Solar Radiation on Mars

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ABSTRACT

Detailed information on solar radiation characteristics on Mars are necessary for effective design of future planned solar energy systems operating on the surface of Mars. In this paper we present a procedure and solar radiation related data from which the diurnally, hourly and daily variation of the global, direct beam and diffuse insolation on Mars are calculated. The radiation data are based on measured optical depth of the Martian atmosphere derived from images taken of the sun with a special diode on the Viking cameras; and computation based on multiple wavelength and multiple scattering of the solar radiation.

INTRODUCTION

NASA, through its Project Pathfinder, has put in place a wide-ranging set of advanced technology programs to address future needs of manned space exploration. Included in the mission sets under study is the establishment of outposts on the surface of Mars. The Surface Power program in Pathfinder is aimed at providing ultralightweight photovoltaic array technology for such an application. Detailed information on solar radiation data on the Martian surface is necessary to allow more accurate estimates of photovoltaic power system size and mass in system analysis and trade-off studies of relevant technology options. Of major concern are the dust storms, which have been observed to occur on local as well as on global scales, and their effect on

*This work was done while the author was a National Research Council - NASA Research Associate; on sabbatical leave from Tel-Aviv University.
solar array output. In general, the assumption has been that global storms would reduce solar array output essentially to zero, because the opacity of the atmosphere may rise to values ranging from 3 to 9, depending on the severity of the storm. Furthermore, such high values of opacity may persist for long periods of time such that the requirement for energy storage quickly becomes much too large to be practical. As shown in Refs. 1 and 2 (and is believed to be published for the first time to the best of our knowledge), there is still an appreciable large diffuse component, even at high opacities, so that solar array operation is still possible.

Calculation of solar radiation incident on the top of Martian atmosphere and on the Martian surface has been previously published (Refs. 3 to 5) taking into account the direct beam component only of the solar radiation. The introduction of the diffuse component in this paper became possible with the normalized net flux function described in the paper. This paper is an extension of the study in Refs. 1 and 2.

As on the planet Earth, the solar radiation on the surface of Mars is composed of two components: the direct beam and diffuse component. The direct beam is affected by scattering and absorption along the path from the top of the Martian atmosphere to the Martian surface. Measurement of the optical depth (Refs. 6 and 7) of the Martian atmosphere allows an estimate of the absorption and scattering out of the beam. These estimates were derived from images taken of the Sun and Phobos with a special diode on the Viking lander cameras.

Earth-terrestrial insolation data are accumulated over many years at different locations around the world and are given as long term average values. The optical depth data for Mars are derived from less than two Mars years. Consequently, the calculated solar insolation, in the present paper, corresponds to short term data. Furthermore, the measured opacities (optical
depth) and the calculated insolation pertain to just two locations on the Mars planet; Viking lander 1 (VL1) is located at 22.3° N latitude and 47.9° W longitude, and Viking lander 2 (VL2) is located at 47.7° N latitude and 225.7° W longitude. However, the similarity in the properties of the dust suspended above the two landing sites suggests that they are also representative of ones at other locations, at least, at latitudes not too far from the lander's sites. Data from lander VL1 may be used for latitudes 40° N to 40° S and data from lander VL2 for higher latitudes. The Martian atmosphere consists mainly of suspended dust particles, the amounts of which vary daily, seasonally, and annually, depending on local and global storm intensities and their duration. The optical depth values given in the section entitled Optical Depth are assumed to be constant throughout the day. Large values of optical depth correspond to global storms, i.e., days with low insolation (dark days).

The albedo of the Martian surface varies in the range of about 0.1 to 0.4. The irradiances derived in the section entitled Solar Radiation correspond to 0.1 albedo, but can be also used for other values of albedo, to the first approximation.

In this paper a normalized net solar flux function is introduced from which, together with the variation of the opacities, characteristics of the solar radiation on the Martian surface are calculated. This includes, among others, the diurnal and hourly variation of the global, beam and diffuse radiation on the horizontal surface. The results are presented in a series of figures and tables. The solar radiation data and the procedure presented in this paper can be used for the calculation of any desired solar radiation quantity in engineering design. New information about Mars may be forthcoming in the future from new analysis of previously collected data, from new
Earth-based observation, or from future flight missions. The Mars solar radiation data will thus be updated accordingly.

**NOMENCLATURE**

Radiation values:

- $G_b$: beam irradiance on Mars surface
- $G_{bh}, G_{dh}$: global, beam and diffuse irradiance on Mars horizontal surface
- $G_{ob}$: beam irradiance at the top of Mars atmosphere
- $G_{obh}$: beam irradiance on a horizontal surface at the top of Mars atmosphere
- $H_{h}, H_{bh}, H_{dh}$: global, beam and diffuse daily insolation on Mars horizontal surface
- $H_{obh}$: daily beam insolation on a horizontal surface at the top of Mars atmosphere
- $I_{h}, I_{bh}, I_{dh}$: global, beam and diffuse hourly insolation on Mars horizontal surface
- $I_{obh}$: hourly beam insolation on a horizontal surface at the top of Mars atmosphere

Subscripts:

- $b$: beam values
- $d$: diffuse values
- $h$: horizontal values
- $o$: values on top of Mars atmosphere

Other values:

- $e$: eccentricity = 0.093377
- $L_s$: areocentric longitude
- $m(z)$: airmass
- $r$: Sun-Mars distance
- $S$: solar constant = 1371 W/m² at the mean Sun-Earth distance of 1 astronomical unit (AU).
- $T$: Mars solar time
The most direct and probably most reliable estimates of opacity are those derived from Viking lander imaging of the Sun. Figures 1 and 2 show the seasonal variation of the normal-incidence of the optical depth at the Viking lander locations VL1 and VL2, respectively. The season is indicated by the value of $L_\text{s}$, areocentric longitude of the Sun, measured in the orbital plane of the planet from its vernal equinox, $L_\text{s} = 0^\circ$. Figures 1 and 2 were derived from references by Pollack (Refs. 6 and 7) and Zurek (Ref. 8) and were discretized for each 5°. As mentioned before, the optical depth is assumed to remain constant throughout the day. Opacities are minimum during the northern spring ($L_\text{s} = 0^\circ$ to $90^\circ$) and summer ($L_\text{s} = 90^\circ$ to $180^\circ$), and maximum during southern spring ($L_\text{s} = 180^\circ$ to $270^\circ$) and summer ($L_\text{s} = 270^\circ$ to $360^\circ$), the seasons during which most local and major dust storms occur. When dust storms are not present, the optical depth is typically about 0.5. Two global dust storms occurred during the periods of each observation as indicated by the high values of the optical depth (they are lower bound values).

Mars has seasons comparable to those of Earth. However, the seasons are on the average about twice as long as on the Earth, corresponding to the
greater length of the Martian year (Table I). Furthermore, they are distinctly unequal in duration as a result of the appreciable eccentricity of the Martian orbit. For that reason, the Martian year is not divided into months. Table I gives the duration of the Martian seasons in terrestrial and Martian days (a Martian day Λ sol, 1 sol = 24.65 hr). Areocentric longitudes $L_s = 0^\circ$ and $180^\circ$ correspond to the spring and fall equinox for the northern hemisphere, respectively, and $L_s = 90^\circ$ and $270^\circ$ correspond to northern and southern summer solstices, respectively.

**GLOBAL AND LOCAL DUST STORMS**

The intensity of Martian global and local dust storms is defined in terms of opacity of the dust it raises. Global dust storms are those which obscure planetary-scale sections of the Martian surface for many Martian days (sols), whereas local dust storms are less intense, and form and dissipate in a few days or less. From a photovoltaic system design point of view, the intensity, frequency, and duration of these storms may be viewed as partially cloudy and cloudy days for which additional energy storage in the photovoltaic system must be taken into account. The characteristics of global and local dust storms are listed below.

**Global Dust Storms**

(1) One, or occasionally two global dust storms of planetary scale may occur each Martian year. The duration may vary from 35 to 70 days or more. Although global dust storms do not occur every year, their occurrence is fairly frequent.

(2) Global dust storms begin near perihelion, when solar insolation is maximum (southern spring and summer) in the southern mid-latitude.

(3) The first global dust storm observed by VL (1977) spread from a latitude of $40^\circ$ S to a latitude $48^\circ$ N in about 5 to 6 days.
(4) The opacity during the global dust storm is greater than 1.

Local Dust Storms

(1) Local dust storms occur at almost all latitudes and throughout the year. However, they have been observed to occur most frequently in the approximate latitude belt 10° to 20° N and 20° to 40° S, with more dust clouds seen in the south than in the north, the majority of which occurred during southern spring.

(2) Based on Viking orbiter observations, it is estimated that approximately 100 local storms occur in a given Martian year.

(3) Local dust storms last a few days.

(4) The opacity of local dust storms may be assumed about 1.

SOLAR RADIATION AT THE TOP OF MARS ATMOSPHERE

The variation of the solar radiation at the top of the Mars atmosphere is governed by the location of Mars in its orbit and by the solar zenith angle, and is of direct beam radiation. The beam irradiance, in W/m², is given by:

$$G_{ob} = \frac{S}{r^2}$$

where $S$ is the solar constant at the mean Sun-Earth distance of 1 AU, i.e., $S = 1371 \text{ W/m}^2$; $r$ is the instantaneous Sun-Mars distance in AU (heliocentric distance) given by (9):

$$r = \frac{a(1 - e^2)}{1 + e \cos \theta}$$

where $a$ is the Mars semimajor axis in AU, and $e$ is the Mars eccentricity, i.e., $e = 0.093377$; and $\theta$ is the true anomaly given by:

$$\theta = L_S - 248^\circ$$

where $L_S$ is the areocentric longitude and $248^\circ$ is the areocentric longitude of Mars perihelion. The Sun-Mars mean distance in astronomical units (AU) is 1.5236915; therefore, the mean beam irradiance at the top of
the Martian atmosphere is: \( \frac{1371}{1.5236915^2} = 590 \text{ W/m}^2 \). The instantaneous beam irradiance is given by Eqs. (1) to (3):

\[
G_{ob} = 590 \frac{[1 + e \cos(L_s - 248^\circ)]^2}{(1 - e^2)^2} \tag{4}
\]

and is shown in Fig. 3.

The beam irradiance on a horizontal surface is:

\[
G_{obh} = G_{ob} \cos z \tag{5}
\]

where \( z \) is the zenith angle of the incident solar radiation given by:

\[
\cos z = \sin \phi \sin \delta + \cos \phi \cos \delta \cos \omega \tag{6}
\]

where

\( \phi \) latitude

\( \delta \) declination angle

\( \omega \) hour angle measured from the true noon westward.

The solar declination angle is given by:

\[
\sin \delta = \sin \delta_0 \sin L_s \tag{7}
\]

where \( \delta_0 = 24.936^\circ \) is the Mars obliquity of rotation axis. The variation of the solar declination angle is shown in Fig. 4. The four seasons pertain here to the northern hemisphere, the reverse is true in the southern hemisphere. The ratio of Mars to Earth length of day is 24.65/24. It is convenient, for calculation purposes, to define a Mars hour by dividing the Martian day into 24 hr. Using the same relationship between the Mars solar time \( T \) and the hour angle as for the Earth, we write:

\[
\omega = 15T - 180 \tag{8}
\]

This is shown in Fig. 5. The final solar radiation results can then be adjusted by the above ratio to correspond to actual (terrestrial) time.
Examples of the solar radiation calculation procedure and results in this paper pertain to $L_s = 69^\circ$, $120^\circ$, $153^\circ$, $249^\circ$ and $299^\circ$ at Viking lander VL1 location $\phi = 22.3^\circ$ N. Areocentric longitude $L_s = 69^\circ$ corresponds to aphelion; $L_s = 249^\circ$ to perihelion; $L_s = 153^\circ$ to mean radiation of 590 W/m$^2$; $L_s = 120^\circ$ corresponds to the lowest atmosphere opacity of 0.4; and $L_s = 299^\circ$ to the highest opacity of 3.25 (Fig. 1). For a given $L_s$ and latitude $\phi$, one can calculate the zenith angle $z$ as function of solar time $T$ using Eqs. (5) to (7). The beam irradiance on a horizontal surface is then determined using Eqs. (4) and (5). The diurnal variation of the beam irradiance on a horizontal surface at the top of the Mars atmosphere for the above mentioned values of $L_s$ is shown in Fig. 6. Because of symmetry, only the afternoon values are shown in the figure. The sunset hour angle is given by:

$$\omega_{SS} = \cos^{-1}(-\tan \phi \tan \delta)$$

and the number of Mars daylight hours is

$$T_d = \frac{2}{15} \cos^{-1}(-\tan \phi \tan \delta)$$

It is of interest to calculate the solar beam insolation on a horizontal surface in watt hours per square meter (Whr/m$^2$), for a desired period of time between hour angles $\omega_1$ and $\omega_2$. This is obtained by integrating Eq. (5), i.e.,

$$I_{obh} = \frac{12*}{\pi} G_{ob} \int_{\omega_1}^{\omega_2} (\sin \phi \sin \delta + \cos \phi \cos \delta \cos \omega) d\omega$$

or

$$I_{obh} = \frac{12*}{\pi} G_{ob} \left[ \frac{2\pi(\omega_2 - \omega_1)}{360} \sin \phi \sin \delta + \cos \phi \cos \delta \sin \omega_2 - \sin \omega_1 \right]$$

Replace the 12 hr by 12.325 in Eqs. (11) and (12) to get the insolation with reference to actual (terrestrial) time.
A commonly used quantity is the hourly insolation, in Whr/m²-hr. In that case \( \omega_1 \) and \( \omega_2 \) define an hour. The daily solar insolation, \( H_{obh} \), on a horizontal surface, in Whr/m²-day, is often needed. This is obtained by integrating Eq. (11) over the period from sunrise to sunset. One gets:

\[
H_{obh} = \frac{24}{\pi} G_{ob} \left[ \frac{2\omega SS}{360} \sin \phi \sin \delta + \cos \phi \cos \delta \sin \omega SS \right] \tag{13}
\]

Table II gives the hourly beam insolation \( I_{obh} \), and the daily beam insolation \( H_{obh} \), on a horizontal surface at the top of the Mars surface. The diurnal variation of the hourly beam insolation data in Table II is shown in Fig. 7. To obtain the terrestrial watt-hours one needs to multiply the values in Table II by the ratio 24.65/24.

SOLAR RADIATION ON THE SURFACE OF MARS

The variation of the solar radiation on the Martian surface is governed by three factors: (1) the Mars-Sun distance, (2) solar zenith angle, and (3) by the opacity of the Martian atmosphere. The global solar radiation is composed of the direct beam and diffuse components. The direct beam irradiance, \( G_b \), on the Martian surface normal to the solar rays is related by Beer's law to the optical depth, \( \tau \), of the intervening atmospheric haze:

\[
G_b = G_{ob} \exp[-\tau m(z)] \tag{14}
\]

where \( m(z) \) is the air mass determined by the zenith angle \( z \), and can be approximated by:

\[
m(z) = \frac{1}{\cos z} \tag{15}
\]

*Replace the 24 hr by 24.65 hr in Eq. (13) to get the insolation with reference to actual (terrestrial) time.
The net solar flux integrated over the solar spectrum on the Martian surface was calculated by Pollack (Ref. 10) based on multiple wavelength and multiple scattering of the solar radiation. Derived data from this calculation are shown in Table III by the normalized net flux function $f(z, \tau)$ where the parameters are the zenith angle $z$ and the optical depth $\tau$. This table pertains to an albedo of 0.1 but can be used for higher albedo values to a first approximation. Using this data we calculated the global solar irradiance. We assumed that the diffuse irradiance is obtained by subtracting the beam from the global irradiance.

The solar irradiance components, on a horizontal Martian surface, are related by:

$$G_h = G_{bh} + G_{dh} \quad (16)$$

where

- $G_h$ : global irradiance on a horizontal surface
- $G_{bh}$ : direct beam irradiance on a horizontal surface
- $G_{dh}$ : diffuse irradiance on a horizontal surface

The diffuse irradiance of the Martian atmosphere may be a result of a different mechanism than that for the Earth atmosphere, nevertheless, to a first approximation, we will apply Eq. (16) as for Earth-terrestrial calculations. The global irradiance $G_h$ on a horizontal surface is given by:

$$G_h = G_{ob} \cos Z \frac{f(z, \tau)}{0.9} \quad (17)$$

The beam irradiance $G_{bh}$ on a horizontal surface is obtained by:

$$G_{bh} = G_{ob} \cos Z \exp \left(-\frac{\tau}{\cos z}\right) \quad (18)$$

*The factor 0.9 comes from the expression $(1 - \text{albedo})$ in the denominator. For an albedo of 0.1, the denominator is 0.9.
The diffuse irradiance on a horizontal surfaces is obtained from Eqs. (16) to (18). Figures 8 to 10 describe the variation of the global, beam and diffuse irradiances, respectively, on a horizontal Martian surface; and are given in pairs as functions of the optical depth $\tau$ and zenith angle $z$. The irradiances were calculated based on Table III data and the mean irradiance of $590 \text{ W/m}^2$. The variation of the global irradiance $G_h$, Eq. (17), is shown in Figs. 8(a) and (b). The beam irradiance $G_{bh}$ is obtained using Eq. (18) and is shown in Figs. 9(a) and (b). The beam irradiance shows a sharp decrease with increasing of the optical depth, and a relative moderate decrease with increasing of the zenith angle. The diffuse irradiance $G_{dh}$ is shown in figure 10(a) and (b). The diffuse irradiance shows a sliding maximum with the variation of the zenith angle.

The solar radiation (global, beam and diffuse) variation (diurnal, hourly and daily) can be calculated based on the preceding equations and the $f(z,\tau)$ table data. The following examples pertain again to the Viking lander VL1 location and areocentric longitudes $L_s = 69^\circ, 120^\circ, 153^\circ, 249^\circ, \text{ and } 299^\circ$. Daily solar insolation are also given for $L_s = 0^\circ, 30^\circ, 60^\circ, 90^\circ, 150^\circ, 180^\circ, 210^\circ, 240^\circ, 300^\circ, \text{ and } 330^\circ$. For a given $L_s$ and $\phi$, one can calculate the variation of the zenith angle $z$ as function of the Mars solar time $T$ using Eqs. (5) to (8). Referring to Fig. 1 for the given $L_s$, the optical depth $\tau$ is determined; with Table III and Eqs. (16) to (18) one can calculate the solar radiation variation for the given day. The results are shown in Figs. 11 to 13. Because of symmetry around 12:00, the graphs in Figs. 12 and 13 are the forenoon or afternoon variation. The figures show clearly that for higher opacities, the diffuse component dominates the solar radiation.

The hourly solar insolation (global, beam and diffuse) on a horizontal surface, in Whr/m$^2$-hr, can be calculated based on Figs. 11 to 13 by
integrating hourly areas. The daily insolation \( H_h \) on a horizontal surface is the summation of the hourly values. The beam insolation, for a desired period of time, can be also calculated by:

\[
I_{bh} = \frac{12}{\pi} G_{ob} \int_{\omega_1}^{\omega_2} \left( \sin \phi \sin \delta + \cos \phi \cos \delta \cos \omega \right) \exp \left[ -\tau / \left( \sin \phi \sin \delta + \cos \phi \cos \delta \cos \omega \right) \right] d\omega \quad (19)
\]

Tables IV to VI give the hourly global \( I_h \), beam \( I_{bh} \) and diffuse \( I_{dh} \) insolation as well as the daily global \( H_h \), beam \( H_{bh} \) and diffuse \( H_{dh} \) insolation. Included in the tables are also the number of Martian daylight hours and the daily mean irradiance. For a day \( L_s = 299 \) with a relative high opacity, the daily mean global irradiance is still appreciable and is about 30 percent of that in a clear day. The diurnal variation of the hourly global insolation data in Table IV is shown in Fig. 14. The percentage of diffuse and beam insolation for the five analyzed \( L_s \) days is shown in Fig. 15. The daily global insolation on a horizontal surface on Mars is shown in Fig. 16 for twelve areocentric longitudes covering a Martian year. Using the procedure outlined, one can calculate the variation of the solar radiation for any desired day to use for any engineering system design.

CONCLUSIONS

Effective design and utilization of solar energy depend to a large extent on adequate knowledge of solar radiation characteristics in the region of solar energy system operation. In this paper we presented a procedure and solar radiation related data from which the diurnally, hourly and daily variation of insolation on Mars were calculated. This includes the global, beam and diffuse insolation on a horizontal surface, from which any desired

*Replace the 12 hr by 12.325 hr in Eq. (19) to get the insolation with reference to actual (terrestrial) time.
solar radiation quantity can be derived for an engineering design. The global
insolation on the surface of Mars was derived based on the normalized net
solar flux function \( f(z, \tau) \); the beam insolation was determined by Beer's
law relating the isolation to the optical depth of the Martian atmosphere; and
the diffuse insolation was calculated as the difference between the global and
the beam insolation. The optical depths were measured at the two Viking
lander locations, but can also be used, to the first approximation, for other
locations. One of the most important results of this study is that there is a
large diffuse component of the solar insolation, even at high optical depth,
so that solar energy system operation is still possible. In absence of long
term insolation data on Mars, the data presented in this paper can be used
until updated data are available from new analysis or future flight missions.

ACKNOWLEDGMENT

We are very grateful to James B. Pollack from the Space Science Division,
NASA Ames Research Center for supplying us with the data from which the \( f(z, \tau) \)
table was derived.

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### TABLE I. - MARS SEASONAL DURATION

<table>
<thead>
<tr>
<th>Areocentric longitude of the sun, $L_s$</th>
<th>Season</th>
<th>Duration of the season Mars</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Northern hemisphere</td>
<td>Southern hemisphere</td>
</tr>
<tr>
<td>0 to 90°</td>
<td>Spring</td>
<td>Autumn</td>
</tr>
<tr>
<td>90 to 180°</td>
<td>Summer</td>
<td>Winter</td>
</tr>
<tr>
<td>180 to 270°</td>
<td>Autumn</td>
<td>Spring</td>
</tr>
<tr>
<td>270 to 360° or 0°</td>
<td>Winter</td>
<td>Summer</td>
</tr>
</tbody>
</table>

### TABLE II. - HOURLY AND DAILY BEAM INSOLATION ON A HORIZONTAL SURFACE AT TOP OF MARS ATMOSPHERE

[VL1: $\phi = 22.3^\circ$N]

<table>
<thead>
<tr>
<th>Hourly* $I_{obh}$ (W/m²-hr) for hours ending at:</th>
<th>Daily* $H_{obh}$, W/m²-day</th>
</tr>
</thead>
<tbody>
<tr>
<td>Day-L$_s$ 13:00 14:00 15:00 16:00 17:00 18:00 19:00</td>
<td></td>
</tr>
<tr>
<td>69°</td>
<td>488</td>
</tr>
<tr>
<td>120°</td>
<td>528</td>
</tr>
<tr>
<td>153°</td>
<td>572</td>
</tr>
<tr>
<td>249°</td>
<td>496</td>
</tr>
<tr>
<td>299°</td>
<td>478</td>
</tr>
</tbody>
</table>

*Multiply by 24.65/24 to obtain actual (terrestrial) W/m²
TABLE III. - NORMALIZED NET FLUX FUNCTION \( f(z, \tau) \) AT THE MARTIAN SURFACE

<table>
<thead>
<tr>
<th>Optical depth</th>
<th>( \tau )</th>
<th>0</th>
<th>10</th>
<th>20</th>
<th>30</th>
<th>40</th>
<th>50</th>
<th>60</th>
<th>70</th>
<th>80</th>
<th>85</th>
</tr>
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<td></td>
<td>0.885</td>
<td>0.883</td>
<td>0.882</td>
<td>0.880</td>
<td>0.876</td>
<td>0.870</td>
<td>0.857</td>
<td>0.830</td>
<td>0.755</td>
<td>0.635</td>
</tr>
<tr>
<td>0.2</td>
<td></td>
<td>0.866</td>
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<td>0.860</td>
<td>0.858</td>
<td>0.851</td>
<td>0.836</td>
<td>0.813</td>
<td>0.758</td>
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<td>0.470</td>
</tr>
<tr>
<td>0.3</td>
<td></td>
<td>0.847</td>
<td>0.846</td>
<td>0.841</td>
<td>0.836</td>
<td>0.826</td>
<td>0.806</td>
<td>0.774</td>
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<td>0.827</td>
<td>0.821</td>
<td>0.815</td>
<td>0.802</td>
<td>0.778</td>
<td>0.740</td>
<td>0.667</td>
<td>0.502</td>
<td>0.373</td>
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<tr>
<td>0.5</td>
<td></td>
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<td>0.810</td>
<td>0.802</td>
<td>0.796</td>
<td>0.778</td>
<td>0.752</td>
<td>0.708</td>
<td>0.628</td>
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<td>0.775</td>
<td>0.755</td>
<td>0.725</td>
<td>0.677</td>
<td>0.593</td>
<td>0.414</td>
<td>0.318</td>
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<tr>
<td>0.7</td>
<td></td>
<td>0.776</td>
<td>0.773</td>
<td>0.766</td>
<td>0.755</td>
<td>0.733</td>
<td>0.700</td>
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<td>0.756</td>
<td>0.750</td>
<td>0.736</td>
<td>0.710</td>
<td>0.675</td>
<td>0.616</td>
<td>0.520</td>
<td>0.360</td>
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<td>0.080</td>
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</table>
## TABLE IV. - HOURLY AND DAILY GLOBAL INSOLATION ON A HORIZONTAL SURFACE AT MARS SURFACE

**[VL1: \( \phi = 22.3^\circ N \)]**

<table>
<thead>
<tr>
<th>Day-L(_s)</th>
<th>( \tau )</th>
<th>13:00</th>
<th>14:00</th>
<th>15:00</th>
<th>16:00</th>
<th>17:00</th>
<th>18:00</th>
<th>19:00</th>
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</thead>
<tbody>
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<td>0.65</td>
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<td>390</td>
<td>338</td>
<td>263</td>
<td>170</td>
<td>78</td>
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<tr>
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<td>477</td>
<td>446</td>
<td>387</td>
<td>306</td>
<td>201</td>
<td>98</td>
<td>15</td>
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<tr>
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<td>508</td>
<td>471</td>
<td>399</td>
<td>302</td>
<td>185</td>
<td>73</td>
<td>3</td>
</tr>
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<td>249°</td>
<td>1.40</td>
<td>307</td>
<td>270</td>
<td>204</td>
<td>122</td>
<td>45</td>
<td>2</td>
<td>--</td>
</tr>
<tr>
<td>299°</td>
<td>3.25</td>
<td>170</td>
<td>149</td>
<td>107</td>
<td>61</td>
<td>24</td>
<td>1</td>
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</tr>
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</table>

*Multiply by 24.65/24 to obtain actual (terrestrial) Whr/m\(^2\) or hours

<table>
<thead>
<tr>
<th>Daily global insol., ( H_2 ) (Whr/m(^2)day)</th>
<th>Daylight hours( T_d ), hr</th>
<th>Daily mean global irradiance, ( W/m^2 )</th>
</tr>
</thead>
<tbody>
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<td>1900</td>
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<td>1024</td>
<td>10.75</td>
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</tr>
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</table>

## TABLE V. - HOURLY AND DAILY BEAM INSOLATION ON A HORIZONTAL SURFACE AT MARS SURFACE

**[VL1: \( \phi = 22.3^\circ N \)]**

<table>
<thead>
<tr>
<th>Day-L(_s)</th>
<th>( \tau )</th>
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<th>15:00</th>
<th>16:00</th>
<th>17:00</th>
<th>18:00</th>
<th>19:00</th>
</tr>
</thead>
<tbody>
<tr>
<td>69°</td>
<td>0.65</td>
<td>252</td>
<td>230</td>
<td>186</td>
<td>128</td>
<td>67</td>
<td>20</td>
<td>3</td>
</tr>
<tr>
<td>120°</td>
<td>0.40</td>
<td>352</td>
<td>322</td>
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<td>190</td>
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<td>33</td>
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</tr>
<tr>
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<td>0.50</td>
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<td>310</td>
<td>244</td>
<td>163</td>
<td>77</td>
<td>15</td>
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<tr>
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<td>50</td>
<td>26</td>
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<td>3</td>
<td>2</td>
<td>1</td>
<td>--</td>
<td>--</td>
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</table>

*Multiply by 24.65/24 to obtain actual (terrestrial) Whr/m\(^2\) or hours

<table>
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<tr>
<th>Daily beam insol., ( H_{bh} ) (Whr/m(^2)day)</th>
<th>Daylight hours( T_d ), hr</th>
<th>Daily mean beam irradiance, ( W/m^2 )</th>
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TABLE VI. - HOURLY AND DAILY DIFFUSE INSOLATION ON A HORIZONTAL SURFACE AT MARS SURFACE
[VL1: = 22.3°N]

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<th>15:00</th>
<th>16:00</th>
<th>17:00</th>
<th>18:00</th>
<th>19:00</th>
<th>Daily diffuse insolation, ( \frac{H_{dn}}{W/m^2 \cdot \text{day}} )</th>
<th>Daylight hours, ( T_d ), hr</th>
<th>Daily mean diffuse irradiance, ( W/m^2 )</th>
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<td>135</td>
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<tr>
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*Multiply by 24.65/24 to obtain actual (terrestrial) Whr/m² or hours
FIGURE 1. - OPTICAL DEPTH AS MEASURED FOR VIKING LANDER VL1 AS FUNCTION OF AREOCENTRIC LONGITUDE.
FIGURE 2. - OPTICAL DEPTH AS MEASURED FOR VIKING LANDER VL2 AS FUNCTION OF AREOCENTRIC LONGITUDE.
FIGURE 3. - BEAM IRRADIANCE AT THE TOP OF MARS ATMOSPHERE AS FUNCTION OF AREOCENTRIC LONGITUDE.

FIGURE 4. - VARIATION OF SOLAR DECLINATION ANGLE $\delta$, WITH AREOCENTRIC LONGITUDE, $L_s$.

FIGURE 5. - SOLAR TIME AND HOUR ANGLE RELATION.
FIGURE 6. - DIURNAL VARIATION OF BEAM IRRADIANCE ON A HORIZONTAL SURFACE AT TOP OF MARS ATMOSPHERE.
FIGURE 7. - DIURNAL VARIATION OF HOURLY BEAM INSOLATION ON A HORIZONTAL SURFACE AT TOP OF MARS ATMOSPHERE.
Figure 8. Variation of global irradiance with optical depth and sun zenith angle on a horizontal surface.
FIGURE 9. - VARIATION OF BEAM IRRADIANCE WITH OPTICAL DEPTH AND SUN ZENITH ANGLE ON A HORIZONTAL SURFACE.
FIGURE 10. - VARIATION OF DIFFUSE IRRADIANCE WITH OPTICAL DEPTH AND SUN ZENITH ANGLE ON A HORIZONTAL SURFACE.
FIGURE 11. - DIURNAL VARIATION OF GLOBAL $G_h$, BEAM $G_{bh}$ AND DIFFUSE $G_{dh}$ IRRADIANCE ON A HORIZONTAL MARS SURFACE AT VIKING LANDER VL1.

FIGURE 12. - DIURNAL VARIATION OF GLOBAL $G_h$, BEAM $G_{bh}$ AND DIFFUSE $G_{dh}$ IRRADIANCE ON A HORIZONTAL MARS SURFACE AT VIKING LANDER VL1.
FIGURE 13. - DIURNAL VARIATION OF GLOBAL $G_h$, BEAM $G_{bh}$ AND DIFFUSE $G_{dh}$ IRRADIANCE ON A HORIZONTAL MARS SURFACE AT VIKING LANDER VL1.
FIGURE 14. - DIURNAL VARIATION OF HOURLY GLOBAL INSOLATION ON A HORIZONTAL SURFACE ON MARTIAN SURFACE.
FIGURE 15. - PERCENT OF DIFFUSE AND DIRECT BEAM INSOLATION ON A HORIZONTAL MARS SURFACE.

FIGURE 16. - DAILY GLOBAL INSOLATION ON A HORIZONTAL MARS SURFACE AT VIKING LANDER VL1.
**Title and Subtitle**
Solar Radiation on Mars

**Author(s)**
Joseph Appelbaum and Dennis J. Flood

**Performing Organization Name and Address**
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Cleveland, Ohio 44135-3191

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National Aeronautics and Space Administration
Washington, D.C. 20546-0001

**Abstract**
Detailed information on solar radiation characteristics on Mars are necessary for effective design of future planned solar energy systems operating on the surface of Mars. In this paper we present a procedure and solar radiation related data from which the diurnally, hourly and daily variation of the global, direct beam and diffuse insolation on Mars are calculated. The radiation data are based on measured optical depth of the Martian atmosphere derived from images taken of the sun with a special diode on the Viking cameras; and computation based on multiple wavelength and multiple scattering of the solar radiation.

**Key Words (Suggested by Author(s))**
Mars; Viking landers; Solar radiation; Irradiance; Insolation; Global; Direct beam and diffuse radiation; Optical depth; Local and global dust storms

**Distribution Statement**
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Subject Category 92

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