

1 **Water Cycle Science Enabled by the GRACE and GRACE-FO Satellite Missions**

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18 **Abstract**

19 **Satellite observations of the time-variable gravity field revolutionized the monitoring of**
20 **large-scale water storage changes beginning with the 2002 launch of the Gravity Recovery**
21 **and Climate Experiment (GRACE) mission. Most hydrologists were skeptical of the satellite**
22 **gravimetry approach at first, but validation studies assuaged their concerns and high-**
23 **profile, GRACE-based groundwater depletion studies caused an explosion of interest. The**
24 **importance of GRACE observations for hydrologic and cryospheric science became so great**
25 **that GRACE Follow On (GRACE-FO) jumped NASA's Earth Science mission queue and**
26 **launched in 2018. A third Mass Change mission is currently under development. Here we**
27 **review key milestones in satellite gravimetry's progression from the fringes of hydrology to**
28 **being a staple of large-scale water cycle and water resources studies and the sole source of**
29 **observations of what is now an 'Essential Climate Variable', terrestrial water storage.**

30

31 Water intersects practically every sub-discipline of Earth science as well as being essential to all
32 life on Earth. The heterogeneities of the water cycle largely determine the distribution of
33 ecological systems and human populations around the world and have important implications for
34 agriculture, natural hazards, and meteorology. Therefore, it may be surprising that there was no
35 viable method for measuring large scale, seasonal to interannual fluctuations of water across the
36 continents prior to the twenty-first century. Direct and indirect observations of precipitation,
37 arguably the most important water cycle observable, have been collected since the middle of the
38 second millennium, if not earlier¹. By the late 1800s, attempts were being made to estimate the
39 balance of precipitation, evapotranspiration, and runoff from the continents. However, despite the
40 physics of infiltration, soil drainage, groundwater flow, and baseflow having been characterized a

41 century or more earlier², implying significant subsurface water storage dynamics, scientists failed
42 to acknowledge the importance of changes in water stored on and in the land as a major component
43 of continental water balance until a 1992 study constructed a global, monthly climatology of soil
44 water content³.

45 Twentieth century hydrologists largely ignored terrestrial water storage (TWS) as a variable for a
46 few reasons. First, at a time when improving general circulation and weather forecast models was
47 a primary goal of numerical modelers, accurate simulation of TWS was not considered important
48 so long as the simulated moisture state of the surface served effectively as a lower boundary
49 condition for the atmosphere. For decades, the land surface was simplified numerically to behave
50 as a 'bucket' of water beneath the atmosphere⁴. Second, save for the terrestrial water budget
51 equation (see Global Water Budget Analysis, below), hydrologists rarely considered the concept
52 of TWS; they instead were concerned with its components. Related to that, measuring large scale
53 changes in TWS using conventional methods has always been a challenge because it comprises
54 multiple components, each with its own considerations. In particular, snow water equivalent (the
55 product of snow depth and snow density) is highly variable in space, and snow tends to be deepest
56 in the least accessible locations: polar regions and on the tops of mountains⁵. Soil moisture is
57 highly variable both spatially and temporally, and measuring wetness below the surface soil layer
58 requires digging a hole or inserting a long probe using a hydraulic press⁶. In situ soil moisture
59 monitoring networks are highly concentrated in the northern midlatitudes, providing very little
60 coverage outside of 30°N to 60°N⁷. Lake and river channel water storage is relatively easy to
61 estimate, but only if measurements are recorded and shared. Outside of the wet tropics, surface
62 water is typically a small component of TWS change⁸. Vegetation water storage changes are
63 typically much smaller than soil moisture changes⁹. Finally, the standard method for monitoring

64 groundwater storage variations is to drill down into an aquifer, install a piezometer, and record
65 water level or piezometric head variations. Owing to cost, logistics, and political impediments
66 (most nations do not share what groundwater measurements they have¹⁰), it would be impossible
67 to establish globally representative, automated networks with sufficient density to monitor any of
68 these variables satisfactorily. Even the USGS's near-real time groundwater monitoring network
69 has major gaps in the western U.S. (Figure 1). The International Groundwater Resources
70 Assessment Centre (IGRAC), which aspires to serve as a clearinghouse for global groundwater
71 data, houses records from fewer than 35 countries representing far less than half the world's land
72 surface (Figure 1). While numerous additional groundwater observational records exist, most are
73 either restricted from distribution or not digitized and centralized. Given these challenges and
74 despite its true significance as an indicator of water resources availability and hydroclimatic
75 change, TWS quantification and understanding would have remained in limbo absent the
76 introduction of a game-changing observational advancement.

77

78 Satellite Gravimetry

79 Geodesists have long recognized that both static features and temporal variations in Earth's gravity
80 field, which reflect heterogeneities in the distribution of solid, liquid, and atmospheric mass,
81 perturb the paths of Earth-orbiting satellites in predictable ways¹¹⁻¹⁴. In the 1960s they conceived
82 the idea of measuring these orbit perturbations precisely in order to infer time varying mass
83 variations in the land, ocean, and atmosphere¹⁵. Satellite laser ranging (i.e., precise tracking of the
84 distances from ground stations to an orbiting satellite) enabled the first global mapping of Earth's
85 gravity field and continues to provide time varying gravity information which is superior to that
86 of GRACE/FO at the largest scales^{14,16}. A compendium of studies¹⁷ laid the scientific foundation

87 for launching a dedicated, time-variable gravimetry mission, demonstrating the potential value for
88 hydrology and other disciplines and describing possible satellite instruments and configurations
89 for such a mission. The preferred architecture comprised identical twin satellites in a polar, low
90 Earth orbit, one following the other with 100-220 km separation. Orbital perturbations caused by
91 gravitational variations would be quantified using micron-level tracking of the intersatellite
92 distance (range) by microwave interferometers aboard both satellites. Non-gravitational
93 accelerations, such as those caused by atmospheric drag, would be measured by highly precise
94 accelerometers. Soon thereafter, GRACE, which was jointly supported by the German space
95 agency, was selected by NASA to be one of its first Earth System Science Pathfinder projects.

96 A seminal pre-launch study¹⁸ introduced a technique for translating gravimetry measurements
97 from GRACE into changes in terrestrial water storage (TWS) and provided the equations necessary
98 to extract region-specific TWS anomaly time series from global gravity field solutions expressed
99 using spherical harmonics (a series of functions that describe anomalies on a sphere, in which
100 higher degrees and orders represent progressively smaller constituents of the anomalies).
101 Subsequent studies^{19,20} explored this potential in depth, comparing anticipated uncertainty in the
102 TWS retrievals with model- and in situ observation-based estimates of TWS changes at various
103 spatial scales. They concluded that the approach would be viable depending on the size of the
104 region and the amplitude of the TWS variations. It was also demonstrated that changes in
105 groundwater storage could be isolated from the TWS data, given modeled or observed estimates
106 of soil moisture and other TWS components as needed²¹.

107

108 GRACE Mission – The Early Years

109 GRACE²² launched on 17 March 2002 to the delight of the geodesy, cryosphere, and ocean science
110 communities. Cryospheric scientists used GRACE to generate startling new estimates of the rates
111 of ice losses from Greenland, Antarctica, and the Alaskan glaciers²³⁻²⁶. Oceanographers used it to
112 improve their understanding and quantification of sea level changes²⁷⁻²⁹. Meanwhile, most
113 hydrologists, if they had heard of the mission at all, fretted that GRACE's spatial (>150,000 km²
114 at mid-latitudes^{30,31}) and temporal (monthly) resolutions were too coarse, and they were befuddled
115 by its non-instantaneous, vertically integrated TWS anomaly data. Further, non-geodesists were
116 concerned that the retrievals could be contaminated by unknown or poorly modeled sources of
117 mass change in Earth's crust and mantle.

118 Discord and controversy in the hydrology community are not unusual, so perhaps the initial
119 resistance to GRACE is not surprising. Multiple sub-disciplines compete for primacy and funding,
120 and GRACE was likely viewed as additional competition, for example by soil moisture enthusiasts
121 who, at the time, were advocating for their own mission. Further, the characteristics of GRACE-
122 based TWS data did not lend themselves to comparisons with or incorporation into grid-based
123 hydrological models. The developers of these models were already busy evaluating and arguing
124 the merits of global versus regional modeling, physics-based versus conceptual modeling, the use
125 of empirically based parameters versus calibration, and complex, high resolution, computationally
126 expensive models versus simplified approaches.

127 While hydrologists were certainly unfamiliar with data products having characteristics like those
128 of GRACE, their concerns were largely unfounded. Mass changes attributable to viscoelastic
129 mantle and crustal deformation occur over long enough timescales that they are non-factors in
130 monthly to interannual terrestrial mass change, with two notable exceptions. First, glacial isostatic
131 adjustment (GIA) can be significant in polar and near-polar regions that were covered by ice during

132 the last glacial maximum. The gravitational effects can be removed reasonably well³² using
133 models of GIA^{33,34} although errors in these models can be an issue for hydrological studies in
134 regions with significant rates of GIA³⁵. Second, major earthquakes (magnitude > 8) produce both
135 step function (coseismic) and exponential (postseismic) signals in mass anomaly time series,
136 whose removal from the GRACE data is further complicated by their wavelike spatial patterns³⁶.
137 Fortunately, only a handful of major earthquakes have occurred in the GRACE era, and their
138 effects are limited in time and space³⁷. In addition to redistribution of TWS, the other main drivers
139 of mass and gravitational change are circulations of the atmosphere and oceans. Both of these are
140 modeled and removed from the observed gravity signal when TWS is the target observable, but
141 errors stemming from the models continue to be a major component of the error budget of the
142 derived TWS anomalies^{18,19,38}. Within the terrestrial mass changes that remain, seasonal and
143 interannual changes in vegetation biomass have been shown to be at or below GRACE's
144 uncertainty threshold⁹. That leaves groundwater, soil moisture, surface waters, snow, ice, and
145 permafrost as the primary components of TWS retrieved from GRACE gravity observations. Of
146 those, seasonal variability is dominated by surface water in the wet tropics, by snow in alpine and
147 high latitude regions, and by soil moisture most elsewhere⁸. Transitioning across interannual to
148 decadal timescales, soil moisture's transient influence wanes and groundwater and ice trends
149 eventually dominate the changes, save for loss of water from major surface water bodies such as
150 the Aral and Caspian Seas^{39,40} and filling of manmade reservoirs⁴¹.

151 Despite the skepticism, a handful of researchers described and evaluated new hydrological
152 applications for GRACE, such as the use of GRACE-based TWS observations for closing the water
153 budget and estimating regional mean evapotranspiration⁴² (ET) and net water flux⁴³ (precipitation
154 minus ET) at the land-atmosphere interface. Also introduced were methods for using GRACE and

155 other data to estimate river discharge^{44,45} and high latitude seasonal snow mass⁴⁶. GRACE TWS
156 data revealed an anti-correlation of TWS variations during 2002-2006 in the Amazon and Congo
157 River basins⁴⁷. They were employed to evaluate the water storage variations simulated by climate
158 models⁴⁸, while numerous other studies evaluated the GRACE TWS time series using ground-
159 based observations and models⁴⁹⁻⁵². During this same period, two studies^{53,54} demonstrated the
160 ability to isolate groundwater storage changes from GRACE TWS data using in situ or modeled
161 soil moisture and snow water equivalent. The first GRACE data assimilation experiments were
162 performed⁵⁵ enabling the TWS components to be separated while also achieving spatial and
163 temporal downscaling of the coarse resolution GRACE observations. Nevertheless, only a tiny
164 fraction of the hydrology community embraced GRACE during its early years.

165

166 Onset of the era of GRACE hydrology

167 A major impediment to the uptake of GRACE data by hydrologists was the necessity of enlisting
168 geodesists to process the level-2 GRACE data (i.e., gravity field solutions) and derive TWS time
169 series. The monthly level-2 gravity field solutions were delivered as sets of spherical harmonic
170 coefficients, and few, if any, hydrologists were familiar with that mathematical construct, capable
171 of applying destriping techniques to remove correlated errors⁵⁶, and able to apply Gaussian
172 averaging kernels⁵⁷ for computing regional time series and associated uncertainty estimates.

173 That situation changed when the first 1° gridded GRACE TWS anomaly fields⁵⁸ began to be
174 distributed. They included numerous caveats and warnings; for example, the method used to
175 derive the fields was optimized for the ocean, not for the land. Further, users were admonished
176 that a TWS time series for an individual 1° grid cell was meaningless on its own; it was necessary

177 to average the gridded data over sufficiently large regions in order to overcome spatial correlation
178 errors associated with the data processing (Gaussian smoothing). In fact, geodesists' fear of such
179 products being misused was one of the main reasons they had not previously been developed.
180 Nevertheless, the original gridded product freed hydrologists to perform GRACE-based studies on
181 their own, and as a result, the number of such studies began to grow.

182 Over time, geodesists became increasingly uncomfortable with the way the first gridded product
183 was being used and the failure of many hydrologists to acknowledge its deficiencies when drawing
184 conclusions. In particular, gravity signal "leakage" from adjacent, large, and imperfectly modeled
185 gravitational phenomena such as ocean tides had not been minimized in the gridded product
186 through advanced geodetic processing or properly accounted in hydrologists' error budgets, and
187 the Gaussian smoothing and destriping caused damping of the dynamic range of TWS that was
188 particularly troublesome for water balance studies. Further, geodesists argued that a consistent
189 comparison between modeled or independently observed hydrological phenomena and GRACE
190 TWS anomalies required the former to be converted to spherical harmonics and Gaussian
191 smoothed in the same way as the latter. Hydrologists saw little practical benefit to this approach
192 beyond validation, and argued that gridded or otherwise preprocessed TWS products were
193 imperative for maximizing GRACE data uptake and enabling the mission to reach its potential as
194 a tool for hydrology. Discussions between these two factions at the Second GRACE Hydrology
195 Workshop⁵⁹ in November 2009 sparked the development of a gridded TWS anomaly product that
196 was better suited for land hydrology and was accompanied by a set of scale factors intended to
197 counteract the signal damping⁶⁰. Just as importantly, data and a short routine were provided for
198 estimating uncertainty in the resulting monthly TWS anomalies when averaged over a region of
199 interest. This uncertainty decreases dramatically as the size of the region increases, and it is also

200 much smaller near the poles, where the satellite ground tracks are close, than near the equator.
201 Since then, many other geodesy teams have introduced their own level-3 TWS products, catering
202 to the needs of hydrologists⁶¹⁻⁶⁵.

203 Two headline-making publications^{66,67} helped to dissolve hydrologists' remaining hesitancy and
204 pique their interest in GRACE. They described shocking rates of groundwater depletion occurring
205 in northern India (Figure 2), driven by groundwater extractions to support irrigated agriculture in
206 a semi-arid climate. While it was known from local reports that groundwater levels were declining,
207 the intensity and scale of declines had not previously been quantified due to the inaccessibility of
208 in situ observations. Considering that over 100 million live in the affected region and that its
209 agricultural output underpins the food supply for all of India, the potential consequences of
210 dewatering the aquifers cannot be overstated. The two publications and the extensive media
211 coverage surrounding them drew worldwide attention to the issue and engendered scores of follow-
212 on studies⁶⁸⁻⁷¹. According to *Web of Science*, in the prior decade (1999-2008) there had been 174
213 publications with both "groundwater" and "India" in the title. There were 832 in the decade that
214 followed. Scientists who were previously unconvinced began to comprehend the power of satellite
215 gravimetry as a tool for seeing "beneath the surface" in a way that no other approach could,
216 providing unique and essential value for the field of hydrology. Figure 3 illustrates the steady
217 growth of journal publications on GRACE and GRACE FO (hereafter GRACE/FO) hydrology
218 since 2002. Clearly, there are far too many to cite here. What cannot be inferred from Figure 3 is
219 that, prior to around 2010, a majority of these publications were written by non-hydrologists.
220 Similarly, while early GRACE-based hydrological studies tended to focus on comparison and
221 evaluation, starting around 2009, as discussed below, an increasing number of studies explored
222 scientific and practical applications.

223

224 Contributions of GRACE to the study of groundwater

225 Following the formulas of the original Indian groundwater depletion studies^{66,67}, researchers
226 investigated other regions where notable TWS declines were seen in the GRACE data, often due
227 to groundwater pumping. In particular, GRACE was applied to provide new estimates of known
228 groundwater depletion in California's Central Valley^{72,73}. Considering that about one quarter of
229 the food consumed in the U.S. comes from the Central Valley, monitoring and preserving its
230 groundwater is crucial. The GRACE-based studies and associated media coverage once again
231 drew renewed attention to groundwater depletion, likely contributing to the passage of California's
232 2014 Sustainable Groundwater Management Act. Several groups focused on another critical
233 source of water for irrigated agriculture in the U.S., the High Plains aquifer⁷⁴⁻⁷⁶. Over the past two
234 decades, this aquifer, which supports "the breadbasket of America", has been stable or gaining
235 water in the north, while water stored in the central and southern portions has been slowly and
236 steadily declining since records began in the middle of the 20th century⁷⁷. Others quantified
237 groundwater losses in the North China plain⁷⁸⁻⁸⁰, another important agricultural region where
238 withdrawals for irrigation have been outpacing recharge. Several studies discussed substantial
239 decreases in surface water and groundwater in the Middle East caused by a combination of drought
240 and irrigated agriculture⁸¹⁻⁸⁴. Scientists also evaluated groundwater storage variations and
241 depletion in the Great Lakes basin⁸⁵, the Colorado River basin⁸⁶, and across Africa⁸⁷⁻⁹⁰, among
242 other regions. Still other authors began to develop global assessments of groundwater storage
243 changes based on GRACE data^{40,91-94}. Figure 4 displays a global map of 20-year trends in TWS
244 based on GRACE and GRACE-FO observations⁶¹. TWS declines in several of the regions, such
245 as those previously noted, can be attributed at least in part to groundwater depletion to support

246 irrigated agriculture, in some cases exacerbated by recent drought⁴⁰. Finally, as the only truly
247 global source of TWS observations, GRACE first made possible a sub-section on groundwater and
248 terrestrial water storage in the *Bulletin of the American Meteorological Society*'s annual "State of
249 the Climate" report in 2011⁹⁵, and it has been included ever since.

250

251 Global Water Budget Analysis

252 Hydrologists have aspired to close the water budget (see equation 1, below) from global to
253 catchment scales for more than a century^{2,96,97}. However, prior to the launch of GRACE, with few
254 exceptions⁹⁸ water budget analyses relied on the assumption that changes in TWS are insignificant
255 on annual and longer timescales. This assumption was enticing because the collocated, in situ
256 observations of groundwater, soil moisture, and, in some cases, snowpack and surface water
257 necessary to estimate interannual changes in TWS were extremely scarce²⁰. Analysis of the first
258 few years of GRACE data provided ample evidence that the assumption of annual steady state of
259 TWS was flawed at both river basin and global scales, and that interannual changes in TWS caused
260 compensating changes in mean sea level⁹⁹⁻¹⁰¹.

261 In addition to the change in storage, ΔTWS , closing the terrestrial water budget requires estimates
262 of the net fluxes into and out of a study region (often a catchment or river basin bounded by
263 topography across which lateral flows are negligible):

264

$$265 \quad \Delta TWS = P - ET - Q \quad (1)$$

266

276 where P, ET, and Q are accumulated precipitation, evapotranspiration, and runoff. Given
277 knowledge of three of the variables, the fourth can be estimated as a residual. Early GRACE-
278 based studies estimated ET, which is difficult to measure directly, as the residual^{42,102}. However,
279 because ΔTWS represents the change between instantaneous TWS values at the start and end of
280 the flux accumulation period, while GRACE/FO provides monthly mean TWS anomalies, the
281 approach requires either complex mathematics⁴², a multi-month moving window approach¹⁰³, or
282 higher tolerance for error. Similarly, a combined atmospheric-terrestrial water budget can be used
283 to estimate atmospheric convergence (P-ET) as a residual:

275

$$276 \quad P - ET = Q + \Delta TWS + \Delta AMC \quad (2)$$

277

278 where ΔAMC is the change in total atmospheric moisture content, which is typically assumed to
279 be negligible⁴³. The resulting ET and P-ET estimates have proven useful for evaluating modeled
280 and remote sensing-based fluxes^{104–109}, at least in terms of magnitude or bias. The approach even
281 helped to resolve a scientific debate. Hydrologists had disagreed on the strength of the seasonal
282 cycle of ET in tropical regions, where water and energy are typically abundant throughout the
283 year^{110,111}. The observed seasonality of TWS provided by GRACE was the key innovation that
284 enabled the matter to be put to rest, as the water budget residual ET indicated that the seasonal
285 cycle in Brazil and central Africa was relatively weak and less dynamic than that depicted by most
286 models¹⁰⁵. Unfortunately, the scarcity of river discharge observations (from basins large enough
287 for GRACE/FO to resolve) limits the applicability of this approach. That's one reason scientists
288 use GRACE/FO TWS data together with atmospheric analysis model-based convergence to

289 estimate river discharge as a water budget residual^{44,45}. However, the long term mean magnitude
290 of the resulting estimates is almost entirely controlled by the simulated convergence, with
291 GRACE/FO TWS primarily acting to modulate the seasonal cycle.

292 Another innovative application of the water budget approach is to estimate accumulated snowfall
293 in high latitude or alpine regions. This can be accomplished with the aid of modeled ET and
294 modeled or observed Q. Both are generally small in comparison with P at high latitudes/altitudes
295 during the winter prior to spring thaw, such that errors in a water budget residual are likewise
296 small. In other words, ΔTWS is mostly attributable to P in equation 1. The resulting regional
297 scale, monthly estimates of snowfall are valuable because precipitation is under-sampled in cold
298 regions and snowfall is notoriously challenging to measure due to gauge undercatch, which itself
299 is difficult to estimate. The study that introduced this approach determined that the undercatch
300 corrections in one popular, global precipitation dataset were too large¹¹². However, other GRACE-
301 based studies have shown that precipitation products underestimate cold season precipitation by
302 30-50% in the Himalayas¹¹³ and that undercatch corrections applied in well-known products are
303 both necessary and highly variable in the Arctic¹¹⁴. Another concluded that four meteorological
304 datasets generally underestimated P even after undercatch corrections had been applied¹¹⁵.

305 When estimates of all the budget variables are available, they can be used within a water balance
306 framework to constrain each other, thereby improving each individual estimate. Among the first
307 to apply this approach with GRACE data, a 2009 study estimated Mississippi River discharge as
308 a water budget residual, showing that it significantly exceeded gauged discharge, and used the
309 results to bias-correct satellite-derived precipitation¹¹⁶. A follow-up study over ten global river
310 basins similarly attributed most of the non-closure error to the precipitation data¹¹⁷. In 2015, a
311 multi-institutional investigation used observation-based datasets including GRACE data within a

312 variational framework to enforce simultaneous water and energy budget closure over continents,
313 ocean basins, and worldwide, on monthly mean and annual mean (2001-2010) timescales^{118,119}.
314 The approach successfully reduced bias in all of the estimated fluxes and storage changes. Results
315 showed that unconstrained annual terrestrial and atmospheric water budgets closed with
316 considerably less than 10% imbalance (relative to precipitation) in most cases, however, on a mean
317 monthly timescale, water budget imbalance approached or exceeded 20% in North America,
318 Eurasia, and Australia. Building on this technique, studies with larger suites of input datasets
319 produced gridded, global results¹²⁰ and results over longer periods¹²¹. Rather than constraining
320 the flux estimates, other investigators have computed the mass imbalance arising from different
321 combinations of inputs to the water budget equation as a way to help select appropriate
322 combinations of datasets for regional hydrological studies¹²².

323 Because the Earth is a closed system, it has long been known that changes in the mass of water
324 stored in the oceans, land, and atmosphere must sum to zero. However, prior to GRACE, there
325 was no way to perform an observation-based accounting. The literature is rich with GRACE-
326 informed studies of the responses of ice sheets and sea level to climate change^{123,124}. GRACE
327 similarly enabled quantification of non-ice sheet, terrestrial water contributions to seasonal and
328 interannual variations in sea level. A 2016 study demonstrated the use of GRACE data over land
329 and ocean in a global least-squares inversion, together with ocean altimetry and steric data from
330 ocean reanalyses and other data sets, to constrain the sea level budget including the TWS
331 contribution to sea level¹²⁵. It has also been shown that non-ice TWS largely controls seasonal
332 and interannual variations of the land-ocean water balance¹²⁶. Another study estimated that, during
333 2002-2014 at the global scale, hydroclimate related increases in non-ice TWS acted to reduce the
334 rate of sea level rise by about 15% after removing the effects of irrigation-enhanced groundwater

335 depletion¹²⁷. It was concluded that this was consistent with regional increases in precipitation that
336 drove flooding and drought recovery events in several locations globally. Ironically, in the year
337 after that study ended there was a huge decrease in non-ice sheet TWS (Figure 5), and it has not
338 recovered to its earlier dynamic range¹²⁸. While it is possible that extreme glacier melt accounted
339 for the sudden decline, recent analyses do not implicate that as the primary driver^{129,130}, leaving
340 liberation of unfrozen forms of TWS as the likely culprit. Supporting that hypothesis, a subsequent
341 study concluded that TWS in endorheic basins (which are often semi-arid to arid) declined by
342 about 106 Gt/yr during 2002-2016¹³¹ – water which would ultimately enter the ocean. TWS
343 fluctuations cause seasonal mass variations in sea level, whose annual range of about 34 mm is
344 comparable to about a decade of long-term barystatic level rise¹³². Non-secular interannual
345 variations in sea level also are largely driven by irregular gains and losses of TWS around the
346 world. For example, a powerful El Nino in 2011 caused an excess of TWS in the southern
347 hemisphere, including record flooding Australia, which in turn produced a temporary decline in
348 global mean sea level^{101,133}. Prior to GRACE, explanations for such a decline would have derived
349 from circumstantial evidence.

350 Because shifts in the water cycle will be among the most visible and consequential effects of
351 climate change, understanding its variations and predicting how it is likely to change in the future
352 are key goals for hydrologists and climate scientists¹¹⁸. A study of the first eight years of GRACE
353 data showed that observed TWS variations were consistent with known wet/dry precipitation
354 patterns associated with El Nino and La Nina¹³⁴. Another study found evidence in the GRACE
355 data for water cycle acceleration, which has long been predicted as a consequence of climate
356 change¹³⁵. Certain regional trends observed by GRACE/FO (Figure 4) have similarly been flagged
357 as possible climate change impacts⁴⁰. Further, global TWS itself seems to govern atmospheric

358 carbon dioxide growth, with faster growth in drier years¹³⁶, revealing one of possibly many TWS-
359 climate feedback loops.

360

361 Integration of GRACE data into hydrological models

362 Early in the mission, hydrological models were frequently used to validate the retrieval of TWS
363 changes from the GRACE gravity observations^{57,137-139} and to refine retrieval algorithms^{140,141}.

364 Later, such models helped identify the processes controlling the observed TWS changes^{8,52,142,143}.

365 As hydrologists became satisfied that GRACE was, in fact, measuring what the geodesists claimed,
366 the retrieved TWS began to be used to validate hydrological models¹⁴⁴⁻¹⁴⁶. For example, it was
367 shown that many land surface models, particularly those lacking a groundwater variable, have
368 insufficient storage capacity to represent the true, full dynamic range of TWS^{147,148}.

369 While most GRACE-based studies have relied on in situ observations, model output, or
370 simplifying assumptions to isolate variations in groundwater or the other components of TWS,
371 scientists recognized that auxiliary data and hydrological models could be used more holistically
372 to extract groundwater information from GRACE observations. In particular, a Kalman smoother
373 based data assimilation approach⁵⁵ was introduced for integrating GRACE derived TWS data into
374 a land surface model, thus enabling spatial and temporal downscaling and vertical disaggregation
375 of the TWS components. Just as importantly for practical, operational applications, data
376 assimilation permitted extrapolation to near-real time based on the meteorological inputs⁵⁵,
377 whereas standard GRACE TWS products were typically released 2-5 months after the time of
378 observation.

379 Subsequent efforts developed techniques for assimilating GRACE data into different land surface
380 models^{149–154}. Others improved on the original Kalman smoother data assimilation by enabling
381 assimilation of gridded (as opposed to basin average) GRACE data^{155,156}. This was important
382 because it minimized spatial artifacts associated with pre-defined regions (i.e., river basin
383 boundaries) and better utilized the spatial information contained in the data products. Still, it was
384 shown that the lack of vertical stratification in the GRACE data could lead to mischaracterization
385 of subsurface processes¹⁵⁷ which further could cause errors in the assimilated output¹⁵⁸.
386 Simultaneous assimilation of GRACE TWS, Soil Moisture and Ocean Salinity (SMOS) or Soil
387 Moisture Active Passive (SMAP) soil moisture, Moderate Resolution Imaging Spectroradiometer
388 (MODIS) or Advanced Microwave Scanning Radiometer (AMSR) snow, and other remote sensing
389 data holds promise for overcoming such issues while also improving horizontal downscaling^{159–}
390 ¹⁶⁴. As an alternative to data assimilation, GRACE data have also been used to inform and calibrate
391 hydrological models^{165–169}.

392

393 Practical Applications

394 Monitoring drought was recognized early in the mission as a practical application of GRACE data,
395 though multiple years of observations were needed to develop a baseline for quantifying wet or
396 dry extremes^{170–173}. GRACE data assimilation was the foundation for the first routinely delivered
397 GRACE-based soil moisture and groundwater drought/wetness indicator maps for the contiguous
398 United States¹⁷⁴. These indicators were presented as wetness percentiles relative to the range of
399 conditions at each specific location and time of year in a 60+ year historical model simulation.
400 This enabled GRACE data to become useful for operational drought monitoring, and the indicator
401 maps began to be consulted by the authors of the U.S. Drought Monitor¹⁷⁵. In addition to their

402 improved spatial resolution, a key to the adoption of GRACE data assimilation products by
403 drought, water resources, and agricultural decision-makers was their weekly, near-real time
404 availability^{175,176}. Few operational applications can make use of data older than about 2 weeks¹⁷⁷.
405 GRACE/FO data assimilation was later expanded to cover all continents except Antarctica, serving
406 as the basis for weekly, global, root zone soil moisture and groundwater wetness/drought indicator
407 maps¹⁰, which were unprecedented (Figure 6). One-to-three month forecasts of the same, for the
408 contiguous U.S., are also available, initialized by GRACE-FO data assimilation¹⁷⁸. Several
409 GRACE-only drought indices have been proposed as well, mostly quantifying drought relative to
410 the historical range of season- and location-specific conditions¹⁷⁹⁻¹⁸⁴.

411 Wet extremes can be detected and quantified in the same ways as dry extremes, and some of the
412 indicators described above are meant to serve both purposes^{10,174,180}. Using GRACE, studies have
413 investigated major flood events around the world¹⁸⁵⁻¹⁸⁷. Further, owing to the importance of
414 elevated water levels as a precursor to severe flooding, scientists have introduced means for
415 assessing flood vulnerability based on the current TWS anomaly relative to the water holding
416 capacity of the land inferred from the GRACE data record^{188,189}. It has also been demonstrated
417 through a case study in the Missouri River basin that snow water equivalent, soil moisture, and
418 groundwater output from a GRACE data assimilating land surface model¹⁷⁴ have the potential to
419 improve the predictability of flood events¹⁹⁰.

420 Ongoing GRACE/FO research may create additional opportunities for practical applications that
421 could be developed in the future for user-specific needs. For instance, application of the water
422 budget approach for ET has been used to estimate water consumption during the summer season
423 in the upper and lower Colorado River Basins¹⁹¹. Remote detection of water usage would be
424 valuable for agricultural monitoring and optimization purposes. As another example, GRACE/FO

425 data assimilation-based soil moisture estimates show promise as potential predictors of fire season
426 intensity and burned area¹⁹² and could perhaps be incorporated into an operational framework for
427 wildfire early warning¹⁹³. Similarly, such estimates have shown promise for predicting slope
428 stability and landslide vulnerability¹⁹⁴. Finally, groundwater assessments that utilize GRACE/FO
429 other remote sensing and in situ observations could be used for operational groundwater
430 monitoring or even groundwater change prediction^{195–197}.

431

432 Future Prospects

433 In contrast to the skepticism that many scientists had about GRACE when it launched in 2002, by
434 the 2010s its importance as a tool for studying the water cycle, cryosphere, and oceans had become
435 so clear that NASA, together with its partner, the German Space Agency, prioritized launching
436 GRACE-FO above several other mission concepts that had been recommended by a 2007 U.S.
437 National Research Council report¹⁹⁸. Moreover, the subsequent decadal report¹⁹⁹ ranked a third
438 Mass Change (satellite gravimetry) mission among its top five Earth science mission priorities for
439 the following decade, in order to avoid a gap in the data record and to potentially improve upon
440 GRACE/FO in terms of spatial resolution, temporal resolution, and/or accuracy. GRACE-FO
441 added an experimental laser interferometer to increase the precision of the inter-satellite range
442 measurements (above that of the primary instrument, a microwave interferometer). The laser has
443 been a success, however, the principal source of error limiting improved resolution at a given level
444 of accuracy is spatial-temporal undersampling of high frequency mass variations in the atmosphere
445 and oceans²⁰⁰ (i.e., “aliasing”). Overcoming aliasing is expected to require a constellation of two
446 or more GRACE-like satellite pairs¹⁷⁷. While that may not materialize in the immediate future,
447 extension of the satellite gravimetry-based TWS data record by GRACE-FO and a subsequent

448 Mass Change mission will support improved understanding of water cycle variability and change
449 associated with natural variations, climate change, direct human impacts, and even implementation
450 of water management policy²⁰¹.

451

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961

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968

969 **Author Contributions**

970 M.R. prepared the review article outline. M.R. and J.T.R. designed the figures and wrote the
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972

973 **Competing Interest Declaration**

974 The authors claim no competing financial or non-financial interests.

975

976 **Additional Information**

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979

980 **Figure Captions**

981 Figure 1. Inadequacy of in situ groundwater observation availability. Global map: locations of
982 groundwater observation wells whose records are archived by the International Groundwater
983 Resources Assessment Centre. Inset map: locations of the U.S. Geological Survey's Groundwater
984 Climate Response Network wells.

985 Figure 2. Terrestrial Water Storage depletion in India. Groundwater withdrawals to support
986 irrigated agriculture in a large region of northern India (corresponding to large rates of TWS
987 decline indicated by warm colors in panel a) have caused TWS to be depleted over the past 20+
988 years (time series in panel b) as observed by the GRACE (blue) and GRACE-FO (dark yellow)
989 satellite missions. Values are reported as equivalent height of water (cm) relative to the long-term
990 mean.

991 Figure 3. Growth of GRACE hydrology publications. Yearly numbers of journal publications
992 containing in their abstracts either "Gravity Recovery and Climate Experiment" or "GRACE
993 satellite" and "water" or one of several other hydrology related terms. Source: Web of Science.

994 Figure 4. Global trends in terrestrial water storage. Mean rates of change of TWS (cm yr⁻¹) based
995 on GRACE and GRACE-FO observations from April 2002 to May 2022. The map was smoothed
996 with a 150 km radius Gaussian filter for the purpose of visualization, but the base product has a
997 native 3° resolution.

998 Figure 5. Global terrestrial water storage anomalies over time. Global average, non-seasonal
999 terrestrial water storage anomalies from GRACE (blue) and GRACE-FO (dark yellow), in cm
1000 equivalent height of water, relative to a 2003-2020 mean baseline. Mascons containing TWS
1001 declines in Antarctica, Greenland, the Gulf Coast of Alaska, and polar islands associated with ice
1002 sheet and glacier losses were excluded from the average.

1003 Figure 6. Global drought and wetness monitoring products. Wetness percentiles on 25 July 2022
1004 (relative to the period 1948-2012) for (a) surface soil moisture, (b) root zone soil moisture, and (c)
1005 shallow groundwater, based on output from a GRACE and GRACE-FO data assimilating land
1006 surface model. As an example, a value of 5 indicates that the location was drier than present only
1007 5% of the time in this week of the year during 1948-2012. The maps are generated weekly and
1008 made available from <https://nasagrace.unl.edu/>.

1009

1010 **Methods**

1011 Figures 2, 4, and 5 were created using the Jet Propulsion Laboratory's (JPL) monthly mass 3-
1012 degree gridded mascon RL06 version 02 dataset⁶¹ from April 2003 to April 2022 (December 2021

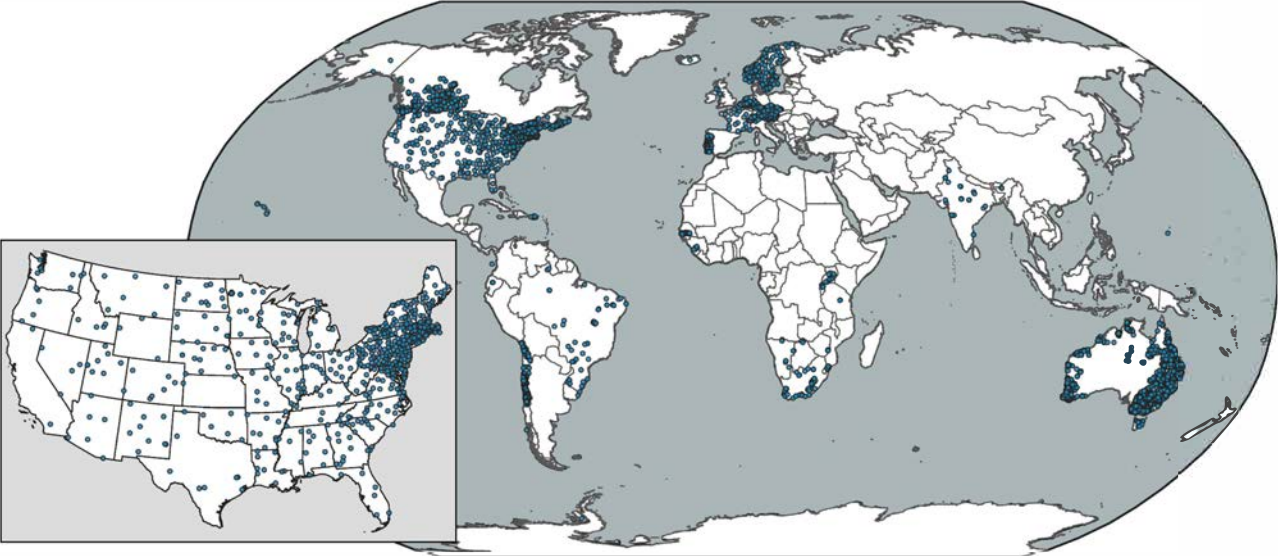
1013 in the case of Figure 5). For Figures 2a and 4, the data were smoothed using a 300 km Gaussian
1014 to improve the visualization. For Figure 2b, the data were averaged over a previously defined area
1015 of TWS decline in northern India⁴⁰ while simultaneously fitting a climatology and a trend to the
1016 timeseries at each pixel. For Figure 5, the data were averaged over the global land surface
1017 excluding Antarctica, Greenland, the Gulf Coast of Alaska, and polar islands, and the mean
1018 seasonal cycle was removed¹²⁸.

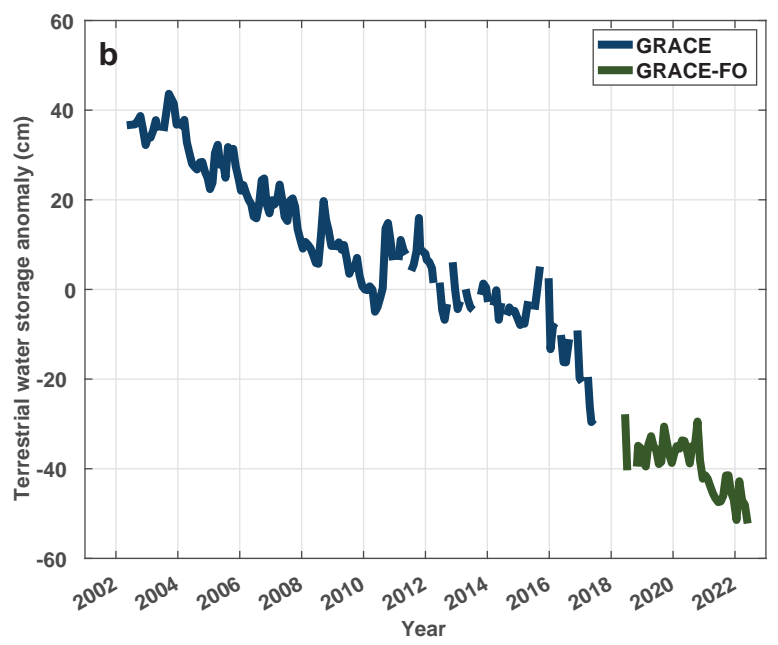
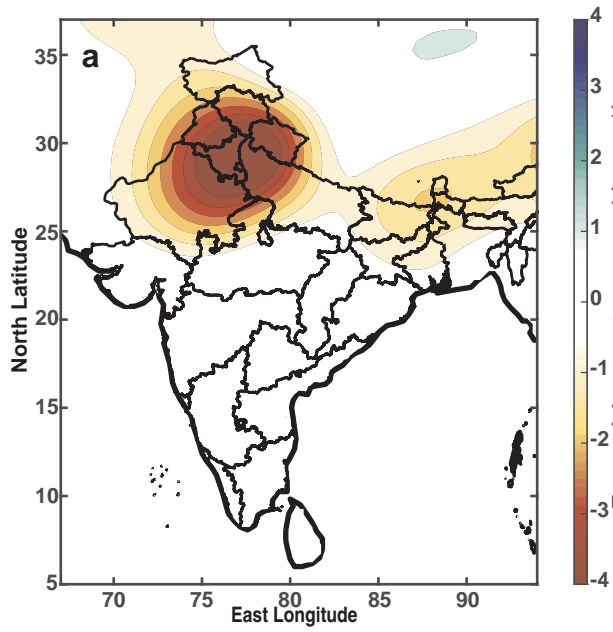
1019

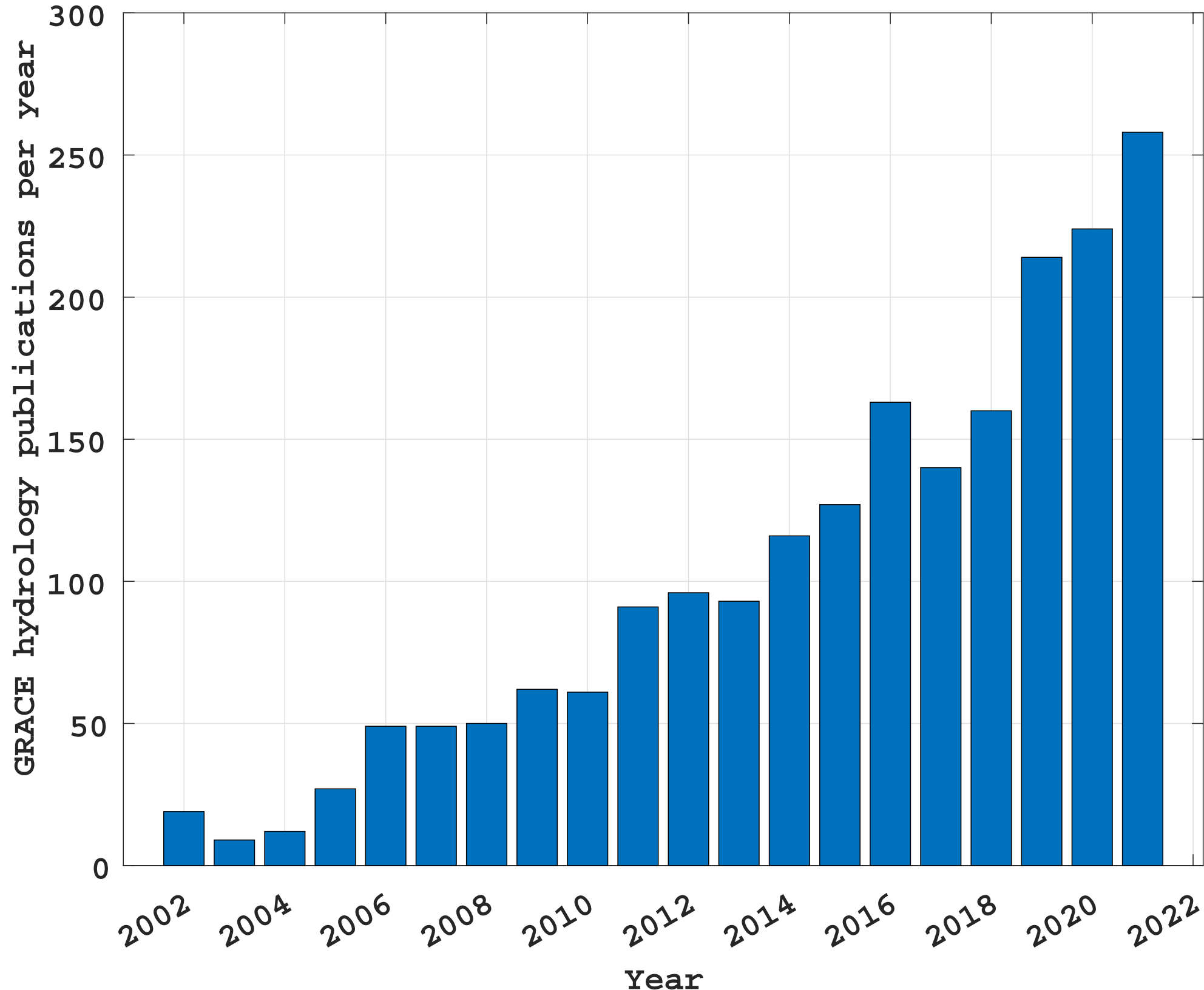
1020 **Data Availability**

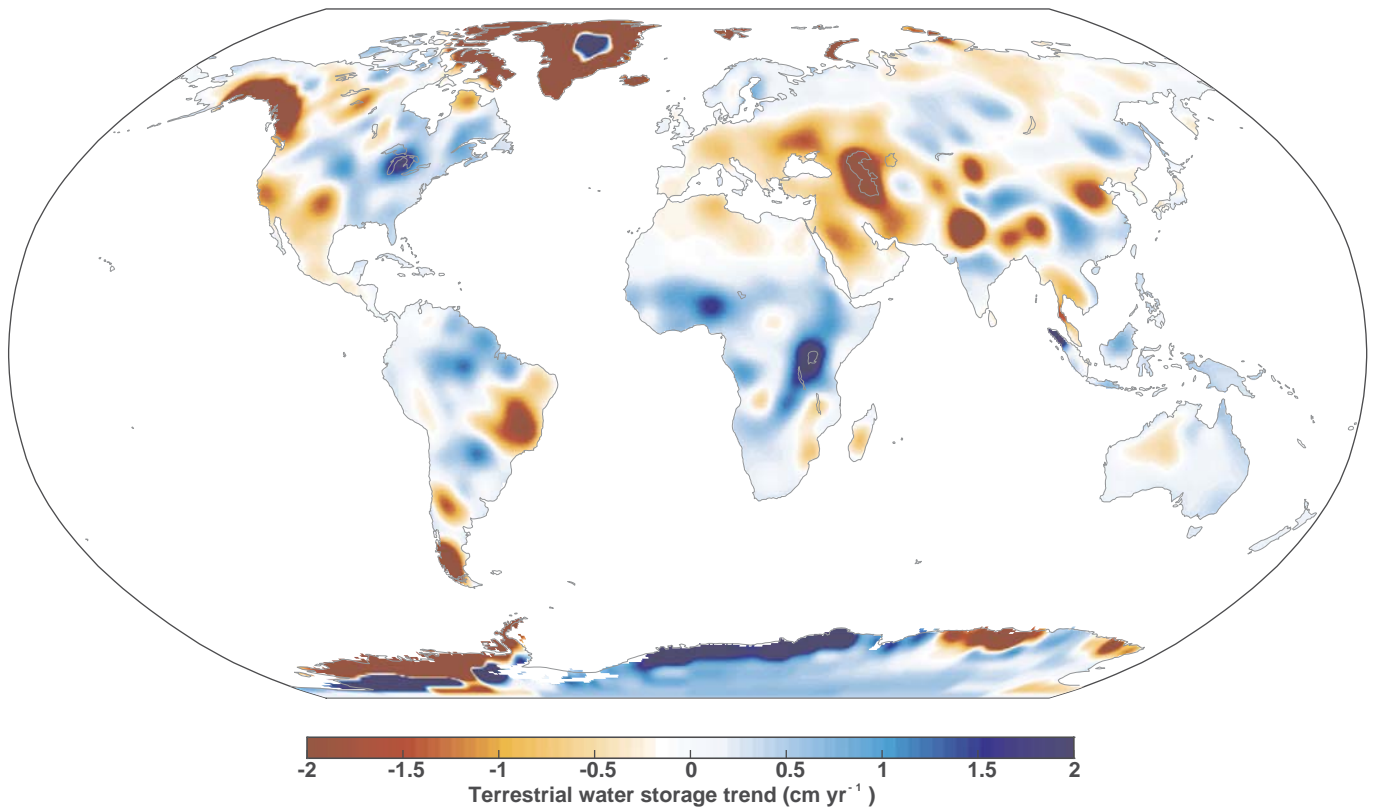
1021 The well locations mapped in Figure 1 are available from the U.S. Geological Survey's
1022 Groundwater Water Climate Response Network website and the International Groundwater
1023 Resources Assessment Centre's website. The JPL RL06_v02 gridded monthly mass data used
1024 Figures 2, 4, and 5 is identical to that which is available from the NASA/JPL GRACE Tellus
1025 website⁶¹. Figure 3 was created from a citation report that searched for journal articles containing
1026 "GRACE satellite" or "Gravity Recovery and Climate Experiment" and one of nine hydrology-
1027 related terms, performed on the Web of Science website. Figure 6 was constructed from images
1028 available on the University of Nebraska-Lincoln National Drought Mitigation Center's NASA
1029 GRACE website^{10,174}.

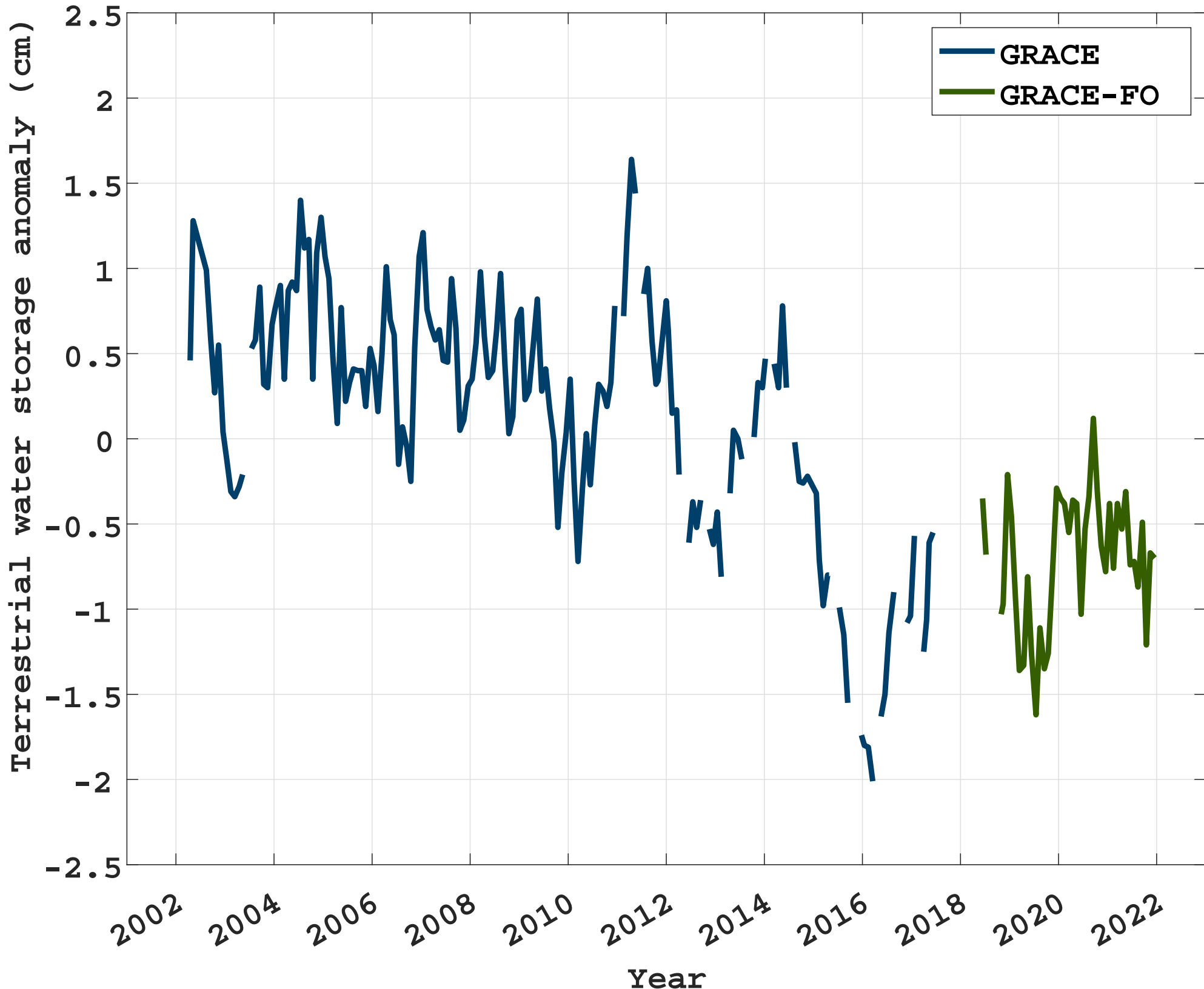
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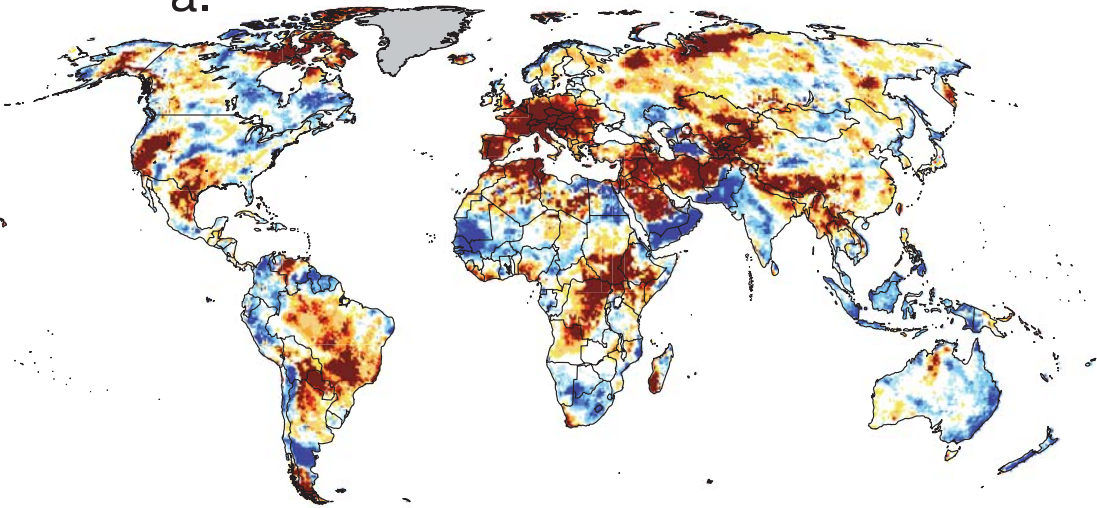




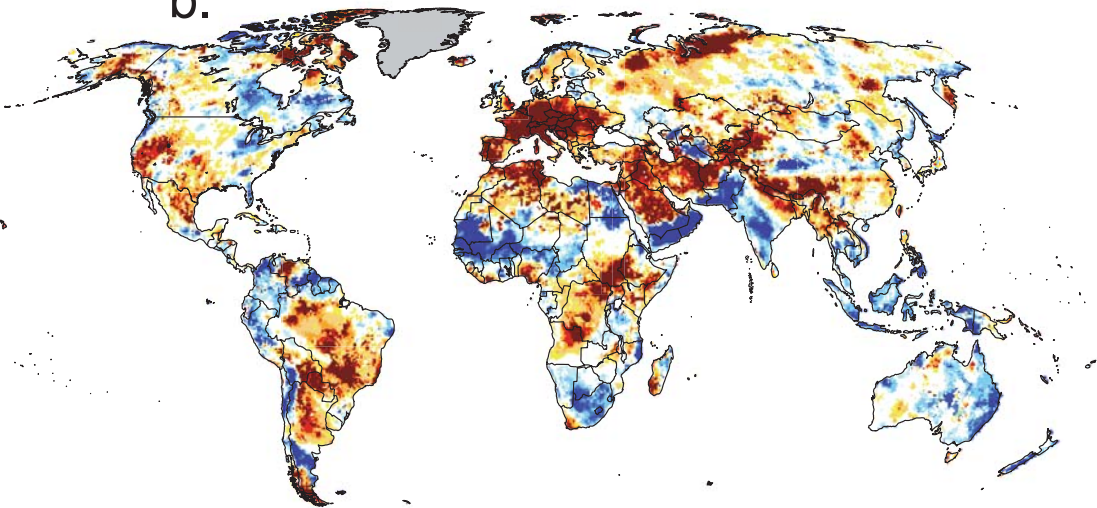




a.



b.



c.

