OZONE FORMATION DURING AN EPISODE OVER EUROPE: A 3-D CHEMICAL/TRANSPORT MODEL SIMULATION

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ABSTRACT

A 3D regional photochemical tracer/transport model for Europe and the Eastern North Atlantic has been developed based on the NASA/GISS CTM. The model resolution is 4x5 degrees latitude and longitude with 9 layers in the vertical (7 in the troposphere). Advevtive winds, convection statistics and other meteorological data from the NASA/GISS GCM are used. An extensive gas-phase chemical scheme based on the scheme used in our global 2D model has been incorporated in the 3D model. In this work ozone formation in the troposphere is studied with the 3D model during a 5 day period starting June 30. Extensive local ozone production is found and the relationship between the source regions and the downwind areas are discussed. Variations in local ozone formation as a function of total emission rate, as well as the composition of the emissions (HC/NO\textsubscript{x} ratio and isoprene emissions) are elucidated. An important vertical transport process in the troposphere is by convective clouds. The 3D model includes an explicit parameterization of this process. It is shown that this process has significant influence on the calculated surface ozone concentrations.

1 Introduction

Ozone formation in the troposphere over Europe and other industrialized regions leading to ambient concentrations well above health standards during periods with high pressure weather situations, is now a well established fact (Guicherit; 1988). The increase is observed at urban as well as rural locations. Measurements (Kley et al.; 88) show that typical ozone values have increased from 5-15 ppbv at turn of the century, to 30-40 ppbv at present most probably due to increased man made emissions of NO\textsubscript{x} and hydrocarbons.

Several model studies have been performed to study this effect. This include box models, 1D and 2D eulerian models. Up to recently, lack of computer resources has put severe restrictions on the chemical scheme as well as the area and resolution of the models. The approach has been either to do regional studies with rather comprehensive chemical schemes (Chang et al.; 87), or global studies with simpler schemes (Crutzen and Zimmerman; 91). This paper presents results from a 3D model with an extensive chemical scheme covering an area from 25°W to 55°E and 24°N to 76°N. The model is intended to be extended to a global scale after development and testing on less computer costly scales.

2 Model description

2.1 Transport

The model was originally developed by Prather et al.(87) to simulate global distribution and temporal variability of CFCs. As consistent global data sets with real-time meteorological data from numerical weather prediction models are not available, the model is set up to use data from the free running NASA/GISS general circulation model (GCM).

In this study calculations are done for a European "window" (25°W to 55°E and 24°N to 76°N) with a resolution of 5° longitude and 4° latitude. In the vertical there are 9 layers (σ-coordinates, σ = (P - P\textsubscript{T})/(P\textsubscript{a} - P\textsubscript{T})) between the surface and 10 hPa. Table 1 gives the vertical resolution.

<table>
<thead>
<tr>
<th>σ</th>
<th>1.0</th>
<th>.95</th>
<th>.87</th>
<th>.73</th>
<th>.55</th>
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<tr>
<td>P (hPa)</td>
<td>984</td>
<td>934</td>
<td>854</td>
<td>720</td>
<td>550</td>
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<tr>
<td>Z (km)</td>
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<td>0.5</td>
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<td>σ</td>
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<td>0.25</td>
<td>0.14</td>
<td>0.061</td>
<td>0.0</td>
</tr>
<tr>
<td>P (hPa)</td>
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<td>255</td>
<td>150</td>
<td>70</td>
<td>10</td>
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<td>Z (km)</td>
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<td>10.3</td>
<td>13.7</td>
<td>18.5</td>
<td>31.2</td>
</tr>
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</table>

Table 1: Horizontal boundaries of the 9 layers in the model. P\textsubscript{T} = 984 hPa is the mean surface pressure in the GCM.

The meteorological data consist of 4 hour integrated values of horizontal mass fluxes and surface pressure, total optical depth, surface precipitation and number of convective events (see Prather et al.;87). Temperature, humidity, precipitation (from cumulus convection and large scale precipitation) are given as 5 day averages from the GCM. Tracer transport due to convection is calculated by a mass flux scheme which divide the convection into three classes (dry, moist/shallow and moist/deep).

The advection of tracers are solved numerically by the method of conservation of second order moments (Prather; 86). A timestep of one hour is applied in calculation of the advective transport.

2.2 Chemistry

The model has been expanded to include an extended ozone chemistry scheme. It includes a gas-phase chemistry with first order scavenging of water soluble species. The scheme follows the basic pattern of the scheme developed for use in our global 2D model (Isaksen and Hov; 87), with recent updates of reaction mechanisms and rates according to IUPAC (89).
Figure 1: Relative distribution of NO emissions. Numbers give percentage per grid square of the total NO-emission which is 9.35 Tg/y as NO.

2.3 Emissions

Reasonable emission rates of primary pollutants are crucial if the model shall be able to simulate the real atmosphere. Table 2 gives the emission rates of the primary pollutants (except isoprene). Emission rates and geographical distribution of NO and hydrocarbons are derived on the basis of data from EMEP (D. Simpson, private com.) and EPA (Watson et al.; 91). In the EMEP data hydrocarbon emissions are given as total man made VOC, while the EPA data divide the VOC emissions into groups such as paraffins, olefins, aromats, etc., but on a much coarser grid (10° x 10°). The total VOC emissions were therefore taken from EMEP, while the speciation was done according to the EPA data. However there was considerable differences between the total emission rates in the two sets of emission data, which means that this is a very significant source to uncertainties in the results. Fig. 1 and 2 show the relative distribution of NO and C₂H₆ emissions applied in the model. In this study no daily variation in the emissions are included (except for isoprene), nor any emissions from sources above the ground (airplanes and lightning).

Isoprene emissions are parameterized as a function of sunlight, surface temperature and deciduous forest cover according to Lubbert and Schopp (1989) and Veldt (1988):

\[ E(\text{kg/h} \times \text{km}^2) = 10^{0.4+T-2.15} \]

E is the hourly emission rate of isoprene in kg km⁻² h⁻¹ during daytime, and \( T \) is the ambient temperature in deg. C.

The data from the NASA/GISS GCM only include 5 day averaged temperatures, so a daily variation in the surface temperature \( T_s \) as a function of the local time \( \text{lt} \), with an amplitude of 10 deg C has been included

\[ T_s = 5.0 \times \sin(\text{MOD}((\text{lt} + 24 - 10) \times 2\pi/24.0) + T_{\text{GCM}}) \]

Fig. 3 shows the calculated relative distribution of isoprene emissions at 15 GMT. The very pronounced maximum results from a combination of very dense forest cover and high surface temperature during the 5 day period. As isoprene emissions are highly uncertain this maximum can at least serve as sensitivity test of the impact of biogenic emissions.

2.4 Boundary conditions and initialization

With such a limited model area, the results will be very sensitive to the fluxes across the vertical boundaries of the model.

The boundary concentration of the species are obtained from results from a 2D channel model (Solberg et al.; 89) covering a zonal band between 30°N and 60°N, and a 2D zonally averaged global model (Isaksen and Hov; 87). The results from the channel model are distributed latitudinally according to the relative latitudinal distribution in the 2D global model to obtain a 3D distribution which can be used to initialize the model, and as boundary conditions. To account for the possibility of a polluted air mass being transported out of the model area and back
3 Results

The model has been run for a 5 day period starting at midnight GMT June 30. Fig. 4 shows the calculated surface concentration of ozone at 13 GMT the 5th day of the simulation. There is a broad maximum of about 70 ppbv over Northwestern Europe and the British Isles, and a secondary maximum over Eastern Europe. Earlier in the day the maximum is shifted towards the east with maximum values at about the same level. Fig. 5 shows the calculated ozone concentrations at 4 different locations during the 5 day period. In the main source regions ozone concentrations show a very pronounced diurnal variation with a maximum concentration occurring usually before noon. The most important reason for the night time reduction is surface deposition (>50% of the loss) and reactions with NOx (R1 > 20% and R2 > 25% in the regions with large NOx emissions).

The maximum surface concentration is very sensitive to the ventilation of the boundary layer through convective transport. As convective transport is only calculated every 4th hour, primary pollutants and their products (ozone, PAN etc.) are allowed to build up during the intervals between the convective events. This effect might give rise to regular over and under predictions, especially in the lowest layer of the model.

The surface ozone concentration and in particular the daily maximum is to a large degree determined by the composition of the emissions. To analyse the sensitivity of the model to variations in the emission rates and composition, several perturbations have been performed. Table 3 shows the maximum ozone concentration during the 5th day of the simulation at 4 locations. The locations were chosen to represent typical source regions.

Again due to a sudden shift in the wind direction, the boundary concentrations approach the upwind concentrations during outflow and return to the original concentrations during a 12 hour relaxation period.

Over Western Europe, represented in Fig. 5 and Table 3 by the grid cell located at 50°N and 5°E, NOx emissions are large and the potential for further photochemical ozone production if NOx emissions increase is low. In some cases it might even be negative, because reaction R1 followed by R3 consumes O3 (and OH) faster than it is produced by R4-R6. The limiting factor for ozone production is then the abundance of RO2 radicals (organic peroxy radicals and RO2) which is produced by oxidation of hydrocarbons and CO by the OH radical.

$\text{NO}_2 + \text{O}_3 + \text{M} \rightarrow \text{O}_3 + \text{M}$  \quad R3

$\text{RO}_2 + \text{NO} + \text{hv} \rightarrow \text{O}_3 + \text{M}$  \quad R4

$\text{NO}_2 + \text{O} + \text{M} \rightarrow \text{O}_3 + \text{M}$  \quad R5

$\text{NO}_2 + \text{O} + \text{M} \rightarrow \text{O}_3 + \text{M}$  \quad R6

In a region with a larger fraction of hydrocarbons in the emissions (42°N, 25°E) there is a large potential for further ozone production if NOx emissions are increased. In this gridcell surface NOx concentration stay around 0.5 ppbv during the period from 10 am to 8 pm, while at the more polluted area over the Netherlands the corresponding NOx concentration is around 2.5 ppbv. Table 3 shows that an increase of 50% in the NOx emissions gives rise to an increase in the calculated surface ozone of 15%.

Over Eastern Europe (50°N, 25°E) the ratio HC/NOx in the emissions are higher than in the west. It is therefore a potential for further ozone production if NOx emissions are increased. A 13% increase in daily maximum ozone at the surface is calculated for a 50% increase in the NOx emissions. At the Scandinavian site (62°N, 15°E) there are small diurnal variations in the ozone concentration after the first day, and the concentration is virtually insensitive to the various perturbations of the emissions. Analysis show that there is net in situ loss of ozone all the time during this period, and a steady-state situation between the local loss processes and transport is established within a few days. However this probably points to an underestimation of the emissions of precursors in this region as observations show evidence of significant ozone production at Scandinavian stations (Grennfelt et al., 87).
results are in general agreement with observations of ozone typical of mid-latitude conditions. The model has been run for a 5 day period in the beginning of July, and the results are in general agreement with observations of ozone typical of this time of the year. Several perturbations have been performed, and the sensitivity of the model is consistent with other model studies. With some refinements of the model (especially with regard to convection) and improvement of the emission database, the model should be well suited to study the processes governing the formation of photochemical oxidants on regional scales.

Table 3: Maximum calculated concentration of ozone (ppbv) in the lowest level during the 5th day at 4 locations.

<table>
<thead>
<tr>
<th>Perturbation</th>
<th>Oz (ppb) at (°N, °E)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Reference</td>
<td>65 79 56 38</td>
</tr>
<tr>
<td>CH₄ conc. = 0</td>
<td>61 78 55 38</td>
</tr>
<tr>
<td>Isoprene em. = 0</td>
<td>64 61 55 38</td>
</tr>
<tr>
<td>NMHC em. * 2</td>
<td>74 83 66 39</td>
</tr>
<tr>
<td>NOₓ em. * 1.5</td>
<td>65 90 63 38</td>
</tr>
<tr>
<td>NOₓ em. * 0.5</td>
<td>63 65 56 38</td>
</tr>
</tbody>
</table>

4 Conclusion

We have developed a 3D chemical tracer model based on the NASA/GISS CTM (Prather et al. 87) for an extended European region with a comprehensive chemical scheme developed to study formation of photochemical oxidants. The model has been run for a 5 day period in the beginning of July, and the results are in general agreement with observations of ozone typical for this time of the year. Several perturbations have been performed, and the sensitivity of the model is consistent with other model studies. With some refinements of the model (especially with respect to convection) and improvement of the emission database, the model should be well suited to study the processes governing the formation of photochemical oxidants on regional scales.

Acknowledgement

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5 References


