The Earth's carbon cycle involves crustal reservoirs that include atmospheric CO₂, oceanic inorganic and organic carbon, carbon in sedimentary organic matter such as that in shales, sedimentary carbonate, and igneous rocks. There also seems to be a relatively low but nontrivial rate of exchange between the carbon reservoir in the Earth's mantle and crustal carbon.

Certain processes moving carbon between these reservoirs are isotope-selective. Photosynthetic carbon fixation produces organic matter which can be between 0.3 and 3 percent depleted in ¹³C relative to sources of inorganic carbon (Buehns, 1979; Deines, 1980). This preference for ¹³C is principally due to catalysis in the first step of carbon dioxide fixation (e.g., the Calvin cycle enzyme ribulose bisphosphate carboxylase, Listop, 1973). The amount of fractionation is often attenuated by processes which transport inorganic carbon to cellular cytoplasm (e.g., O'Leary, 1981). The respiration of organic carbon is also associated with a discrimination against ¹³C. The carbon dioxide produced is typically depleted between 0 and 1.2 percent in ¹³C relative to the organic carbon source (de Niro, 1977; Kaplan and Rittenberg, 1964).

The organic carbon buried in sediments can be isotope-selectively fractionated by sulfate reduction (Buehns, 1979). Work in our laboratory has shown that even fermentation can produce striking discrimination against the ¹³C isotope.

More deeply buried organic carbon becomes progressively more depleted by thermal decomposition. If this carbon is heated to several hundred degrees Celsius, it can be converted to graphite. The ¹³C content of this residual organic matter changes very little until its elemental carbon to hydrogen ratio falls below 0.2 (Schopf, 1983). Below 0.2 the residual carbon becomes more ¹³C-enriched, because the ¹³C value of sedimentary carbon is so well preserved, we are assured that the ¹³C values in the better preserved sedimentary rocks as old as 3.5 billion years reflect ancient microbial processes.

Carbon issuing from the midocean ridges is about 0.5 percent depleted in ¹³C relative to marine carbonates, and about 2 percent enriched in ¹³C relative to sedimentary organic carbon (Des Marais and Noone, 1984). This midocean ridge carbon represents the largest single carbon flux from the mantle, and its isotopic composition is the same as the average of all the crustal reservoirs. The crust/mantle exchange of carbon was more intense during the early Precambrian than and, very likely, the crust had a larger carbon inventory then than now (see Des Marais' article in Sundquist and Broeker, 1984).
Carbon isotope fractionation values have been used to understand the history of the biosphere. For example, plankton analyses confirmed that marine extinctions at the end of the Cretaceous period were indeed severe (see Hsu's article in Sundquist and Broeker, 1984). Variations in the isotopic compositions of carbonates and evaporitic sulfates during the Paleozoic might reflect the relative abundances of euxinic (anoxic) marine environments and organic deposits from terrestrial flora (Berner and Raissell, 1983). The carbon isotopic composition of Precambrian sediments suggest that the enzyme ribulose bisphosphate carboxylase has existed for perhaps 3.5 billion years (Schopf, 1983). Future work in our laboratory seeks to elucidate the relationship between the carbon isotopic composition of stromatolites, atmospheric CO₂, and oxygen inventories.


Figure I-3. Metabolic carbon isotope values for classes of biological compounds as determined by mass spectrometry.

METABOLIC CARBON ISOTOPE FRACTIONATION

\[
\begin{align*}
\delta^{13}C, \% & \quad -15 \\
-20 & \\
-25 & \\
-30 & \\
0 & \\
-5 & \\
-10 & \\
Ribulose bio-phosphatase carbonylase & \\
\end{align*}
\]
Figure 1-4. Carbon isotope fractionation values in live organisms and sediments. The $^{13}C$ values are a function of temperature in live marine organisms, and these values are altered by temperature alteration in sediments.
Figure I-5. Carbon isotopic fractionation ranges for extant carbon samples from terrestrial sources. Delta $^{13}$C scale at left, boxes represent major reservoirs, arrows represent major processes that transport and convert the carbon between the reservoirs.
Figure I-6. Reconstruction of the extent of depositional environments based on carbon and sulfur isotope fractionation measurements from ancient sedimentary rocks.

A. Major extant environments where sulfur and carbon compounds are sedimented indicating the relative distribution now.

B. Environmental reconstruction of the Cambrian Period based on interpretations of carbon and sulfur isotope data.

C. Environmental reconstruction of the Permian Period based on interpretations of carbon and sulfur isotope data.
**CAMBRIAN TIME 540 MILLION YEARS AGO**
Sulfates with high delta E34
Carbonates with low delta C13

**ORIGINAL PLANTS OF POOR QUALITY**

- Extensive production of epicontinental seas
- No land plants
- Euxinic basins
- Moderate marine productivity

**PERMIAN TIME 190 MILLION YEARS AGO**
Sulfates with low E34
Carbonates with high delta C13

**LOW MARINE PRODUCTIVITY**
- Freshwater rain forest and swamps

**EXTENSIVE HIGHLANDS**
- High organic carbon
- Low carbonate and sulfate

**HIGH SULFATE**
- Low sulfate and organic carbon

**RELATIVELY LEAST CONTINENTAL AREA**
- Marine life flourishes

**MAJOR LAND AND MARINE EXTINCTIONS**

Des Marais based on hypothesis of Veror and Russell. 1983