

# 1 **Tropospheric nitrogen dioxide levels vary diurnally in Asian cities**

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33

34 **Abstract**

35 Satellite measurements of nitrogen dioxide have been used to infer nitrogen oxides emissions,  
36 a critical component in tropospheric chemistry and pollution. New observations from the  
37 Geostationary Environmental Monitoring Spectrometer offer a breakthrough by providing a  
38 daytime record of nitrogen dioxide over Asia. Here we present the summertime diurnal patterns  
39 of nitrogen dioxide at major cities, power plant regions, and Strait of Malacca. The  
40 Geostationary Environmental Monitoring Spectrometer data across various regions show high  
41 nitrogen dioxide in the morning that decrease in the afternoon, with varying hourly peaks,  
42 troughs, and amplitudes reflecting diurnal characteristics of local emissions and chemistry.  
43 Nitrogen oxides emissions inferred from Geostationary Environmental Monitoring  
44 Spectrometer and the Weather Research and Forecasting model coupled with Chemistry also  
45 show distinct patterns among regions: early morning peaks occur over Hanoi, Guangzhou, and  
46 Bangkok; mid-to-late morning peaks appear over Seoul and Beijing; and late afternoon peaks  
47 are noted in the Yangtze River Delta region. Top-down emissions incorporating temporal  
48 changes in the Geostationary Environmental Monitoring Spectrometer nitrogen dioxide yield  
49 the most accurate nitrogen dioxide simulations.

50

## 51 **Introduction**

52  $\text{NO}_x$  (=  $\text{NO} + \text{NO}_2$ ) is easily converted into nitrate aerosols and plays a key role in the  
53 formation of tropospheric ozone and acid rain<sup>1</sup>. Regular  $\text{NO}_2$  monitoring with improved  
54 capabilities can help control  $\text{NO}_x$  emission levels. Near-surface concentrations of air pollutants,  
55 including  $\text{NO}_2$ , show variable diurnal changes depending on their source type and chemical  
56 and meteorological conditions<sup>2</sup>. Hourly tropospheric  $\text{NO}_2$  monitoring afforded by space-borne  
57 Geostationary Earth Orbit (GEO) spectrometers promise to play a crucial role in next  
58 generation air quality monitoring. These GEO observations should have the benefit of diurnal  
59 estimating top-down  $\text{NO}_x$  emissions using more complete observations than based on twice-  
60 a-day measurements from Low Earth Orbit (LEO) sensors<sup>3-5</sup>.

61 Over the past two decades, satellite data have revealed substantial regional variations  
62 in tropospheric  $\text{NO}_2$  vertical column densities (VCDs) in industrial regions and cities  
63 worldwide. From 2000 to 2019, approximately 71% of urban regions experienced an increase  
64 in annual mean  $\text{NO}_2$  concentrations, averaging +0.4% per year<sup>6-7</sup>. However, trends differed by  
65 region. North America and Europe showed notable declines in  $\text{NO}_2$  levels, reflecting the impact  
66 of stringent air quality polices and emission reduction efforts<sup>6-7</sup>. In contrast, sub-Saharan Africa,  
67 the Middle East, and South Asia recorded increasing  $\text{NO}_2$  concentrations, largely attributed to  
68 urban expansion and industrialization. Among these regions, South Asia recorded the highest  
69  $\text{NO}_2$  exposures, which remained largely constant over the past two decades despite regional  
70 variations. In this study, we refer to vertical column density (VCD) as simply “column”.  
71 Although  $\text{NO}_2$  columns in some Asian countries have declined in recent years, they tend to  
72 remain higher than those in Europe and North America<sup>6-9</sup>. However, many Asian countries  
73 have relatively few air quality sites that can continuously monitor  $\text{NO}_2$  ambient concentrations.  
74 Until recently, this lack of spatial coverage of the surface monitoring networks has been

75 complemented by twice-a-day space-borne observations from LEO spectrometers such as  
76 Global Ozone Monitoring Experiment-2 (GOME-2) in the morning and Ozone Monitoring  
77 Instrument (OMI) and TROPOspheric Monitoring Instrument (TROPOMI) in early afternoon.  
78 This has recently changed with the launch of the two pioneering geostationary spectrometers:  
79 the Geostationary Environmental Monitoring Spectrometer (GEMS), which started hourly  
80 observations over Asia in 2020 and the Tropospheric Emissions: Monitoring Pollution  
81 (TEMPO) started hourly observations in 2023 over North America. The upcoming  
82 ESA/Copernicus Sentinel-4 geostationary spectrometer is planned for launch in 2025 for  
83 hourly observations over Europe and northern Africa.

84         Hyperspectral sensors on sun-synchronous polar-orbiting satellites have contributed  
85 substantially to identifying regional and global NO<sub>2</sub> spatial distribution characteristics and  
86 long-term trends, improving NO<sub>x</sub> emissions estimates and elucidating atmospheric NO<sub>x</sub>  
87 chemistry<sup>10-18</sup>. These instruments include the Global Ozone Monitoring Experiment  
88 (GOME<sup>12</sup>), the Global Ozone Monitoring Experiment-2 (GOME-2<sup>13</sup>), the SCanning Imaging  
89 Absorption spectroMeter for Atmospheric CHartography (SCIAMACHY<sup>14</sup>), the Ozone  
90 Monitoring Instrument (OMI<sup>15,16</sup>), and the TROPOspheric Monitoring Instrument  
91 (TROPOMI<sup>17</sup>). These data have been retrieved twice-a-day (morning and early afternoon)  
92 from the constellation of LEO measurements<sup>19</sup>. However, the hourly daytime observations  
93 from GEO spectrometers such as GEMS and TEMPO facilitate investigating diurnal processes  
94 that determine the atmospheric composition<sup>20-22</sup>. In addition, top-down NO<sub>x</sub> emissions have  
95 been estimated using tropospheric NO<sub>2</sub> column data available at limited times in the morning  
96 and/or afternoon from LEO measurements. When a single observation per day is available to  
97 infer emissions, it is necessary to assume that the existing knowledge of diurnal variations in  
98 chemistry and emissions are correct<sup>23</sup>. There is thus a need to evaluate previous assumptions  
99 about diurnal variations in NO<sub>x</sub> emissions using tropospheric NO<sub>2</sub> data from geostationary

100 measurements. Studies have already examined the diurnal variations of observed NO<sub>2</sub> columns  
101 over limited regions using ground-based and airborne observations<sup>24-26</sup> as well as the  
102 constellation of LEO measurements<sup>27,28</sup>. These studies demonstrate diverse diurnal changes in  
103 column NO<sub>2</sub> concentrations, as well as a significant month-to-month and day-to-day variability  
104 in diurnal changes at a single site. These variations are collectively influenced by emissions,  
105 chemistry, and transport. Daytime spatial NO<sub>2</sub> distributions are now available, for the first time,  
106 at hourly resolution from Geostationary Environmental Monitoring Spectrometer (GEMS)  
107 observations. Detailed field of regard (FOR) and times of GEMS measurements are provided  
108 in Supplementary Fig. 1 and Supplementary Table 1. The unprecedented high spatiotemporal  
109 resolution of GEMS measurements can provide spatially complete information on diurnal  
110 variability in vertical column NO<sub>2</sub> densities over Asia, which is often lacking in both in-situ  
111 and remote-sensing observations.

112         Several papers regarding GEMS hourly data have been published recently<sup>29-31</sup>. One of  
113 them<sup>29</sup> was focused on understanding the diurnal variation in the GEMS tropospheric NO<sub>2</sub>  
114 column retrievals over Northeast Asia and Seoul in South Korea using a Multi-Scale  
115 Infrastructure for Chemistry and Aerosols (Version 0, MUSICAv0) 3D chemical transport  
116 model in January and June 2023. Another study<sup>30</sup> shows similar variability and diurnal  
117 variation of the NO<sub>2</sub> column in Beijing and Seoul using GEMS data, GEOS-Chem chemical  
118 transport model and ground-based Pandora observations. In addition, they examined the role  
119 of transport in the diurnal variation of NO<sub>2</sub> column. The other study<sup>31</sup> presents the correction  
120 for the operational GEMS L2 NO<sub>2</sub> columns using a machine learning model to train with  
121 TROPOMI NO<sub>2</sub> columns measured in early afternoon around 1:30 pm local solar time, because  
122 TROPOMI NO<sub>2</sub> retrievals are more mature. In addition, in the previous studies<sup>32,33</sup>, the surface  
123 NO<sub>2</sub> concentration has been estimated using the GEMS L2 NO<sub>2</sub> product, and uncertainty has

124 been quantified. As presented in these studies the hourly data from GEMS can be used to gain  
125 new insights in diurnal evolution in atmospheric composition.

126         Herein, we identify the diurnal characteristics of tropospheric NO<sub>2</sub> column not only  
127 East Asia but also major cities and hotspots over Asia, where a considerable portion of global  
128 NO<sub>x</sub> emissions occur. Three emissions source types are considered: cities, power plants, and  
129 ship tracks. This study focuses on the data collected during summer when the measurement  
130 data are available for a longer time interval of the day than in other seasons. Top-down NO<sub>x</sub>  
131 emissions for each source are estimated utilizing a chemical transport model combined with  
132 GEMS observations. We also demonstrate the ability of GEMS to detect the transport of NO<sub>2</sub>  
133 plumes across South Korea within a single day.

134

## 135 **Results and Discussion**

### 136 **Diurnal dependency of tropospheric NO<sub>2</sub> spatial distributions**

137 To identify the dependency of the distribution on the time of the day, for the first time,  
138 we investigated diurnal cycles of multiple source regions over Asia as observed by GEMS for  
139 available times. We selected Seoul, Beijing, the Yangtze River Delta (including Shanghai),  
140 Guangzhou, Bangkok, and Hanoi as cities; power plant complexes in China named “Wuhai”  
141 and “Baotou” and clusters of power plants in India labeled “India P1”, “India P2”, and “India  
142 P3”; and one ship track region, the Strait of Malacca (see Fig. 1 and 2). In this study, in order  
143 to investigate the characteristics of diurnal variation in NO<sub>2</sub> column and estimate the NO<sub>x</sub>  
144 emissions at hotspots over Asia, we focused on major cities and major hotspots that have huge  
145 anthropogenic source regions over GEMS field of regards. Some of major cities such as Tokyo  
146 were excluded due to insufficient observations. Locations of each source box and power plants,  
147 along with the number of GEMS data for each source, are detailed in Supplementary Tables 2-  
148 3 and Supplementary Figs. 3-6. Fig. 1 shows GEMS tropospheric NO<sub>2</sub> columns during  
149 weekdays and weekends for all source regions. The GEMS NO<sub>2</sub> values collected at a swath  
150 passing through a specific source box area with cloud fraction  $\leq 0.3$ , solar zenith angle  $\leq 70^\circ$ ,  
151 and viewing zenith angle  $\leq 70^\circ$  for summer of 2021 (June 1–August 31) and 2022 (June  
152 1–August 31) are used to generate the mean and standard deviation. With the exception of  
153 Seoul, there are no significant differences in NO<sub>2</sub> columns between weekdays and weekends  
154 throughout the day across all source regions. A previous study indicated no significant  
155 weekday-weekend differences in satellite NO<sub>2</sub> columns in China<sup>34</sup>. In Seoul, there are large  
156 reductions in columns (up to 30%) on weekends compared to weekdays. There are some  
157 weekend NO<sub>2</sub> reductions observed in Bangkok and Hanoi (Fig. 1 and Supplementary Figs. 4-  
158 7).

159 In Fig. 1, the WRF-Chem NO<sub>2</sub> columns using constant NO<sub>x</sub> emissions are compared  
160 with the GEMS columns to investigate how the satellite columns vary if there were no diurnal  
161 changes in the emissions. The model results were normalized to have the same average column  
162 values as the GEMS tropospheric columns for available times. Overall, the GEMS NO<sub>2</sub>  
163 columns show peaks in the morning, decrease after the peaks, and increase in the late afternoon  
164 in the cities (Fig. 1 and 2). The time of maximum in the morning ranges from 08:00 LT to  
165 11:00 LT and the time of minimum in the afternoon varies from 13:00 LT to 16:00 LT. The  
166 minimum column values in the afternoon are 17% to 40% lower than the maximum values.  
167 Diurnal pattern of GEMS NO<sub>2</sub> and inferred NO<sub>x</sub> emissions for each region is presented below.

168

169 - *Seoul*

170 In Seoul, the average tropospheric NO<sub>2</sub> column was low in the early morning (08:00–  
171 09:00 LT), peaked in the late morning (11:00 LT), decreased in the early afternoon, and then  
172 started to increase again in the late afternoon (15:00 LT). The minimum GEMS column value  
173 in the afternoon is approximately 30% lower than the maximum value in the late morning. The  
174 WRF-Chem model results using a constant emission over Seoul exhibit different diurnal  
175 patterns, particularly before noon. The model shows the highest value in 08:00 LT, decreasing  
176 until the early afternoon, and then increasing again after 14:00 LT. The minimum model  
177 column value in the afternoon is approximately 40% lower than the maximum value at 08:00  
178 LT (Supplementary Table 4). The observation that GEMS columns are lower than WRF-Chem  
179 columns in the early morning, but higher from noon to early afternoon, suggests that NO<sub>x</sub>  
180 emissions are lower than the assumed constant emissions during nighttime and higher than the  
181 constant emissions during daytime.

182 In Fig. 2, GEMS shows broadly enhanced NO<sub>2</sub> columns over the Seoul-Gyeonggi  
183 region, peaking at its center, Seoul, where the column values reach approximately  $20 \times 10^{15}$

184 molec (molecular)  $\text{cm}^{-2}$ . The plot clearly shows increased columns over a large area in Seoul  
185 around 10:00–12:00 LT, a shrinkage of these plumes at 14:00 LT, and slight increases again at  
186 16:00 LT (follow dark red to black colored areas).

187 The  $\text{NO}_2$  columns from Pandora, a direct-sun spectrometer that retrieves the  $\text{NO}_2$  with  
188 a spectral resolution of 0.5 nm using wavelength ranges from 270 to 530 nm, show a significant  
189 peak from mid to late morning at several locations across the Seoul-Gyeonggi region, similar  
190 to the GEMS columns over Seoul<sup>25</sup>. Additionally,  $\text{NO}_2$  columns measured from the Airborne  
191 Compact Atmospheric Mapper (ACAM), which uses the ultraviolet and visible band from 304  
192 to 520 nm with spectral resolution of 0.8 nm onboard the UC-12 aircraft, did not display an  
193 early morning peak, which is consistent with the GEMS data<sup>26</sup>. Further validation<sup>35</sup> of the  
194 GEMS L2  $\text{NO}_2$  product using 1 year of the tropospheric  $\text{NO}_2$  columns IUP-UB (Institute of  
195 Environmental Physics at the University of Bremen) product, TROPOMI product, and ground-  
196 based measurements has demonstrated a correlation coefficient of 0.76, despite a general  
197 overestimation in most regions. The diurnal variation captured by GEMS observation also  
198 showed good agreement with ground-based observations, with variations depending on local  
199 pollution levels. These findings further support the reliability of GEMS in monitoring  $\text{NO}_2$   
200 diurnal variations and emphasize the importance of regional emission characteristics in  
201 interpreting early morning peaks.

202 In Fig. 3, the top-down emissions in Seoul, estimated from GEMS data under a steady-  
203 state assumption (TD\_S) using a simple mass balance approach, exhibit a peak at 12:00 LT.  
204 However, the model results using TD\_S did not reproduce the observed diurnal variations of  
205 GEMS  $\text{NO}_2$  columns (correlation coefficient = -0.23) (Fig. 3). Another top-down emission  
206 estimation based on a non-steady-state assumption (TD\_NS) reproduces the observed columns;  
207 it shows high emissions during the early to mid-morning (peaking between 9:00 and 11:00 LT)  
208 and reduced emissions from noon to the afternoon. The recent studies also addressed the need

209 for non-steady state approaches<sup>36,37</sup>. The model results using TD\_NS closely replicated the  
210 GEMS observations, which peaked between 10:00 and 12:00 LT, achieving the highest  
211 correlation coefficient (0.88) and the lowest root-mean-square error (RMSE). Our study  
212 highlights the importance of morning emission accumulation in driving maximum NO<sub>2</sub>  
213 columns between 10:00 and 12:00 LT in Seoul.

214 Top-down emission estimations in this study are explained in the Methods section and  
215 Supplementary Figs. 8-12 and Supplementary Table 5. Details on the top-down emission  
216 estimation method with non-steady state assumption will be discussed later for certain source  
217 regions in South Asia.

218

#### 219 - *Beijing*

220 Similar to Seoul, GEMS NO<sub>2</sub> columns in Beijing exhibits a peak at 10:00 LT and trough  
221 at 16:00 LT. The model diurnal pattern, assuming constant emissions, shows higher NO<sub>2</sub>  
222 columns than GEMS from 08:00 LT to 10:00 LT, and lower values from 10:00 LT to 14:00  
223 LT. High nighttime emissions in the constant emission case cause large model columns in the  
224 early morning (Fig. 3). The top-down NO<sub>x</sub> emissions with a steady state assumption (TD\_S)  
225 in Beijing peak between 11:00 LT and 12:00 LT (Fig. 3). The top-down NO<sub>x</sub> emissions under  
226 a non-steady-state assumption (TD\_NS) in Beijing show a higher peak at 11:00 LT and lower  
227 trough values in the late afternoon compared to those under a steady-state assumption (Fig. 3).  
228 The model results using the control emissions and both TD\_S and TD\_NS show higher NO<sub>2</sub>  
229 column values than GEMS in the early morning but agree more closely with GEMS  
230 observations in the mid-day. In the late afternoon, the model using TD\_S and TD\_NS shows  
231 higher NO<sub>2</sub> columns again, but TD\_NS compared to TD\_S led to smaller biases. Among these,  
232 the TD\_NS results exhibit the highest correlation and the lowest RMSE value. The discrepancy  
233 between the model and GEMS values in the early morning may stem from elevated nighttime

234 NO<sub>x</sub> emissions used in this study. These findings highlight the importance of accurately  
235 representing nighttime NO<sub>x</sub> emissions and their associated chemistry to improve predictions of  
236 NO<sub>2</sub> columns in the early morning. The ratio of NO<sub>x</sub> to NO<sub>2</sub> is a parameter that is needed to  
237 derive NO<sub>x</sub> emissions from NO<sub>2</sub> observations. This parameter may add an uncertainty in  
238 estimating top-down emissions in the morning. Beijing shows that emissions are higher in mid-  
239 to-late morning than during early morning commuting hours in the top-down emissions with a  
240 non-steady-state assumption. In this city, the minimum GEMS (model) column value in the  
241 afternoon is approximately 30% (45%) lower than the maximum value in the late morning  
242 (Supplementary Table 4).

243

244 - *Yangtze River Delta (Shanghai and its surrounding areas)*

245 The Yangtze River Delta region in this study is a large area including Shanghai and  
246 nearby cities (Fig. 2). In Fig. 1, GEMS NO<sub>2</sub> columns in this area peaks at 08:00 LT and reach  
247 a minimum at 13:00 LT, with a relatively small difference between maximum and minimum  
248 (the ratio of minimum to maximum is 0.83). The model column also shows a peak at 08:00 LT,  
249 but decreases more rapidly over time compared to the GEMS data (Fig. 1). Interestingly, the  
250 top-down NO<sub>x</sub> emissions (both TD\_S and TD\_NS) over the Yangtze River Delta exhibit the  
251 lowest value at 08:00 LT and increases until 16:00 LT (Fig. 3). Note that GEMS NO<sub>2</sub> columns  
252 at 16:00 LT is high and comparable to those at 10:00 LT (Fig. 2). This may reflect unique  
253 traffic patterns in the Yangtze River Delta cities or be related to port activities. Due to minimal  
254 temporal variations in GEMS NO<sub>2</sub> columns over the Yangtze River Delta, the results of TD\_S  
255 and TD\_NS are similar. The model results based on top-down emissions demonstrate the  
256 highest correlation with observations and the lowest RMSE, underscoring the value of GEMS  
257 data. The increase in NO<sub>x</sub> emissions during the day may impact photochemistry relevant to

258 surface ozone and aerosols. Further study using in-situ observations and chemical transport  
259 models is needed to quantify the effects of diurnal variations of NO<sub>x</sub> emission over the Yangtze  
260 River Delta.

261

#### 262 - *Guangzhou, Bangkok, and Hanoi*

263 In Guangzhou, Bangkok, and Hanoi, GEMS NO<sub>2</sub> columns are slightly higher than the  
264 model columns in the morning but slightly lower in the afternoon (Fig. 1). The spatial  
265 distribution of GEMS NO<sub>2</sub> columns highlights large GEMS NO<sub>2</sub> plumes over these cities,  
266 particularly in the early morning (Fig. 2). In contrast to Hanoi, Haiphong and Nam Dinh,  
267 located east and south of Hanoi respectively, show persistently high NO<sub>2</sub> columns throughout  
268 the day (Fig. 2 and Supplementary Figs. 6-7). This pattern may be attributed to traffic routes  
269 connecting Hanoi, Haiphong, and Nam Dinh. In Guangzhou and Bangkok, the minimum values  
270 in GEMS columns show approximately a 30% reduction from the maximum values. This  
271 reduction is even more significant in Hanoi, where it reaches 40% (Supplementary Table 4).  
272 Guangzhou, Bangkok, and Hanoi exhibit notably high top-down NO<sub>x</sub> emissions early in the  
273 morning, followed by a rapid decrease toward noon under the non-steady-state assumption  
274 (TD\_NS) (Fig. 3). Model results using TD\_NS achieve the highest correlation (R = 0.92–0.96)  
275 and the lowest RMSE, whereas bottom-up emissions and top-down emissions under a steady-  
276 state assumption show poor correlations with GEMS observations.

277 Unexpectedly, the NO<sub>x</sub> chemical lifetime ( $\tau_c$ ) in Hanoi, calculated based on chemical  
278 loss through HNO<sub>3</sub> formation via reaction with OH or N<sub>2</sub>O<sub>5</sub> heterogeneous uptake, is the  
279 longest among all source regions (Supplementary Table 5). In Hanoi, the modeled OH  
280 concentration is the lowest among all source regions. Further analysis suggests that extremely  
281 high isoprene emissions in this region deplete OH<sup>38,39</sup>, thereby potentially increasing the NO<sub>x</sub>

282 chemical lifetime (Supplementary Figs. 10-11<sup>40,41</sup>). A comparison of WRF-Chem HCHO and  
283 TROPOMI HCHO columns reveals significant positive model biases south of 30°N in the  
284 domain, particularly in Hanoi, power plant regions in India, and countries near the Strait of  
285 Malacca. These biases are partly attributed to overestimated biogenic volatile organic  
286 compound emissions, including isoprene (Supplementary Fig. 11). This model deficiency  
287 affects the accuracy of top-down emission estimates. Consequently, in this study, we utilized  
288  $\tau_r$ , a surrogate for the directly calculated chemical lifetime ( $\tau_c$ ).  $\tau_r$  is indirectly derived from a  
289 non-steady-state mass balance equation based on WRF-Chem results (Methods section). Under  
290 non-steady-state assumptions,  $\tau_r$  yields non-negative top-down NO<sub>x</sub> emissions in Hanoi, unlike  
291  $\tau_c$ .

292

### 293 - *Chinese Power Plant Regions (Wuhai and Baotou)*

294 Besides cities, the GEMS NO<sub>2</sub> column was also analyzed for power plants and shipping  
295 routes. Fig. 2 and Supplementary Fig. 3 provide enhanced GEMS columns near power plants  
296 located in Wuhai and Baotou and their proximity in north China and northeastern India. Details  
297 about power plants in each location are provided in Supplementary Fig. 2 and Supplementary  
298 Table 3. Fig. 2 also illustrates NO<sub>2</sub> plumes from ship tracks detected by GEMS in the Strait of  
299 Malacca. In Wuhai and Baotou, similar to Seoul and Beijing, GEMS NO<sub>2</sub> columns peak in the  
300 late morning around 11:00 LT and decrease over time (Fig. 1). The minimum GEMS values  
301 are only 20% lower than the maximum values in Wuhai and Baotou, whereas the model using  
302 constant emission shows approximately a 40-50% decrease. The model columns are  
303 significantly higher than GEMS columns in the early morning and lower in the afternoon. Top-  
304 down NO<sub>x</sub> emissions with a steady-state assumption (TD\_S) peak at noon, but show small  
305 variations throughout the day in the regions around Chinese power plants (Fig. 3). Top-down

306 NO<sub>x</sub> emissions under a non-steady-state assumption (TD\_NS) exhibit a moderate peak in the  
307 late morning over Wuhai and a sharp peak in the early morning over Baotou. In the studied  
308 region, both power plants and other area source regions play significant roles, contributing to  
309 the diurnal variations in GEMS NO<sub>2</sub> columns over power plants in China, which to some extent  
310 resemble patterns observed in cities. The model results using both TD\_S and TD\_NS show  
311 good correlation and low RMSE for Wuhai. For Baotou, the results using TD\_NS exhibit the  
312 highest correlation and the lowest RMSE.

313

#### 314 - *Indian Power Plant Regions*

315 In Indian power plant regions, the GEMS NO<sub>2</sub> columns typically peak around 08:30  
316 LT, gradually decreasing thereafter (Fig. 1). These GEMS columns are slightly higher than the  
317 model columns in the early morning but diminish in the afternoon. The GEMS minimum values  
318 in the India P1, P2, and P3 regions experience significant reductions, with drops of 34%, 32%,  
319 and 37%, respectively, from their maximum values. In contrast, the model exhibits smaller  
320 reductions of 17%, 23%, and 25% for the India P1, P2, and P3 regions, respectively  
321 (Supplementary Table 4). Top-down NO<sub>x</sub> emissions with a steady-state assumption for these  
322 power plants in India peak in the early morning and gradually decrease over time, as shown in  
323 Supplementary Fig. 12. In contrast, top-down NO<sub>x</sub> emission calculations based on a non-  
324 steady-state assumption yield negative emissions for most power plant regions in India even  
325 using  $\tau_r$ . This is attributed to the significant negative temporal change terms and smaller  
326 chemical sink terms, which result from longer chemical lifetimes and lower OH concentrations  
327 compared to the power plant regions in China (Supplementary Table 5 and Supplementary  
328 Figs. 10-12).

329

330 - ***Strait of Malacca (shipping source region)***

331 GEMS NO<sub>2</sub> along the ship routes in the Strait of Malacca reach their peak at 09:00 LT,  
332 declining over time (Fig. 1). The minimum GEMS value in the late afternoon is 40% lower  
333 than the maximum observed in the early morning. The model exhibits an even more substantial  
334 reduction, at 60%, in the Strait of Malacca (Supplementary Table 4). Interestingly, as shown  
335 in Supplementary Fig. 12, the derived top-down NO<sub>x</sub> emissions with a steady-state assumption  
336 increase steadily throughout the day. Top-down NO<sub>x</sub> emission calculations based on a non-  
337 steady-state assumption yield negative emissions for the Strait of Malacca, similar to those  
338 observed in power plant regions in India (Supplementary Table 5). Both model results using  
339 bottom-up and available top-down emission show strong correlations (correlation coefficient  
340 = ~1) with GEMS data (Supplementary Fig. 12).

341

342 **Transport of NO<sub>2</sub> plumes**

343 When atmospheric species are transported to a receptor area, the transported species  
344 can influence changes in both the temporal and spatial distribution of the species and their  
345 products in the receptor area<sup>42,43</sup>. GEMS have, for the first time, captured the hourly transport  
346 of NO<sub>2</sub> plumes from Seoul and several other hotspots in the Korean peninsula to receptor  
347 regions several hundred kilometers away (Fig. 4). In this study, in order to present the  
348 possibility of hourly scale observation for NO<sub>2</sub> transport, we could briefly touch on this long-  
349 range transport case study. Therefore, we just focused on two long-range transport events that  
350 occurred in the same period (March 22, 2021) which were determined by method adopted from  
351 the previous study<sup>44</sup>. The difference from this previous study is that we used a simple approach  
352 with the ECMWF Reanalysis v5 (ERA5) wind data<sup>45</sup> to compare with the movement of the  
353 NO<sub>2</sub> plume, rather than using the more precise and advanced hybrid single particle Lagrangian

354 integrated trajectory (HYSPLIT) model for event detection. Therefore, this brief case study can  
355 be extended with more profound interpretations in future studies. On March 22, 2021, a large  
356 NO<sub>2</sub> plume over Seoul and its surroundings at 09:45 LT (Fig. 4, red box; case (A)) was  
357 transported southeast until 14:45 LT, reaching Jecheon, Yeongju, and Uljin, where it was  
358 widespread. These three receptor cities had no substantial NO<sub>x</sub> emission sources, and their NO<sub>2</sub>  
359 column values were low at 09:45 LT on March 22 but substantially higher between 11:45 and  
360 14:45 LT (Supplementary Fig. 13). NO<sub>x</sub> emissions in Jecheon, Yeongju, and Uljin are 12%,  
361 6%, 1% of the emissions in Seoul, respectively. NO<sub>2</sub> was transported from hotspots including  
362 Gwangyang, Busan, and Pohang (southeastern coastline, Korean Peninsula; blue box; case (B),  
363 Fig. 4) to the Korea Strait between South Korea and Japan, with several distinct long NO<sub>2</sub> trails  
364 (Fig. 4, blue box; case (B)). The plume transport directions observed by GEMS were consistent  
365 with the wind directions (based on ECMWF ERA5 data) at 1,000 hPa (Supplementary Fig. 14).  
366 The winds are weak and varies in direction during the morning (10:00–12:00 LT) in the regions  
367 marked as red box. However, in the afternoon (13:00–15:00), the wind blows strongly toward  
368 the east. In addition, winds blow from the southern part of the Korean Peninsula toward  
369 Kagoshima, Japan at speeds greater than 5 m/s. In Fig. 4 (g and h), we included the spatial  
370 distributions of NO<sub>2</sub> at 13:45 and 14:45 to compare the long-range transport event on March  
371 22 with March 23, when no long-range transport occurred.

372         The transported NO<sub>2</sub> contributed to the enhancement at the receptor areas, which led to  
373 substantial hourly changes in the NO<sub>2</sub> spatial distributions over the southern parts of the Korean  
374 Peninsula and its surroundings. The NO<sub>x</sub> emissions estimated based on satellite measurements  
375 can be biased depending on the transported amount and transport frequency in the receptor  
376 areas. It is important to understand the temporal and spatial distribution of NO<sub>2</sub> in each area  
377 because it can help establish an appropriate NO<sub>x</sub> control strategy by correcting the temporally  
378 biased NO<sub>x</sub> levels available at a single time per day in each area.

## 379 **Conclusions**

380           The findings from the GEMS reported in the present study reveal that diurnal  
381 tropospheric NO<sub>2</sub> characteristics vary widely across various source regions. The top-down NO<sub>x</sub>  
382 emissions obtained from tropospheric NO<sub>2</sub> column data derived from twice-a-day LEO  
383 measurements can be improved using GEO observations, such as GEMS. This improvement is  
384 possible because the emissions were calculated based on the assumption that both the chemistry  
385 and the diurnal variation of emissions are accurate, due to the absence of diurnally available  
386 NO<sub>2</sub> spatial distribution information. In addition, hourly tropospheric NO<sub>2</sub> column data, which  
387 can constrain both emissions and chemistry, can be utilized to assess diurnal NO<sub>x</sub> emissions  
388 and improve the current NO<sub>x</sub> emissions estimations from bottom-up emission inventory. In  
389 contrast to previous simple mass balance methods with a steady-state assumption, our study  
390 emphasizes the importance of including the temporal change in GEMS NO<sub>2</sub> columns, which  
391 reflects the carry-over of NO<sub>2</sub> over time in top-down emission estimations. Top-down emission  
392 estimations with non-steady-state assumption show an increase in NO<sub>x</sub> emissions at a given  
393 time in the morning, followed by an increase in NO<sub>2</sub> column at later times. Conversely, a  
394 decrease in NO<sub>x</sub> emissions at noon or in the early afternoon is followed by a decrease in NO<sub>2</sub>  
395 columns at later times. For example, in Seoul, our optimal top-down NO<sub>x</sub> emission peaks  
396 between 9-11 LT, followed by NO<sub>2</sub> columns peaking between 10-12 LT. Meanwhile, the  
397 minimum NO<sub>x</sub> emissions occur at 14 LT, followed by the minimum NO<sub>2</sub> columns at 15 LT.  
398 For most source regions, top-down NO<sub>x</sub> emission estimations based on non-steady-state  
399 assumption utilizing hourly varying GEMS observations results in the most accurate NO<sub>2</sub>  
400 column simulations.

401           This study also highlights the challenges associated with this top-down approach.  
402 Accurate estimations of chemical lifetime, derived from relevant chemical reactions in the

403 chemical transport model, can be difficult, as observed in Hanoi, power plants in India, and the  
404 Strait of Malacca, where large biogenic emissions in the source region or nearby areas rapidly  
405 deplete OH and, thereby increasing the chemical lifetime of NO<sub>x</sub>. In such cases, evaluating the  
406 biogenic emission estimates in the model becomes crucial. Additionally, uncertainties  
407 surrounding nighttime NO<sub>x</sub> emissions and the chemistry influencing early morning NO<sub>2</sub>  
408 columns, which can vary by location, present another critical factor for improving model NO<sub>2</sub>  
409 simulations in the morning. In general, uncertainties in the chemical transport model, stemming  
410 from factors such as anthropogenic and biogenic emissions, chemical mechanisms, spatial  
411 resolution, and other physical processes, need further investigation.

412         The possibility of increasing uncertainties in GEMS tropospheric NO<sub>2</sub> column  
413 retrievals rise with increasing solar zenith angle, such as in the early morning or late afternoon.  
414 Therefore, a more thorough evaluation of GEMS data and model results using independent in-  
415 situ and remote sensing observations from surface and aircraft (e.g., Airborne and Satellite  
416 Investigation of Asian Air Quality, ASIA-AQ, [https://www-air.larc.nasa.gov/missions/asia-  
417 aq/index.html](https://www-air.larc.nasa.gov/missions/asia-aq/index.html)) is necessary to accurately assess diurnal profiles in bottom-up emission  
418 inventories and to develop precise top-down NO<sub>x</sub> emission estimates utilizing GEMS  
419 measurements. This study lays the groundwork and provides direction for future research  
420 efforts.

421

## 422 **Methods**

### 423 **Geostationary Environmental Monitoring Spectrometer**

424         GEMS build on the technology and algorithms of heritage instruments and augments  
425 various applications through unprecedented hourly observations. GEMS is the first UV–visible

426 hyperspectrometer in geostationary orbit and was launched aboard GEO-KOMPSAT-2B on  
427 February 18, 2020. It provides daytime measurements of solar radiance reflected at the surface  
428 and scattered in the atmosphere. Measurements are made in the 300–500 nm range, with a  
429 mean spectral resolution of 0.6 nm in full width at half maximum (FWHM)<sup>20</sup>. The  
430 backscattered solar radiance contains information on Rayleigh and Mie scattering, surface  
431 albedo, and atmospheric trace gas absorption, including that of NO<sub>2</sub>. These measurements  
432 allow the observation of key air quality components, such as NO<sub>2</sub>, sulfur dioxide (SO<sub>2</sub>), ozone  
433 (O<sub>3</sub>), formaldehyde (HCHO), glyoxal (CHOCHO), and aerosols, at a spatial resolution of 3.5  
434 × 7.7 km<sup>2</sup>. GEMS is the first and only space-based instrument to provide daily time-resolved  
435 (hourly) observations of these species. The field of regard of GEMS covers 20 countries in  
436 Asia, extending E–W from Japan to India, and N–S from Mongolia to Indonesia  
437 (Supplementary Fig. 1 and Supplementary Table 1)<sup>20</sup>.

438

### 439 **Retrieval methods**

440 This study introduces new GEMS primary products: diurnal total and tropospheric NO<sub>2</sub>  
441 column variations retrieved in real-time over cities, industrial areas, and ports in Asia. In  
442 addition, diurnal patterns are discussed. The NO<sub>2</sub> slant column density (SCD), i.e., the NO<sub>2</sub>  
443 concentration integrated along the light path through the atmosphere, was retrieved from the  
444 GEMS radiance and irradiance spectra using the differential optical absorption spectroscopy  
445 (DOAS)<sup>46</sup>. Spectral fitting was applied over the wavelength range of 432–450 nm, with the  
446 absorption cross-sections convoluted using the GEMS instrument function (FWHM = 0.6 nm).  
447 The absorption cross sections used were NO<sub>2</sub> (220 K<sup>47</sup>), O<sub>3</sub> (243 and 293 K<sup>48</sup>), O<sub>4</sub><sup>49</sup>, H<sub>2</sub>O<sup>50</sup>,  
448 and ring spectrum<sup>51</sup>. The SCD was converted to the NO<sub>2</sub> column by dividing it by the air mass  
449 factor (AMF; VCD = SCD/AMF). Here, the AMF represents the ratio of the NO<sub>2</sub> column

450 integrated along the light path to the vertical NO<sub>2</sub> column for a selected spectral fitting  
451 window<sup>20</sup>. The AMF was calculated using the linearized pseudo-spherical scalar and vector  
452 discrete ordinate radiative transfer model<sup>52</sup>, *a priori* NO<sub>2</sub> vertical profile shapes from the  
453 regional chemical transport model with the horizontal resolution of 28 x 28 km<sup>2</sup> (WRF-Chem<sup>53</sup>  
454 and the community atmosphere model with chemistry<sup>54</sup>), and ancillary data such as aerosol  
455 optical depth, aerosol effective height, surface reflectance, total ozone column, pressure, and  
456 geometries. Details of the GEMS NO<sub>2</sub> column density retrieval method has been published  
457 elsewhere<sup>20</sup>.

458         The tropospheric NO<sub>2</sub> column density was calculated by subtracting the stratospheric  
459 NO<sub>2</sub> column density from the total NO<sub>2</sub> column density. The stratospheric NO<sub>2</sub> column density  
460 was estimated by scaling the stratospheric NO<sub>2</sub> column density from the single-layer isentropic  
461 model of chemistry and transport (SLIMCAT)<sup>55,56</sup> using GEMS NO<sub>2</sub> column densities  
462 observed over the Pacific Ocean, where negligible tropospheric NO<sub>2</sub> concentrations are  
463 expected. Over polluted areas, total NO<sub>2</sub> column diurnal variations appear to be dominated by  
464 tropospheric column variations. The mean value of the stratospheric NO<sub>2</sub> column over the  
465 GEMS domain, based on SLIMCAT estimation, was  $0.1 \times 10^{16}$  molec cm<sup>-2</sup> (Supplementary  
466 Fig. 15).

467

## 468 **Assessment of GEMS NO<sub>2</sub> data**

### 469 ***GEMS vs. TROPOMI and OMI***

470         GEMS total and tropospheric NO<sub>2</sub> column data were compared with TROPOMI and  
471 OMI data for observations at 04:45 coordinated UTC and 13:45 LT for June–August 2021  
472 (Supplementary Table 6). TROPOMI, aboard the Sentinel-5P satellite, is the most recently

473 launched instrument capable of monitoring NO<sub>2</sub> column density from space with high  
474 temporospatial resolution (up to 5.5 × 3.5 km<sup>2</sup> pixels with daily revisit). Extensive validation  
475 of TROPOMI data has been published<sup>56–58,61</sup>. The correlation coefficients between the GEMS  
476 and TROPOMI (OMI) total NO<sub>2</sub> column densities were 0.78 (0.58), slopes were 1.87 (0.58),  
477 and intercepts were  $-0.36 (0.23) \times 10^{16}$  molec cm<sup>-2</sup>; for the GEMS and TROPOMI (OMI)  
478 tropospheric NO<sub>2</sub> column densities, the correlation coefficients were 0.63 (0.76), slopes were  
479 1.27 (0.83), and intercepts were  $0.06 (0.13) \times 10^{16}$  molec cm<sup>-2</sup> for the period June–August 2021  
480 (Supplementary Table 6). These differences in statistics between the comparisons of GEMS  
481 and OMI, as well as GEMS and TROPOMI, can be influenced by the differences in spatial  
482 resolution of each satellite<sup>62</sup>. In addition, the differences in parameters between the total  
483 columns of GEMS and TROPOMI can be partly attributed to differences in the input data used  
484 to calculate their respective AMFs<sup>20,57,59,60</sup>. Differences between stratospheric and tropospheric  
485 partition algorithms are likely to contribute to differences in the tropospheric columns. GEMS  
486 tropospheric columns were obtained by subtracting SLIMCAT-based stratospheric NO<sub>2</sub> data  
487 from GEMS total columns, whereas those of TROPOMI were obtained using a data  
488 assimilation system<sup>60</sup> based on the TM5-MP model data<sup>63</sup>. Although both the total and  
489 tropospheric NO<sub>2</sub> column data observed from GEMS tended to be higher than those of  
490 TROPOMI, NO<sub>2</sub> column data agreed well<sup>29,31,35</sup> with those of TROPOMI, which have already  
491 been validated using ground-based MAX-DOAS and Pandora measurements<sup>56,58,61</sup>.

492

### 493 ***GEMS vs. ground-based Pandora and MAX-DOAS***

494 GEMS NO<sub>2</sub> data were compared with Pandora and MAX-DOAS data available for the  
495 two metropolitan cities of Seoul (South Korea), Xianghe (China), and a suburban area of  
496 Seosan (South Korea). The ground-based Pandora instrument, which uses direct sunlight as the

497 light source, produces NO<sub>2</sub> total column data<sup>64,65</sup>. Pandora measurements have been used to  
498 evaluate NO<sub>2</sub> column densities retrieved by TROPOMI and OMI<sup>56,58,61,64–68</sup>. GEMS total NO<sub>2</sub>  
499 column data were compared with the Pandonia Global Networks (PGN) NO<sub>2</sub> column data  
500 (<https://pandonia-global-network.org/>, last accessed March 16, 2024) in Seoul for the period  
501 June–August 2021. The correlation coefficient, slope, and intercept against Pandora are 0.78,  
502 0.93, and  $0.13 \times 10^{16}$  molec cm<sup>-2</sup>, respectively; (Supplementary Table 6). The GEMS total NO<sub>2</sub>  
503 column data were also compared with those of Pandora at Seosan, a suburban area of South  
504 Korea, for the first GEMS map of the air pollution campaign period from October 12 to  
505 November 27, 2020. The spectral fitting interval and absorption cross-sections adopted by  
506 PGN<sup>69</sup> were used to retrieve the NO<sub>2</sub> column data from the Pandora measurements during the  
507 campaign. The correlation coefficient, slope, and intercept were 0.74, 0.54, and  $0.37 \times 10^{16}$   
508 molec cm<sup>-2</sup>, respectively; (Supplementary Table 6). The GEMS NO<sub>2</sub> column densities tended  
509 to be lower than those of Pandora in both Seoul megacity and Seosan suburban areas. GEMS  
510 tropospheric NO<sub>2</sub> column data were also compared with those of the multi-axis differential  
511 optical absorption spectroscopy (MAX-DOAS) at Xianghe city<sup>70</sup> for the period June–August  
512 2021. The correlation coefficient, slope, and intercept were 0.51, 0.13, and  $1.42 \times 10^{16}$  molec  
513 cm<sup>-2</sup>, respectively, between GEMS tropospheric NO<sub>2</sub> columns and those of MAX-DOAS  
514 (Supplementary Table 6). This tendency of underestimation in GEMS NO<sub>2</sub> columns can be due  
515 to differences in vertical sensitivities in the lower atmosphere<sup>71</sup> and the relatively large GEMS  
516 footprint<sup>72</sup>. A similar underestimation tendency was also found in the comparisons of  
517 TROPOMI and OMI NO<sub>2</sub> data with those of Pandora and MAX-DOAS<sup>56,58,61,64–74</sup>. GEMS  
518 provides a representative value for an entire pixel, whereas MAX-DOAS measures NO<sub>2</sub>  
519 concentrations at a specific location. This fundamental difference in spatial resolution can lead  
520 to discrepancies in observed NO<sub>2</sub> variations, particularly in highly polluted areas. In such  
521 regions, localized emission sources and meteorological conditions can cause significant within

522 pixel variability, which MAX-DOAS captures but GEMS averages over a broader area. As a  
523 result, this differences in observation may contribute to differences in the timing and magnitude  
524 of NO<sub>2</sub> peaks observed by the two observations<sup>56,71,72</sup>.

525 To understand vertical sensitivities of the GEMS retrieval over the course of the day,  
526 we investigated the discrepancy in either diurnal or measurement geometry variability between  
527 GEMS NO<sub>2</sub> and those measured by MAX-DOAS. Especially, MAX-DOAS is known to have  
528 a high sensitivity to surface NO<sub>2</sub><sup>70,71</sup>. Supplementary Figs. 16-18 show tropospheric NO<sub>2</sub>  
529 column measured from the GEMS and MAX-DOAS observation, SZA and VZA, and aerosol  
530 optical depth (AOD) and cloud fraction (CF) as a function of local time (LT). The diurnal  
531 variations captured in GEMS and MAX-DOAS are quite similar. Absolute difference between  
532 GEMS and MAX-DOAS tropospheric NO<sub>2</sub> columns tend to be large from 8:45 to 10:45 LT  
533 when the SZA is large (from 24.0 to 49.6°) and decreases at noon when the SZA is small (from  
534 13.5 to 19.7°), and then increases again after 14:45 LT when the SZA becomes larger (from  
535 32.5 to 46.1°). CF does not have a significant effect on the difference between tropospheric  
536 NO<sub>2</sub> from GEMS and MAX-DOAS, and AOD is higher in Seoul than in Kasuga, which leads  
537 to the overall difference between GEMS and MAX-DOAS being larger in Seoul than in  
538 Kasuga. These results mean that different geometry conditions and aerosols may affect the  
539 sensitivities to retrieve the surface gases under certain conditions.

540

#### 541 **Bottom-up emission inventory**

542 The bottom-up anthropogenic emissions inventory for Asia used in this study is based  
543 on the Air Quality in Northeast Asia (AQNEA) emission inventory for Northeast Asia and the  
544 Comprehensive Regional Emissions inventory for Atmospheric Transport Experiment  
545 (CREATE) for the rest of Asian domain<sup>75,76</sup>. The AQNEA is a research project studying air

546 quality under various future carbon mitigation and air pollution control strategies using  
547 multiple Integrated Assessment Models (IAMs) from participating regional and international  
548 institutions. Emissions of six countries – China, South Korea, Japan, North Korea, Mongolia,  
549 and Russia – in Northeast Asia were estimated in the AQNEA inventory. The inventory covers  
550 eight criteria air pollutants - CO, SO<sub>2</sub>, NO<sub>x</sub>, non-methane volatile organic compounds  
551 (NMVOCs), NH<sub>3</sub>, PM<sub>10</sub>, PM<sub>2.5</sub> - across eight emission sectors; power generation, industry,  
552 residential, on-road/nonroad mobile, agriculture, solvent use, and waste. Annual emissions in  
553 AQNEA for the year 2019 are as follows: 126.8 Tg for CO, 8.2 Tg for SO<sub>2</sub>, 20.5 Tg for NO<sub>x</sub>,  
554 26.1 Tg NMVOCs, 10.6 Tg for NH<sub>3</sub>, 11.6 Tg for PM<sub>10</sub>, and 8.3 Tg. for PM<sub>2.5</sub>. The inventory  
555 was further processed using the SMOKE-Asia<sup>75</sup> emission processing model to generate spatio-  
556 temporally allocated and chemically speciated emission fields in support of chemical transport  
557 modeling. NO<sub>x</sub> emissions in 2021 were reduced by 20% from AQNEA that represents the year  
558 2019, following the trends in surface and satellite observations in Seoul, South Korea<sup>77</sup>. This  
559 change does not represent whole region of interest.

560

### 561 **Top-down NO<sub>x</sub> emission estimations with a steady-state assumption (TD\_S)**

562 Top-down emissions of NO ( $E$ ) can be estimated from NO<sub>2</sub> column ( $\Omega$ ) by dividing it  
563 by an effective time scale  $\tau_{\text{eff}}$ <sup>78</sup>. In this case, temporal change of column is ignored. The  
564 observed column at a certain time is associated with the emission at the same time. The  
565 effective time scale encapsulates the chemical loss of NO<sub>2</sub>, and the conversion of NO<sub>2</sub> to NO,  
566 depending on the special circumstances of the case. The advection effect is negligible  
567 concerning a mass balance approach over a large source box. This effect may not be negligible  
568 for small source regions or for spring, fall, and winter when NO<sub>x</sub> lifetime is long<sup>59</sup>.

569

$$E = \frac{\Omega}{\tau_{eff}} \quad (1)$$

571

572         The effective time scale  $\tau_{eff}$  can be derived from chemical transport models. In this  
573 study, we utilized the Weather Research Forecast-Chemistry (WRF-Chem) model results with  
574 various emission option to determine  $\tau_{eff}$ . We use the same version and configuration of WRF-  
575 Chem as described in previous study and references therein<sup>79</sup> for July 2021. The bottom-up  
576 emission inventory combines data from AQNEA and CREATE. The model has horizontal  
577 resolution of 28 by 28 km<sup>2</sup> and 60 vertical layers up to 50 hPa. The Whole Atmosphere  
578 Community Climate Model (WACCM) forecast data is used for the chemical boundary/initial  
579 conditions.

580

581

$$\tau_{eff} = \frac{\Omega_{NO2\_WRFChem}}{E_{NO\_BU}} \quad (2)$$

583

584 , where  $\Omega_{NO2\_WRFChem}$  is NO<sub>2</sub> columns calculated from WRF-Chem results and  $E_{NO\_BU}$  is  
585 bottom-up emission inventories used for the model runs. The top-down NO emission,  $E_{NO\_TD}$   
586 is calculated from GEMS tropospheric NO<sub>2</sub> columns,  $\Omega_{NO2\_GEMS}$  and  $\tau_{eff}$  from the chemical  
587 transport model.

588

589 
$$E_{NO\_TD} = \frac{\Omega_{NO2\_GEMS}}{\tau_{eff}} \quad (3)$$

590

591 In this study, we ran WRF-Chem from June 24 to July 31 using four different bottom-  
 592 up emission inventory values to account for uncertainties in the estimated effective lifetimes  
 593 and top-down NO emission (Table 1). Three emissions data used the same daily NO emissions  
 594 but adopt different diurnal variabilities, while one applied the minimum nighttime emission  
 595 value  $E_{NO\_BU\_BG}$  throughout the entire day to estimate background NO<sub>2</sub> column value  
 596  $\Omega_{NO2\_WRFChem\_BG}$ . In practice, in the total GEMS tropospheric NO<sub>2</sub> columns, the deviated part  
 597 from the background NO<sub>2</sub> columns is used to calculate the top-down NO emissions minimizing  
 598 the impacts of background value:

599

600 
$$E_{NO\_TD} = \frac{(\Omega_{NO2\_GEMS} - \Omega_{NO2\_WRFChem\_BG})}{\tau_{eff}} + E_{NO\_BU\_BG} \quad (4)$$

601

602 We estimated four different top-down NO emissions and primarily present the averaged  
 603 value. Details about four bottom-up and top-down emissions, corresponding WRF-Chem  
 604 simulations, and difference between top-down and bottom-up emissions are provided in  
 605 Supplementary Figs. 8, 9 and 12.

606

607 Table 1. Bottom-up emission inventories used for WRF-Chem simulations for development  
 608 of TD\_S.

Bottom-Up (BU) Emission Inventory	Description
--------------------------------------	-------------

(WRF-Chem case name)	
BU1 (WRF-Chem1)	AQNEAV2 constant emissions using nighttime minimum value
BU2 (WRF-Chem2)	AQNEAV2 constant emissions (default value)
BU3 (WRF-Chem3)	AQNEAV2 with diurnal factor from EDGAR emission inventory
BU4 (WRF-Chem4)	AQNEAV2 with diurnal factor from Kim et al. (2016)

609

610 **Top-down NO<sub>x</sub> emission estimations with non-steady-state assumption (TD\_NS)**

611 The chemical lifetime of NO<sub>2</sub> ( $\tau_c$ ) against a reaction with OH during the daytime in  
612 summer is relatively short, but generally exceeds one hour. As a result, the emissions at a given  
613 time can influence NO<sub>2</sub> columns at later times, even in summer. The temporal change term  
614 ( $\frac{\partial \Omega}{\partial t}$ ) helps account for this memory effect when estimating NO<sub>x</sub> emissions based on GEMS  
615 hourly varying observations. With a non-steady-state assumption, emissions can be defined as:

616

$$617 \quad E = \alpha \left( \frac{\partial \Omega}{\partial t} + \frac{\Omega}{\tau_c} \right) \quad (5)$$

618

619 where  $\alpha$  is a conversion factor (NO<sub>x</sub>/NO<sub>2</sub>) that can be obtained from the model results.  $\tau_c$  can  
620 be directly calculated with reaction coefficients and reactant concentrations (e.g. [OH]).  
621 Alternatively, this time scale can be derived from the model results as a residual of the equation,  
622 using known bottom-up emission values, such as:

623

$$624 \quad \tau_r = \frac{\Omega_{NO_2, WRFChem}}{\left( E_{NO\_BU} - \frac{\partial \Omega_{NO_2, WRFChem}}{\partial t} \right)} \quad (6)$$

625

626 Here,  $\tau_r$  is referred to as a residual time scale and is acquired from the WRF-Chem model  
627 results. With known  $\tau_r$  values, the top-down emission is calculated with GEMS tropospheric  
628 NO<sub>2</sub> columns observations at a given time and the previous time, as follows:

629

$$630 \quad E_{\text{NO\_TD}} = \alpha \left( \frac{\partial \Omega_{\text{NO}_2 \text{GEMS}}}{\partial t} + \frac{\Omega_{\text{NO}_2 \text{GEMS}}}{\tau_r} \right) \quad (7)$$

631

632 The temporal change term is calculated as an instantaneous gradient at the overpass time. In  
633 this way, we estimated another set of top-down NO<sub>x</sub> emissions based on the non-steady state  
634 assumption using the equation above and WRF-Chem model results (TD\_NS). In the  
635 Supplementary Table 5,  $\tau_c$  and  $\tau_r$  and OH concentrations averaged from 11 to 13 LT are  
636 summarized for each source.  $\tau_c$  and  $\tau_r$  are comparable for most sources, except for those with  
637 very low OH values such as Hanoi and the Strait of Malacca. For these two regions,  $\tau_c$  is much  
638 larger than  $\tau_r$ . In these cases, using  $\tau_c$  instead of  $\tau_r$  can result in negative top-down NO<sub>x</sub>  
639 emissions. Therefore, in this study, we utilized  $\tau_r$  values for estimating TD\_NS. Even with  $\tau_r$   
640 values, the estimated top-down emissions can still be negative for power plants in India and  
641 the Strait of Malacca. As such, TD\_NS for India P1, P2, and P3 and the Strait of Malacca are  
642 not presented in the main text in this manuscript. The results for these source regions are  
643 summarized in Supplementary Fig. 12.

644

645 **Detection of long-range transport of NO<sub>2</sub> plumes**

646 To determine the event case, we adopted a method from the previous study<sup>44</sup> that  
647 focused on the sulfur dioxide (SO<sub>2</sub>) long-range transport detection and long-range transported  
648 SO<sub>2</sub> flowrate calculation using in-situ observation data, OMI satellite data, and the hybrid  
649 single particle Lagrangian integrated trajectory (HYSPLIT) model. However, in this study, we  
650 used satellite observation data and ECMWF Reanalysis v5 (ERA5) data<sup>45</sup> to detect the NO<sub>2</sub>  
651 long-range transportation event. First, we selected several cities (Jaechon, Yeongju, and Uljin),  
652 located on the east side of Seoul, which have much lower NO<sub>2</sub> column densities than Seoul.  
653 The cities east of Seoul were considered for this study because the westerlies influenced the  
654 Korean Peninsula. Second, we identified the specific times when the NO<sub>2</sub> column densities  
655 increased by more than twice compared to the NO<sub>2</sub> column at the trough. This approach was  
656 used because the daily variability of NO<sub>2</sub> is not significant in the background region. Then, we  
657 found the dates when the peak of NO<sub>2</sub> columns was more than twice the trough of the day NO<sub>2</sub>  
658 column at Jaechon, Yeongju, and Uljin. Lastly, we used ERA5 model wind field data to verify  
659 whether the movement of NO<sub>2</sub> plumes observed in GEMS data aligned with the wind direction  
660 in ERA5, and we selected the one event that was shown clearly visible in the satellite images  
661 as a case study of this study. In this study, we employed a simpler detection method to suggest  
662 the possibility of long-range transport. Therefore, a more precise approach would involve using  
663 backward trajectory models such as the FLEXible PARTicle dispersion (FLEXPART) model  
664 or HYSPLIT for long-range transport detection. Thus, a more detailed study is needed in the  
665 future.

666

## 667 **References**

- 668 1. Menut, L., et al. Impact of realistic hourly emissions profiles on air pollutants  
669 concentrations modelled with CHIMERE. *Atmos. Environ.* **49**, 233–244 (2012). doi:  
670 10.1016/j.atmosenv.2011.11.057.
- 671 2. Stevenson, D. S., et al. Multimodel ensemble simulations of present-day and near-  
672 future tropospheric ozone. *J. Geophys. Res. Atmos.* **111**, D08301 (2006). doi:  
673 10.1029/2005JD006338.
- 674 3. Goldberg, D. L., et al. A top-down assessment using OMI NO<sub>2</sub> suggests an  
675 underestimate in the NO<sub>x</sub> emissions inventory in Seoul, South Korea, during KORUS-AQ.  
676 *Atmos. Chem. Phys.* **19**(3), 1801–1818 (2019). doi: 10.5194/acp-19-1801-2019.
- 677 4. Qu, Z., et al. Impacts of global NO<sub>x</sub> inversions on NO<sub>2</sub> and ozone simulations. *Atmos.*  
678 *Chem. Phys.* **20**(21), 13109–13130 (2020). doi: 10.5194/acp-20-13109-2020.
- 679 5. Fioletov, V., et al. Quantifying urban, industrial, and background changes in NO<sub>2</sub>  
680 during the COVID-19 lockdown period based on TROPOMI satellite observations. *Atmos.*  
681 *Chem. Phys.* **22**(6), 4201–4236 (2022). doi: 10.5194/acp-22-4201-2022.
- 682 6. Sicard, P., et al. Trends in urban air pollution over the last two decades: A global  
683 perspective. *Sci. Total Environ.* **858**, 160064 (2023). doi: 10.1016/j.scitotenv.2022.160064.
- 684 7. Georgoulias, A. K., et al. Trends and trend reversal detection in 2 decades of  
685 tropospheric NO<sub>2</sub> satellite observations. *Atmos. Chem. Phys.* **19**(9), 6269–6294 (2019). doi:  
686 10.5194/acp-19-6269-2019.
- 687 8. Krotkov, N. A., et al. Aura OMI observations of regional SO<sub>2</sub> and NO<sub>2</sub> pollution  
688 changes from 2005 to 2015. *Atmos. Chem. Phys.* **16**(7), 4605–4629 (2016). doi: 10.5194/acp-  
689 16-4605-2016.
- 690 9. Cooper, M. J., et al. Global fine-scale changes in ambient NO<sub>2</sub> during COVID-19  
691 lockdowns. *Nature* **601**(7893), 380–387 (2022). doi: 10.1038/s41586-021-04229-0.

- 692 10. Boersma, K. F., et al. Improving algorithms and uncertainty estimates for satellite NO<sub>2</sub>  
693 retrievals: results from the quality assurance for the essential climate variables (QA4ECV)  
694 project. *Atmos. Meas. Tech.* **11**(12), 6651–6678 (2018). doi: 10.5194/amt-11-6651-2018.
- 695 11. Lamsal, L. N., et al. Ozone Monitoring Instrument (OMI) Aura nitrogen dioxide  
696 standard product version 4.0 with improved surface and cloud treatments. *Atmos. Meas.*  
697 *Tech.* **14**(1), 455–479 (2021). doi: 10.5194/amt-14-455-2021.
- 698 12. Burrows, J. P., et al. The global ozone monitoring experiment (GOME): Mission  
699 concept and first scientific results. *J. Atmos. Sci.* **56**(2), 151–175 (1999). doi: 10.1175/1520-  
700 0469(1999)056<0151:TGOMEG>2.0.CO;2.
- 701 13. Callies, J., et al. GOME-2-Metop’s second-generation sensor for operational ozone  
702 monitoring. *ESA bull.* **102**, 28–36 (2000).
- 703 14. Bovensmann, H., et al. SCIAMACHY: Mission objectives and measurement modes. *J.*  
704 *Atmos. Sci.* **56**(2), 127–150, (1999). doi: 10.1175/1520-  
705 0469(1999)056<0127:SMOAMM>2.0.CO;2.
- 706 15. Levelt, P. F., et al. The ozone monitoring instrument. *IEEE Trans. Geosci. Remote Sens.*  
707 **44**(5), 1093–1101 (2006). doi: 10.1109/TGRS.2006.872333.
- 708 16. Levelt, P. F., et al. The Ozone Monitoring Instrument: overview of 14 years in  
709 space. *Atmos. Chem. Phys.* **18**(8), 5699–5745 (2018). doi: 10.5194/acp-18-5699-2018.
- 710 17. Veefkind, J. P., et al. TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission  
711 for global observations of the atmospheric composition for climate, air quality and ozone layer  
712 applications. *Remote Sens. Environ.* **120**, 70–83 (2012). doi: 10.1016/j.rse.2011.09.027.
- 713 18. Richter, A., et al. Increase in tropospheric nitrogen dioxide over China observed from  
714 space. *Nature* **437**(7055), 129–132 (2005). doi: 10.1038/nature04092.
- 715 19. A geostationary satellite constellation for observing global air quality: An international  
716 path forward. CEOS Atmospheric Composition Constellation, Version 4, 41.

717 [https://ceos.org/observations/documents/AC-VC\\_Geostationary-Cx-for-Global-AQ-](https://ceos.org/observations/documents/AC-VC_Geostationary-Cx-for-Global-AQ-)  
718 [final\\_Apr2011.pdf](#) (2011).

719 20. Kim, J., et al. New era of air quality monitoring from space: Geostationary Environment  
720 Monitoring Spectrometer (GEMS). *Bull. Am. Meteorol. Soc.* **101**(1), E1–E22 (2020). doi:  
721 10.1175/BAMS-D-18-0013.1.

722 21. Zoogman, P., et al. Tropospheric emissions: Monitoring of pollution (TEMPO). *J.*  
723 *Quant. Spectrosc. Radiat. Transf.* **186**, 17–39 (2017). doi: 10.1016/j.jqsrt.2016.05.008.

724 22. Fishman, J., et al. The United States’ next generation of atmospheric composition and  
725 coastal ecosystem measurements: NASA’s GEO stationary Coastal and Air Pollution Events  
726 (GEO-CAPE) mission. *Bull. Am. Meteorol. Soc.* **93**, 1547–1566 (2012). doi: 10.1175/BAMS-  
727 D-11-00201.1.

728 23. Lin, J-T., et al. Constraint of anthropogenic NO<sub>x</sub> emissions in China from different  
729 sectors: a new methodology using multiple satellite retrievals *Atmos. Chem. Phys.*, **10**, 63–78  
730 (2010), doi: 10.5194/acp-10-63-2010.

731 24. Knepp, T., et al. Estimating surface NO<sub>2</sub> and SO<sub>2</sub> mixing ratios from fast-response total  
732 column observations and potential application to geostationary missions, *J. Atmos. Chem.*, **72**,  
733 261–286, (2015), doi: 10.1007/s10874-013-9257-6.

734 25. Judd, L. M., et al. The Dawn of Geostationary Air Quality Monitoring: Case Studies  
735 From Seoul and Los Angeles, *Front. Environ. Sci.*, **6**, (2018), doi: 10.3389/fenvs.2018.00085.

736 26. Li, J., et al., Comprehensive evaluations of diurnal NO<sub>2</sub> measurements during  
737 DISCOVER-AQ 2011: Effects of resolution-dependent representation of NO<sub>x</sub> emissions,  
738 *Atmos. Chem. Phys.*, **21**, 11133-11160, (2021), doi: 10.5194/acp-21-11133-2021.

739 27. Boersma, K. F., et al. Intercomparison of SCIAMACHY and OMI tropospheric NO<sub>2</sub>  
740 columns: Observing the diurnal evolution of chemistry and emissions from space. *J. Geophys.*  
741 *Res. Atmos.* **113**(D16) (2008). doi: 10.1029/2007JD008816.

- 742 28. Boersma, K. F., et al. Validation of urban NO<sub>2</sub> concentrations and their diurnal and  
743 seasonal variations observed from the SCIAMACHY and OMI sensors using in situ surface  
744 measurements in Israeli cities, *Atmos. Chem. Phys.*, **9**, 3867–3879, (2009), doi: 10.5194/acp-  
745 9-3867-2009.
- 746 29. Edwards, D. P., et al. Quantifying the diurnal variation in atmospheric NO<sub>2</sub> from  
747 Geostationary Environment Monitoring Spectrometer (GEMS) observations. *Atmos. Chem.*  
748 *Phys.* **24**(15), 8943–8961 (2024). doi: 10.5194/acp-24-8943-2024.
- 749 30. Yang, L. H., et al. Interpreting Geostationary Environment Monitoring Spectrometer  
750 (GEMS) geostationary satellite observations of the diurnal variation in nitrogen dioxide (NO<sub>2</sub>)  
751 over East Asia. *Atmos. Chem. Phys.* **24**(12), 7027–7039 (2024). doi: 10.5194/acp-24-7027-  
752 2024.
- 753 31. Oak, Y. J., et al. A bias-corrected GEMS geostationary satellite product for nitrogen  
754 dioxide using machine learning to enforce consistency with the TROPOMI satellite instrument.  
755 *Atmos. Meas. Tech.* **17**(17), 5147–5159 (2024). doi: 10.5194/amt-17-5147-2024.
- 756 32. Ahmad, N., et al. Estimation of ground-level NO<sub>2</sub> and its spatiotemporal variations in  
757 China using GEMS measurements and a nested machine learning model. *Atmos. Chem.*  
758 *Phys.* **24**(16), 9645–9665 (2024). doi: 10.5194/acp-24-9645-2024.
- 759 33. He, Q., et al. Quantifying uncertainty in ML-derived atmosphere remote sensing:  
760 Hourly surface NO<sub>2</sub> estimation with GEMS. *Geophys. Res. Lett.* **51**(18), e2024GL110468  
761 (2024). doi: 10.1029/2024GL110468.
- 762 34. Stavrou, T., et al. Satellite evidence for changes in the NO<sub>2</sub> weekly cycle over large  
763 cities. *Sci Rep* **10**, 10066 (2020). doi: 10.1038/s41598-020-66891-0.

- 764 35. Lange, K., et al. Validation of GEMS tropospheric NO<sub>2</sub> columns and their diurnal  
765 variation with ground-based DOAS measurements. *Atmos. Meas. Tech.* **17**(21), 6315-6344  
766 (2024). doi: 10.5194/amt-17-6315-2024.
- 767 36. Koene, E. F. M., et al. On the theory of the divergence method for quantifying source  
768 emissions from satellite observations. *J. Geophys. Res. Atmos.* **129**(12), e2023JD039904  
769 (2024). doi: 10.1029/2023JD039904.
- 770 37. Cifuentes, F., et al. Accurate space-based NO<sub>x</sub> emission estimates with the flux  
771 divergence approach require fine-scale model information on local oxidation chemistry and  
772 profile shapes. *Geosci. Model Dev.* **18**(3), 621-649 (2025). doi: 10.5194/gmd-18-621-2025.
- 773 38. Lelieveld, J., et al. On the role of hydroxyl radicals in the self-cleansing capacity of the  
774 troposphere. *Atmos. Chem, Phys.* **4**(9/10), 2337-2344 (2004). doi: 10.5194/acp-4-2337-2004.
- 775 39. Kim, S. W., et al. Impact of turbulent mixing on isoprene chemistry. *Geophys. Res. Lett.*  
776 **43**(14), 7701-7708 (2016). doi: 10.1002/2016GL069752.
- 777 40. De Smedt, I., et al. CCI+P HCHO tropospheric column L3 data from TROPOMI, v1  
778 (Version 1) [Data set], Royal Belgian Institute for Space Aeronomy (2024).  
779 <https://doi.org/10.18758/2IMQEZ32>.
- 780 41. Copernicus Sentinel-5P (processed by ESA), 2018, TROPOMI Level 2 Formaldehyde  
781 Total Column products. Version 01. European Space Agency (2018).  
782 <https://doi.org/10.5270/S5P-tjlxfd2>.
- 783 42. Qu, Y., et al. An overview of emissions of SO<sub>2</sub> and NO<sub>x</sub> and the long-range transport  
784 of oxidized sulfur and nitrogen pollutants in East Asia. *J. Environ. Sci.* **44**, 13-25 (2016). doi:  
785 10.1016/j.jes.2015.08.028.

- 786 43. Jeong, U., et al. Assessing the effect of long-range pollutant transportation on air quality  
787 in Seoul using the conditional potential source contribution function method. *Atmos. Environ.*  
788 **150**, 33-44 (2017). doi: 10.1016/j.atmosenv.2016.11.017.
- 789 44. Park, J., et al. Long-range transport of SO<sub>2</sub> from continental Asia to northeast Asia and  
790 the northwest Pacific ocean: flow rate estimation using OMI data, surface in situ data, and the  
791 HYSPLIT model. *Atmosphere*, **7**(4), 53 (2016). doi: 10.3390/atmos7040053.
- 792 45. Hersbach, H., et al. ERA5 hourly data on single levels from 1940 to present, Copernicus  
793 Climate Change Service (C3S) Climate Data Store, 2023. doi: 10.24381/cds.adbb2d47.
- 794 46. Platt, U. & Stutz, J. *Differential Optical Absorption Spectroscopy: Principles and*  
795 *Applications*, Springer, Berlin, Heidelberg, 135–174 (2008).
- 796 47. Vandaele, A. C., et al. High-resolution Fourier transform measurement of the NO<sub>2</sub>  
797 visible and near-infrared absorption cross sections: Temperature and pressure effects. *J.*  
798 *Geophys. Res.* **107**(D18), ACH-3 (2002). doi: 10.1029/2001JD000971.
- 799 48. Bogumil, K., et al. Measurements of molecular absorption spectra with the  
800 SCIAMACHY pre-flight model: instrument characterization and reference data for  
801 atmospheric remote-sensing in the 230–2380 nm region. *J. Photochem. Photobiol., A.* **157**(2–  
802 3), 167–184 (2003). doi: 10.1016/S1010-6030(03)00062-5.
- 803 49. Thalman, R. & Volkamer, R. Temperature dependent absorption cross-sections of O<sub>2</sub>–  
804 O<sub>2</sub> collision pairs between 340 and 630 nm and at atmospherically relevant pressure. *Phys.*  
805 *Chem. Chem. Phys.* **15**(37), 15371–15381 (2013). doi: 10.1039/C3CP50968K.
- 806 50. Rothman, L. S., et al. The HITRAN 2008 molecular spectroscopic database. *J. Quant.*  
807 *Spectrosc. Radiat. Transf.* **110**(9–10), 533–572 (2009). doi: 10.1016/j.jqsrt.2009.02.013.
- 808 51. Chance, K. V. & Spurr, R. J. D. Ring effect studies: Rayleigh scattering, including

809 molecular parameters for rotational Raman scattering, and the Fraunhofer spectrum. *Appl.*  
810 *Opt.* **36**(21), 5224–5230 (1997). doi: 10.1364/AO.36.005224.

811 52. Spurr, R. & Christi, M. On the generation of atmospheric property Jacobians from the  
812 (V) LIDORT linearized radiative transfer models. *J. Quant. Spectrosc. Radiat. Transf.* **142**,  
813 109–115 (2014). doi: 10.1016/j.jqsrt.2014.03.011.

814 53. Grell, G. A., et al. Fully coupled “online” chemistry within the WRF model. *Atmos.*  
815 *Environ.* **39**(37), 6957–6975 (2005). doi: 10.1016/j.atmosenv.2005.04.027.

816 54. Emmons, L. K., et al. The chemistry mechanism in the community Earth system model  
817 version2 (CESM2). *J. Adv. Model. Earth Syst.* **12**(4), e2019MS001882 (2020). doi:  
818 10.1029/2019MS001882.

819 55. Chipperfield, M. P. New version of the TOMCAT/SLIMCAT off-line chemical  
820 transport model: Intercomparison of stratospheric tracer experiments. *Q. J. R. Meteorol.*  
821 *Soc.*, **132**(617), 1179–1203 (2006). doi: 10.1256/qj.05.51.

822 56. Verhoelst, T., et al. Ground-based validation of the Copernicus Sentinel-5p TROPOMI  
823 NO<sub>2</sub> measurements with the NDACC ZSL-DOAS, MAX-DOAS and Pandonia global  
824 networks. *Atmos. Meas. Tech.*, **14**(1), 481–510 (2021). doi: 10.5194/amt-14-481-2021.

825 57. Van Geffen, J., et al. S5P TROPOMI NO<sub>2</sub> slant column retrieval: Method, stability,  
826 uncertainties and comparisons with OMI. *Atmos. Meas. Tech.* **13**(3), 1315–1335 (2020). doi:  
827 10.5194/amt-13-1315-2020.

828 58. Ialongo, I., et al. Comparison of TROPOMI/Sentinel-5 Precursor NO<sub>2</sub> observations  
829 with ground-based measurements in Helsinki. *Atmos. Meas. Tech.* **13**(1), 205–218. (2020). doi:  
830 10.5194/amt-13-205-2020.

831 59. Seo, S., et al. Diurnal variations of NO<sub>2</sub> tropospheric vertical column density over the  
832 Seoul metropolitan area from the Geostationary Environment Monitoring Spectrometer

833 (GEMS): seasonal differences and the influence of the a priori NO<sub>2</sub> profile. *Atmos. Meas.*  
834 *Tech.* **18**(1), 115-128 (2025). doi: 10.5194/amt-18-115-2025.

835 60. van Geffen, J. H. G. M., et al. TROPOMI ATBD of the total and tropospheric NO<sub>2</sub> data  
836 products, Tech. Rep. S5P-KNMI-L2-0005-RP, Koninklijk Nederlands Meteorologisch  
837 Instituut (KNMI), CI-7430-ATBD, issue 1.4.0 (2019).  
838 [https://sentinel.esa.int/documents/247904/2476257/sentinel-5p-tropomi-atbd-no2-data-](https://sentinel.esa.int/documents/247904/2476257/sentinel-5p-tropomi-atbd-no2-data-products)  
839 [products](https://sentinel.esa.int/documents/247904/2476257/sentinel-5p-tropomi-atbd-no2-data-products)

840 61. Lambert, J., et al. *Quarterly Validation Report of the Copernicus Sentinel-5 Precursor*  
841 *Operational Data Products# 06: April 2018–February 2020*. Belgian Institute for Space  
842 Aeronomy: Uccle, Belgium (2020).

843 62. Judd, L. M., et al. Evaluating the impact of spatial resolution on tropospheric NO<sub>2</sub>  
844 column comparisons within urban areas using high-resolution airborne data. *Atmos. Meas.*  
845 *Tech.* **12**(11), 6091-6111 (2019). doi: 10.5194/amt-12-6091-2019.

846 63. Williams, J. E., et al. The high-resolution version of TM5-MP for optimized satellite  
847 retrievals: description and validation. *Geosci. Model Dev.* **10**, 721–750 (2017). doi:  
848 10.5194/gmd-10-721-2017.

849 64. Herman, J., et al. NO<sub>2</sub> column amounts from ground-based Pandora and MFDOAS  
850 spectrometers using the direct-Sun DOAS technique: Intercomparisons and application to OMI  
851 validation. *J. Geophys. Res. Atmos.* **114**(D13) (2009). doi: 10.1029/2009JD011848.

852 65. Herman, J., et al. Underestimation of column NO<sub>2</sub> amounts from the OMI satellite  
853 compared to diurnally varying ground-based retrievals from multiple PANDORA spectrometer  
854 instruments. *Atmos. Meas. Tech.* **12**(10), 5593–5612 (2019). doi: 10.5194/amt-12-5593-2019.

855 66. Zhao, X., et al. Assessment of the quality of TROPOMI high-spatial-resolution NO<sub>2</sub>  
856 data products in the Greater Toronto Area. *Atmos. Meas. Tech.* **13**(4), 2131–2159 (2020). doi:  
857 10.5194/amt-13-2131-2020.

- 858 67. Judd, L. M., et al. Evaluating Sentinel-5P TROPOMI tropospheric NO<sub>2</sub> column  
859 densities with airborne and Pandora spectrometers near New York City and Long Island  
860 Sound. *Atmos. Meas. Tech.* **13**(11), 6113–6140 (2020). doi: 10.5194/amt-13-6113-2020.
- 861 68. Reed, A. J., et al. Effects of local meteorology and aerosols on ozone and nitrogen  
862 dioxide retrievals from OMI and pandora spectrometers in Maryland, USA during  
863 DISCOVER-AQ 2011. *J. Atmos. Chem.* **72**(3), 455–482 (2015). doi: 10.1007/s10874-013-  
864 9254-9.
- 865 69. Cede, A., et al. (2019). TN on PGN products “correct use” guidelines, Pandonia Global  
866 Network, available at: [https://www.pandonia-global-network.org/wp-](https://www.pandonia-global-network.org/wp-content/uploads/2020/01/LuftBlick_FRM4AQ_PGNUUserGuidelines_RP_2019009_v1.pdf)  
867 [content/uploads/2020/01/LuftBlick\\_FRM4AQ\\_PGNUUserGuidelines\\_RP\\_2019009\\_v1.pdf](https://www.pandonia-global-network.org/wp-content/uploads/2020/01/LuftBlick_FRM4AQ_PGNUUserGuidelines_RP_2019009_v1.pdf) (la  
868 st access: 22 April 2022).
- 869 70. Hendrick, F., et al. Four years of ground-based MAX-DOAS observations of HONO  
870 and NO<sub>2</sub> in the Beijing area. *Atmos. Chem. Phys.* **14**(2), 765–781 (2014). doi: 10.5194/acp-14-  
871 765-2014.
- 872 71. Wagner, T., et al. Inversion of tropospheric profiles of aerosol extinction and HCHO  
873 and NO<sub>2</sub> mixing ratios from MAX-DOAS observations in Milano during the summer of 2003  
874 and comparison with independent data sets. *Atmos. Meas. Tech.* **4**(12), 2685-2715 (2011). doi:  
875 10.5194/amt-4-2685-2011.
- 876 72. Chan, K. L., et al. MAX-DOAS measurements of tropospheric NO<sub>2</sub> and HCHO in  
877 Munich and the comparison to OMI and TROPOMI satellite observations. *Atmos. Meas. Tech.*  
878 **13**(8), 4499-4520 (2020). doi: 10.5194/amt-13-4499-2020.
- 879 73. Wang, C., et al. Comparison and Validation of TROPOMI and OMI NO<sub>2</sub> Observations  
880 over China. *Atmosphere*. **11**(6), 636 (2020). doi: 10.3390/atmos11060636.
- 881 74. Park, J. U., et al. Spatiotemporal inhomogeneity of total column NO<sub>2</sub> in a polluted urban  
882 area inferred from TROPOMI and Pandora intercomparisons. *GIScience & Remote Sens.* **59**(1),

- 883 354–373 (2022). doi: 10.1080/15481603.2022.2026640.
- 884 75. Woo, J.-H., et al. Development of an anthropogenic emissions processing system for  
885 Asia using SMOKE. *Atmos. Environ.* **58**, 5-13 (2012). doi: 10.1016/j.atmosenv.2011.10.042.
- 886 76. Woo, J.-H., et al. Development of a global integrated assessment model for climate-  
887 air quality management focused on Northeast Asia (GUIDE-Global): A Framework Design.  
888 <https://pure.iiasa.ac.at/id/eprint/19501/> (2023).
- 889 77. Seo, S., et al. Reductions in NO<sub>2</sub> concentrations in Seoul, South Korea detected from  
890 space and ground-based monitors prior to and during the COVID-19 pandemic. *Environ. Res.*  
891 *Commun.* **3**, 051005 (2021). doi: 10.1088/2515-7620/abed92.
- 892 78. Martin, R. V., et al. Global inventory of nitrogen oxide emissions constrained by  
893 space-based observations of NO<sub>2</sub> columns. *J. Geophys. Res. Atmos.* **108** (2003). doi:  
894 10.1029/2003JD003453.
- 895 79. Kim, K.-M., et al. Sensitivity of the WRF-Chem v4. 4 simulations of ozone and  
896 formaldehyde and their precursors to multiple bottom-up emission inventories over East Asia  
897 during the KORUS-AQ 2016 field campaign. *Geosci. Model Dev.* **17**, 1931-1955 (2024). doi:  
898 10.5194/gmd-17-1931-2024.

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957 Contributions

958 J.P., H.H., H.L., and S.W.K. designed the study. J.P., H.H., H.L., S.W.K., J.K., M.V.R.,

959 D.J.J., T.W., A.R., N.A.K., and L.N.L. contributed to writing the original manuscript. M.V.R.

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962 D.J.J., S.W.K., S.S., and K.M.K. contributed to the interpretation of chemistry and model

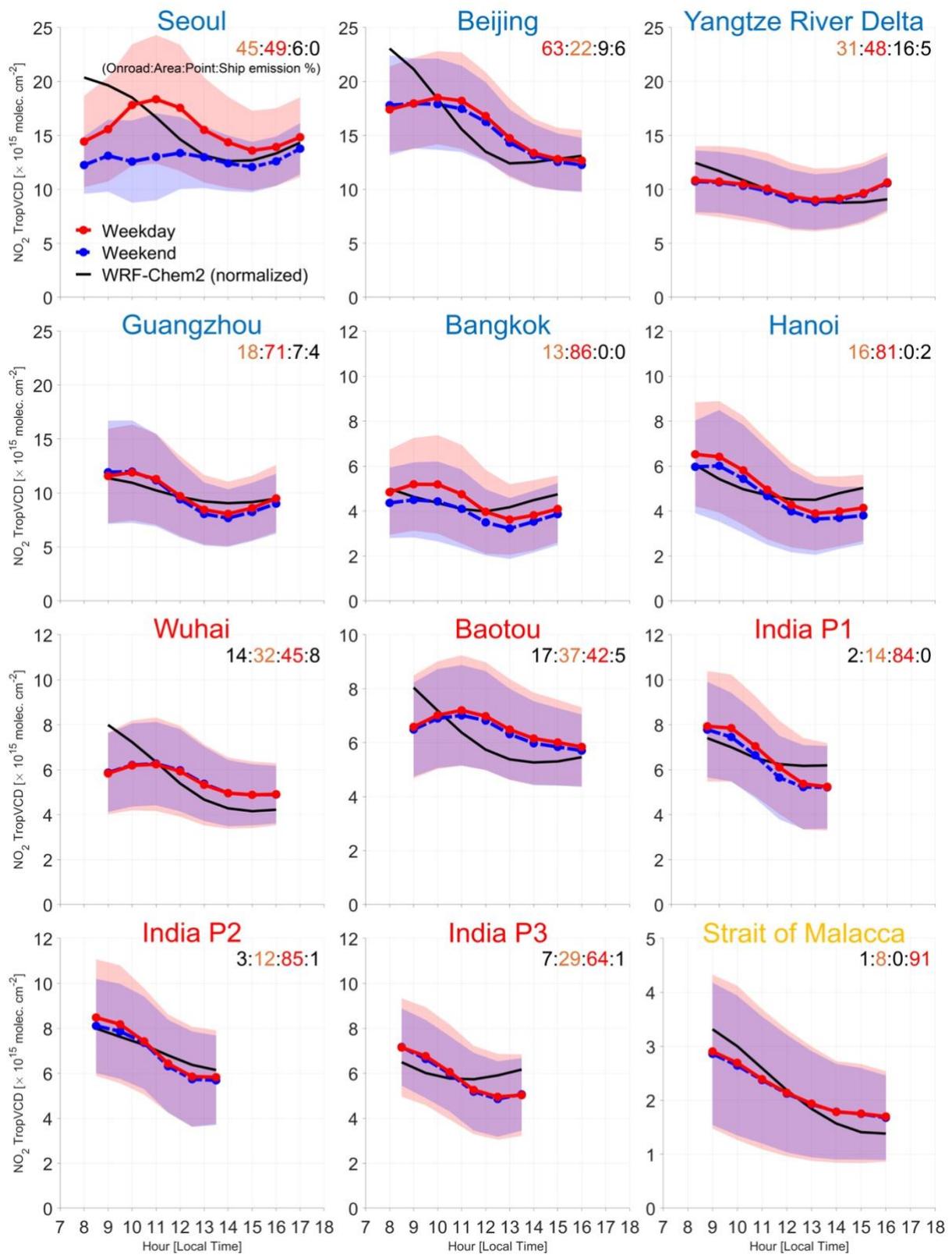
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966 the review, editing, and improvement of the manuscript.

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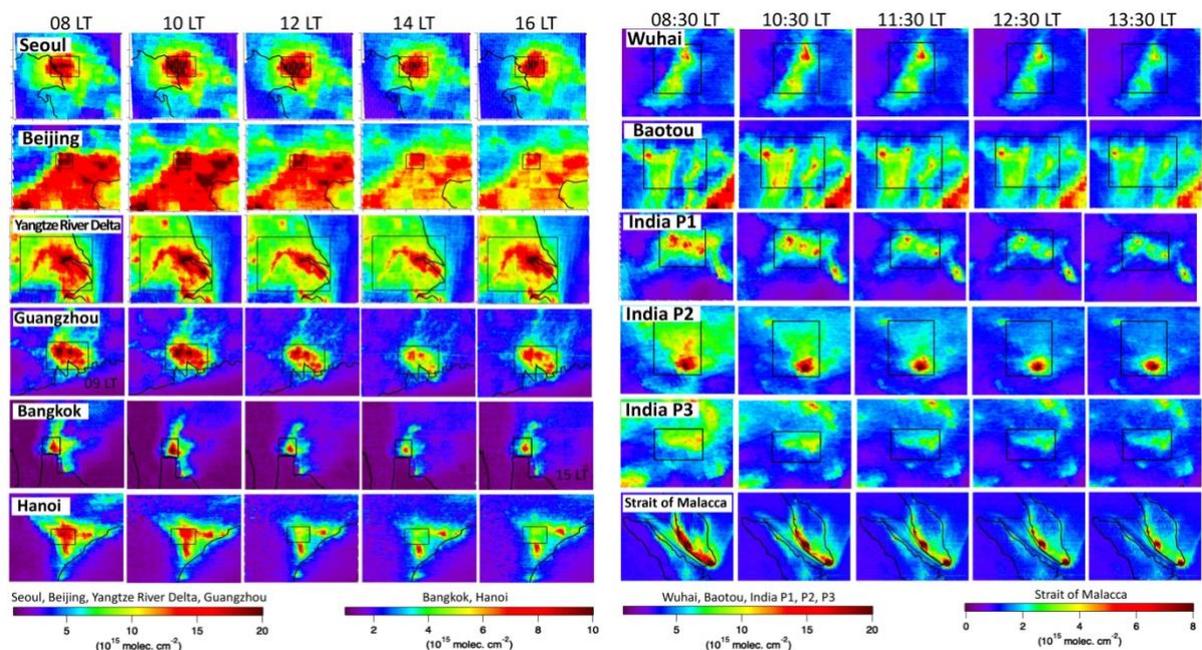


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970 **Fig. 1. Hourly variations in GEMS tropospheric NO<sub>2</sub> columns over major source regions**

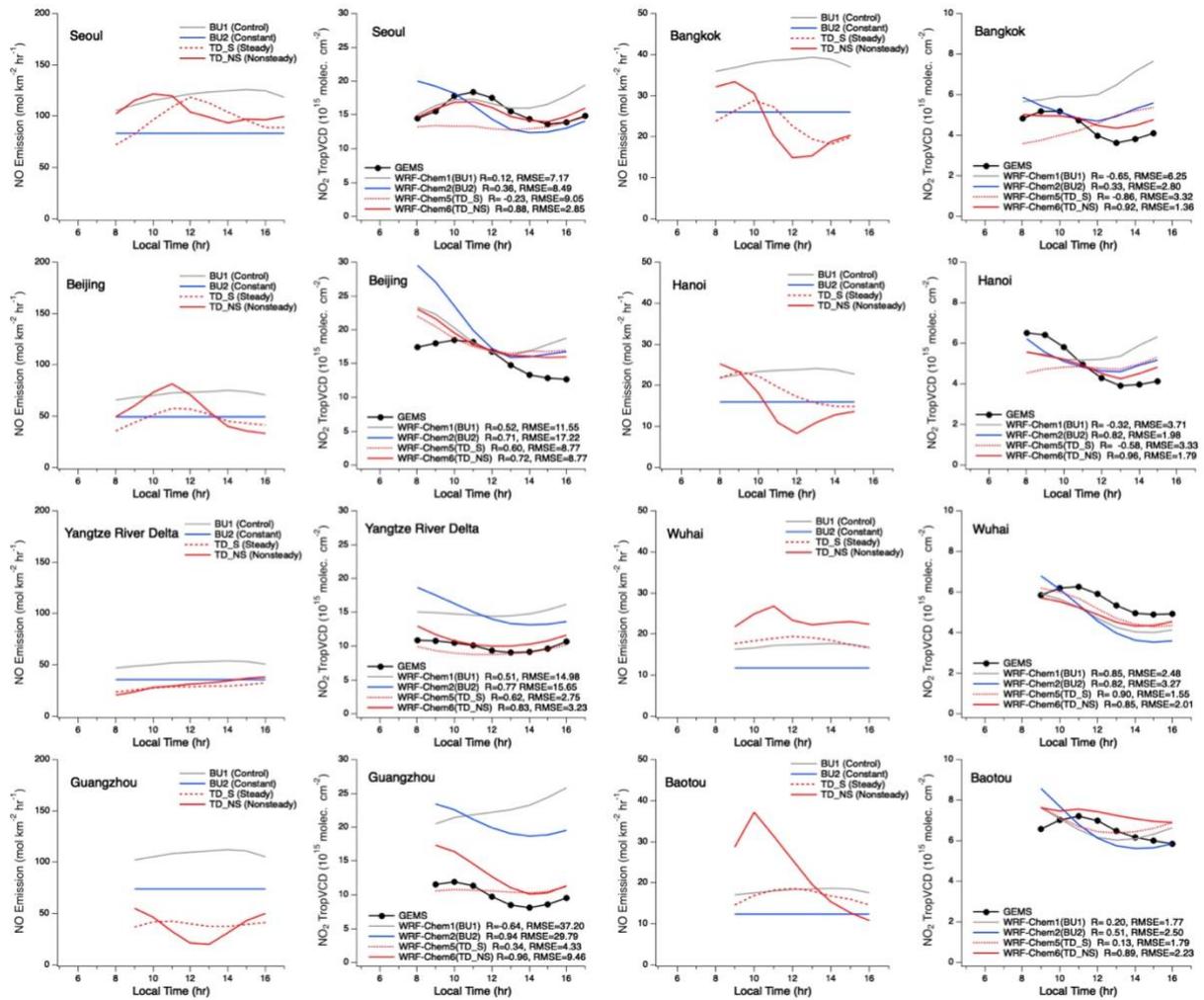
971 **during the study period (June, July, August 2021 and 2022). One standard deviation is**

972 shaded. Red (blue) lines show weekday (weekend) values. Black line is NO<sub>2</sub> columns from  
973 WRF-Chem using a constant emission over time that are normalized to have the same daytime  
974 mean value as GEMS. Blue, red, and yellow labeled regions denote urban, power plant, and  
975 ship track sources, respectively. The source attribution from the bottom-up emission inventory  
976 is shown in top right corner. Contributions from on road, area, point, and ship emissions to the  
977 total NO<sub>x</sub> emission (%) as in bottom-up emission inventory are provided. The largest and the  
978 second largest sectors are noted in red and orange, respectively. Here, area source represents  
979 mainly nonroad (offroad) and residential sources.  
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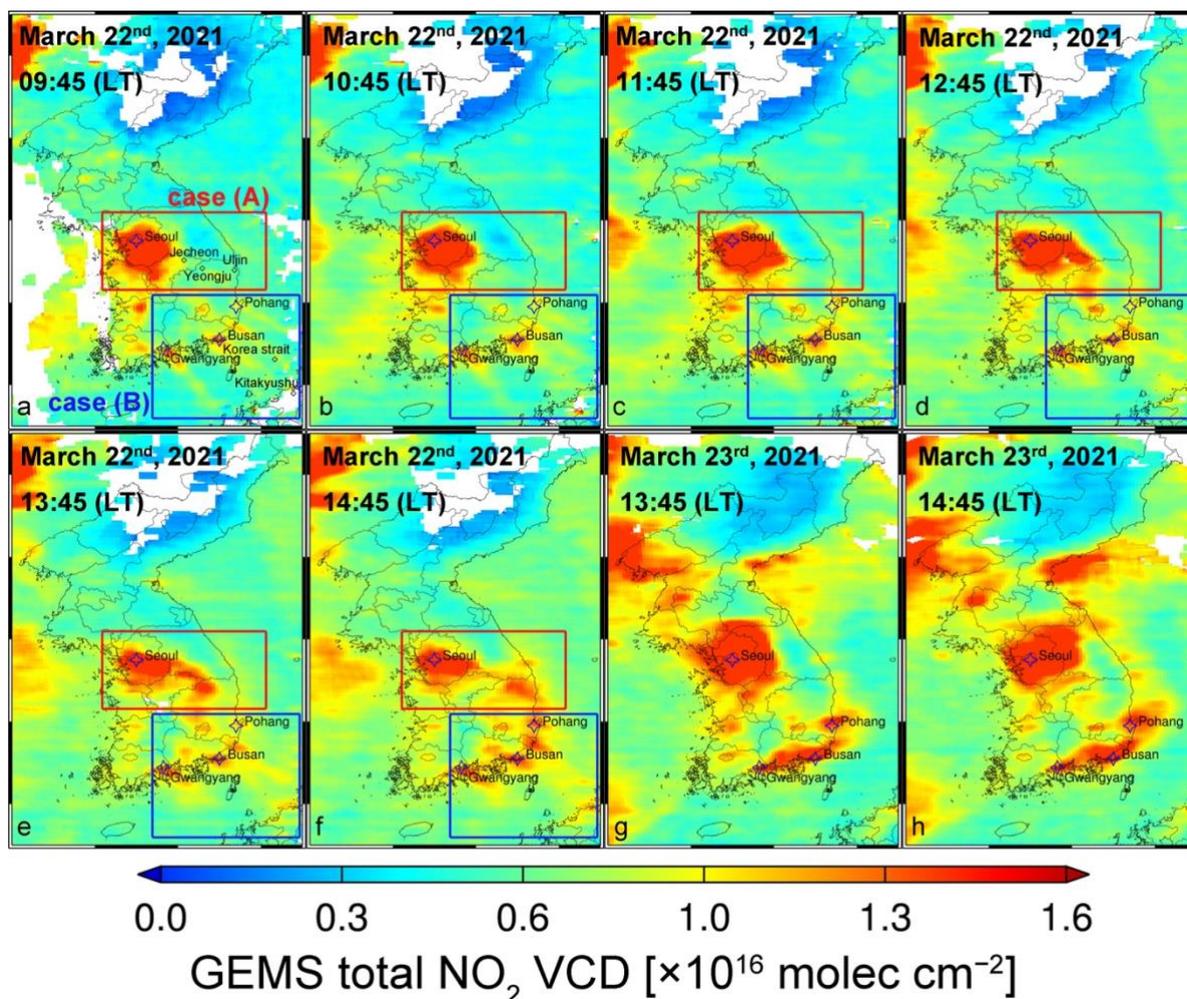
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**Fig. 2. Hourly spatial distribution of average GEMS tropospheric NO<sub>2</sub> columns over major source regions.** The observation times are indicated on the top of the plots. Different times are noted in black on the map of Guangzhou (09 LT instead of 08 LT) and Bangkok (15 LT instead of 16 LT) according to the GEMS observation schedule. The latitude and longitude information for each map is provided in Supplementary Data. Approximate local times are shown for each source. The source boxes used for averaging data are exhibited with black rectangles.



990

991 **Fig. 3. Hourly variations in bottom-up and top-down NO emissions and tropospheric NO<sub>2</sub>**  
 992 **columns from GEMS and model.** Hourly variations in bottom-up and top-down NO  
 993 emissions (left) and GEMS tropospheric NO<sub>2</sub> columns and model NO<sub>2</sub> columns (right) over  
 994 major source regions during weekdays, in the study period. Correlation coefficient and RMSE  
 995 values are shown on the right panel. The unit of RMSE is 10<sup>15</sup> molec cm<sup>-2</sup>.  
 996



997  
 998 **Fig. 4. Transport of NO<sub>2</sub>, revealed by GEMS hourly column analysis over the Korean**  
 999 **Peninsula, Korea Strait, and Japan. a–f: Hourly GEMS NO<sub>2</sub> observations for 09:45–14:45**  
 1000 **LT, March 22, 2021. g and h: GEMS NO<sub>2</sub> observations for 13:45–14:45 LT, March 23, 2021**  
 1001 **(these panels represent a day without NO<sub>2</sub> long-range transport). Red and blue boxes indicated**  
 1002 **the regions for NO<sub>2</sub> transport case (A) and case (B), respectively.**