

1 **Physical retrieval of surface emissivity spectrum from**
2 **hyperspectral infrared radiances**

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7 **Abstract**

8 Retrieval of temperature, moisture profiles and surface skin temperature from
9 hyperspectral infrared (IR) radiances requires spectral information about the surface
10 emissivity. Using constant or inaccurate surface emissivities typically results in large
11 retrieval errors, particularly over semi-arid or arid areas where the variation in emissivity
12 spectrum is large both spectrally and spatially. In this study, a physically based algorithm
13 has been developed to retrieve a hyperspectral IR emissivity spectrum simultaneously
14 with the temperature and moisture profiles, as well as the surface skin temperature. To
15 make the solution stable and efficient, the hyperspectral emissivity spectrum is
16 represented by eigenvectors, derived from the laboratory measured hyperspectral
17 emissivity database, in the retrieval process. Experience with AIRS (Atmospheric
18 InfraRed Sounder) radiances shows that a simultaneous retrieval of the emissivity
19 spectrum and the sounding improves the surface skin temperature as well as temperature
20 and moisture profiles, particularly in the near surface layer.

1 **1. Introduction**

2

3 Accurate retrieval of atmospheric temperature and moisture profiles, as well as
4 surface skin temperature from hyperspectral infrared (IR) radiance measurements, is
5 needed for climate research, as well as medium range and short-range forecast
6 applications. Hyperspectral IR sounders onboard polar orbiting satellites, such as the
7 Atmospheric InfraRed Sounder (AIRS) (Chahine et al. 2006) on EOS (Earth Observing
8 System) Aqua, the Interferometer Atmospheric Sounding Instrument (IASI)
9 (<http://smc.cnes.fr/IASI/index.htm>) on European METOP-A, and the Cross-track
10 Infrared Sounder (CrIS) (http://www.ipc.noaa.gov/Technology/cris_summary.html) on
11 the next generation National Polar-orbiting Operational Environmental Satellite System
12 (NPOESS), are developed for global temperature and moisture sounding observations
13 with high vertical resolution and high accuracy. Although hyperspectral IR radiances
14 have been successfully assimilated in a global forecast model (LeMarshall et al. 2006),
15 challenges remain over land due to the uncertainty in emissivity.

16 Since the top of atmosphere radiance (TOA) contains a surface IR emissivity (ϵ)
17 contribution (see Figure 1), especially for a channel within the atmospheric window
18 regions, knowledge of surface emissivity is critical for atmospheric temperature and
19 moisture profile retrieval from radiance measurements. The impact of IR emissivity on
20 sounding or surface temperature retrievals has been studied using the GOES
21 (Geostationary Operational Environmental Satellite) Sounder (Plokhenko and Menzel
22 2000) and MODIS (Moderate Resolution Imaging Spectroradiometer) (Wan and Li 1997;
23 Ma et al. 2002; Seemann et al. 2003; Wan et al. 2004). Handling IR surface emissivities

1 in the retrieval process is essential for deriving accurate temperature and boundary layer
2 moisture profiles, as well as surface skin temperature, especially over land. This is
3 equally true for IR radiance assimilation in Numerical Weather Prediction (NWP).
4 Surface emissivity (ϵ) for a given channel is often fixed in the physical retrieval process,
5 for example, using fixed emissivities from a regression approach (Li et al. 2000; Zhou et
6 al. 2006; Zhou et al. 2007).

7 Some physical algorithms also retrieve emissivities together with the sounding,
8 but only at selected channels and spectral bands. For example, Hayden (1988) retrieved
9 emissivities at two spectral bands (longwave and shortwave IR bands) in GOES sounding
10 processing, Zhou et al. (2007) and Susskind et al. (2003) used approximately 40 channels
11 for emissivity retrieval in AIRS retrieval processing. It is difficult to retrieve emissivities
12 of all channels directly in the sounding step, this is due to a large number of unknowns in
13 the inverse equations and the instability of the solution. Retrieving the whole emissivity
14 spectrum is possible if emissivity eigenvectors (EVs) are derived. The hyperspectral
15 emissivity spectrum can be represented in the retrieval process by its EVs derived from
16 laboratory measured hyperspectral emissivity database. Using EVs to represent radiances
17 or parameters to be retrieved has been suggested and attempted by numerous researchers
18 (Smith and Woolf 1976; Huang 1998; Zhou et al. 2006; Liu et al. 2006).

19 Knowledge of surface IR emissivity is also very important for creating a climate
20 forecasts (Jin and Liang 2006). Data from a satellite based IR imager such as MODIS
21 provide global emissivity distribution at a few IR spectral bands (Wan et al. 2004). With
22 hyperspectral IR data available, a global map of hyperspectral IR emissivity spectra is
23 possible.

1 Based on a physical iterative approach, this study demonstrates that a
2 hyperspectral emissivity spectrum can be retrieved simultaneously along with
3 temperature and moisture soundings, as well as surface skin temperature from a
4 hyperspectral IR radiance spectrum by using EV representation. This approach has been
5 successfully tested using both simulated and measured AIRS radiances, and is expected
6 to help improve the hyperspectral IR radiance assimilation in forecast models over land.
7 For example, one can use a variational (1DVAR) approach to derive emissivity properties
8 and other atmospheric parameters, and use a four dimensional variational (4DVAR)
9 approach to directly assimilate those derived products in a forecast model (Weng et al.
10 2007). A further goal of this research is to study the sounding improvement in the
11 physical method over the regression technique in handling surface IR emissivities.

12

13 **2. Methodology**

14

15 With pre-determined surface IR emissivities, algorithms for retrieving the
16 atmospheric temperature and moisture profiles, as well as surface skin temperature, have
17 been developed to process single field-of-view radiance measurements (Li and Huang
18 1999; Ma et al. 1998, Li et al. 2000, Zhou et al. 2003). Since emissivity is wavenumber
19 dependent, it is difficult to retrieve emissivities at all channels together with temperature
20 and moisture profiles due to a large number of unknowns. To take advantage of spectral
21 correlations, the emissivity spectrum can be represented by its EVs (e.g., the first 6 EVs)
22 in the retrieval process, leaving only a few unknowns (emissivity EV coefficients) to be
23 added together with the temperature profile ($T(p)$), moisture profile ($q(p)$) and surface

1 skin temperature (Ts) in the 1DVAR process. In addition to the regular unknowns (T(p),
 2 q(p), Ts), the emissivity spectrum $\vec{\varepsilon} = (\varepsilon_1, \varepsilon_2, \dots, \varepsilon_N)$, where N is the total number of
 3 channels used, is expressed by it EVs

$$4 \quad \vec{\varepsilon} = \sum_i^l \varphi_i a_i = \phi \vec{a} \quad (1)$$

5 where φ_i is the i th EV and a_i is the associated EV coefficient, and l is the number of
 6 EVs used. ϕ and \vec{a} are the corresponding EV matrix and EV coefficient vector, respectively.
 7 Figure 1 (lower panel) shows the first 6 EVs for the AIRS spectrum derived from
 8 laboratory measurements of hyperspectral emissivity spectra. Our study shows that the
 9 first 6 EVs (6 pieces of independent emissivity information) are representative of the
 10 emissivity spectrum information in a simultaneous retrieval process. The Jacobian matrix
 11 of the radiance with respect to the eigenvector coefficient can be derived

$$12 \quad J_a = J_\varepsilon * \phi \quad (2)$$

13 where J_a is the Jacobian matrix of the radiance with respect to the emissivity EV
 14 coefficient, while J_ε is the diagonal matrix with Jacobians corresponding to the
 15 emissivity spectrum, and the diagonal values can be calculated approximately by an
 16 analytical method (Li et al. 1994). Figure 1 shows the AIRS brightness temperature (BT)
 17 spectrum calculated from the U.S. standard atmosphere (top panel) and associated
 18 emissivity Jacobian spectrum (middle panel). A Jacobian value of 50 means that a
 19 change in emissivity of 0.01 results in a 0.5 K change in BT. The longwave IR window
 20 region has a larger emissivity signal than the shortwave IR window region, which is
 21 important to note since a good signal-to-noise ratio is required to retrieve emissivity

1 spectrum according to the Jacobian analysis. The convoluted Jacobian from Eq.(2) then
2 can be used in the physical retrieval process.

3

4 **3. Experiment with simulated AIRS radiances**

5

6 The algorithm has been tested with both simulated and measured AIRS radiances.
7 In the simulation study, a global set of training profiles (Seemann et al. 2007) was used.
8 Each profile contains a temperature profile, water vapor mixing ratio profile, ozone
9 profile and surface skin temperature; emissivities at 10 spectral points have been assigned
10 to each profile based on the combination of global MODIS emissivity measurements
11 (Wan and Li, 1997; Wan et al., 2004) and laboratory emissivity measurements
12 (<http://www.ices.ucsb.edu/modis/EMIS/html/em.html>; <http://speclib.jpl.nasa.gov/>). By
13 using a similar approach and applying emissivity EVs derived from the laboratory, each
14 profile of the training dataset is assigned a hyperspectral emissivity spectrum (e.g., at
15 AIRS full spectral coverage). Figure 2 shows the emissivities assigned to ocean (upper
16 left), grassland (upper right), cropland (lower left) and desert (lower right) regions. In the
17 simulation study, an AIRS radiance spectrum is calculated using the fast and accurate
18 Stand-Alone Radiative Transfer Algorithm (SARTA) developed by University of
19 Maryland Baltimore County (UMBC) for each training profile. The AIRS instrument
20 noise plus 0.2 K forward model errors are added to the simulated radiances. The
21 temperature and moisture retrieval algorithm is a two-step approach: regression followed
22 by a physical iterative approach (Li et al. 2000). The regression technique provides a
23 reasonable hyperspectral emissivity spectrum retrievals. For example, Zhou et al. [2006]

1 have applied regression for NASTI emissivity retrievals, and Zhou et al. [2007] have
2 used the regression for AIRS emissivity retrievals. Physical retrieval of sounding and
3 surface IR emissivities at the selected channels in a sequential way was performed in the
4 operational AIRS product generation (Susskind et al. 2003). In this study, the
5 simultaneous retrieval of a sounding and the whole emissivity spectrum in a physical
6 iterative approach is developed in an attempt to improve statistical results. The following
7 three configurations are examined:

- 8 (1) Use a constant emissivity of 0.98 in the physical retrieval, and the emissivity is
9 not changed in each physical iteration;
- 10 (2) Use a regression emissivity spectrum in the physical retrieval, and the emissivity
11 is not changed in physical each iteration;
- 12 (3) Use a regression emissivity spectrum as the initial guess in the physical retrieval,
13 and the emissivity is updated in each physical iteration.

14 In the simulation, 90% of the profiles are used as training for the regression coefficients,
15 while the remaining 10% of the profiles are used as independent testing. The temperature
16 and water vapor relative humidity (RH, 0 – 100%) root mean square errors (RMSE) are
17 calculated for the above configurations; the RMSE is based on the absolute difference
18 between the truth and the retrieval.

19 Figure 3 shows the retrieved RMSE for the above three configurations along with
20 the first guess (from the regression). The first guess provides a reasonable profile with an
21 accuracy of approximately 10% for water vapor RH and 1 K above 500 hPa; the accuracy
22 for temperature is limited in the boundary layer from the regression. With a fixed
23 constant emissivity, the physical retrieval significantly degrades the first guess for both

1 temperature and water vapor since the assumed emissivity of 0.98 is not accurate. As
2 expected, when the regression based emissivity spectrum is fixed in the physical
3 iterations, the temperature and moisture are improved from the first guess, especially for
4 water vapor, due to the nonlinear contribution of IR radiances to the temperature and
5 water vapor. With a simultaneous retrieval of the sounding and emissivity spectrum, the
6 temperature and moisture retrievals are the best in all three configurations, especially in
7 the boundary layer where emissivity has significant contributions. Configuration 3
8 improves over configuration 2 significantly. The retrieval simulation illustrates that
9 estimating emissivity in the physical iteration is necessary and helpful for sounding
10 retrievals, especially in desert regions where emissivity variation is large both spectrally
11 and spatially. In addition, the emissivity RMSE from both the regression and the
12 physical retrieval are also shown in the upper panel (from configuration 3), which
13 demonstrates that the physical approach improves the regression. However, the
14 shortwave physical retrieval still has a retrieval error of 0.02 due to the limited emissivity
15 information (see Figure 1). The surface skin temperature retrieval indicates similar
16 results to the boundary layer temperature as shown in Table 1.

17

18 **4. Experiment with measured AIRS radiances**

19

20 The algorithm has also been tested with AIRS radiance measurements using
21 granule 011 for 08 September 2004. The MODIS cloud mask is used to identify the
22 AIRS clear footprints (Li et al. 2004). The AIRS granule contains various surface types
23 (cropland, desert, ocean etc.). Figure 4 shows the emissivity spectrum retrieval from the

1 regression (upper left) and physical (upper right) approaches at 1227 cm^{-1} or $8.15\text{ }\mu\text{m}$.
2 The difference between the physical and first guess (regression) can be seen, especially
3 over the desert region. The lower panel shows one example of an emissivity spectrum
4 retrieval over the desert; the physical approach changes the regression in both longwave
5 and shortwave window region. Three emissivity spectrum references derived from the
6 laboratory database, representing the surface types of desert, cropland, and ocean
7 respectively, are also shown. Accurate surface properties captured by hyperspectral
8 measurements over land, especially in the vicinity of the Sahara Desert, are clearly
9 evident. Sounding in the boundary layer leads to greater improvement in physical
10 retrievals over regression retrievals when compared with the ECMWF analysis (not
11 shown).

12

13 **5. Summary**

14

15 Handling surface IR emissivity is very important for sounding retrieval and
16 radiance assimilation. The emissivity uncertainty has a significant impact on the retrieval
17 of boundary layer temperature and moisture, especially over desert regions where surface
18 IR emissivity has large variations both spectrally and spatially. This study shows that
19 simultaneous retrieval of hyperspectral IR emissivity spectrum and sounding is helpful in
20 the sounding retrieval process. The emissivity spectrum can be retrieved together with the
21 profile through an EV representation of the spectrum; a representative laboratory
22 hyperspectral IR emissivity measurement data set containing various ecosystem types are
23 crucial for EVs. With such a technique the global IR emissivity spectrum product can be

1 derived through composite clear hyperspectral IR radiance measurements. The derived
2 hyperspectral IR emissivity product is very useful for processing broad IR spectral band
3 radiances such as from the Advanced Baseline Imager (ABI) (Schmit et al. 2005)
4 onboard the next generation of Geostationary Operational Environmental Satellite
5 (GOES-R) and beyond (e.g., using retrieved emissivity spectra from polar orbiting
6 hyperspectral IR radiances for processing ABI IR radiances). The global emissivity
7 product is also very important for improving the global climate forecast. The algorithm
8 can similarly be applied to process IASI and CrIS.

9

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1 **Figure captions**

2 Figure 1. The AIRS brightness temperature (BT) spectrum calculated from the U.S.
3 standard atmospheric profile (top panel) and associated emissivity Jacobian spectrum
4 (middle panel). The lower panel shows the first 6 EVs of emissivity spectra derived from
5 hyperspectral laboratory measurements.

6

7 Figure 2. The selected emissivity spectra assigned to ocean (upper left), grassland (upper
8 right), cropland (lower left) and desert (lower right) regions from the training data set.
9 The dark black lines are the means for each regional data set.

10

11 Figure 3. The root mean square errors (RMSE) of retrievals for three configurations
12 described in the text along with the first guess (from regression) results. The first guess
13 provides a reasonable profile with an accuracy of approximately 10% for water vapor RH
14 and 1 K above 500 hPa.

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16 Figure 4. The emissivity retrieval from the regression (upper left) and physical (upper
17 right) approaches at 1227 cm^{-1} or $8.15\text{ }\mu\text{m}$. The lower panel shows one example of an
18 emissivity spectrum retrieval over the desert region along with three reference emissivity
19 spectra from laboratory data corresponding to ocean, cropland and desert surface types,
20 respectively.

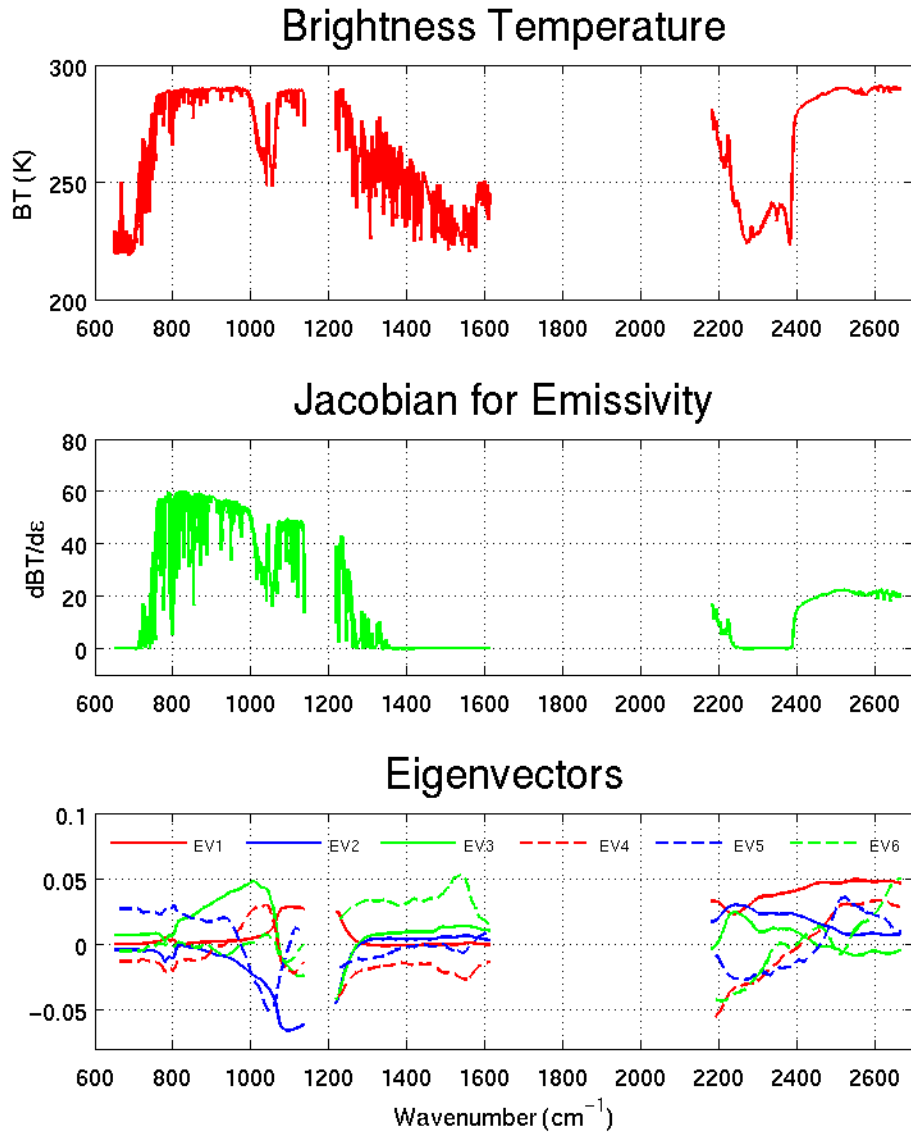
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1

2 **Table captions**

3 Table 1. The retrieved surface skin temperature root mean square error for three
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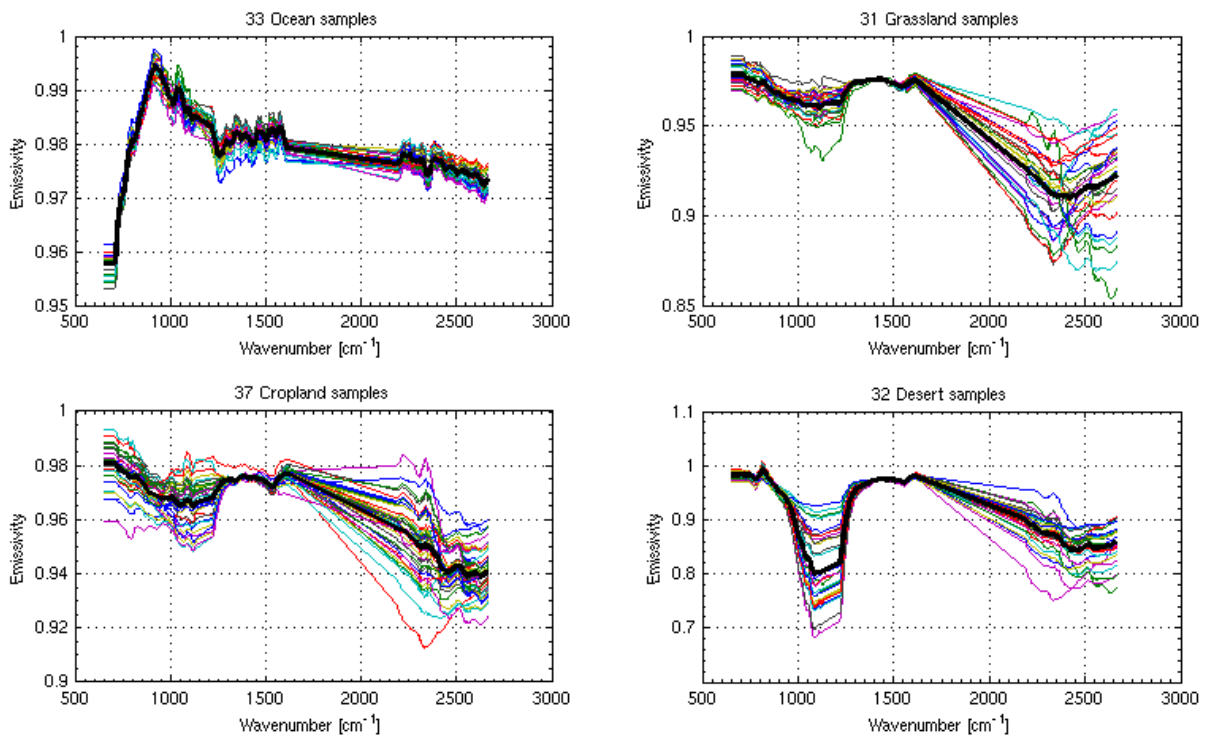
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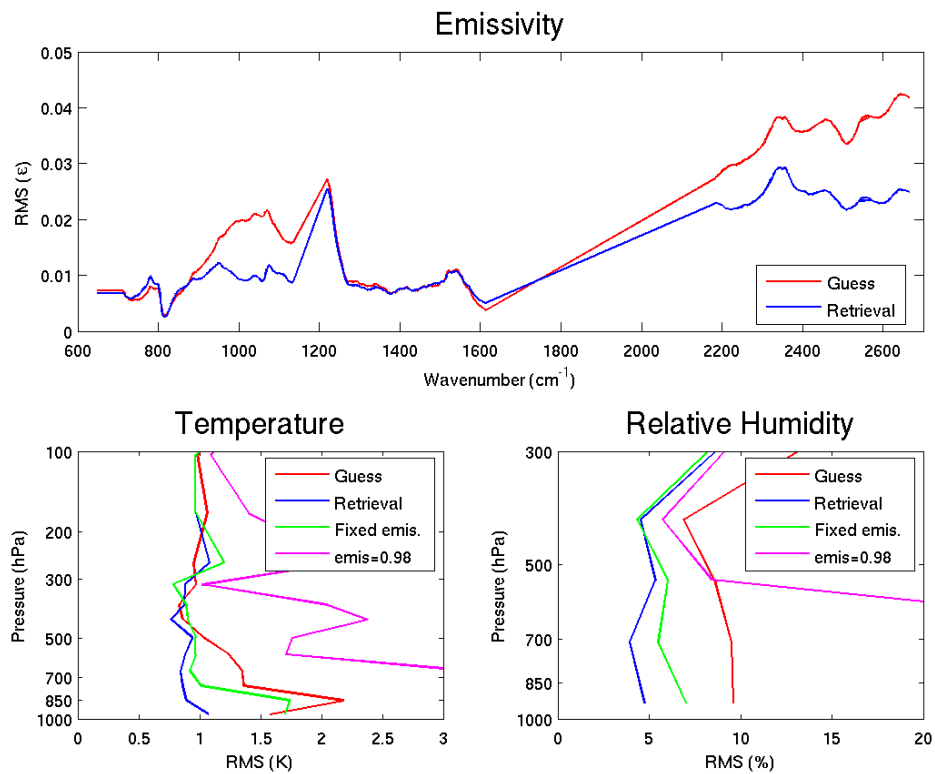
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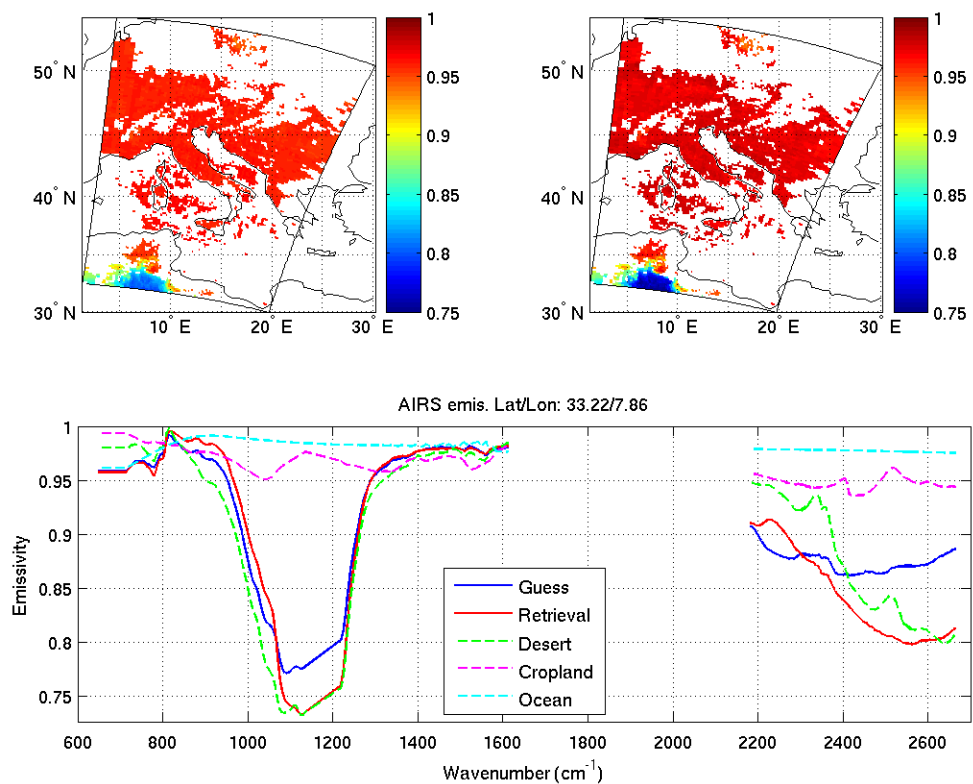


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6 respectively.

7



8

1 Table 1. The retrieved surface skin temperature root mean square error for three
2 configurations described in the text along with the regression results.

3

Method	Cropland RMS (K)	Desert RMS (K)	Grassland RMS (K)	Ocean RMS (K)
Reg	0.485	0.624	0.461	0.703
Rtv (configuration 3)	0.327	0.540	0.316	0.472
Fixed emis (configuration 2)	0.360	0.822	0.421	0.563
Emis=0.98 (configuration 1)	0.686	9.544	0.877	0.409

4