

Using Soil Moisture Information to Better Understand and Predict Wildfire Danger: A Review of Recent Developments and Outstanding Questions

1 **Running Head:** SOIL MOISTURE-WILDFIRE DANGER REVIEW

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16 **Keywords:** Soil moisture, wildfire, review, in situ, remote sensing, modeling

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17 **Abstract**

18 Soil moisture conditions are represented in existing fire danger rating systems mainly through
19 relatively simple drought indices based on meteorological variables, even though better sources
20 of soil moisture information are increasingly available and have been shown to help improve
21 predictions of fuel loads, fuel moisture, wildfire probability, and wildfire size. Without
22 operational use of this soil moisture information, the potential for more accurate and timely fire
23 danger warnings is unrealized, while increasing wildfire activity harms human and natural
24 systems in various regions around the world. This review summarizes a growing body of
25 evidence indicating that greater utilization of in situ, remotely sensed, and modeled soil moisture
26 information in fire danger rating systems could lead to better estimates of dynamic live and dead
27 fuel loads, more accurate live and dead fuel moisture predictions, earlier warning of elevated
28 wildfire danger, and more precise forecasts of wildfire occurrence and severity. Although
29 important research questions remain, several of which are identified here, the path forward is
30 clear. Soil moisture information can and should be used to improve fire danger rating systems
31 and contribute to more effective fire management for the protection of communities and
32 ecosystems worldwide.

33

34 **Summary:** Soil moisture is an underused resource for improving fire danger rating systems and
35 fire management worldwide. We review key studies describing relationships between wildfires
36 and in situ, remotely sensed, and modeled soil moisture; describe the potential to incorporate soil
37 moisture into wildfire danger assessments; and identify outstanding challenges and opportunities.

38

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39 Introduction

40 At 6:33 a.m. on the morning of 8 November 2018, a small fire was reported under
41 electrical power lines near Camp Creek Road outside the town of Pulga in northern California,
42 USA. Dry conditions and strong downslope winds with gusts $>25 \text{ m s}^{-1}$ (Brewer and Clements
43 2020) rapidly transformed that small fire into the deadliest and most costly wildfire in
44 California's history. The Camp Fire burned $>62,000 \text{ ha}$, destroyed $>18,000$ structures, and
45 resulted in 85 fatalities (California Department of Forestry and Fire Protection 2019). This
46 tragedy powerfully illustrates the importance of fire danger rating systems and the need to
47 provide earlier and more accurate warnings for fire management agencies and the public. Toward
48 that end, this review explores recent developments, data gaps, and challenges in applying
49 previously underutilized soil moisture information to better understand, assess, and predict
50 wildfire danger. Up until now, the incorporation of soil moisture information into existing fire
51 danger rating systems has been limited to simplistic models or indices which use standard
52 weather variables to estimate soil moisture, even though such information is becoming
53 increasingly available via in situ measurements, remote sensing, and more sophisticated
54 modeling. One week prior to the tragic Camp Fire, for example, satellite observations showed
55 strong negative soil moisture anomalies across northern California (Fig. 1), conditions that are
56 known to substantially increase the probability of large wildfires (Krueger et al. 2015; Krueger et
57 al. 2016; Sazib et al. 2021), **but the tools needed to effectively put this information into**
58 **action are currently lacking.**

59 **That soil moisture conditions are important for fire danger rating is not a recent**
60 **revelation. Prominent fire danger rating systems in Canada, Australia, and the United**
61 **States use approximations of the moisture of mineral or organic soil horizons to quantify**

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62 wildfire danger (Kumar and Dharssi 2015). For example, the Canadian Forest Fire
63 Danger Rating System (CFFDRS) (Stocks et al. 1989; Wotton 2009) includes three
64 moisture indices, termed moisture codes, to represent moisture stored in the organic layers
65 of the forest floor. The Fine Fuel Moisture Code (FFMC) represents fuel moisture of fine
66 surface litter with a depth of 10-20 mm, the Duff Moisture Code (DMC) represents fuel
67 moisture of loosely compacted duff with a nominal depth of 50-100 mm, and the Drought
68 Code (DC) represents fuel moisture of deep organic materials having a nominal depth of
69 100-200 mm (de Groot 1987) (Fig. 2). While intended to represent moisture of surface
70 organic layers, DMC and DC are strongly correlated to soil moisture of mineral horizons
71 near the surface (D'Orangeville et al. 2016; Pellizzaro et al. 2007), likely in part because of
72 capillary and vapor flow between mineral and organic soil layers (Zhao et al. 2022). When
73 considering soils with deep organic layers at the surface, i.e., deep O horizons in soil science
74 terminology, the water stored in those layers may be viewed as either soil moisture or fuel
75 moisture because the organic layer itself can become combustible at low water contents.

76 In the recently modified Australian Fire Danger Rating System (Matthews 2022),
77 fire danger ratings for dry eucalypt forests are dependent in part on soil moisture deficit
78 estimated using the Keetch-Byram Drought Index (KBDI, Keetch and Byram 1968). KBDI
79 uses temperature and precipitation data to estimate the moisture deficit in the upper soil
80 layers (mineral and organic, if present) using a water balance approach. KBDI was
81 designed to represent approximately the top 760-890 mm for a fine-textured soil and
82 greater depths for coarser-textured soils (Keetch and Byram 1968) (Fig. 2). Similarly, in
83 the National Fire Danger Rating System (NFDRS) used in the United States (Bradshaw et
84 al. 1983; Burgan 1988; Deeming et al. 1972; Jolly 2018), KBDI helps determine fire danger

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85 ratings through its influence on fuel load. The inclusion of these moisture indices in widely
86 used fire danger rating systems makes it clear that their developers recognized the
87 importance of soil moisture for understanding wildfire danger. But at the time these
88 systems were developed, large-scale soil moisture measurement systems and physically-
89 based hydrologic models were not sufficiently developed, thus moisture indices were
90 instead estimated from commonly measured weather data. Given advances in soil moisture
91 measurement and modeling systems in recent decades, there is a need to reassess how to
92 best represent the moisture of organic and mineral soil layers in fire danger ratings systems
93 and to better understand the effects of those representations on the accuracy of fire danger
94 ratings.

95 The effectiveness of fire danger rating systems can be determined through retrospective
96 analyses of the relationship between fire danger ratings **and important wildfire metrics**
97 **including** occurrence and size. For example, a recent analysis of the NFDRS showed generally
98 positive correlations between fire danger ratings and fire sizes across the contiguous US, but
99 there were important spatial inconsistencies. Notably, there was poorer performance in the
100 eastern half of the country compared to the western half, possibly due to regional differences in
101 soil-vegetation-climate interactions and in the timing and length of the fire season (Walding et al.
102 2018). **Furthermore, large areas in the central US lacked the necessary data to generate fire**
103 **danger ratings because those areas contained no reporting stations for the Weather**
104 **Information Management System, which provides weather data for the NFDRS** (Walding et
105 al. 2018). Thus, improvements to the NFDRS will likely need to consider both model structural
106 improvements, as well as new and better sources and types of input data. Currently, the NFDRS
107 and most other fire danger rating systems in use around the world rely on a relatively standard

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108 set of input variables that are routinely measured at weather stations, chief among them being air
109 temperature, relative humidity, precipitation, and wind speed (de Groot et al. 2015). **However,**
110 **as described in this review, a growing body of research provides strong evidence that new**
111 **sources of soil moisture information can be important predictors of wildfire danger, but**
112 **this valuable information has not yet been effectively integrated into fire danger rating**
113 **systems.**

114 Historically, a major hindrance to such integration has been the limited availability of soil
115 moisture information with adequate duration and spatial extent. That situation is rapidly
116 changing as a variety of new sources of soil moisture information are becoming available, **each**
117 **with unique strengths (Fig. 3). These new data sources include 1) soil moisture measured in**
118 **situ, 2) soil moisture measured remotely by satellites, and 3) soil moisture data that is**
119 **generated using physically-based models.** This groundswell of information began with the
120 advent of state and national automated soil moisture monitoring networks in the US in the late
121 **1990s** and the subsequent emergence of similar networks in other countries around the world
122 (Dorigo et al. 2021). In parallel, satellite missions capable of monitoring soil moisture and
123 closely-related variables have been developed and launched by NASA and other space agencies,
124 with substantial increases in daily coverage of **the** Earth's surface since the late **1990s**
125 (Karthikeyan et al. 2017). These advances in soil moisture measurements have occurred
126 alongside advances in numerical soil moisture models, which can now provide soil moisture
127 estimates for large domains with sub-km resolution (Holden et al. 2019). Using **these three**
128 **types of soil moisture information,** researchers began to generate first glimpses of the strong
129 relationships between wildfire and in situ **soil moisture** (Krueger et al. 2015), remotely-sensed
130 **soil moisture** (Bartsch et al. 2009), and modeled **soil moisture** (Slocum et al. 2010). Subsequent

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131 studies have provided new insights into the **relationships between soil moisture and fuel**
132 **characteristics including fuel loads (e.g., Ellsworth et al. 2013; Sharma et al. 2018), curing**
133 **(e.g., Sharma et al. 2021; Wittich 2011), and live (e.g., Bianchi and Defossé 2015; Fan et al.**
134 **2018) and dead fuel moisture (e.g., Masinda et al. 2021; Rakhmatulina et al. 2021). Other**
135 **studies have directly related soil moisture to fire occurrence (e.g., Jensen et al. 2018;**
136 **Vinodkumar and Dharssi 2019) and fire size (e.g., Forkel et al. 2012; Krueger et al. 2015;**
137 **Slocum et al. 2010), while still others have identified the impact of vegetation type on soil**
138 **moisture-wildfire relationships (e.g., Rigden et al. 2020; Schaefer and Magi 2019). These**
139 **and other important contributions to our understanding of soil moisture—wildfire relationships**
140 **have emerged across a wide variety of scientific disciplines, which often are not well-connected,**
141 **making the accelerating progress difficult to track and synthesize.**

142 **A further roadblock complicating the use of soil moisture information for fire**
143 **danger ratings is that soil moisture conditions can be expressed in a variety of ways,**
144 **making it more difficult to compare results across studies. For example, soil moisture can**
145 **be expressed simply as soil volumetric water content (e.g., Ambadan et al. 2020; Schaefer**
146 **and Magi 2019; Vinodkumar et al. 2021) or water content summed over some soil depth,**
147 **i.e., soil water storage (e.g., Chikamoto et al. 2015; Krawchuk and Moritz 2011; Slocum et**
148 **al. 2010). Alternatively, soil moisture can be formulated to represent the amount of soil**
149 **moisture available to plants (Krueger et al. 2019), and it may also be normalized to allow**
150 **for comparison across sites or across different soil moisture metrics. This normalization**
151 **procedure may be based on the physical properties of the soil (e.g., Krueger et al. 2015;**
152 **Vinodkumar et al. 2017; Waring and Coops 2016) or use statistical techniques (Lyons et al.**
153 **2021). To further complicate the situation, soil moisture may be expressed across different**

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154 soil depths (Fan et al. 2018; Vinodkumar et al. 2021) and as absolute values or anomalies
155 (O et al. 2020; Sazib et al. 2021). These varied formulations of soil moisture must be
156 understood to properly interpret the growing body of literature establishing the important
157 relationships between soil moisture and wildfire.

158 Therefore, our objectives are to 1) summarize the rapidly growing body of research on
159 soil moisture—wildfire relationships, 2) broaden the community of researchers aware of and
160 engaged in this line of research, and 3) make a convincing case for more widespread use of soil
161 moisture information in operational fire danger rating systems. This review is organized into four
162 primary sections. The first three sections summarize what is known about the relationships of
163 wildfire and fuel bed properties to 1) in situ soil moisture measurements, 2) remotely sensed soil
164 moisture, and 3) modeled soil moisture. The fourth section explains potential links between soil
165 moisture information and existing fire danger rating systems, using NFDRS as one specific
166 example. We conclude by describing primary challenges and opportunities for using soil
167 moisture information to better understand and predict wildfire danger, including the
168 identification of key areas of needed future research.

169

170 **In situ soil moisture measurements**

171 In situ soil moisture measurements are the gold standard of soil moisture information
172 (Levi et al. 2019) against which remote sensing and modeled values are evaluated (Fig. 3), and in
173 some geographic areas, in situ soil moisture data are available at sufficient spatial and temporal
174 resolutions to inform wildfire management. The International Soil Moisture Network houses
175 publicly available data from nearly 2700 in situ soil moisture monitoring stations across 65
176 networks worldwide, a number that is steadily growing (Dorigo et al. 2021). The United States

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177 has an especially prolific collection of in situ soil moisture monitoring networks, and the ongoing
178 National Coordinated Soil Moisture Monitoring Network initiative aims to produce harmonized
179 data products from in situ soil moisture measurements from approximately 2000 sites across the
180 nation (Cosh et al. 2021). One of the longest running and densest large-scale soil moisture
181 monitoring networks in the US, and in the world, is the Oklahoma Mesonet (McPherson et al.
182 2007; Ochsner et al. 2013). Oklahoma is also consistently among the top 10 states in the US for
183 wildfire risk (III 2021); accordingly, data from Oklahoma has proven valuable for understanding
184 soil moisture-wildfire relationships.

185 A striking example of the connections between soil moisture, fuel bed properties, and
186 wildfire comes from the Marena, Oklahoma, In Situ Sensor Testbed (MOISST) located in north-
187 central Oklahoma. The MOISST site was established in 2010 to compare in situ soil moisture
188 sensing technologies (Cosh et al. 2016) and measure vegetation dynamics in tallgrass prairie
189 (PhenoCam 2021), with fuel bed properties measured at and around the site (Sharma et al. 2018).
190 PhenoCam images collected at the site show markedly different vegetation conditions during
191 August of 2012 and 2013 (Fig. 4). Drought conditions from May through July 2012 resulted in a
192 fuel moisture content for mixed live and dead fuels of only 27% in early August when the photo
193 on the left was taken. The severity of the drought was reflected in the measured soil moisture,
194 expressed as fraction of available water capacity (FAW). **FAW is a measure of plant-available**
195 **water that is calculated based on measured volumetric water content and the available**
196 **water capacity of the soil (Krueger et al. 2015), and it can be determined for any landscape**
197 **(e.g., grassland, forest, cropland) for which these variables are known.** It is defined as the
198 ratio of measured plant available water to the maximum plant available water capacity of the
199 soil, and it typically ranges from 0 (no plant available water) to 1 (maximum plant available

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200 water). From May through July 2012, FAW averaged only 0.23 (**i.e., plant available water was**
201 **at 23% of its possible maximum**), levels indicative of severe drought (Sridhar et al. 2008). In
202 contrast, FAW averaged 0.82 (**i.e., 82% of possible maximum**) over the same period in 2013,
203 which corresponded with green vegetation in August 2013 (photo on the right) and a mixed fuel
204 moisture content of 101%. The low fuel moisture contents in August 2012 contributed to
205 extreme wildfire danger and the devastating Freedom Hill Fire, which ignited approximately 80
206 km east of the MOISST site the same day the photo was taken. This fire burned nearly 24,000
207 hectares of mostly prairie, savanna, and woodland over a two-week period; destroyed more than
208 300 homes; and resulted in Federal Emergency Management Agency assistance claims totaling
209 more than \$7 million.

210 The qualitative soil moisture-fuel bed relationships that are clear in Figure 4, and may be
211 intuitive to fire managers, have been described in detail by recent research based on in situ soil
212 moisture measurements. The soil moisture-fuel moisture relationship was quantified for various
213 shrub species in Italy by Pellizzaro et al. (2007), who found that soil moisture was a better
214 predictor of live fuel moisture than weather variables or weather-derived drought indices. Their
215 finding was corroborated by Qi et al. (2012), who found that soil moisture explained 66% of the
216 variability in live fuel moisture for oak and sagebrush in northern Utah, and soil moisture was
217 more strongly correlated with live fuel moisture than were remotely sensed vegetation indices.
218 Similar linear relationships between soil moisture and fuel moisture have also been reported for
219 grassland fuels in South Africa (McGranahan et al. 2016).

220 These findings have been corroborated by a series of studies in Oklahoma, the key results
221 of which are summarized in Figure 5. Sharma et al. (2021), using data from a grassland field
222 study close to the MOISST site, reported that when soil moisture was plentiful (FAW values of

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223 at least 0.59), mixed fuel moisture was not related to soil moisture, but mixed fuel moisture
224 declined as FAW decreased below this threshold. When FAW dropped below 0.40, the
225 transpiration and growth rate of grassland live fuels declined, reflecting the intensification of
226 drought stress (Krueger et al. 2021). When FAW declined below 0.36, the greenness of the
227 vegetation, as indicated by the normalized difference vegetation index (NDVI), began to
228 decrease (Sharma et al. 2021) (Figure 5). At a still lower FAW threshold of 0.30, the transition
229 of live fuel to dead (i.e., curing rate) increased rapidly, from near $0 \text{ g m}^{-2} \text{ day}^{-1}$ when FAW was $>$
230 0.30 to more than $10 \text{ g m}^{-2} \text{ day}^{-1}$ as FAW approached 0.20 (Sharma et al. 2021). This drought-
231 induced curing is vividly depicted in Figure 4, with extremely low soil moisture corresponding
232 with vegetation that was almost completely cured by early August 2012, while little curing had
233 occurred by the same time in 2013 when soil moisture was plentiful. A perhaps subtler
234 distinction in fuel bed characteristics between these years is that the live fuel load in 2013 was
235 more than double that in 2012, which portended potentially high wildfire activity if dry and
236 windy conditions prevailed during the subsequent dormant season. These findings offer a
237 physical explanation for the observed dependence of growing season wildfire size and
238 probability on soil moisture conditions (Fig. 5).

239 In a different study that used in situ soil moisture data from the entire state of Oklahoma,
240 Krueger et al. (2015) showed that 90% of large growing season wildfires across all Oklahoma
241 landscapes (**forest, shrubland, grassland**) occurred when FAW was < 0.40 , which matches the
242 threshold for transpiration reduction due to moisture stress in grassland vegetation (Krueger et al.
243 2021). These soil moisture-wildfire relationships were further described using probabilistic
244 models in a subsequent study (Krueger et al. 2016). When plant available soil moisture was near
245 its maximum, the probability of a large growing season wildfire across all Oklahoma landscapes

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246 was near zero even when temperature, wind speed, and relative humidity conditions were ripe for
247 wildfires (Fig. 4 and Krueger et al. 2016). As FAW decreased to 0.59, the soil moisture
248 threshold below which grassland fuel moisture decreases, wildfire probability increased to 0.10,
249 and for a FAW value of 0.30, the threshold for rapid fuel curing, wildfire probability more than
250 tripled to 0.44 (Fig. 5). These results suggest that soil moisture and weather conditions work in
251 concert to support high growing season wildfire probability. Low soil moisture is associated
252 with decreased fuel moisture and accelerated curing, while high temperatures, low relative
253 humidity, and high wind speed facilitate fire ignition and spread.

254 When vegetation is dormant, however, current FAW levels were not a strong predictor of
255 the probability of large wildfires in Oklahoma (Krueger et al. 2016), likely because in grasslands
256 dead fuel moisture content was not strongly dependent on soil moisture (Sharma et al. 2021). But
257 dormant season wildfire probability was increased by high soil moisture during the previous
258 growing season. For example, when FAW during the growing season was at least 0.40, the
259 probability of a large wildfire during the subsequent dormant season was approximately double
260 compared with growing season FAW values near 0.20 (Krueger et al. 2016). Vegetation
261 productivity, at least for Oklahoma grasslands, is maximized when $FAW > 0.40$ (Krueger et al.
262 2021), contributing to increased fine fuel loads in the subsequent dormant season.

263 Although there is a lack of evidence for soil moisture effects on dead fuel moisture in
264 grasslands, in situ measurements from a diverse array of sites around the world reveal important
265 links between soil moisture and dead fuel moisture **for surface fuels** in forests. In Australia, the
266 influence of soil moisture on the fuel moisture content of fine dead fuels, i.e., leaf litter, was
267 observed in plantations of Monterey pine (*Pinus radiata*) approximately three decades ago (Pook
268 and Gill 1993). The fuel moisture content for the pine needle litter on the surface was positively

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269 correlated with measured soil moisture in the 0-40 cm soil layer, and the correlation was stronger
270 for un-thinned and un-pruned stands ($r = 0.91$) than in thinned and pruned stands ($r = 0.45$). The
271 fuel moisture content of the surface leaf litter was predicted more accurately when soil moisture
272 data were included along with temperature and humidity data in a multiple regression model
273 compared to a similar model without soil moisture data. More recently in Australia, in situ soil
274 moisture measurements have been linked to the fuel moisture content of the **surface and**
275 **subsurface** litter layer under various *Eucalyptus* species (Zhao et al. 2021). A follow-up
276 experiment showed that dry soil had a limited influence on the fuel moisture content of the litter,
277 primarily through vapor flow between the soil and the litter (Zhao et al. 2022). In contrast, wet
278 soil had a stronger influence on litter moisture content, with evidence for both vapor and
279 capillary flow between the soil and the litter. Similarly, in situ measurements from forested sites
280 in the foothills of the Sierra Nevada in central California showed that soil moisture had a
281 stronger influence than any other environmental or meteorological factor on fuel moisture of 10-
282 h fuels (6 to 25 mm diameter dead fuels) for wet soil conditions (Rakhmatulina et al. 2021). A
283 dominant influence of soil moisture on the moisture content of dead fine fuels was also
284 documented through in situ measurements in Korean pine (*Pinus koraiensis*) and Scots pine
285 (*Pinus sylvestris*) stands in northeastern China (Masinda et al. 2021). **These reports of the**
286 **connection between moisture of mineral soils and that of overlying organic layers**
287 **corroborate previous studies correlating soil moisture measurements with moisture codes**
288 **from the CFFDRS. For example, correlation coefficients ranging from 0.6-0.8 were**
289 **reported between measured soil moisture and soil moisture estimated from the DC index in**
290 **Canadian forests (D'Orangeville et al. 2016).**

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291 **These findings from diverse ecosystems and geographies** highlight the dependencies
292 of fuel loads, fuel moisture content, and wildfire probability on soil moisture. They have also laid
293 the groundwork for a new generation of wildfire danger assessment tools that use in situ soil
294 moisture information. However, even with expanding national and regional scale soil moisture
295 monitoring networks, using in situ data for wildfire danger monitoring and management
296 decisions is still constrained by the limited number of measurement sites in some locations [e.g.,
297 **the boreal forest, and most of South America, Africa, and Australia (Dorigo et al. 2021)**].
298 And because soil moisture can vary greatly across even small distances (Famiglietti et al. 2008),
299 point measurements of soil moisture are not necessarily representative of soil moisture at the
300 landscape scale (Fig. 3). Therefore, there is a clear need for supplemental strategies for
301 quantifying soil moisture, which include remotely sensed and modeled soil moisture information.

302

303 **Remotely sensed soil moisture**

304 Remote sensing technology has advanced rapidly since the first photograph of Earth was
305 taken from space in 1946. Since that time, improvements in sensor fidelity, satellite and rocket
306 launch technology, data storage, and aperture development have enabled many new capabilities,
307 including near real-time operations related to earth sciences and hydrology (McCabe et al. 2017).
308 The ability to characterize the land surface using strategic regions of the electromagnetic
309 spectrum has resulted in opportunities to remotely monitor and assess near-surface soil moisture
310 and vegetation dynamics (Kumar et al. 2020; Mladenova et al. 2020), which are key to
311 understanding the risks and impacts of wildfires. With the advent of refined satellite-based
312 microwave sensors such as the European Space Agency's Soil Moisture Ocean Salinity (SMOS)
313 mission, which launched in 2009 (Kerr et al. 2010), and NASA's Soil Moisture Active-Passive

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314 (SMAP) mission (Entekhabi et al. 2010), which launched in 2015, **evidence is beginning to**
315 **emerge that** satellite-based soil moisture data **can provide** value for understanding and
316 predicting wildfire danger **in many ecosystems** (O et al. 2020).

317 Remotely sensed soil moisture data have proven useful for assessing fuel bed properties
318 including biomass accumulation (i.e., fuel production) and fuel moisture content. For example, in
319 southern France live fuel moisture measurements for Mediterranean shrub species were
320 significantly correlated with the preceding 15-day average remotely sensed soil moisture from
321 the European Space Agency's Climate Change Initiative Soil Moisture dataset (ESA CCI SM,
322 formerly known as ESV SM) (Fan et al. 2018). A subsequent study used soil moisture data from
323 SMAP to estimate live fuel moisture of chamise (*Adenostoma fasciculatum*) at 12 chaparral sites
324 in southern California (Jia et al. 2019). At those sites, a statistical model using weighted,
325 accumulative soil moisture and growing degree days outperformed models using vegetation
326 optical depth or other optical indices. There is also some evidence that remotely sensed soil
327 moisture might be useful for estimating dead fuel moisture. Burapapol and Nagasawa (2016)
328 reported that remotely sensed soil moisture based on Landsat and MODIS was closely linked
329 with fuel moisture **of dead leaves** in dipterocarp and deciduous forests in Thailand. Soil
330 moisture based on microwave remote sensing may be preferable to optical reflectance indices
331 commonly used to characterize fuel moisture [see reviews by Gale et al. (2021); Yebra et al.
332 (2013); and Arroyo et al. (2008)] because microwave sensors are less prone to disturbances from
333 unfavorable weather (e.g., clouds) and because soil moisture is physiologically linked to plant
334 processes (Nolan et al. 2020).

335 The results of the above regional studies (Fan et al., 2018; Jia et al., 2019) were supported
336 by a nationwide analysis of the ESA CCI SM data and live fuel moisture at >1000 sites across

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337 the contiguous US (Lu and Wei 2021). That analysis spanned numerous vegetation types and
338 climate zones and revealed that the correlations between soil moisture and live fuel moisture
339 were typically strongest when soil moisture was measured 10-50 days in advance. Important
340 vegetation types showing a relatively high sensitivity to soil moisture included pine, redcedar,
341 sagebrush, oak, manzanita, chamise, mesquite, and juniper. The SMAP Level-4 surface and root
342 zone soil moisture products, which result from assimilation of SMAP observations into a land
343 surface model, and in situ soil moisture measurements at selected sites both showed somewhat
344 stronger correlations with live fuel moisture than did the ESA CCI SM data.

345 The links between remotely sensed soil moisture data and fuel bed characteristics make
346 those data useful for assessing wildfire danger. For example, positive soil moisture anomalies
347 observed by Earth Resources Satellite 1 and 2 corresponded with a lower burned area of forest
348 fires in the **boreal forest** of Siberian (Bartsch et al. 2009). **Furthermore, extreme fire events in**
349 **this region were more closely associated with remotely sensed soil moisture [AMSR-E**
350 **(Njoku et al. 2003)] than precipitation anomalies or fire danger indices (Forkel et al. 2012).**
351 More recently, SMOS observations over boreal Canada revealed that wildfires occurred more
352 frequently in anomalously low soil moisture conditions (Ambadan et al. 2020). At more
353 southerly latitudes, models using SMOS-derived soil moisture, in conjunction with temperature
354 and site specific variables such as land cover type, explained 68% of variability of maximum fire
355 extent on the Iberian Peninsula (Chaparro et al. 2016). And the inclusion of SMAP soil moisture
356 observations increased skill in predicting wildfire occurrence in the western US relative to the
357 use of vapor pressure deficit alone, particularly in grasslands (Rigden et al. 2020).

358 Because current soil moisture conditions can influence future fuel moisture and fuel load,
359 soil moisture observations may be particularly helpful for forecasting wildfire danger. In

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360 Australia and California, for example, Sazib et al. (2021) found that soil moisture from SMAP
361 was negatively correlated with wildfires at 1-2 month lead times in moist regions where fuels are
362 typically plentiful, and it was positively correlated with wildfires in drier regions where fuel is
363 scarce. These trends were attributed to a decrease in moisture of surface fuels in moist regions
364 and increased biomass accumulation in dry regions. In an analysis that spanned the globe, O et
365 al. (2020) found that soil moisture from ECV-SM was an important early predictor of wildfires.
366 They reported that in arid regions positive soil moisture anomalies corresponded with increased
367 biomass accumulation followed by wildfire outbreaks at lead times of 5 months. In humid
368 regions, negative soil moisture anomalies were related to wildfires at lead times of four months,
369 presumably because of decreased moisture of surface fuels. Likewise, soil moisture inferred from
370 NASA's Gravity Recovery and Climate Experiment (GRACE) mission was often positively
371 correlated with wildfire occurrence **in herbaceous vegetation, shrublands, and forests at**
372 **seasonal lead times**, indicating that a wetter pre-fire-season can lead to increased **plant (i.e.,**
373 **fuel) production in these landscapes** (Jensen et al. 2018).

374 The large spatial extent of remote sensing datasets provides natural opportunities to
375 explore how soil moisture-wildfire relationships vary across different land cover types. Schaefer
376 and Magi (2019) used satellite-based fire counts from NASA (Giglio et al. 2018), land-use and
377 land cover maps (Hurt et al. 2020), and the ESA CCI SM product (Dorigo et al. 2017) with a
378 biome map (Levvasseur et al. 2012) to study how fires behave relative to soil moisture
379 variability within land cover types and across biomes. They found that the fire-productivity curve
380 shape, which describes resource and climate limits surrounding a zone of optimal fire conditions
381 (Krawchuk and Moritz 2011), was captured within the phase-space of fire and soil moisture. Fire
382 counts were generally greatest when remotely sensed average monthly soil moisture was

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383 relatively low, often around $0.1 \text{ m}^3 \text{ m}^{-3}$. At lower soil moisture levels, the average number of
384 fires decreased with decreasing soil moisture, presumably because of resource limitations (i.e.,
385 decreasing fuel availability). At higher soil moisture levels, the average number of fires
386 decreased with increasing soil moisture, likely due to increased fuel moisture contents. But the
387 shape of the fire-soil moisture curve differed as a function of biomes and land cover types. **For**
388 **example, the occurrence of fires in boreal forests (Fig. 6, lower panel), which have a**
389 **shallower rooting depth than forests in other biomes (Fan et al. 2017), relates to soil**
390 **moisture availability in a way that is similar to grasslands (Fig. 6, upper panel), which also**
391 **have shallow root depths.** This apparent effect of root depth on the sensitivity of fire
392 occurrence to soil moisture under different biomes reinforces the value of soil moisture as a
393 predictor of fire danger. Consistent with these results, Forkel et al. (2017) showed that across the
394 world, biophysical models of fire activity (e.g., Rabin et al. 2015) performed better when
395 remotely sensed soil moisture (and moisture state in general) was considered.

396 These global scale analyses are possible because, unlike in situ soil moisture
397 measurements, remotely sensed measurements provide data on soil moisture conditions across
398 large spatial domains. However, remotely sensed measurements typically represent soil moisture
399 conditions in only the top few centimeters of the soil (Abbaszadeh et al. 2021) and have lower
400 temporal resolution compared to in situ networks. **Furthermore, remotely sensed soil**
401 **moisture measurements have historically shown a limited ability to accurately monitor soil**
402 **moisture conditions where a dense vegetative canopy is present (Djamai et al. 2015; Dorigo**
403 **et al. 2015). But recent advances provide unequivocal evidence that remote sensing**
404 **measurements are sensitive to soil moisture under forest canopies (Ayres et al. 2021;**
405 **Colliander et al. 2020).** There is a clear need to focus future research on remotely sensed soil

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406 moisture-wildfire relationships at higher spatial resolution and for specific land cover types.
407 Such studies may enhance the relevance of satellite-based soil moisture data to fire managers.
408 These types of studies may also be particularly well-suited for linking with model-based
409 approaches as described in the next section.

410

411 **Modeled soil moisture**

412 Given the historical lack of in situ and satellite measurements, proxies and estimates of
413 soil moisture conditions have long been used in the context of wildfire danger. Approaches have
414 ranged from drought indices based on simplified soil water balance models (e.g., Keetch and
415 Byram 1968; Mount 1972; Palmer 1965), to actual soil moisture values simulated using more
416 complex process-based models (Carrega 1991; Holden et al. 2019), to hybrid approaches that
417 incorporate measured soil moisture data into plant growth models (Krueger et al. 2021). These
418 approaches have been applied across widely-varying time horizons, with some showing the
419 possibility to facilitate predictions of soil moisture, and subsequently wildfire, for time frames
420 potentially spanning decades (Chikamoto et al. 2015). The Keetch-Byram Drought Index
421 (KBDI) (Keetch and Byram 1968), in particular, has been used extensively to address the
422 challenges of representing moisture deficits and their influence of wildfire danger. For example,
423 KBDI has been used in the McArthur Mark 5 forest fire danger index (Holgate et al. 2017), the
424 Fosberg fire weather index (Goodrick 2002), and the US National Fire Danger Rating System
425 (Burgan 1988).

426 Developed in the southern United States in the **1960s** to predict moisture deficits in **duff**
427 **and mineral soil layers**, KBDI is a unitless index ranging from 0-800. The KBDI calculation
428 attempts to address important physical processes such as canopy interception of precipitation and

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429 the effects of biomass on rates of soil water loss. However, it has significant limitations. For
430 example, it does not include humidity, wind, or radiation in its estimate of soil water loss. The
431 model also uses climatological average precipitation as a surrogate for both leaf area and canopy
432 interception, based on the assumption that wetter sites support more vegetation. Finally, KBDI
433 does not consider variability in soil properties, instead assuming a water holding capacity of 8
434 inches for all soils. Given these limitations, it is not surprising that in situ and remotely sensed
435 soil moisture are more strongly related to wildfires than KBDI in grasslands in the western US
436 (Rigden et al. 2020) **and grasslands, shrublands, and forests** in Oklahoma (Krueger et al.
437 2017).

438 When compared to in situ soil moisture observations in Australia, KBDI showed a large
439 wet bias relative to measurements in the 0-30 cm and 0-90 cm soil layers, had correlations with
440 measured soil moisture that vary widely across climate zones and were sometimes negative,
441 tended to dry down too slowly after wet periods, and performed more poorly than simulations
442 from a physically-based land surface model (Holgate et al. 2017; Vinodkumar et al. 2017).
443 Although KBDI can be calibrated to represent temporal variations in live fuel moisture at
444 specific sites, it is unable to accurately represent spatial variations in live fuel moisture, and thus
445 is not recommended for use in operational fire management (Ruffault et al. 2018). Replacing
446 drought indices like KBDI with more robust soil moisture models has been noted as a priority for
447 improving fire danger rating in the US (Jolly 2018) and is well underway in Australia
448 (Vinodkumar and Dharssi 2019; Vinodkumar et al. 2021).

449 Process-based models link vegetation growth and functioning with soil properties and
450 climate information and are sometimes referred to as land surface models or soil-vegetation-
451 atmosphere-transfer models (Moran et al. 2004). These models can be particularly useful because

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452 they represent plant physiological processes, allowing vegetation to be modelled accurately over
453 space and time, and thus capture many of the vegetation fuel attributes that are relevant for fire
454 spread models (Landsberg et al. 2003). One example is the TOPOFIRE model, which was
455 recently developed to provide high spatial resolution, daily estimates of soil moisture, fuel
456 moisture, and fire danger data and maps for the conterminous US (Holden et al. 2019) (Figure 7).
457 Another recent example is the modeling system developed by the Australian Bureau of
458 Meteorology based on the Joint UK Land Environment Simulator (JULES) called the JULES-
459 based Australian Soil Moisture Information (JASMIN) system (Vinodkumar and Dharssi 2019).
460 The JASMIN system was specifically designed for application in operational fire prediction and
461 risk management.

462 Such models hold promise, not only for wildfire decision support, but also for revealing a
463 new foundational understanding of soil moisture—wildfire relationships. For example, modeled
464 soil moisture values from the US National Oceanic and Atmospheric Administration Climate
465 Prediction Center have been used with remotely sensed active fire data to understand global
466 patterns in the constraints of **fine** fuel loads and fuel moisture on wildfire occurrence (Krawchuk
467 and Moritz 2011). Fuel moisture content strongly influences fire ignition and spread, and recent
468 simulations from a physics-based model showed the close interaction of soil moisture with the
469 fuel moisture of the litter layer **in shrubland, woodland, and forests** (Zhao et al. 2021).
470 Likewise, soil moisture modeled using TOPOFIRE has been shown to be a better predictor of
471 canopy water content across the western US than is atmospheric vapor pressure deficit (Lyons et
472 al. 2021). In fact, gridded 5-km resolution live fuel moisture estimates in **grasslands,**
473 **shrublands, and forests** have been generated for Australia based on soil moisture values
474 simulated with the JASMIN system (Vinodkumar et al. 2021). These live fuel moisture

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475 predictions used soil moisture as a leading indicator with a 14-d lag period. The 0-35 cm soil
476 layer was determined to be the best layer for live fuel moisture prediction. This is similar to the
477 0-40 cm layer used for in situ soil moisture measurements in several prior wildfire-related studies
478 (Krueger et al. 2015; Krueger et al. 2016; Krueger et al. 2017; Pook and Gill 1993; Sharma et al.
479 2021). There is a clear need for further development and refinement of process-based models
480 specifically designed to capture soil moisture-fuel load-fuel moisture-fire danger relationships
481 and for the application of those models for greater scientific understanding and improved fire
482 danger ratings.

483 One limitation to process-based modeling approaches relative to simple drought indices
484 is the increased complexity of model inputs and sometimes intensive calibration needs.
485 Necessary inputs typically include gridded data sets for climate conditions, soil properties, and
486 vegetation type and condition. Obtaining these input data at the necessary spatial and temporal
487 scale and resolution can be challenging. For example, soil maps are often compiled at broad
488 spatial scales, often do not cross political boundaries, and sometimes use inconsistent
489 nomenclatures (Mulder et al. 2011; Zheng et al. 1996). Some critical soil attributes, for example
490 soil depth and available water capacity, can be hard to derive using traditional soil mapping
491 techniques, making these even more challenging to input into fire danger models. Levi and
492 Bestelmeyer (2018) summarize available spatial soil information datasets for fire modeling in the
493 US and suggest that advances in soil modeling can lead to improved soil property maps and
494 therefore more accurate fire predictions.

495 One promising approach is to use process-based models in an inverse fashion to estimate
496 the necessary soil physical properties. For example, forest leaf area index has been shown to be
497 indicative of soil properties, with increases in leaf area associated with increases in fertility and

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498 available soil water (Waring 1983; White et al. 1997). Therefore, by modeling tree growth and
499 variables such as leaf area index, soil physical properties can be inferred, which can then be input
500 into land surface models. Following this approach with the 3PG model (Landsberg and Waring
501 1997), Coops et al. (2012) predicted available soil water storage capacity across western North
502 America and compared model predictions of leaf area with those observed from satellite
503 observations. They developed soil maps at a much finer scale (1 km) than those previously
504 available over the area. When these updated soil maps were integrated into modeling, the forest
505 model predictions more closely matched the anticipated growth of a key forest species in the area
506 compared to models driven solely from pre-existing soil map information.

507 There appears to be great potential for these types of hybrid approaches that incorporate
508 in situ or satellite soil moisture measurements into process-based models. For example,
509 predictions of grassland fuel loads can be improved by direct insertion of in situ soil moisture
510 observations into a simulation model's soil water balance routine (Krueger et al. 2021), or a soil
511 moisture model can be improved by assimilating satellite-based soil moisture observations, as
512 demonstrated in Bolten et al. (2010). Hybrid approaches have also proven useful when predicting
513 areas of vegetation stress, which may be more prone to wildfires. For example, areas of
514 increased land surface temperature and decreased greenness are likely to be subject to lower
515 vegetation growth and increased stress (Nemani et al. 1996). If prolonged, these stresses can
516 result in increased litter fall, increased non-photosynthetic vegetation, and drier soil, which in
517 turn correspond with increased fuel load. Based on this concept, Mildrexler et al. (2009)
518 developed a global disturbance index using remotely sensed land surface temperature and
519 greenness and demonstrated that this index could identify areas of broad scale vegetation stress.
520 Waring et al. (2011) applied this index over western North America and demonstrated that

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521 increases in the area designated as stressed were positively correlated with the areas of increased
522 simulated soil water stress and wildfire. Waring and Coops (2016) then compared simulated soil
523 moisture with satellite derived area burned (Fig. 8, left). Using a decision tree approach, they
524 identified four seasonal combinations of current and antecedent soil moisture conditions that
525 predicted where forest fires $>1 \text{ km}^2$ occurred with 69% accuracy (Fig. 8, right).

526 These studies add to the growing body of evidence that an accurate accounting of soil
527 moisture status, either by in situ measurements, remote sensing, or modeling, can improve our
528 ability to anticipate when and where wildfires will occur. While soil moisture models can suffer
529 from errors caused by inaccuracies in input data and the model structure, they are appealing
530 because of their capability to incorporate diverse data sources including measured soil moisture
531 and vegetation condition (Fig. 3). Yet as described in the following section, soil moisture
532 information has thus far been largely absent from major fire danger rating systems.

533

534 **Potential for inclusion of soil moisture information into fire danger rating systems**

535 In this section we explore the potential for integration of soil moisture information into
536 fire danger rating systems. We begin with a brief review of some of the leading fire danger
537 rating systems and how they incorporate weather and other information to estimate fuel bed
538 properties, estimates that could potentially be improved by incorporating soil moisture
539 information.

540 *National fire danger rating systems*

541 Fire danger rating systems integrate inputs representing multiple fire danger factors, often
542 via a model, into one or more qualitative or numerical indices of fire danger. Some systems also

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543 model physical characteristics of the fire, such as fire intensity, rate of spread, and flame length.
544 Fire danger rating systems provide assessments of fire danger over broad geographical areas,
545 encompassing up to millions of hectares, and are typically not designed to provide detailed fire
546 danger information at the field scale. Spatially, when calculated over a grid, the fire danger for
547 each grid cell represents the average fire danger at a given time over that cell, assuming
548 homogeneous fuels, weather, and topography within the cell. Such systems are used to provide
549 public warnings, set preparedness levels, provide a good indication of the difficulty of fire
550 suppression over a wide range of conditions, and to help wildfire managers make wise tactical
551 and strategic management decisions (NWCG 2002).

552 While there are a number of fire danger ratings systems across the world, ranging from
553 national to regional to local scales, it is instructive to look at three national systems that have
554 been widely used for many decades, those of Australia, Canada, and the United States. While a
555 new Australian system is becoming operational in 2022 (AFAC 2022), the previous system
556 consisted of two fire danger indices, each with six fire danger categories: the McArthur Forest
557 Fire Danger Index (FFDI) and the McArthur Grassland Fire Danger Index (GFDI) (McArthur
558 1966, 1967; Noble et al. 1980). The FFDI and GFDI each required temperature, relative
559 humidity, wind speed, and rainfall as weather inputs. The FFDI assumed a standard eucalyptus
560 forest in its calculations, while the GFDI assumed a standard grassland. For each, the fuel type
561 and total fuel load (live + dead) were constant. For the FFDI, fuel availability (drought factor)
562 was calculated from soil moisture deficit, time since last rain, and rainfall amount (Matthews
563 2009). The soil moisture deficit used KBDI or the Soil Dryness Index (SDI, Mount 1972). For
564 the GFDI, the degree of curing was also an input, which was typically estimated by ground-
565 based visual observations, satellite imagery, or a combination of both.

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566 **The Canadian Forest Fire Danger Rating System (CFFDRS, Stocks et al. 1989) has**
567 **been in its current form since 1992, with a series of improvements planned for release**
568 **beginning in 2025 (CFSFDG 2021). It consists of two major subsystems, the Fire Weather**
569 **Index (FWI) System (Van Wagner 1987) and the Fire Behavior Prediction (FBP) System**
570 **(FCFDG 1992).** The FWI System uses a standard jack pine forest and consists of six
571 components: three moisture indices (**fine fuel moisture code, duff moisture code, and drought**
572 **code**) that represent three organic layers at or beneath the forest floor (Fig. 2), and three fire
573 danger indices, including FWI itself. Weather inputs are temperature, relative humidity, wind
574 speed, and rainfall. The **FBP System** consists of 16 available fuel models, including a grass
575 model that requires degree of curing as an input (De Groot 1993). The system incorporates three
576 outputs from the FWI System and uses topography in its calculations. Foliar (live) fuel moisture
577 is modeled using elevation, geographical location, and date; thus, the foliar moisture remains the
578 same on a given date and location from year to year. Outputs of the **FBP System** include
579 physical characteristics of the wildfire (e.g., rate of spread and fire intensity).

580 The National Fire Danger Rating System (NFDRS) of the United States was originally
581 released in 1972 (Deeming et al. 1972) with major updates in 1978 (Bradshaw et al. 1983) and
582 1988 (Burgan 1988). The latest version (NFDRS2016) includes five standard fuel models,
583 reduced from 20 in the 1978 and 1988 versions (Jolly 2018). As with the two former versions,
584 NFDRS2016 separately calculates live and dead fuel moisture as well as the dynamic fuel load
585 transfer (i.e., curing or **green-up**) between 1-hour dead (< 6 mm diameter dead fuels) and live
586 herbaceous fuels (Fig. 9). The live fuel moisture and dynamic fuel load transfer calculations in
587 NFDRS2016 are a function of Growing Season Index (GSI), which is a function of temperature,
588 relative humidity, and photoperiod (Jolly et al. 2005). As with the 1988 NFDRS, the new system

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589 uses KBDI as a drought surrogate to linearly increase the dead fuel loads when KBDI increases
590 above a threshold of 100. The inputs to NFDRS2016 are temperature, relative humidity, rainfall,
591 wind speed, solar radiation, and photoperiod (based on latitude and day of year). The outputs
592 from NFDRS consist of four components describing the wildfire danger: Spread Component,
593 Energy Release Component, Burning Index, and Ignition Component.

594 *Potential pathways for inclusion of soil moisture information*

595 Within fire danger rating systems like those described above, there are at least five
596 potential uses for soil moisture information: 1) as a replacement or supplement for drought
597 indices; 2) as an input for live and 3) dead fuel moisture modeling; 4) as an input to estimate
598 curing for herbaceous fuels; and as 5) as an input for estimation of fuel loads. We now briefly
599 discuss each of these potential uses within the context of NFDRS2016 (Fig. 9), which provides a
600 representative example for how soil moisture could potentially be used in fire danger rating
601 systems worldwide.

602 First, soil moisture measurements or simulations from process-based models could be
603 used to replace drought indices in fire danger rating systems. Moisture indices **that can**
604 **represent** soil moisture have been used in fire danger rating systems across the world, including
605 KBDI and SDI in the Australian FFDI system, the drought code in the Canadian FWI System,
606 and KBDI in the US NFDRS system. **A growing body of evidence indicates that new sources**
607 **of soil moisture information are useful for predicting wildfire danger across a variety of**
608 **landscapes including grasslands, shrublands, and temperate and boreal forests, and soil**
609 **moisture information may be more closely related to wildfire danger than traditional**
610 **drought indices (e.g., Ambadan et al. 2020; Bartsch et al. 2009; Chaparro et al. 2016;**
611 **Forkel et al. 2012; Krueger et al. 2015; Rigden et al. 2020; Schaefer and Magi 2019). For**

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612 example, in Oklahoma in situ soil moisture measurements provided an average of ten days
613 earlier warning than KBDI for the largest growing-season wildfires (Krueger et al. 2017). For
614 one of the largest wildfires in that study, the Chester Fire, soil moisture reached dangerously low
615 levels ($FAW \leq 0.2$) more than 3 weeks before the actual fire, while KBDI never reached levels
616 considered dangerous (≥ 600), thus providing no advance warning at all. When FAW is < 0.2 , as
617 it was leading up to the Chester Fire, grassland curing rates of $\sim 13 \text{ g m}^{-2} \text{ d}^{-1}$ can occur (Sharma et
618 al., 2021), which could result in the accumulation of $> 270 \text{ g m}^{-2}$ of dead fuel in 3 weeks, or near
619 100% curing given typical grassland fuel loads for the region (Krueger et al., 2021). Analyses
620 from Australia indicate that SDI is on average more strongly correlated with in situ soil moisture
621 measurements than is KBDI, but like KBDI, SDI exhibits slower dry downs than in situ soil
622 moisture and greater variation in performance across regions than more advanced process-based
623 models (Holgate et al. 2016). These results suggest that soil moisture measurements or
624 simulations from process-based models could effectively supplement or replace drought indices
625 in fire danger rating systems.

626 A second potential use of soil moisture information is for live **fuel moisture modeling in**
627 **fire danger rating systems. Soil moisture has been shown to be a strong predictor of live**
628 **fuel moisture in grasslands, (Sharma et al. 2021), shrublands (Pellizzaro et al. 2007; Qi et**
629 **al. 2012), and forest understory (Bianchi and Defossé 2015). In fact, soil moisture**
630 **observations have shown stronger correlations with live fuel moisture than drought indices**
631 **in some Mediterranean shrub species (Pellizzaro et al. 2007) and stronger than remotely**
632 **sensed vegetation indices in Gambel oak and sagebrush (Qi et al. 2012). In NFDRS2016,**
633 live fuel moisture is estimated using GSI, a simple empirical index for vegetation phenology
634 based on photoperiod, vapor pressure deficit, and air temperature (Jolly et al. 2005 and Fig. 8).

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635 We are not aware of any peer-reviewed evaluations of the accuracy of live fuel moisture
636 estimates based on GSI, **although GSI shows temporal trends similar to live fuel moisture**
637 **content in sagebrush and chamise** (Jolly 2018). Based on the evidence from the literature, we
638 hypothesize that inclusion of soil moisture information as an additional input variable in the GSI
639 calculation would lead to improved live fuel moisture estimates. Alternatively, live fuel moisture
640 could be directly estimated from soil moisture information as has been successfully demonstrated
641 in Australia (Vinodkumar et al. 2021).

642 Third, soil moisture could also be useful for dead fuel moisture estimation in fire danger
643 rating systems. Soil moisture influences near-surface air temperature and humidity (McKinnon et
644 al. 2021), and water movement between the soil and dead surface fuels has been observed **in**
645 **shrubland and eucalyptus forests** (Zhao et al. 2022; Zhao et al. 2021) **and aspen forests**
646 (Samran et al. 1995). NFDRS2016 estimates dead fuel moisture using the Nelson model, which
647 uses temperature, relative humidity, solar radiation, and precipitation as inputs (Nelson Jr 2000).
648 The Nelson model has shown reasonable accuracy in estimating dead fuel moisture with r^2
649 values ranging from 0.51-0.79 (Carlson et al. 2007), but for some landscapes like **conifer forests**
650 there is evidence that dead fuel moisture models incorporating soil moisture information provide
651 better estimates than those that omit soil moisture information (Masinda et al. 2021; Pook and
652 Gill 1993; Rakhmatulina et al. 2021). These studies highlight the potential to improve fire danger
653 rating systems by using soil moisture information for estimating dead fuel moisture, particularly
654 for dead surface fuels at forested sites.

655 Improving representation of the curing of herbaceous fuels is a fourth promising use of
656 soil moisture information. Few studies have directly considered the relationship between soil
657 moisture and curing, but the limited available data suggest a strong relationship between soil

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658 moisture conditions and the rate of curing in grasslands (Sharma et al. 2021). Likewise, positive
659 correlations between the fuel moisture content and soil moisture content in grasslands have been
660 observed during the curing period from the end of the growing season into mid-winter
661 (McGranahan et al. 2016). **The degree of curing in herbaceous fuels can be determined**
662 **through direct measurements, visual estimates, remote sensing, or soil moisture deficit or plant**
663 **phenology models (Duff et al. 2019). For example, in NFDRS2016 the dead herbaceous fuel**
664 **load transfer is estimated as a function of the degree of curing, which is estimated from the**
665 **GSI plant phenology model.** Unpublished data show a negative relationship between GSI and
666 grass curing ($r^2 = 0.41$) for one site in North Dakota, USA (Jolly, 2018), but few, if any,
667 published studies have compared GSI with curing levels measured in situ. Given that soil
668 moisture deficits enhance curing (Wittich 2011), soil moisture information could perhaps be used
669 as an additional input for the estimation of GSI and therefore curing, or the curing rate could be
670 directly estimated from soil moisture observations.

671 A fifth potential use of soil moisture information in fire danger rating systems is for the
672 estimation of fuel loads. Current fire danger rating systems assume a constant fuel load (live +
673 dead) regardless of the differences in weather from one growing season to the next. But loads
674 can vary substantially year-to-year, especially for herbaceous fuels. For example, incorporation
675 of soil moisture observations into a simple, process-based plant growth model can provide
676 improved predictions of grassland productivity and fuel loads (Krueger et al., 2021). Likewise,
677 soil moisture is a significant predictor of live fine fuel loads at guinea grass (*Megathyrsus*
678 *maximus*) dominated sites in Hawaii (Ellsworth et al. 2013). There is also evidence for an
679 important role of soil moisture conditions in regulating the growth rates of shrubland fuels in the
680 Arctic (Ackerman et al. 2017; Martin et al. 2017; Myers-Smith et al. 2015). Accounting for soil

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681 moisture effects on fuel production rates could lead to better approaches to represent dynamic
682 fuel loads and could potentially improve the performance of fire danger rating systems.

683 While research provides evidence for these proposed uses of soil moisture in wildfire danger
684 rating, the supporting studies have been relatively few and often of limited geographic scope.
685 Substantiating research across diverse geographic locations and biomes is essential to support
686 implementation on a large scale. Furthermore, the usefulness of soil moisture information in fire
687 danger rating systems is dependent on the way such information is generated. In situ soil
688 moisture measurements can monitor conditions deep into the soil profile rather than just the top
689 few centimeters. Therefore, these in situ measurements effectively represent root zone
690 conditions, and they can be located in diverse vegetation types (e.g., grasslands, shrublands,
691 forests). A main limitation of in situ measurements is that each measurement typically represents
692 only a small area and may not adequately reflect heterogenous soil moisture conditions in the
693 surrounding landscape. Unlike in situ observations, which are lacking in many regions, satellite
694 remote sensing is available globally and can provide useful large-scale estimates of soil moisture
695 conditions. But remotely sensed soil moisture measurements have limited capacity to monitor
696 conditions below the 5-cm depth and limited accuracy beneath dense forest canopies, and less
697 resolution in time as compared to in situ measurements. In contrast, simulated soil moisture
698 information from process-based models can represent the entire root zone, can be extended to
699 almost any land cover and land use type, and have flexible spatial and temporal resolution. Yet,
700 the accuracy of these simulated values is limited by the availability and quality of the necessary
701 soil, vegetation, and weather input data and by uncertainties in the model structure and
702 parameters. Another limiting factor is the sometimes large computational requirements for
703 running the simulations. The use of soil moisture information in fire danger rating systems may

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704 need to rely on a combination of all three sources to represent the best available information
705 across a range of relevant scales.

706

707 **Challenges and opportunities**

708 We have described a steadily growing body of evidence indicating the need for and
709 potential benefits of using soil moisture for wildfire danger assessments. While this research is
710 promising, many questions remain. First, and perhaps most important, while the current body of
711 research supports a litany of *potential* uses of soil moisture in fire danger rating systems, the
712 practical benefits of these uses remain largely untested and logistical challenges likely remain.
713 Pioneering efforts in the operational use of soil moisture information in fire danger rating
714 systems include the use of in situ soil moisture measurements in OK-FIRE, a weather-based
715 decision support system for wildland fire managers in Oklahoma that produces maps of growing-
716 season wildfire danger, updated every 30 minutes, based on soil moisture (Oklahoma Mesonet
717 2021). These maps supplement similar maps based on KBDI for operational fire management
718 decisions. Similarly, the operational use of remotely sensed soil moisture data is being explored
719 by the Barcelona Expert Centre (BEC), which downscales SMOS soil moisture data to create
720 near-real time fire risk maps (BEC Team 2018) that are currently used by Barcelona Provincial
721 Council to provide wildfire early warning (Chaparro et al. 2016). And in Australia, modeled soil
722 moisture values are being used to generate dynamic nationwide live fuel moisture estimates
723 designed for use in operational fire danger ratings (Vinodkumar and Dharssi 2019; Vinodkumar
724 et al. 2021). Further research specifically aimed at techniques for incorporating soil moisture into
725 wildfire danger systems is critically needed, as well as evaluation of fire danger ratings with and
726 without soil moisture information.

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727 Other important research needs and opportunities abound in this context. Some key
728 research questions include: 1) What representations of soil moisture (absolute values, scaled
729 values like FAW, anomalies, percentiles, etc.) are best suited for wildfire danger assessment? 2)
730 What are the soil depths for which moisture conditions are most strongly related to fuel
731 production rates, fuel moisture, and wildfire occurrence and size? 3) How can the various
732 sources of soil moisture information (in situ, remotely sensed, modeled, or a combination of
733 these) best be leveraged for improving operational fire danger assessments? 4) How can soil
734 moisture information be used to produce accurate dynamic estimates of live and dead fuel loads
735 in fire danger rating systems? 5) How are soil moisture conditions related to and predictive of
736 wildfire occurrence and severity in organic soil layers, where the soil itself is the fuel (Elmes et
737 al. 2018; Prat-Guitart et al. 2016; Reardon et al. 2007; Rein et al. 2008)? 6) How do pre-fire soil
738 moisture conditions influence burn severity, soil heating, and post-fire impacts of both wildfire
739 and prescribed fire across different landscapes? These questions must all be answered in parallel
740 with continued research aimed at refining and expanding in situ, remotely sensed, and modeled
741 soil moisture products.

742 After clearing these scientific hurdles, there remains the further challenge of convincing
743 wildfire professionals of the importance of soil moisture compared with more familiar wildfire
744 danger metrics. For example, the importance of KBDI has been engrained in generations of
745 wildfire professionals, and it benefits from widespread familiarity and is inherently understood.
746 It is critical that soil moisture be distinguished from other drought indicators, or it risks being
747 overlooked as just another drought metric. The challenge for scientists is to formulate soil
748 moisture information into a form that is easily understood and used by fire managers.
749 Acceptance may be gained through a “snowball approach”, where use of soil moisture

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750 encourages acceptance, and greater acceptance encourages greater use. If available soil moisture
751 information had been included in operational fire danger rating systems in the US, would it have
752 resulted in earlier warning of extreme fire danger prior to the Camp Fire in 2018? Would it have
753 helped save the lives of any of the 85 people who died from the fire? We do not know. But we
754 know that we now have sufficient soil moisture information and adequate scientific evidence to
755 begin using that information to improve fire danger rating systems around the world. So, let's
756 begin.

757

758 **Acknowledgements**

759 We thank the NOAA National Integrated Drought Information System for co-sponsoring the
760 symposium ([link](#)) which brought together many of the coauthors and launched the development
761 of this review paper. We also thank Marina Skumanich (NOAA NIDIS) and Stephanie Connolly
762 (USFS) for their contributions as co-organizers for that symposium.

763

764 **Declaration of Funding:** This work was supported in part by the US Department of Interior
765 South Central Climate Adaptation Science Center (grant no. G18AC00278). This research was
766 also supported in part by the USDA National Institute of Food and Agriculture through Hatch
767 Project OKL03123 and by the Division of Agricultural Sciences and Natural Resources at
768 Oklahoma State University.

769

770 **Conflicts of Interest:** The authors identify no conflicts of interest.

771

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Review of Recent Developments and Outstanding Questions**

772 **Data Availability Statement:**

773 This is a review manuscript and contains no original data.

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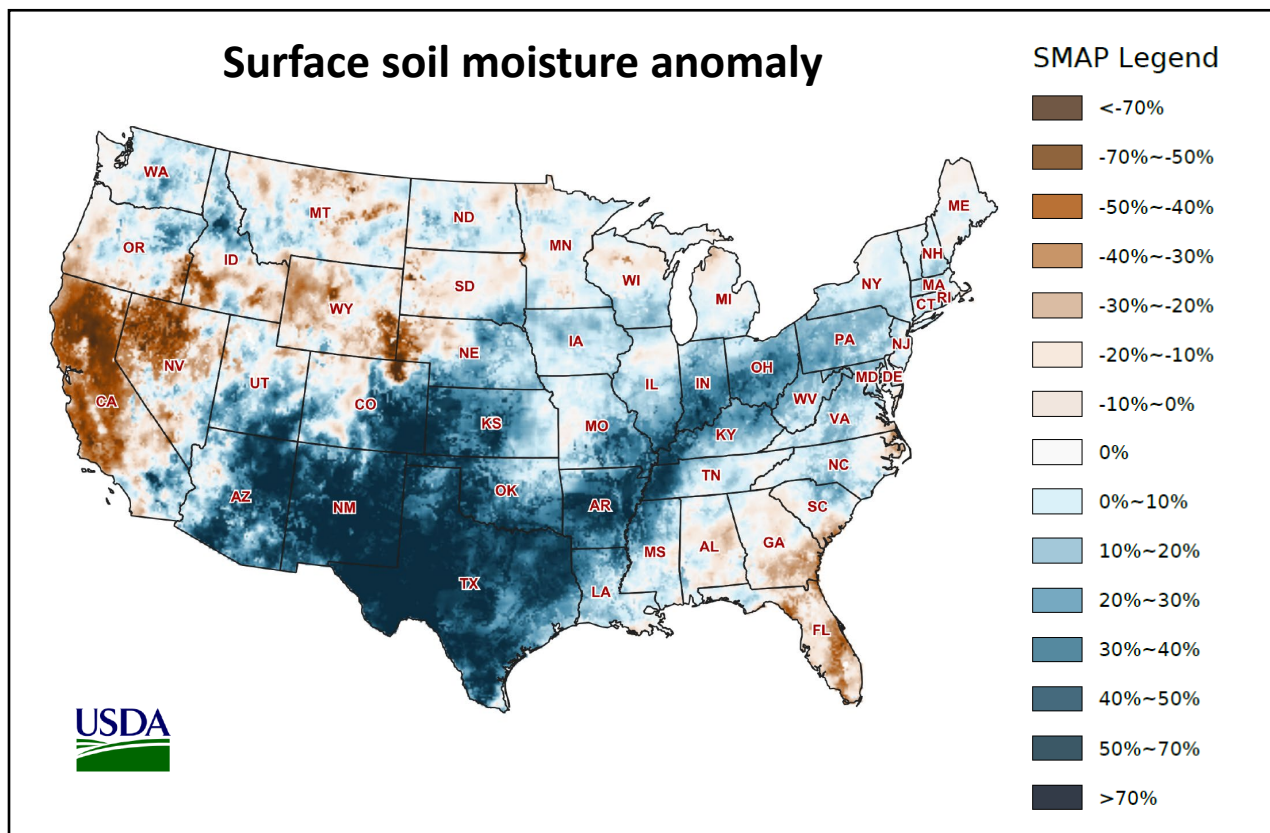
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1178 Figures

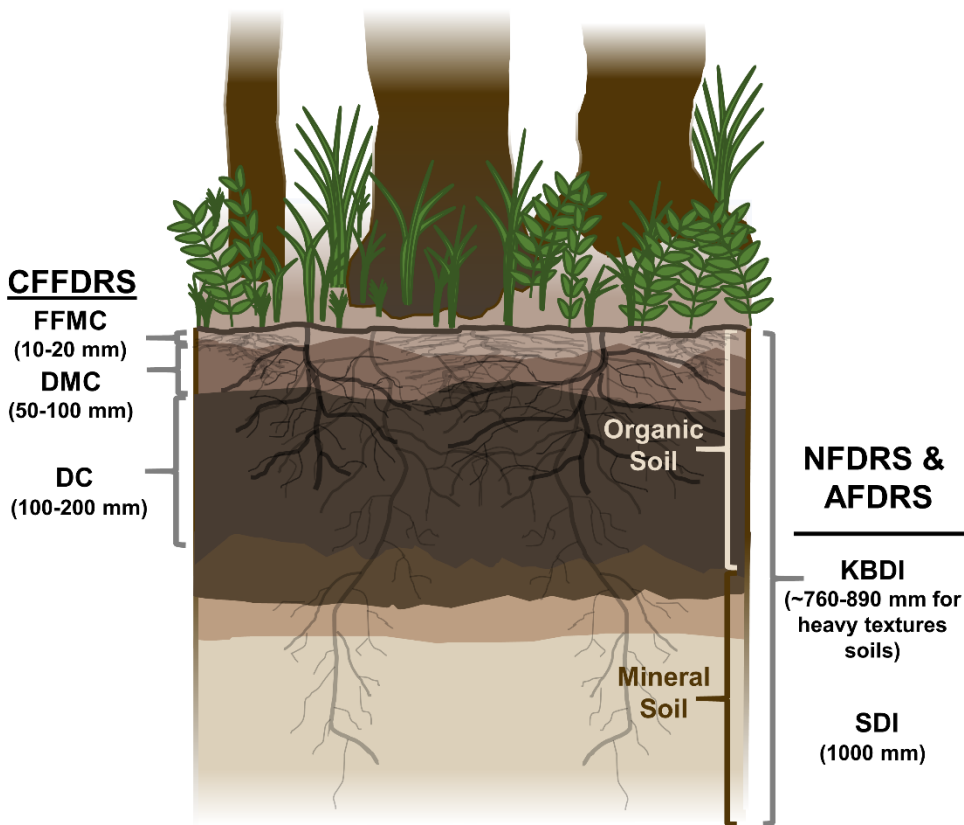
1179 **Fig. 1.** Soil moisture conditions on 1 Nov. 2018, one week prior to the Camp Fire in northern
1180 California, the deadliest and most destructive wildfire in the state's history. The map shows the
1181 surface soil moisture anomaly as reported by NASA's SMAP satellite mission, indicating
1182 exceptionally dry soil conditions conducive to high fire danger in northern California (image:
1183 USDA NASS [Crop Condition and Soil Moisture Analytics](#) system).

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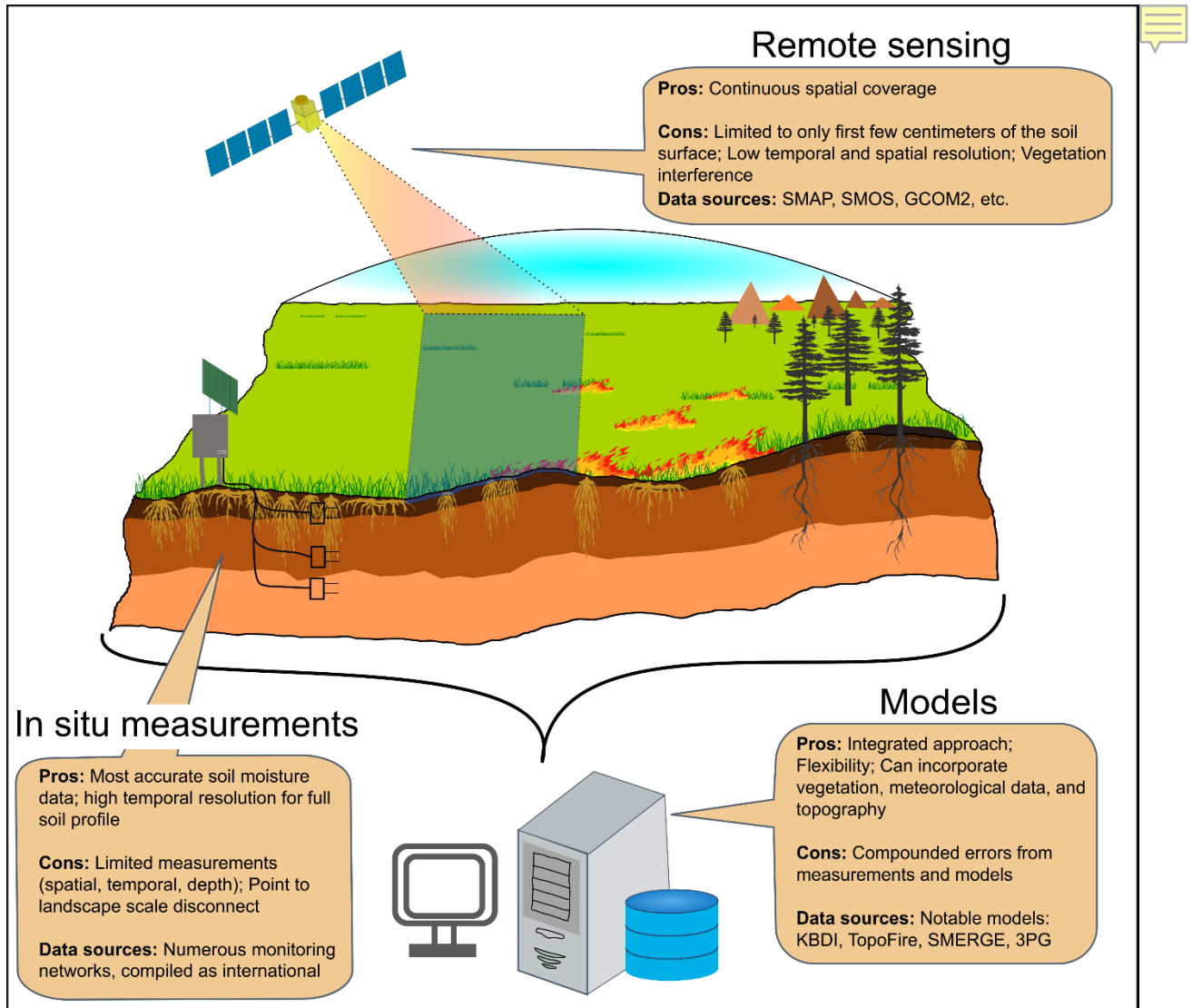
1185 **Fig. 2.** Diagram of a hypothetical forest soil profile with a thick layer of organic soil or duff at
1186 the surface. The diagram approximates the relationships of the soil layers to moisture indices
1187 used in fire danger rating systems in Canada (CFFDRS), the United States (NFDRS), and
1188 Australia (AFDRS). These indices do not use measured soil moisture, account for physical
1189 properties of the soil, or directly account for impacts of overlying vegetation. Instead, moisture
1190 content is calculated using simplistic water balance approaches based on commonly measured
1191 weather variables (e.g., temperature, relative humidity, wind speed, and rainfall).



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1193 **Fig. 3.** A variety of in situ, remotely sensed, and modeled soil moisture data sources have been
1194 ntly developed, with each having unique qualities making them well suited for wildfire
1195 danger modeling.

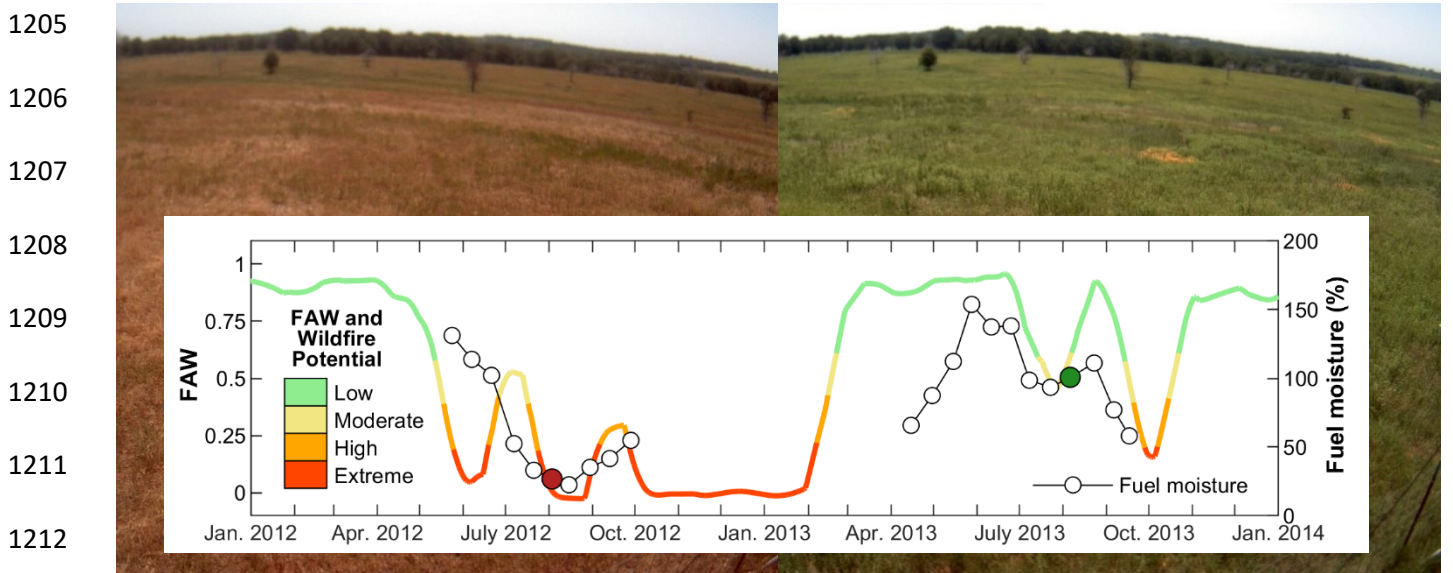


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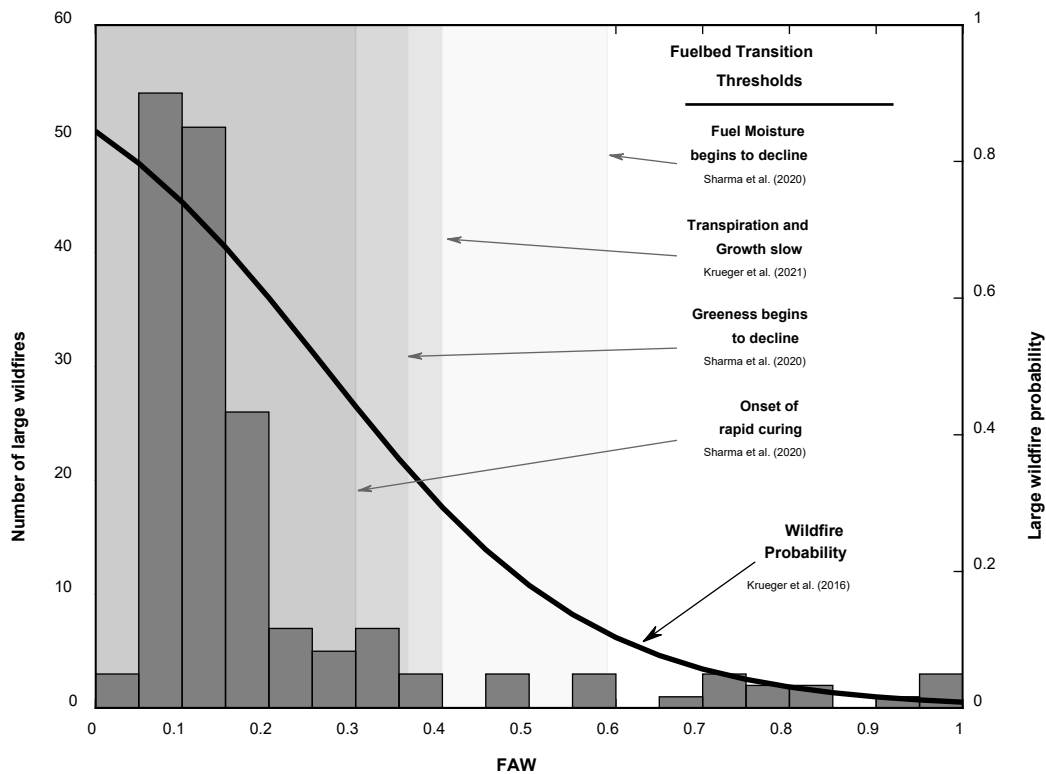
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1198 **Fig. 4.** PhenoCam images collected over grassland near Stillwater, Oklahoma on 2 August 2012
1199 (left) and 6 August 2013 (right), show the influence of soil moisture on vegetation, and by
1200 extension, fire danger. The graph shows the measured fraction of available water capacity
1201 (FAW) at the image location, with colors indicating relative wildfire danger and solid circles and
1202 diamonds representing the mixed (live + dead) fuel moisture and fuel load on days the images
1203 were collected. Photo Credits: University of New Hampshire PhenoCam Network (adapted from
1204 Levi et al., 2019)



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1213 **Fig. 5.** Frequency distribution (histograms) and probabilistic relationship (solid black curve)
1214 between fraction of available water capacity (FAW) and large growing-season wildfires in
1215 Oklahoma from 2000–2012, adapted from Krueger et al. (2015) and Krueger et al. (2016).
1216 Subsequent research provided physical explanations and thresholds for empirical soil moisture-
1217 wildfire links (Krueger et al. 2021; Sharma et al. 2021). These thresholds describe how live
1218 grassland fuels transition to dead fuels as soil moisture declines, beginning with a drop in live
1219 fuel moisture (FAW = 0.59) followed by decreased transpiration and growth (FAW = 0.40).
1220 Next, vegetative greenness declines (FAW = 0.36), which culminates in rapid fuel curing as soil
1221 moisture conditions continue to deteriorate (FAW = 0.30).



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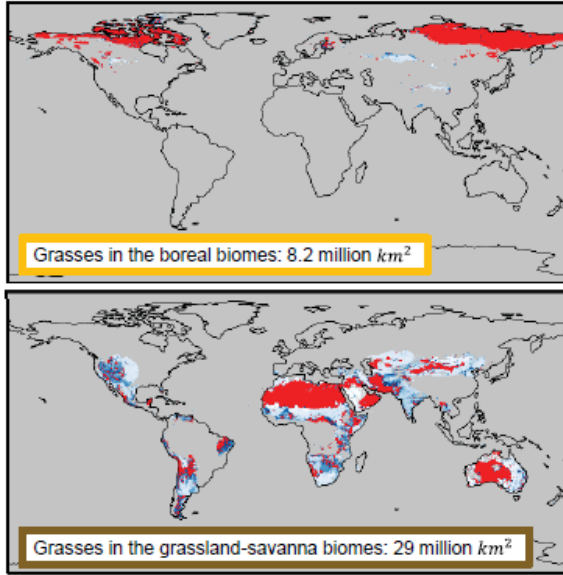
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1224 **Fig. 6.** Fire-soil moisture curves for different land cover types (grasses on the left, and forests on
1225 the right) in different biomes (boreal, grassland-savanna, temperate, and tropical) showing the
1226 resource and climate limits at low and high soil moisture values, respectively. The curves are
1227 derived from monthly-averaged soil moisture for the 0-5 cm soil layer from the European Space
1228 Agency Climate Change Initiative (version 4.2) and monthly fire counts from MODIS Collection
1229 6. The area analyzed in each biome is shown on the maps in red, where this shading denotes
1230 where greater than 75% of the grid cell is a single land cover type. The area covered by grasses
1231 in temperate and tropical biomes (about 5.7 million km², or 13% of global grasses) and area
1232 covered by forests in grassland-savanna biomes (about 4.0 million km², or 11% of global
1233 forests) were excluded due to weaker statistical signal. The shape of the fire-soil moisture curves
1234 varies across land cover types and across biomes, suggesting that soil moisture may be a viable
1235 predictor of biome-scale fire danger for different land cover types (adapted from Schaefer and
1236 Magi, 2019).

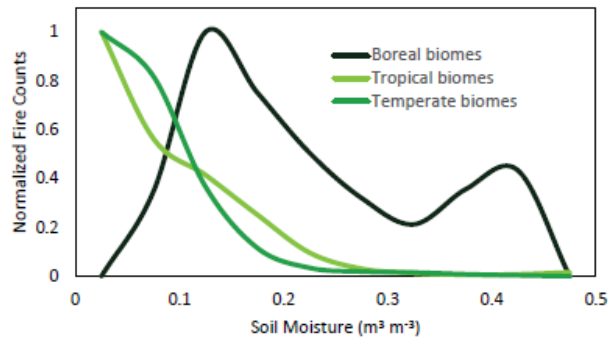
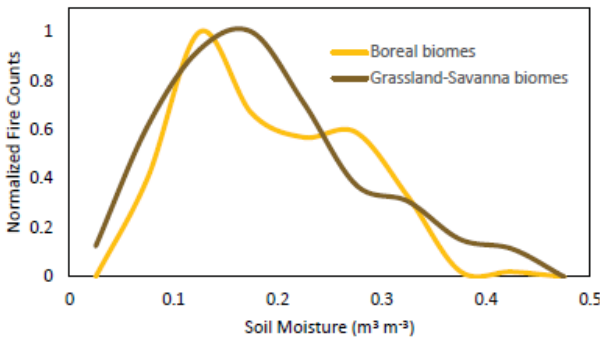
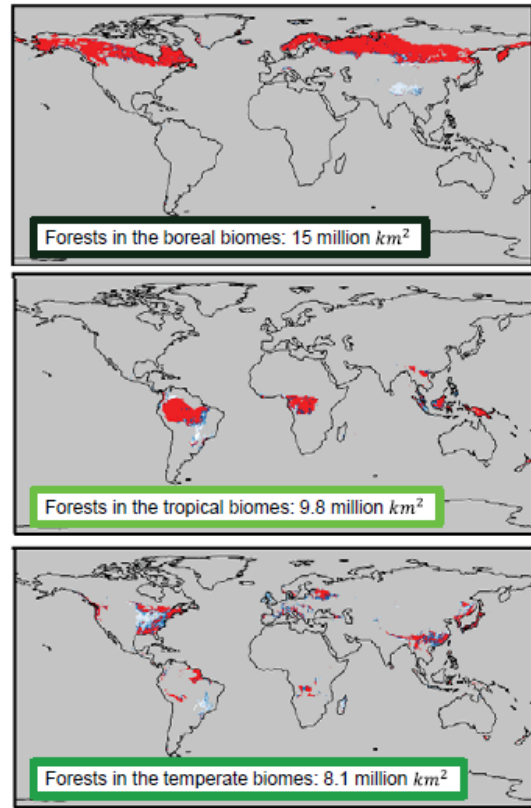
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(a)



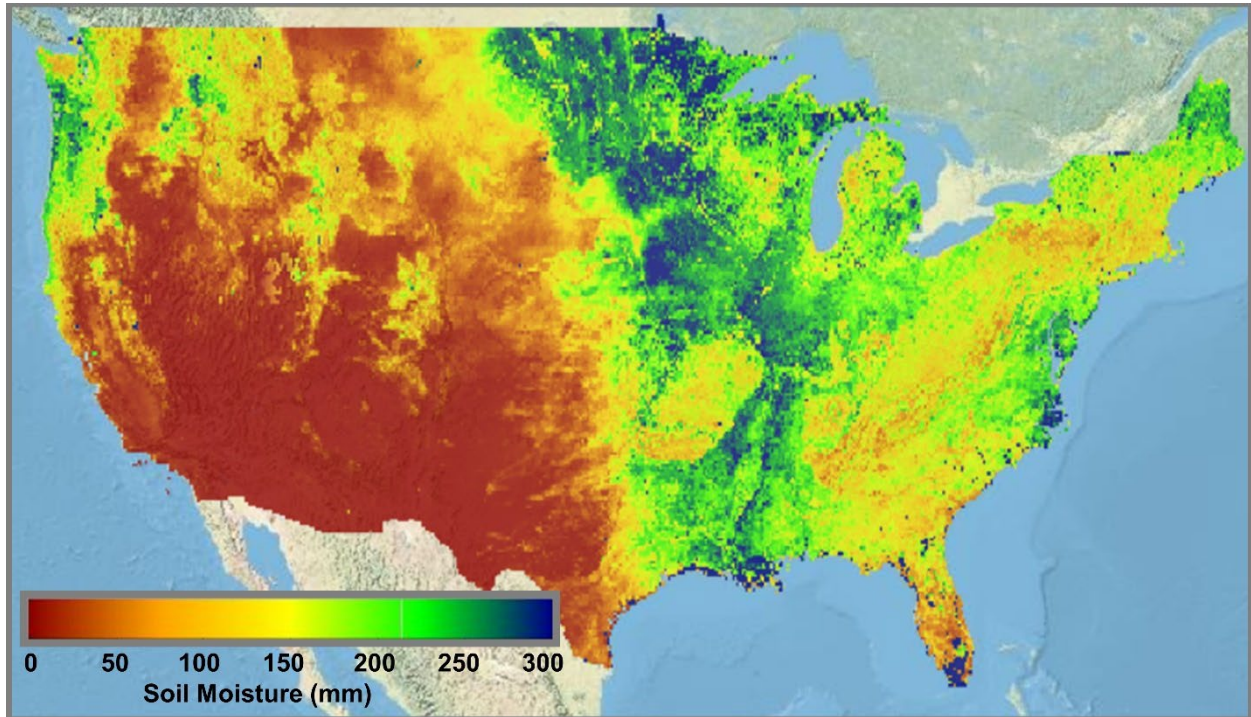
(b)



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1239 **Fig. 7.** Example soil moisture output from TOPOFIRE for the 3 March 2019. Soil moisture maps
1240 for the conterminous United States are produced daily as part of the TOPOFIRE processing
1241 chain (Holden et al., 2019).

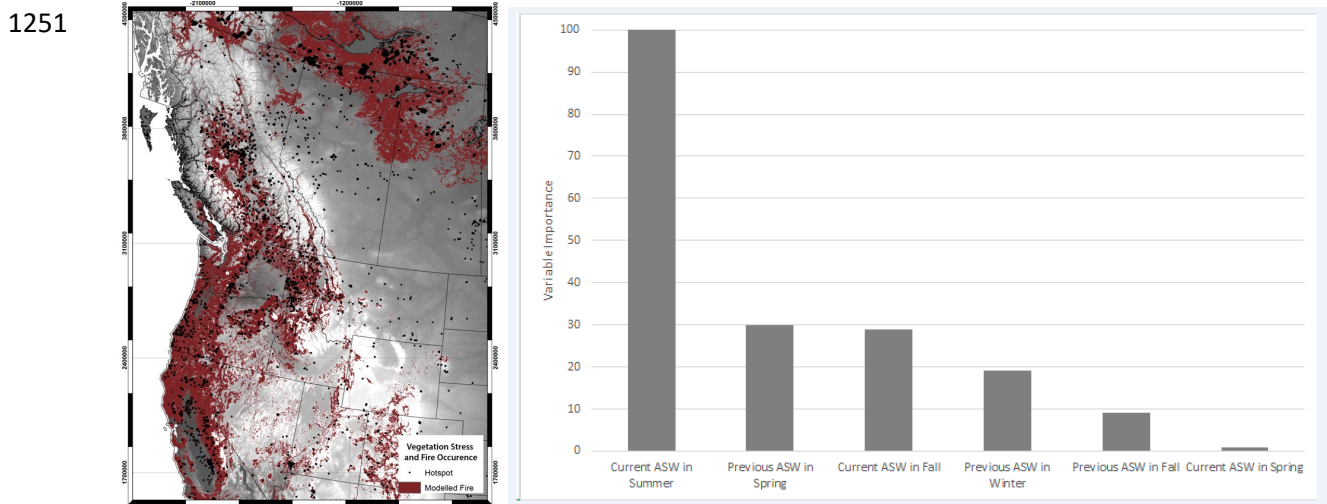
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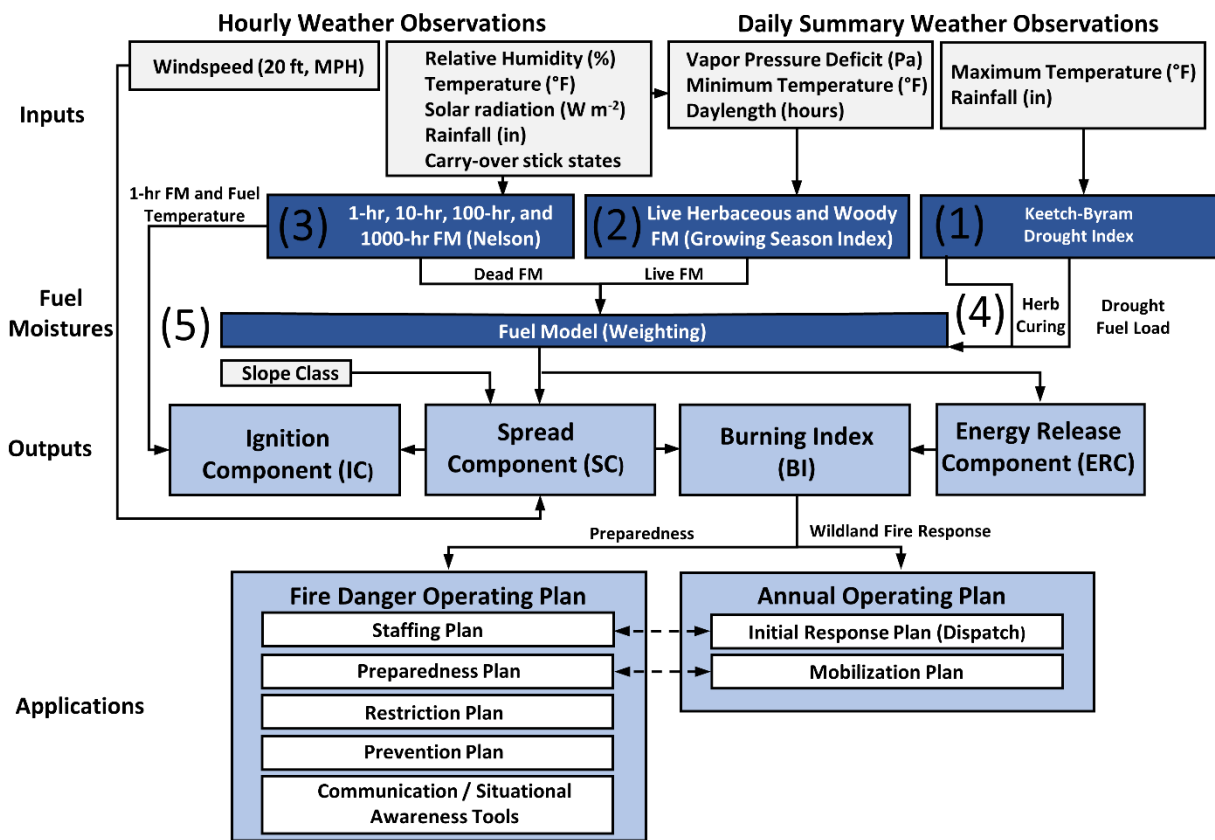
Using Soil Moisture Information to Better Understand and Predict Wildfire Danger: A Review of Recent Developments and Outstanding Questions

1245 **Fig. 8.** Example of model predictions of wildfires (red) based on available soil water (ASW) for
1246 forested portions of western North America in 2004, along with the locations of MODIS active
1247 fire hotspots (black dots) for the same period (left panel). The right panel shows the relative
1248 importance of different seasonal functions of available soil water used to predict MODIS Active
1249 hotspot occurrence of wildfires in 2001, 2004, and 2007 (adapted from Waring and Coops,
1250 2016).



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1252 **Fig. 9.** Structure of the United States National Fire Dander Rating System NFDRS2016 adapted
 1253 from Jolly (2018). Possible uses of soil moisture information in NFDRS2016 are numbered and
 1254 in dark blue boxes, and potential downstream effects of the inclusion of soil moisture
 1255 information are in light blue boxes. These potential uses include (1) supplementing or replacing
 1256 KBDI, (2) live fuel moisture modeling, (3) dead fuel moisture modeling, (4) to estimate
 1257 herbaceous curing, and (5) fuel load modeling.



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