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**SURFACE WAVE TOMOGRAPHY**

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Surface waves are now being used by several groups to map lateral heterogeneity (Nakanishi and Anderson, 1982, 1983, 1984a,b, Woodhouse and Dziewonski, 1984) and anisotropy (Tanimoto and Anderson, 1984a,b, Nataf, Nakanishi and Anderson, 1984) of the upper mantle on a global basis. The method involves measuring the phase and/or group velocity over hundreds of small arcs and long arcs connecting earthquakes and seismic stations. These averages are then converted to three-dimensional images of the seismic velocity structure and, hence, this is a form of tomography. The large amount of data processing required is made feasible by the existence of long-period digital seismic networks including IDA (International Deployment of Accelerometers), SRO (Seismic Research Observatories) and GDSN (Global Digital Seismic Network). These instruments are operated by a variety of university and government groups including U.C. San Diego, U.S.G.S., D.A.R.P.A. and D.O.E. with the cooperation of many countries. The global coverage is still very sparse compared to the analog W.W.S.S.N. (World Wide Standardized Seismic Network) but preliminary results are very encouraging. The possibility of an expanded global digital network of broad-band seismic stations is now being pursued actively by the United States and several other countries. Because of the sparseness of the present network, mantle structure can only be mapped with fairly low resolving power. Only features with half wavelength of the order of 2,000 km can be detected.

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The maps on the cover show the results from one recent study (Nataf, Nakanishi and Anderson, 1984). Shown are the parameters VSV and XI at two depths, 250 and 350 km. VSV is the velocity of vertically polarized shear waves, determined primarily from fundamental mode Rayleigh waves, and XI is related to the difference between VSH, the velocity of horizontally polarized shear waves, and VSV. The parameter XI is therefore a measure of anisotropy. Positive XI means  $VSH > VSV$ . Aggregates composed of a-axis horizontal olivine crystals, for example, are expected to have  $XI > 0$ . This situation is probably diagnostic of horizontal flow.  $XI < 0$  is expected for vertical a-axis orientation of olivine and implies vertical flow. Blue areas in the maps represent fast regions or  $VSH > VSV$ . Orange areas are slower than average or  $VSH < VSV$ . The maps are spherical harmonic representations including coefficients up to and including order and degree six.

In spite of the lack of short wavelength information, there is much important information in these maps. Midoceanic ridges and regions of recent volcanic activity are generally slow at 250 km. The regions near the Tasman Sea - New Zealand, Red Sea - African rift and western North America are slow. The central Pacific, between Hawaii and Tahiti is also slow at this depth. Fast regions include the Canadian and Fennoscandian shields and the Siberian platform, as expected, but also the North Pacific and the eastern Indian Ocean. Many hotspots are on the edges of low-velocity regions rather than centrally located.

The parameter XI is negative over regions of upwelling (East Pacific Rise, Antarctic-Pacific Rise, South Indian Rises, Mid-Atlantic Rise and the Red Sea region and areas of presumed downwelling (Japan, Philippines, Mariannas, Sumatra). Plate interiors are generally positive XI, suggesting horizontal flow at 250 km.

The maps at 350 km are similar, as they should be since a high degree of correlation between 220 km and 400 km was assumed in the inversion. The differences are therefore particularly instructive. The central Pacific, Red Sea and Mid-Atlantic ridge slow anomalies persist, suggesting that these are relatively deep seated. Most shields are no longer evident. The thermal anomaly associated with the East Pacific Rise appears to be displaced. At 450 km depth, not shown, most ridges are fast and most subduction regions are also fast. The velocity and XI parameter at 350 km are consistent with upwelling flow along the East Pacific Rise, the central and southern Mid-Atlantic Rise and the Red Sea area. These parameters are consistent with downwelling in the western Pacific. An upwelling is implied in the south central Pacific.

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#### References

- Nakanishi, I., and D. L. Anderson, World-wide distribution of group velocity of mantle Rayleigh waves as determined by spherical harmonic inversion, Bull. Seismol. Soc. Am., 72, p. 1185-1194, 1982.
- Nakanishi, I., and D. L. Anderson, Measurement of mantle wave velocities and inversion for lateral heterogeneity and anisotropy, Part I., J. Geophys. Res., 88, p. 10,267-10,283, 1983.

Nakanishi, I., and D. L. Anderson, Aspherical heterogeneity of the mantle from phase velocities of mantle waves, *Nature*, 307, 117-121, 1984a.

Nakanishi, I., and D. L. Anderson, Measurements of mantle wave velocities and inversion for lateral heterogeneity and anisotropy, Part II. Analysis by the single-station method, *Geophys. J. R. astr. Soc.*, in press, 1984b.

Nataf, H.-C., I. Nakanishi, and D. L. Anderson, Anisotropy and shear velocity heterogeneities in the upper mantle, *Geophys. Res. Lett.*, in press, 1984.

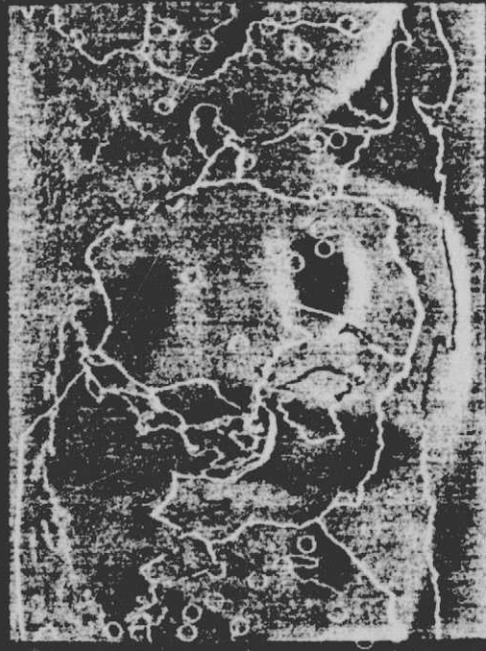
Tanimoto, T., and D. L. Anderson, Lateral heterogeneity and azimuthal anisotropy of the upper mantle: Love and Rayleigh waves 100-250 sec. *J. Geophys. Res.*, 1984a (submitted)

Tanimoto, T., and D. L. Anderson, Mapping convection in the mantle, *Geophys. Res. Lett.*, 1984b (submitted).

Woodhouse, J. H., and A. M. Dziewonski, Mapping the upper mantle: three dimensional modelling of earth structure by inversion of seismic waveforms, *J. Geophys. Res.* 1984 (submitted).

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35MM PRINTS

Model NNAB, VSV Depth=250 km



Model NNAB, VSV Depth=350 km



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Model NNAB, XI Depth=250 km



Model NNAB, XI Depth=350 km



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