



# Retrieval of ambient aerosol properties from In-situ data

Joseph Schlosser<sup>1,2</sup> and Snorre Stamnes<sup>1</sup>

<sup>1</sup>Langley Research Center, <sup>2</sup>Oakridge Associated Universities

With direct collaboration from the HSRL-2 and the LARGE teams



# Objective and Significance



- The objective of the in-situ aerosol retrieval algorithm (ISARA) is to derive ambient aerosol optical and microphysical products from a suite of direct (i.e., in-situ) measurements
- In particular, the algorithm produces ambient aerosol size and complex refractive index, for both fine-mode and coarse-mode aerosol size distributions
- The ISARA aerosol products can be directly compared to polarimeter aerosol products and to high spectral resolution lidar aerosol measurements
- The ISARA aerosol products can also provide insight into aerosol behavior in complicated atmospheric environments and near clouds

# PACE-PAX In-situ Measurements Summary



- ❖ Dry aerosol particle size resolved number concentrations
  - DMT UHSAS (optical particle diameters from 0.06 to 1.1  $\mu\text{m}$ )
  - TSI-3321 APS (aerodynamic particle diameters from 1.1 to  $\sim 5 \mu\text{m}$ )
- ❖ Dry aerosol particle scattering coefficient at 450, 550, 700 nm wavelengths
  - TSI Scatter Nephelometer
- ❖ Dry aerosol particle absorption coefficient at 467, 530 and 660 nm wavelengths
  - PSAP
- ❖ Hygroscopic growth function
  - TSI Scatter Nephelometer
- ❖ Ambient relative humidity (RH)
  - DLH

## Additional Analyses and Measurements

- Inlet efficiency and cut-off based on LARGE/CIRPAS analysis
- Real refractive index (RRI) constrained by LI-Nephelometer

# Methods Summary

Step 1 (novel): Alignment of UHSAS and APS (Salibas et al., 20019)

Step 2: Retrieve dry complex refractive index (CRI)

- Spectral dry scattering and absorption coefficients
- Dry aerosol size resolved number concentration ( $n$ )

Step 3: Retrieve hygroscopicity ( $\kappa$ )

- Hygroscopic growth function ( $f(\text{RH})$ )

Step 4: Calculate ambient aerosol properties (Schlosser et al, 2025)

Step 5: Collocation (Schlosser et al., 2024)

Note: All aerosol properties calculated with the Modeled optical properties of ensembles of aerosol particles (MOPSMAP) package

- Schlosser et al., Closing the gap between in-situ and remotely sensed aerosol particle properties, TBD, Atmos. Meas. Tech.
- Schlosser, J. S., and Coauthors, 2024: Maximizing the Volume of Collocated Data from Two Coordinated Suborbital Platforms. J. Atmos. Oceanic Technol., 41, 189–201, <https://doi.org/10.1175/JTECH-D-23-0001.1>.
- Gasteiger, J. and Wiegner, M.: MOPSMAP v1.0: a versatile tool for the modeling of aerosol optical properties, Geosci. Model Dev., 11, 2739–2762, <https://doi.org/10.5194/gmd-11-2739-2018>, 2018.
- G. Saliba, et al., Factors driving the seasonal and hourly variability of sea-spray aerosol number in the North Atlantic, Proc. Natl. Acad. Sci. U.S.A. 116 (41) 20309-20314, <https://doi.org/10.1073/pnas.1907574116> (2019).

# Algorithm

Step 1: Alignment of UHSAS and APS.

$$D_{APS}(\rho) = \frac{D_{UHSAS}}{\sqrt{\rho}}$$

$$\rho: < 0.9: 0.1: 3.0 > \text{ g cm}^{-3}$$

Measured dry size distributions  
 $(n_{UHSAS}(D_{UHSAS})$   
 and  $n_{APS}(D_{APS})$

Fit  $n_{UHSAS}(D_{UHSAS})$  a the  
 power function  
 $y^p(D_{UHSAS}) \approx C1 \cdot D_{UHSAS}^{-C2}$

Iteratively compute objective function for  
 range of  $\rho$  given the changing number of  
 overlapping bins ( $n_{bins}$ ) at each step.

$$O(\rho) = n_{bins}^{-1} \sum_{i=1}^{n_{bins}} \log^2 \frac{y^p(D_{UHSAS})}{n_{APS}(D_{APS})}$$

$\bar{\rho} = \rho$  where  $O(\rho)$   
 $= \min(O(\rho))$   
Note: 1100 nm  
 stitching point

# Algorithm

Step 2: Retrieve dry complex refractive index (CRI)

$$\text{CRI} = \text{RRI} + \text{IRI} * i$$

**RRI: 1.55**

**IRI: <0.00,0.001:0.001:0.08>**

Measured dry scattering and absorption coefficient ( $C_{\text{dry}}$ )

Measured dry size distribution

Compute scattering and absorption coefficient ( $C_{\text{calc}}$ )

$$\zeta_{\text{sca}} = \left| \frac{C_{\text{calc}} - C_{\text{dry}}}{C_{\text{dry}}} \right|_{450,550,700}^{\text{sca}} \times 100\%$$

$$\zeta_{\text{abs}} = \left| C_{\text{calc}} - C_{\text{dry}} \right|_{470,532,660}^{\text{abs}}$$

**IRI** = mean(**IRI**), For all **IRI** where  $\zeta_{\text{sca}} < 20\%$  and  $\zeta_{\text{abs}} < 1 \text{ Mm}^{-1}$  in all three wavelengths

# Algorithm

## Step 3: Retrieve hygroscopicity ( $\kappa$ )

$$\left(\frac{D_{\text{amb}}}{D_{\text{dry}}}\right)^3 = g^3 = 1 + \kappa * \frac{\text{RH}}{100 - \text{RH}}$$

$\kappa$ :  $\langle 0.01:0.01:1.40 \rangle$   
CRI

Humidified scattering coefficient  
( $C_{\text{wet}} = C_{\text{dry}} * f(\text{RH})$ )

Measured dry  
size distribution  
Ambient RH

Compute  
scattering  
coefficient ( $C_{\text{calc}}$ )

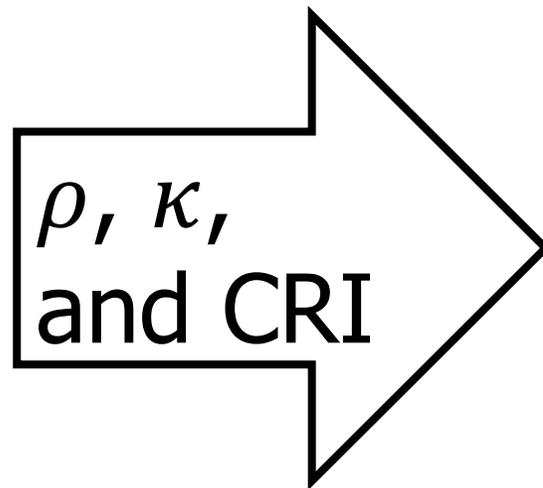
$$\zeta_{\text{sca}} = \left| \frac{C_{\text{calc}} - C_{\text{wet}}}{C_{\text{wet}}} \right|_{550}^{\text{sca}} \times 100\%$$

$\bar{\kappa} = \kappa$ , for smallest  
 $\kappa$  where  $\zeta_{\text{sca}} < 1\%$

## Inputs

- $n_{\text{UHSAS}}(D_{\text{UHSAS}})$
- $n_{\text{APS}}(D_{\text{APS}})$
- Ambient RH

## Retrieved



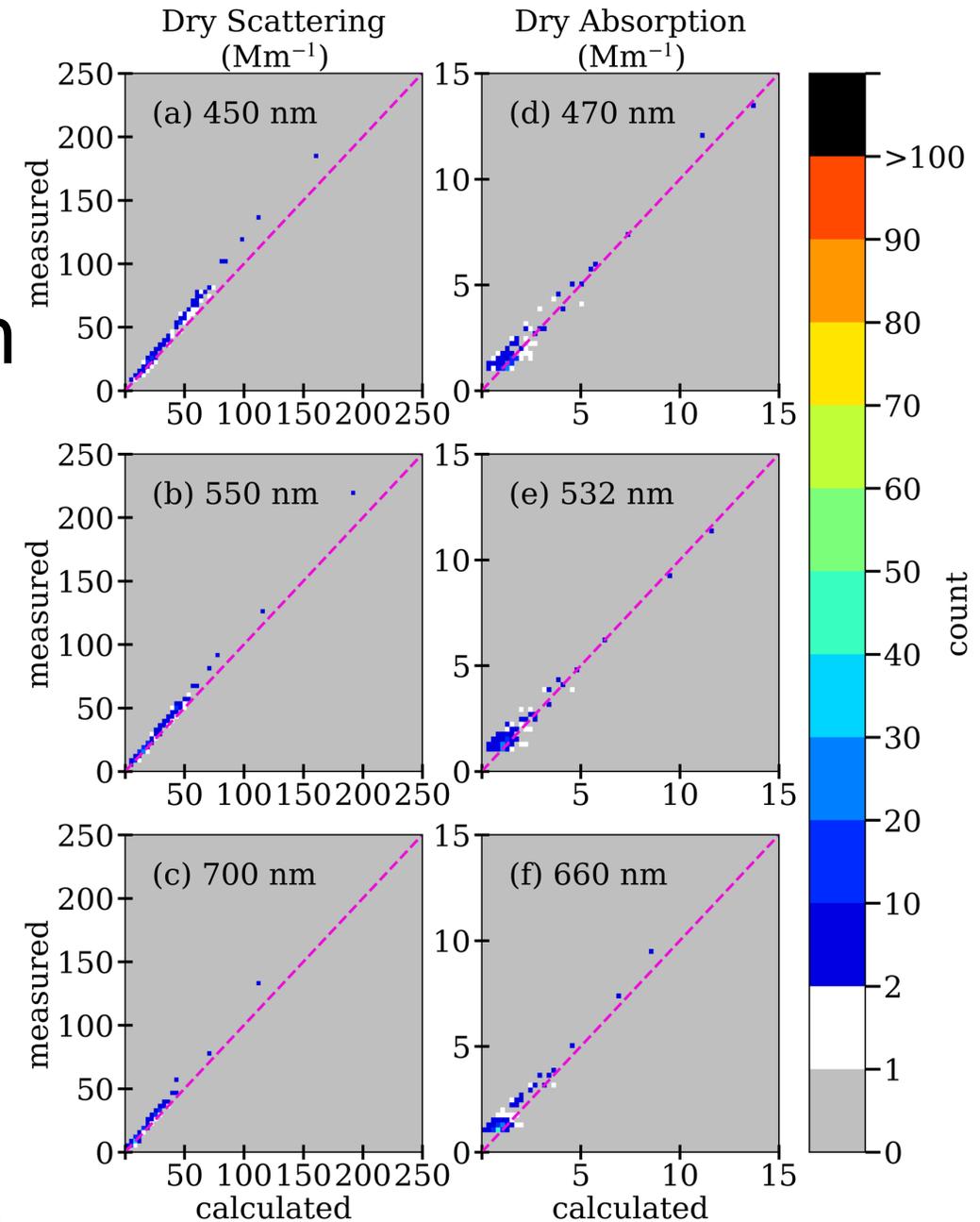
## Ambient aerosol properties:

- Number, surface and volume concentrations
- Spectral extinction coefficient and cross section
- Spectral single scattering albedo

# Internal Closure (Dry)

- All results from PACE-PAX
- Scattering coefficients agree to within 2.5%
- Absorption coefficients agree to within 5%

	Scattering			Absorption		
Wavelength (nm)	450	550	700	470	532	660
NMAD (%)	2.1	1.8	2.4	2.3	2.8	4.3



# Internal Closure (Dry)

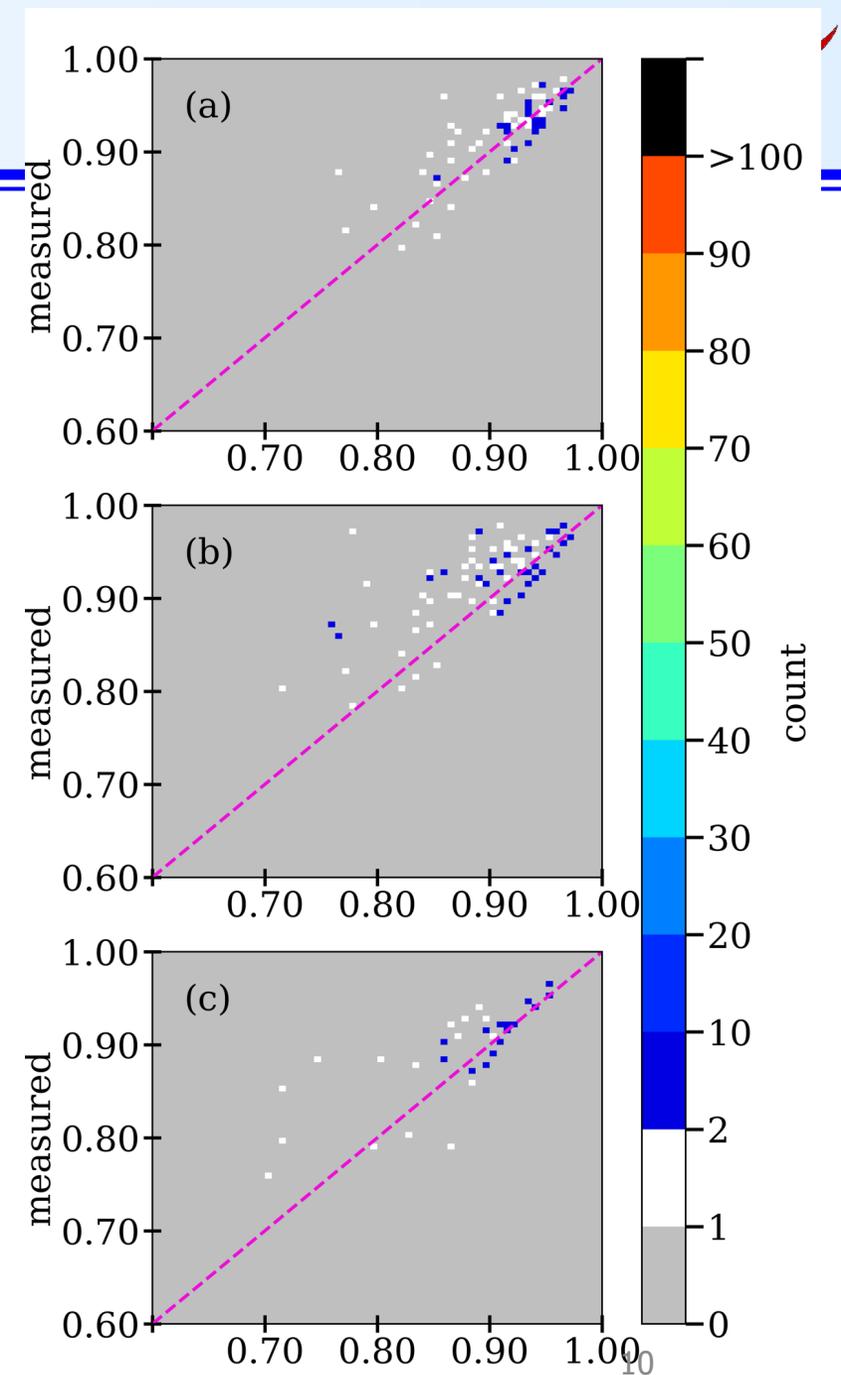
- Single scattering albedo (SSA) to within 15%.

Normalized-range mean absolute deviation (NMAD)

$$\sum_{j=1}^n \frac{|Y_j - X_j|}{n} \cdot \frac{100\%}{\max(X) - \min(X)}$$

where  $Y_j$  is set of calculated data,  $X_j$  is set of measurement data, and  $n$  is the total number of points.

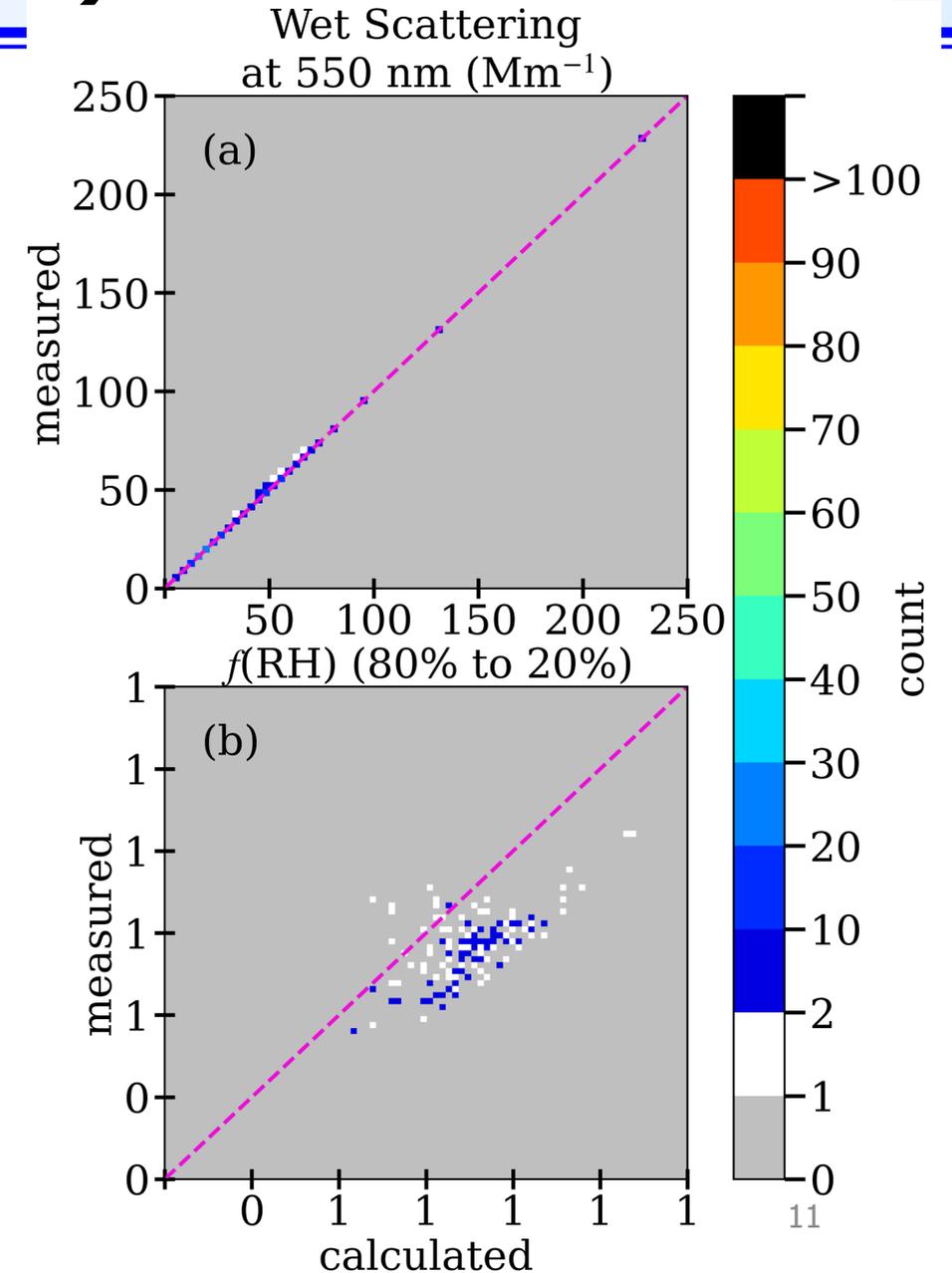
	SSA		
Wavelength (nm)	450	550	700
NMAD (%)	9.2	13.1	10.8



# Internal Closure (Humidified)

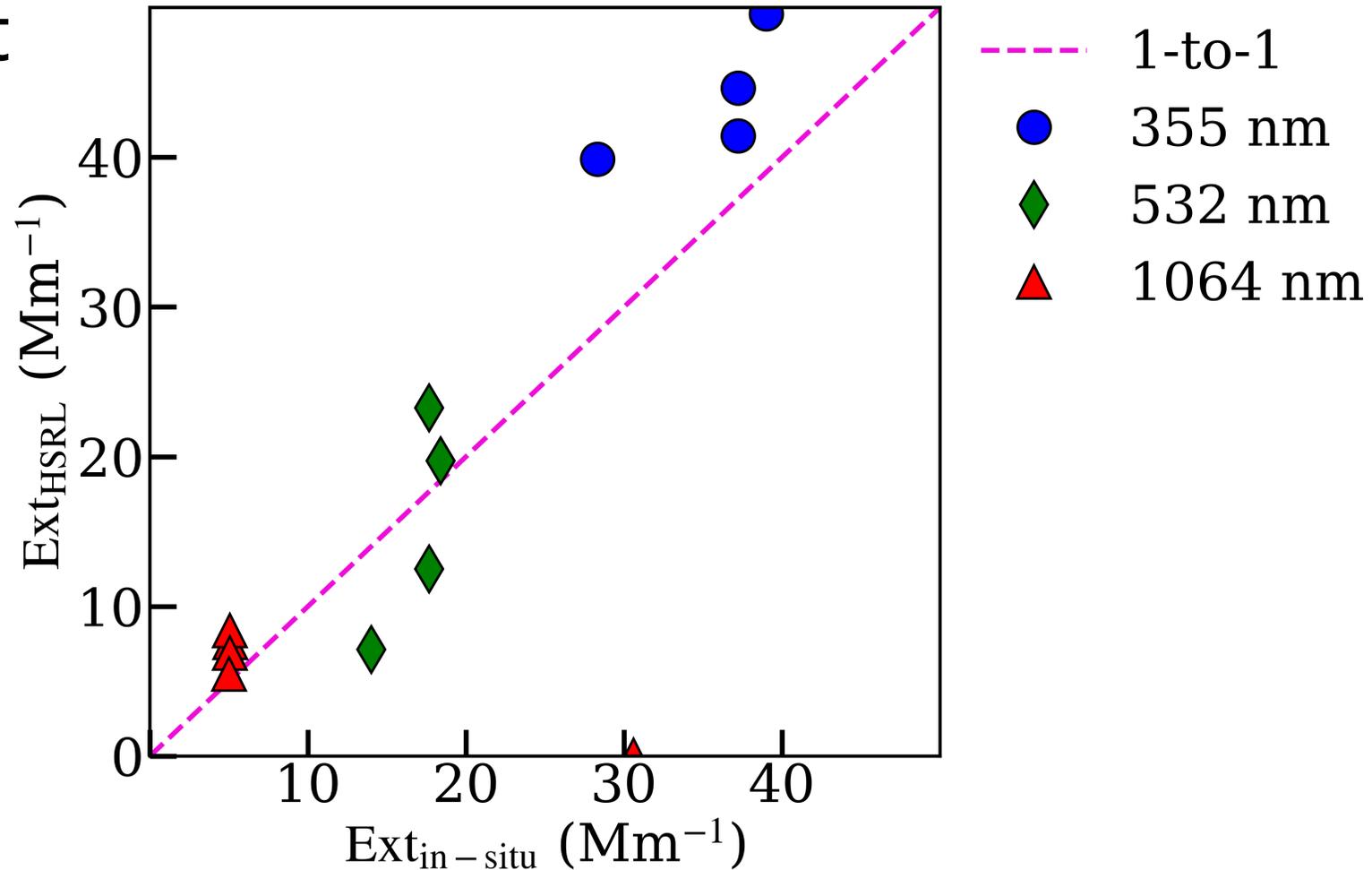
- Humidified (wet) scattering at 550 nm agrees to within 0.5%
- $f(\text{RH})$  agrees to within 7%

	Extinction	$f(\text{RH})$
NMAD (%)	0.2	21



# External Closure

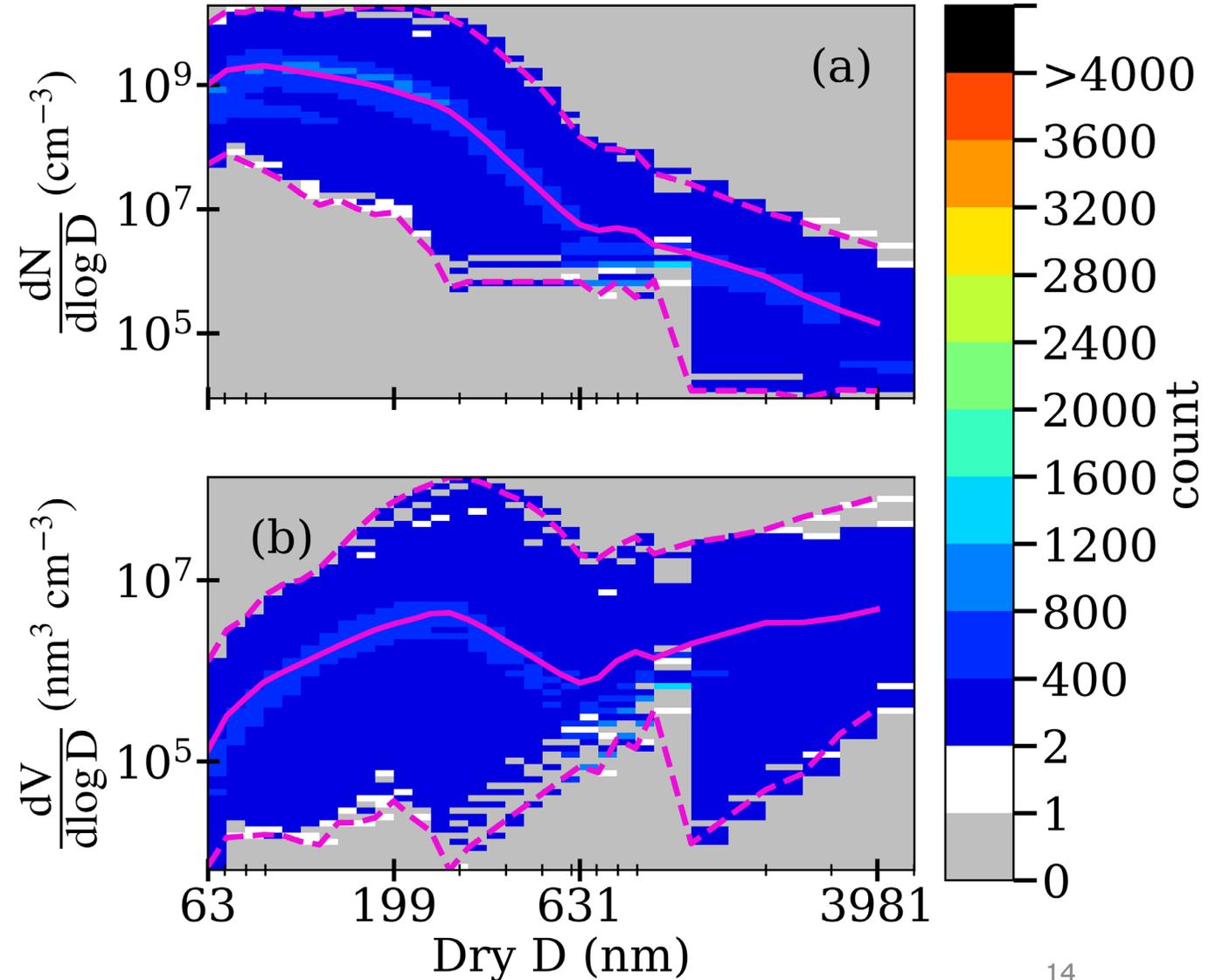
- HSRL-2 derived vs. in-situ derived spectral ambient aerosol extinction
- 532nm (NMAD: 26%, Count: 5, p-value: 0.01)



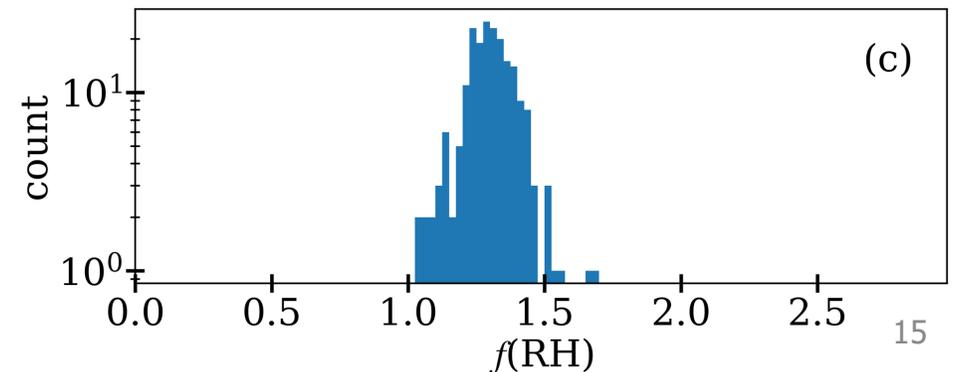
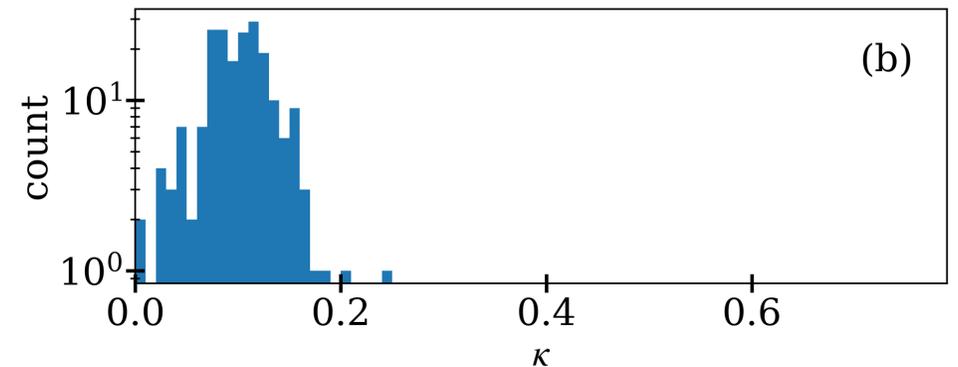
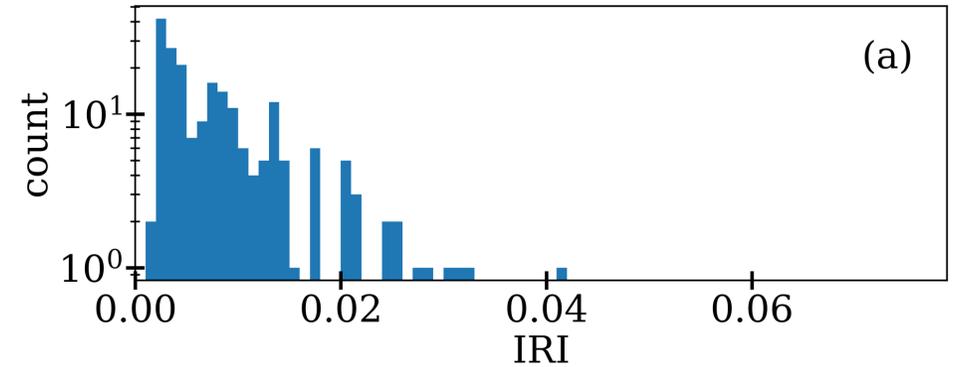
An aerial photograph of a coastline. The left side shows brown, hilly land with a road and some buildings. The right side is a vast expanse of blue ocean under a cloudy sky. A semi-transparent grey rectangle is overlaid in the center, containing the text "Questions?".

Questions?

- There is a small coarse mode present in low number concentrations but even for PACE-PAX, this is not common.

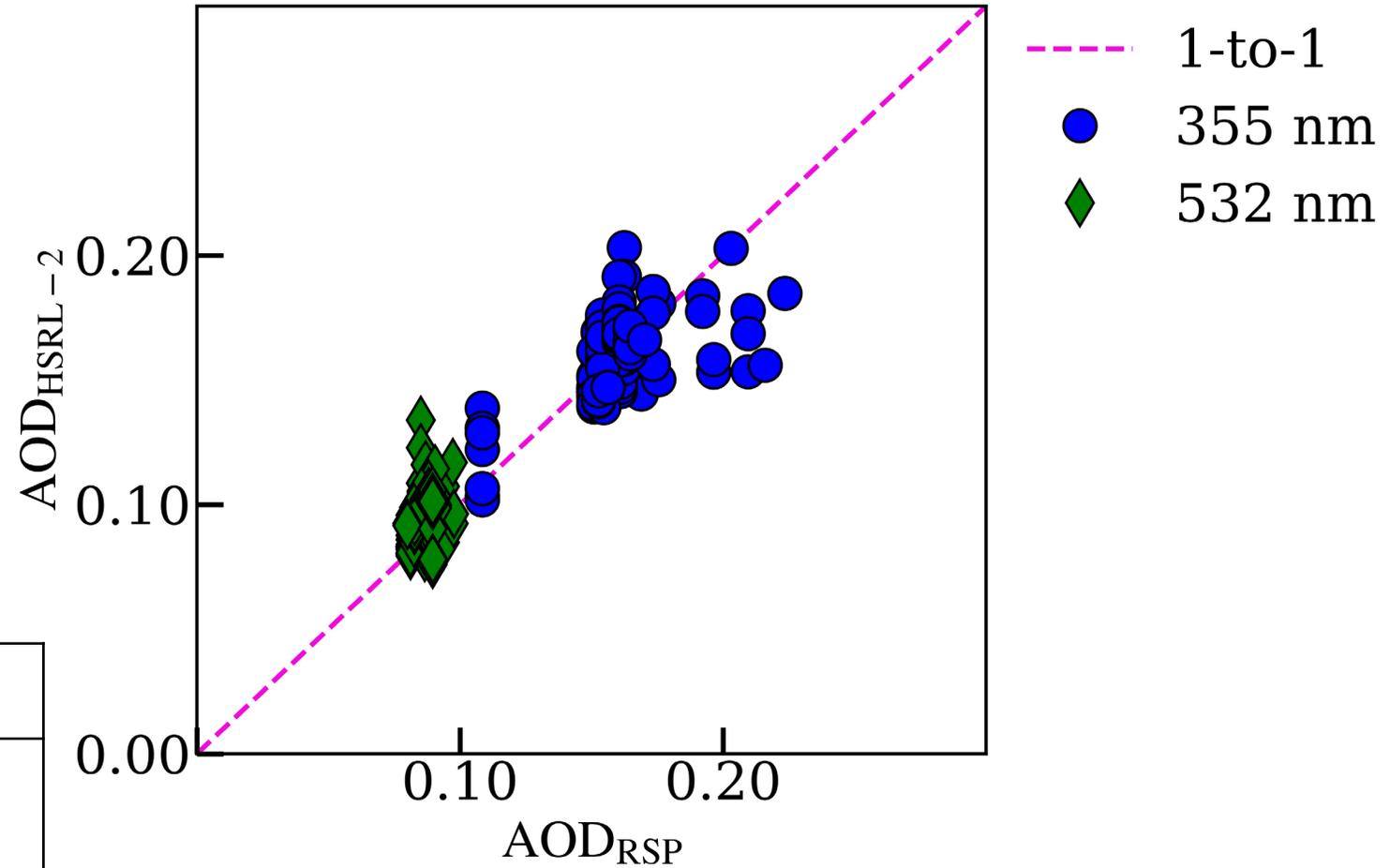


- Retrieved IRI and hygroscopicity ( $\kappa$ ) follow expected trends.
- Low-absorption and generally low hygroscopicity.



# HSRL-2 and RSP AOD Comparison

- HSRL-2 to RSP derived spectral AOD (count = 125).



Wavelength (nm)	AOD	
	355	532
NMAD (%)	20	3