

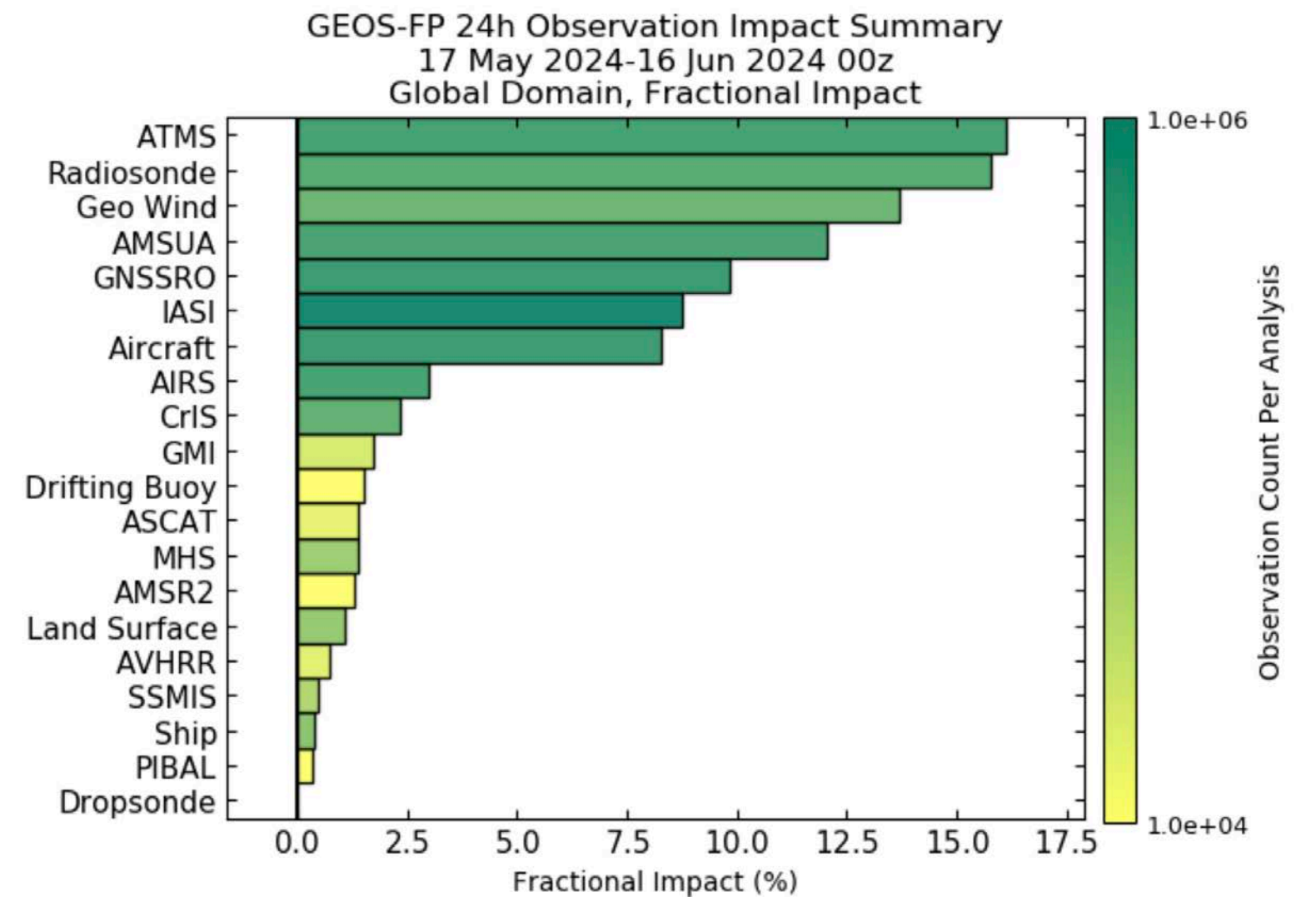
An aerial photograph of a coastline. On the left, a rugged, rocky shore meets the water. The water near the shore is a deep teal color, while the ocean further out is a darker blue with white, turbulent waves. The sky is not visible, as the image is focused on the sea and land.

# The GEOS-ADAS' Preparedness for GeoXo

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- Quick overview of observations used in GEOS-ADAS
- Deep-ish dive into Cloud detection and why we can't have nice things...
- Initial results from IASI using reconstructed radiances from PC scores
- Future Ideas

- Suite of Global Microwave, Infrared, GPS-RO, Conventional, Aircraft, Ship, etc
- Relevant for GeoXo
  - GEO imagers → Only assimilate AMVs
  - GEO imager Clear Sky Radiances → Don't Assimilate Directly Marginal impact
  - IASI – METOP B and METOP C
  - CrIS – NOAA20
    - Don't assimilate NPP or N21
    - NPP may be possible to turn back on with SW/MW cloud det work in-progress (Jones et al. 2024)
    - Several Issues with implementation (a few with unimplemented fixes)
    - Suboptimal CrIS implementation blocking in part turning on N21



Only Real Knob

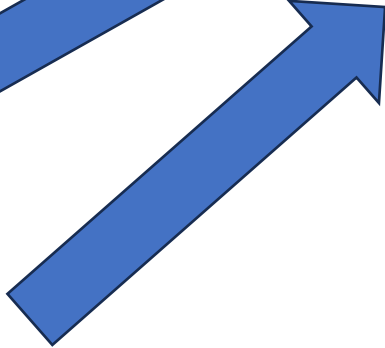
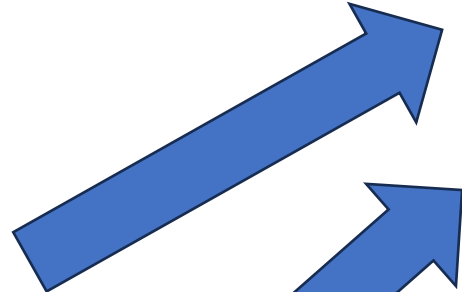
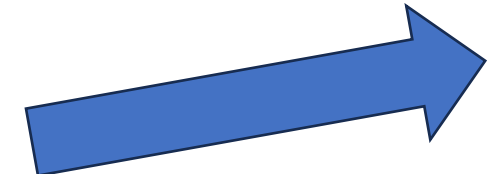
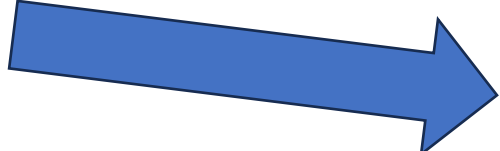
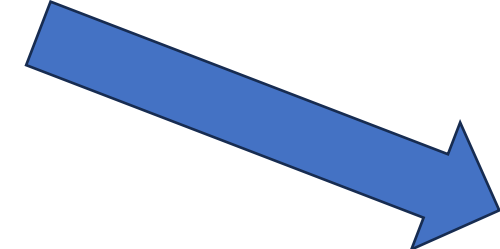
Observed Channel Radiances

Simulated Clear Channel Radiances

Channel Noise Level/Error Variance

Overcast Channel Radiance at Pressure level (From Jacobians)

Tropopause Height



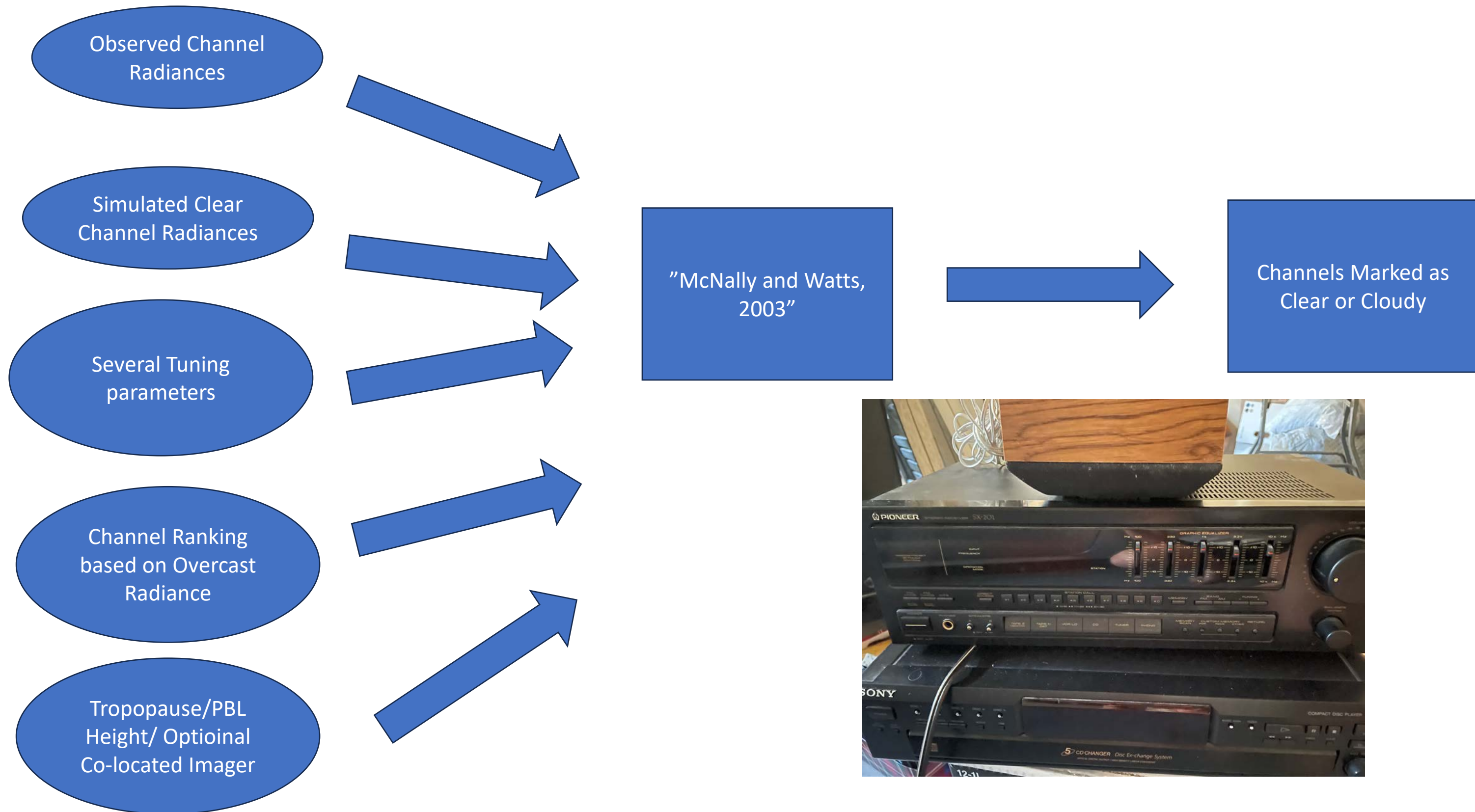
"Eyre and Menzel, 1989"



Channels Marked as Clear or Cloudy

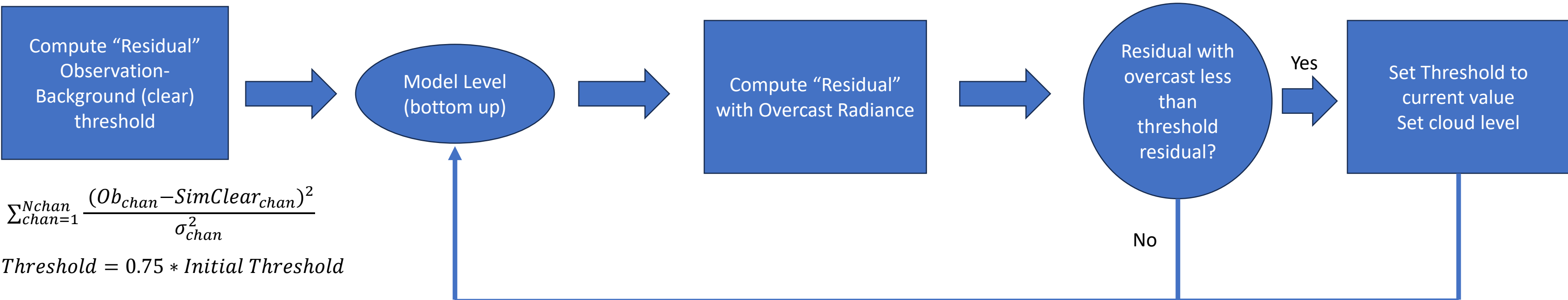
Imager Data Used for Thinning Not in Detection itself (Pick Clearest FOV)





$$N = \frac{\sum_{chan=1}^{Nchan} (Ob_{chan} - SimClear_{chan})(Overcast_{chan}(plevel) - SimClear_{chan})}{\sum_{chan=1}^{Nchan} (Overcast_{chan}(plevel) - SimClear_{chan})^2}$$

$$\sum_{chan=1}^{Nchan} \frac{(Ob_{chan} - SimClear_{chan})^2 - N^2 (Overcast_{chan}(plevel) - SimClear_{chan})^2}{\sigma_{chan}^2}$$



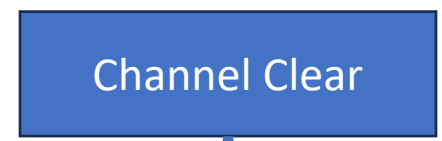
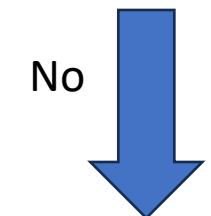
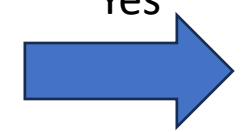
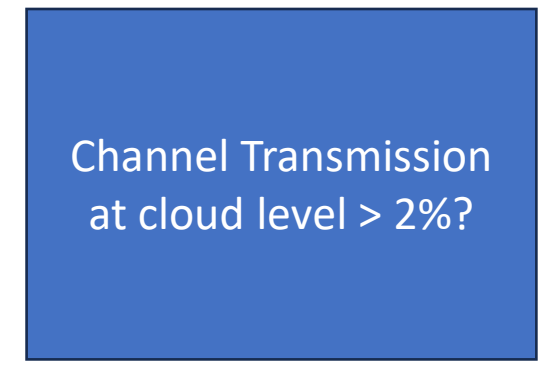
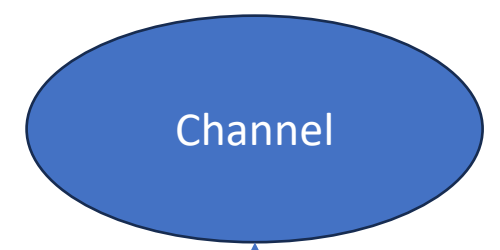
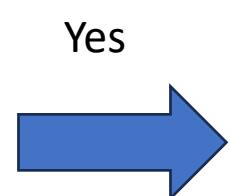
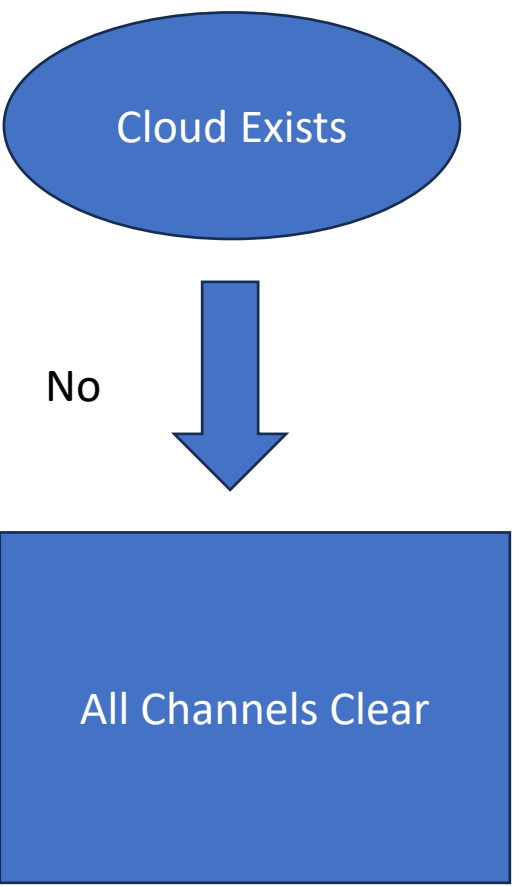
$$\sum_{chan=1}^{Nchan} \frac{(Ob_{chan} - SimClear_{chan})^2}{\sigma_{chan}^2}$$

*Threshold = 0.75 \* Initial Threshold*

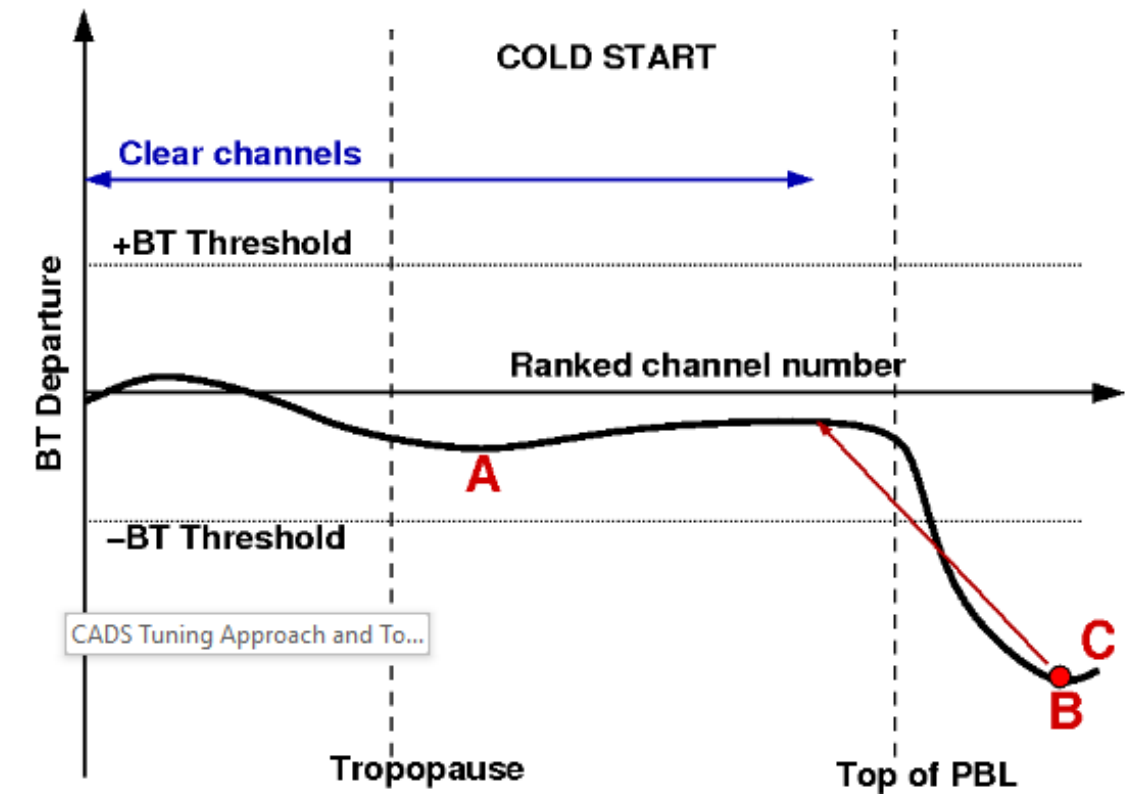
Next level until tropopause

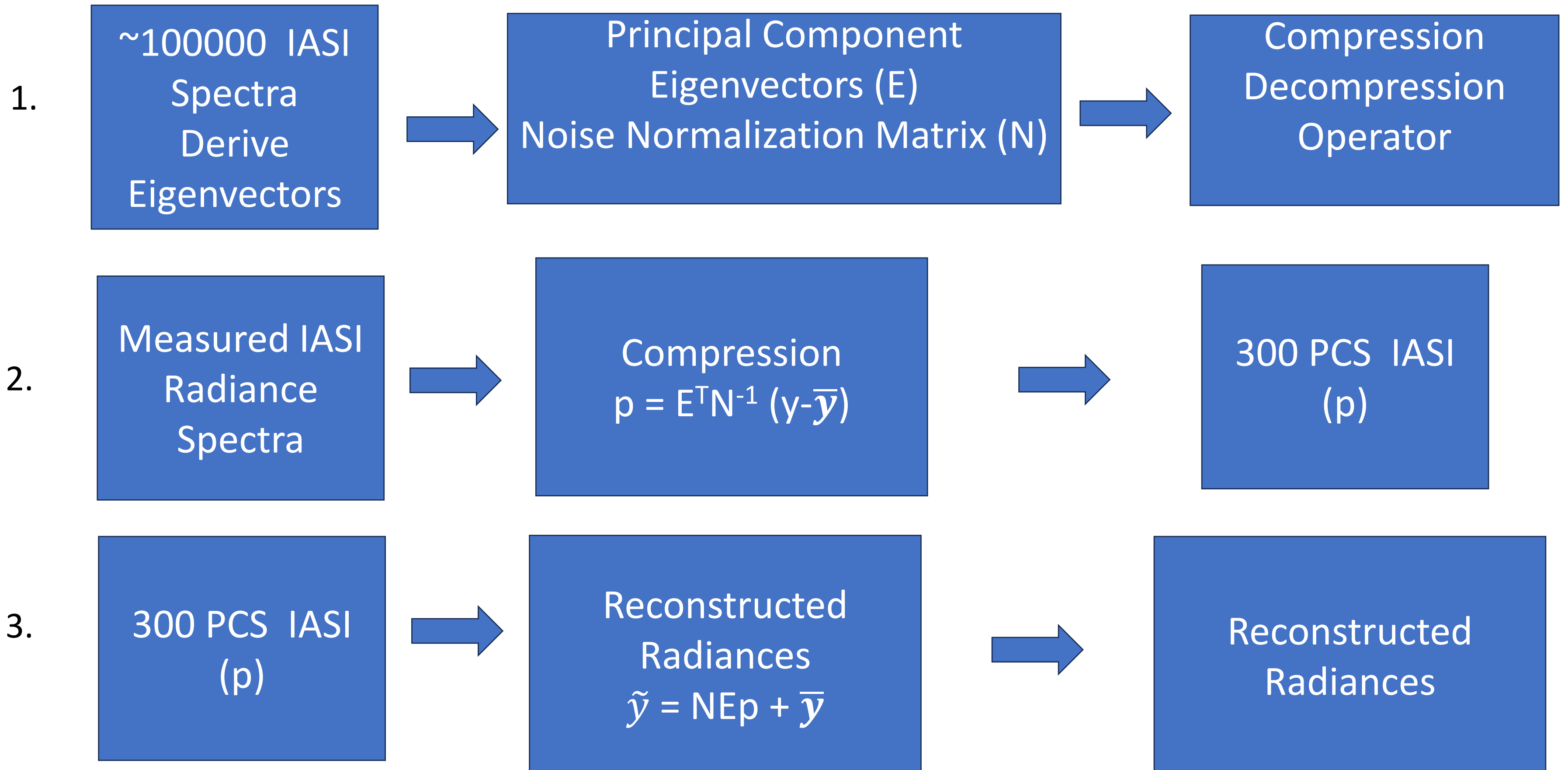
$$\sum_{chan=1}^{Nchan} \frac{((Ob_{chan} - SimClear_{chan}) - N (Overcast_{chan}(plevel) - SimClear_{chan}))^2}{\sigma_{chan}^2}$$

? Typo somewhere

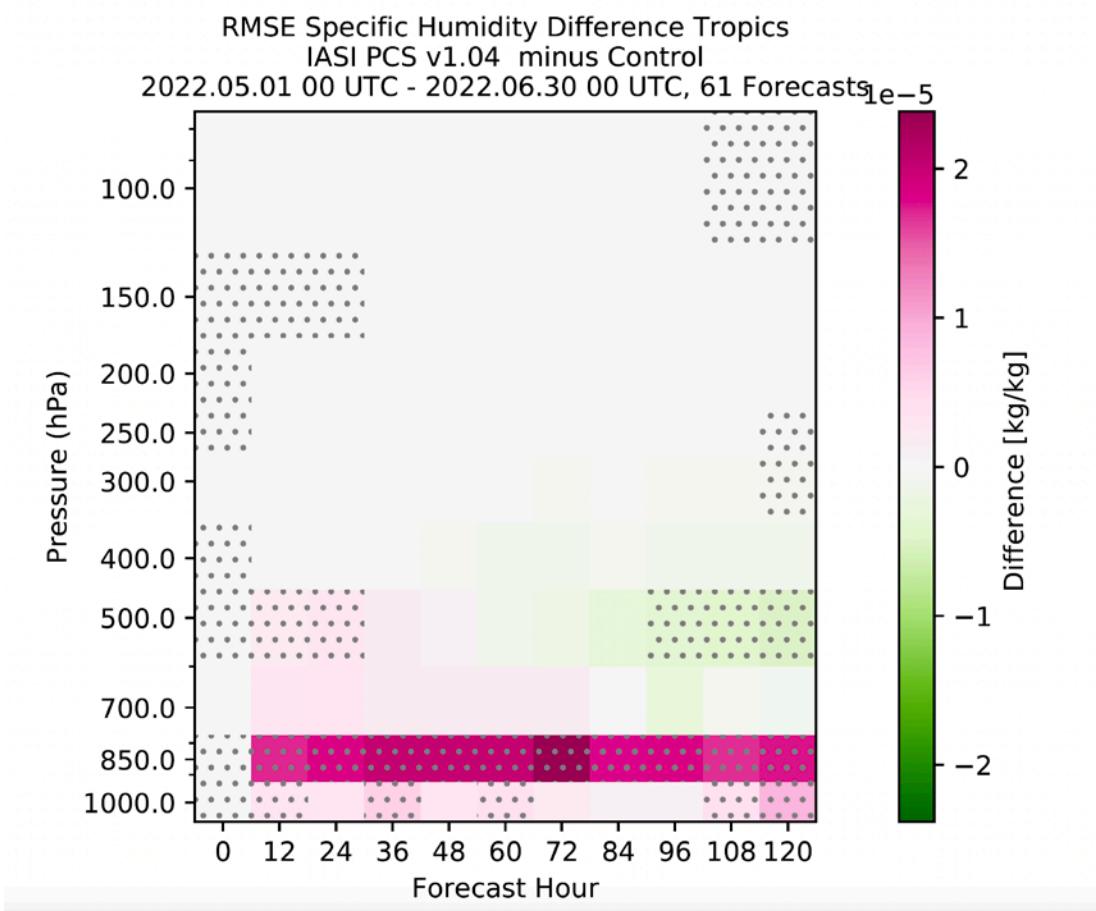


- Essentially based on McNally and Watts, 2003
- Cloud signal is assumed to be a smooth signal between vertically ranked channels
- The algorithm smooths the O-B of vertically ranked channels, and detects a cloud based off departures/gradients in sounding channels and gradients in window channels
- If a cloud is detected in one channel, all channels ranked below it vertically are flagged as cloudy
- Three scenarios warm start, cold start, quick exit
- Several knobs, smoothing window, gradient threshold, absolute threshold
- Where available co-located imager data incorporated into the process
- Minimal use of Overcast Radiance, only used to determine channel rank/height





- Make no modifications to the DAS, only difference is replacing the data source to use reconstructed radiances
- Generally Neutral Impact except for 850 hPa
- Why?
  - Bias correction?
  - Cloud Detection?
  - Combination of both.



### x48\_pci GEOS Scorecard

Comparison of scores for x48\_cln (Control) and x48\_pci (Experiment) experiments for the period of May 14, 2022 to June 29, 2022.

### Legend

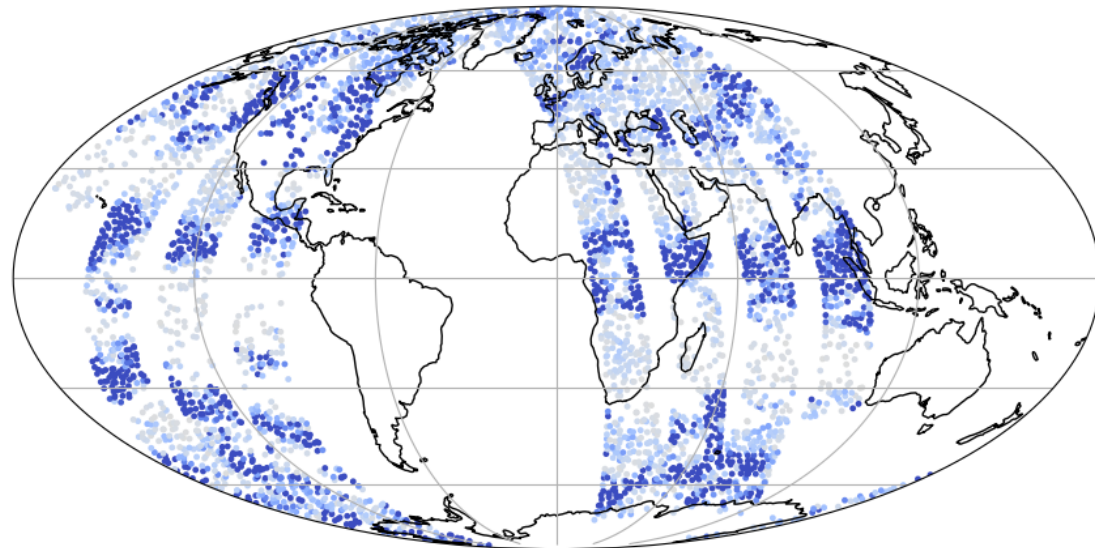
- ▲ far better, significant (99.99% confidence)
- △ better, significant (99% confidence)
- ⊞ slightly better, significant (95% confidence)
- no significant difference
- ⊞ slightly worse, significant (95% confidence)
- ▽ worse, significant (99% confidence)
- ▼ far worse, significant (99.99% confidence)

Northern Hemisphere					Southern Hemisphere					Tropics						
Variable	Pressure Level	COR					RMS									
	Forecast Day	1	2	3	4	5	1	2	3	4	5	1	2	3	4	5
Geopotential Height	10						▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
	70		▼									▼	▼	▼	▼	▼
	100											▲	▲	▲	▲	▲
	250											▲	▲	▲	▲	▲
	500											▲	▲	▲	▲	▲
	700											▲	▲	▲	▲	▲
SLP	850		▼									▼	▼	▼	▼	▼
	1000						▼	▼				▼	▼			
	10		⊞	⊞								▲	▲	▲	▲	▲
	70											▲	▲	▲	▲	▲
	100											▲	▲	▲	▲	▲
	250											▲	▲	▲	▲	▲
Specific Humidity	500											⊞	⊞	⊞	⊞	⊞
	700											⊞	⊞	⊞	⊞	⊞
	850		▼	▼			▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
	10											▲	▲	▲	▲	▲
	70											▲	▲	▲	▲	▲
	100											▲	▲	▲	▲	▲
Temperature	250											⊞	⊞	⊞	⊞	⊞
	500											⊞	⊞	⊞	⊞	⊞
	700											⊞	⊞	⊞	⊞	⊞
	850		▼	▼			▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
	10											▲	▲	▲	▲	▲
	70											▲	▲	▲	▲	▲
U-Wind	100											⊞	⊞	⊞	⊞	⊞
	250											⊞	⊞	⊞	⊞	⊞
	500											⊞	⊞	⊞	⊞	⊞
	700											⊞	⊞	⊞	⊞	⊞
	850											⊞	⊞	⊞	⊞	⊞
	10											▲	▲	▲	▲	▲
V-Wind	70											⊞	⊞	⊞	⊞	⊞
	100											⊞	⊞	⊞	⊞	⊞
	250											⊞	⊞	⊞	⊞	⊞
	500											⊞	⊞	⊞	⊞	⊞
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	850											⊞	⊞	⊞	⊞	⊞

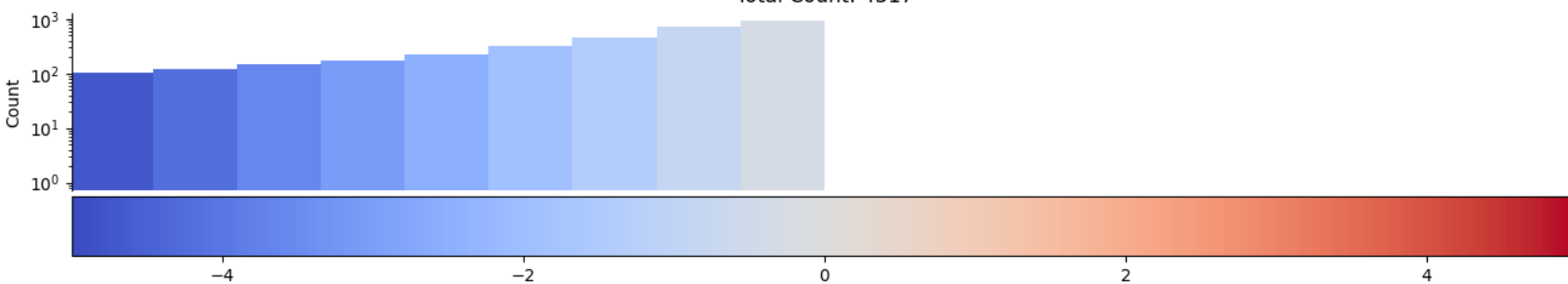
- GSI minimum residual method relies on normalization of by the expected “noise”
- PC compression effectively “drops” noise making it less noisy
- Using the original noise estimate results in cutting off cold end of distribution on surface channels, causing the bias to shift positive, results in a feedback
- Similar phenomenon exists for CrIS after inclusion of Correlated observation error → surface channel drift (also arbitrary 0.75 scaling of threshold questionable)

IASI OmB 426

Mean= -4.7393 Standard Deviation= 7.6323 Min= -62.4337 Max= -0.0012



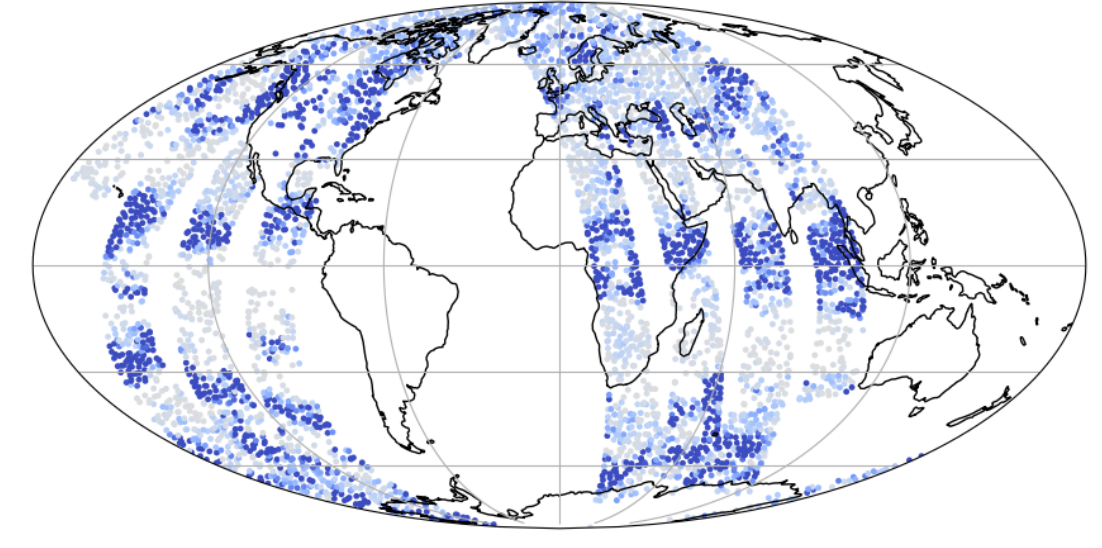
Total Count: 4317



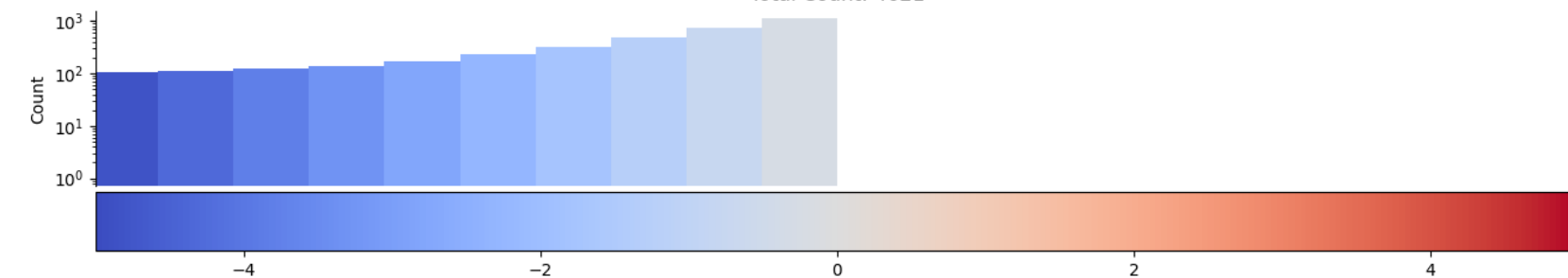
Cold Cloud Flagged Pixels Control

IASI OmB 426

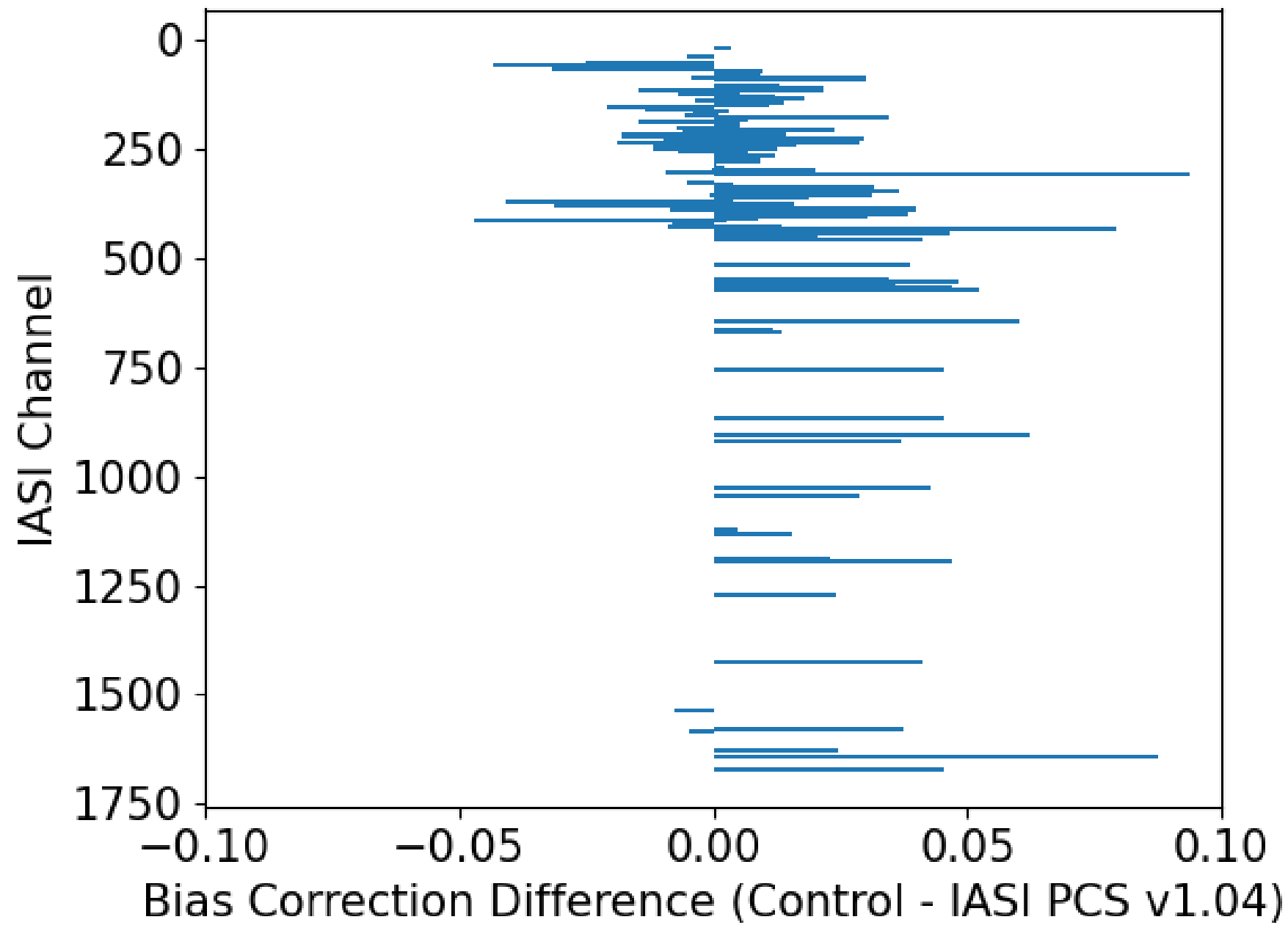
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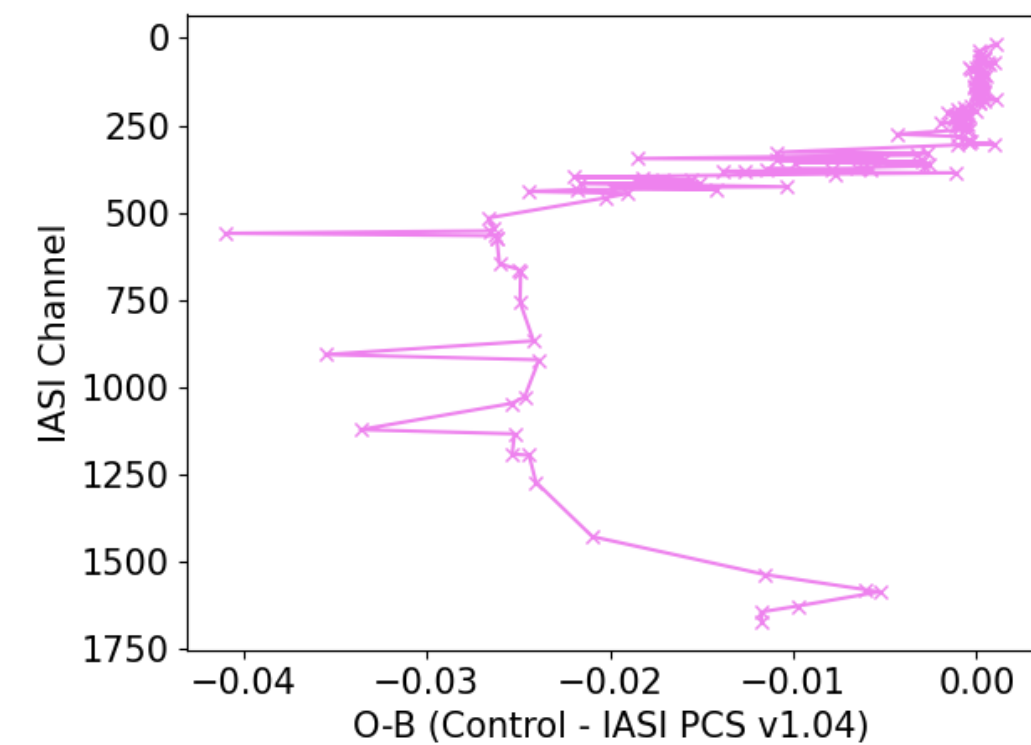
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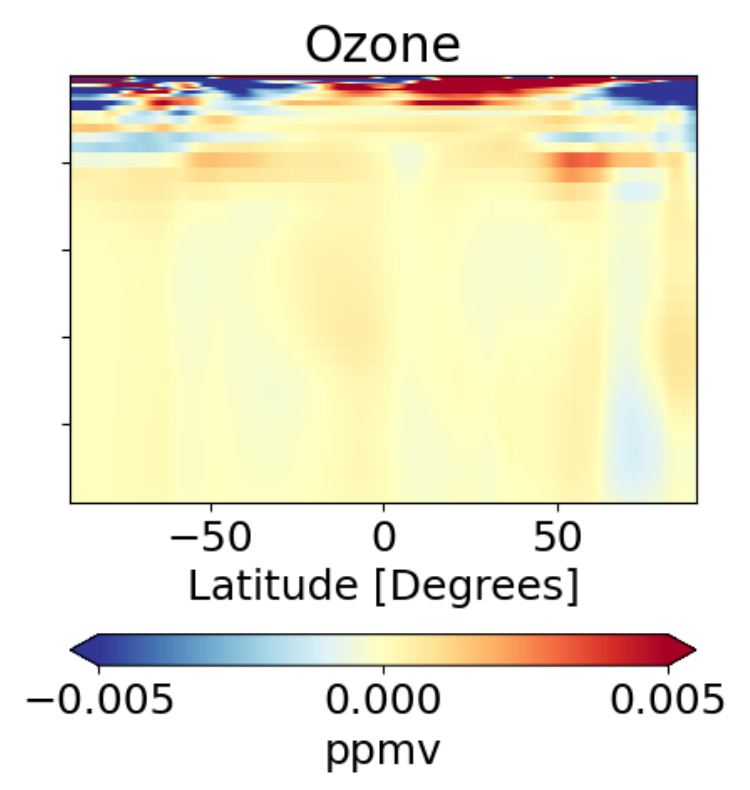
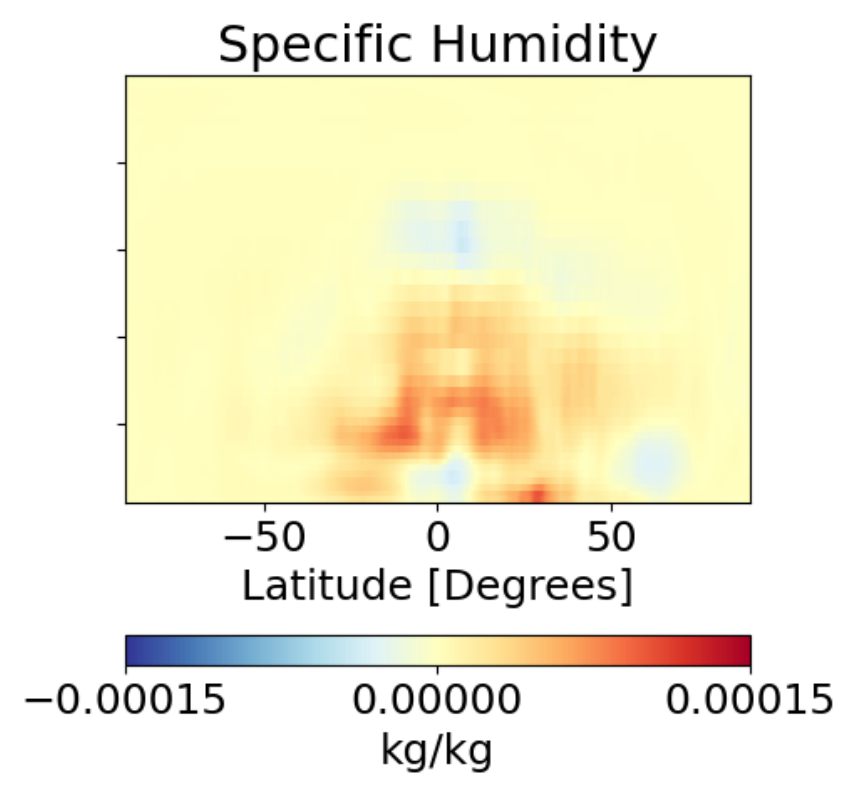
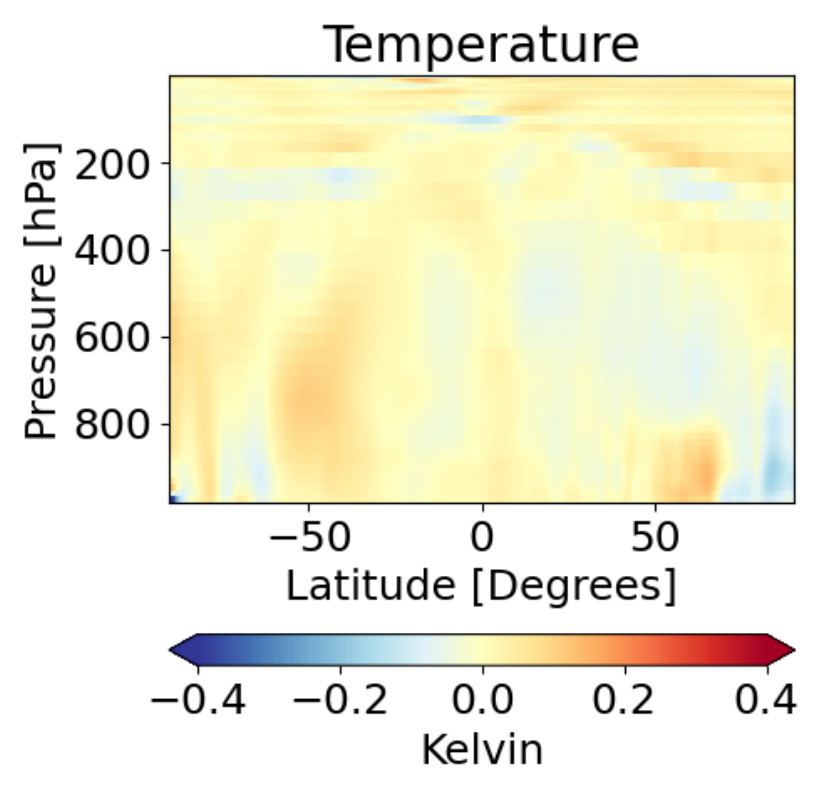
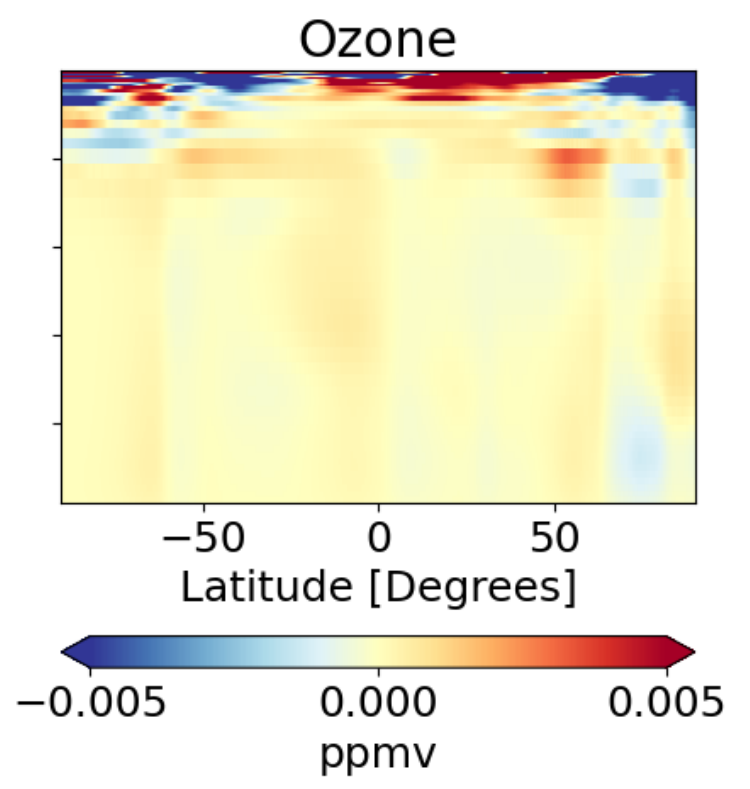
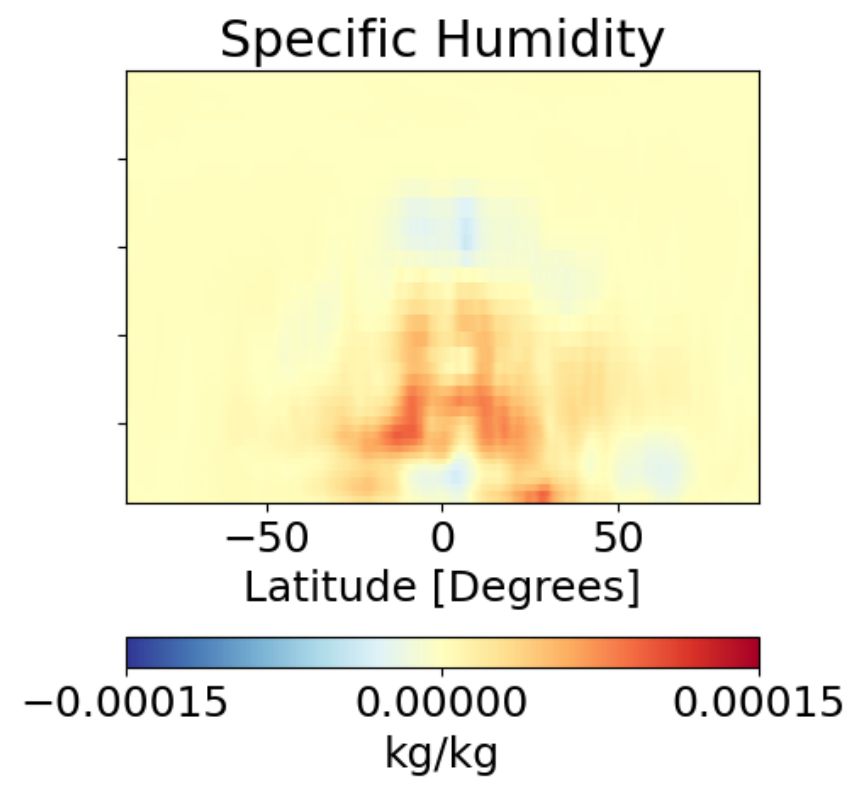
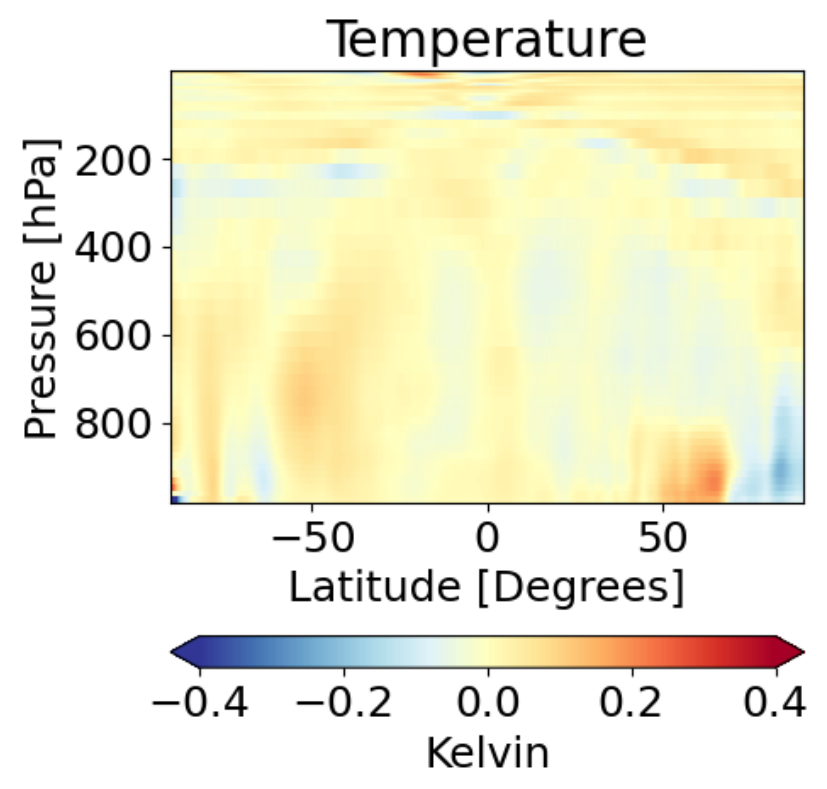


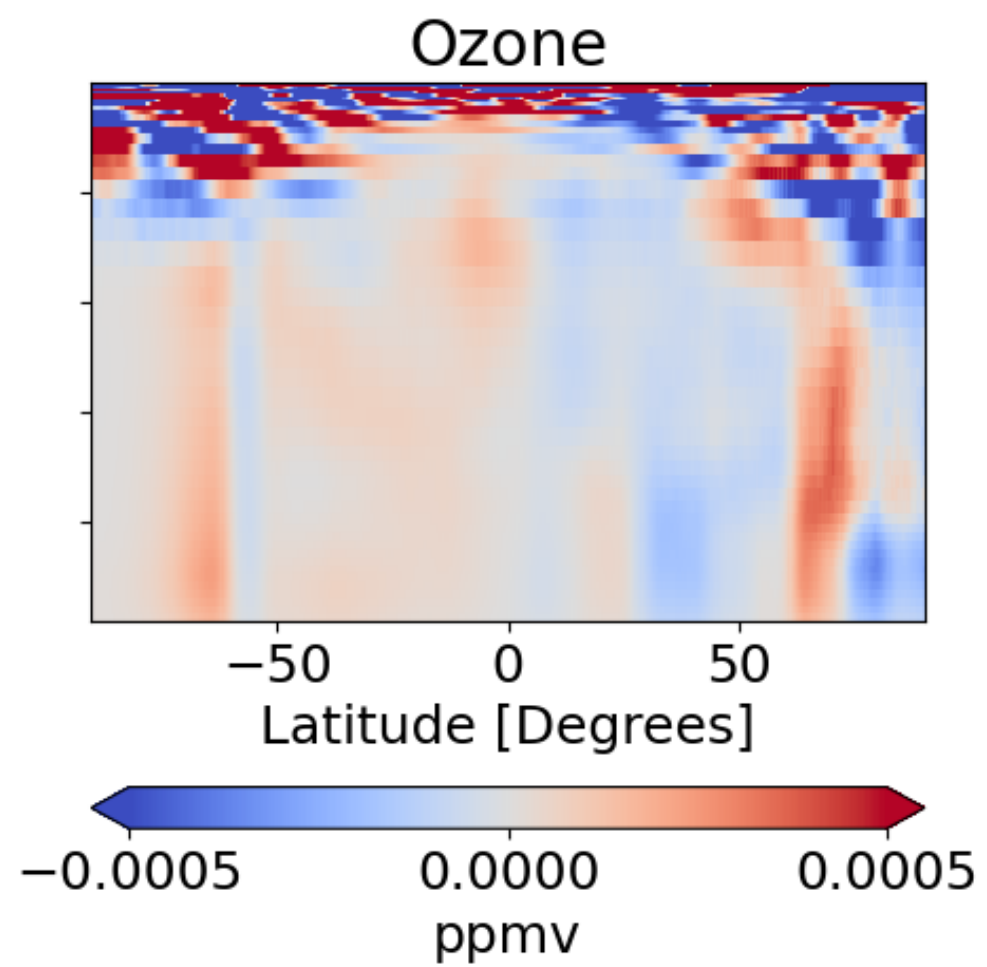
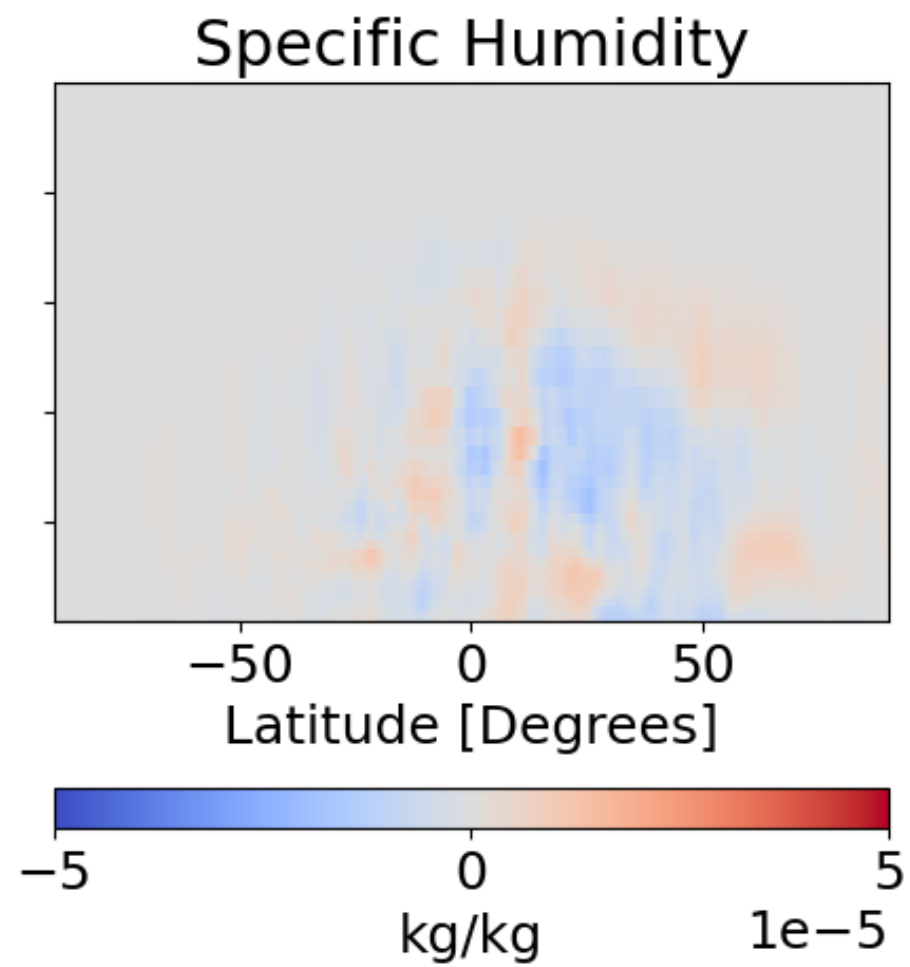
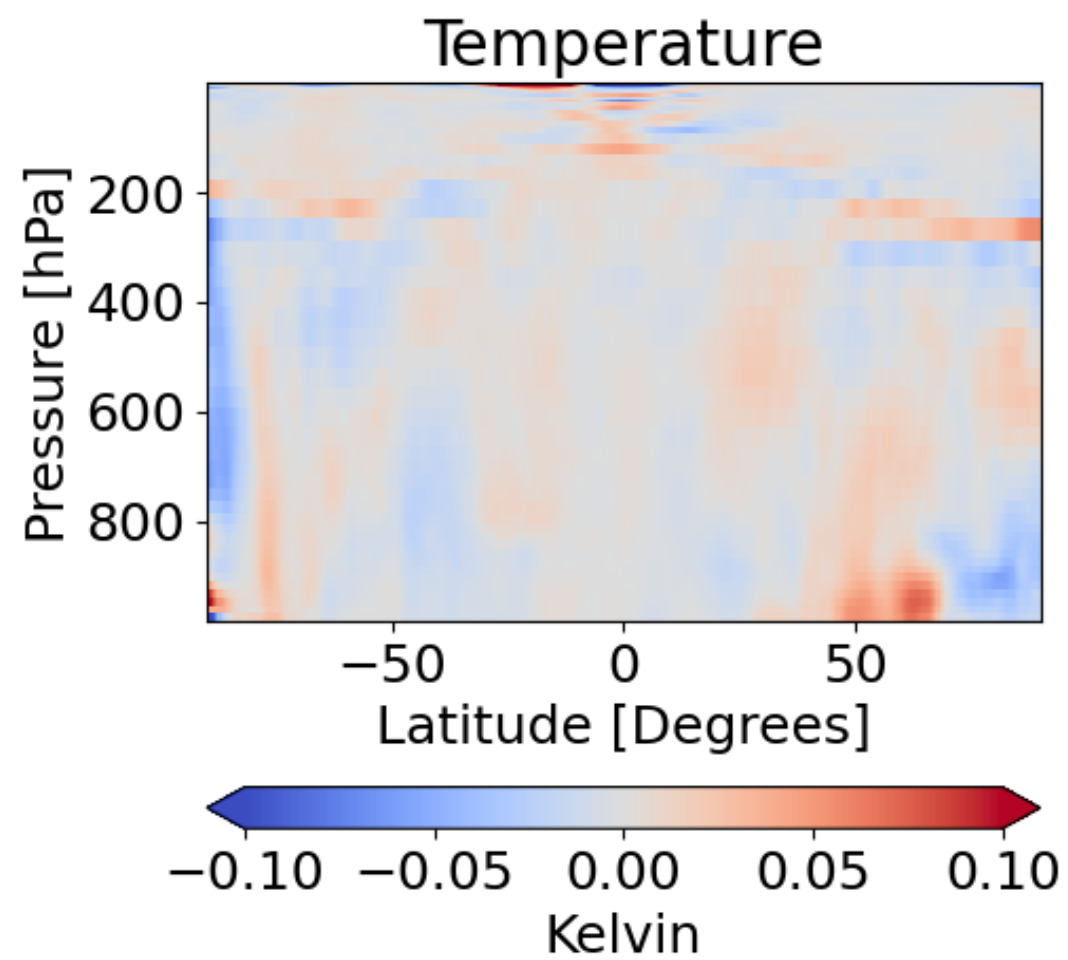
Cold Cloud Flagged Pixels Experiment

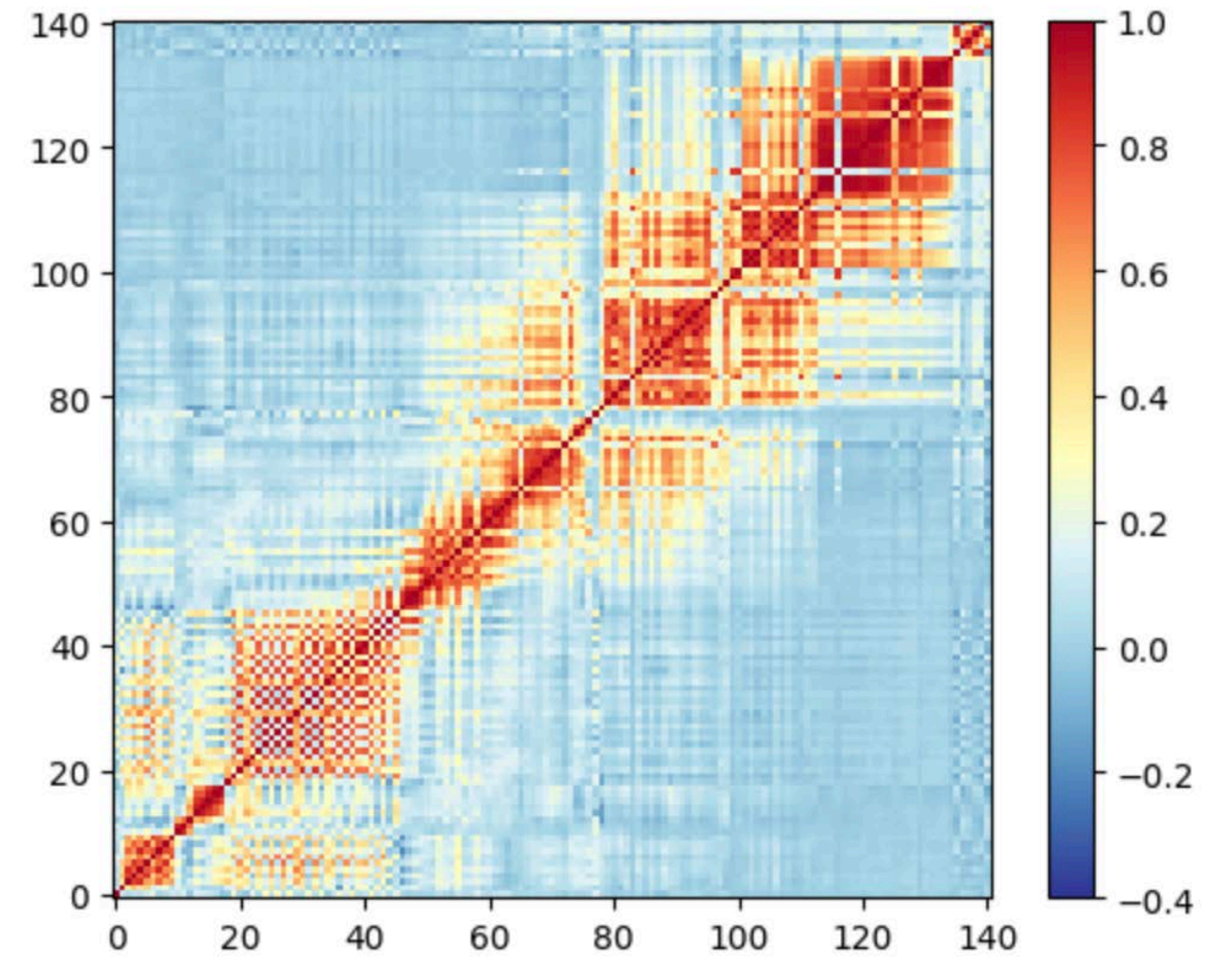
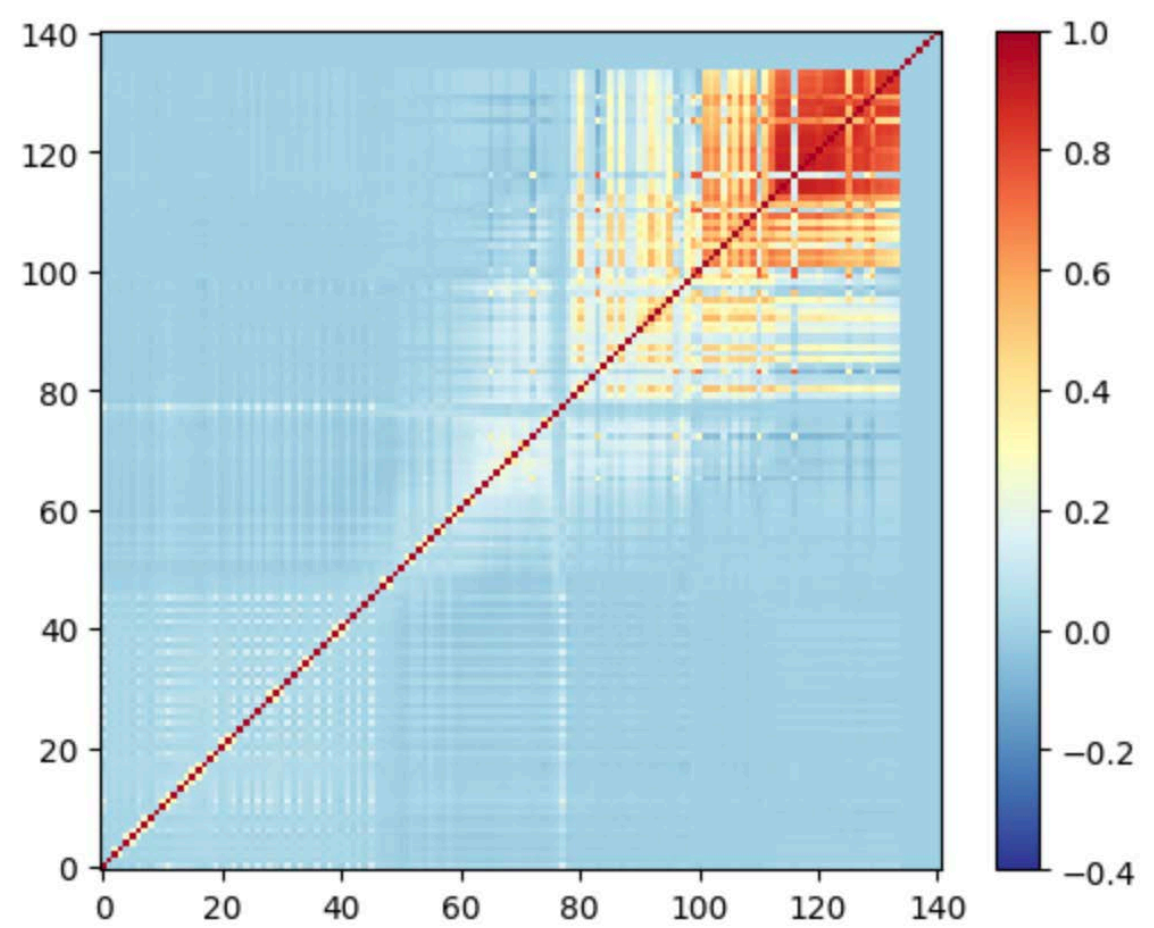
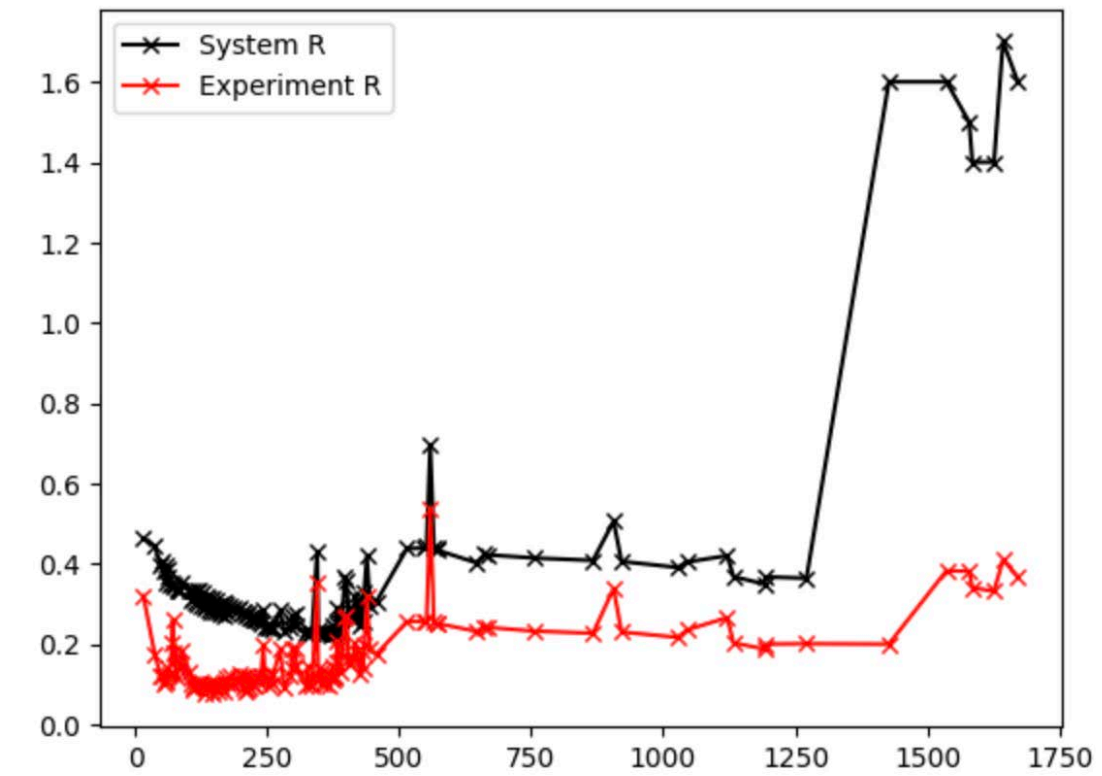


- The over enthusiastic cloud detection causes the bias correction to shift the distributions positive for more surface sensitive channels
- Most likely what is causing the undesirable features at 850 hPa forecast error
- Also average OMB difference for surface sensitive channels

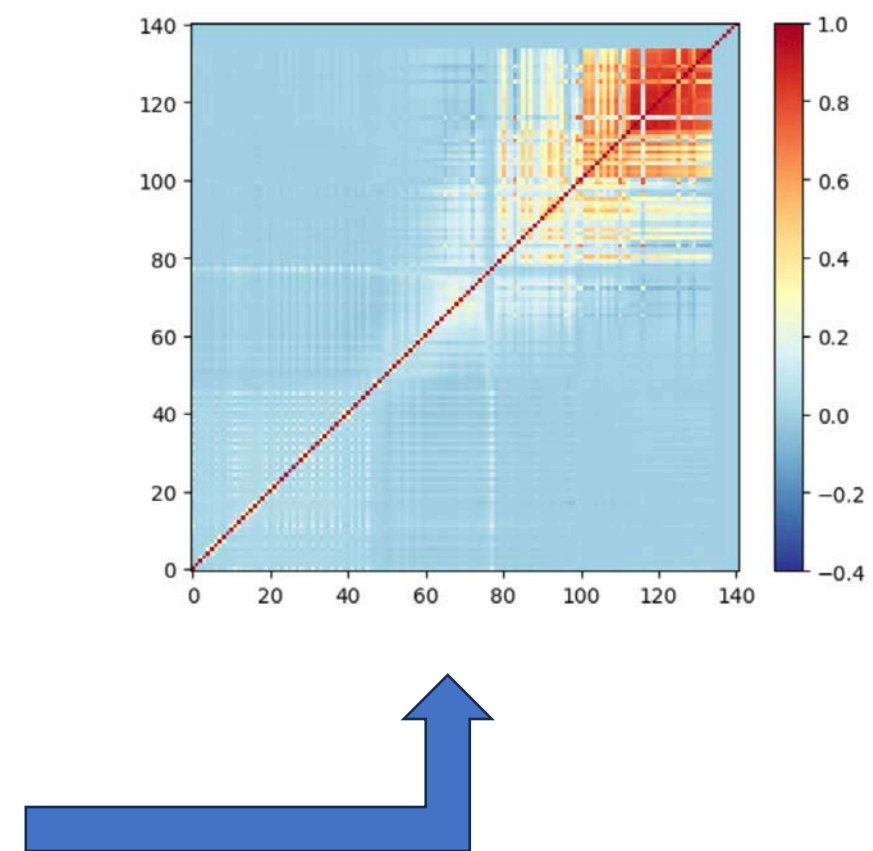
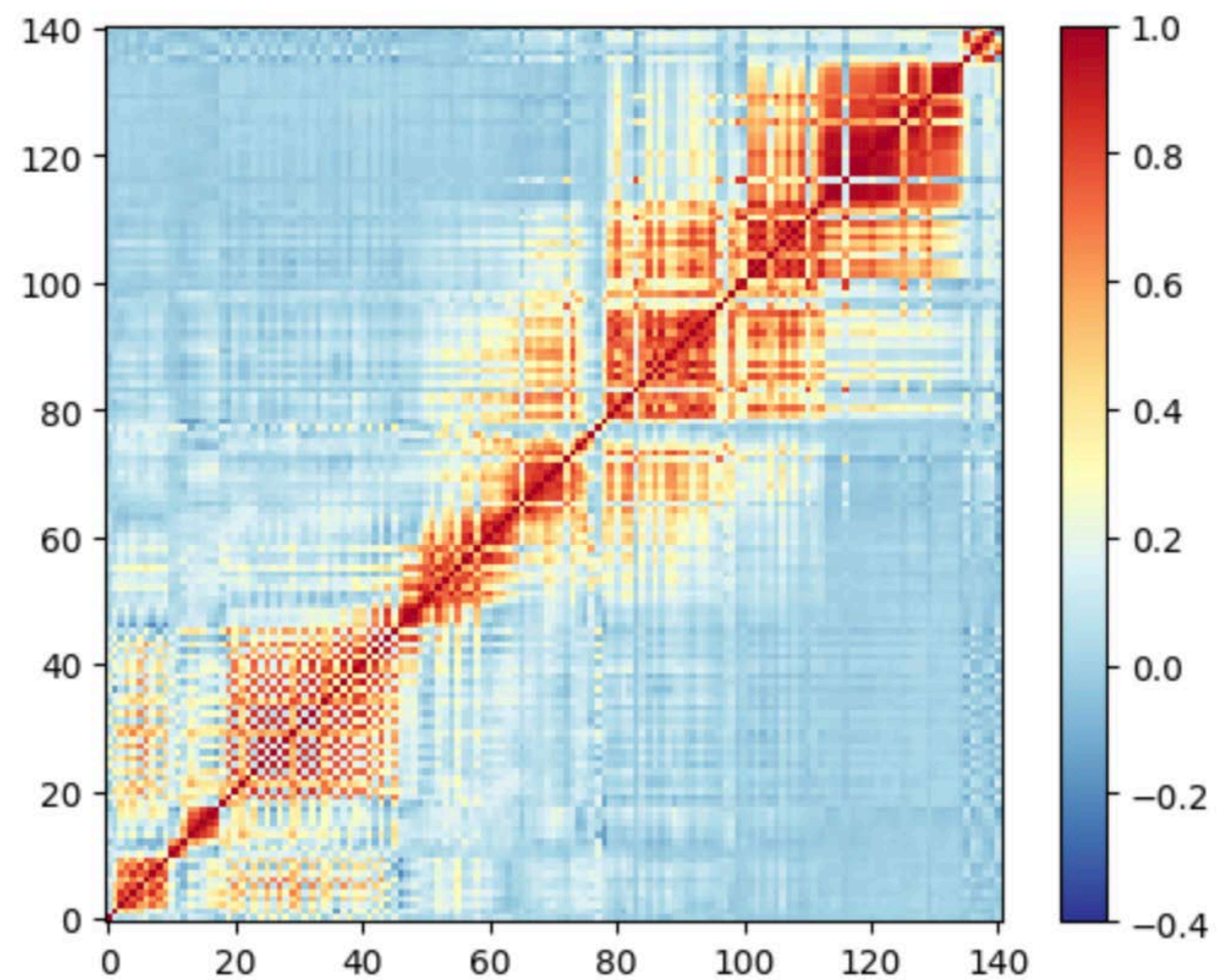
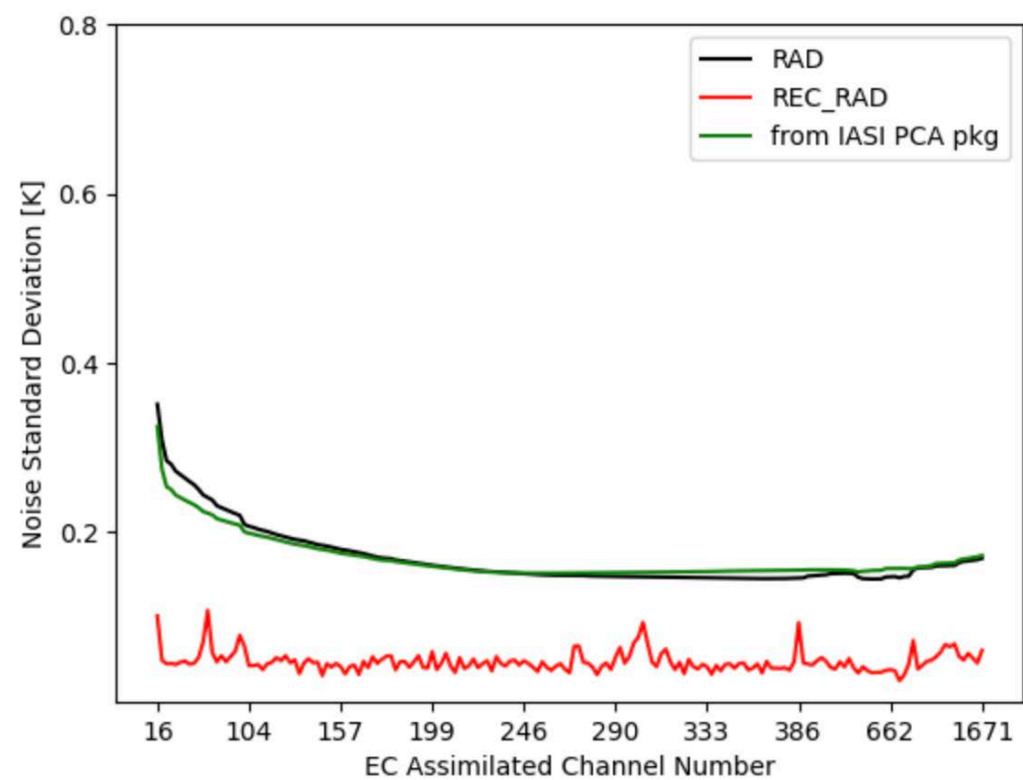




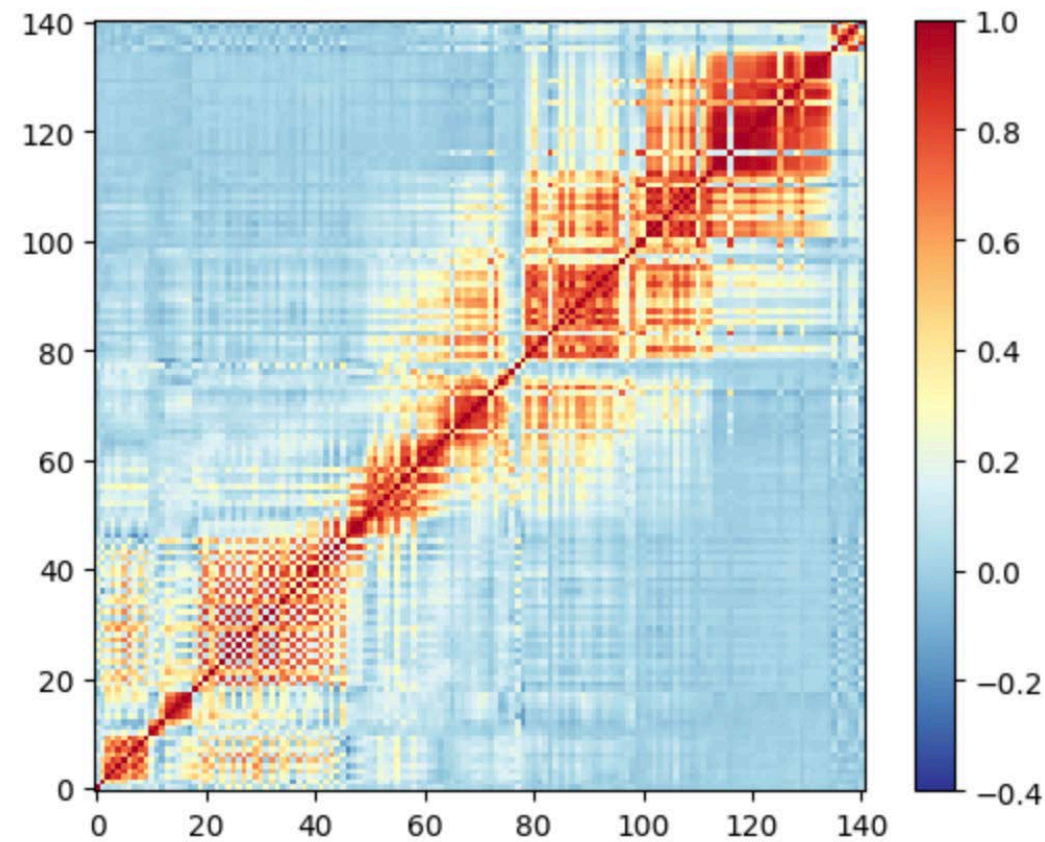




- Noise filtering performed by PC compression also generates very complex correlation structures
- Open question: How best to estimate the error (that will work in a DAS) when you don't have an existing instrument without PC compression (e.g., MTG-IRS, GeoXo, etc)?



**Desrozier Estimate + Difference in instrument Noise matrix and Noise that remains in reconstructed radiances**  
**Gives observation error covariance matrix similar to what is used operationally for conventional radiance assimilation**



- Rerun Reconstructed Radiance case with cloud detection consistent with background error estimate for reconstructed radiances
- Run experiment with using inflated Desroziers estimate based upon reconstructed radiances → pretend we're assimilating a "new instrument"
- Look at similar principal component products such as CrIS
- Prepare for assimilation of MTG-IRS
- PCA reconstructed radiances should be readily assimilated with some modifications to cloud detection
- Direct assimilation of PC scores with PCRTM a possibility
- Use of more observations over land being explored
- All-Sky cloud sensitive radiances → possibility, but a ways off (at least at GMAO)
- Improved cloud detection in GEOS-ADAS will be key