

1 **Effect of Baseline Period on Quantification of Climate Extremes over the United States**

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8 **Key Points:**

- 9 ● Updating the baseline period from 1981-2010 to 1991-2020 leads to significant changes
10 in percentile-based extreme climate indices in the US
- 11 ● Temperature indices show generally increased cold extremes and decreased warm
12 extremes across the US when the baseline period is updated
- 13 ● For precipitation indices, the later baseline period indicates fewer but more intense
14 extreme events in the south and central US

15 **Abstract**

16 Extreme climate events are societally harmful and have increased in frequency and intensity in
17 recent decades. Indices based on temperature and precipitation are a valuable way to quantify
18 climate extremes. Certain indices are defined relative to percentiles, which are dependent on a
19 climatological baseline period. In this study, indices computed using temperature and
20 precipitation from the Modern Era Retrospective Analysis for Research and Applications,
21 Version 2 (MERRA-2) are calculated using percentiles from three baseline periods: 1981-2010,
22 1991-2020 and 1981-2020. Updating the baseline period from 1981-2010 to 1991-2020 leads to
23 significant changes in the quantification of temperature and precipitation extremes over the
24 United States over 1980-2021. Using the later baseline period indicates more cold extremes,
25 fewer warm extremes, and fewer but more intense precipitation extremes throughout the US,
26 with regional variation. Changing the baseline period can mislead the public and decision
27 makers, potentially undermining the appropriate response to climate-related health risks.

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29 **Plain Language Summary**

30 Indices computed using 2-meter air temperature and precipitation are used to represent extreme
31 climate events such as heat waves, cold waves, heavy precipitation, and drought. Some indices
32 are defined relative to percentile-based thresholds, which are computed using a baseline
33 climatology period. The baseline climatology is typically a thirty-year period and is updated
34 every ten years. This study examines how updating the baseline climatology period from 1981-
35 2010 to 1991-2020 affects the quantification of climate extremes in the United States over 1980-
36 2021. In general, since the 1991-2020 period is warmer than 1981-2010 throughout the United
37 States, there are fewer warm extremes detected and more cold extremes detected when it is used
38 as the baseline. The differences are most notable in the southwest and northeast United States.
39 The changes in the precipitation indices vary throughout the country, but in certain parts of the
40 southern and central United States, updating the baseline period leads to the detection of fewer
41 but heavier extreme precipitation events. It is important to communicate the choice of baseline
42 climatology period to prevent misinterpretation of the extreme climate indices and the
43 comparison of different studies.

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51 **1 Introduction**

52 Extreme climate events, including heatwaves, heavy precipitation, and drought, have a
53 large impact on society through human health, destruction of infrastructure, ecological change,
54 and economic losses. Indices where daily temperature or precipitation is compared to a threshold
55 are a valuable tool for the monitoring and quantification of extremes across different regions
56 (Zhang et al., 2011; Alexander et al., 2019; Dunn et al., 2022). Some indices use a percentile-
57 based threshold, and thus are dependent on the choice of baseline period used to define the
58 percentiles (Zhang et al., 2005; Dunn & Morice, 2022). As global and regional climate continues
59 to change, the interpretation of extreme events is increasingly reliant on this baseline period, and
60 this can be a source of confusion and ambiguity for the policy making community.

61 To have the best representation of the current climate, operational centers typically use a
62 30-year climate baseline period that shifts in time every ten years (Arguez et al., 2012), also known
63 as a normal. However, due to the non-stationarity of climate, alternatives to the 30-year climate
64 normal have been suggested (Livezey et al., 2007; Wilks, 2013; Wilks & Livezey, 2013). The
65 World Meteorological Organization (WMO) suggests that the maximum amount of data should be
66 included for the detection of extreme events due to their rare occurrence (Trewin, 2007). The
67 appropriate baseline period may differ based on the application (i.e., Schreck et al., 2021).

68 The climate changes between long-term means (e.g., Kendon et al., 2020), so the shifting
69 of the baseline period can affect the magnitude and interpretation of climate anomalies (Scherrer
70 et al., 2006; Arguez & Vose, 2011). This issue has the potential to be exacerbated in the situation
71 of climate extremes. Previous studies have shown linear trends in percentile-based extreme
72 temperature indices to vary significantly with different baseline periods (Yosef et al., 2020; Dunn
73 & Morice, 2022). Conversely, the transition to a new baseline was found to affect a drought index
74 only marginally (Cammalleri et al., 2021).

75 In this study, we examine how updating the baseline period from 1981-2010 to either 1981-
76 2020 or 1991-2020 affects the quantification and classification of climate extremes across the
77 continental United States. We employ indices defined using 2-m temperature and precipitation
78 from NASA's Modern Era Retrospective Analysis for Research and Applications, version 2
79 (MERRA-2; Gelaro et al., 2017). Changing the baseline period can affect the perception of the
80 public and decision makers, so it is crucial to understand and communicate how to interpret this
81 change. This evaluation expands on the above-mentioned studies by focusing on distinct regions
82 within the United States, including heatwave and precipitation indices, and examining seasonal
83 variability of changes in the indices. This manuscript also serves to document differences between
84 Version 1 (GMAO, 2020) and Version 2 (GMAO, 2022) of the MERRA-2 Monthly Extremes
85 Detection Indices dataset. Data and methods are described in section 2, while section 3 outlines
86 the changes in temperature and precipitation extreme climate indices with the updated baseline.
87 Conclusions follow in section 4.

88

89 **2 Data and Methods**

90 **2.1 MERRA-2**

91 Data used in this study is from the MERRA-2 reanalysis (Gelaro et al., 2017) and is akin to the
92 extremes detection indices file collection (Collow et al., 2022; GMAO 2020; GMAO 2022). Daily
93 2-m temperature and precipitation data from MERRA-2 are available at a spatial resolution of
94 0.625° longitude by 0.5° latitude from January 1980 to present (GMAO 2015a, b), though the

95 current analysis is for 1980-2021. Precipitation used to generate the climate statistics is the model
96 generated output, and not the observation corrected land-forcing precipitation (Reichle et al.,
97 2017). An evaluation of the climate of MERRA-2 can be found in Bosilovich et al. (2015).

98 2.2 Percentile Calculation

99 Percentiles and extreme indices were derived using daily mean fields of precipitation and
100 2-m temperature, as well as daily minimum and maximum 2-m temperature from MERRA-2.
101 Percentiles for each calendar day of the year were computed with the multi-year daily running
102 percentile values (*ydrunpctl*) function from the Climate Data Operators toolbox (Schulzweida,
103 2022) with a 15-day running window. This differs from the 5-day window recommended by the
104 Expert Team on Climate Change Detection and Indices (ETCCDI), because this shorter window
105 resulted in too much day-to-day variability in the percentiles across the United States. The 15-day
106 window has been utilized in the past by Collow et al. (2016) and Thomas et al. (2020). Depending
107 on the location, this may result in additional exceedances of warm extreme thresholds during the
108 summer and fewer exceedances during the winter. There is minimal influence during the shoulder
109 seasons. Zhang et al. (2005) evaluated differences between a 5-day and 25-day window and
110 demonstrated that the 25-day window results in a smaller bias within the baseline period but could
111 complicate the interpretation of more intense extreme events. Only days with at least 1 mm of
112 precipitation were included in the percentile calculation for precipitation. Three baseline periods,
113 1981-2010, 1981-2020, and 1991-2020, were used for the percentile calculations to determine the
114 dependency on the climatology period used.

115 2.3 Indices Calculation

116 Daily exceedances of the percentiles were detected in the MERRA-2 dataset for the years
117 of 1980 through 2021 using the three sets of percentiles and aggregated into monthly indices
118 representing extreme temperature and precipitation events, as well as heatwaves. These indices are
119 analogous to those included in the MERRA-2 monthly extremes detection indices file collection
120 (GMAO 2020; GMAO 2022), and most have been recommended for use by the ETCCDI
121 (Alexander et al., 2006). More specific details pertaining to the extreme indices are given in Table
122 1. The selected indices are included in the MERRA-2 extremes detection indices data product and
123 are also available for visualization on the Global Modeling and Assimilation Office's Framework
124 for Live User-Invoked Data (FLUID) webpage,
125 https://fluid.nccs.nasa.gov/reanalysis/extreme_merra2/. The heatwave related indices (HWD,
126 HWF, and HWM) are based on Perkins and Alexander (2013) in which a heatwave occurs if the
127 mean 2-meter temperature exceeds the calendar day 90th percentile for at least three consecutive
128 days. The frequency of extreme precipitation events, R90d, R95d, and R99d, are given as a count
129 of the number of events as opposed to the percentage of the total precipitation that is considered
130 extreme as included in the Climpact list of indices (<https://climpact-sci.org/indices/>). The
131 frequency of 99th percentile precipitation events was previously used to evaluate the underlying
132 general circulation in MERRA-2 with respect to teleconnection patterns (Collow et al., 2017). The
133 dependence on baseline period is assessed using the difference between a given index computed
134 using two baseline periods over the entire MERRA-2 period (1980-2021). Significance of
135 differences is assessed using a two-sample student's t-test at the 90% confidence level.

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137 **Table 1.** Percentile-based indices included in this study. *Recommended by the Expert Team on
 138 Sector-Specific Climate Indices (ET-SCI; <https://climpact-sci.org/indices/>).

Index	Name	Calculation	Units
HWD*	Heat wave duration	Maximum length of consecutive days that satisfy heat wave criteria (daily mean 2 m temperature exceeds the 90 th percentile for at least three consecutive days)	days
HWF*	Heat wave frequency	Count of days that satisfy heat wave criteria (see HWD)	count
HWM*	Heat wave magnitude	Mean temperature anomaly on days that satisfy heat wave criteria (see HWD)	K
R90p	Wet day precipitation	Mean precipitation on days that exceed the 90 th percentile of precipitation	mm day ⁻¹
R90d	Wet days	Count of days that exceed the 90 th percentile of precipitation	count
R95p*	Very wet precipitation	Mean precipitation on days that exceed the 95 th percentile of precipitation	mm day ⁻¹
R95d	Very wet days	Count of days that exceed the 95 th percentile of precipitation	count
R99p*	Extremely wet precipitation	Mean precipitation on days that exceed the 99 th percentile of precipitation	mm day ⁻¹
R99d	Extremely wet days	Count of days that exceed the 99 th percentile of precipitation	count
TN10p*	Cold Nights	Percent of days with a minimum temperature below the 10 th percentile	%
TX10p*	Cold Days	Percent of days with a maximum temperature below the 10 th percentile	%
TN90p*	Warm Nights	Percent of days with a minimum temperature above the 90 th percentile	%
TX90p*	Warm Days	Percent of days with a maximum temperature above the 90 th percentile	%

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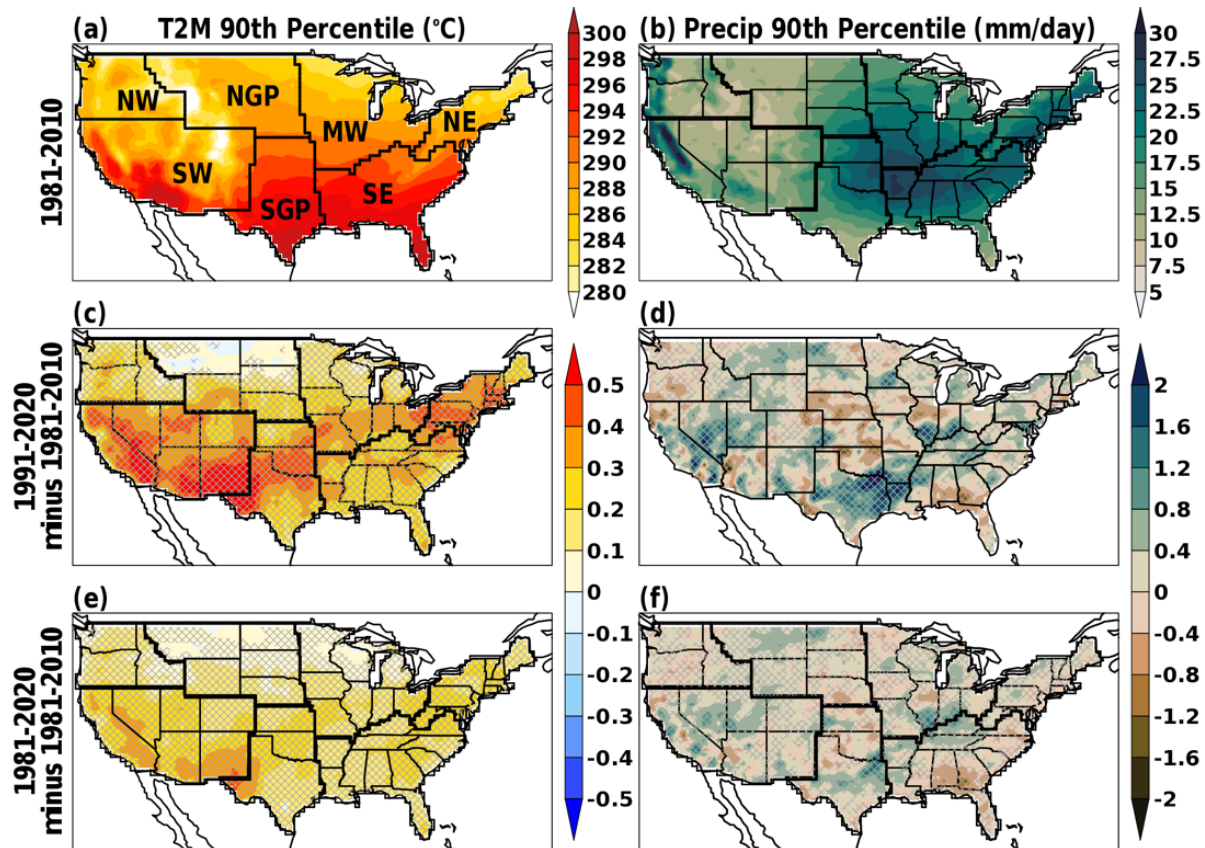
140 **3 Results**

141 3.1 Percentile changes with changing baseline period

142 The average over all calendar days of the 90th percentile of daily mean 2-meter temperature
 143 and precipitation is shown in Fig. 1. During the 1981-2010 period, the 90th percentile of 2-meter
 144 temperature is greatest in the Southern Great Plains and smallest in the high-elevation areas in the
 145 Rocky Mountain range (Fig. 1a). When the baseline period is updated to 1991-2020 (Fig. 1c), the
 146 90th percentile of 2-meter temperature increases throughout the US, with largest differences in the
 147 Southwest. Differences are significant everywhere except for a small region in the Northern Great
 148 Plains. The spatial pattern is similar to the differences in 30-year normals produced by the National
 149 Centers for Environmental Information (NCEI 2021). The differences when the baseline period is
 150 1981-2020 are smaller, but still statistically significant throughout the continental US (Fig. 1e).

151 For precipitation, the 90th percentile averaged over all calendar days over the 1981-2010
 152 period (Fig. 1b) shows higher values in the south-central US, eastern US, and Pacific Northwest,
 153 and lower values in the intermountain west. With the updated baseline period of 1991-2020,
 154 changes in the percentiles for precipitation are less spatially consistent than for temperature, but
 155 still significant in many regions. Parts of the Southern Great Plains through the Midwest and
 156 Southwest US show significantly larger precipitation percentiles with the updated climatology
 157 (Fig. 1d). This differs from the change in the NCEI precipitation normals, which shows a decrease

158 over the Southwest US with the later period (NCEI 2021). With the 1981-2020 baseline period,
 159 the changes are more muted, but still significantly positive in these regions (Fig. 1f).



160
 161 **Figure 1.** (a) Average of all calendar-day 90th percentiles of 2-m temperature computed using
 162 1981-2010, (c) average difference over all calendar days between percentiles computed using
 163 1991-2020 and 1981-2010, (e) average difference over all calendar days between percentiles
 164 computed using 1981-2020 and 1981-2010. (b,d,f) as in (a,c,e) but for the 90th percentile of
 165 precipitation. Grey hatching indicates where differences are significant at the 90% confidence
 166 level. Labels in (a) denote the regions used in Figure 2.

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168 3.2 Changes in extreme climate indices

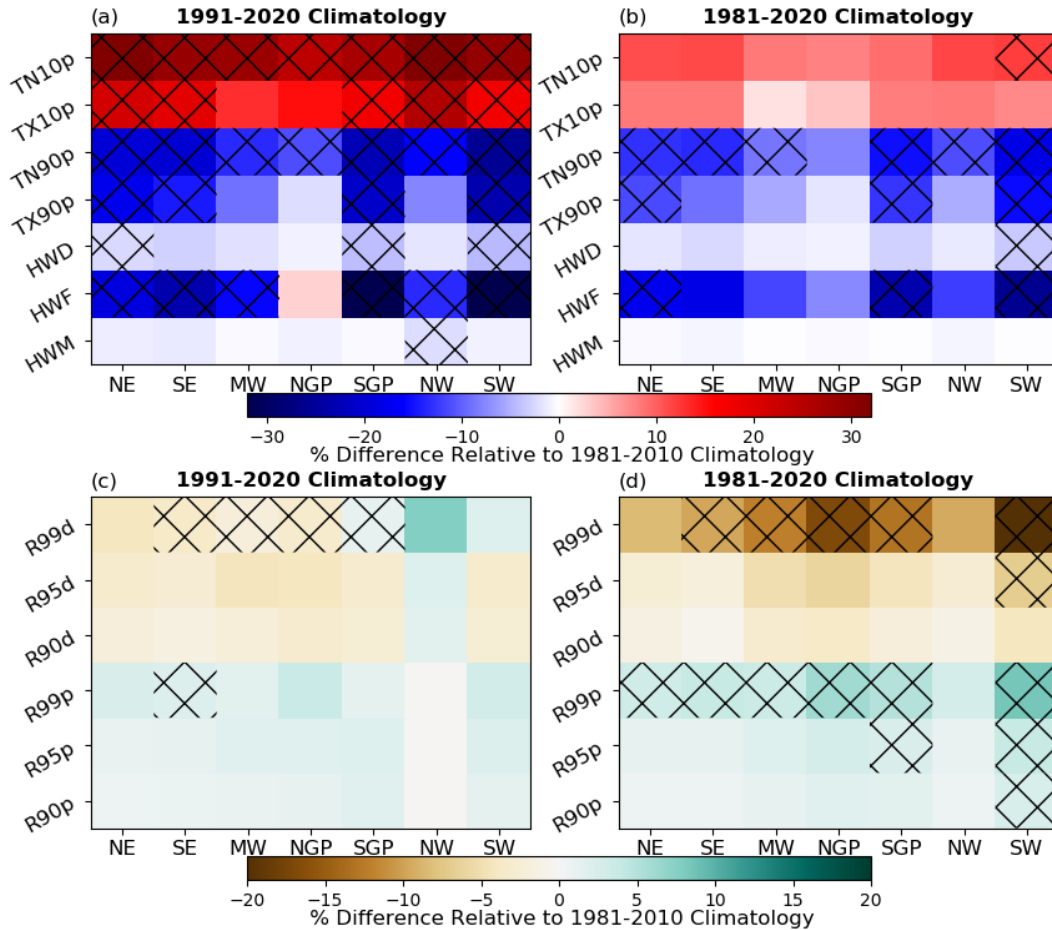
169 The identified changes in the temperature and precipitation percentiles will lead to changes
 170 in the extreme climate indices that are defined relative to them. Fig. 2 shows a summary of each
 171 of the percentile-dependent indices and how they change, on average, when the baseline period for
 172 percentiles is changed. Fig. 2 separates the percent differences relative to the 1981-2010
 173 climatology period for each index into the regions of the US used in the National Climate
 174 Assessment (NCA; Wuebbles et al., 2017) and denoted in Fig. 1. In general, changes with baseline
 175 period are largest for the temperature indices using the 1991-2020 climatology (Fig.
 176 2a). Temperature extremes defined using the 10th percentile (TN10p and TX10p) are more
 177 frequent with the updated climatology – with the later climatology period, there are more days and
 178 nights identified below the 10th percentile. The opposite is true for indices defined using the 90th

179 percentile (TN90p and TX90p) - the 1991-2020 climatology results in fewer identified extreme
180 warm days and nights. Changes in heat wave frequency (HWF) indicate that the 1991-2020
181 climatology leads to fewer heat wave days on average in most regions, except in the Northern
182 Great Plains, where changes in the 90th percentile of temperature were weak and insignificant when
183 updating the baseline period (Fig. 1c). In general, the changes when using the 1981-2020
184 climatology (Fig. 2b) result in the same sign, but weaker in magnitude and significance, as is
185 expected under a warming climate when moving to a longer reference period that includes more
186 recent years (i.e., Fig. 1c, e).

187 For the precipitation indices, in general the updated climatology periods result in fewer
188 days with extreme precipitation (R99d, R95d, R90d) but more precipitation on extreme days
189 (R99p, R95p, R90p), with differences most significant for the 99th percentile indices (Fig. 2c-
190 d). Unlike the temperature indices, the differences for R99d and R99p are larger when using the
191 1981-2020 climatology rather than the 1991-2020, likely due to the very rare nature of these
192 events.

193 While the focus here is on the annual changes, variations across the seasonal cycle are also
194 an important consideration and are included in the supplementary material (Figs. S1-S4). In
195 general, changes in the minimum temperature indices (TN10p and TN90p) are stronger and more
196 significant in summer (Fig. S3a) while the maximum temperature indices (TX10p and TX90p) are
197 stronger and more significant in winter (Fig. S1a). The increase in detected heat wave frequency
198 (HWF) in the Northern Great Plains is primarily a spring phenomenon (Fig. S2a). For precipitation
199 indices, there are no significant changes during the winter or spring seasons when updating to the
200 climatology period of 1991-2020 (Figs. S1c, S2c). Changes are most significant during summer,
201 when the Southeast and Southwest regions show a decrease in the frequency of extreme
202 precipitation events and increase in the amount of precipitation from an event (Figure S3c).

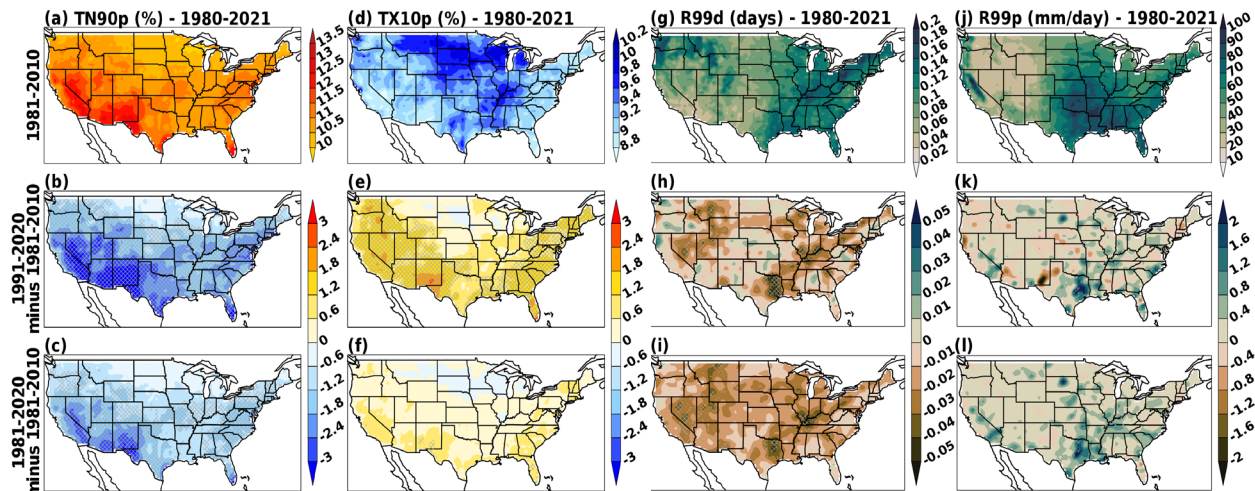
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204
 205 **Figure 2.** Average percent difference relative to the baseline climatology of 1981-2010 in area
 206 averaged over regions of the United States for (a) temperature indices using a baseline climatology
 207 of 1991-2020, (b) temperature indices using a baseline climatology of 1981-2020, (c) precipitation
 208 indices using a baseline climatology of 1991-2020, and (d) precipitation indices using a baseline
 209 climatology of 1981-2020. Hatching denotes the two climatologies result in statistically significant
 210 differences at 90% confidence.

211
 212
 213 Based on the regionally area-averaged changes shown in Fig. 2, spatial variability of
 214 changes in selected indices are shown in Fig. 3 (the other indices are shown in Figs. S5-S8). Here,
 215 differences between indices defined with the two baseline periods are averaged over all months in
 216 1980-2021. Figure 3a shows the spatial variability of warm nights (TN90p) averaged over all
 217 months in 1980-2021 using the 1981-2010 climatology period. On average, there are relatively
 218 more warm nights detected in the Southwest US and fewer in the Northern Great Plains. When
 219 the baseline climatology is updated to 1991-2020, TN90p is reduced everywhere throughout the
 220 United States—strongest in the Southwest, and weakest in the Northern Great Plains (Fig. 3b). This
 221 spatial pattern is a result of the change in the 90th percentile of temperature (Fig. 1c) with the 1991-
 222 2020 baseline. Differences are similar when the baseline period is 1981-2020, but with smaller
 223 magnitude throughout the United States (Fig. 3c). For cool days (TX10p), there are
 224 climatologically more in the Northern Great Plains and Midwest, and fewer in the Southwest (Fig.
 225 3d). TX10p increases on average throughout the United States when the baseline period is 1991-

226 2020 (Fig. 3e), except for the Northern Great Plains where differences are small and
 227 insignificant. When the baseline period is 1981-2020, the changes are smaller and not significant
 228 in most regions of the United States, likely due to the thirty-year overlap of the two baseline periods
 229 (Fig. 3f).
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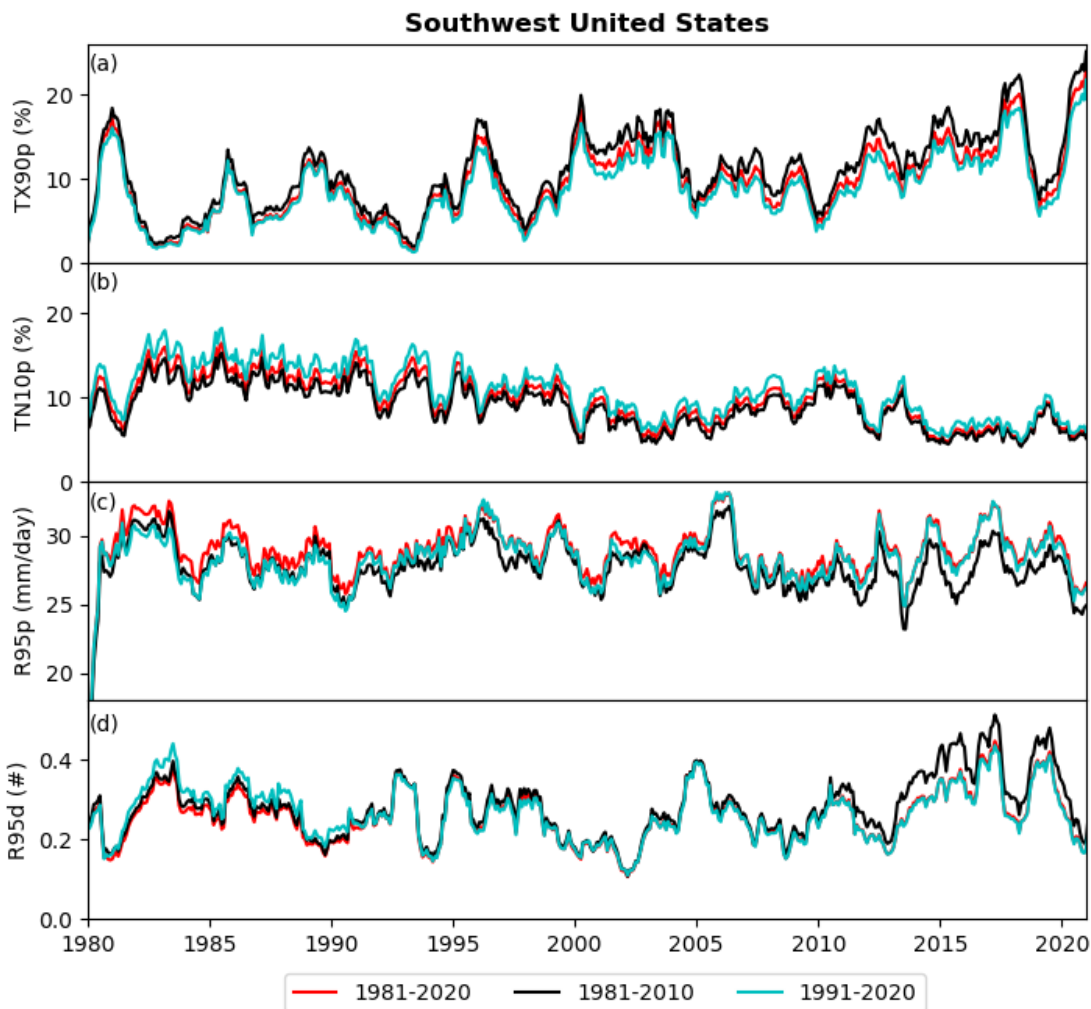


232
 233 **Figure 3.** (a) TN90p defined using percentiles from the 1981-2010 baseline period, averaged over
 234 all months in 1980-2021, (b) the difference between TN90p defined using percentiles from the
 235 1991-2020 baseline period and TN90p defined using percentiles from the 1981-2010 baseline
 236 period, averaged over all months 1980-2021; grey hatching indicates where difference is
 237 significant at the 90% confidence level, (c) the difference between TN90p defined using
 238 percentiles from the 1981-2020 baseline period and TN90p defined using percentiles from the
 239 1981-2010 baseline period, averaged over all months 1980-2021, (d,e,f), (g,h,i), and (j,k,l) as in
 240 (a,b,c) but for TX10p, R99d, and R99p. For readability, panels h-i, k-l are plotted with a 9-point
 241 smoother, i.e., a weighted average of the values of the grid point and the 8 surrounding ones.
 242
 243

244 R99d (days with precipitation above the 99th percentile; Fig. 3g) and R99p (precipitation
 245 on these days; Fig. 3j) are both largest, on average, in the eastern US and pacific northwest. When
 246 the climatology period is updated, R99d is decreased over much of the US, i.e., fewer days with
 247 precipitation above the 99th percentile. Differences are largest and most significant over eastern
 248 Texas, parts of the west and the Midwest US. The changes are larger when the climatology period
 249 is 1981-2020 (Fig. 3i) than 1991-2020 (Fig. 3h). The differences in R99p are less consistent across
 250 the country, and mostly consist of increases in eastern Texas and parts of the Southeast and
 251 Midwest (Fig. 3k-l).

252 Finally, Fig. 4 shows the monthly time series of select indices averaged over the Southwest
 253 region of the US (as shown in Fig. 1a). The Southwest is chosen due to the relatively large changes
 254 observed in this region when updating the baseline period (Fig. 2). Monthly indices are shown as
 255 computed from the three baseline climate periods: 1981-2020 (red line), 1981-2010 (black line)
 256 and 1991-2020 (blue line). For the index representing warm days (TX90p; Fig. 4a), values are
 257 consistently lower when the index is defined using the 1991-2020 percentiles than the 1981-2010
 258 percentiles. The difference between them increases later in the time series, indicating implications
 259 for trends in the indices; Dunn and Morice (2022) showed that positive trends in warm indices

260 such as TX90p were reduced when a later baseline period was used. The opposite is true for cold
 261 nights (TN10p; Fig. 4b), where the newer climatology period results in higher values for the index
 262 throughout the period. For the precipitation indices shown (Fig. 4c-d), differences become most
 263 apparent after 2010, when the 1991-2020 baseline period results in fewer very wet days (Fig. 4d),
 264 but more precipitation on very wet days (Fig. 4c).



265

266 **Figure 4.** Time series of 12-month running means for (a) TX90p, (b) TN10p, (c) R95p, and (d)
 267 R95d area averaged over the Southwest region using a base climatology of 1981-2020 (red
 268 lines), 1981-2010 (black lines), and 1991-2020 (cyan lines).

269

270 4 Conclusions

271 Defining a climatological baseline period is necessary for the computation of percentile-based
 272 extreme climate indices. However, in a non-stationary climate, updating this baseline period leads
 273 to significant changes in the quantification of these climate extremes. In summary, over the United

274 States, updating the baseline period from 1981-2010 to 1991-2020 (or 1981-2020) generally leads
275 to more days identified as a cold temperature extreme, fewer days identified as a warm temperature
276 extreme, and extreme precipitation events that are classified as being less frequent but more
277 intense. There is regional variability in these changes: temperature indices are most affected by the
278 baseline period in the Southwest and Northeast, and least affected in the Northern Great Plains;
279 precipitation changes are localized but typically greatest in the Southeast, Midwest, and inter-
280 mountain west. This work has focused on the United States. However, the effect of the baseline
281 period on the definition of extremes could be even more pronounced in other regions more
282 sensitive to climate change.

283 The goal of this study has been to quantify the changes in detected temperature and
284 precipitation extremes with an updated baseline. The cause of the differences in temperature and
285 precipitation percentiles between the baseline periods is potentially related to several factors. The
286 relative roles of human-induced climate change and multidecadal variability need to be assessed,
287 especially to better quantify how extreme climate indices will change in the future (Sillmann et al.
288 2013a). The changing observing system of the reanalysis (e.g., McCarty et al., 2016) will also be
289 explored in future work. While the focus of this work has been on MERRA-2, future work should
290 involve analysis of extreme indices in other data sets, as Sillmann et al. (2013b) showed these can
291 vary among reanalysis datasets.

292 While it is standard practice to use a 30-year climate baseline period that shifts in time every
293 ten years (Arguez et al., 2012), the results here suggest that it may be useful to consider alternatives
294 for defining climate extremes. The baseline period could be updated more frequently than every
295 ten years, though it is not an easy task for operational centers to update their climatology period
296 every year. To minimize the effect of multidecadal climate variability on extremes, the climatology
297 could be extended to consist of the longest-record possible (Trewin, 2007). However, if one
298 considers that society may adapt to the impacts of extremes over time, a shorter, more recent
299 climatology may be a more logical comparison point. Furthermore, if using an observational record
300 to define a baseline period, it is important to note whether in situ observational sites reported data
301 within the reference period used.

302 The most appropriate baseline likely depends on the application, so data centers could create
303 versions of indices using multiple baseline periods (e.g., Dunn et al., 2020), or provide users with
304 the option to develop their own baseline climatology best suited to their purpose. Some users may
305 need a more frequently updated baseline, while other users may need older baselines retained. It
306 should also be noted that extreme climate events can be defined without percentile-based
307 thresholds, such as using indices with a fixed threshold, though these have limited regional
308 relevance. Methods based on return periods or time of emergence (Lewis et al., 2017) could also
309 be used. Regardless of the approach, it is important to clearly communicate how extremes are
310 defined and interpreted as this choice and the unique statistics produced can influence public
311 perception.

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323

324 **Open Research**

325 CDO is available for download at

326 <https://code.mpimet.mpg.de/projects/cdo/#:~:text=Climate%20Data%20Operators,more%20than%20600%20operators%20available.>

328 MERRA-2 data is publicly available through the GES DISC at

329 <https://disc.gsfc.nasa.gov/information/glossary?title=MERRA-2.>

330 **References**

331 Alexander, L.V., Fowler, H.J., Bador, M., Behrangi, A., Donat, M.G., Dunn, R.D., et al. (2019)
 332 On the use of indices to study extreme precipitation on sub-daily and daily timescales.

333 *Environmental Research Letters*, 14, 125008. <https://doi.org/10.1088/1748-9326/ab51b6>

334

335 Alexander, L. V., Zhang, X., Peterson, T.C., Caesar, J., Gleason, B., Klein Tank, A.M.G., et al.
 336 (2006) Global observed changes in daily climate extremes of temperature and precipitation. *J.*
 337 *Geophys. Res.*, 111, D05109. <https://doi.org/10.1029/2005JD006290>

338

339 Arguez, A., Durre, I., Applequist, S., Vose, R.S., Squires, M.F., Yin, X., et al. (2012) NOAA's
 340 1981-2010 U.S. Climate Normals: An Overview. *BAMS*, 93, 1687-1697.

341 <https://doi.org/10.1175/BAMS-D-11-00197.1>

342

343 Arguez A. & Vose, R.S. (2011) The Definition of the Standard WMO Climate Normal. *BAMS*
 344 92, 699-704. <https://doi.org/10.1175/2010BAMS2955.1>

345

346 Bosilovich, M.G., Akella, S., Coy, L., Cullather, R., Draper, C., Gelaro, R., et al. (2015)
 347 MERRA-2: Initial Evaluation of the Climate. NASA Tech. Memo. NASA/TM-2015-
 348 104606/Vol. 43, 145 pp. [Available online at
 349 <https://gmao.gsfc.nasa.gov/pubs/docs/Bosilovich803.pdf>]

350

351 Cammalleri, C., Spinoni, J., Barbosa, P., Toreti, A. & Vogt, J.V. (2021) The effects of non-
 352 stationarity on SPI for operational drought monitoring in Europe. *International Journal of*
 353 *Climatology*, 42, 1-13. <https://doi.org/10.1002/joc.7424>

354

- 355 Collow, A., Bosilovich, M. G., & Koster, R. D. (2016) Large Scale Influences on Summertime
356 Extreme Precipitation in the Northeastern United States. *J. Hydrometeorol*, *17*, 3045-3061.
357 <https://doi.org/10.1175/JHM-D-16-0091.1>
358
- 359 Collow, A., Thomas, N., Bosilovich, M., Dezfuli, A., & Lucchesi, R. (2022) File Specification
360 for MERRA-2 Climate Statistics Products. GMAO Office Note No. 19 (Version 1.2), 15 pp,
361 available from http://gmao.gsfc.nasa.gov/pubs/office_notes.
362
- 363 Dunn, R.J.H. & Morice, C.P. (2022) On the effect of reference periods on trends in percentile-
364 based extreme temperature indices. *Environmental Research Letters*, *17*,
365 034026. <https://doi.org/10.1088/1748-9326/ac52c8>
366
- 367 Dunn, R.J.H., Alexander, L.V., Donat, M.G., Zhang, X., Bador, M., Herold, N., et al. (2020)
368 Development of an Updated Global Land In Situ-Based Data Set of Temperature and
369 Precipitation Extremes: HadEX3. *J. Geophys. Res. Atmos.*, *125*.
370 <https://doi.org/10.1029/2019JD032263>
371
- 372 Dunn, R.J.H., Donat, M.G., & Alexander, L.V. (2022) Comparing extremes indices in recent
373 observational and reanalysis products. *Front. Clim.*, *4*, 989505. doi: 10.3389/fclim.2022.989505
374 Gelaro, R., McCarty, W., Suarez, M.J., Todling, R., Molod, A., Takacs, L., et al. (2017) The
375 Modern-Era Retrospective Analysis for Research and Applications, Version 2 (MERRA-2). *J.*
376 *Clim.*, *30*, 5419-5454. <https://doi.org/10.1175/JCLI-D-16-0758.1>
377
- 378 Global Modeling and Assimilation Office (GMAO), 2015a: MERRA-2 statD_2d_slv_Nx: 2d,
379 Daily, Aggregated Statistics, Single-Level, Assimilation, Single-Level Diagnostics V5.12.4,
380 Greenbelt, MD, USA, Goddard Earth Sciences Data and Information Services Center (GES
381 DISC), doi: 10.5067/9SC1VNTWGWV3.
382
- 383 Global Modeling and Assimilation Office (GMAO), 2015b: MERRA-2 tavg1_2d_flux_Nx: 2d, 1-
384 Hourly, Time-Averaged, Single-Level, Assimilation, Surface Flux Diagnostics V5.12.4,
385 Greenbelt, MD, USA, Goddard Earth Sciences Data and Information Services Center (GES
386 DISC), doi: 10.5067/7MCPBJ41Y0K6.
387
- 388 Global Modeling and Assimilation Office (GMAO) (2020) MERRA-2 statM_2d_edi_Nx: 2d,
389 Single-Level, Monthly Extremes Detection Indices V1, Greenbelt, MD, USA, Goddard Earth
390 Sciences Data and Information Services Center (GES DISC), doi: 10.5067/QFJ13GEGDI99.
391
- 392 Global Modeling and Assimilation Office (GMAO) (2022) MERRA-2 statM_2d_edi_Nx: 2d,
393 Single-Level, Monthly Extremes Detection Indices based on 1991-2020 V2, Greenbelt, MD,
394 USA, Goddard Earth Sciences Data and Information Services Center (GES DISC),
395 doi: 10.5067/O8AX56DO60MI.
396
- 397 Kendon, M., McCarthy, M., Jevrejeva, S., Matthews, A., Sparks, T. & Garforth, J. (2020) State
398 of the UK Climate 2020. *International Journal of Climatology*, *41*, 1-76.
399 <https://doi.org/10.1002/joc.7285>
400

- 401 Lewis, S.C., King, A.D., & Perkins-Kirkpatrick, S.E. (2017) Defining a New Normal for
402 Extremes in a Warming World. *BAMS*, 98, 1139-1151. <https://doi.org/10.1175/BAMS-D-16->
403 0183.1
- 404
- 405 Livezey, R.E., Vinnikov, K.Y., Timofeyeva, M.M., Tinker, R. & van den Dool, H.M. (2007)
406 Estimation and Extrapolation of Climate Normals and Climatic Trends. *Journal of Applied*
407 *Meteorology and Climatology*, 46, 1759-1776. <https://doi.org/10.1175/2007JAMC1666.1>
- 408
- 409 McCarty, W., Coy, L., Gelaro, R., Huang, A., Merkova, D., Smith, E.B., et al. (2016) MERRA-2
410 input observations: Summary and initial assessment. Technical Report Series on Global
411 Modeling and Data Assimilation, Vol. 46, NASA Tech. Rep. NASA/TM–2016–104606, 61 pp.
412 [Available online at
413 <https://gmao.gsfc.nasa.gov/pubs/docs/McCarty885.pdf>.] [https://www.ncei.noaa.gov/products/lan-](https://www.ncei.noaa.gov/products/land-based-station/us-climate-normals)
414 [d-based-station/us-climate-normals](https://www.ncei.noaa.gov/products/land-based-station/us-climate-normals)
- 415
- 416 National Centers for Environmental Information (NCEI) (2021, May 4) NOAA Delivers New
417 U.S. Climate Normals, <https://www.ncei.noaa.gov/news/noaa-delivers-new-us-climate-normals>,
418 last accessed 29 June 2023.
- 419
- 420 Perkins, S. E., & Alexander, L. V. (2013). On the Measurement of Heat Waves, *Journal of*
421 *Climate*, 26(13), 4500-4517. <https://doi.org/10.1175/JCLI-D-12-00383.1>.
- 422
- 423 Reichle., R.H., Liu, Q., Koster, R.D., Draper, C.S., Mahanama, S.P.P. & Partyka, G.S. (2017)
424 Land Surface Precipitation in MERRA-2. *Journal of Climate*, 30, 1643-1664.
425 <https://doi.org/10.1175/JCLI-D-16-0570.1>
- 426
- 427 Scherrer, S.C., Appenzeller, C., & Liniger, M.A. (2006) Temperature trends in Switzerland and
428 Europe: Implications for Climate Normals. *International Journal of Climatology*, 26, 565-580.
429 <https://doi.org/10.1002/joc.1270>
- 430
- 431 Schreck III, C.J., Klotzbach, P.J., & Bell, M.M. (2021) Optimal climate normal for North
432 Atlantic Hurricane activity. *Geophysical Research Letters*, 48, 1-9.
433 <https://doi.org/10.1029/2021GL092864>
- 434
- 435 Schulzweida, U. (2022) CDO User Guide (2.1.0). Zenodo.
436 <https://doi.org/10.5281/zenodo.7112925>
- 437
- 438 Sillmann, J., Kharin, V. V., Zwiers, F. W., Zhang, X., and Bronaugh, D. (2013a), Climate
439 extremes indices in the CMIP5 multimodel ensemble: Part 2. Future climate projections, *J.*
440 *Geophys. Res. Atmos.*, 118, 2473- 2493, doi:10.1002/jgrd.50188
- 441
- 442 Sillmann, J., Kharin, V. V., Zhang, X., Zwiers, F. W., and Bronaugh, D. (2013b), Climate
443 extremes indices in the CMIP5 multimodel ensemble: Part 1. Model evaluation in the present
444 climate, *J. Geophys. Res. Atmos.*, 118, 1716- 1733, doi:10.1002/jgrd.50203
- 445
- 446 Trewin, B. (2007) The role of climatological normals in a changing climate. WCDMP No. 61 –
447 WMO/TD-No. 1377. https://library.wmo.int/doc_num.php?explnum_id=4546

- 448
449 Thomas, N. P., Bosilovich, M. G., Collow, A. B. M., Koster, R. D., Schubert, S. D., Dezfuli, A.,
450 & Mahanama, S. P. (2020) Mechanisms associated with Daytime and Nighttime Heat Waves
451 over the Contiguous United States. *J. Appl. Meteorol. Clim.*, 59 (11), 1865-1882.
452 <https://doi.org/10.1175/JAMC-D-20-0053.1>
453
- 454 Wilks, D.S. (2013) Projecting “Normals” in a Nonstationary Climate. *Journal of Applied*
455 *Meteorology and Climatology*, 52, 289-302. <https://doi.org/10.1175/JAMC-D-11-0267.1>
456
- 457 Wilks, D.S. & Livezey, R.E. (2013) Performance of Alternative “Normals” for Tracking Climate
458 Changes, Using Homogenized and Nonhomogenized Seasonal U.S. Surface
459 Temperatures. *Journal of Applied Meteorology and Climatology*, 52, 1677-
460 1687. <https://doi.org/10.1175/JAMC-D-13-026.1>
461
- 462 Wuebbles, D.J., Fahey, D.W., Hibbard, K.A., Dokken, D.J., Stewart, B.C., & Maycock, T.K.
463 Eds. (2017) Climate Science Special Report: Fourth National Climate Assessment. U.S. Global
464 Change Research Program. Vol. I. 470 pp., <https://doi.org/10.7930/J0J964J6>
465
- 466 Yosef, Y., Agilar, E., & Alpert, P. (2021) Is it possible to fit extreme climate change indices
467 together seamlessly in the era of accelerated warming? *Int. J. Climatol.*, 41.
468 <https://doi.org/10.1002/joc.6740>
469
- 470 Zhang, X., Hegerl, G., Zwiers, F. W., & Kenyon, J. (2005) Avoiding inhomogeneity in
471 percentile-based indices of temperature extremes. *J. Climate*, 18, 1641–1651.
472 <https://doi.org/10.1175/JCLI3366.1>
473
- 474 Zhang, X., Alexander, L., Hegerl, G.C., Jones, P., Tank, A.K., Peterson, T.C., & Zwiers, F.W.
475 (2011) Indices for monitoring changes in extremes based on daily temperature and precipitation
476 data. *Wiley Interdiscip. Rev.: Clim Change*, 2, 851-70. <https://doi.org/10.1002/wcc.147>