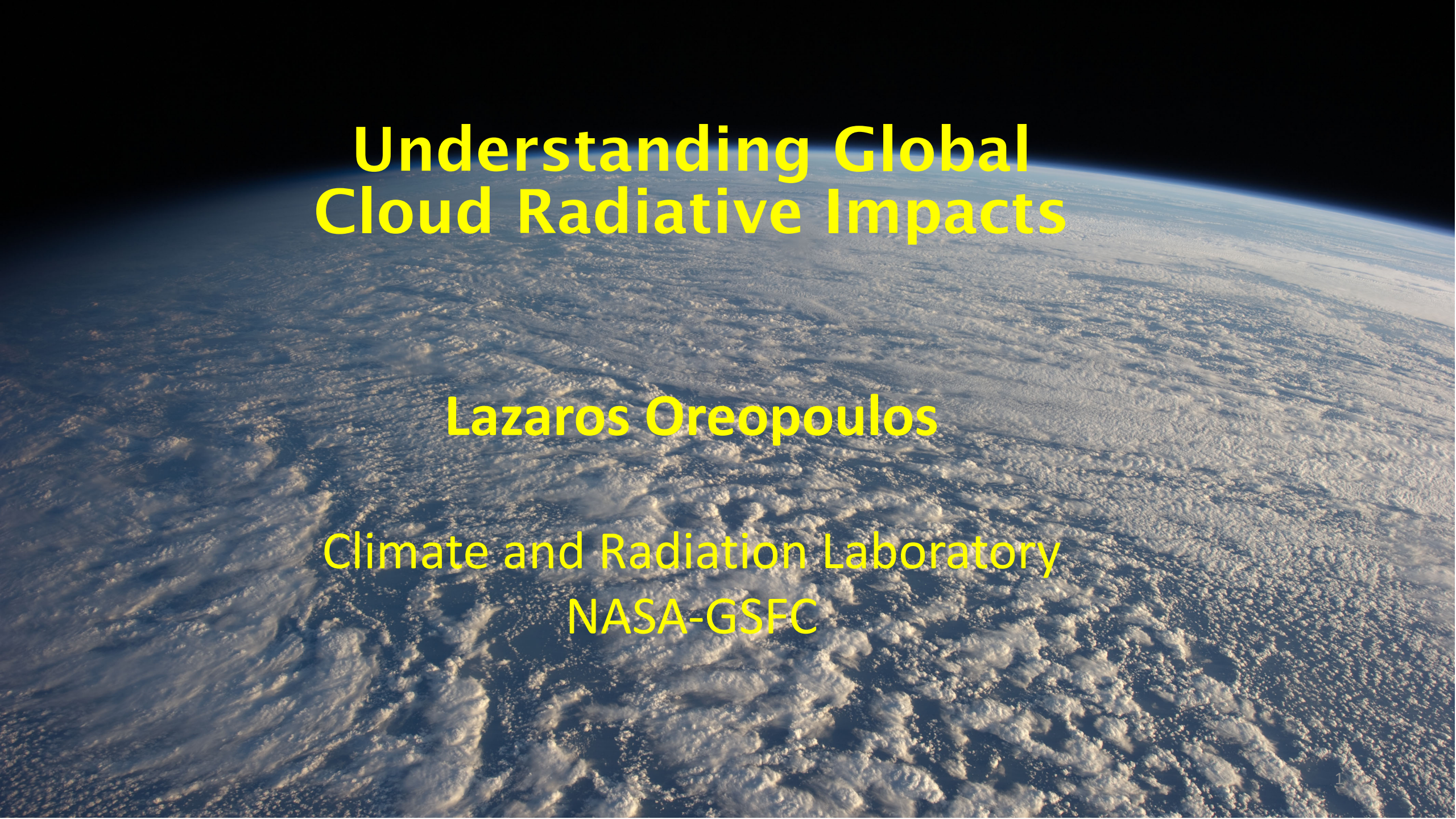


A satellite image of Earth from space, showing a vast expanse of white and grey cloud cover over a dark blue ocean. The horizon of the Earth is visible in the upper right, with a thin blue line of the atmosphere. The text is overlaid on the left side of the image.

Understanding Global Cloud Radiative Impacts

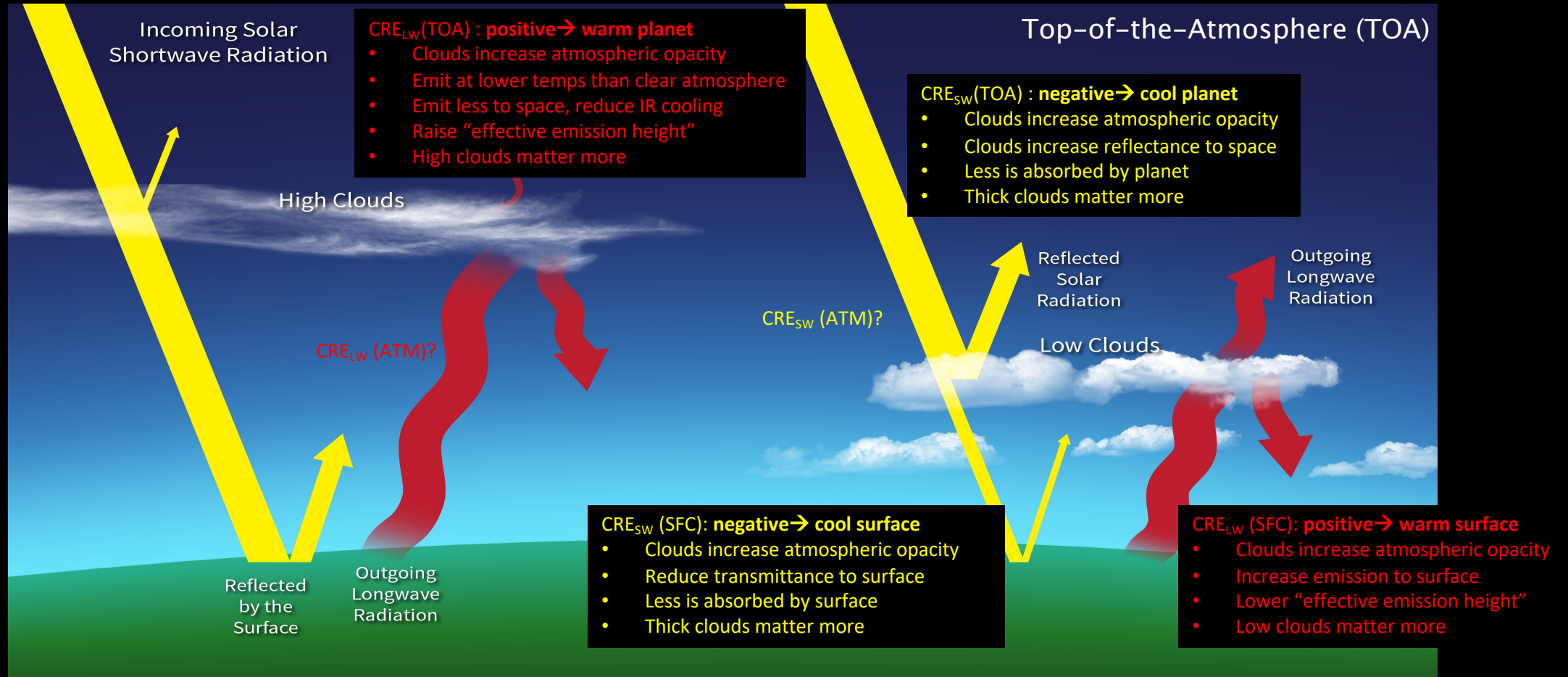
Lazaros Oreopoulos

Climate and Radiation Laboratory
NASA-GSFC

How clouds affect the radiation budget: Fundamentals

Solar (SW): Scatter (reflect up) and absorb (remaining is transmitted)

Thermal IR (LW): Absorb and emit (up and down) thermal (scattering is small)



A standard measure of cloud impact:

Cloud Radiative Forcing Effect (CRE) = (All-Sky Net Flux) - (Clear-Sky Net Flux)
Net Flux = Down Flux - Up Flux

$$\begin{aligned} \text{CRE}_{\text{TOT}} &= \text{CRE}_{\text{SW}} + \text{CRE}_{\text{LW}} \\ \text{CRE}(\text{TOA}) &= \text{CRE}(\text{SFC}) + \text{CRE}(\text{ATM}) \end{aligned}$$

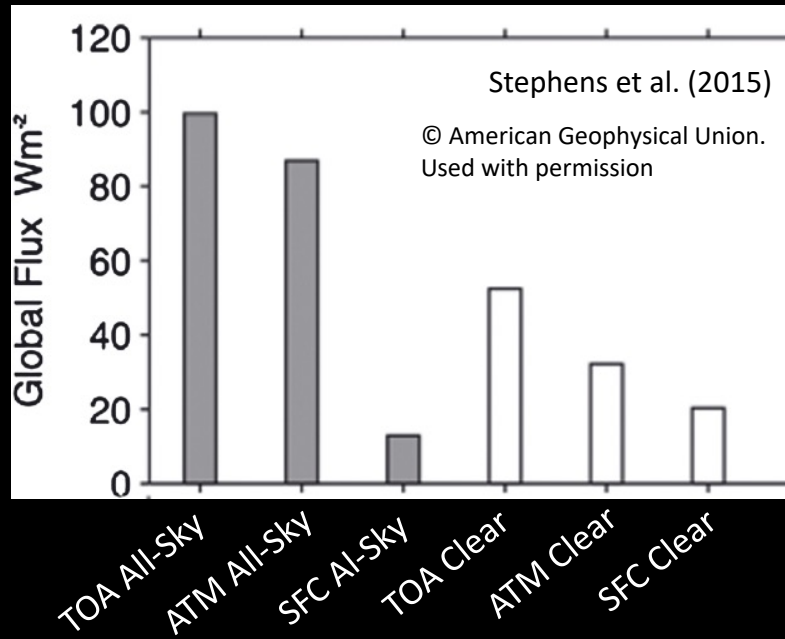
Q1: How much brighter is the Earth because of clouds? (space/TOA perspective)

Increase in albedo $\sim 81\%$: 0.29 vs 0.16

- Even for clear skies atmosphere contributes more than surface (see below).

In energy units:

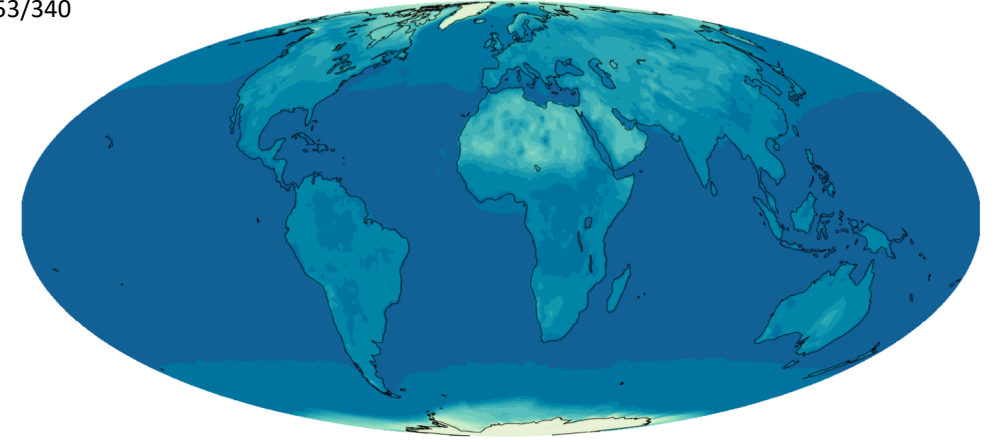
- $\text{CRE}_{\text{SW}}(\text{TOA}) = (\text{All-Sky Flux}) - (\text{Clear-sky Flux}) \approx -46 \text{ Wm}^{-2}$
more reflected to space (less absorbed by the planet)



Flux=Down-Up
 $46/340 \approx 0.13$; $53/340 \approx 0.16$

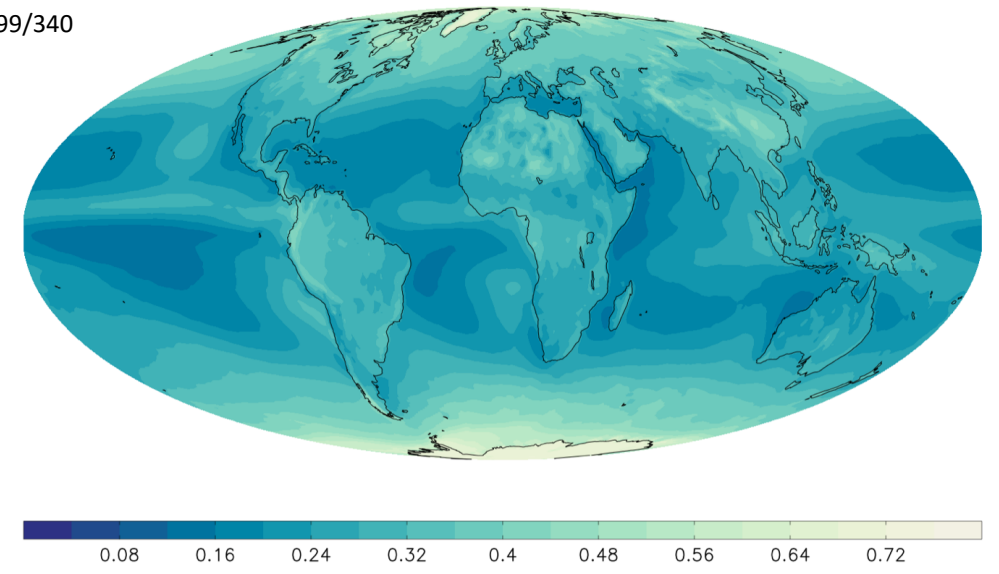
Clear-sky albedo (0.16)

53/340



All-sky albedo (0.29)

99/340



Q2: How much dimmer is the Earth's surface because of clouds? (SFC perspective)

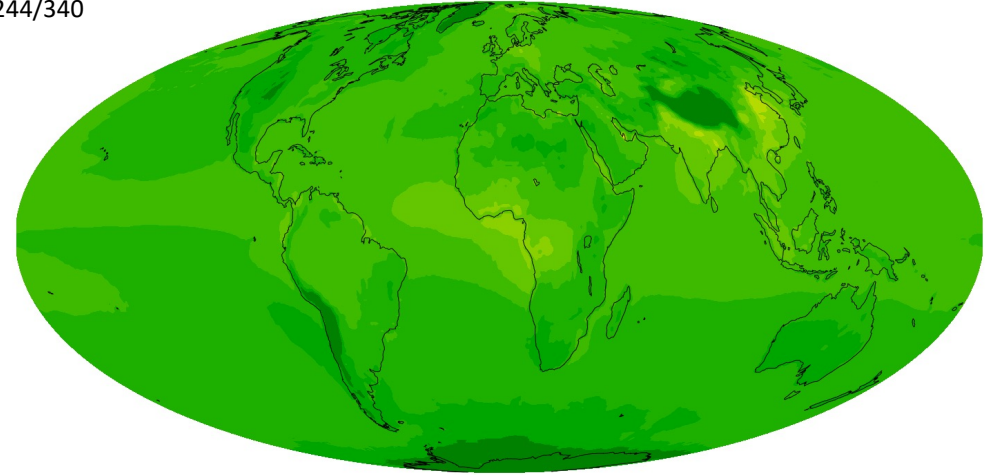
Clouds decrease solar reaching (and absorbed by) the surface by ~23%: 0.55 vs 0.72 (0.48 vs 0.63)

In energy units: $CRE_{SW}(SFC) =$
(All-Sky Flux absorbed@SFC) - (Clear-sky Flux absorbed@SFC)
 $\approx 164 - 214 = -50 \text{ Wm}^{-2}$ less absorbed by the surface

$CRE_{SW}(ATM) = -46 - (-50) \approx 4 \text{ Wm}^{-2}$, i.e., ~5% additional atmospheric absorption due to clouds (77 Wm^{-2} vs 73 Wm^{-2})

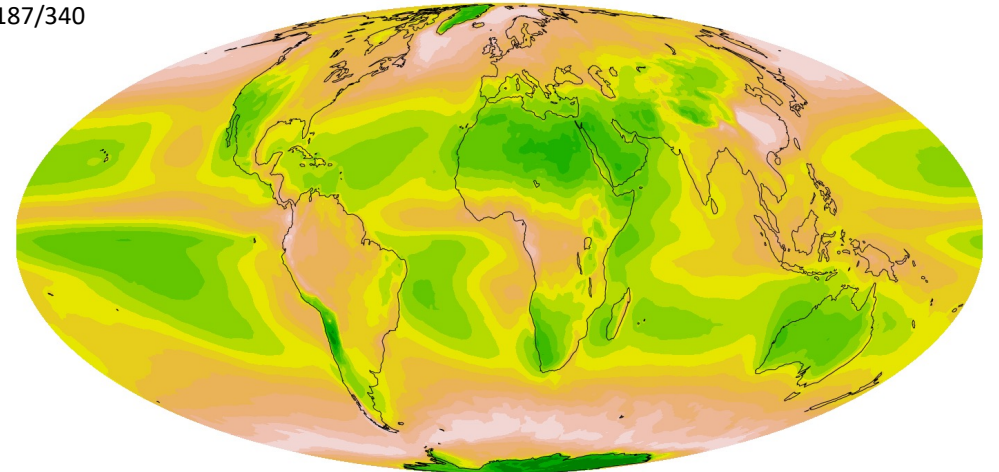
Clear-sky transmittance (0.72)

244/340



All-sky transmittance (0.55)

187/340



Q3: How much less thermal radiation escapes to space because of clouds? (TOA perspective)

Greenhouse factor: $\frac{F_{up}^{TOA}}{F_{up}^{SFC}}$; compare $\frac{F_{up}^{TOA}(all_sky)}{F_{up}^{SFC}(all_sky)}$ to $\frac{F_{up}^{TOA}(clear)}{F_{up}^{SFC}(clear)}$

$0.60/0.67 \approx 0.89$, i.e., ~11 % less escapes to space → ~11% stronger greenhouse with clouds

Much simpler: $\frac{F_{up}^{TOA}(all_sky)}{F_{up}^{TOA}(clear)} \approx \frac{240}{268} \approx 0.89$

also 11% reduction (surface upwelling is on average same under clear and cloudy conditions)

In energy units:

$CRE_{LW} = (All-Sky \text{ Flux}) - (Clear-sky \text{ Flux}) \approx 28 \text{ Wm}^{-2}$ less emitted to space

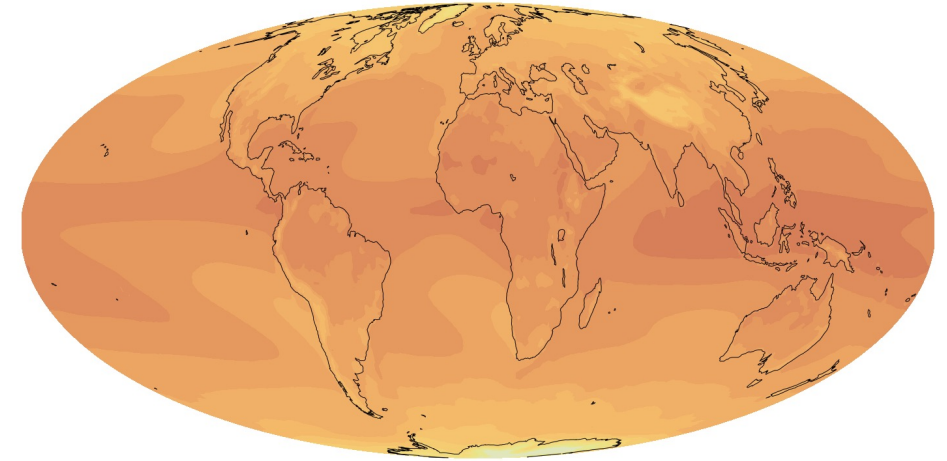
Flux=Down-Up

How much colder does the Earth appear from space because of clouds? (TOA perspective)

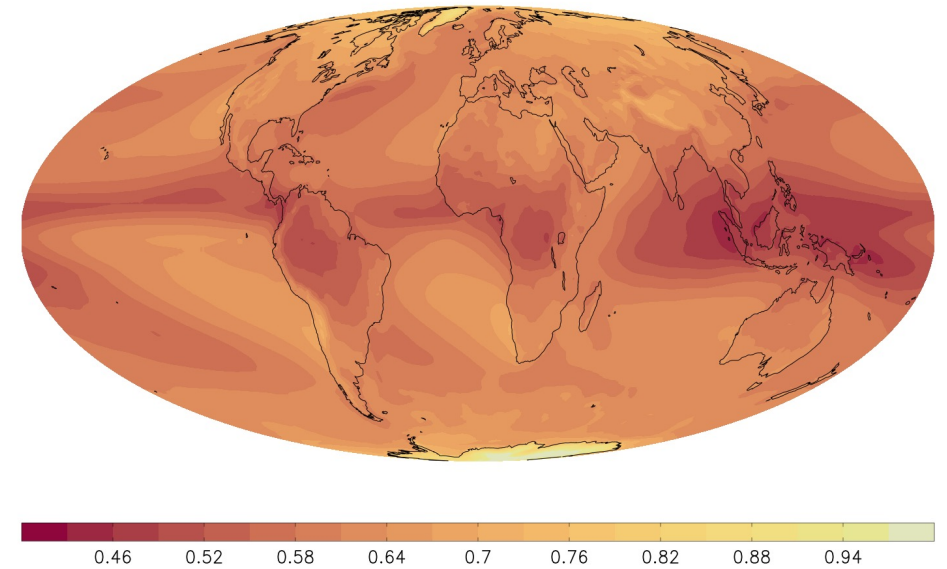
$$T_{eff}^{all_sky, TOA} = \left(\frac{F_{up}^{TOA}(all_sky)}{\sigma} \right)^{1/4} \approx 255 \text{ K} \quad T_{eff}^{clear, TOA} = \left(\frac{F_{up}^{TOA}(clear)}{\sigma} \right)^{1/4} \approx 262 \text{ K}$$

(if Earth is assumed a blackbody)

Clear-sky Greenhouse Effect Space Perspective (0.67)



All-sky Greenhouse Effect Space Perspective (0.6)



Q4: How much more downwelling does the Earth's Surface receive because of clouds? (SFC perspective)

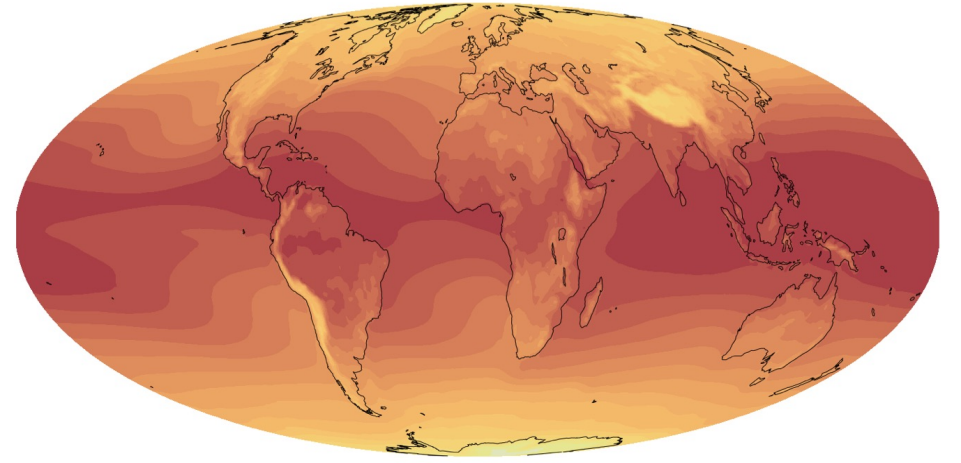
- 345/315, ~10% more to surface, due to clouds
- $CRE_{LW}(SFC) = (All-Sky\ Flux) - (Clear-sky\ Flux) \approx 30\ Wm^{-2}$ more to surface
- $CRE_{LW}(ATM) = 28 - 30 \approx -2\ Wm^{-2}$, very small global decrease in atmospheric LW absorption due to clouds, but with substantial uncertainty

How much warmer does the atmosphere appear from the surface because of clouds? (SFC perspective)

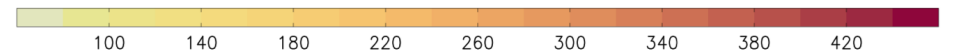
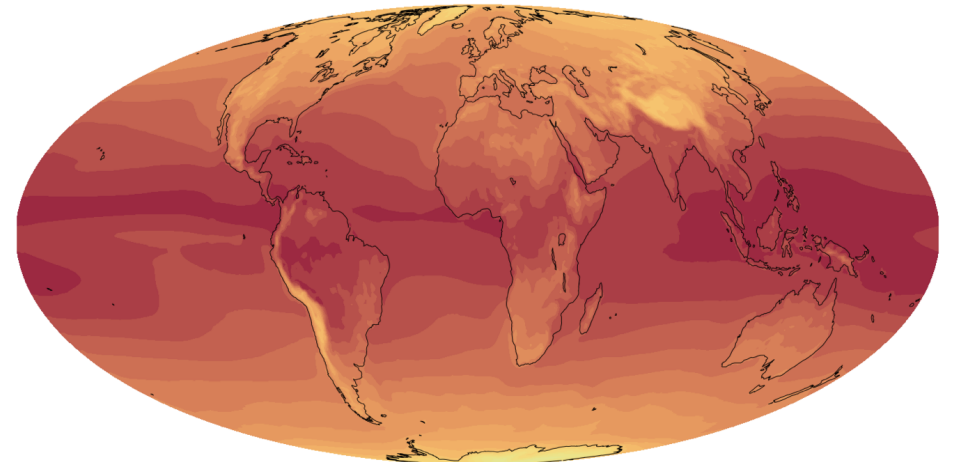
$$T_{eff}^{all_sky,SFC} = \left(\frac{F_{down}^{SFC}(all_sky)}{\sigma} \right)^{1/4} \approx 279\ K \quad T_{eff}^{clear,SFC} = \left(\frac{F_{down}^{SFC}(clear)}{\sigma} \right)^{1/4} \approx 273\ K$$

(if atmosphere is assumed a blackbody)

Clear-sky Greenhouse Effect Surface Perspective (315)



All-sky Greenhouse Effect Surface Perspective (345)



Summary of global annual Cloud Radiative Effects

(annual, according to CERES EBAF Ed4.0)

Can be positive in the winter hemisphere

“Isolated” low clouds
(L, liquid, CR7-CR11)
biggest contributor

Very few clouds have warming
effect on annual basis

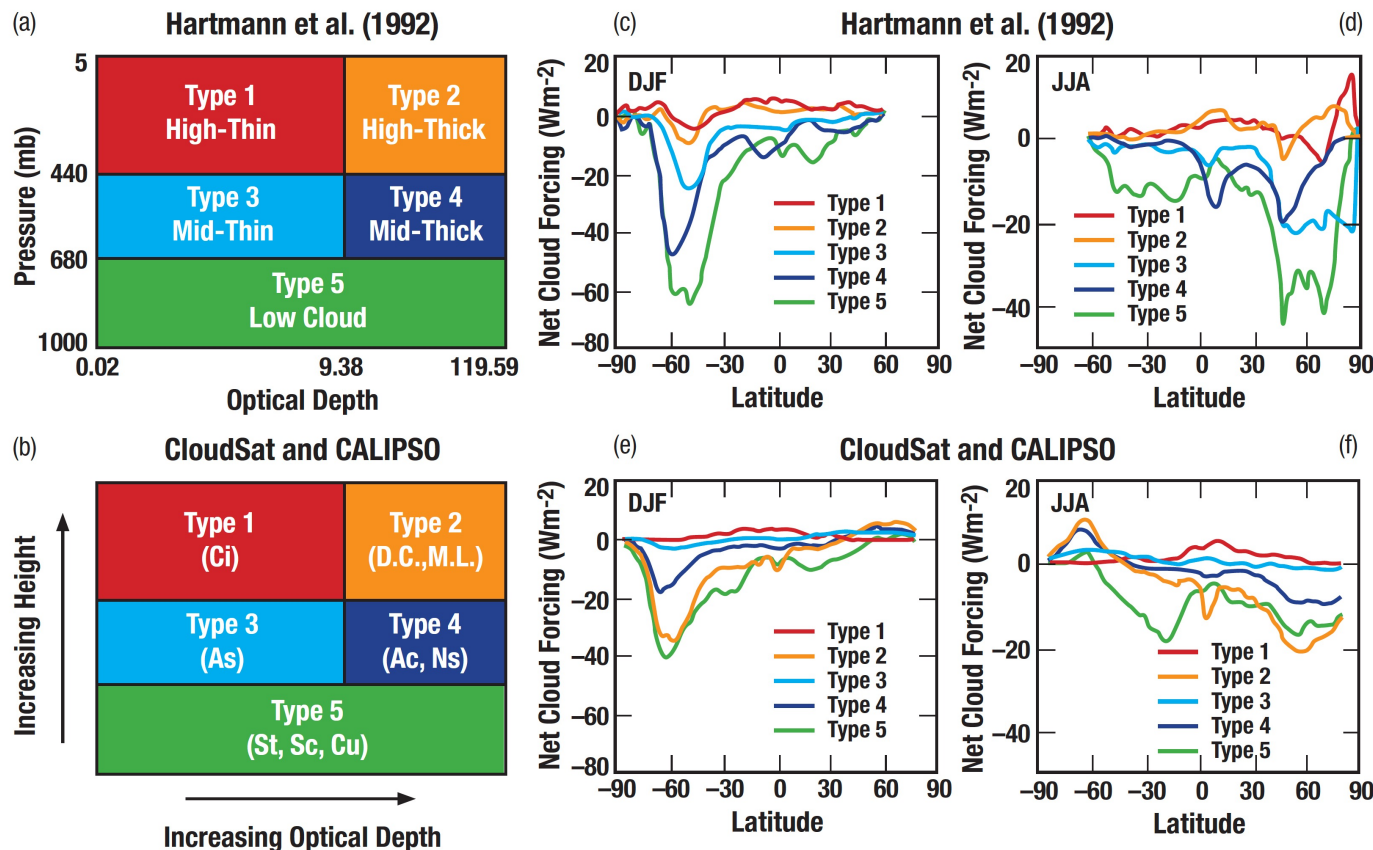
Perspective	LW	SW	Total	Impact
TOA (“Planet”, “Climate”)	+28	-46	-18	Cooling
Surface (SFC, BOA)	+30	-50	-20	Cooling
Atmosphere (ATM)	-2	+4	+2	Slight Warming

substantial uncertainty

Outcome of many cancellations, and
large uncertainty: high clouds warm,
low clouds cool

Virtually no clouds have
warming effect on annual basis

CRE by cloud type



ERBE-ISCCP

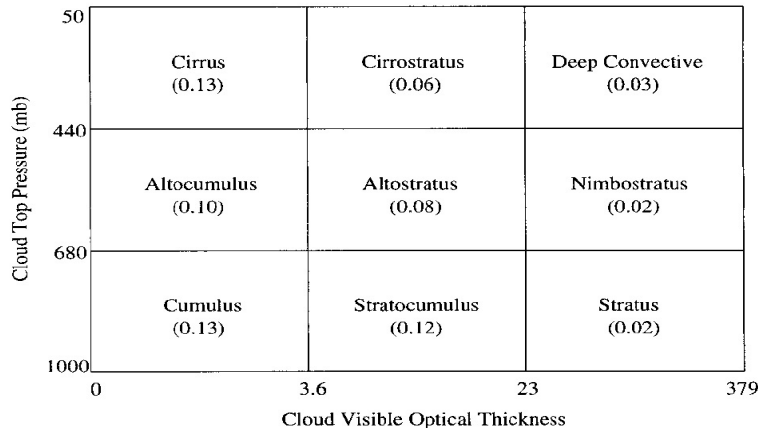
CloudSat-CALIPSO

Stephens et al. (2018)

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Current climate: $CRE_{TOT} < 0$
Relevant for future: CRE anomaly

ISCCP

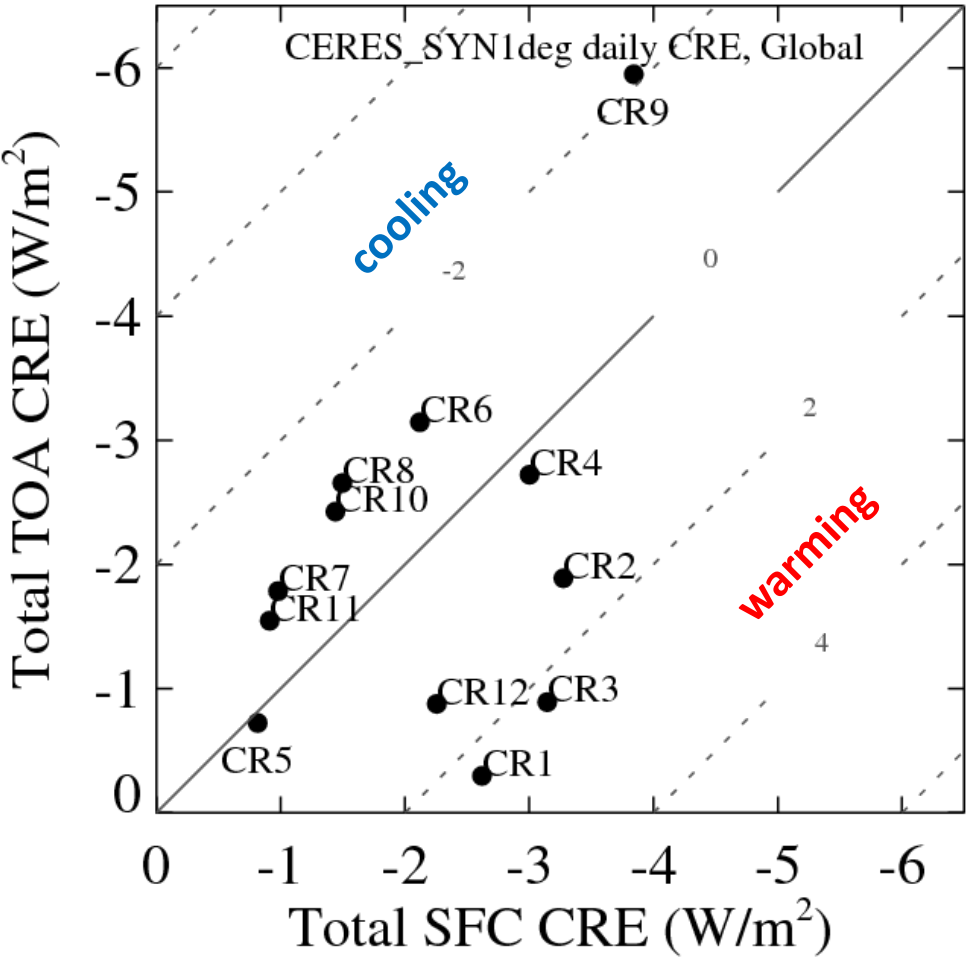


Cloud type	TOA		
	SW	LW	TL
Cirrus	-4.2	5.5	1.3
Cirrostratus	-7.9	5.5	-2.4
Deep convective	-6.2	2.9	-3.3
Altostratus	-3.2	1.5	-1.7
Altostratus	-8.3	2.0	-6.3
Nimbostratus	-3.4	0.7	-2.7
Cumulus	-5.2	0.6	-4.6
Stratocumulus	-12.7	1.2	-11.5
Stratus	-2.4	0.2	-2.2
Sum (true)	-53.5	20.1	-33.4

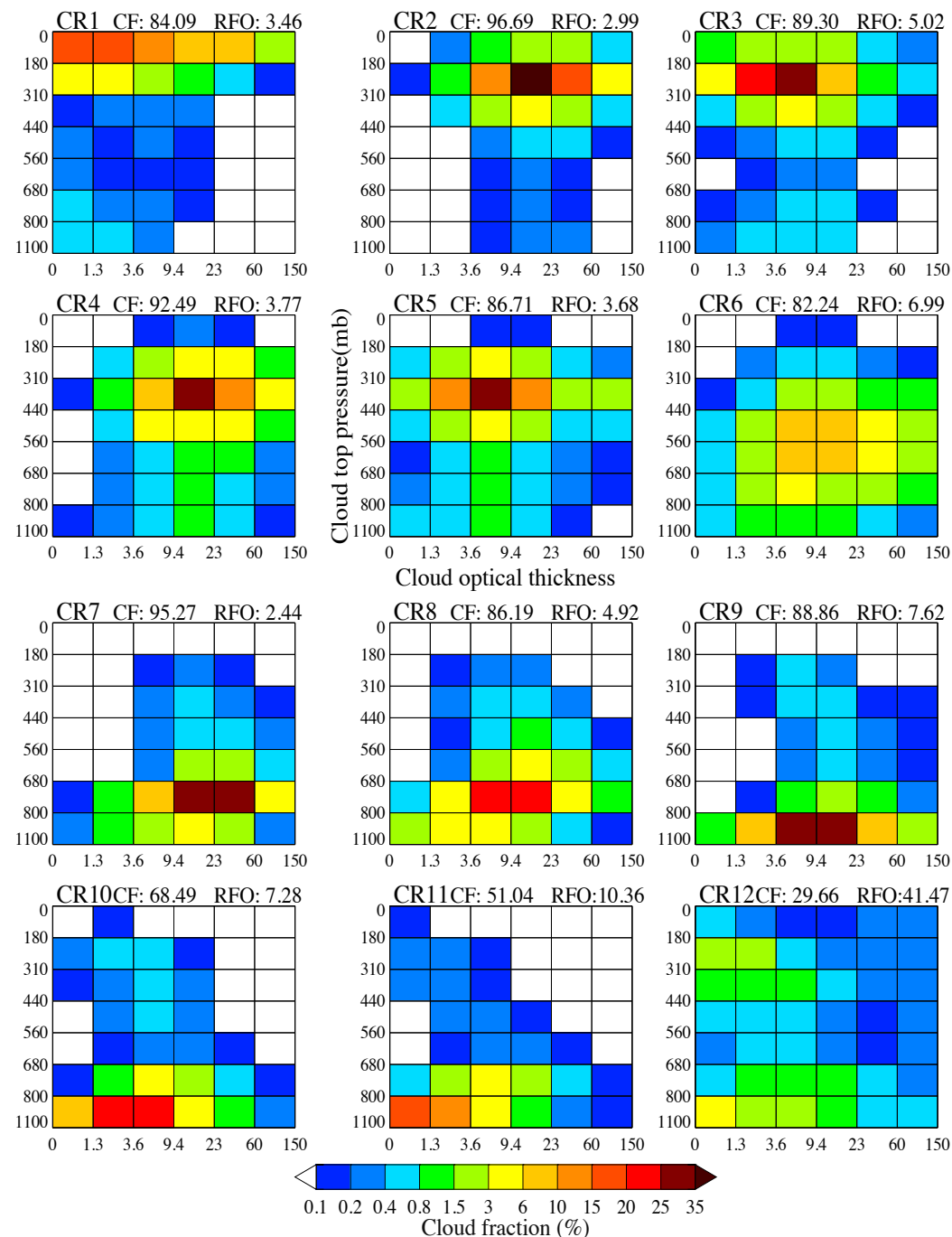
Chen et al. 2000

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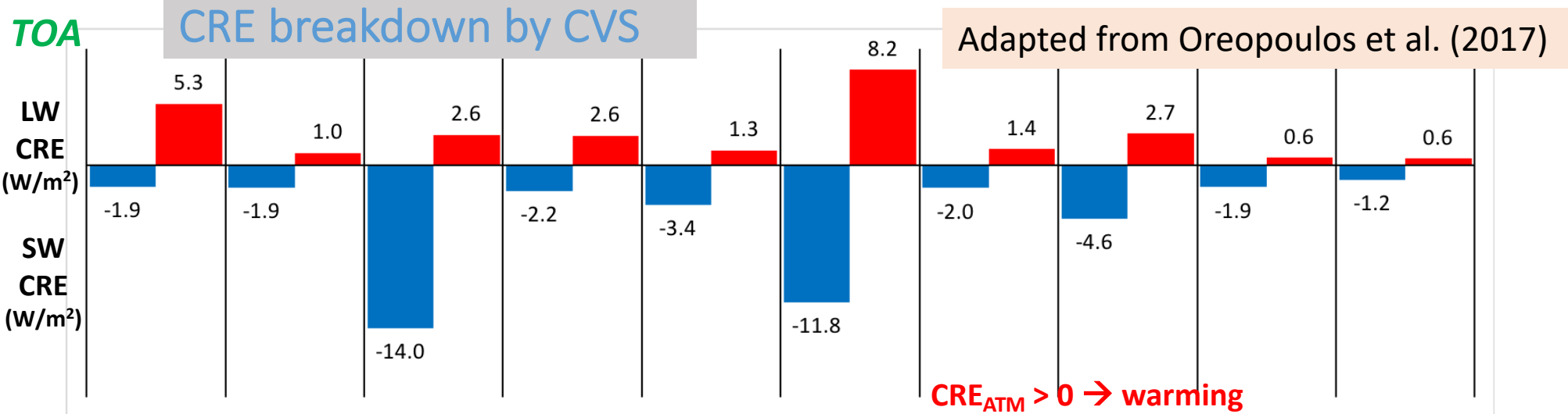
CRE breakdown by MODIS Cloud Regimes



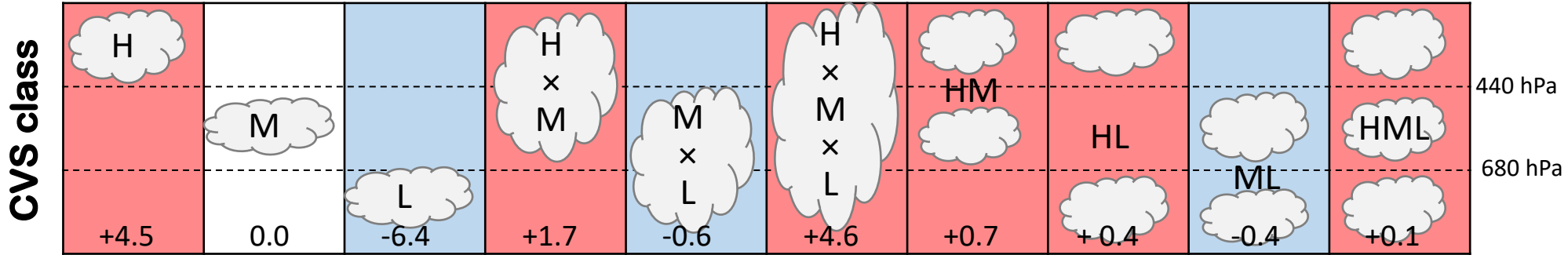
Adapted from Oreopoulos et al. (2016)



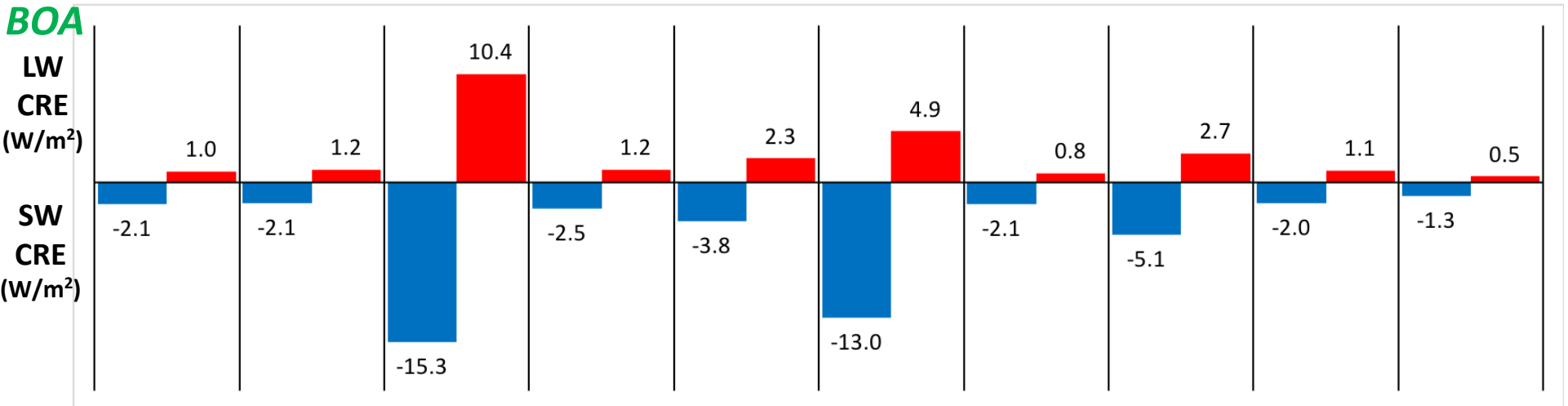
Only “H” and “HxM”
warm the planet



Total CRE (W/m²) in the atmosphere

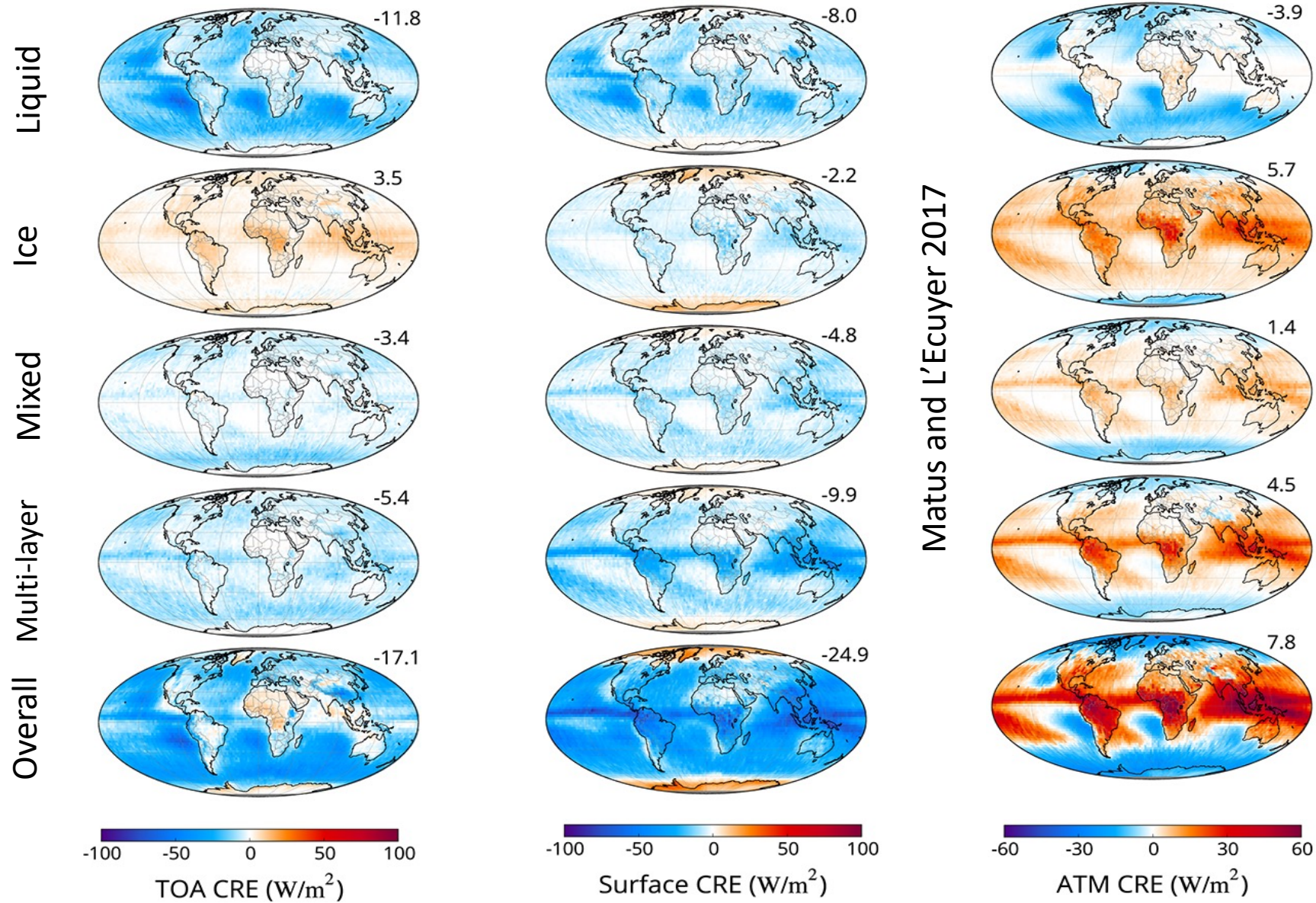


All CVS cool surface



CloudSat-CALIPSO
2B-CLDCLASS-LIDAR
2B-FLHXR-LIDAR

Breakdown of total planetary cloud radiative effect by phase

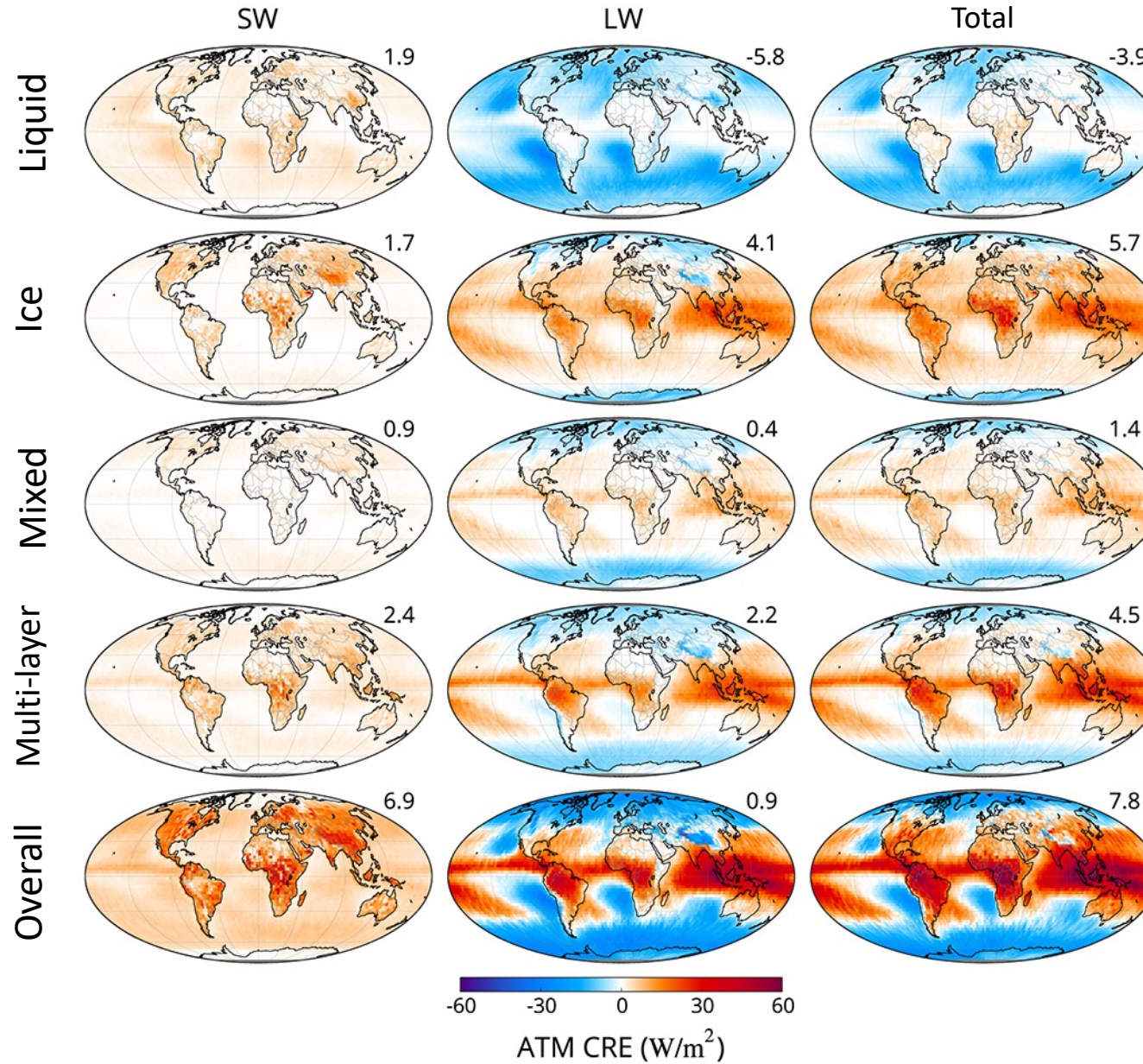


Matus and L'Ecuyer 2017

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"Planet"

Breakdown of atmospheric cloud radiative effect by phase



Matus and L'Ecuyer 2017

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Summary of Global Cloud Radiative Impacts

- Overall, current clouds radiatively cool the planet and (a bit more) the surface; radiative impact on atmosphere is much smaller (slight warming, from SW, within uncertainties)
 - Atmospheric effect is result of lots of cooling/warming cancellation in LW
- Clouds can have extensive warming TOA/SFC effect in the winter hemisphere
- Thanks to modern satellite observations and radiative transfer modeling, the radiative impact of clouds can be decomposed by cloud class:
 - Cloud classes with low/liquid clouds cool the surface and atmosphere
 - Cloud classes with high clouds still cool the surface, but warm the atmosphere
- Of climate change relevance (cloud feedback):
 - Positive feedback: Fewer clouds that cool and/or more that warm planet
 - Negative feedback: More clouds that cool and/or fewer that warm planet