiting our Earth in 2021 (<https://www.statista.com/statistics/897719/number-of-active-satellites-by-year/>), but only 20-25% collect Earth observations, and fewer than 1% are regularly used
 90 for land DA. Gravity measurements from the Gravity Recovery and Climate Experiment (GRACE)
 91 and GRACE Follow-On missions directly sense changes in total water storage but at a very coarse
 92 scale. Optical sensors (onboard the Terra, Landsat, and Sentinel-2 missions, among others) measure
 93 fine-scale water content proxies, e.g., snow cover extent, open water extent, vegetation, and soil color
 94 or temperature. Microwave sensors (onboard the Soil Moisture Ocean Salinity -SMOS-, Soil Moisture
 95 Active Passive -SMAP- and Sentinel-1 missions, the Advanced Microwave Scanning Radiometer
 96 onboard Aqua, and the Advanced SCATterometer -ASCAT- onboard Metop, among many more) are
 97 used to retrieve water amounts in the soil, vegetation and snow. The passive radiometer sensors collect
 98 brightness temperature data at a coarse resolution (~40 km), whereas active synthetic aperture radar
 99 (SAR) instruments can collect backscatter data at finer resolutions (< 1 km). Microwave sensors exploit
 100 the fact that the presence of water directly affects the dielectric properties of the soil, vegetation and
 101 snow, and it strongly influences the emission and scattering of microwave radiation (Ulaby et al.,
 102 2014). Insight into how radiation interacts with water in different land compartments is summarized in
 103 radiative transfer models, which can be used in two ways: (i) to invert the observed radiance into
 104 geophysical “retrieval” products (e.g., soil moisture, vegetation or snow water content), or (ii) as so-
 105 called observation operators to map simulated land surface variables to satellite-observed signals (e.g.,
 106 brightness temperature or backscatter).
 107

108 Many land DA systems have used microwave observations to estimate surface and deeper soil moisture
 109 (de Rosnay et al., 2014, De Lannoy et al., 2016, Reichle et al., 2019), and related variables such as
 110 discharge (Lievens et al., 2015, De Santis et al., 2021), turbulent fluxes (Lu et al., 2020), and even
 111 groundwater in peatlands (Bechtold et al., 2020). With the activation of dynamic vegetation models,
 112 the assimilation of optical vegetation indices (e.g., leaf area index) and microwave vegetation optical
 113 depth (Fairbairn et al., 2017, Kumar et al., 2020, Mucia et al., 2022) has gained interest, including to
 114 improve evapotranspiration (ET) and runoff. DA of thermal satellite data has also been popular for ET
 115 and soil moisture estimation (Crow et al., 2008), but studies on the intersection between the water and
 116 energy cycle will not be further discussed, to keep the focus on water cycle variables. At the finer scale,
 117 optical and radar satellite data have been assimilated in crop models to update canopy or soil state
 118 variables and ultimately estimate transpiration, agricultural biomass and yield (Jin et al., 2018, Lu et
 119 al., 2022). Under frozen conditions, the assimilation of optical snow cover fraction or microwave-based
 120 snow depth has been explored (Helmert et al., 2018, Giroto et al., 2021).

121 In practice, land DA systems are developed by merging the theoretical insights in DA, which provide
 122 a portfolio of algorithms, with the operational and physical constraints of land surface observations
 123 and modeling. An overview of regional and global land DA systems is given by Xia et al. (2019). The
 124 observations consist either of satellite retrieval products (soil moisture: Dharssi et al., 2011, Liu et al.,
 125 2011, Rodriguez-Fernandez et al., 2019; vegetation: Albergel et al., 2017, Kumar et al., 2020; snow:
 126 De Lannoy et al., 2010) or direct satellite signals (related to soil moisture: De Lannoy and Reichle,
 127 2016, Lievens et al., 2017, Muñoz-Sabater et al., 2019, Reichle et al., 2019; snow: Larue et al., 2018,
 128 Xue et al., 2018), and most land DA systems consider far fewer observations than state variables (this
 129 characterizes DA in the geosciences at large). For example, one surface soil moisture retrieval every
 130 few days can update soil moisture in multiple soil layers and possibly vegetation, or one weekly snow
 131 cover fraction observation can update the water amount in different snow layers, while the model state
 132 evolves at sub-hourly time steps. Furthermore, most land DA systems are one-dimensional, i.e., they
 133 update each soil-vegetation-snow column (grid cell) independently and the analysis update is strictly
 134 limited to the observed columns. This formulation does not exploit the capability of many DA
 135 approaches to propagate information across the model domain from observed to unobserved areas. If
 136 communication among different columns is made possible via the physics-based model or via spatial
 137 error correlations, thus making the DA system spatially distributed, then state variables in neighboring
 138 (observed or unobserved) columns within the influence radius of a given observation are analyzed
 139 together (Reichle and Koster, 2003, De Lannoy et al., 2010, Magnusson et al., 2014, Reichle et al.,
 140 2019).

141 The above studies all aim at state estimation via particle or Kalman filtering variants (other DA
 142 methods such as variational DA or direct insertion have also been used) to correct the land surface state
 143 for short-term and interannual errors in the meteorological forcings (in offline systems, i.e., not coupled
 144 to an atmospheric model) or other unmodeled temporary deviations in some water compartments. In
 145 this process, only a few DA systems effectively assign the DA corrections to the source of errors, such
 146 as for example snowfall or precipitation input to obtain good snow depth or total water storage
 147 estimates (Winstral et al., 2019, Giroto et al., 2021). Most DA systems do not conserve mass, unless
 148 the water budget is explicitly constrained (Pan et al., 2012).

149 To correct land surface estimates for longer-term or systematic deviations, and to minimize water
 150 budget imbalances, satellite data can be used more effectively for parameter estimation. These
 151 parameters can be part of the prognostic model (Han et al., 2014, Kolassa et al., 2020), the diagnostic
 152 radiative transfer model (De Lannoy et al., 2013, Rains et al., 2022), or represent a bias factor for
 153 meteorological input (Wrzesien et al., 2022). Long-term model calibration could be seen as a form of

154 long-term DA or history matching. Alternatively, DA for sequential parameter updating (with or
 155 without simultaneous state updating) allows to account for time-varying parameters (Montzka et al.,
 156 2011, Magnusson et al., 2016).

157 **3 Perspective on Current Observing, Modeling and DA Systems**

158 **3.1 Observations**

159 The spaceborne observations of many water cycle variables have been improving in radiometric,
 160 spatial, and temporal resolution, but dedicated missions are not yet available for all parts of the water
 161 budget. Soil moisture is now routinely measured at a coarse resolution by dedicated L-band satellite
 162 missions (SMOS, SMAP, Kerr et al., 2010, Entekhabi et al., 2014), and can also be inferred from
 163 shorter wavelength C-band sensors onboard meteorological satellite missions (ASCAT, Figa-Saldana
 164 et al., 2002). Finer-scale estimates can be obtained from current C-band SAR or optical sensors, and
 165 the upcoming NASA-ISRO L-band SAR (NISAR, Rosen and Kumar, 2021) and ESA High Priority
 166 Candidate Mission Radar Observation System for Europe in L-band (ROSE-L, Pierdicca et al., 2019)
 167 are expected to improve fine-scale soil moisture estimates.

168 There is currently no mission devoted to SWE, but various passive microwave sensors have been
 169 combined to produce coarse-scale SWE products (Luojus et al., 2021). The complexity of snow itself
 170 and its presence in complex terrain require more insight on how different types of radiation interact
 171 with snow to support the development of a dedicated mission (e.g. Ku and X-band) for fine-scale SWE
 172 observation. Multi-frequency missions such as the planned Copernicus Imager Microwave Radiometer
 173 (CIMR) will become relevant for SWE remote sensing in the future. Meanwhile, existing sensors have
 174 been used in an opportunistic way (e.g., snow depth from Sentinel-1 radar, Lievens et al., 2022), and
 175 upcoming missions such as NISAR and ROSE-L will further help to estimate high resolution SWE.

176 The water stored in vegetation is also not yet fully observed from space. Several optical vegetation
 177 indices (e.g., leaf area index) approximate the vegetation health and transpiration (Bannari et al., 2009).
 178 More recently, the microwave-based vegetation optical depth (VOD) products have shown promise to
 179 represent biomass, vegetation structure and water (Steele-Dunne, 2017, Wigneron et al., 2021,
 180 Chaubell et al., 2020). The upcoming BIOMASS Earth Explorer mission (Quegan et al., 2019)
 181 promises to explore long wavelength (P-band) measurements to estimate the total biomass in whole
 182 forest layers. Recent studies also aim at the estimation of plant transpiration from novel solar induced
 183 fluorescence (SIF) retrievals (Maes et al., 2020). The upcoming FLEX Earth Explorer mission (Drusch
 184 et al., 2017) will collect SIF data to serve agricultural purposes. Ultimately, advancing VOD and SIF-
 185 based retrievals and gaining insights in how vegetation affects microwave radiation or fluorescence
 186 will lead to better estimates of the water, carbon and energy cycle when combined with dynamic
 187 vegetation and crop yield modeling.

188 Spaceborne observation of water fluxes such as total ET and discharge remains a challenge.
 189 Intermittent satellite-based discharge estimates can be derived from optical and altimeter data (Abdella
 190 et al., 2021, Tarpanelli et al., 2021). The Surface Water and Ocean Topography (SWOT) mission will
 191 soon enable frequent spaceborne observations of river stage for large rivers to allow inference of
 192 discharge (Biancamaria et al., 2016; Frasson et al., 2021). Currently, no mission is specifically
 193 dedicated to ET measurements (Fisher et al., 2017), and ET is most typically inferred from satellite-
 194 observed surface or skin temperature (related to sensible heat) as the residual of a simple energy
 195 balance model (Anderson et al., 2021), or indirectly obtained via soil moisture and VOD DA in a land
 196 surface model (Martens et al., 2017). Most high-resolution ET methods based on optical sensors suffer
 197 from low coverage (clear-sky conditions, low revisit times) and from large discrepancies among the

198 various products. The ECOsystem Spaceborne Thermal Radiometer Experiment on Space Station
 199 (ECOSTRESS, Fisher et al., 2020) mission helped evaluating the use of thermal infrared observations
 200 at fine spatial and temporal resolutions to define future ET mission requirements. The future Land
 201 Surface Temperature Monitoring (LSTM, or Sentinel-8) and Thermal infraRed Imaging Satellite for
 202 High-Resolution Natural resource Assessment (TRISHNA, Roujean et al., 2021) missions promise to
 203 advance ET measurements in the coming decade. The quantification of other water fluxes, such as
 204 irrigation fluxes, from satellite observations is still in its infancy (Kumar et al., 2015, Massari et al.,
 205 2021, Dari et al., 2022).

206 Most DA systems use satellite observations that are directly related to land surface state variables (e.g.,
 207 soil moisture, temperature, vegetation, snow) to improve subsequent state and flux forecasts.
 208 Conversely, the *assimilation of satellite-based flux observations* (e.g., ET, runoff, irrigation) is
 209 relatively less explored and limited to regional applications (Hartanto et al., 2017, Gavahi et al., 2020),
 210 because only a few global flux products are available (mainly ET) and they heavily depend on model
 211 background which might be inconsistent with the assimilation model. Furthermore, the diagnostic flux
 212 DA requires a careful design to link flux observations to prognostic state or parameter updates that are
 213 memorized in the system for improved forecasts. The latter can be achieved e.g., via selecting particles
 214 with a self-consistent combination of parameter, state and flux values in particle filters or via an
 215 adequate observation operator in Kalman filter-based techniques (Pauwels et al., 2006).

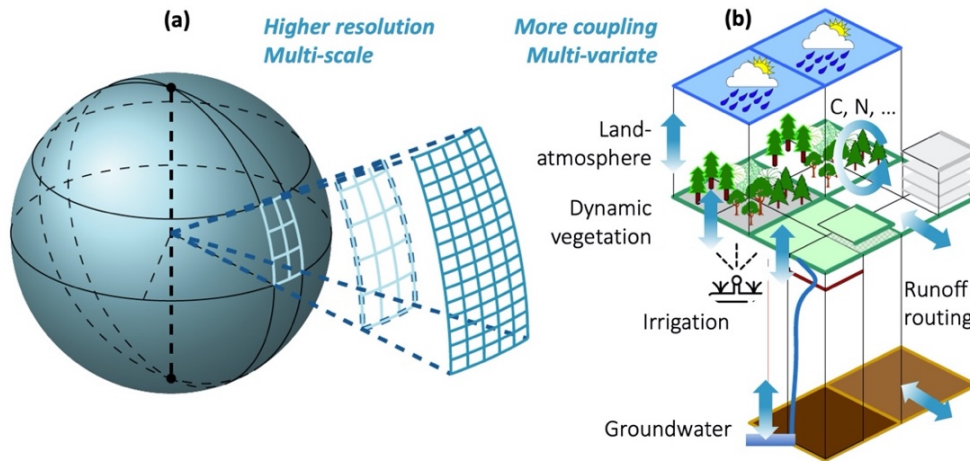
216 *New sensor technologies have not only helped to observe more variables, but also increased the*
 217 *resolution of data.* For example, SAR data can regularly monitor soil moisture and surface water at
 218 km-scale resolution, albeit with more noise, longer revisit times or smaller coverage than coarser-scale
 219 data. Even if some (mainly commercial) sensors are indeed able to measure with high levels of detail,
 220 observations with meter-scale resolution are unlikely to make it into equally fine-scale land DA
 221 systems for global applications any time soon (**Section 4.2**).

222 *The level of satellite observation processing desirable for land DA* is the subject of a debate that should
 223 strengthen the collaboration between geophysical retrieval and DA communities in the future. Land
 224 DA uses satellite observations either in the form of gridded radiances collected by the sensor or as the
 225 associated geophysical retrievals. Just like retrieval DA, radiance DA has been used to update the land
 226 surface state (examples below) and parameters in the land surface or radiative transfer model (Han et
 227 al., 2014). Radiance DA requires a forward model to relate the land surface state (soil moisture,
 228 temperature, snow, vegetation water content) and parameters (clay fraction, vegetation scattering
 229 albedo) to the satellite radiance signals as part of the observation operator (Reichle et al., 2014). The
 230 observation operator can also deal with the difference in spatial support of the observations and
 231 simulations in a multiscale DA system, e.g., to downscale coarse-scale observations to a finer
 232 resolution. Some studies report little DA skill difference between radiance and geophysical retrieval
 233 assimilation (De Lannoy and Reichle, 2016, Aires et al., 2021), and other studies show that radiance
 234 DA can circumvent biases associated with retrievals. For example, for deep mountain snow, SWE
 235 retrievals can be significantly biased (e.g. Wrzesien et al., 2017), but microwave radiance DA allowed
 236 both Li et al. (2017) and Kim et al. (2019) to achieve unbiased SWE estimates. Furthermore, DA of
 237 radiances facilitates the simultaneous updating of multiple state variables (e.g., soil moisture,
 238 temperature and vegetation) more elegantly than DA of the various associated individual retrieval
 239 products with cross-correlated errors. Radiance DA is also physically more self-consistent than
 240 retrieval DA, because retrievals are constrained by background information that may deviate from that
 241 of the model. E.g., soil moisture retrieval may use temperature information, and soil or vegetation
 242 parameters from data sources that are different from those of the model. The physical consistency
 243 makes radiance DA particularly attractive for coupled land-atmosphere DA (de Rosnay et al., 2022).

244 Finally, the observation error characterization is more traceable for radiance DA. In the realm of DA
 245 algorithms, the use of (nonlinear) observation operators enables solving DA as a nonlinear optimization
 246 problem, without (or with limited) relying on linearity assumptions. In short, *satellite observations*
 247 *should be provided along with good observation operators that can support land DA.*

248 The spatio-temporal characterization of the observation error (that is, retrieval or instrument error, plus
 249 representativeness error) is a key element to successful DA systems. New sensor developments would
 250 thus ideally be preceded by a synthetic *observing system simulation experiment* (Crow et al., 2005) to
 251 quantify the tolerable levels of uncertainty for efficient DA. Furthermore, observations and model
 252 estimates typically have distinct biases, which are ideally resolved, explained, or removed prior to state
 253 updating (see **Section 3.3** and **4.1**). *This requires that satellite missions span enough years to quantify*
 254 *climatological biases in observation space*, and this has so far limited the use of short-lived exploratory
 255 missions onboard new platforms (e.g., drones, cubesats) in DA systems.

256 **3.2 Models**



257
 258 **Figure 2.** Land surface model sophistication increased in the 21st century towards (a) higher
 259 resolutions, by resolving multi-scale hydrological processes and improving model parameterizations,
 260 and (b) coupling of more processes, by replacing simplistic parameterizations and including more
 261 interactions between variables in multiple compartments.

262 The beauty of nature is that it is intelligible and can be captured in general physical laws, despite its
 263 complexity in the details. This knowledge is indispensable to add value to observations, and to inter-
 264 and extrapolate them to unobserved variables. In the last decades, a slow but steady increase in
 265 sophistication of large-scale land surface modeling and DA systems has been achieved (Fisher and
 266 Koven, 2020) by (i) improving model parameterizations (Balsamo et al., 2009) and resolving
 267 multiscale processes (**Figure 2a**), (ii) improving prognostic representations of hydrological processes,
 268 such as e.g., lateral subsurface flows in aquifers (Shrestha et al., 2014), snow processes (Bartelt and
 269 Lehning, 2002, Deschamps-Berger et al., 2022), peatland-specific processes (Bechtold et al., 2020),
 270 (iii) improving prognostic representations of vegetation (Clark et al., 2011), biogeochemical cycles
 271 including the nitrogen cycle (Oleson et al., 2013) and phosphorus cycle (Goll et al., 2017), (iv)
 272 activating anthropogenic processes, such as irrigation (Lawston et al., 2017), or by (v) land-atmosphere
 273 coupling (**Figure 2b**).

274 *By shifting from parameterized to physically resolved modeling (e.g., static parameterized to*
 275 *prognostic dynamic vegetation) and by coupling more processes, the DA impact of a single observation*

276 *can reach more unobserved, but model-resolved, compartments.* For example, snow depth DA can
 277 improve discharge and low-level atmospheric estimates (Griessinger et al., 2019, Lahmers et al., 2022,
 278 Rudisill et al., 2021), and backscatter DA can update dynamic vegetation and soil moisture, to
 279 eventually update irrigation (Modanesi et al., 2022). Efforts are ongoing to advance land DA in coupled
 280 land-atmosphere models (de Rosnay et al., 2014, Boussetta et al., 2015, Carrera et al., 2019, Reichle et
 281 al., 2021b) to make good on the promise to improve NWP and subseasonal to seasonal predictions
 282 (Kumar et al., 2022). As a matter of fact, the use of physics-based models has also been pivotal to the
 283 success of atmospheric DA in NWP to propagate information to unobserved areas (Kalnay, 2002). At
 284 the same time, several studies with current state-of-the-art land surface models also reported limited
 285 success (Crow et al., 2020, Hung et al., 2022) in propagating information from one compartment to
 286 another, which suggests that the modeling (parameterization) of the coupling and fluxes between land
 287 compartments as well as DA strategies need further research.

288 Surprisingly few new prognostic physics-based models (or model components) have been developed
 289 in response to the growing number of satellite data. This might be because we have reached the
 290 maximal desired structural complexity for large-scale applications, or because the coarse resolution of
 291 many satellite data integrates too much spatial variability, complicating a clean local physical
 292 interpretation of processes. As both model simulations and satellite observations become available at
 293 higher resolutions and for longer time spans, more spatial and temporal scales get resolved (**Figure**
 294 **2a**). This might possibly deepen our process understanding, limit parameterizations and ultimately help
 295 hyper-resolution modeling (Wood et al., 2011) and DA.

296 Alternative ways of model development are emerging, which in fact have the potential to use the
 297 growing amount of (possibly coarse-scale) data and artificial intelligence rather than our human
 298 intelligence to build a model. Specifically, *machine learning* (ML) holds promise (Nearing et al., 2020)
 299 to develop models for multiple variables directly from multiple types of observations. E.g., ML can be
 300 used to diagnose how satellite-observed signals can be related to a set of land surface variables via
 301 complex interactions. Especially for microwave-based observation operators (Xue et al., 2018, Shan et
 302 al., 2022), ML might currently be more efficient than trying to fully understand and parameterize all
 303 radiation interactions. It is however unclear if ML is capable of entirely replacing prognostic land
 304 surface models in Earth system models, given that ML is not well-suited for nonstationary systems
 305 (e.g., under climate change), or to support the inference of unobserved land variables, because ML
 306 typically employs supervised learning that requires the existence of observations prior to training. More
 307 pragmatically and potentially more successfully, ML might complement physically based descriptions
 308 in a hybrid fashion (Reichstein et al., 2019). Note that in this subsection, ML is presented as a tool for
 309 model development. **Section 3.3** discusses how ML can be used for DA.

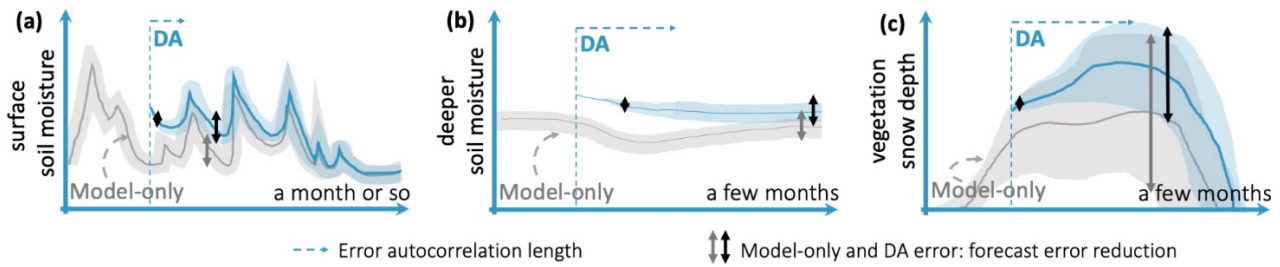
310 Ideally, models offer a framework to propagate observations to unobserved variables, but models are
 311 imperfect, and their uncertainties originate from errors in the numerical schemes, unresolved scales,
 312 parameters, initial conditions, meteorological input (in offline systems) or missing processes. Via DA,
 313 the model state, parameters or forcing inputs will be updated to correct the model trajectory. If
 314 parameterizations are replaced with physically resolved process descriptions and the associated
 315 parameters would become physically measurable, then the need to update parameters should reduce in
 316 favor of more state updating. Similarly, when offline forcing inputs are replaced with coupled land-
 317 atmosphere modeling and constrained by atmospheric observations, then the need to update
 318 meteorological input in land surface models should reduce in favor of more state updating.

319 **3.3 DA Methods**

320 The choice of DA method for a given application is arguably often driven by the research group's
 321 repository of source code, and it is rarely optimal in a mathematical sense (Carrassi et al., 2018).
 322 However, the discontinuity (e.g., via activation thresholds) and non-differentiability of land surface
 323 processes (including prognostic soil-water-vegetation-snow and diagnostic radiation interactions) is a
 324 valid reason to *favor ensemble Kalman or particle-based techniques* (Evensen et al., 2022), instead of
 325 variational methods that require model adjoints, which are difficult to obtain and maintain.
 326 Furthermore, ensemble- and particle-based DA can diagnose flow-dependent forecast error estimates
 327 for nonlinear land surface models. Like in other areas of climate science, filters dominate operational
 328 land DA systems because they naturally support the sequential inclusion of satellite observations,
 329 provided they are available to describe an optimal current state for subsequent forecasts. For longer-
 330 term re-analysis solutions, or for slowly varying variables, smoothers (Dunne and Entekhabi, 2006,
 331 Margulis et al., 2015) gather observations over a sliding retrospective time window to obtain the best
 332 historical solution.

333 The key to any DA method is in the *treatment of the forecast and observation errors*. State estimation
 334 assumes random errors. In the ensemble Kalman filter (EnKF) or particle filter (PF), the distribution
 335 of the random forecast error accumulated between assimilation time steps is diagnosed from an
 336 ensemble of realizations, and ensemble generation is an art by itself (choice of perturbations, variable
 337 transformations to obtain Gaussianity, covariance inflation, localization; Carrassi et al., 2018). The
 338 observation errors are typically set to a constant standard deviation parameter that reflects the
 339 instrument or retrieval error, increased by the representativeness error that also includes observation
 340 operator error (Tijana et al., 2018). The forecast and observation error estimates are typically
 341 hyperparameters optimized by manual tuning of DA diagnostics (Reichle et al., 2017), because
 342 automated adaptive filtering (Crow and Reichle, 2008, De Lannoy et al., 2009) remains too inefficient.
 343 Most DA methods rely on the assumption of unbiased sources of information, and thus biases are
 344 typically removed prior to DA via, e.g., cumulative distribution function matching between the
 345 assimilated observations and the model simulations (Reichle and Koster, 2004, Kumar et al., 2012);
 346 consequently, DA analyses are consistent with the (potentially erroneous) model climatology. Ideally,
 347 biases are disentangled to estimate (De Lannoy et al., 2007, Pauwels et al., 2013) and possibly remove
 348 forecast or observation bias, or perhaps to identify the impact of water management or other human
 349 activity (e.g., unmodeled groundwater pumping and irrigation; Kumar et al., 2015, Giroto et al., 2017).

350 The nature of the errors associated with different land variables is very different. **Figure 3** illustrates
 351 how soil moisture has a more bounded error growth than snow or vegetation, and that a single DA
 352 update reduces the forecast error for a longer time in variables with longer error autocorrelation lengths.
 353 *More research is needed on how to best address these various types of errors for different variables*
 354 *via different DA methods and different bias treatments*. For example, state-only updating without
 355 observation bias correction was advantageous to correct the accumulated snow and the associated river
 356 discharge in Smyth et al. (2019) and Lahmers et al. (2022), but in other studies, snow observation bias
 357 correction (De Lannoy et al., 2012, Liu et al., 2013) or bias correction to snowfall (Magnusson et al.,
 358 2016) was preferred. Similarly, Albergel et al. (2017) and Kumar et al. (2020) used a bias-blind filter
 359 for vegetation updating, but omitting bias correction for vegetation observations can possibly cause
 360 undesirable sawtooth timeseries and inferior ET and runoff estimates when assimilating intermittent
 361 observations, when the model is pushed out of its statistical equilibrium. The need for observation bias
 362 correction might depend on the boundedness of variables and the coupling between variables in
 363 different models, i.e., whether there is strong circular coupling equilibrium (vegetation-transpiration-
 364 soil moisture-vegetation) or rather a dominant one-way coupling (snowpack-discharge).



365

366 **Figure 3.** Different error characteristics of (a) surface soil moisture, (b) deeper soil moisture, and (c)
 367 vegetation or snow depth, for (gray) model-only and (blue) DA simulations. The semi-transparent
 368 plumes represent ensemble uncertainties.

369 Finally, most of the above considerations hold for DA in the traditional sense of merging physics-based
 370 model variables with satellite observations. New, data-driven methods such as ML offer an alternative
 371 to DA and can in some ways be similar to four-dimensional variational DA (by including the time
 372 dimension as in smoothers). Like DA (Geer et al., 2021), ML can be used to obtain better state
 373 estimates, bias estimates (Pan et al., 2021) or parameter estimates (Mudunuru et al., 2022). *Novel*
 374 *hybrid DA-ML methods* (Bonavita et al., 2020) are showing success in discovering and emulating
 375 unresolved-scale processes, whenever a chronic lack of data makes the task extremely difficult for pure
 376 ML. In the context of coupled atmosphere-ocean modeling, DA-ML has shown promising results
 377 (Brajard et al., 2021) and its future use in a land-atmosphere context could be attractive.

378 4 Perspective on Future DA Development

379 4.1 Land DA Goals of the Future: Priorities

380 From its origin in atmospheric and ocean sciences, DA for state updating provides the best-possible
 381 initial conditions for subsequent forecasts. Properly estimating the initial state is critically important in
 382 chaotic systems (Carrassi et al, 2022), where small errors can grow exponentially in time and where
 383 the characteristics of such growth are themselves unpredictable. By contrast, land systems are usually
 384 asymptotically stable. Therefore, initial errors are typically internalized in the state (memory) until the
 385 system reaches an equilibrium after some time. Nevertheless, in a coupled land-atmosphere system, a
 386 small land initialization error could result in exponential error growth in the atmosphere. Seemingly
 387 small improvements obtained via land DA are therefore critical for NWP and seasonal predictions,
 388 provided the coupling mechanisms for long-term predictability are well represented. In land-only
 389 applications, state updating is essential to reset cumulative vegetation or snow variables for seasonal-
 390 scale yield or discharge forecasts, or to adjust soil moisture or input forcings for more accurate short-
 391 term hazard predictions of landslides (Felsberg et al., 2021), fires (Jensen et al., 2018), floods (Massari
 392 et al., 2018), and droughts (Li et al., 2019).

393 However, DA is not equally effective in all circumstances. For example, soil moisture updating can
 394 generally improve streamflow predictions (Mahanama et al., 2012, Reichle et al., 2021a), but might
 395 not be effective to reduce errors in the fast runoff component which are dominated by rainfall errors
 396 (Mao et al., 2020). Similarly, the influence of soil moisture on ET depends on the seasons, the coupling
 397 strength between soil moisture and ET in different climate regimes (Dong et al., 2020) and the ability
 398 of the assimilation model to accurately capture that coupling (Crow et al., 2020). To use data and
 399 resources most efficiently in a century when ever more data are becoming available (**Section 4.2**), one
 400 should wonder which specific type of observations at what time and location has the largest impact on
 401 land DA analyses and beyond. A suggestion for future research is thus to explore *targeted land DA*.
 402 This requires that we first determine which type of observations are most useful to improve the forecast

403 skill of particular land or atmosphere variables under the given circumstances, via sensitivity studies,
 404 forecast sensitivity-based observation impact studies, or coupling strength analyses. The land DA
 405 community can learn from the NWP community, which already has a strong grasp on how much
 406 various observations contribute to forecast skill (Eyre et al., 2021). Thereafter, we can efficiently
 407 assimilate those observations that likely have the most impact. The limitation is that satellite
 408 observations are collected in fixed orbits and not necessarily at the strategically most optimal location
 409 or time, so the main goal of targeting observations is a careful selection of the available observations.

410 Apart from state updating, satellite DA should be further explored for parameter estimation to (i)
 411 improve inherited static global soil and vegetation parameter databases that served older model
 412 generations, and to (ii) assign values to newly resolved parameters that will emerge from the
 413 sophistication of land surface models, e.g. to parameterize dynamic vegetation growth or water table
 414 dynamics. Parameter estimation is in principle possible using the same DA framework used for state
 415 estimation. Nevertheless, the success of parameter updating depends on the model sensitivity to that
 416 specific parameter and to its correlation with the observed quantities. The latter could be automatically
 417 estimated via the ensemble (for the EnKF) or the particles (for the PF), thus further promoting the use
 418 of this family of methods. Recently, hybrid EnKF-PF methods have been developed precisely to use
 419 the EnKF for the “more linear” state update and the PF for the “more nonlinear” parameter updates
 420 (van Leeuwen et al., 2019). Those methods could prove effective in land DA as well. Finally, parameter
 421 updating is particularly relevant for long-term applications, because DA frameworks for state-only
 422 updating rely on the assumption of the system being autonomous and stationary and are thus not
 423 theoretically suitable if a system is subject to climate change or human activities that cause changes in
 424 the system’s equilibrium. This is a broader issue for DA and goes well beyond the realm of land DA.
 425 Usually, the assumption is that by sequentially updating the system state and parameters we drive the
 426 conditional posterior probability toward the new equilibrium, yet rigorous mathematical results along
 427 these lines are still missing.

428 DA can be used to correct the state, parameters or forcings for unmodeled or poorly modeled processes,
 429 such as e.g. human activities. For example, Saharan dust deposited on snow should result in a sudden
 430 update of the parameterized or simulated albedo to ensure correct snow melt estimates. A forest fire,
 431 deforestation, land use change, or crop rotation within or across years (Boas et al., 2021) all require
 432 updates of vegetation model parameters or states. Such events will be followed by a gradual adjustment
 433 to a new soil moisture equilibrium both in the model and reality, but the transition time might differ,
 434 because some (unobserved) model parameters that determine the transition time are not in line with
 435 reality. The same is true for land systems in the presence of climate change, which might necessitate
 436 gradually changing model parameters. How to *combine long-term updates for poorly modeled*
 437 *processes via parameter updating with short-term state updating* should be explored in the future.

438 DA diagnostics of observation-minus-forecast and analysis-minus-forecast residuals allow an
 439 evaluation of the optimality of the DA system (Desroziers, 2005, Reichle et al., 2017). These
 440 diagnostics could in the future also help to identify (and improve) times and locations of poorly
 441 modeled processes, or system transitions from steady state to a new equilibrium.

442 DA aims at blending multiple sources of information seamlessly. However, in one-dimensional DA
 443 systems, no horizontal information propagation is achieved, which can result in artificial spatial
 444 patterns (e.g., swath edges or cloud screening imprinted in the DA analysis). When the land DA is
 445 coupled to an atmospheric model, such spatial discontinuities could lead to undesirable triggering of
 446 turbulence (Alapaty et al., 1997). Furthermore, only a part of the model variables might be included in
 447 the DA state vector. E.g., only a few soil moisture layers might be updated out of all soil-vegetation

448 variables, or only the land variables and no atmospheric variables might be updated in a coupled DA
 449 system. To *avoid unphysical discontinuities* at the border between domains (e.g., land vs. atmosphere,
 450 or observed vs. non-observed land) or at the interface between variables, *spatially distributed and*
 451 *multivariate DA methods* are recommended, where multiple state variables of the land surface and
 452 coupled processes are updated.

453 As an extension of multivariate DA, the use of coupled DA is seen as another key area of desired DA
 454 development (de Rosnay et al., 2022). Strongly-coupled DA intends to inform one component of the
 455 climate system (e.g., the land) by using observations of the other (e.g., the atmosphere) and vice-versa
 456 (Penny and Hamill, 2017). This contrasts with the so called “weakly-coupled DA” in which the analysis
 457 update only affects the model compartment where data are collected, but then a coupled model is used
 458 in the forecast step. The model usually acts as a dynamical way of propagating information from the
 459 observed to the unobserved component, and weakly-coupled DA is usually developed first towards the
 460 ultimate goal of strongly-coupled DA. The spatio-temporal difference between processes in the
 461 coupled media (e.g., land-atmosphere) make it extremely difficult to construct a suitable error
 462 covariance across them (Tondeur et al., 2020). *The sophistication of DA techniques will need to grow*
 463 *with a stronger coupling* of the simulated water, energy and biogeochemical cycles (Baatz et al., 2021)
 464 in land surface, terrestrial ecosystem and atmospheric modeling and with the use of multivariate
 465 constraints across the various compartments of these coupled systems.

466 **4.2 Increased Dimensions of Future Land DA: Challenges and Opportunities**

467 Most visions for future land DA include multisensor DA (Durand et al., 2021), multivariate DA (Kumar
 468 et al., 2022), and multi-scale DA with a push towards finer resolutions. Our priorities above should be
 469 viewed against the backdrop of these foreseen developments, and here we highlight some associated
 470 opportunities. A multisensor approach is recommended to constrain more water cycle variables
 471 (Giroto et al., 2019) and obtain finer spatial and temporal resolutions, e.g., to benefit from the higher
 472 accuracy of coarse-scale observations and from the spatial detail in fine-scale observations (De Lannoy
 473 et al., 2012, Lievens et al., 2017). The use of multiple independent observations also has the *potential*
 474 *to mitigate equifinality problems, i.e., to identify the state variable, input or parameter correction, or*
 475 *combination thereof, that results in the most effective constraint (e.g., particle selection)*. As discussed
 476 above, multivariate DA is needed for physical consistency and to reach more unobserved variables in
 477 more sophisticated systems. Higher-resolution (km-scale) DA systems promise to better resolve local
 478 land details for improved NWP and land-atmosphere reanalysis products. Higher resolutions for
 479 coupled land surface-subsurface models also better represent runoff processes at the hillslope scale,
 480 and narrow valleys with underlying groundwater bodies, which affect the simulation of ET (Shrestha
 481 et al., 2018). Furthermore, high-resolution estimates are needed for agricultural applications and hazard
 482 estimation.

483 From the viewpoint of system theory, these desires for high-resolution multivariate and multisensor
 484 DA translate into *larger dimensions of state and observation error covariances, which necessitates*
 485 *practical and computationally affordable solutions*. Larger updated state vectors require larger
 486 ensemble sizes to mitigate the sampling error in the ensemble-based error covariances, or beg for
 487 alternative solutions to partition the state into less-dependent groups of correlated variables that can be
 488 updated sequentially (thus making the ensemble covariance essentially block-diagonal). Indeed, the
 489 latter grouping of state variables conceptually mimics the idea of localization to filter out spurious error
 490 correlations in spatially distributed DA systems, and is also the essence of weakly-coupled DA systems.

491 Assimilating more observations from multiple sensors, multiple products, or high-resolution datasets
 492 increases the dimensions of the observation error covariance matrix. Spatially neighboring
 493 observations, or joint soil moisture and vegetation retrievals from the same microwave sensor, cannot
 494 be assimilated independently due to associated error correlations. *Solutions can be found in directly*
 495 *assimilating radiances rather than multiple derived retrieval products, targeting only those*
 496 *observations that have most impact on the forecast (Section 4.1), and thinning the observations* (Waller
 497 et al., 2018). A second problem with assimilating multiple observations is that they each might have
 498 their own bias, represent something different than the model variables, and/or might cause
 499 contradicting updates (Giroto et al., 2019). Appropriate uncertainty estimates and bias removal partly
 500 solve this problem, and allow DA to update the temporal variability, while preserving the model's
 501 climatological water distribution as a strong constraint (Pan et al., 2012). *Ideally, when observation*
 502 *biases get resolved and we find adequate ways to relate modeled and observed land estimates in*
 503 *absolute terms, then the multitude of observations should be used to also correct the modeled terrestrial*
 504 *water partitioning, and thereby create the correct climatological land conditions to support the correct*
 505 *coupling regimes.*

506 Finally, using models, observations and DA at ever finer resolutions inevitably requires advanced
 507 computational infrastructure, more background information, e.g., on land surface processes, soil and
 508 land use parameters, high-resolution meteorological information (for off-line land simulations), and a
 509 DA method that can address the problem complexity (Carrassi et al., 2018). Furthermore, fine-scale
 510 estimates are by their nature more uncertain than the aggregated counterparts. In the future, we will
 511 have to balance the advantages of resolving more detail against the curse of dimensionality.

512 **5 Conclusion**

513 Satellite-based land DA is an interdisciplinary field of research that yields the most complete and
 514 consistent estimates of terrestrial water cycle variables. The growing amount of satellite data and the
 515 sophistication of modeling systems in the 21st century require efficient land DA systems to fuse
 516 observations and models into meaningful information for end users. Land DA can convert the
 517 intermittent swaths of satellite signals into temporally and spatially complete, gridded fields of soil
 518 moisture, snow or vegetation estimates and related variables, including land surface fluxes such as ET
 519 and runoff. By coupling the land with atmosphere or groundwater processes, and by resolving
 520 vegetation or snow parameterization schemes with physics-based processes, observations have the
 521 potential to update more unobserved variables, and to have an impact beyond the land surface. This is
 522 especially the case for NWP, crop monitoring, hazard (landslides, fires, floods, droughts) assessment,
 523 and carbon management.

524 Large, dynamical modeling systems that include more resolved or coupled processes, require
 525 sophisticated DA techniques (perhaps supplemented with ML) to optimally distribute the observed
 526 information into improved estimates of the multivariate state, parameters, or boundary conditions. The
 527 exponential growth of satellite data will support improved constraints of the advanced modeling
 528 frameworks, but the growing dimensions in land DA will also necessitate the development of efficient
 529 DA algorithms. It will thus become increasingly important to select the most suitable levels of
 530 observation processing and the most impactful observations for assimilation, because not all
 531 observations are equally efficient all the time in DA systems. We can curb the growth of state and
 532 observation dimensions in the DA problem by considering targeted DA, rather than a mass integration
 533 of all data.

534

535 **6 Conflict of Interest**

536 *The authors declare that the research was conducted in the absence of any commercial or financial*
537 *relationships that could be construed as a potential conflict of interest.*

538 **7 Author Contributions**

539 GDL developed the paper structure and wrote the initial draft, all co-authors contributed to the shaping
540 of the paper contents and to the writing and editing of the paper.

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546

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