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12 **Keywords: Wildfire, Debris Flows, Mass Wasting, Remote Sensing, Machine Learning, IMERG**

13 **Abstract**

14 Wildfire is a global phenomenon that has dramatic effects on erosion and flood potential. On steep
15 slopes, burned areas are more likely to experience significant overland flow during heavy rainfall
16 leading to post fire debris flows (PFDFs). Previous work establishes methods for PFDF hazard
17 assessment, often relying on regional-scale parameterizations with in-situ rainfall measurements to
18 categorize hazard as a function of meteorological and surface properties. We present a globally
19 scalable approach to extend the benefit these models provide to new areas. Our new model relies on
20 publicly available satellite-based inputs with a global extent to provide first order hazard assessments
21 of recently burned areas. Our results show it is possible to identify the conditions relevant for
22 PFDF-initiation processes across a variety of physiographic settings. Improvements to satellite-
23 borne rainfall intensity data and increased availability of PFDF occurrence data worldwide are
24 expected to enhance model skill and applicability further.

25 **Plain Language Summary**

26 Wildfires are destructive hazards whose subsequent effects can be long lasting. This includes the
27 increased risk of flooding and post fire debris flows (PFDFs). We present a framework that uses
28 freely available satellite-based data to inform where PFDF activity may be elevated. We use a
29 standardized set of resources available globally such that it is easier to make direct comparisons
30 between regions. This work lays the groundwork to more thoroughly investigate why PFDFs are
31 more common in some regions over others, as well as demonstrate the use of satellite data for timely
32 assessments of the hazards that follow wildfire.

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37 **1. Introduction**

38 Driven by the effects of wildfire, burned steplands are vulnerable to the cascading hazard of
39 runoff-induced post fire debris flows (PFDFs). These hazards are a consequence of the reduced
40 infiltration capacity of burned soils (Debano, 2000; Letey, 2001) and an increase in available
41 sediment for transport (Florsheim et al., 1991, 2016; Gabet, 2003; Lamb et al., 2011, 2013). The
42 mixing of runoff from heavy rainfall and loose debris can create flows capable of destroying
43 infrastructure and taking the lives of those in the path downslope. As wildfires are already occurring
44 at higher frequencies and severities (Abatzoglou & Williams, 2016; Cannon & DeGraff, 2009;
45 Dennison et al., 2014; Keeley, 2009; Miller & Safford, 2012; Mueller et al., 2020; Westerling et al.,
46 2006), there is a continued need to understand how differences in topography, burn severity, and
47 soil properties may affect intensity thresholds for PFDF initiation.

48 Previous work details several methods for PFDF hazard assessment at the regional or local scale.
49 These approaches include physically-based rainfall intensity-duration threshold delineation via
50 dimensionless discharge and Shields stress criterion (Tang et al., 2019), and data-driven empirical
51 approaches that assign a probability of PFDF initiation and/or estimate of sediment delivery from
52 field-based or remote sensing-derived preconditions (e.g., Gartner et al., 2014; Nikolopoulos et al.,
53 2018; Nyman et al., 2015; Staley et al., 2016). The United States Geologic Survey (USGS)
54 implements a version of the logistic regression model from Staley et al. (2016) as an operational tool
55 to derive rainfall intensity thresholds for PFDF initiation within the Contiguous United States
56 (CONUS). This model is trained on a database compiled by the USGS that documents PFDF
57 occurrence (or the lack thereof) and the associated storm characteristics concurrent with that
58 observation since 2000. Empirical models, such as the one utilized by the USGS, remain a popular
59 approach for PFDF hazard assessment as numerical or physically based models traditionally rely on
60 site-specific parameterizations. These data-driven methods link heterogenous empirical observations

61 across physiographic regions to assign spatially variable probabilities of PFDF occurrence. From an
62 operational standpoint, empirical approaches provide a computationally efficient and transferrable
63 method appropriate for rapid hazard assessment, even if the entirety of PFDF initiation processes
64 cannot be fully derived from the chosen model inputs alone.

65 Since 2016, NASA's Goddard Space Flight Center has run the near-global *LHASA* (Landslide
66 Hazard Assessment for Situational Awareness) model (Kirschbaum & Stanley, 2018; Stanley et al.
67 2021). This model provides near-real time assessments of landslide hazards worldwide with
68 approximately 3hrs latency. To date, this model has not explicitly incorporated the impacts of
69 wildfires on mass wasting processes into its assessment framework. Following the ever-increasing
70 importance of wildfires and their subsequent effects, this work describes a complementary model
71 for PFDF hazard susceptibility to provide physically interpretable assessments to burned steeplands
72 at the same scale and timeframe.

73 Further, this work details the first PFDF hazard assessment model solely from publicly available,
74 global scale remote sensing data and demonstrates the utility that it can provide for vulnerable
75 communities and infrastructure. Although PFDF activity is well documented and actively monitored
76 in the United States, PFDFs also occur in other regions (e.g, García-Ruiz et al., 2013; Jin et al., 2022;
77 Nyman et al., 2019). To our knowledge, no work has attempted to reconcile the entire spatial extent
78 of these events to quantify and dynamically monitor PFDF hazard potential in all areas where they
79 occur. We thus propose a framework to provide situational awareness of these events in near real-
80 time.

81 **2. Methods**

82 Our proposed model relies on PFDF occurrence data provided by the USGS. Staley et al. (2016)
83 highlight a comprehensive database documenting the relevant slope, burn severity, and soil

84 properties of several burned watersheds across the Western United States. This database further
85 records the intensity-duration characteristics of select storms that occurred in each area, indicating
86 whether a debris flow initiated from each storm. This is the most extensive database of PFDF
87 activity ($n = \sim 1500$) that is publicly available, and it serves as the basis for the current model in use
88 by the USGS. As this database consists of field observations, local rain gauge data, and remote
89 sensing using CONUS-specific datasets, we utilize the information contained therein as a guide for
90 replicating the data collection process using only satellite-derived resources with a worldwide extent.
91 Taken together, these publicly available products help recreate the training database provided by the
92 USGS in a format sufficient for the needs of this study. Notably, we do not include soil properties
93 used in the original USGS model. While global soil property products exist, they are inconsistent in
94 both accuracy and compilation method. Accordingly, we elect only to include the extensively tested
95 datasets provided by NASA and the USGS, as these products must meet a consistent set of
96 standards before publication ([https://earthdata.nasa.gov/collaborate/open-data-services-and-](https://earthdata.nasa.gov/collaborate/open-data-services-and-software/data-information-policy/data-levels)
97 [software/data-information-policy/data-levels](https://earthdata.nasa.gov/collaborate/open-data-services-and-software/data-information-policy/data-levels)).

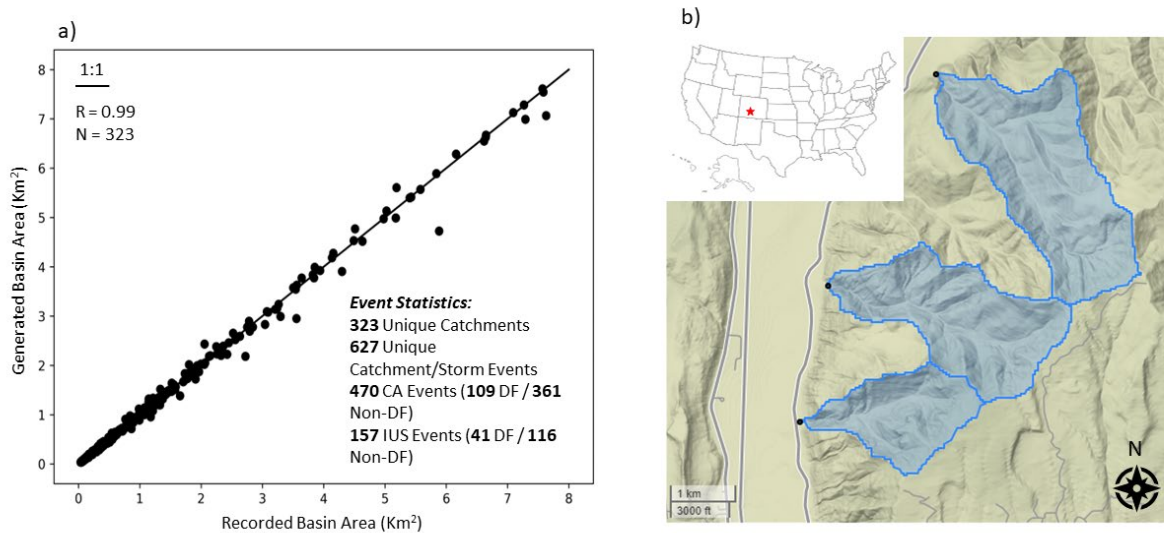


Figure 1. (a) Comparison of catchment areas recorded by Staley et al. (2016) and the catchment areas generated in this work. (b) Example catchments derived in this work (blue) and their corresponding point reference via Staley et al (2016) (black) located near Durango, Colorado. IUS is defined as the Interior United States; CA is defined as Southern California.

99 An important consideration in the database recreation process is the availability of precise spatial
 100 reference information of each documented catchment. The data made publicly available by Staley et
 101 al. (2016) link detailed catchment and storm occurrence data to coordinates typically located near
 102 basin and/or channel outlets. To relate new observations to the same original assessment areas, it is
 103 therefore necessary to delineate catchment boundaries based on the location of each datapoint. We
 104 infer catchment boundaries using flow routing software provided by Bartos (2020) and the NASA
 105 DEM 30m Digital Elevation Model (NASA JPL, 2020) as a source dataset. The contributing area of
 106 each newly generated catchment is compared with the originally recorded value assigned to each
 107 point, where both values must be within approximately 20% agreement. This threshold is chosen
 108 from manual inspection to ensure that the inferred areas are a reasonable estimation of the original
 109 assessment area. Figure 1a provides a quantitative comparison of the resulting catchments areas

110 within this similarity threshold ($n = 323$). Visual analysis suggests that these catchments are realistic
111 (Figure 1b).

112 Following delineation, each catchment provides the spatial footprint for a data collection process
113 that replicates the original structure of that provided by the USGS database, and follows a modified
114 approach suitable for the automated, near real-time workflows required by the LHASA model. This
115 includes estimating burn severity based on cloud-free composites derived from Landsat Surface
116 Reflection (SR) products (U.S. Geological Survey) based on the methodology proposed by Parks et
117 al. (2018) using the Google Earth Engine (GEE) platform (Gorelick et al., 2017). Burn severity is
118 quantified by the differenced or “delta” Normalized Burn Ratio (dNBR) (Key & Benson, 2006).
119 Higher dNBR values indicate higher burn severities, while lower or negative values indicate a low
120 severity burn or vegetation regrowth, respectively. For topographic assessment, we use the NASA
121 DEM 30m product as specified earlier.

122 The most significant departure from the original methodology of Staley et al (2016) is the use of
123 satellite derived rainfall intensity estimate data in contrast to the gauge data in the original database.
124 This also represents one of the major scientific contributions of our work, as we provide the first
125 comparison between satellite and gauge-based intensity thresholds for PFDF initiation in the
126 Western United States. Specifically, we collect 30min max intensity estimate data for each of the
127 storms in the recreated database from the Integrated Multi-Satellite Retrievals for GPM (IMERG)
128 Late Run v06B dataset (Huffman et al., 2019). This 30min peak intensity window is the highest
129 temporal resolution available to us. While 30min data is shown to be a relevant window for PFDF
130 hazard assessment (Liu et al., 2022), we note that PFDF initiation processes are frequently
131 documented to be more effectively delineated by peak intensities at smaller intervals such as 15min
132 or lower (e.g. Cannon et al., 2008; Kean et al., 2011; Staley et al., 2013; Staley et al., 2020). We

133 acknowledge this difference and provide a direct comparison of model performance in the following
134 section.

135 Following the collection of catchment burn severity, topography, and storm intensity characteristics,
136 we train our probabilistic model using the machine learning algorithm *XGBoost* (Chen & Guestrin,
137 2016)—a tree-based ensemble model which uses a series of shallow decision trees to iteratively
138 correct residual errors made by the trees before it. This choice contrasts with the logistic regression
139 model detailed in Staley et al. (2016); however, we select the XGBoost algorithm due to its non-
140 linear modeling capabilities and flexibility for setting monotonic constraints that dictate how the
141 probability of PFDF occurrence should correspond to an increase or decrease of a particular
142 variable. These constraints provide physically consistent results such that an increase in rainfall
143 intensity should indeed correspond to an increase in the probability of debris flow initiation. The
144 combination of these features allows for a robust, data driven modeling process to derive complex
145 physical relationships from potentially noisy observations.

146 To train and validate our model, we rely on a stratified K-Fold cross validation process (He & Ma,
147 2013) to minimize sampling bias and preserve original class balance. We also recreate the original
148 training/testing scenario documented by Staley et al (2016) for closer comparison between the two
149 models. This approach uses the data collected in California as a training dataset, and model
150 performance is tested on data collected in the Interior United States (IUS). For both training
151 scenarios we report model skill using the Area Under the Curve (AUC) metric, and the Threat Score
152 (TS) which served as the central evaluation metric in Staley et al. (2016), defined as:

$$153 \quad TS = \frac{\textit{True Positives}}{\textit{True Positives} + \textit{False Positives} + \textit{False Negatives}}.$$

154 A model with a high TS will demonstrate skill for correctly identifying the conditions most relevant
155 to PPDFs while minimizing classification error.

156 We finally demonstrate the potential for real time implementation with a case study to evaluate how
157 the model performs on data collected outside of the initial database. This process accomplishes the
158 important tasks of establishing that the model is adequately learning from the database itself, in
159 addition to evaluating model potential in burned steplands outside of the United States.

160 3. Results

161 Figure 2 provides a direct comparison between satellite and gauge derived intensity thresholds for
162 our available training data ($n = 627$). For IMERG-derived rainfall intensity estimates, the full

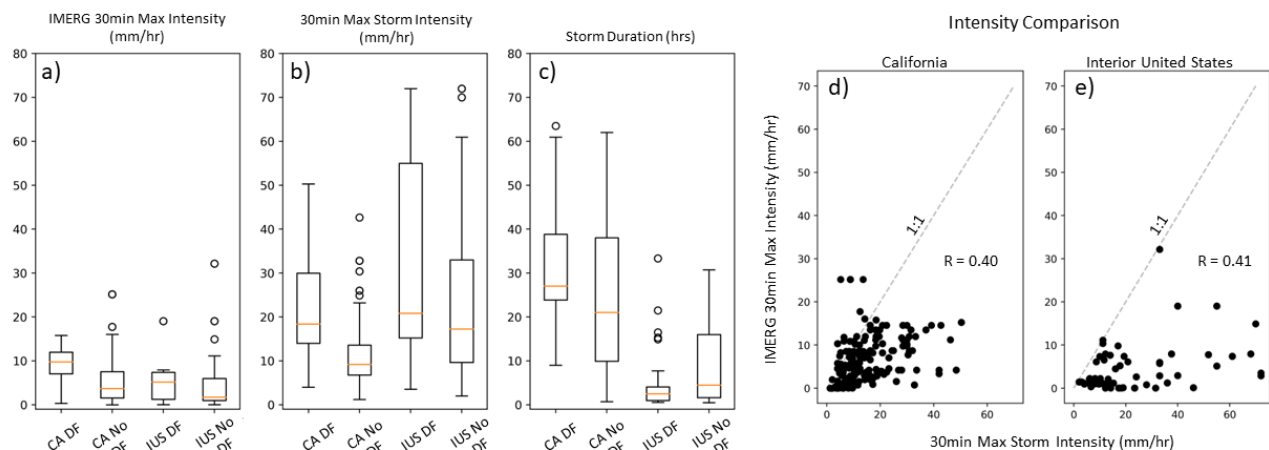


Figure 2. (a) IMERG peak 30min intensity values for DF and Non-DF events separated by region. (b) Gauge-recorded peak 30min intensity values for the same catchments. (c) Storm duration comparison for the intensity values in (a) and (b). Storm durations are notably longer in California than in the Interior United States, which is a function of local climatology (see Staley et al., 2020). (d & e) Direct comparison between gauge and IMERG Peak 30min intensities. Values are weakly correlated.

163 distributions of DF and Non-DF intensity groups are statistically different as reported by the Mann-
164 Whitney U test ($p \approx 0$). Despite strong class difference, IMERG 30min peak intensity values are
165 underestimated in both regions, with the highest errors associated with storms in the Interior United

166 States (IUS), where short duration, high intensity convective events are observed to be a frequent
167 mechanism for PFDF initiation (Cannon et al., 2008; Staley et al., 2020). Errors are lower in
168 California (CA) where PFDFs are typically—although not exclusively—induced from short bursts of
169 intense rainfall within longer duration atmospheric rivers (Kean et al., 2011; Oakley et al., 2017;
170 Staley et al., 2020). This includes PFDFs resulting from convection-driven high intensity rainfall
171 within a larger atmospheric river event (e.g., Oakley et al., 2018).

172 For both regions, storm duration plays an important role for effective precipitation estimate. The
173 IMERG algorithm uses a Kalman filter to combine the more accurate, yet periodic, passive
174 microwave (PMW) measurements with near constant infrared measurements from geostationary
175 satellites (GEO-IR) to provide globally available precipitation data at 0.1° resolution in 30min
176 intervals (Huffman et al., 2019). Product error is largely the result of the frequency of PMW
177 overpasses (~ 12 - 16 times daily for a given location in CONUS) as temporally propagated PMW
178 measurements are reported to be less correlated than calibrated GEO-IR after ± 90 minutes
179 (Huffman et al., 2019; Joyce & Xie, 2011; Khan & Maggioni, 2019). As such, short duration storms
180 may occur between PMW overpasses, leading to increased error. In contrast, longer storms allow for
181 increased PMW overpasses for improved interpolation between individual measurements.

182 For further comparison, recurrence intervals (RIs) are frequently used to normalize variable intensity
 183 duration thresholds across different climatic regions. Staley et al. (2020) provide a detailed analysis of
 184 the RIs of PFDf generating storms across a variety of climatic regimes. However, a direct
 185 comparison of their RI data with IMERG-derived intensity values is challenging due to the
 186 methodology and available gauge records used in the NOAA Atlas 14 products (e.g., Perica et al.,
 187 2013) referenced by Staley et al (2020), and the comparatively short ~20-year record of IMERG data
 188 available for manual RI calculation. Nonetheless, a comparison of RI data for both sources yields

<i>Standard Cross-Val</i>		Staley et al. (2016) Data	*This Study				
Features			# Model Runs	Mean AUC	Mean Threat Score	Train Size	Test Size
dNBR/1000	PropHM23	Peak 115 mm/hr	50	0.81	0.43	499	125
dNBR/1000	PropHM23	Peak 130 mm/hr	50	0.8	0.45	495	124
dNBR/1000	PropHM23	IMERG Max 30min Intensity*	50	0.77	0.40	498	125
dNBR/1000*	PropHM23*	IMERG Max 30min Intensity*	50	0.74	0.37	494	124
<i>Split: Train CA / Test IUS</i>							
Features				AUC	Threat Score	Train Size	Test Size
dNBR/1000	PropHM23	Peak 115 mm/hr		0.7	0.37	467	154
dNBR/1000	PropHM23	Peak 130 mm/hr		0.68	0.35	465	157
dNBR/1000	PropHM23	IMERG Max 30min Intensity*		0.65	0.32	469	154
dNBR/1000*	PropHM23*	IMERG Max 30min Intensity*		0.58	0.27	469	123

Table 1. Results from the cross-validation process. This includes training on subsets of data collected from both regions (CA and IUS), and testing on a group not included in the training process. This process repeats for 50 unique train/test split iterations that preserve the original class balance of the dataset. Results using the train/test split from Staley et al. (2016) are included for a comparison to those reported by the previous study. We include the variable names used in their database for easier reference. “dNBR/1000” refers to the catchment averaged Differenced (or Delta) Normalized Burn Ratio. “PropHM23” is the proportion of slope pixels >23° burned at moderate to high severity to all slope pixels >23° within the catchment.

189 values of <5yrs, which suggests agreement with the general frequency of these storms (Staley et al.
 190 2020) (See supplement).

191 We report model performance metrics in Table 1 and highlight the effect of IMERG 30min peak
192 intensity values on model performance when used as the exclusive source of rainfall intensity data in
193 the original USGS database. This provides the closest comparison for changes in model
194 performance and potential utility as a data source when it stands in for gauge-collected threshold
195 information. With both standard cross-validation and recreating the CA training/IUS testing split
196 from Staley et al. (2016), model performance decreases from threat score values of 0.43-0.45 to 0.4,
197 and 0.35-0.37 to 0.32, respectively. This suggests that while peak 30min intensity values are
198 underestimated by the IMERG data, the relative differences between DF and non-DF classes are
199 significant enough to build a model of comparable performance. Model results following this same
200 validation process using all globally available satellite derived inputs are provided in the last row of
201 each section in Table 1. Results are accordingly lower with the inclusion of burn severity and
202 topographic metrics following our modified database recreation workflow. Differences here are
203 largely the result of methodology: data provided by Staley et al. (2016) come from field-validated
204 burn severity class data provided by the Burned Area Emergency Response (BAER) program. In
205 contrast, we modify the methodology provided by Parks et al (2018) to rapidly approximate soil
206 burn severity conditions using dNBR values provided by Landsat imagery. It is thus reasonable to
207 assume that field validated burn severity data lead to a stronger model and Table 1 highlights the
208 tradeoff in model skill between on-the-ground assessments and a fully automated method at the
209 global scale.

210 While these results document our model performance within the scope of the original database, it is
211 important to understand how the proposed framework would perform in the context of a near real
212 time system, and especially on observations outside of the original training database within the
213 United States. To demonstrate its applicability, we highlight model performance on data collected
214 following the 2009 Beechworth Fire in Victoria, Australia, where Nyman et al. (2011, 2019) reported

215 evidence of PFDF activity following heavy rainfall. The exact timing of the storms that initiated the
216 majority of debris flows remain unreported; however, local news ties at least one PFDF inducing
217 storm event to March 14th, 2009 which also led to a road closure (Dulhunty, 2009). As such, we map
218 the location of a proximal debris flow following the wildfire recorded by Nyman et al. (2011) and
219 reference it with the timing and nearby location of the events documented by Dulhunty (2009). We
220 then apply the same modified data collection process used for the catchments referenced in the
221 USGS database to gather the relevant catchment/storm characteristics for this event. The final
222 model is trained on the fully recreated USGS database to make equivalent assessments for this area.

223 Figure 3 highlights PFDF probability as a function of daily IMERG max 30min rainfall intensity
224 estimates from the approximate time of fire conclusion to several days past the time of recorded
225 PFDF activity. As burn severity and slope properties are held constant during this time, changes in
226 probability are dependent on rainfall. Notably the largest spike in model probability occurs during
227 the same day which PFDF activity is recorded. Model probabilities remain consistently low for the
228 rest of the days included in this time series. The threshold line at 31% in Figure 3 represents the
229 mean probability for DF classification which maximized the Threat Score over 50 cross-validated
230 model runs. In the event that the recorded debris flow occurred during this time, a system using this
231 threshold would have correctly identified the PFDF in this area while avoiding false positives or
232 false negatives during the prior month.

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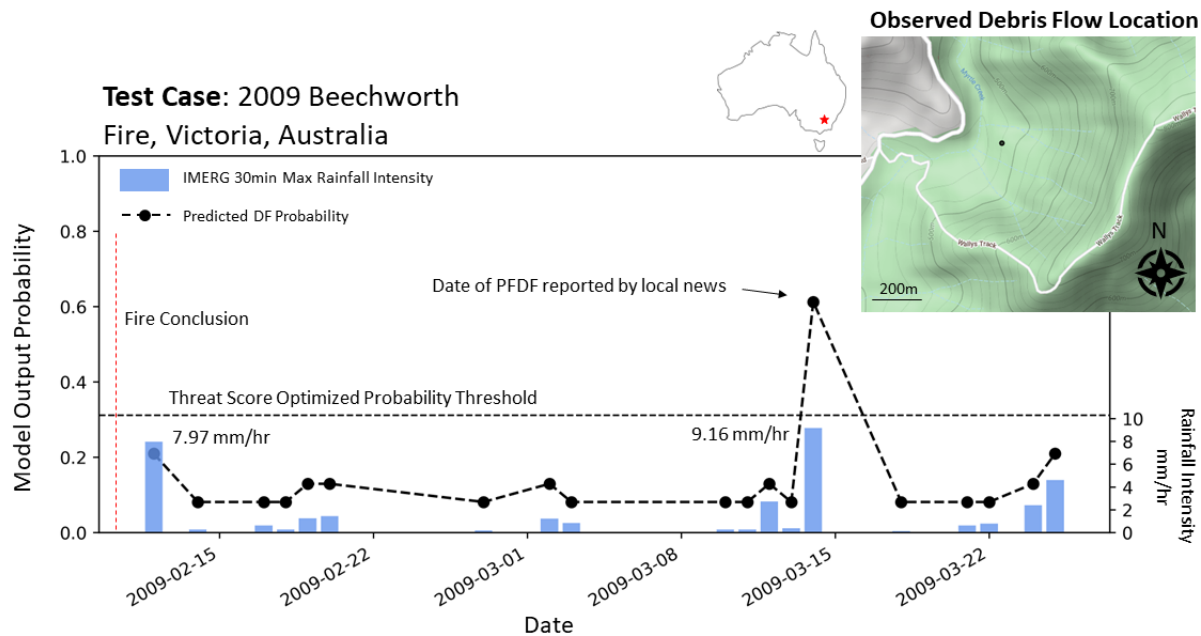


Figure 3. Daily PFDF probability as a function of IMERG-recorded 30min peak intensity rainfall for that day. Our model demonstrates nonlinear sensitivity to rainfall with a large jump in probability above 8.22 mm/hr, which is the median DF-inducing rainfall intensity value from our training dataset. The magnitude of this spike is nonetheless dependent on the burn and slope characteristics of the catchment (i.e., lower slope and burn severity values lead to probability increases at lower rates). Gauge-recorded accumulation values for this event vary between 12-40mm total (Dulhunty 2009). As the exact timing of this event was inferred from reporting by Nyman et al (2011, 2019) and Dulhunty (2009), there remains some uncertainty as to whether the debris flow occurred during the recorded storm time. Inset basemap image provided by Google Earth Engine.

234

235 4. Discussion

236 This work documents the development of a probabilistic PFDF initiation model that uses publicly
 237 available satellite-derived inputs. It is designed to operate in near real time and at the global scale. A
 238 significant difference between the proposed framework and that of others (e.g. Staley et al. 2016) is
 239 the intended use. Previous studies have focused on providing field-validated measurements of burn
 240 severity and gauge-collected rainfall in the Western United States to provide detailed insight into the
 241 physical constraints of PFDF initiation across spatially heterogeneous terrain. Comparatively, our

242 proposed framework employs an analogous remote sensing-based methodology that can provide
243 routine situational awareness of these hazards. It does so by assessing the relevant combinations of
244 slope, burn severity, and rainfall that may lead to the heightened likelihood of PFDF activity based
245 on empirically correlated relationships of these characteristics derived from the Western United
246 States. It appropriately demonstrates varied sensitivity to rainfall as a function of burn severity and
247 slope parameters such that it produces reasonable model performance for its intended use cases.

248 As PFDFs are documented outside of the US (e.g. García-Ruiz et al., 2013; Jin et al., 2022; Nyman
249 et al., 2019), model assessments to catchments on the global scale will represent an exciting
250 opportunity for future research. Additional PFDF inventories would provide further insight as to
251 which catchment characteristic thresholds are most relevant for PFDF activity in burned areas
252 worldwide, with a strong focus on understanding which catchments simply do not have the
253 appropriate characteristics to facilitate PFDF initiation. This is specifically an issue highlighted by
254 Parise and Cannon (2008, 2012), who speculate that the lack of significant activity in other
255 Mediterranean ecosystems might be a result of a combination of less severe fires, differences in
256 watershed morphology, and/or patterns in rainfall. The proposed framework provides an
257 opportunity to quantitatively evaluate differences between these regions and those where PFDFs are
258 more commonly observed. Similarly, each false alarm or miss provides valuable information that
259 represents an opportunity to improve future hazard estimates. We note that the validation of future
260 work may be limited to observations near populated areas—as debris flows occurring in remote
261 regions often go unrecorded—but hope that this system will bring more attention to the locations
262 and frequencies of these events where they are observed.

263 Despite adequate model performance, triggering rainfall intensity estimates in the IMERG dataset
264 are indeed lower than their gauge-recorded counterpart. This behavior is expected. For instance,

265 Khan et al. (2016) record lower IMERG performance in continental climates with dry summers
266 compared to those with year-round precipitation. Further, performance of IMERG data has been
267 extensively studied in several locations worldwide, where IMERG typically underestimates high
268 intensity rainfall in complex terrain (e.g., Lu et al., 2019; Maggioni et al., 2017; Mayor et al., 2017;
269 Nascimento et al., 2021; Navarro et al., 2020; Nepal et al., 2021; Rojas et al., 2021). Decreased
270 IMERG performance due to orography is a known issue and remains an area for future work (Tan
271 et al., 2019). Despite these limitations, recent versions of IMERG still help characterize patterns of
272 orographic rainfall (e.g., Rojas et al., 2021; Sharma et al., 2020) and precipitation patterns on daily to
273 seasonal timescales (e.g., Mayor et al., 2017; Nascimento et al., 2021; Nepal et al., 2021). Figure 2
274 demonstrates that PFDF triggering intensities are indeed higher than the peak intensities for the
275 storms during which no PFDFs are recorded. This suggests that while the absolute magnitude of
276 DF-inducing storm intensities will not be accurately captured, IMERG can differentiate these
277 common—yet still significant—storms from the events during which PFDF initiation is statistically
278 less likely.

279 **5. Conclusion**

280 We introduce a model for PFDF hazard assessment capable of near real time observations at the
281 global scale. This model is trained on a modified database of historic PFDF activity and relies on
282 globally available data products to increase situational awareness of these hazards globally. While the
283 model does not increase lead-time for hazard warning, it is a resource for situational awareness
284 where gauges and/or radar are not available. As our retrospective analysis demonstrates, this model
285 can differentiate between DF and Non-DF inducing peak rainfall intensities across variable
286 topographic and climatic regimes—even when hazard estimates are limited by the underestimations
287 of intensity from IMERG data. There are still components we wish to incorporate into future

288 iterations of the model, which include an updated training database with more events outside of the
289 United States, and more recent data products such as IMERG Version 7 and Landsat 9. Once
290 implemented, this model will provide a framework for near real time hazard assessments of PFDF
291 processes on a global scale.

292 Acknowledgements

293 This research was supported by NASA's Disasters program through the solicitation for Earth
294 Science Applications: Disaster Risk Reduction and Response (NNH18ZDA001N). We would like to
295 thank Dennis Staley, Jason Kean, Matt Thomas, Francis Rengers and three anonymous reviewers for
296 their valuable feedback on this work. We gratefully acknowledge the data provided by the USGS,
297 which served as the basis for our model training, cited as Staley et al. (2016). Basemap data for
298 Figures 1b and 3 provided by Stamen Design, under CC BY 3.0 (Data by OpenStreetMap, under
299 ODbL) and Google Earth, respectively.

300 Data Availability Statement:

301 A developmental branch of the operational component of this work is available at the following
302 Github repository: <https://github.com/nasa/LHASA/tree/master/pfdf>. Training data provided by
303 the USGS is available at the following link: <https://doi.org/10.3133/ofr20161106>. Model training
304 files are available at <https://doi.org/10.5281/zenodo.7058363>.

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