

An agenda for land data assimilation priorities: Realizing the promise of terrestrial water, energy, and vegetation observations from space

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Key Points:

- Land data assimilation has shown significant promise for short-term forecasting applications

- 30 • Significant gaps remain in the current land data assimilation systems related to
31 observation utilization and models
- 32 • Coordinated development with the modeling and observational community and
33 adoption of technological enhancements are needed in the future

34 **Abstract**

35 The task of quantifying spatial and temporal variations in terrestrial water, energy, and
36 vegetation conditions is challenging due to the significant complexity and heterogene-
37 ity of these conditions, all of which are impacted by climate change and anthropogenic
38 activities. To address this challenge, Earth Observations (EOs) of the land and their
39 utilization within data assimilation (DA) systems are vital. Satellite EOs are partic-
40 ularly relevant, as they offer quasi-global coverage, are non-intrusive, and provide uni-
41 formity, rapid measurements, and continuity. The past three decades have seen un-
42 precedented growth in the number and variety of land remote sensing technologies
43 launched by space agencies and commercial companies around the world. There have
44 also been significant developments in land modeling and DA systems to provide tools
45 that can exploit these measurements. Despite these advances, several important gaps
46 remain in current land DA research and applications. This paper discusses these gaps,
47 particularly in the context of using DA to improve model states for short-term numer-
48 ical weather and sub-seasonal to seasonal predictions. We outline an agenda for land
49 DA priorities so that the next generation of land DA systems will be better poised to
50 take advantage of the significant current and anticipated shifts and advancements in
51 remote sensing, modeling, computational technologies, and hardware resources.

52 **Plain Language Summary**

53 Satellite remote sensing measurements have enabled the monitoring of the Earth's
54 land surface with unprecedented scale and frequency. These measurements allow us
55 to monitor the changes on the land surface and understand the contribution of hu-
56 man activities toward them. The information from such observations is combined with
57 the modeled estimates through data assimilation (DA) algorithms. This article discusses
58 the progress made in the development of land DA systems and the major gaps that
59 remain. The paper also outlines priorities that we need to consider in the develop-
60 ment of next generation land DA systems so that the potential of land remote sens-
61 ing measurements can be fully realized.

1 Introduction and premise

Warmer global temperatures associated with climate change have been shown to lead to more uncertainties in water availability both regionally and globally (Barnett et al., 2005; Milly et al., 2005; Berg et al., 2016; Konapala et al., 2020; Wu et al., 2020). There is increasing evidence for an uneven distribution of changes in key water flux rates such as precipitation and evaporation leading to more intense and frequent extreme events, with the lack of predictability in water availability exacerbating water scarcity (Oki & Kanae, 2006; Trenberth et al., 2014; Cook et al., 2018), water excess and natural disasters. The scientific literature shows that changes in atmospheric moisture transport will lead to consequences such as longer droughts in arid and semi-arid regions, more intense precipitation and snowfall, and earlier snow melt, among other changes (Trenberth et al., 2003; J. Huang et al., 2016; Ingram, 2016; Wainwright et al., 2021). Additionally, anthropogenic factors such as agricultural water use and industrial expansion, reservoir management, deforestation, and wetland drainage have not only quickened the pace of land transformation, but have also contributed to, often unpredictable, changes in the water, energy, and carbon cycles (Wood et al., 1997; McDonald et al., 2011; Vorosmarty & Sahagian, 2000; Haddeland et al., 2014; Mehran et al., 2017). Given the urgent need to understand the variability in water cycle fluxes, there have been several efforts to develop inventories of key water cycle fluxes (Trenberth et al., 2007; L'Ecuyer et al., 2015; Rodell et al., 2015; Vargas Godoy et al., 2021). The combined use of modeling and remote sensing estimates of water and energy states and fluxes is a key feature of these studies, which typically report closure of water and energy budgets with relatively small residuals and uncertainty ($\sim 4-8\%$) at coarse (monthly or larger) temporal scales. However, these studies also note that the imbalances and closure errors increase with increasing spatial and temporal resolution; consequently, concurrent high-quality Earth observations are needed to further constrain the modeled estimates and enable reliable assessment of water, energy, and carbon cycle budgets.

Equally important is the vegetation state, which is also a key factor in the coupling of the water, energy, and carbon cycles. Vegetation exerts key controls on surface water and energy partitioning and the associated response of terrestrial water, energy, and carbon fluxes to climatic drivers (Powell et al., 2013). Changes in vegetation impact surface albedo and surface roughness, interception of precipitation, and

95 reduction in surface runoff (L. Zhang et al., 2015) and can explain up to 30% of the
96 variance in precipitation and surface radiation (Green et al., 2017). Accurately repre-
97 senting vegetation processes in land surface models (LSMs) is thus critical for captur-
98 ing the efficiency of carbon uptake and transpiration and – more importantly – how
99 this efficiency is expected to change in a future climate.

100 Despite the importance of biospheric processes, their representation in many LSMs
101 does not accurately capture observed vegetation behavior. The modeled biosphere-
102 atmosphere coupling is much weaker than in observations, in part because the mod-
103 els generally underestimate the observed response of the biosphere to climate (Green
104 et al., 2017; Papagiannopoulou et al., 2017). This can lead to substantial biases in near-
105 surface states, such as air temperature and relative humidity, and can contribute to
106 the misrepresentation of soil moisture dynamics (Koster et al., 2009) – problems that
107 persist at the monthly and longer timescales associated with such vegetation anoma-
108 lies.

109 In addition to their role in hydrological and climate prediction, coupled land surface-
110 atmosphere models are integral to numerical weather prediction (NWP) and sub-seasonal
111 to seasonal (S2S) forecasts, with their ability to represent heterogeneity in land use,
112 soil type, vegetation types, soil moisture, the presence of snowpack, and the impact
113 of these factors on the land-atmosphere exchange of energy and water. For example,
114 through the soil-plant-atmosphere continuum, the land surface conditions can con-
115 trol the surface heat fluxes (both latent and sensible) with consequent impacts on air
116 temperature, the height and stability of the atmospheric boundary layer, and occur-
117 rence of precipitation (Eltahir, 1998; Dirmeyer & Halder, 2016). This means land sur-
118 face initialization can influence NWP and S2S forecast performance when the land-
119 atmosphere coupling is strong and when there is sufficient variability, and also sig-
120 nificant memory in land surface states (Koster et al., 2004, 2010, 2021)). Over recent
121 years, regions of strong land-atmosphere coupling have been shown to have expanded,
122 to now cover regions that never before exhibited strong land-atmosphere coupling (Dirmeyer
123 et al., 2021). For example, soil moisture conditions have been shown to directly im-
124 pact temperature and precipitation forecast skill, provided that the forecast model prop-
125 erly represents the processes linking land anomalies, surface fluxes, and atmospheric
126 physics. Models with incorrect land initialization and/or poor land-atmosphere cou-
127 pling behavior will have biases in temperature, humidity, and precipitation forecasts

128 during periods of drought (Dirmeyer & Halder, 2016; Dirmeyer, 2018). Observational
129 constraints are needed to improve such deficiencies (Dirmeyer et al., 2016).

130 Data assimilation (DA) is the typical approach used in Earth system modeling
131 to combine information from observational data sources with dynamical process rep-
132 resentations in numerical process models to develop estimates that are superior to those
133 obtained by using just the data or model alone (Reichle, 2008a). The resulting estimate
134 from DA is typically called the ‘analysis’ and is calculated according to the relative
135 errors and uncertainties in the model and the observations. There is a wide variety
136 of methods that are employed to generate such DA analyses, with the mathematical
137 objective of minimizing the posterior error of the DA analysis, whilst being constrained
138 by the available prior information about the system being modeled. The majority of
139 such methods are founded in optimal control or Bayesian estimation theory (Nichols,
140 2010), which incorporates observations into models in a sequential or non-sequential
141 manner (Bouttier & Courtier, 2002). While the atmosphere and ocean modeling com-
142 munities (Ghil & Malanotte-Rizzoli, 1991; Malanotte-Rizzoli, 1997; Navon, 2009; La-
143 hoz & Schneider, 2014) have led the charge in the methodological development and
144 applications of DA, it is being increasingly adopted in land surface and hydrology mod-
145 eling, particularly during the past two decades (Reichle et al., 2009; Houser et al., 2012;
146 de Rosnay et al., 2014).

147 Given the importance of land surface initialization and land-atmosphere coupling
148 to applications such as NWP and S2S forecasting, it is perhaps surprising that devel-
149 opment of LSMs, and land-focused DA activities in particular, have been a relatively
150 low priority for NWP and S2S development in the past. This prioritization is perhaps
151 a reflection of how land surface modeling has developed in relation to atmospheric
152 modeling. In an editorial piece in the American Geophysical Union’s newsletter “EOS”,
153 Dirmeyer (2018) addressed this question directly, suggesting that a lack of observa-
154 tions of land surface states or land-atmosphere fluxes resulted in the development of
155 land modeling schemes “to ameliorate systematic atmospheric model errors... addressing one
156 error by introducing additional ones” (Dirmeyer, 2018). In such a circumstance, the real-
157 istic representation of land surface states, as obtained from land DA, may lead to only
158 minimal or disadvantageous impacts on atmospheric forecasts (Crow et al., 2020). This
159 inconsistency was more easily overlooked when high-quality observations of land states
160 were not available. However, the latest generation of satellite and in-situ observations

161 now allow for more rigorous verification of process-model improvements and initial-
162 ization of forecasts.

163 In recent decades, there has been a tremendous expansion in the number of re-
164 mote sensing platforms that provide observations of the land surface (Kimball, 2008;
165 Lakshmi, 2014; McCabe et al., 2017; Balsamo et al., 2018). Over the years, monitoring
166 capabilities from these platforms have encompassed a broad range of land surface/hydrology
167 variables including land use, soil moisture, snow cover, snow water equivalent, land
168 surface temperature, vegetation change/dynamics, terrestrial water storage, and sur-
169 face water. The technological and scientific success of these spaceborne platforms have
170 now yielded multidecadal observations for many variables (e.g., terrestrial water stor-
171 age, soil moisture, leaf area index, snow, land surface temperature) enabling the pro-
172 duction of long-term climate data records (Lettenmaier et al., 2015). Similarly, for some
173 variables such as vegetation dynamics, soil moisture and precipitation, observations
174 are now provided by multiple platforms, allowing the possibility of increased spatio-
175 temporal coverage from remote sensing. More recent advances in land remote sens-
176 ing have also enabled complementary remote sensing observations of key water and
177 carbon cycle variables, e.g. precipitation from the Global Precipitation Measurement
178 (GPM; Hou et al. (2014)) mission, surface soil moisture from Soil Moisture Active Pas-
179 sive (SMAP; Entekhabi et al. (2010a)) and Soil Moisture Ocean Salinity (SMOS; Kerr
180 et al. (2010)) missions, terrestrial water storage from Gravity Recovery and Climate
181 Experiment Follow-On (GRACE-FO; Flechtner et al. (2014); Kornfeld et al. (2019)) mis-
182 sion, Solar-induced chlorophyll fluorescence from the Orbiting Carbon Observatory-
183 2 (OCO2; Sun et al. (2017)) mission, and greenhouse gases observing satellite (GOSAT;
184 Doughty et al. (2022)).

185 Together, these observations facilitate the development of observation-only based
186 water, energy, and carbon budget closure estimates (Trenberth et al., 2014). In addi-
187 tion to conventional space agency driven science missions, sensing platforms from com-
188 mercial companies have also started to become more commonplace (Dash & Ogutu,
189 2016; Houborg & McCabe, 2016; Tollefson, 2017). There has been a growing recogni-
190 tion in the community that modeling systems that can effectively exploit such multi-
191 sensor data are key to answering fundamental questions about global and regional
192 water, energy, and carbon cycle changes (Durand et al., 2021). In particular, it is im-
193 perative that DA systems be designed to take advantage of this presumably “golden

194 age” of land surface remote sensing (Stavros et al., 2017; Sellars et al., 2013; Schimel
195 et al., 2019) and be designed and positioned to take advantage of upcoming missions
196 as soon as data become available, such as the surface water storage from the upcom-
197 ing Surface Water Ocean Topography (SWOT; Biancamaria et al. (2016)) mission, Sur-
198 face Biology and Geology-designate observable (Cawse-Nicholson et al., 2021) or ESA’s
199 BIOMASS (Quegan et al., 2019).

200 The primary objective of this commentary article is to provide an agenda for the
201 next generation land DA systems given the significant current and anticipated shifts
202 and advancements in remote sensing, modeling, computational technologies, and hard-
203 ware resources. Note that this article discusses these priorities primarily in the con-
204 text of space-borne remote sensing observations of the land surface. Accurate land char-
205 acterization has major impacts on atmospheric forecasts, water availability quantifi-
206 cation, management of water and agricultural resources, and developmental issues
207 facing the World’s population growth (Collins et al., 2013). It is also critical for cap-
208 turing the long-term carbon-cycle feedbacks of the biosphere to climate and meteo-
209 rology and thus making more accurate projections of vegetation, water, and energy
210 states with elevated CO₂ conditions (Canadell et al., 2021). Effective DA systems have
211 a critical role to play in realizing the full information potential and promise of Earth
212 science remote sensing measurements. The specific goals of this agenda paper are to:

- 213 • Define the most significant gaps in and priorities for the development of regional
214 to global scale land DA systems for water, energy, and vegetation predictions
215 at NWP and S2S lead times.
- 216 • Identify priorities for coordination between the modeling, remote sensing, and
217 technology communities to improve the effectiveness of land DA.
- 218 • Provide a framework for future enhancements of DA systems.

219 Note that the present paper focuses on DA in the context of state estimation. Re-
220 mote sensing observations can also be used to reduce errors associated with model
221 parameters and structure. Challenges and priorities related to LSM calibration and pa-
222 rameter estimation, with a particular focus on coupled carbon-vegetation-water in-
223 teractions related to longer-term biosphere-climate feedbacks, will be addressed in a
224 separate companion paper (Quaife et al., in preparation). In parallel, a perspective is

225 offered (De Lannoy et al., 2022) on how to address the increased complexity of land
226 DA for either state or parameter estimation (or both) in the future.

227 **2 Background**

228 As described in Reichle (2008a), the primary objectives of DA include extend-
229 ing the spatio-temporal coverage of infrequent observational data, science translation
230 of raw measurements from satellites (or other data sources), and optimal merging of
231 information from multiple sources while preserving the basic constraints on the phys-
232 ical system. By incorporating observational inputs, DA also helps in reducing the un-
233 certainty of estimated land states and fluxes within Earth systems models. This is par-
234 ticularly important given the proliferation of models and modeling systems, leading
235 to large uncertainties purely from differences in conceptual formulations and param-
236 eterizations (Clark et al., 2017). The observational constraints from DA are a practi-
237 cal approach to reduce such uncertainties, so that reliable estimates of the water and
238 carbon cycles, and their feedbacks, can be developed.

239 Taking advantage of the increased availability of environmental datasets, atmo-
240 spheric DA has played fundamental roles in steadily improving regional and global
241 environmental prediction capabilities (Benjamin et al., 2018; Hersbach et al., 2020). In
242 particular, the prediction of precipitation, short-term extreme events, and applications
243 of nowcasting (Kain et al., 2010; Mass, 2012) have hugely benefitted from the advances
244 in DA techniques to the point that DA is considered an integral component of NWP
245 systems (Kalnay, 2003). Similarly, DA has been key in global and regional ocean pre-
246 diction and reanalysis systems (Moore et al., 2019; Zuo et al., 2019). These commu-
247 nities have also pioneered the development of the majority of variational and ensem-
248 ble DA techniques used in operational and reanalysis environments. These method-
249 ological advances have been adapted by the land/hydrology/biogeochemistry com-
250 munities to develop similar improvements in the land surface water and carbon cy-
251 cle predictions (Reichle, 2008a; Rayner, 2010; de Rosnay et al., 2014; Lahoz & De Lan-
252 noy, 2014; Rayner et al., 2019; Fairbairn et al., 2019; Smith et al., 2020).

253 In the past several decades, there has been a significant increase in the devel-
254 opment and application of DA to LSMs, with the majority using sequential filtering
255 approaches. The use of variational DA approaches, which are popular in the atmo-

spheric and ocean modeling communities, is less common in NWP and S2S land DA applications. Variational DA approaches are more widely used for optimizing vegetation and carbon cycle related parameters LSMs, with the goal of better estimating longer-term (annual to centennial) carbon stocks and fluxes (Kaminski et al., 2013; Pinnington et al., 2020; MacBean et al., 2022), although these approaches have also been used to improve LSM soil moisture predictions (Scholze et al., 2016; Pinnington et al., 2018; Raoult et al., 2021). Variational methods provide updates - by minimizing a specified cost function - that constrain the model forecasts with a set of observations. They require the use of linearized observation operators, and in some cases, a tangent linear or adjoint model that represents the sensitivity of model states backward in time (Courtier et al., 1993). Some variational applications exist for land DA that use tangent linear or adjoint models (e.g. Peylin et al. (2016); Raoult et al. (2016); Schurmann et al. (2016)) or adopting a linearity assumption (Hess, 2001; Balsamo et al., 2004) but, in general, the development of such models for the land surface has been more challenging due to the tendency for LSMs to contain non-differentiable threshold processes (Reichle et al., 2002a). In variational DA systems, the temporal evolution of the background error covariance is typically not modeled, whereas Kalman and particle filtering approaches allow dynamic updates of these estimates (Rabier & Liu, 2003).

Variants of Kalman filters are often used in sequential DA systems. The Extended Kalman Filter (EKF), the nonlinear version of the Kalman filter, is used in some land surface DA environments (Rudiger et al., 2010; Ghent et al., 2010; de Rosnay et al., 2013), by employing linearized observation operators and error variance propagation formulations (Reichle et al., 2002a). The ensemble Kalman filtering (EnKF) methods (Evensen, 2003), which are Monte-Carlo approximations of the Bayesian filtering process, provide a more convenient approach that is suitable for weakly non-linear models. In addition, due to its Monte Carlo formulation, the EnKF is easy to implement and provides a convenient means for representing the wide variety of errors impacting LSMs. The use of an ensemble of model realizations in the EnKF also allows for the implicit propagation of model errors, making it more computationally tractable and scalable for large systems. As a result, the EnKF has become widely used in the land surface community. Non-sequential extensions of the EnKF that allow information updates backwards over a time window have also been developed (Evensen & Leeuwen, 2000). Detailed methodological descriptions of various DA approaches are discussed in sev-

289 eral prior review papers, including (Reichle, 2008a; van Leeuwen, 2009; Lahoz & De Lan-
290 noy, 2014; Lahoz & Schneider, 2014; Houtekamer & Zhang, 2016; van Leeuwen et al.,
291 2019).

292 Land DA applications have been examined for a wide range of land surface wa-
293 ter, energy, and carbon variables. Given the developmental legacy of LSMs as the bound-
294 ary conditions to atmospheric models, a majority of the land DA efforts are focused
295 on constraining variables such as soil moisture and snow that have relatively long-
296 memory and a direct impact on the NWP initialization skill (de Rosnay et al., 2014;
297 Santanello et al., 2019; Gomez et al., 2020; Benjamin et al., 2022). Many NWP centers
298 currently employ some form of screen level assimilation of synoptic measurements
299 of air temperature and relative humidity close to the surface (Mahfouf, 1991), updat-
300 ing soil moisture states within DA. Motivated by the need for improved land char-
301 acterization for short-term forecasts and hydrology applications, a wide variety of stud-
302 ies have focused on the assimilation of more direct information of soil moisture from
303 ground and spaceborne platforms, with a range of DA techniques (Houser et al., 1998;
304 Margulis et al., 2002; Parajka et al., 2006; Drusch, 2007; Sabater et al., 2007; Kumar et
305 al., 2012; Draper et al., 2012; Renzullo et al., 2014; Lievens et al., 2015; De Lannoy &
306 Reichle, 2016a; Lievens et al., 2016; Kolassa et al., 2017a, 2017b; Lievens et al., 2017;
307 Reichle et al., 2019; Rodriguez-Fernandez et al., 2019; De Lannoy et al., 2019). These
308 studies encompass the assimilation of soil moisture measurements and retrievals from
309 passive microwave satellite platforms ranging from Special Sensor Microwave Imager/Sounder
310 (SSMIS) in the early 1990s to more recent missions such as SMOS and SMAP that
311 rely on L-band microwave sensors. While less numerous, other studies have focused
312 on the assimilation of root-zone soil moisture proxy information derived from thermal-
313 infrared remote sensing (e.g., Hain et al. (2012)).

314 In addition to soil moisture, assimilation of snow cover and snow water equiv-
315 alent (SWE) relevant measurements have also been examined with approaches rang-
316 ing from direct insertion to variants of Kalman filtering and smoothing. Snow cover
317 observations from optical sensors, which do not directly provide quantitative infor-
318 mation on the amount of snow in the snowpack are often assimilated using rule-based
319 update methods (Rodell & Houser, 2004; Zaitchik & Rodell, 2009; Arsenault et al., 2013;
320 De Lannoy et al., 2012). A few studies have also used the EnKF to assimilate snow
321 cover observations (Clark et al., 2006; H. Su et al., 2008; Y.-F. Zhang et al., 2014; Toure

322 et al., 2018). These studies typically demonstrate modest improvements for estimat-
323 ing ephemeral SWE and little benefit over deeper snowpacks, owing to the fact that
324 the information about SWE in instantaneous snow cover observations tends to dimin-
325 ish for larger values of SWE, as snow cover saturates at 100%. To address the weak
326 instantaneous correlation between snow cover and SWE, a number of studies (Durand
327 et al., 2008; Giroto et al., 2014; Margulis et al., 2015; Oaida et al., 2019) employed smooth-
328 ing approaches that assimilate all snow cover observations within a time window. Es-
329 timates of SWE from ground measurements, microwave, and lidar platforms have also
330 been widely used for DA in snow hydrology studies (Andreadis & Lettenmaier, 2006;
331 Slater & Clark, 2006; Liu et al., 2013; Magnusson et al., 2014; Margulis et al., 2015; Ku-
332 mar et al., 2015a; C. Huang et al., 2017; Margulis et al., 2019; Smyth et al., 2019). Ef-
333 forts to assimilate brightness temperature observations from microwave instruments
334 and radar backscatter measurements using machine learning and radiative transfer
335 model operators for improving snow estimation have also been reported over differ-
336 ent regions of the world (Phan et al., 2014; Xue & Forman, 2017; Xue et al., 2018; Kwon
337 et al., 2019; Park et al., 2021). Consequently, some operational centers around the world
338 have enabled a form of land analysis to incorporate remote sensing information for
339 NWP initialization (Barrett, 2003; Pullen et al., 2011; de Rosnay et al., 2013, 2014; Car-
340 rera et al., 2019; Wegiel et al., 2020; Gomez et al., 2020). These analyses also range from
341 optimal interpolation to Kalman filter methodologies.

342 In addition to work within NWP and with offline LSMs, there exists a large set
343 of applications that rely on accurate estimates of land surface conditions. Water re-
344 source planning efforts require reliable estimates of water fluxes and their variabil-
345 ity to account for water availability issues and hydrological extremes. Consequently,
346 these inputs are also becoming critically important for agricultural and food security
347 assessments. Military agencies rely on accurate land characterization of surface states
348 for mobility and trafficability assessments. Water storage in peatlands controls large
349 carbon fluxes to the atmosphere. Estimates of river flow are important for reservoir
350 management as well as transboundary water applications. Given these significant needs
351 for land surface hydrology information, there have been land-focused systems (some-
352 times called 'offline' environments) developed in the past several decades. The Data
353 Assimilation Research Testbed (DART; Anderson et al. (2009)), developed at the Na-
354 tional Center for Atmospheric Research has enabled ensemble DA for the land sur-

355 face schemes of several Earth system models. The Global Land Data Assimilation Sys-
356 tem (GLDAS; Rodell et al. (2004)) and the North American Land Data Assimilation
357 System (NLDAS; Mitchell et al. (2004)) were the two pioneering efforts that established
358 software frameworks for driving LSMs with observationally constrained meteorology,
359 with the eventual goal of enabling the assimilation of land surface states. Other ex-
360 amples of such land-only simulation products include MERRA-Land (Reichle et al.,
361 2011), ERA-Interim/Land (Balsamo et al., 2015), and ERA5-Land (Munoz-Sabater et
362 al., 2021), which supplement major atmospheric reanalysis products. Another land-
363 only product is LDAS-Monde (Albergel et al., 2017, 2020), which jointly assimilates
364 surface soil moisture and leaf area index. Moreover, the land-only SMAP Level-4 Soil
365 Moisture product focuses on the assimilation of SMAP radiance observations (Reichle
366 et al., 2019). A flexible environment called the Land Information System (LIS; Kumar
367 et al. (2006)) was developed at NASA that allows for the interoperable use of multi-
368 ple land and hydrology models and formal DA capabilities around them (Kumar et
369 al., 2008). The capabilities of LIS are not only used in both GLDAS and NLDAS, but
370 have also fostered many LDAS configurations and instances in different parts of the
371 world (Goncalves et al., 2009; McNally et al., 2017; Kumar et al., 2018; Erlingis et al.,
372 2021). The development of similar LDAS environments has also been reported in other
373 organizations around the world (Lewis et al., 2012; Reichle et al., 2014; Sawada et al.,
374 2015; Carrera et al., 2015; Albergel et al., 2017, 2020).

375 The development of land-only DA environments has enabled the assimilation
376 of other land hydrology variables, including some that are not viable for operational
377 NWP needs because of their long latency. For example, Terrestrial Water Storage anoma-
378 lies from gravity missions such as GRACE have been used in DA configurations to
379 develop inferences on subsurface changes (Zaitchik et al., 2008; Syed et al., 2008; B. Li
380 et al., 2012; Kumar et al., 2016; Giroto et al., 2017; Zhao & Yang, 2018) The use of GRACE
381 data has provided unprecedented information about the changes in Earth's ground-
382 water storage and has uncovered unsustainable exploitation from groundwater pump-
383 ing in many areas of the world (Rodell et al., 2009, 2018; Lo et al., 2016; Thomas &
384 Famiglietti, 2019; Nie et al., 2019). The use of GRACE-based constraints also has en-
385 abled critical inputs for drought monitoring efforts, by providing an observation-informed
386 estimate of groundwater changes, a critical water resource that varies at longer timescales
387 than surface water (Houborg et al., 2012; B. Li et al., 2019; Getirana et al., 2020).

388 With the availability of prognostic representations of carbon processes within the
389 third and fourth generation LSMs (Pitman, 2003; Fisher & Koven, 2020), there have
390 been several efforts to utilize remotely sensed vegetation datasets through DA (e.g.
391 in the aforementioned LDAS-Monde system). There is a long legacy of high-resolution
392 measurements of canopy states from optical sensors such as leaf area index, normal-
393 ized difference vegetation index, fraction of photosynthetically active radiation, and
394 vegetation biomass. Assimilation of leaf area index to improve the estimation of crop
395 yields, vegetation biomass, root zone soil moisture, and carbon fluxes has been demon-
396 strated in several prior studies (Demarty et al., 2007; Quaife et al., 2008; Dente et al.,
397 2008; Sabater et al., 2008; Jarlan et al., 2008; Barbu et al., 2011; Nearing et al., 2012; Fox
398 et al., 2018; Kumar et al., 2019; Raczka et al., 2021). At coarser resolution, analogs of
399 vegetation changes can also be obtained from passive microwave radiometry (Konings
400 et al., 2016; Feldman et al., 2021). While optical sensor measurements are subject to
401 coverage limitations from cloud obscuration, microwave measurements are nearly all-
402 weather. The assimilation of vegetation optical depth estimates from microwave mea-
403 surements has also shown promising results in improving land surface states, and it
404 provides the opportunity to extend the spatio-temporal coverage of higher-resolution
405 optical measurements (Kumar et al., 2020a; Mucia et al., 2021). Many of these stud-
406 ies have explicitly shown the significant promise of vegetation assimilation in captur-
407 ing human management impacts such as agriculture, where anthropogenic activities
408 lead to seasonal changes in vegetation growth (Mocko et al., 2021). Vegetation assim-
409 ilation has also shown effectiveness in capturing vegetation disturbance features from
410 fires, which often lead to changes in regional hydrological response (Kumar et al., 2021,
411 2022). DA offers a more practical approach to representing these changes on the land
412 surface, which are difficult to characterize with process-based models alone (Pongratz
413 et al., 2018).

414 Other reported land DA instances involve variables such as land surface tem-
415 perature, albedo, and water surface elevation. Land surface temperature and albedo
416 from thermal and optical sensors are variables at the heart of the surface energy bal-
417 ance and influence the latent and sensible heat fluxes at the land-atmosphere bound-
418 ary. Therefore, assimilation of these variables should have a direct influence on weather
419 and climate model forecast accuracy. Unlike variables such as soil moisture, snow, ter-
420 restrial water storage, and vegetation, the response timescale of land surface temper-

421 ature is more rapid, with associated diurnal and seasonal cycles. Additional challenges
422 related to bias correction at these finer timescales are described in studies of land sur-
423 face temperature assimilation, which have employed both variational and filtering tech-
424 niques (Castelli et al., 1999; Lakshmi, 2000; Bosilovich et al., 2007; Ghent et al., 2010;
425 Reichle et al., 2010; Han et al., 2013; Draper et al., 2015; Lu et al., 2017). Comparatively,
426 fewer studies examine the assimilation of albedo, despite the classification of albedo
427 as an essential climate variable (Hollmann et al., 2013), particularly over large spatial
428 scales. Most of the reported studies employ direct insertion approaches due to the lack
429 of prognostic representation of albedo variables within LSMs and have reported some
430 positive impacts on characterizing surface fluxes and snow states (Malik et al., 2012;
431 Yin et al., 2016; Boussetta et al., 2015; Kumar et al., 2020b).

432 In addition to NWP and S2S, improving the initial conditions for hydrological
433 forecasting applications through DA has also been the subject of many studies (e.g.,
434 Liu et al. (2012); Baugh et al. (2020)). Most of these focus on improving soil moisture
435 (e.g., Loizu et al. (2018)) and snow states (e.g., D. Li et al. (2019)), though a few stud-
436 ies have also examined the assimilation of hydrological states and fluxes such as dis-
437 charge (Clark et al., 2008; D.-J. Seo et al., 2009; Pauwels & De Lannoy, 2009; Thirel et
438 al., 2010; DeChant & Moradkhani, 2011; Fairbairn et al., 2017). Estimates of water stage
439 or elevation have been derived from radar altimeters since the early 1990s, provid-
440 ing nearly all-weather observations (Silva et al., 2010). The use of DA within hydro-
441 dynamic models has been examined to extend the spatio-temporal utility of these typ-
442 ically narrow swath altimeter measurements. These studies include efforts to assim-
443 ilate river discharge, velocity, and water level (Andreadis et al., 2007; Neal et al., 2007;
444 Ricci et al., 2011; Paiva et al., 2013; Ercolani & Castelli, 2017; Emery et al., 2020; El Gharamti
445 et al., 2021). The improved surface water characterization over data-poor transbound-
446 ary basins from the assimilation of remote sensing data is a highly relevant hydro-
447 logical modeling challenge not routinely employed in hydrological prediction systems.

448 **3 Major gaps in land data assimilation**

449 As described in the previous section, there have been significant advances in the
450 development of methodologies for land DA, the utilization of land surface measure-
451 ments from various spaceborne platforms as well as in-situ observations, and the use
452 of land DA for routine applications. Despite this progress, there are several signifi-

453 cant remaining gaps and challenges, which are described in this section. Figure 1 shows
454 a summary of the key topics described in this section.

455 **3.1 Reliance on the assimilation of retrieval products**

456 All DA methods require an “observation operator” or forward model to trans-
457 late variables from model space to observation space, and vice versa. The simplest ob-
458 servation operator is the identity matrix where the assimilated value is either an ob-
459 servation (in-situ soil moisture, atmospheric temperature) or a retrieval from (typically
460 remotely sensed) observations (e.g., soil moisture, snow mass, or leaf area index) of
461 a variable that is also estimated by the LSM. To date, most of the reported land DA
462 studies and applications have focused on the assimilation of retrieved products, whereas
463 the atmospheric and ocean communities have made far more progress in the direct
464 use of radiance measurements. The retrievals are obtained using radiative transfer or
465 empirical models with their own underlying assumptions of parameter, land surface,
466 and climate characteristics. These assumptions are often inconsistent with those si-
467 multaneously made in LSMs that are utilized in DA instances. In many cases, retrieval
468 models are developed and calibrated with limited airborne and field campaign data.
469 Subsequently, their application over different or large spatial extents leads to errors.
470 The assimilation of satellite radiances as in the SMAP L4.SM algorithm (Reichle et al.,
471 2019) is likely to reduce errors stemming from such inconsistencies and potentially
472 make observation errors more easily quantifiable (Lewis et al., 2012; Lievens et al., 2017;
473 de Rosnay et al., 2022), though forward radiative transfer models also suffer from un-
474 certainties related to conceptualization and parameterization. As noted earlier, given
475 the complexity of forward modeling for snow mass, machine learning models have
476 also been used as observation operators for snow DA studies.

477 The historical reliance on retrieval products is a significant obstacle for the more
478 widespread adoption of land DA in near-real-time or operational modeling environ-
479 ments. This is partly due to the fact that latency associated with retrieval products of-
480 ten does not meet the constraints of operational environments. Radiance assimilation
481 also provides an opportunity to update states impacting the forward modeling of land
482 surface emission (e.g., vegetation states controlling canopy opacity). To realize the more
483 common use of radiance measurements within land DA environments, relevant for-
484 ward modeling tools, whether they use physically-based radiative transfer models (e.g.,

485 Quaife et al. (2008); De Lannoy et al. (2013)) or employ statistical approaches (e.g., Aires
486 et al. (2021); Kolassa et al. (2018)), are needed. Compared to the breadth of LSMs, only
487 a handful of land relevant community forward modeling systems and tools are cur-
488 rently available (Parrens et al., 2014; Royer et al., 2017).

489 **3.2 Limitations from developmental legacies**

490 The legacy of LSMs has also significantly influenced the application of land DA
491 instances. As mentioned earlier, the first generation LSMs were conceptualized as the
492 boundary conditions to atmospheric models with the primary objective of improving
493 atmospheric prediction. The focus of ensuring physical realism of land surface states
494 has historically been a secondary priority. For example, while most of the major NWP
495 centers around the world maintain detailed scorecards of prediction skill of atmospheric
496 variables (Ebert et al., 2013; Haiden et al., 2018; Brown et al., 2021), such routine eval-
497 uations for land surface variables are uncommon. Consequently, process representa-
498 tions, and therefore the climatologies of variables such as soil moisture, have evolved
499 specifically for each associated NWP system. It is well documented that soil moisture
500 climatologies from different models are starkly different (Reichle et al., 2004), to the
501 point that they are essentially model-specific representations of soil wetness (and should
502 not even be called soil moisture; Koster et al. (2009)).

503 As noted earlier, the screen level assimilation approaches used in many NWP
504 centers update soil moisture states instead of atmospheric states assuming a causative
505 relationship between soil moisture boundary conditions and low-level atmospheric
506 forecasts. Since these assumptions are only valid when weakly forced atmospheric con-
507 ditions are present, the screen level DA scheme to adjust soil moisture often has be-
508 come an approach that compensates for true underlying errors in other parts of the
509 model (Draper et al., 2009). It is recognized that the resulting improvements in the low-
510 level atmospheric forecasts are not necessarily due to improved representation of land
511 surface states (Douville et al., 2000; Drusch & Viterbo, 2007), and may in fact degrade
512 the surface states and hydrological consistency (e.g., Zsoter et al. (2019)). Consequently,
513 the assimilation of observations or retrievals which have a direct relation to soil mois-
514 ture becomes problematic in such DA environments. For example, studies have demon-
515 strated that assimilation of satellite soil moisture data and screen level temperature
516 and humidity observations often lead to diverging estimates of root zone soil mois-

517 ture updates (Draper et al., 2011), though satellite soil moisture assimilation itself im-
518 proves land surface states (Carrera et al., 2019; Munoz-Sabater et al., 2019). As we noted
519 above, with the evolution of LSMs, they are also increasingly being used in environ-
520 ments where there is a direct need for accurate characterization of land surface, with-
521 out being necessarily connected to atmospheric models and requirements. As the need
522 to improve the physical realism of land surface states increases and direct measure-
523 ments of land surface states become increasingly available, methodologies that treat
524 the land surface states as an ‘error sink’ are fundamentally undercutting the poten-
525 tial of land DA. Localization strategies (Anderson, 2007) that selectively adjust por-
526 tions of the model states most closely related to the observation may also be needed
527 to circumvent such limitations.

528 **3.3 Bias correction and efficiency issues**

529 Another key issue related to the lack of physical realism of LSMs is the current
530 practice of bias correction within land DA environments. In real data assimilation sys-
531 tems, systematic errors (including biases) between model estimates and observations
532 are unavoidable, and they are typically caused by a combination of errors from in-
533 strument noise, retrieval issues, and model deficiencies, related to model structure and
534 uncertainties in parameters and inputs. The model structural deficiencies are also tied
535 to the legacy issues mentioned in Section 3.2. Additionally, the lack of physical real-
536 ism in LSM estimates also stems from the difficulties of the complex heterogeneity of
537 the land surface, which makes it difficult to develop spatially distributed assessments
538 of biases from limited in-situ measurements (that are not spatially representative). Prior
539 studies have explored both bias-aware and bias-blind approaches in land DA (Kumar
540 et al., 2012). Bias-aware approaches typically attribute the source of the biases exclu-
541 sively to the model or the observations and use the analysis increments to progres-
542 sively estimate the bias (De Lannoy et al., 2007; Bosilovich et al., 2007; Reichle et al.,
543 2010). While these approaches are useful for systems with transient changes in bias,
544 the lack of realism in the model estimates or observations makes their application dif-
545 ficult. Bias-blind approaches are more common in land DA systems, where the rel-
546 ative biases between model estimates and observations are removed before assimi-
547 lation, essentially focusing on the correction of errors in short-term or interannual vari-
548 ations. Approaches such as Cumulative Distribution Function (CDF)-matching and

549 standard normal deviate scaling are commonly used in soil moisture assimilation sys-
550 tems to rescale the observations into the model climatology prior to assimilation (Reichle
551 & Koster, 2004; Crow et al., 2005a; de Rosnay et al., 2020). The observation rescaling
552 approaches require long data records of observation and modeled soil moisture es-
553 timates to develop scaling parameters, which are difficult with new satellite missions.
554 Similarly, long model integrations with a consistent driving meteorology are required
555 to establish the model climatology, which also often do not exist for operational en-
556 vironments because of frequent system updates. Moreover, regional modeling centers
557 also often change their domains of focus, which makes the reliance on an established
558 soil moisture climatology specific to each modeling domain impractical. The rescal-
559 ing approaches are also documented to reduce the information transfer from retrievals
560 by about 10% due to errors in empirical CDF computations (Nearing et al., 2018) and
561 violation of theoretical assumptions underlying the matching of higher-order statis-
562 tical moments in the CDF approach (Yilmaz & Crow, 2013).

563 Further, recent studies have shown that the rescaling approaches are problem-
564 atic when stationarity assumptions about the model-observational biases do not hold
565 true, dynamic changes in bias occur, or when human management factors or unmod-
566 eled processes are present in the observational signal and missing from the models
567 (Kumar et al., 2015b; Girotto et al., 2017). In this regard, the land surface is unique com-
568 pared to the atmosphere and ocean, as the direct influence of human activities (e.g.,
569 irrigation, vegetation disturbances, groundwater pumping, urbanization, reservoir man-
570 agement) is ubiquitous on the land. Because many of these processes are subjective
571 in nature, their accurate representation within models is hard. For example, though
572 irrigation representations have been developed within LSMs, the determination of ir-
573 rigation onset, duration, and magnitude is difficult to specify accurately (Massari et
574 al., 2021). The use of remote sensing observations through DA offers a more practi-
575 cal approach to incorporating such impacts, without the reliance on heuristic rules.
576 However, as the unmodeled signals are often manifested as systematic errors or bi-
577 ases, new approaches are needed that can identify the root causes of bias, including
578 bias from systematic instrument errors, model parameter and parameterization de-
579 ficiencies, and unmodeled processes; otherwise, it will remain difficult if not impos-
580 sible to preserve important observational signals (Kumar et al., 2015b).

3.4 Methodological limitations

As noted earlier, owing to the difficulties in deriving adjoint models for highly non-linear LSMs, Kalman filtering approaches have become the most commonly used methods for land DA, providing a computationally efficient alternative to variational DA systems with comparable performance (Reichle et al., 2002b; Caparrini et al., 2004; Kotsuki et al., 2022). One of the key assumptions in the serial applications of the EnKF is that model and observation errors are Gaussian and mutually and serially uncorrelated (Katzfuss et al., 2016), which are often not realistic in land DA systems. The particle filter (Moradkhani et al., 2005; Smyth et al., 2019) overcomes this limitation, by approximating the model posterior distribution with Monte Carlo sampling, but requires large ensemble sizes (Weerts & El Serafy, 2006; Dong et al., 2015). Another possibility are hybrid ensemble-variational techniques (Bannister, 2017), which solve the variational problem by approximating required properties of the adjoint and tangent linear using a model ensemble and are being increasingly used in NWP. Hybrid approaches also show promise for overcoming some of the limitations inherent in other ensemble DA techniques (such as limited sample size), and have been applied to land surface modeling problems (e.g., Pinnington et al. (2021)).

However, all of these ensemble DA systems can suffer from issues of filter collapse and degeneracy when adequate ensemble spread does not exist, which is sometimes addressed with techniques such as covariance inflation (Anderson & Anderson, 1999; Fox et al., 2018; Gharamti, 2018, 2021). These techniques also require significant tuning and knowledge of observation error covariances (Miyoshi & Yamane, 2007). The reliance on the ensemble to derive the model error covariances can be limiting in some instances when the state variable is bound by theoretical limits. For example, since fractional snow cover is constrained by zero as the lower limit and unity as the upper limit, the effective generation of spread in the ensemble becomes problematic when the model estimate is close to these bounds and has been shown to cause disadvantageous impacts within DA (Arsenault et al., 2013). Reliance on a single choice of surface meteorological forcing has also been shown to underestimate the land model error covariance and reduce the efficiency of land DA (Kumar et al., 2017). Studies have also shown that DA can lead to water balance errors (which are particularly important in coupled land-atmosphere systems) when the analysis increments are not explicitly constrained to conserve the water balance (Giroto et al., 2021). The use of

614 weak constraints in land DA systems has been explored to reduce these inefficiencies
615 (Pan & Wood, 2006; Yilmaz et al., 2011). Finally, most DA systems are designed to work
616 with errors that are strictly random. Since systematic errors are more dominant in land
617 hydrology processes, the primary focus on improving short-term errors through DA
618 often limits the level of improvements possible through land DA (Section 3.3).

619 **3.5 Lack of information transfer in land data assimilation efforts**

620 A key appeal of DA is the potential to extend the information in observations
621 to other model states that are connected to the observations. Land DA efforts have
622 yielded mixed results in effectively realizing this goal, due to a variety of reasons. For
623 example, the assimilation of surface soil moisture retrievals has been widely used to
624 develop improvements in deeper soil moisture, with studies reporting varying degrees
625 of success. The low skill of soil moisture retrievals from older sensors was a key fac-
626 tor in early soil moisture DA studies that reported small improvements in root zone
627 estimation. Larger improvements in subsurface soil moisture estimation with DA have
628 been reported with data from newer L-band sensors such as SMOS and SMAP (De Lan-
629 noy & Reichle, 2016a; Reichle et al., 2017). Similarly, larger improvements from DA
630 are often reported with studies involving older or simpler soil moisture models (F. Li
631 et al., 2010; Mladenova et al., 2019), as the information from remote sensing data com-
632 pensates for the lack of skill in those models. When provided with high-quality bound-
633 ary conditions, modern LSMs are fairly skillful in their soil moisture estimates. As a
634 result, the added improvements from soil moisture DA featuring such models and high
635 quality precipitation inputs are relatively small (Kumar et al., 2014; Kolassa et al., 2017b;
636 E. Seo et al., 2021; Reichle et al., 2021a). The strength of the surface-subsurface soil mois-
637 ture coupling within LSMs is a significant factor in the level of improvements obtained
638 in the root zone through surface soil moisture assimilation (Kumar et al., 2009). Given
639 that the true surface-subsurface coupling strength is largely unknown and significantly
640 heterogeneous, these studies suggest that there is a strong model configuration de-
641 pendence on the level of improvements in root zone soil moisture realized through
642 surface soil moisture assimilation.

643 Attempts to extract information from the assimilation of remote sensing soil mois-
644 ture observations about associated key water cycle variables such as evapotranspira-
645 tion (ET), runoff, and river discharge have also produced mixed results. For exam-

646 ple, most soil moisture and vegetation assimilation studies that examined the impact
647 on ET and/or river discharge either report marginal improvements or degradations
648 (Peters-Lidard et al., 2011; Fairbairn et al., 2017; Albergel et al., 2017; Kumar et al., 2018;
649 Munoz-Sabater et al., 2019; Bonan et al., 2020). Similar to the earlier example on root
650 zone soil moisture, the information transfer issues with other variables such as ET are
651 also related to the inherent assumptions of coupling between the relevant variables
652 (Crow et al., 2018, 2020). In reality, such coupling relationships are dependent on a
653 number of factors including land surface characteristics, surface meteorology, and cli-
654 mate conditions (Dong et al., 2020a). For example, soil moisture has a strong influ-
655 ence on ET in a water limited domain but may not be the primary controlling factor
656 in an energy limited domain (Kumar et al., 2020a). Similarly, LSMs often underesti-
657 mate the coupling between simulated soil moisture and subsequent surface runoff and
658 thus potentially squander valuable streamflow predictability associated with soil mois-
659 ture assimilation (Crow et al., 2017, 2018, 2019). Consequently, the results from DA
660 will be misleading if the LSM does not adequately represent such coupling relation-
661 ships accurately, even when high quality observations are assimilated. The impact of
662 such inherent model coupling features is also reflected in studies assimilating obser-
663 vations related to vegetation that have shown stronger impact on root zone soil mois-
664 ture than from the assimilation of surface soil moisture retrievals (Albergel et al., 2017,
665 2020; Kumar et al., 2020a).

666 In other cases, the lack of success in information transfer is related to a lack of
667 adequate model prognostics or limited physics. For example, the assimilation of land
668 surface temperature within LSMs has been problematic because there are large dis-
669 crepancies in the representation of surface temperature in models and the skin tem-
670 perature derived from remote sensing instruments (Bosilovich et al., 2007). Partly due
671 to the lack of an equivalent model prognostic, the land surface temperature assimi-
672 lation results have yielded little improvement in surface fluxes (Reichle et al., 2010).
673 Similarly, though terrestrial water storage retrievals from the GRACE mission have
674 provided valuable insights on subsurface water storage changes around the world,
675 their assimilation may produce erroneous trends because most LSMs only include a
676 shallow groundwater representation (whereas GRACE may be observing deeper ground-
677 water storage signals; Girotto et al. (2017)).

678 Another key factor in the lack of information transfer is the limited skill in ob-
679 servations themselves. Though there is often a direct connection between streamflow
680 and melt from the snowpack, little success in this area is reported from snow DA stud-
681 ies that solely rely on remote sensing observations (Andreadis & Lettenmaier, 2006;
682 Kumar et al., 2014; Zsoter et al., 2019). On the other hand, the use of more reliable data
683 from airborne or ground measurements of snow has shown greater promise in im-
684 proving streamflow (C. Huang et al., 2017; Lahmers et al., 2022), suggesting that sig-
685 nificant improvements in remote sensing snow retrievals are needed to improve their
686 utility. As noted above, despite the availability of multi-decadal observations of land
687 surface variables such as land surface temperature, albedo, and leaf area index, these
688 observations were used in relatively few DA studies, mainly because LSMs have large
689 model biases or have not considered and developed appropriate prognostic represen-
690 tations of such variables.

691 **3.6 Spatio-temporal characterization of error specifications**

692 The specification of model and observation errors is a critical factor in the per-
693 formance of land DA systems and remains a significant challenge (Crow & Loon, 2006).
694 Observation errors must take into account instrument errors, errors in observation op-
695 erators or retrieval algorithms (considered synonymous with representativeness er-
696 rors (van Leeuwen, 2015)); consequently, observation error parameters are difficult to
697 measure and specify in a spatially distributed manner. Though spatial correlations are
698 present in the observation errors, they are often ignored in DA environments (Ying,
699 2020). Model error specification must consider limitations in model physics and un-
700 certainties in boundary conditions and parameters, among other factors. It is gener-
701 ally recognized that it is more difficult to specify model errors than observation er-
702 rors because the general lack of physical realism in most LSM outputs makes it dif-
703 ficult to perform direct comparisons against available ground/truth measurements (Reichle,
704 2008a). Additionally, the significant heterogeneity and complexity of land processes
705 makes it difficult to develop spatially distributed estimates of model errors. Because
706 of these limitations, the derivation of model error covariance specifications is often
707 done through idealized experiments and largely remains a subjective process (Kumar
708 et al., 2008, 2009; Carrera et al., 2015; Fox et al., 2018; Raczka et al., 2021). In fact, most
709 DA studies to-date employ domain-wide uniform error specifications. DA diagnos-

710 tics such as innovations and analysis increments are routinely used to evaluate error
711 assumptions in DA configurations, by examining their deviations from the theoret-
712 ical ideal values (e.g., Reichle et al. (2017)). These methods have also been employed
713 within DA systems to adaptively diagnose and improve error variance specifications
714 (Reichle et al., 2008a; Crow & Reichle, 2008). Similarly, the application of the triple col-
715 location approach has also been shown to improve the estimation of soil moisture re-
716 trieval errors by removing the problematic assumption – required by innovation anal-
717 ysis – that errors in assimilated observations are serially uncorrelated (Crow & Berg,
718 2010). These theoretical approaches, however, become less effective when observational
719 signals are dominated by unmodeled processes, which can be especially true for land
720 DA systems (Kumar et al., 2015b). In land-ensemble filtering, the forecast ensemble
721 spread is typically created by adding small perturbations sampled from randomly gen-
722 erated noise to the model states at regular time intervals. Recent studies have shown
723 that this approach creates unrealistic estimates of forecast uncertainty, since the en-
724 semble spread reflects the local model persistence of the applied perturbations, which
725 is not necessarily an accurate representation of the model uncertainty (Draper, 2021).
726 Explicit consideration of uncertainty from several contributing sources for more re-
727 alistic model spread representation has been reported in some studies (Raeder et al.,
728 2021; Dokoohaki et al., 2022). Additionally, the model background error can be un-
729 derestimated due to the limitations of the specific model configurations (Kumar et al.,
730 2017).

731 **3.7 Limitations in the use of metrics and evaluation strategies**

732 The typical approach in DA studies is to use the available (often point-scale) ground
733 measurements as the truth for evaluating the impact of DA, with the inherent assump-
734 tion being that the spatio-temporal scale of the model and remote sensing data is suf-
735 ficiently fine to contain information about the in-situ measurements. This assumption
736 may not always be valid, as shown in studies that have quantified information use
737 efficiency within DA systems (Nearing et al., 2018). The assessments versus in-situ mea-
738 surements also overlook the larger utility of remote sensing datasets to provide spa-
739 tially distributed coverage, offering the possibility to capture spatial features and het-
740 erogeneity that cannot be obtained with in-situ measurements. Finally, the geographic
741 distribution of in-situ stations is skewed toward data-rich regions, where model per-

742 performance is generally good, whereas the impact of land DA is generally larger in data-
743 sparse regions, which was demonstrated using instrumental variable approaches (e.g.,
744 Dong et al. (2019); Reichle et al. (2021a)), including the aforementioned triple collo-
745 cation method (Gruber et al., 2020). Therefore, spatially distributed evaluation strate-
746 gies and metrics must be emphasized in demonstrating the potential utility of remote
747 sensing datasets. For example, the impacts of agricultural activity and man-made dis-
748 turbances such as fires are hard to represent in models or capture through ground mea-
749 surements. Assimilating vegetation datasets, on the other hand, has been effective in
750 representing such features over large spatial scales (Kumar et al., 2020a, 2021; Mucia
751 et al., 2021; Kumar et al., 2022).

752 Additionally, when representativeness differences are significant with variables
753 such as soil moisture, DA evaluations mainly focus on the use of temporal anomaly
754 metrics (Entekhabi et al., 2010b). The restricted focus on the evaluation of temporal
755 variability has been shown to underestimate errors in soil moisture datasets (Dorigo
756 et al., 2010) and miss out on capturing the impact of unmodeled processes (Kumar
757 et al., 2015b). Evaluation strategies based on approaches such as information theory
758 (Shannon, 1948), as used for instance in NWP context (Cardinali, 2009), that allow for
759 the spatially distributed quantification of the information transfer and noise reduc-
760 tion can be considered to better demonstrate the utility of land remote sensing and
761 DA (e.g. Balsamo et al. (2007)). Since the land is particularly impacted by anthropogenic
762 impacts, land DA must also focus on demonstrating the characterization of such pro-
763 cesses, beyond simply improving processes driven by natural variability. It must be
764 stressed that some of these evaluation limitations are due to a profound lack of ob-
765 servational capability for many land surface variables. For example, the point-to-satellite
766 footprint representativeness issues for soil moisture are more severe for the represen-
767 tation of spatial patterns than for temporal anomaly statistics (Chen et al., 2019). Datasets
768 or observational techniques that allow for a direct evaluation of soil moisture patterns
769 are generally lacking.

770 **3.8 Multivariate and land reanalysis efforts**

771 While there have been numerous efforts to develop consistent multi-decadal es-
772 timates of atmospheric and ocean states through reanalysis of all available remote sens-
773 ing datasets, such efforts are largely non-existent in the land and hydrology commu-

774 nity. This is partly because land DA development has historically lagged behind DA
775 development in the atmosphere and ocean communities, despite the availability of long
776 land surface and hydrology observational data records. In fact, multivariate and con-
777 current DA of land variables have only been reported in more recent years. An ob-
778 vious beneficial impact from multi-sensor and multivariate DA setups has been in im-
779 proving the temporal coverage of observational inputs within the DA environment.
780 For example, since 2010, soil moisture retrievals from multiple passive and active mi-
781 crowave instruments have become available, which improves the temporal observa-
782 tional coverage from remote sensing. Similarly, vegetation conditions are available from
783 both optical/thermal infrared and microwave sensors, with each contributing unique
784 advantages (Section 2). Combining both types of vegetation measurements can help
785 in providing a more continuous observational constraint with a DA environment (Kumar
786 et al., 2020a; Mucia et al., 2021). This can be particularly important when temporally
787 continuous coverage of land surface conditions is needed for monitoring of hazards
788 such as fires and related vegetation disturbances (Kumar et al., 2021).

789 The concurrent handling of complementary information from multiple observa-
790 tion types is a significant challenge in multivariate DA environments (Montzka et al.,
791 2012). Because soil moisture and snow are important in warm and cold seasons, re-
792 spectively, the assimilation is, for the most part, mutually exclusive and thus relatively
793 straightforward. The development of a reanalysis that incorporates soil and snow depth
794 from multiple sensors across three decades has demonstrated improvements in land
795 surface states, particularly with more modern sensors (Kumar et al., 2018). Resolving
796 the disparities in temporal and spatial resolution differences across observational prod-
797 ucts is another challenge in multivariate DA systems (Tian et al., 2017). Generally, stud-
798 ies have reported added improvements from multivariate DA than those from cor-
799 responding univariate configurations (H. Su et al., 2010; Hain et al., 2012; Zhao & Yang,
800 2018; Abbaszadeh et al., 2020). Multivariate observational constraints may be neces-
801 sary in instances where observations provide unique insights into related processes.
802 For example, univariate assimilation of terrestrial water storage or soil moisture in places
803 where groundwater pumping is prevalent may produce incorrect results, while their
804 joint assimilation provides more consistent improvements across relevant water stor-
805 age terms (Giroto et al., 2019; Tangdamrongsub et al., 2020). Given the simultaneous
806 impact from anthropogenic processes and the interlinked nature of water cycle pro-

807 cesses, emphasis on multivariate DA environments will be needed for highlighting
808 the utility of land DA. There is also a significant need to develop land reanalysis prod-
809 ucts so that an observation-informed spatially and temporally continuous climate record
810 of water, energy, and carbon cycle changes on the land surface can be established.

811 **3.9 Coupled DA environments**

812 Since the initial development of land DA was motivated by the need to improve
813 NWP, most applications of land DA are limited to providing better initial conditions
814 prior to a forecast. As noted earlier, such independent development of approaches can
815 lead to divergent developmental priorities and inconsistent results with land and at-
816 mosphere systems. Coupling between the respective DA systems is needed to reduce
817 the imbalance between them. Generally, there are two types of coupled DA systems
818 (Penny & Hamill, 2017). In a weakly-coupled DA system, communication and data
819 exchange between the land and the atmosphere only occurs through the physics within
820 a coupled forecast model, which is the land DA approach commonly used in the con-
821 text current NWP and reanalysis development (de Rosnay et al., 2013; Draper & Re-
822 ichle, 2019; Carrera et al., 2019; Gomez et al., 2020; Reichle et al., 2021b). A strongly
823 coupled DA environment, on the other hand, facilitates cross-component interactions
824 during DA, to allow the use of cross-model error covariances during DA (Shahabadi
825 et al., 2019). Though there is a recognition that strongly coupled DA systems enable
826 increased exploitation of the information from observations and are better at address-
827 ing the imbalance issues, no such operational systems exist currently (Sluka et al., 2016).
828 Coupled DA approaches are going to be pursued for their consistent surface-atmosphere
829 initialization (for a “spatial-consistency” target), beneficial for NWP prediction or within
830 a climate reanalysis. They, however, carry a substantially higher computational bur-
831 den. Land-only DA systems are, therefore, likely to thrive and provide valuable and
832 complementary solutions for frequent land reanalysis and permitting to initialize past
833 reforecasts in support of anomaly-based predictions, such as those produced in the
834 S2S framework (for a “time-consistency” target). While land DA and coupled DA ap-
835 proaches may have important methodological differences, they do face common chal-
836 lenges related to model and observation errors, characterization, handling of biases,
837 quality control, compensating coupling errors (Vitart et al., 2019). Therefore several
838 of the priorities identified hereafter are crosscutting.

839 **4 Future priorities, new opportunities, and recommendations**

840 Despite the challenges and gaps described above, DA methodologies are vital
841 tools for realizing the promise of land surface observations. In this section, we out-
842 line the priorities for advancing land DA including the need for improvements in mod-
843 els and observations, co-development of models with the DA community, and embrac-
844 ing technological advancements, so that land DA environments can not only help to
845 realize the full potential of land surface observations, but also help define the next gen-
846 eration observational needs. These enhancements will likely require an iterative pro-
847 cess, and coordination across several interdisciplinary areas, as shown in Figure 2. A
848 number of specific recommendations to realize this vision is outlined below.

849 **4.1 Enable multivariate, multi-sensor DA**

850 As noted earlier, the number and type of space-borne observations relevant to
851 the land surface and hydrology continue to increase. This increased availability of ob-
852 servations presents unique challenges and opportunities for significantly advancing
853 land and hydrologic prediction (McCabe et al., 2017). The simultaneous availability
854 of multiple observation types of land surface processes provides concurrent water and
855 carbon budget constraints, and has the potential to reduce errors in water and veg-
856 etation state estimates beyond what is achievable with numerical process models alone.
857 This is particularly relevant in regions of the world where water cycle processes are
858 dominated by anthropogenic impacts and univariate assimilation may provide mis-
859 leading results.

860 The wealth of observational information available from multiple platforms is a
861 unique opportunity to embrace a “systems approach” (Jenkins, 1969) to land DA that
862 truly acknowledges the interconnected nature of water cycle processes and the hu-
863 man influence by enabling multivariate and multi-sensor environments. Focused site
864 level analyses (e.g., Fox et al. (2018)) can be used to systematically diagnose the util-
865 ity of specific observational inputs and to minimize representativeness errors. Such
866 strategies can also be used to configure appropriate localization approaches to spec-
867 ify the extent of information transfer, and to reduce sampling errors from ensemble
868 structures.

869 **4.2 Develop coupled DA**

870 In strongly coupled DA, multiple components from a coupled model are updated
871 in a single DA step. Strongly coupled land-atmosphere DA expands the possibilities
872 for multivariate and multisensor DA, and its potential advantages are being actively
873 explored. There have been significant advances in the development of coupled Earth
874 system models, through efforts such as the Earth System Modeling Framework (ESMF;
875 www.earthsystemmodeling.org) that provide interoperable standards and tools for the
876 community. Similar, concerted efforts to enable such paradigms (e.g. Joint Effort for
877 Data Assimilation Infrastructure; JEDI; Tremolet and Auligne (2020)) and provide flex-
878 ible and computationally scalable software infrastructure are needed to realize the vi-
879 sion of coupled Earth system DA environments (de Rosnay et al., 2022). Recent stud-
880 ies examining strongly coupled land-atmospheric DA systems have also stressed the
881 challenges of addressing inconsistencies in spatial and temporal error correlation scales
882 (across the land and atmosphere), the need for reliable forward modeling tools, and
883 the recognition of the memory and timescale differences in land and atmospheric dy-
884 namics (Lin & Pu, 2020; Draper, 2021).

885 **4.3 Improve the physical realism of LSMs**

886 Many of the limitations of land DA environments are closely related to LSM de-
887 ficiencies, as discussed in Section 3.2. Therefore, there needs to be a strong empha-
888 sis on improving the physical realism of LSM process descriptions and of retrieval prod-
889 ucts so that interpretation errors and efficiency loss issues can be reduced. Calibra-
890 tion of model parameters can be used to reduce systematic errors in models prior to
891 DA - typically performed using optimization algorithms and treating the parameters
892 as time-invariant (Zhou et al., 2020). Online DA tools have also been applied in this
893 regard, simultaneously adjusting the model parameters and states. While computa-
894 tional cost is often prohibitive in employing these algorithms over large spatial ex-
895 tents and fine resolutions, more computationally efficient approaches founded in ma-
896 chine learning have been developed more recently (Tsai et al., 2021). Embracing these
897 technological advances and improving the physical realism of the model and obser-
898 vations should reduce the reliance on rescaling approaches and enable the direct in-
899 corporation of observations within land DA systems. In addition to the parameters,
900 LSM development must also include the refinement of the inherent coupling repre-

901 presentations between relevant energy, water, and carbon cycle processes. Diagnostic stud-
902 ies that identified inconsistencies in coupling strength representations (Section 3.5) should
903 serve as a guideline for such improvements. The same observational advances that
904 enable the assimilation of higher-quality state and flux estimates into LSMs also pro-
905 vide opportunities for better assessing, and correcting, coupling strength biases in LSMs.
906 If the objective of information transfer to related variables through DA is to be real-
907 ized, such model process and parameter improvements are essential. Finally, LSM for-
908 mulations must evolve to include relevant prognostics that can be employed in land
909 DA environments. More broadly, there is a significant need to consider the require-
910 ments from assimilation needs in LSM development.

911 **4.4 Enhance spaceborne observations**

912 There is also a need for co-producing land observation datasets in coordination
913 with the land DA community so that a consistent characterization of spatio-temporal
914 errors of observations can be established. Given the prevalence and heterogeneity of
915 bias in land modeling and assimilation, bias-aware DA algorithms that are built to di-
916 agnose and incorporate systematic errors during DA (as opposed to addressing them
917 *a priori*) are needed (Dee, 2005; Pauwels et al., 2013; Ridler et al., 2017). However, such
918 bias-aware systems are only practical if the climatology of the observations can be con-
919 sidered as the reference (Dong et al., 2020b). Efforts to reduce biases in land remote
920 sensing retrievals or forward modeling approaches should be a priority to achieve these
921 goals (Grant et al., 2008; Gao et al., 2020). Further, coordination across mission science
922 teams and data producers to reduce errors and inconsistencies across products will
923 be beneficial in improving the quality of observational products available for assim-
924 ilation. Such data homogenization and harmonization efforts can reduce interpreta-
925 tion errors and foster the development of universally accepted observation benchmarks
926 (McCabe et al., 2017). DA systems, in fact, offer natural environments for the harmo-
927 nization of satellite data from multiple platforms, by appropriately handling the re-
928 spective sources of errors and uncertainties. Land reanalysis and multivariate assim-
929 ilation efforts incorporating multiple sources of data should be considered important
930 goals for the land DA community. Such efforts are necessary to establish climate data
931 records, which are important benchmarks for climate studies (C.-H. Su et al., 2016).

932 **4.5 Exploit ground measurements**

933 As noted in the introduction, the focus of this article is primarily on the assim-
934 ilation of space-borne, spatially distributed measurements. Ground-based measure-
935 ments of land surface variables such as soil moisture, snow, ET, and river flow are avail-
936 able from several measurement networks and focused field campaigns. These mea-
937 surements have played and will continue to play an important role in enabling the
938 systematic evaluation and optimization of DA analyses. Further coordination with the
939 ground measurement community (e.g. the International Soil Moisture Network (ismn.geo.tuwien.ac.at),
940 FLUXNET (fluxnet.org), and the U.S. National Ecological Observatory Network (NEON;
941 www.neonscience.org)) to communicate the specific DA relevant needs is required to
942 overcome some of the gaps discussed in Section 3. For example, reliable, co-located
943 in-situ measurements of land surface states and fluxes (soil moisture, ET, runoff, LAI)
944 can help in validating satellite data (e.g., SMOSREX (de Rosnay et al., 2006) and VA-
945 LERI (Weiss et al., 2001)) and establishing true surface-subsurface coupling relation-
946 ships, which are needed to improve the realism of such formulations in models and
947 to ultimately improve the efficiency of DA methods. Similarly, focused ground-truthing
948 efforts to characterize human management impacts (e.g., irrigation impact on soil mois-
949 ture) would also be helpful in validating space-borne measurements over such areas.
950 Much like the anticipated growth in the number of space-based platforms and mea-
951 surements, the advent of new technologies is fostering an era of crowd-sourced ob-
952 servations from platforms such as mobile phones and inexpensive sensors that can
953 be placed within a virtual network through them. For example, crowd-sourcing via
954 smartphone apps and GPS devices has revolutionized how traffic information is col-
955 lected and updated in real-time. While there are still challenges in ensuring the re-
956 liability and quality of such measurements for science applications, they offer new pos-
957 sibilities not just for evaluation, but also for DA, particularly if their availability be-
958 comes routine and spatially distributed.

959 **4.6 Expand forward modeling capabilities**

960 Advances in forward modeling capabilities are needed to reduce the reliance on
961 retrieval products and to enhance the exploitation of observations in coupled DA sys-
962 tems. In addition to radiative transfer models, the land DA community could also lever-
963 age advances in machine learning to develop observation operators. Even if the rel-

964 evant radiative transfer model is available for the translation of a certain type of ob-
965 servation, ensuring their reliability over different types of terrestrial, biosphere, and
966 climatic conditions often requires significant calibration and validation efforts. The mod-
967 ern class of machine learning methods such as deep learning has shown great promise
968 in its ability to perform hierarchical feature extraction and could effectively function
969 as a forward operator within DA systems (e.g., Shan et al. (2022)). Generally, machine
970 learning models are often criticized for their non-transparent nature, a hindrance to
971 developing physical understanding and insights into the results generated by them.
972 Recent advancements have also enabled ‘explainable AI’ methods (Montavon et al.,
973 2018) that have specifically focused on increasing transparency of the decisions and
974 inner workings of the machine learning algorithms. These methods are useful for pro-
975 viding improved understanding of the sensitivities of the inputs, which can lead to
976 improved information exploitation to produce more skillful retrieval products or us-
977 ing such translations directly within DA systems. These methods can also be applied
978 to understand when and where remote sensing measurements provide relevant in-
979 formation content, which can then be used in DA environments to screen observations
980 and refine the quality control procedures to improve the observational information
981 utilization. In addition to providing a statistically robust model, machine learning-based
982 emulators can help in reducing the computational expense of retrieval and/or forward
983 modeling (Rodriguez-Fernandez et al., 2015; Kolassa et al., 2018).

984 **4.7 Incorporate human-driven land processes**

985 A unique challenge and opportunity for land surface DA is to account for hu-
986 man activities for which we either do not have enough knowledge to implement pro-
987 cess representations in models or which are inherently impossible to implement due
988 to their subjective nature. As discussed earlier, the identification and inclusion of such
989 signals and their attribution to the systematic differences stemming from them are needed
990 in such instances. The use of machine learning methods can be potentially useful in
991 this regard. Machine learning methods are also increasingly used for object detection
992 and feature extraction applications (LeCun et al., 2015; Bengio, 2009), where the al-
993 gorithm automatically learns and characterizes features from the data that are presented
994 to it. As these methods are less ‘feature engineering’ dependent, they can be an ef-
995 fective tool for identifying unmodeled features and characterizing the sources of bias

996 in land observations and models. Taking advantage of such capabilities and incorpo-
997 rating machine learning-based modernizations (Reichstein et al., 2019) are likely to re-
998 sult in better characterization of anthropogenic signals.

999 **4.8 Exploit new space-borne observation technologies**

1000 With the availability of remote sensing opportunities afforded by smallsats and
1001 cubesats, commercial platforms for Earth observations are becoming more prevalent
1002 (Sandau, 2010; Llop et al., 2014). While space-agency missions provide long-term cov-
1003 erage for critical Earth processes, they tend to be costly and often require decades of
1004 development. The smallsats are less costly and can provide continuous coverage at
1005 fine resolutions, when launched in multiples. Additionally, other sensing opportuni-
1006 ties from smart phones, unmanned aerial vehicles, and other close-to-Earth platforms
1007 are going to increase, fostering the era of “big data” (Sellars et al., 2013). While these
1008 measurements could be exploited within DA environments, they also likely present
1009 additional technology hurdles. As the size of the datasets become larger, there is an
1010 increasing shift to using cloud-based resources away from local computing resources.
1011 Effective and consistent quality control of these measurements from such distributed
1012 platforms can also be challenging. Modeling and DA systems must anticipate and adapt
1013 to such changing technologies if the significant potential of such datasets is to be re-
1014 alized.

1015 **4.9 Employ land DA in mission planning efforts**

1016 Land DA efforts have a critical role to play in the definition of the next gener-
1017 ation of observational types and needs. Given the considerable resources required to
1018 implement Earth observing missions, their utility must be systematically assessed through
1019 observing system simulation experiments (OSSEs), prior to instrument development.
1020 DA is key to enabling OSSEs, which can help assess the relative utility of competing
1021 mission designs (e.g. Tan et al. (2007); Kaminski and Mathieu (2017)). Mature land DA
1022 environments can be used in this regard for the sub-selection of mission configura-
1023 tions and technologies and to define the level of accuracy required from the observa-
1024 tions to meet science utility. Comprehensive multi-observation OSSEs can also help
1025 to identify the spatio-temporal gaps in the types of observations that are needed to
1026 improve the characterization of the water cycle. For example, snow and soil moisture

1027 processes may be seasonally dominant for water availability characterization in a re-
1028 gion, but the accuracy, revisit, and spatial resolution requirements of those observa-
1029 tions are likely to be different for water balance closure or water availability charac-
1030 terization needs. By enabling ‘what-if’ studies, OSSEs can formalize such requirements
1031 and influence the recommendations and definitions of future land hydrology observ-
1032 ing systems. In a pioneering effort, the SMAP mission was informed by land DA-based
1033 OSSEs in its formulation phase (Crow et al., 2005b; Reichle et al., 2008a) and has been
1034 operationally generating a soil moisture assimilation product (Reichle et al., 2019). OSSEs
1035 have been used to quantify the anticipated utility of snow measurements from dif-
1036 ferent sensors and technologies to guide the development of future snow missions (Garnaud
1037 et al., 2019; Kwon et al., 2021; Wrzesien et al., 2022). Similar land DA environments
1038 and efforts must be considered part of the lifecycle of future land surface satellite mis-
1039 sions.

1040 **5 Summary and Conclusions**

1041 Land DA has been an active area of research for some decades now and has shown
1042 significant promise for improving land surface characterization and benefiting NWP,
1043 S2S and operational hydrology forecasting environments. However, to realize its full
1044 potential, the land DA community, in conjunction with the relevant observational and
1045 modeling communities should address the significant gaps and challenges laid out
1046 in this article. Separate configurations of land DA environments specifically tailored
1047 for different applications will likely be needed to overcome the constraints of devel-
1048 opmental legacies, lack of model realism, and to foster the development and widespread
1049 acceptance of land DA as a necessary capability. We hope this agenda article will serve
1050 as the pathway for future land DA developmental priorities so that land DA systems
1051 will evolve into viable environments for addressing the critical science and water re-
1052 source problems facing society through the exploitation of observational information.

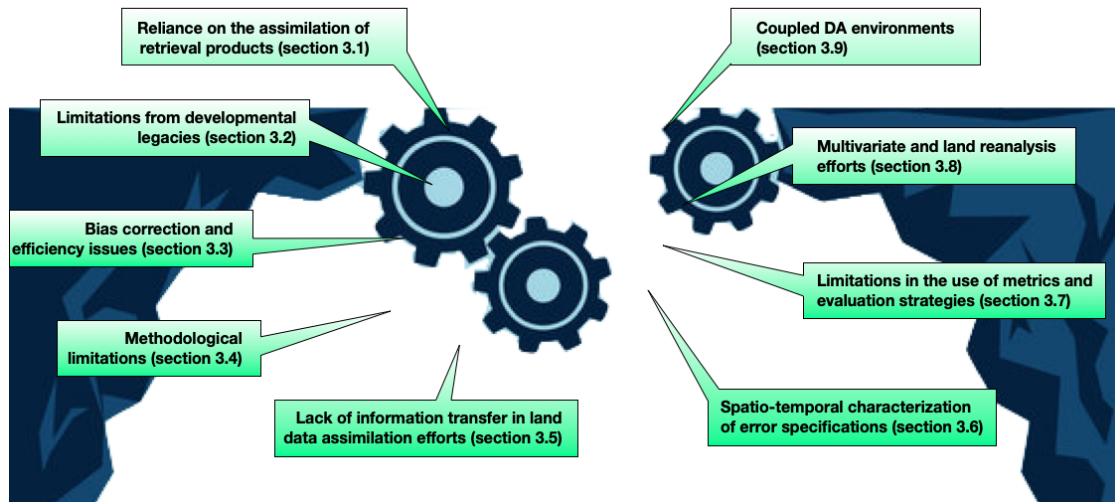


Figure 1. List of the major challenges and gaps in land data assimilation discussed in this article.

1053 **Acronyms**

1054 **DA** Data Assimilation

1055 **EO** Earth Observation

1056 **LSM** Land Surface Model

1057 **NWP** Numerical Weather Prediction

1058 **S2S** Sub-seasonal to Seasonal

1059 **GPS** Global Positioning System

1060 **GPM** Global Precipitation Measurement

1061 **SMAP** Soil Moisture Active Passive

1062 **SMOS** Soil Moisture Ocean Salinity

1063 **GRACE-FO** Gravity Recovery And Climate Experiment - Follow On

1064 **GRACE** Gravity Recovery And Climate Experiment

1065 **OCO-2** Orbiting Carbon Observatory-2

1066 **SWOT** Surface Water and Ocean Topography

1067 **EKF** Extended Kalman Filter

1068 **EnKF** Ensemble Kalman Filter

1069 **SSM/I/S** Special Sensor Microwave Imager/Sounder

1070 **SWE** Snow Water Equivalent

1071 **GLDAS** Global Land Data Assimilation System

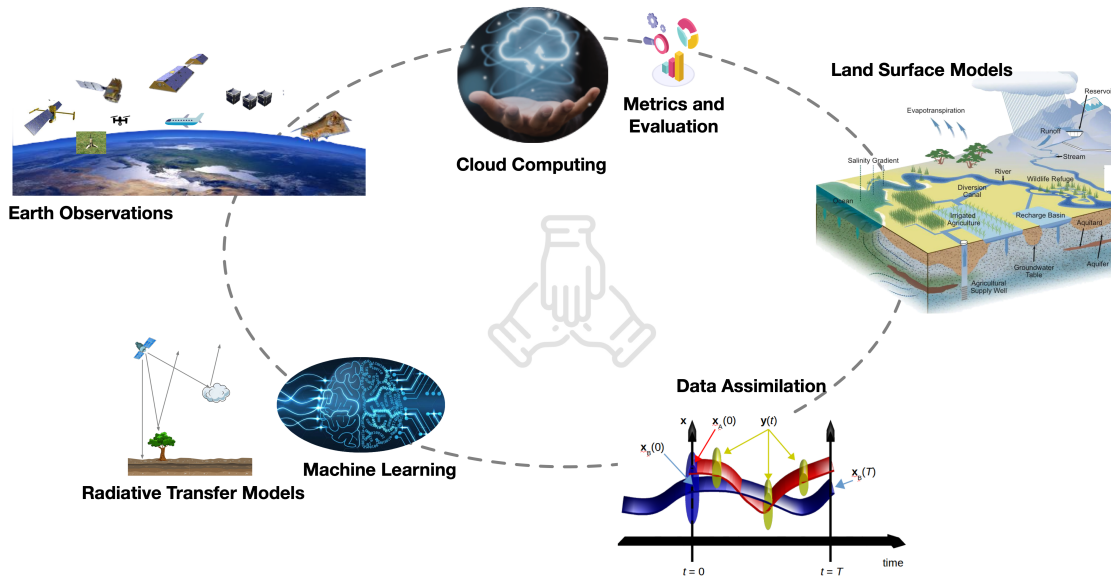


Figure 2. A vision for the next generation land DA systems, with closer developmental coordination with the LSM and observation communities, adoption of technological enhancements in machine learning and cloud computing, and advancement in the use of metrics and evaluation strategies to realize the potential of land surface observations.

1072 **NLDAS** North American Land Data Assimilation System

1073 **LIS** Land Information System

1074 **CDF** Cumulative Distribution Function

1075 **ET** Evapotranspiration

1076 **ESMF** Earth System Modeling Framework

1077 **JEDI** Joint Effort for Data Assimilation Infrastructure

1078 **OSSE** Observing System Simulation Experiment

1079 **NASA** National Aeronautics and Space Administration

1080 **USDA-ARS** United States Department of Agriculture - Agricultural Research Service

1081 **NOAA** National Oceanic and Atmospheric Administration

1082 **6 Open Research**

1083 This article is focused on discussing the challenges and priorities in the field of
 1084 land data assimilation and does not include the specific use of any particular software
 1085 or results involving specific data products.

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